

Editorial Manager(tm) for Marine Geophysical Research  
Manuscript Draft

Manuscript Number:

Title: Seafloor mapping for geohazard assessment: state of the art

Article Type: SI Seafloor Mapping

Keywords: geological risks, multibeam bathymetry, natural hazards, seafloor morphology, submarine landslides

Corresponding Author: Francesco Latino Chiocci, Ph.D.

Corresponding Author's Institution: University of Rome "La Sapienza"

First Author: Francesco Latino Chiocci, Ph.D.

Order of Authors: Francesco Latino Chiocci, Ph.D.;Antonio Cattaneo, PhD;Roger Urgeles

Abstract: During the last two decades, increasing use of full-coverage sonic mapping of the seafloor, has made us more aware of the large and different number of seafloor processes and events bearing significant geohazard potential. This awareness combines with the increasing use of seafloor for infrastructures and with the high density of population and settlement on the coast.

Seafloor mapping is the first step to make a census of the geohazard-bearing features present in a given offshore area. It often provides the only tool for a comprehensive seafloor geohazard assessment over large areas, scarcely groundtruthed by acoustic prospection and seafloor sampling. Indeed, the characterization of geohazard features on a morphological basis alone is however limited, as and more detailed investigations are needed to define the character and state of activity of potentially hazardous features. Such investigations include the use of deep-tow or autonomous platforms designed to acquire HR data at depth as well as in situ measurements, both being very expensive activities not applicable over large areas. This is the reason why seafloor mapping is often not only the first and the main but also the only tool for a comprehensive seafloor geohazard assessment over large areas, often scarcely groundtruthed by acoustic prospection and seafloor sampling.

This special issue represents an example of the diversity of approaches to seafloor geohazard assessment and summarizes the present state of this discipline. Both the diverse technologies applied and the specific aims of offshore geohazard assessment brought different communities to deal with the study of seafloor processes/events from remarkably distinct viewpoints. We identified three end members in offshore geohazard assessment: 1) geohazard assessment "sensu stricto", 2) "engineering" geohazard assessment, 3) "non-specific" geohazard assessment. These are being conducted by industry, academia and public agencies in charge of civil protection and land planning and management. Understanding the needs and geohazard perception of the different groups is a necessary step for a profitable collaboration in such an interesting and rapidly developing field of marine geology.

Suggested Reviewers: none none  
none@none.it

none2 none2  
none2@none.it

none3 none3

none3@none.it

none4 none4  
none4@none.it

# 1                    **Seafloor mapping for geohazard assessment:** 2                    **state of the art**

3                    Francesco L. Chiocci <sup>1,2</sup>, Antonio Cattaneo <sup>3</sup>, Roger Urgeles <sup>4</sup>

4                    1 University of Rome "La Sapienza", Italy

5                    2 National Reserach Council IGAG, Rome, Italy

6                    3 Ifremer, GM-LES, Plouzané, BP70 29280, France

7                    4 Institut de Ciències del Mar (CSIC), Barcelona, Spain

8  
9                    francesco latino chiocci, Dip. Scienze della Terra, Università di Roma "La Sapienza", tel 0039-06-4994-  
10                    4308 (tel. e fax), 0039-338-5942311 (mob.); francesco.chiocci@uniroma1.it

## 11 12                    **Keywords**

13                    geological risks, multibeam bathymetry, natural hazards, seafloor morphology,  
14                    submarine landslides

## 15                    **Abstract**

16                    During the last two decades, increasing use of full-coverage sonic mapping of the seafloor, has made  
17                    us more aware of the large and different number of seafloor processes and events bearing significant  
18                    geohazard potential. This awareness combines with the increasing use of seafloor for infrastructures  
19                    and with the high density of population and settlement on the coast.

20                    Seafloor mapping is the first step to make a census of the geohazard-bearing features present in a  
21                    given offshore area. It often provides the only tool for a comprehensive seafloor geohazard  
22                    assessment over large areas, scarcely groundtruthed by acoustic prospection and seafloor sampling.  
23                    Indeed, the characterization of geohazard features on a morphological basis alone is however limited,  
24                    as and more detailed investigations are needed to define the character and state of activity of  
25                    potentially hazardous features. Such investigations include the use of deep-tow or autonomous  
26                    platforms designed to acquire HR data at depth as well as in situ measurements, both being very  
27                    expensive activities not applicable over large areas. This is the reason why seafloor mapping is often  
28                    not only the first and the main but also the only tool for a comprehensive seafloor geohazard  
29                    assessment over large areas, often scarcely groundtruthed by acoustic prospection and seafloor  
30                    sampling.

31                    This special issue represents an example of the diversity of approaches to seafloor geohazard  
32                    assessment and summarizes the present state of this discipline. Both the diverse technologies applied  
33                    and the specific aims of offshore geohazard assessment brought different communities to deal with  
34                    the study of seafloor processes/events from remarkably distinct viewpoints. We identified three end  
35                    members in offshore geohazard assessment: 1) geohazard assessment "sensu stricto", 2)  
36                    "engineering" geohazard assessment, 3) "non-specific" geohazard assessment. These are being  
37                    conducted by industry, academia and public agencies in charge of civil protection and land planning  
38                    and management. Understanding the needs and geohazard perception of the different groups is a  
39                    necessary step for a profitable collaboration in such an interesting and rapidly developing field of  
40                    marine geology.

41

42 **1. Introduction**

43 In May 2009, an international conference on “Seafloor Mapping for Geohazard  
44 Assessment” was held in the charming location of the Giardini “La Mortella” in the  
45 volcanic island of Ischia (Naples Gulf, Italy) within the umbrella of the IGCP project  
46 511 and the MaGIC Italian National Project. Ischia island was affected in the recent  
47 geological past by a massive sector collapse and debris avalanche (Chiocci and De  
48 Alteriis 2006, De Alteriis et al. 2010), whose deposits and scar morphology were the  
49 object of a pre-conference fieldtrip.

50 The conference brought together some hundred participants from 12 countries,  
51 representing the academic, industry and public administration communities. The  
52 participants presented several case studies and discussed on the state of the  
53 knowledge of geohazard assessment in the submarine domain, focusing especially  
54 on the use of seafloor mapping by multibeam echosounder data complemented by  
55 other geophysical and in situ data and measurements for the purpose of geohazard  
56 assessment(Chiocci et al. eds., 2009).

57 Noticeably, about one third of the presentations were directly or indirectly linked to,  
58 and given by, representatives of industry and private companies working for  
59 geohazard offshore surveys, witnessing the applied nature of the topic. It was made  
60 evident early during the conference that industry and academia do not share the  
61 same point of view as to what represents a geohazard and what is simply a  
62 constraint for offshore activities (see section 3). The Special Issue resulting from the  
63 conference contribution aims, among others, to bridge the gap between academic  
64 research (not much funding, relative long-term development research themes) and  
65 the industry practice (spectacular, often very costly dataset focused on specific,  
66 applied issues, with little time to carry out bibliographic comparisons and a  
67 comprehensive view of the study area). Both the Ischia conference and this volume  
68 cover different areas worldwide, with higher density in the Mediterranean Sea  
69 (Fig.1).

70

71 **2 Seafloor morphology exploration in the recent decades: seafloor mapping**  
72 **programs and offshore geohazard investigations**

73 Because of the widespread and always increasing use of multibeam and other  
74 mapping techniques, marine geology is undergoing a very stimulating and active  
75 historical moment as the scientific community is flooded by a massive amount of  
76 extremely high-resolution morpho-bathymetric datasets that allow an  
77 unprecedented look at the seafloor. The impact of these data is tremendous,  
78 sometimes comparable with the effect that the production of the first bathymetric  
79 map of ocean floor (Heezen and Tharp 1954, Heezen et al. 1959) exerted on the

80 understanding of global tectonic processes that lead to the definition of plate  
81 tectonics. Thanks to multibeam mapping techniques, our ability to produce a near-  
82 continuous bathymetric surface of the seafloor has revolutionized our understanding  
83 of marine morphodynamics. This has significantly changed the detail with which we  
84 can interpret seafloor tectonic and sedimentary process at the origin of geohazards  
85 (Hughes Clarke et al. 1996).

