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12
13 **Summary**

14
15 Fire affects large parts of the dry Mediterranean shrubland, resulting in erosion and
16 losses of plant nutrients. We have attempted to measure these effects experimentally on
17 a calcareous hillside representative of such shrubland. Experimental fires were made on
18 plots (4 m x 20 m) in which the fuel was controlled to obtain two different fire
19 intensities giving mean of soil surface temperature of 439°C and 232°C with
20 temperatures exceeding 100°C lasting for 36 minutes and 17 minutes. The immediate
21 and subsequent changes induced by fire on the soil's organic matter content and other
22 soil chemical properties were evaluated, together with the impact of water erosion.

23 Seven erosive rain events, which occurred after the experimental fires (from
24 August 1995 to December 1996) were selected, and on them runoff and sediment

25 produced from each plot were measured. The sediments collected were weighed and
26 analysed. Taking into account the variations induced by fire on the soil properties and
27 their losses by water erosion, estimates of the net inputs and outputs of the soil system
28 were made. Results show that the greatest losses of both soil and nutrients took place in
29 the four months immediately after the fire. Plots affected by the most intense fire
30 showed greater losses of soil (4077 kg ha^{-1}) than those with moderate fire intensity
31 (3280 kg ha^{-1}). The unburned plots produced the least sediment (72.8 kg ha^{-1}). Organic
32 matter and nutrient losses by water erosion were related to the degree of fire intensity.
33 However, the largest losses of N-NH_4^+ and N-NO_3^- by water erosion correspond to the
34 moderate fire (8.1 and 7.5 mg N m^{-2} respectively).

35

36

37 **Introduction**

38

39 During the last two decades, environmental scientists have come to realise that fire is a
40 dominant control on the forest landscapes in Southern Europe, especially in the
41 Mediterranean basin where it is been considered a major environmental problem.

42 The temperatures of the soil and the burned vegetation produce changes in soil
43 chemical and physical properties. The magnitude of these changes greatly depends on
44 fire and soil characteristics, vegetation cover and the weather (Blank & Zamudio, 1998).
45 Fires not only cause changes in soil and environmental properties, they could also
46 increase net losses of nutrients by volatilization, leaching and water erosion. These
47 losses depend on the intensity of fires (peak temperatures reached on the soil surface
48 and their duration). In fires of medium and high intensity, the disappearance of the
49 vegetal cover and the partial combustion of the soil organic matter could alter soil

50 structure, thereby affecting the porosity and other hydrological characteristics (Andreu
51 *et al.*, 1994; Giovannini & Lucchesi, 1997), and enhance the accumulation of
52 hydrophobic substances at several centimetres depth and thereby reduce infiltration and
53 increase runoff (De Bano, 1971; Imeson *et al.*, 1992). These conditions enhance the
54 susceptibility of soil to erosion and they generally cause an increase in runoff and soil
55 loss. Nutrients are then lost as they are transported with the solids by runoff (Andreu *et*
56 *al.*, 1996; Rubio *et al.*, 1996). Andreu *et al.* (1996) showed after a so called ‘intense
57 fire’, erosion was more intense than for a ‘light or moderate fire’, producing significant
58 losses in chemical elements removed with the sediment or dissolved in runoff. Thus, the
59 effect of fire on soil erosion depends mainly on its intensity and on the intrinsic
60 characteristics of any subsequent rain event (intensity, duration, etc.) (Rubio *et al.*,
61 1996).

62 This problem becomes important in the Mediterranean basin where the impact of
63 continuous and repeated fires, mainly in summer, followed by torrential rains in autumn
64 results in intense erosion (Giovannini *et al.*, 1990; Andreu *et al.*, 1996).

65 We have studied the immediate changes in soil organic matter content, total and
66 mineral soil nitrogen, available phosphorus and exchangeable cations induced by fire.
67 The temporal variation of erosion rates on burned plots, with varied intensities of fire
68 under similar environmental conditions was also studied. Sediments generated by runoff
69 in the erosive rain events are analysed to determine the soil organic matter and nutrient
70 losses by water erosion. Taking into account the variations of these soil losses induced
71 by fire, an analysis of net inputs and outputs of organic matter and nutrients is made.

72

73

74 **Materials and methods**

75

76 *Study site*

77

78 The study area of ‘La Concordia’ is in the municipality of Liria (Valencia, Spain), 50
79 km NW of Valencia city. It is 575 m above the sea level, on land ceded by the Forestry
80 Services of the Valencian Government (Generalitat Valenciana). The experimental fires
81 were made under field conditions on a forested hillside facing South South East, with a
82 sclerophyllous shrub cover regenerated after a previous wildfire occurred in 1978. The
83 dominant vegetation type belongs to the *Rhamno lycioidis-Quercetum cocciferae*
84 association, which is typical of semi-arid Mediterranean areas. The most abundant
85 species include *Rosmarinus officinalis*, *Ulex parviflorus*, *Quercus coccifera*, *Rhamnus*
86 *lycioides*, *Stipa tenacissima*, *Globularia alypum*, *Cistus clusii* and *Thymus vulgaris*.