86 Multibeam sonars underwent rapid development during the '80s (Farr 1980, Brown  
87 and Blondel 2008), although the number of surveys at sea using this technique was  
88 limited. Since the '90s, more performing systems have been used and technology  
89 improved to integrate acquisition of echosoundings including backscatter (de  
90 Moustier 1988, Lurton 2002). Today, the modern swath mapping sonar systems  
91 collect concurrent swath bathymetry and acoustic backscatter data. When processed  
92 together, these data can be used for remote sensing of seafloor characteristics, with  
93 several applications in geotechnical and geohazard surveys (e.g., offshore drilling and  
94 mining, dredging and disposal, subsea cable and pipeline routes; de Moustier et al.  
95 2010).

96 It is only during the last 10-15 years that multibeam echosounders become a  
97 common tool in any kind of subaqueous survey, from very shallow coastal and  
98 lacustrine environments to full ocean depths, providing increasing amounts of high  
99 resolution data. These data constitute nowadays the unavoidable base for any  
100 geohazard assessment study. The use of AUVs and other near seafloor platforms has  
101 extended high-resolution surveys to the deep water. In addition, bathymetric data  
102 extracted from 3D seismic volumes rival in horizontal resolution with multibeam  
103 echosounders and allow the comparison of seafloor and buried seismic surfaces with  
104 a seismic geomorphology approach (Posamentier 2007).

105 As far as seafloor geohazard is concerned, multibeam bathymetry is not only able to  
106 identify features representing the trace of hazardous geological processes, but may  
107 also precisely measure morphometric parameters and their variation through time, to  
108 enhance the monitoring of seafloor changes representing active processes.  
109 Multibeam bathymetry also provides boundary conditions and inputs to numerical  
110 modeling aiming to reproduce geological processes to make predictions for future  
111 scenarios (e. g. run out of failed sediment, tsunami modeling).

112 Obtaining a comprehensive seafloor map of homogeneous resolution at the regional,  
113 EEZ and oceanic level has been the focus of numerous initiatives. In the pre-  
114 multibeam era, large scale seafloor mapping has been first carried-out using low-  
115 frequency, long-range side scan sonar systems such as GLORIA (Geological Long-  
116 Range Inclined Asdic, Rusby and Revie 1975). A long-standing initiative to merge,  
117 collect, homogenize and edit existing datasets of mainly single-beam soundings such  
118 as GEBCO (General Bathymetric Chart of the Oceans, Carpine-Lancre et al. 2003, Hall

119 2006) is being carried out with the support of international organizations (UNESCO,  
120 IHO) and hydrographic national offices.

121 In the multibeam era, several projects aimed to acquire full -overage multibeam data  
122 for different purposes such as the Norwegian Mareano (Thorsens 2009, Dolan et al.  
123 2009), the Californian CSMC (California Seafloor Mapping Program, OPC, 2007 ), the  
124 Irish INSS-INFOMAR (Irish National Seabed Survey), the mapping of the Spanish EEZ  
125 (Muñoz et al. 1998), the European MESH (Mapping European Seabed Habitats) and  
126 the Italian MAGIC (Marine Geohazard Along the Italian Coasts). At a Mediterranean  
127 scale, an effort to compile existing data in deep water is being made by CIESM and  
128 Ifremer (MediMap Group 2005).

129

130

### 131 **3. Geohazards in the marine realm**

132 Kvalstad (2007) defined geohazard in the offshore domain, as “local and/or regional  
133 site and soil conditions having a potential of developing into failure events causing  
134 loss of life or damage to health, environment or field installations”. Such definition is  
135 broad enough to include most of the geological processes shaping the seafloor, as no  
136 time frame is defined and potentially any modification of the seafloor may be  
137 damaging for structures resting on it.

138 Geohazards are one of the elements in the equation of geological risk. The geological  
139 risk is in fact the product of the occurrence of a given geological event or process  
140 (geohazard), the vulnerability of a given area or region to this geohazard and the  
141 possible damaging consequences on humans, structures and/or the environment.

142 In marine and coastal environments, the major geological hazards are linked to the  
143 occurrence of events such as earthquakes, volcanic eruptions, submarine landslides  
144 or rapid processes that are able to modify the morphology and character of the  
145 seafloor such as gravity-driven sediment flows, fluid emissions, bedform migration,  
146 retrogressive erosion at canyon heads, etc (Fig. 2). Secondary effects such as  
147 tsunamis (either triggered by earthquakes or landslides) also need to be considered,  
148 as both their genesis and propagation is strongly controlled by seafloor morphology.

149 Despite the definition of geohazards, though debatable, is relatively straightforward,  
150 their assessment is relatively complex. In fact, for the marine realm, it is often  
151 difficult if not impossible to define:

152 **1) the time scale of the hazard.** Seafloor sector collapses and related debris  
153 avalanches are common processes in the millennial-scale geologic evolution of  
154 volcanic islands, and thus an event bearing an extremely high damage potential

155 might reoccur in the future. However, do they have to be considered in a geohazard  
156 assessment if no signs of activity are present? In a similar way, retrograding canyon  
157 heads, that seldom affect the coastline, often show a clear match between subaerial  
158 and submarine erosional morphologies, showing that landslides shaped the coast  
159 inshore of the canyon head. The question is therefore: should this be considered a  
160 geohazard if no signs of instability are detected but there is only the knowledge that  
161 a geological process could cause mass wasting in an undefined future?

162 **2) the capability of a given process to produce dangerous effects.** Our knowledge  
163 on submarine processes is in many cases very limited. Do we have to assume a  
164 conservative approach (worst-case scenario) in any instance? For example, for  
165 landslide-generated tsunamis there are physical models linking the volume of the  
166 mobilized sediment and water depth to the tsunami wave height. Is it reasonable to  
167 consider all slide scars, caused by different failures (rotational, translational, complex  
168 etc.), as a result of catastrophic events, when we do not know precisely the failure  
169 dynamics? What would be the result of the assessment if many or most of the  
170 failures were slow enough not to produce significant effects on the overlying water  
171 column?

172 **3) the recurrence time of most of the hazardous events.** Sometimes only erosional  
173 scars witness the occurrence of mass wasting in a given setting, so hindering any  
174 possible definition of age and recurrence time of landslide events. For instance mass  
175 wasting at canyon heads is the main process for their genesis and retrogression, as  
176 witnessed by the multiple, complex and overlapping landslide scars that form it.  
177 However mass wasting there is usually occurring with the complete disintegration of  
178 the failed mass and its transformation into gravity flows dispersing down canyon. In  
179 this case we know that the process is frequent, but what does “frequent” mean?

180 **4) the frequency – magnitude relationship of observed geohazards.** In addition to  
181 frequency-dependant attenuation, hull mounted multibeam systems have an  
182 inherent loss of resolution with increasing water depth. Therefore, the capability to  
183 depict and characterize geohazard-bearing features is water depth-dependant. This  
184 problem is shown in Figure 3 with an example of size distribution of submarine  
185 landslides in the Mediterranean basin with data compiled from the public literature  
186 (Camerlenghi et al. 2010). The fact that landslides appear to have a characteristic  
187 size-magnitude, probably reflects our inability to image small-sized landslides  
188 (hundreds of m<sup>2</sup> or lower), or/and to a subjective underestimate by the operator  
189 (small=not hazardous?).

190 The impact of offshore geohazards for coastal communities is also depth-dependent,  
191 i.e. the deeper the geohazard generating feature, the smaller the potential impact,  
192 so this somehow counterbalances our inability to depict small geohazard features in

193 deep water. However, the impact of geohazard processes for offshore infrastructure,  
194 can be very high, also in deep water, no matter the size of the geohazard generating  
195 feature. To confront this, industry is massively using AUVs and ROV-mounted  
196 multibeam systems for seafloor mapping to achieve the high resolution of hull-  
197 mounted coastal surveys also in very deep waters.

198 **5) the spatial distribution of hazardous events.** In open slope settings, it is common  
199 practice to highlight features such as submarine landslides as evidence for  
200 geohazard, and pipeline and cable routes are often set to avoid these features.  
201 However, how likely is that the next failure occurs over the same area where failure  
202 already occurred? Isn't it more likely that the next failure occurs in a nearby area  
203 that shows no current signs of seafloor disturbance? In other words, is it a wiser  
204 decision to lay a pipeline through a submarine landslide scar or through an area that  
205 apparently shows no signs of recent failure but is near to another area that failed?

206 Despite the above mentioned limitations, geohazards have to be defined, depicted  
207 and evaluated because the use of the seafloor for settling structures continues to  
208 increase as is the coastal population. Therefore, assessment of marine geohazards  
209 must be a key element in coastal and seafloor management. For instance the always  
210 increasing use of the seafloor for cable route and drilling facilities (more and more  
211 extending in deep water), let marine geohazard to be a major concern for industries  
212 and public agencies dealing with marine infrastructures. Also onshore, for coastal  
213 communities and structures, marine geohazards may produce either direct effect as  
214 for submarine landslides retrogressively propagating onshore such as Finneidfjord in  
215 1996 (Longva et al. 2003) or Stromboli in 2002 (Chiocci et al. 2008) or indirect effects  
216 as for submarine landslide generating or contributing to tsunamis such as Great  
217 Banks in 1929 (Piper et al. 1999) or Nice in 1979 (Malinverno et al. 1988).

218 For geologically active regions (such as the Mediterranean, the Caribbean and many  
219 regions around the Pacific belt of fire), which are often highly touristic regions, the  
220 risk is extremely high as the scenic coasts are often volcanic or fault-controlled and  
221 tourists and coastal settlements are packed in narrow coastal belts and constricted  
222 pocket beaches. In such a setting, even limited coastal landslides or tsunamis may  
223 have enormous effects, as witnessed by the Nice 1979 landslide and subsequent  
224 tsunami (Sultan et al. 2010).

#### 225 **4. Diverse approaches and characterizations of geohazard**

226 One of the main outcomes of the Ischia conference and of this volume is the  
227 evidence that there are **three end-member approaches in geohazard**  
228 **characterization** that we hereafter refer as 1) "geohazard assessment sensu-stricto";  
229 2) "engineering geohazard assessment" and 3) "non specific geohazard assessment"  
230 (Fig. 4).



231 The **“geohazard assessment sensu-stricto”** is the one conforming to the classical  
232 definition (see beginning of section 2) and is typical of applicative surveys as it is  
233 aimed to precisely identify one or more hazards in a specific site and the possibility  
234 this hazard might occur in a given time span (see for instance L’Heureux et al. this  
235 volume, Strasser et al. this volume). This approach needs a precise knowledge of the  
236 character of the process/event considered hazardous, its recurrence time, the  
237 present state of the seafloor and subseafloor and the factors controlling it. For  
238 instance, for a submarine landslide, the knowledge of the state of stability of the  
239 slope (based on geomechanical characters of the potentially failing mass) and the  
240 definition of possible triggers of the instability is needed.

241 The study should therefore rely on a suite of geophysical data and in situ  
242 measurements that made it very costly and only possible for small areas and limited  
243 number of features/hazardous processes. The outcome of the study however is not  
244 only a truthful assessment of the hazard but may also define the character of the  
245 infrastructure in order to reduce the vulnerability to the hazard. In this case, a real  
246 risk assessment is possible.

247 The **“engineering geohazard assessment”** considers the presence of any uneven  
248 feature at the seafloor that constitutes a geohazard in itself as it is a development  
249 constraint that may impact the seafloor structure, if not taken into the due account.  
250 As an example the presence of a rock outcrop or a slope-break at the seafloor is  
251 considered by the industry a geohazard to be avoided or carefully considered for  
252 operational planning of a cable or pipeline route to avoid free spans. In strict terms  
253 these uneven features (ridges, boulders, ...) should not be considered as real  
254 “geohazards” because by themselves they have no “potential of developing into  
255 failure events” (Kvalstad 2007). In other words, the hazard results from human  
256 action at the location of this feature without the feature itself necessarily develops  
257 into a failure event. Furthermore, in this approach, the time span considered for the  
258 assessment is usually comparable with the life span of the infrastructure at the  
259 seafloor (e.g., pipeline, submarine cable), generally not longer than a few decades  
260 and thus shorter in time of the return period of most geological processes (see for  
261 example Dyer this volume, Cecchini et al. this volume).

262 Finally the **“non-specific geohazard assessment”** involves the census of potentially  
263 hazardous features present in a whole region without targeting at any specific  
264 process/event and at any specific effect on exposed good. Such assessment is aimed  
265 to define in a general perspective the presence and character of all the hazardous  
266 marine geological processes/events and possibly the seafloor and subseafloor  
267 predisposition to these processes/events. This approach does not attempt to define  
268 the precise age and recurrence period of the hazard; it rather evaluates the spatial  
269 occurrence of hazardous events and highlight the possibility that similar events

270 occur in the future, given the similarity of the morpho-structural and  
271 lithostratigraphic setting of the surroundings.

272 In strict terms the “non-specific geohazard assessment” should not be considered as  
273 a real geohazard assessment but, in our perception and given the present  
274 technology, is the only possible way to define at a regional scale the location, type  
275 and characteristics of the geohazards present in a given region, with possible  
276 indication on their state of activity. Cost-efficient multibeam mapping is in this  
277 respect the fundamental tool for such assessment, that therefore relies essentially  
278 on geomorphic interpretation of high-resolution bathymetric data. Examples of this  
279 type of geohazard assessment are provided in Chiocci and Ridente (this volume),  
280 Larroque et al. (this volume) and Lo Iacono et al. (this volume).

281 Assessment of geohazards according to these 3 end-members is very clearly divided  
282 by the approach that different communities have when studying geohazards (Fig. 4).  
283 In fact, Academia (universities and research institutes) moves between the  
284 “geohazard assessment sensu-stricto” and the “non-specific geohazard assessment”  
285 approaches, as it is mainly interested in understanding the processes and events  
286 both at small and large scale. Academic studies are often focused in small areas that  
287 are investigated in detail to characterize the seafloor and subseafloor. Academia is  
288 also interested in the overall geological evolution of continental margins and  
289 therefore the “non-specific geohazard assessment”, is seen as a way to understand  
290 the processes shaping the seafloor and the relationship between processes and long-  
291 term morphological evolution.

292 On the contrary, industry moves between the “engineering geohazard assessment”  
293 and the “geohazard assessment sensu-stricto”, and often these two are performed  
294 in sequential order. The need to find the safest route or site for laying and installing  
295 infrastructures at the lowest possible economic cost often means that even small  
296 seafloor features have to be considered in geohazard assessment, as they constrain  
297 the deployment of the structure.

298 Finally, stakeholders and public administrations (including geological surveys) are  
299 interested in the “non-specific geohazard assessment”, needed to wholly assess the  
300 risk for public goods in a given region and the “engineering geohazard assessment”  
301 approach needed for public works and land management.

## 302 **5. Special Issue outline**

303 Here follows a succinct outline of the contributions of the Special Issue, grouped by  
304 thematic affinity. Due to the location of the venue, the majority of the contributions  
305 show case studies from European, and in particular Mediterranean, continental  
306 margins (Fig. 1). This fact also reflects the increased attention towards offshore  
307 geohazard assessment in this densely populated region and towards the

308 management of marine areas and marine spatial planning, which is for the moment  
309 limited to the coastal areas (e.g., Schaefer 2009), but will impact in the near future  
310 entire regional seas or marine basins for activities where seafloor geohazards are  
311 relevant, including installations of offshore eolian plants, offshore oil and gas  
312 activities (COM 2010).

313

### 314 **5.1 Methodological aspects of offshore geohazard mapping**

315 The first group of papers is devoted to generic aspects of seafloor mapping for  
316 geohazard purposes. Chiocci and Ridente (this volume) describe the efforts made by  
317 the Italian community in trying to produce regional maps of geohazard features, and  
318 the need to standardize cartographic representation amongst the diverse settings  
319 and parties involved in such effort. Mosher (this volume) highlights the limitations of  
320 multibeam data and the risks of data over-interpretation resulting from system  
321 resolution problems. Along a similar line, Dyer (this volume) discusses a few cases of  
322 inappropriate data processing and how this had an impact on offshore installations.

323

### 324 **5.2 Seafloor morphology and implications for geohazards**

325 The second group of papers provides a few examples of how seafloor mapping  
326 techniques are useful in delineating zones of active geohazards, or excluding certain  
327 seafloor features as the source of potential geohazard. Urgeles et al. (this volume)  
328 perform a detailed analysis of prodeltaic bedforms that had been previously  
329 interpreted as indicative of early seafloor deformation. Cecchini et al. (this volume)  
330 discuss the seaforms on the eastern Sardinia-Corsica continental shelf that could  
331 represent a hazard for pipeline routing. Boudillon et al. (this volume), Ercilla et al.  
332 (this volume), Lo Iacono et al. (this volume) and Morelli et al. (this volume) depict the  
333 complex geomorphology and stratigraphic architecture of various continental  
334 margins and the implications of the structural and sedimentary-erosive features for  
335 geohazard assessment. Finally, Hough et al (this volume) present an integrated and  
336 systematic map-based approach for the assessment and mitigation of seabed  
337 geohazards and risk to proposed deepwater development.

338

### 339 **5.3 Faults and fluids**

340 This section illustrates the ability of multibeam mapping, together with other  
341 geophysical techniques, in delineating active faults and fluid seepage structures.  
342 Larroque et al. (this volume) and Nomikou et al. (this volume) provide two examples  
343 of recent fault activity in the Ligurian margin and Aegean Sea and show how

344 multibeam mapping can help in redefining regional seismic hazard. Dalla Valle et al.  
345 (this volume) provide evidence of interaction between fluid seepage, faulting and  
346 mass-wasting. Léon and Somoza (this volume) present a GIS application to help  
347 identify marine geohazards derived from gas hydrate dissociation using seafloor  
348 mapping data.

349

#### 350 **5.4 Landslide identification, modeling and possible tsunami implications**

351 The last section in this special issue presents a series of papers characterizing  
352 submarine slope failures and slope failure potential and how the morphometric  
353 parameters extracted from seafloor mapping and subsurface geophysical data help  
354 define the post-failure dynamics of these events and their consequent tsunamigenic  
355 potential. Migeon et al (this volume) and Casas et al. (this volume) define various  
356 types of mass failure events and deduce the controlling factors in the development  
357 of the observed failures. Casalbore et al (this volume) show the advantage of  
358 differential bathymetric surveys for depicting seafloor dynamic processes at the  
359 origin of geohazards. Mazzanti et al. (this volume) use constraints from an historical  
360 event and observed seafloor morphology to test landslide propagation and tsunami  
361 generation models. Tinti et al. (this volume) and Argnani et al (this volume) use  
362 morphometric parameters from a large sector collapse on the island of Ischia and  
363 the Eastern Adriatic Basin, respectively, to test tsunami worst-case scenarios for the  
364 adjacent Italian coasts. L'Heureux et al. (this volume) use multibeam bathymetric  
365 data together with additional geophysical and geotechnical data to clarify the  
366 sequence of events around the 1888 landslide and tsunami in the bay of Trondheim.  
367 Finally, Strasser et al. (this volume) present a new concept for evaluating basin-wide  
368 slope stability through time as a potential tool for regional seismic and tsunami  
369 hazard assessment.

370

#### 371 **6. Challenges in offshore geohazard mapping**

372 Multibeam bathymetry probably offers the most cost-effective way to chart the  
373 ocean floor, yet this only portrays a static view of the seafloor. Marine sedimentary  
374 processes often occur over large time spans and/or at very low recurrence rates, and  
375 therefore this static view of the ocean floor depicts a series of potentially hazardous  
376 phenomena that are not active anymore or not representative of present day  
377 processes, particularly in formerly glaciated margins (see also cautionary note in  
378 Mosher this volume). In active geodynamic settings or in areas close to sources of  
379 high sediment supply, the seafloor offers a much more dynamic environment, and  
380 differs greatly from one season to another. This also offers the opportunity to  
381 witness the processes at the origin of geohazards at work with **repetitive surveys**.

382 Programs are just being established to monitor seafloor dynamic systems using  
383 multibeam surveys. For example Duffy and Hughes-Clarke 2005, Smith et al., 2005  
384 and Hughes Clarke et al. 2009 have shown that using repetitive multibeam surveys it  
385 is possible to measure seafloor bedform migration. Knowledge of migration rate  
386 together with bedform height theoretically enable the calculation of sediment  
387 transported within the migrating bedforms or “bedload transport” (Duffy and  
388 Hughes-Clarke 2005). The repetitive surveys have also wide application to  
389 monitoring mass-wasting phenomena, both in sedimentary and volcanoclastic  
390 environments as already shown in Chiocci et al. 2008, Hughes-Clarke et al.2009,  
391 Casalbore et al (this volume), or affecting man-made structures (Dan et al. 2007, Li et  
392 al. 2009). A series of backscatter multibeam surveys has also been used to assess the  
393 impact on benthic ecosystems of hyperpycnal flows and related turbiditic deposition  
394 associated with major flood events (Urgeles et al. 2002).

395 Yet, for monitoring morphological change, the typical scale of spatial change must be  
396 greater than the total survey accuracy, which significantly limits our ability to  
397 monitor geohazard processes in deep sea environments. Due to survey accuracy  
398 issues, the scale of the apparent vertical difference is usually proportional to the  
399 slope of the seafloor (Hughes-Clarke et al. 2009).

400 Repetitive surveys are often planned in an opportunistic manner, constrained by the  
401 availability of previous surveys and the occurrence of catastrophic events such as  
402 floods (e.g. Urgeles et al. 2002, Casalbore et al. this volume) or seafloor failures  
403 (Chiocci et al. 2008). However, monitoring seafloor dynamic processes requires  
404 careful planning and some idea of the rate of seafloor change. Duffy and Hughes-  
405 Clarke 2005 and Hughes Clarke et al 2009 already indicate that in order to track  
406 moving dunes, the surveying period needs to be close enough for the dunes not to  
407 migrate more than half their spacing.

408 The common practice for geohazard assessment in the domain of offshore  
409 exploration includes a reiterative approach where geohazard are identified with  
410 repeated and more and more focussed surveys. This include the use of **AUV-ROV**  
411 **seafloor surveys**, also because of the need to identify potentially hazardous features  
412 at the scale of the planned seafloor infrastructures (Kvalstad 2007). Also academic  
413 research may take advantage from these new technologic developments. Migeon et  
414 al. (2011b) propose a closer look at the seafloor down the area of the Nice 1979  
415 event with the use of an AUV with spatial resolution of 2 m (compared to the 25 m  
416 of the Simrad EM300 echosounder). Apart from the increased number of submarine  
417 landslide scar detected, the main advantage is the possibility to identify clearly the  
418 signature of retrogressive erosion and the traces of gravity flow transformations at  
419 meter scale, with implication for the understanding of the causal sedimentary  
420 processes.

421 The importance of **groundtruthing** remains in any case a key point in most common  
422 practice of geohazard assessment, especially when quantitative data are needed to  
423 support modelling. The search for the crucial site to obtain in situ data is in any case  
424 based on the use of swath bathymetry and seismic profiles: indirect geophysical data  
425 may actually guide the selection of significant measures, for example of sediment  
426 pore pressure to evaluate the possible effects of a recent earthquake (Sultan et al.  
427 2008).

428 The increasing amount of data and the diversity of dataset in the offshore domain  
429 has already propelled the integration of large and complex datasets in Geographic  
430 Information Systems as a current practice to evaluate the areal and size **distribution**  
431 **of geohazards**. Concerning the distribution of submarine landslides along entire  
432 continental margins, for example, a limited number of studies exist at large scale to  
433 evaluate if there is a scaling relationship between submarine landslide area/volume  
434 and frequency of occurrence, as it is the case for earthquakes and their magnitude  
435 (e.g., Chaytor et al 2009). In some cases mapping is supported by other analyses such  
436 as landslide susceptibility linked to seismic activity (Strasser et al. this volume), and  
437 susceptibility maps of submarine landslides and of fluid flow features related to  
438 possible controlling parameters (Leon and Somoza, this volume).

439 Finally, another need related to large seafloor datasets, especially when data have to  
440 be summarized in synthetic views, is the rapid analysis of the data with the  
441 application of automatic **seafloor classifications** (e.g., Atallah and Smith 2004). These  
442 approaches have the advantage of allowing to make some predictions on the state of  
443 the seafloor that can be tested with successive surveys integrating additional  
444 techniques. The literature on this subject is sizeable and rapidly expanding (e.g.,  
445 Hamilton 2005). Some attempts to compare different methods show that the  
446 success of a method in recognizing known patterns depends on the character of the  
447 seafloor morphology and its dominant grain size (Müller et al. 2007). The use of  
448 these techniques may serve as a guide to interpretation and a basis to formulate  
449 hypotheses, but it needs to be applied with caution and to be accompanied by  
450 control from independent data.

451 The increasing number of seafloor mapping studies at local and regional scale will  
452 certainly contribute to capture any inherent pattern in the signature of potentially  
453 catastrophic processes and thus provide enhanced seafloor geohazard assessment.

454

## 455 **7. Conclusion**

456 Seafloor geohazards are one of the increasing concerns worldwide because of the  
457 increasing development of offshore and coastal facilities and growth of marine

458 services' exploitation. Assessment of seafloor geohazards involves several  
459 communities facing different interests and needs.

460 Technological development is offering marine geology powerful seafloor mapping  
461 tools, multibeam bathymetry above all; they are essential for geohazard  
462 characterization and, through repetitive surveys, monitoring and understanding of  
463 ongoing geological processes. Does not only high-resolution seafloor mapping  
464 enhance the assessment of marine geohazard, but it is also boosting the scientific  
465 knowledge on seafloor tectonic, erosional and depositional processes.

466 The study of geohazard-bearing features provides an opportunity for collaborative  
467 activities between academia and industry as well as with public authorities in charge  
468 of land planning or civil defense. However, for the different communities, geohazard  
469 assessment has different objectives, time span and required precision in forecast.  
470 Accordingly, we identify three types of geohazard assessment: sensu-strictu,  
471 engineering and non-specific. Regardless of the different approach, seafloor  
472 mapping remains the unavoidable if not the main step in the geohazard assessment  
473 process.

474 For industry, focused in short time span and small areas, seafloor mapping is the first  
475 step in zoning the area of interest, in order to define the physical environment  
476 where the infrastructure will lay and define geohazards that may recur during the  
477 life-span of the infrastructure. Sea floor mapping is also the base to plan further  
478 investigations and in-situ measurements aimed to identify type, location and  
479 possible timing of hazardous event.

480 For public authorities, seafloor mapping is the main tool to obtain a homogeneous  
481 product over vast offshore areas. Given the impossibility to achieve a precise  
482 (definite) assessment of geohazard for any seafloor feature, the assessment mainly  
483 relies in the identification of geomorphic features linked to hazardous  
484 processes/events that might re-occur in the future over the same area, without any  
485 specific indication of time and precise location. This assessment does not involve a  
486 specific indication of time or precise location of geohazard occurrence.

487 Finally, scientific research, is interested in defining in detail the physical processes  
488 that shape the seafloor with the maximum possible resolution. However seafloor  
489 mapping is also the primary tool to define the amount and type processes occurring  
490 in a given area that are the key to interpret the long-term morphostructural and  
491 stratigraphic evolution of the continental margin and ocean basins.

492

493 **Acknowledgments**

494 We acknowledge the Chief Editors of Marine and Geophysical Researches for their  
495 willing to publish a joint academic-industry contribution on seafloor mapping, and  
496 for their patience in following all the steps towards publication.

497 This volume was partially supported by the UNESCO and IUGS through project IGCP-  
498 511 “Submarine Mass Movements and Their Consequences”, currently superseded  
499 by IGCP project 585, “Earth’s continental MARGins: aSsessing geoHAzard from  
500 submarine Landslides (E-MARSHAL; <http://www.igcp585.org>) and of MAGIC project  
501 “Marine Geohazards along the Italian Coasts” ([www.magicproject.it](http://www.magicproject.it)).

502 Finally, National Civil Protection Department of Italy and namely Prof. B. de  
503 Bernardinis, are acknowledged not only for the aid they gave to the Ischia  
504 conference and to the volume but for supporting the MaGIC Project (Marine  
505 Geohazards along the Italian Coasts, [www.magicproject.it](http://www.magicproject.it)) that is boosting a new  
506 wave of marine geology researches among the Italian marine geology community.

## 507 **References**

508 Argnani A, Tinti S, Zaniboni F, Pagnoni G, Armigliato A, Panetta D, Tonini R (2011) The  
509 eastern slope of the southern Adriatic basin: a case study of submarine landslide  
510 characterization and tsunamigenic potential assessment. Marine Geophysical Researches,  
511 this volume, DOI 10.1007/s11001-010-9100-2

512 Atallah L, Smith PJ (2004) Automatic seabed classification by the analysis of sidescan sonar  
513 and bathymetric imagery. IEE proceedings. Radar, Sonar and Navigation 151: 327-336.

514 Boudillon F., The Bulgheria canyon-fan: a small scale, proximal, fluvial-like system in the  
515 eastern Tyrrhenian Sea. Marine Geophysical Researches, this volume, DOI 10.1007/s11001-011-9138-9

516 Brown CJ, Blondel P (2008) Developments in the application of multibeam sonar backscatter  
517 for seafloor habitat mapping. Appl Acoust, 70-12:1242:1247

518 Camerlenghi A, Urgeles R, Fantoni L (2010) A Database on Submarine Landslides of the  
519 Mediterranean Sea. In Mosher DC, Moscardelli L, Shipp RC, Chaytor JD, Baxter CDP, Lee HJ,  
520 Urgeles R (Eds.) Submarine Mass Movements and Their Consequences, Advances in Natural  
521 and Technological Hazards Research, 28, Springer, Dordrecht (The Netherlands), pp. 491-501.

522 Carpine-Lancre, J, Fisher R, Harper B, Hunter P, Jones M, Kerr A, Laughton A, Ritchie S, Scott  
523 D, Whitmarsh M (2003) The History of GEBCO 1903-2003: the 100-year story of the General  
524 Bathymetric Chart of the Oceans. 140 pp, GITC bv, Netherlands, ISBN 90-806205-4-8.

525 Casalbore D, Chiocci FL, Mugnozza GS, Tommasi P, Sposato A (2011) Flash-flood hyperpycnal  
526 flows generating shallow-water landslides at Fiumara mouths in Western Messina Straits  
527 (Italy) . Marine Geophysical Researches, DOI: 10.1007/s11001-011-9128-y

528 Casas, D., Ercilla, G., Yenes, M. Estrada, F., Alonso, B., García, M., Somoza, L (2011) The  
529 Baraza slide. Model and dynamics, Marine Geophysical Researches, this volume, DOI  
530 10.1007/s11001-011-9132-2

531 Cecchini S, Taliana D, Giacomini L, Herisson C, Bonnemaire B (2011) Preliminary results on  
532 submarine geo-hazards in the eastern Sardinia-Corsica continental margin. Marine  
533 Geophysical Researches, this volume, DOI: 10.1007/s11001-011-9126-0



- 534 Chaytor JD, ten Brink US, Solow AR, Andrews BD (2009) Size distribution of submarine  
535 landslides along the U.S. Atlantic margin. *Marine Geology* 264: 16-27.
- 536 Chiocci F.L., De Alteriis G. (2006). The Ischia debris avalanche: first clear submarine evidence  
537 in the Mediterranean of a volcanic Island pre-historic collapse. *TERRA NOVA*, 18 (3):162-180,
- 538 Chiocci FL, Ridente D (2011) Regional-scale geohazard assesment on Italian continental  
539 margins. The MaGIC project experience (Marine Geohazards along the Italian Coasts. *Marine*  
540 *Geophysical Researches*, this volume, DOI: 10.1007/s11001-011-9120-6
- 541 Chiocci FL, Romagnoli C, Bosman A (2008) Morphologic resilience and depositional processes  
542 due to the rapid evolution of the submerged Sciara del Fuoco (Stromboli Island) after the  
543 December 2002 submarine slide and tsunami, *Geomorphology* 100: 356-365.
- 544 Chiocci F.L., Romagnoli C., Tommasi P., Bosman A. (2008). The Stromboli 2002 tsunamigenic  
545 submarine slide: Characteristics and possible failure mechanisms. *Journal of Geophysical*  
546 *Research*, 113 B10: 151-181
- 547 Chiocci FL, Ridente D, Casalbore D, Bosman A Eds (2009) International Conference on  
548 Seafloor Mapping for Geohazard Assessment, Extended Abstracts. *Rendiconti Online Società*  
549 *Geologica Italiana*, 7: 152 pp. ISSN 2035-8008
- 550 COM (2010) 771 Maritime spatial planning in the EU - Achievements and future  
551 developments. [http://ec.europa.eu/maritimeaffairs/pdf/com\\_2010\\_771\\_en.pdf](http://ec.europa.eu/maritimeaffairs/pdf/com_2010_771_en.pdf) 10pp.
- 552 Dalla Valle G, Gamberi F (2011) Pockmarks and seafloor instability in the Olbia continental  
553 slope (Northeastern Sardinian margin, Tyrrhenian sea) . *Marine Geophysical Researches*, this  
554 volume, DOI 10.1007/s11001-011-9133-1
- 555 Dan G, Sultan N, Savoye B (2007) The 1979 Nice harbour catastrophe revisited: Trigger  
556 mechanism inferred from geotechnical measurements and numerical modelling. *Marine*  
557 *Geology* 245: 40-64.
- 558 De Alteriis G., D. Insinga, S. Morabito, V. Morra, Chiocci F.L., F. Terrasi, C. Lubritto, C. Di  
559 Benedetto, M. Pazzanese (2010). Age of submarine debris avalanches and  
560 tephrostratigraphy offshore Ischia Island, Tyrrhenian sea, Italy. *Marine Geology*, 278 (1-4): 1-  
561 18
- 562 de Moustier C (1988) State of the Art in Swath Bathymetry Survey Systems, *International*  
563 *Hydrographic Review* 65: 25-54.
- 564 de Moustier C (2010) Seafloor Acoustic Backscatter Measurements With Multibeam  
565 Echosounders, *Proceedings Offshore Technology Conference*, Paper Number 20866-MS.  
566 DOI: 10.4043/20866-MS
- 567 Dolan MFJ, Buhl-Mortensen P, Thorsens T, Buhl Mortensen L, Bellec VK, Boe R (2009)  
568 Developing seabed nature-type maps offshore Norway: initial results from MAREANO  
569 programme. *Norwegian Journal of Geology* 89: 17-28
- 570 Duffy GP, Hughes-Clarke JE (2005) Application of Spatial Cross-Correlation to Detection of  
571 Migration of Submarine Sand Dunes. *Journal of Geophysical Research - Earth Surface*  
572 110(F4), DOI: 10.1029/2004JF000192.
- 573 Dyer JM (2011) Geohazard identification: the gap between the possible and reality in  
574 geophysical surveys for the engineering industry. *Marine Geophysical Researches*, this  
575 volume, DOI 10.1007/s11001-011-9137-x
- 576 Ercilla G, Casas D, Vázquez JT, Iglesias J, Somoza L, Juan C, Medialdea T, León R, Estrada F,  
577 García-Gil S, Bohoyo F, García M, Maestro A, Farran, M (2011) Imaging the recent sediment

578 dynamics of the Galicia Bank region (Atlantic, NW Iberian Peninsula). *Marine Geophysical*  
579 *Researches*, this volume, DOI 10.1007/s11001-011-9129-x

580 Farr HK (1980) Multibeam bathymetric sonar: Sea beam and hydro chart. *Marine Geodesy*  
581 *4/2*, 77-93.

582 Hall JK (2006) GEBCO Centennial Special Issue – Charting the secret world of the ocean floor:  
583 the GEBCO project 1903–2003. *Marine Geophysical Researches* 27: 1-5.

584 Hamilton LJ (2005) A Bibliography of Acoustic Seabed Classification. DSTO, Defence Science  
585 and Technology Organisation, Australia, Technical report 27: 11 pp.

586 Heezen BC, Tharp M (1954) Physiographic diagram of the western North Atlantic. *GSA*  
587 *Bulletin* 65:1261.

588 Heezen BC, Ewing M, Tharp M (1959) The Floors of the Oceans: Part I. The North Atlantic.  
589 *GSA Special Paper* 65.

590 Hough G, Green J, Fush P, Mills A, Moore R (2011) A Geomorphological Mapping Approach  
591 for the Assessment of Seabed Geohazards and Risk. *Marine Geophysical Researches*, this  
592 volume, DOI: 10.1007/s11001-010-9111-z

593 Hughes Clarke JE, Mayer LA, Wells DE (1996) Shallow-water imaging multibeam sonars : A  
594 new tol for investigating seafloor processes in the coastal zone and on the continental shelf :  
595 *Marine Geophysical Researches*, 18: 607-629.

596 Hughes Clarke JE, Brucker S, Hill P, Conway K (2009) Monitoring evolution of fjord deltas n  
597 temperate and Arctic regions, *Rendiconti Online Società Geologica Italiana* 7: 145-150.

598 Kvalstad TJ (2007) What is the Current "Best Practice" in Offshore Geohazard Investigations?  
599 A State-of-the-Art Review. *Offshore Technology Conference OTC 18545*, 14pp.

600 Larroque C, Mercier de Lépinay B, Migeon S (2011) Morphotectonic and faults-earthquakes  
601 relationship along the northern Ligurian margin (Western Mediterranean) based on high  
602 resolution multibeam bathymetry and multichannel seismic-reflection profiles. *Marine*  
603 *Geophysical Researches*, this volume, DOI: 10.1007/s11001-010-9108-7

604 Léon R, Somoza L (2011) GIS-based mapping for marine geohazards derived from gas  
605 hydrate dissociation. *Marine Geophysical Researches*, this volume, DOI 10.1007/s11001-011-9135-z

606 L'Heureux J-S, Glimsal S, Longva O, Hansen L, Harbitz CB (2011) The 1888 shoreline landslide  
607 and tsunami in Trondheimsfjorden, central Norway. *Marine Geophysical Researches*, this  
608 volume, DOI: 10.1007/s11001-010-9103-z

609 Li MZ, Parrott DR, Yang Z (2009) Sediment Stability and Dispersion at the Black Point  
610 Offshore Disposal Site, Saint John Harbour, New Brunswick, Canada, *Journal of Coastal*  
611 *Research* 25:1025-1040

612 Lo Iacono C, Sulli A, Agate M, Lo Presti V, Pepe F (2011) Submarine Canyons and slope  
613 failures in the Palermo Gulf (Southern Tyrrhenian Sea): implications for geo-hazard  
614 assessment. *Marine Geophysical Researches*, this volume, DOI: 10.1007/s11001-011-9118-0

615 Longva, O., Janbu, N., Blikra, L.H., Boe, R., 2003. The 1996 Finneidfjord Slide, seafloor failure  
616 and slide dynamics. In: Locat, J., Mienert, J. (Eds.), *Submarine Mass Movements and Their*  
617 *Consequences*. Kluwer Acad. Publ., Dordrecht, The Netherlands, 531– 538.

618 Lurton X (2002) *An Introduction to Underwater Acoustics Principles and applications*.  
619 London, Springer, 347 pp.

620 Malinverno, A., Ryan, W.B.F., Auffret, G., Pautot, G., 1988. Sonar images of the path of  
621 recent failure events on the continental margin off Nice, France. *Spec. Pap.-Geological*  
622 *Society of America* 229, 59–75

623 Mazzanti P, Bozzano F (2011) Revisiting the 6th February 1783 Scilla (Calabria, Italy) landslide  
624 and tsunami by numerical simulation. *Marine Geophysical Researches*, this volume, DOI:  
625 10.1007/s11001-011-9117-1

626 MediMap Group, Loubrieu B, Mascle J et al. (2005) Morpho-bathymetry of the  
627 Mediterranean Sea. CIESM/ Ifremer special publication, *Atlases and Maps*, two maps at  
628 1:2,000,000 scale.

629 Migeon S, Cattaneo A, Hassoun V, Larroque C, Corradi N, Fanucci F, Dano A, Mercier de  
630 Lepinay B, Sage F, Gorini C (2011a) Morphology, distribution and origin of recent submarine  
631 landslides of the Ligurian Margin (North-western Mediterranean): some insights into  
632 geohazard assessment. *Marine Geophysical Researches*, this volume, DOI:10.1007/s11001-  
633 011-9123-3

634 Migeon S, Cattaneo A, Hassoun V, Dano A, Ruellan E (2011b) Failure processes and gravity-  
635 flow transformation revealed by high-resolution AUV swath bathymetry on the Nice  
636 continental slope (Ligurian Sea). In: Yamada Y, Kawamura K., Ikehara K., Ogawa Y., Urgeles  
637 R., Mosher D., Chaytor J, Strasser M (eds) *Submarine Mass Movements and Their*  
638 *Consequences*. in press.

639 Morelli D, Cuppari A, Colizza E, Fanucci F (2011) Geomorphic setting and geohazard-related  
640 features along the Ionian Calabrian margin between Capo Spartivento and Capo Rizzuto  
641 (Italy). *Marine Geophysical Researches*, this volume, DOI 10.1007/s11001-011-9130-4

642 Morgan JK, Silver E, Camerlenghi A, Dugan B, Kirby S, Shipp C, Suyehiro K (2009) Addressing  
643 Geohazards Through Ocean Drilling. *Scientific Drilling* 7: 15-30.

644 Mosher DC (2011) Geohazard mapping with multibeam sonar: resolution and the need for 3-  
645 D. *Marine Geophysical Researches*, this volume, DOI: 10.1007/s11001-010-9104-y

646 Müller RD, Eagles S, Hogarth P, Hughes M (2007) Automated textural image analysis of  
647 seabed backscatter mosaics: A comparison of four methods. In Todd BJ and Greene HG (eds)  
648 *Mapping the Seafloor for Habitat Characterization*. Geological Association of Canada, Special  
649 Paper 47: 43-61.

650

651 NGI (2005) Offshore geohazards. Summary Report 20021023-2. 72 pp.

652 Muñoz A, Acosta J, Palomo C, Pardo de Domlebum M (1998) Hydrographic and  
653 Oceanographic Programme for the Spanish EEZ. *EEZ Technology* 2: 71-75.

654 Nomikou P, Papanikolaou D (2011) The Active Mandraki Fault Zone of Nisyros Volcano based  
655 on onshore and offshore data. *Marine Geophysical Researches*, DOI: 10.1007/s11001-011-  
656 9119-z

657 OPC, 2007, California Ocean Protection Council,  
658 [http://www.opc.ca.gov/webmaster/ftp/pdf/agenda\\_items/20071025/07\\_seafloor](http://www.opc.ca.gov/webmaster/ftp/pdf/agenda_items/20071025/07_seafloor)  
659 [mapping/1007COPC07\\_seafloor mapping.pdf](http://www.opc.ca.gov/webmaster/ftp/pdf/agenda_items/20071025/07_seafloor_mapping/1007COPC07_seafloor_mapping.pdf)

660 Piper, D.J.W., Cochonat, P., Morrison, M.L., 1999. The sequence of events around the  
661 epicenter of the 1929 Grand Banks earthquake: initiation of debris flows and turbidity  
662 current inferred from sidescan sonar. *Sedimentology* 46: 79–97.

- 663 Posamentier H.W., R. J. Davies, J. A. Cartwright and L. Wood (2007) Seismic geomorphology  
664 - an overview, Geological Society, London, Special Publications, 277: 1-14
- 665 Rusby JSM, Revie J, (1975) Long-range sonar mapping of the continental shelf, Marine  
666 geology 19: 41-49
- 667 Schaefer N (2009) The Integrated EU Maritime Policy and Maritime Spatial Planning - The  
668 Way Ahead. EMSAGG Conference, 7-8 May 2009. 6 pp.
- 669 Smith WH, Sandwell DT (1997) Global Sea Floor Topography from Satellite Altimetry and  
670 Ship Depth Soundings. Science Magazine 277/5334,
- 671 Smith DP, Ruiz , Kvitek R, Iampietro PJ (2005) Semiannual patterns of erosion and deposition  
672 in upper Monterey Canyon from serial multibeam bathymetry, GSA Bulletin 117: 1123-1133.
- 673 Strasser M, Hilbe M, Anselmetti FS (2011) Mapping basin-wide subaquatic slope failure  
674 susceptibility as a tool to assess regional seismic and tsunami hazards. Marine Geophysical  
675 Researches, this volume, DOI: 10.1007/s11001-010-9100-2
- 676 Sultan N, Cattaneo A Sibuet J-C, Schneider J-L, the Sumatra Aftershocks team (2008) Deep  
677 sea in situ excess pore pressure and sediment deformation off NW Sumatra and its relation  
678 with the December 26, 2004 Great Sumatra-Andaman Earthquake. Int J Earth Sci (Geol  
679 Rundsch) DOI 10.1007/s00531-008-0334-z
- 680 Sultan N, Savoye B, Jouet G, Leynaud D, Cochonat P, Henry P, Stegmann S, Kopf A (2010)  
681 Investigation of a possible submarine landslide at the Var delta front (Nice continental slope,  
682 southeast France). Canadian Geotechnical Journal 47 , 486-496.
- 683 Thorsnes T. (2009) MAREANO- An introduction, Norwegian Journal of Geology, 89, 3
- 684 Tinti S, Chiocci FL, Zaniboni F, Pagnoni G, de Alteriis G (2011) Numerical simulation of the  
685 tsunami generated by a past catastrophic landslide on the volcanic island of Ischia, Italy.  
686 Marine Geophysical Researches, this volume, DOI: 10.1007/s11001-010-9109-6
- 687 Urgeles R, Cattaneo A, Liqueste C, De Mol B, Amblàs D, Sultan N, Trincardi F (2011) A review  
688 of undulated sediment features on Mediterranean prodeltas: distinguishing sediment  
689 transport structures from sediment deformation. Marine Geophysical Researches, this  
690 volume, DOI: 10.1007/s11001-011-9125-1
- 691 Urgeles R, Locat J, Schmitt T, Hughes-Clarke J (2002) Spatial and temporal backscatter  
692 variations in the Saguenay Fjord: causes and relation with the July, 1996 Saguenay flood,  
693 Marine Geology 184: 41-60.

694

## 695 **Figure Captions**

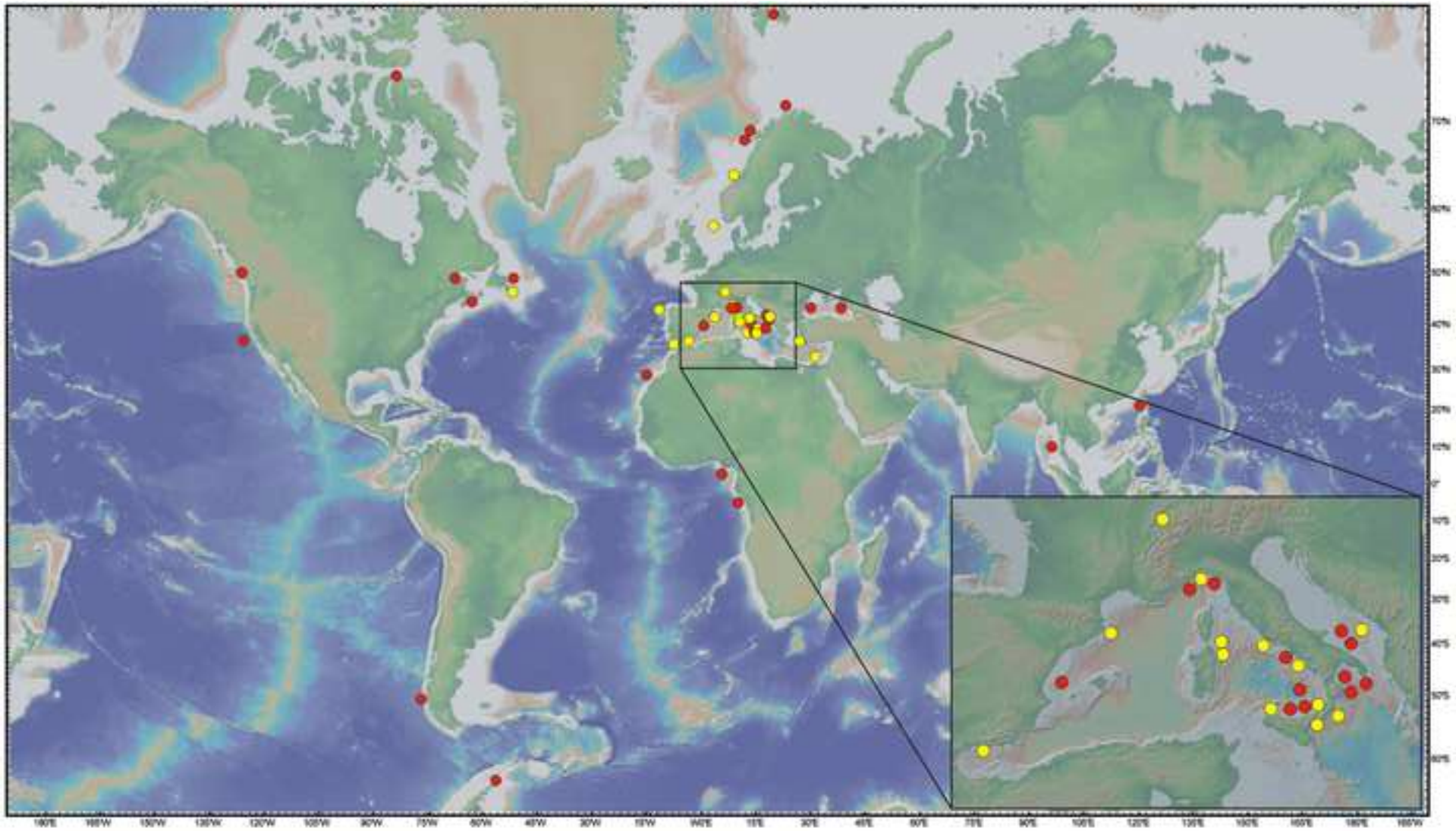
696 Figure 1. Location of the study areas described in the articles of this volume (yellow  
697 dots). Further case histories presented at the Ischia Conference in May 2009 but not  
698 included in this volume are reported as red dots. Their extended abstracts are  
699 collected in Chiocci et al. eds. (2009).

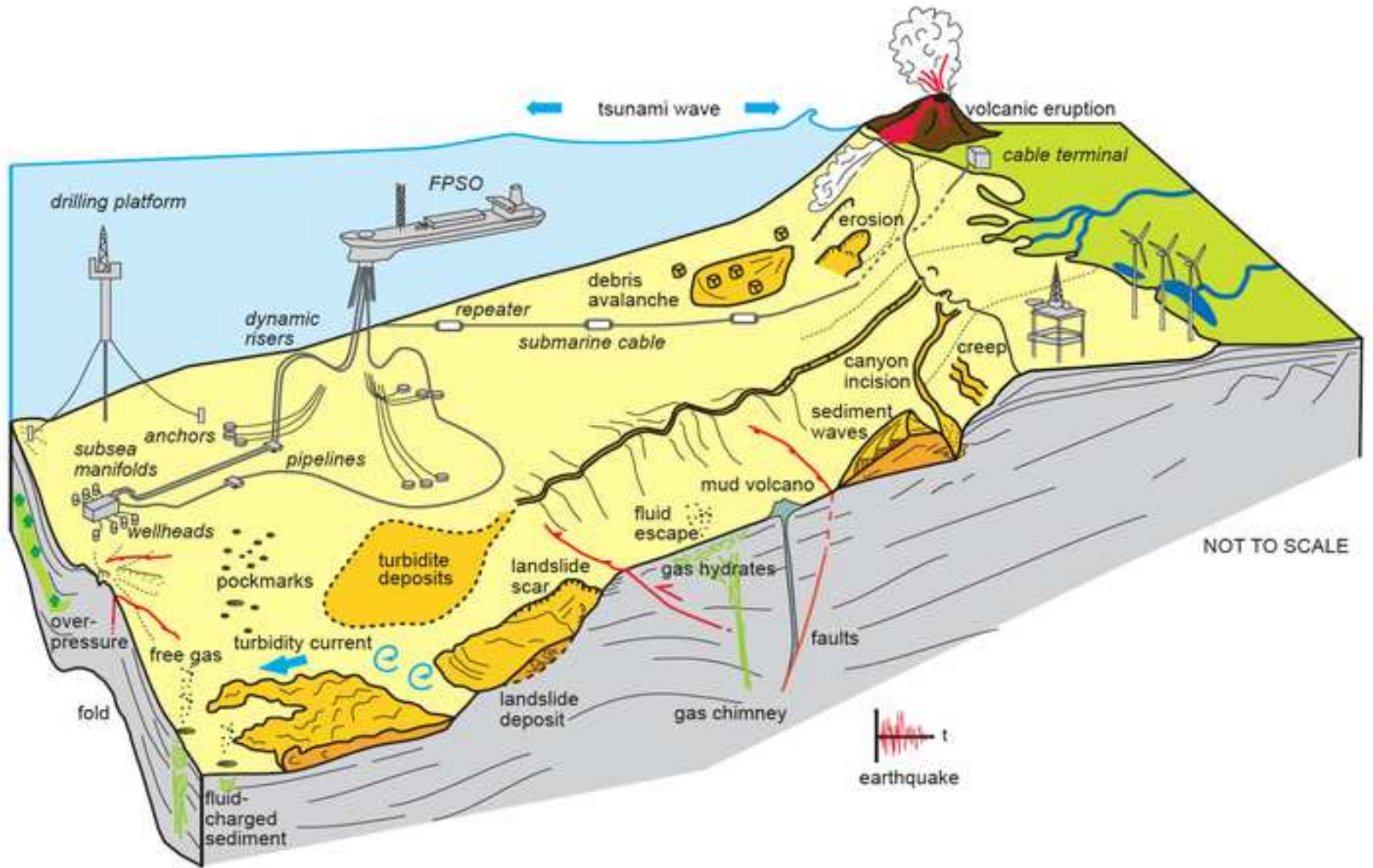
700 Figure 2. Cartoon summarizing the seafloor features linked to potentially hazardous  
701 geological processes. This figure depicts an idealized continental margin with both  
702 natural geohazard-bearing features and main anthropogenic structures lying on the  
703 seafloor.

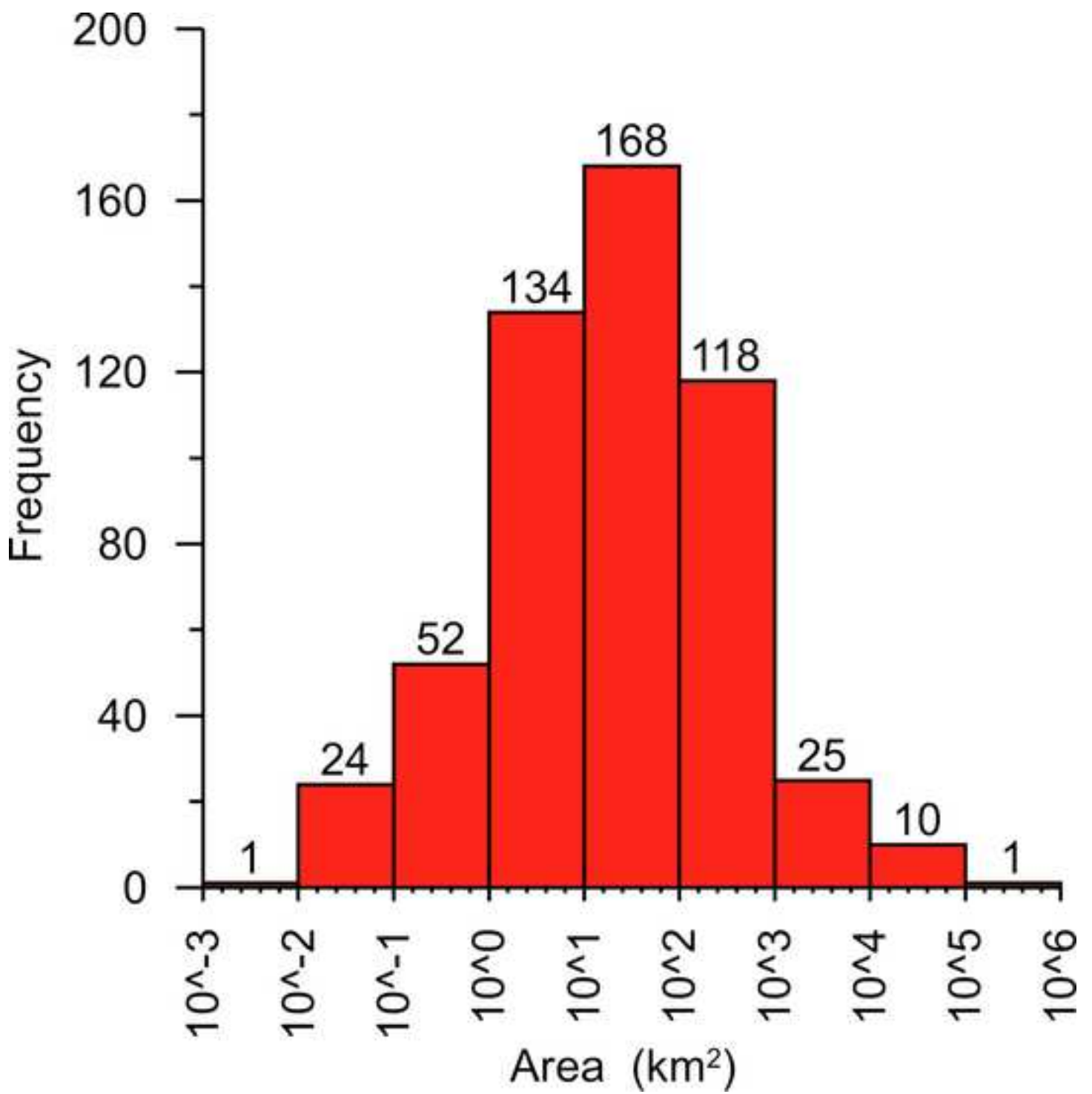
704 Figure 3. Frequency of magnitude (log-area) for submarine landslides from the  
705 Mediterranean landslide database (see Camerlenghi et al. 2010) in the  
706 Mediterranean basin showing incomplete view of distribution (i.e. smaller landslides  
707 are not picked in the distribution).

708 Figure 4. Triangular diagram summarizing the different type of geohazard  
709 assessment and how the different communities relate to them. See text for details.

710









*Geohazard  
assessment  
sensu-strictu*

*industry*

*academia*

*Engineering  
geohazard  
assessment*

**public agencies**

*Non-specific  
geohazard  
assessment*