87 Climatically the area belongs to the dry ombroclimate of the lower
88 mesomediterranean belt, according to Thornthwaite’s classification. The average annual
89 precipitation is around 400 mm with two maximums, autumn and spring, and a dry
90 period from June to September. Mean monthly temperatures range from 13.3°C in
91 January to 25.8°C in August.

92 The soil is a Rendzic Leptosol (FAO-UNESCO, 1988) developed on Jurassic
93 limestone. This soil has a variable depth, always less than 40 cm, many stones (\cong 40%),
94 good drainage, a sandy-loam texture, and an alkaline pH (7.4).

95

96 *Experimental plots and fire treatments*

97

98 The experimental set-up consists of nine plots, each 20 m long by 4 m wide, with
99 similar morphology, slope gradient, rock outcrops, soil and vegetation cover. The
100 location of each plot was made after intensive survey of the vegetation, soil and
101 morphology patterns, based on across-slope transects every 2 m.

102 The plots were oriented parallel to the slope and bounded by bricks. At the foot
103 of each plot a 2-m wide collector ran into a 1500-l tank to collect all the runoff and
104 sediment produced during each rain event. Inside this tank a 30-l tank facilitates the
105 collection of sediments produced.

106 A design of two different fire intensity treatments, with three plots each, was
107 used. Contrasting amounts of fuel, obtained from the surrounding shrub, were added to
108 obtain two different fire intensities. One consists in the addition of 20 t biomass ha⁻¹ (F1
109 treatment). The second consisted in the addition of 40 t biomass ha⁻¹ (F2 treatment). The
110 fuel was spread uniformly on the plots. The remaining three plots were used as control
111 (F0 treatment). The assignation of fire treatment to each plot was made by lot. We
112 assigned the treatments to the plots completely at random without blocking.

113 The composition and the spatial distribution of the vegetation in each plot were
114 determined by intensive field survey, counting the individuals identified to species and
115 measuring their size (height, maximum and minimum diameter) as well as the
116 percentage of soil covered by plants on a 1m x 1m basis. This information was used to
117 map dry biomass and vegetation cover (visual estimation) present in each 1m² and to
118 calculate the mean dry biomass present in the plots. The biomass was estimated by
119 using a non-destructive method similar to the proposed by Etienne & Legrand (1994).
120 Moreover, under the two species that cover large percentage of soil surface (*Ulex*
121 *parviflorus* and *Rosmarinus officinalis*), eight litter samples on a 25 cm x 25 cm were

122 collected, and the biomass was also directly estimated. The biomass present in the plots
123 range from 5 t ha⁻¹ to 8 t ha⁻¹. Approximately the 50% of the plots surface have a
124 quantity of natural biomass ≤ to 5 t ha⁻¹. The most abundant species and their mean
125 frequencies are: *Rosmarinus officinalis* 34%, *Ulex parviflorus* 21%, *Globularia alypum*
126 27%, *Rhamnus lycioides* 4%, *Cistus clusii* 3%, *Quercus coccifera* 3.5% and *Thymus*
127 *vulgaris* 6%.

128 To measure the temperatures on the soil surface and their duration,
129 thermosensitive paints and thermocouples were used. Six thermocouples (type K
130 Inconel 600-insulated) per plot were installed at ground level along parallel lines
131 running downslope and separated from one another by 3 m. From these measurements
132 direct estimates were made of the time that temperature exceeded 100°C. We chose this
133 value because it seemed to be that it is which the most significant changes begin,
134 starting with the water evaporation.

135 To obtain the spatial distribution of temperatures on the soil surface, a set of
136 twenty-four thermosensitive paints were used (Omega Stik Crayons) ranging between
137 100°C and 677°C, in increments of approximately 25°C. They were applied on iron rods
138 each covered with another identical rod, but not painted, to protect them from ashes and
139 flames. The system was tied with two pieces of wire. Just before the experimental fire
140 one iron rod per square metre was placed (a total of 80 iron rods per plot) with the
141 painted side in contact with soil. Immediately after the passage of fire the iron rods were
142 collected and they were read.

143 We lit the fires on 20 and 21 June 1995. Weather and the characteristics of the
144 different rainfall events were monitored by a logging system of sensors placed close to
145 the plots. The weather at the time of burning is recorded in Table 1.

146 When the thermocouples and the thermosensitive paints were in place and the
147 extra biomass had been spread on the plots, a small amount of fuel oil was applied at the
148 bottom of the plots and fire started. The fires progressed upslope, and their behaviours
149 were uniform in all the plots, except in plot 6 that suffered repeated changes in wind
150 direction. In all cases, the fires progressed from their starts to the middle parts of the
151 plots, faster in the centre of fire front than on their flanks. Once the fires had passed the
152 half-way marks they progressed uniformly thereafter.

153 Immediately before the fires we took 36 soil samples (four per plot) 0 to 5 cm
154 depth. Due to the distribution of vegetation in a patchy shrub mosaic we chose two
155 different soil microenvironments for sampling: underneath plant canopies and in the
156 centres of openings. Openings with a minimum diameter greater than 0.5 m were
157 sampled and openings located less than 1 m from shrubs were excluded. Litter was
158 removed prior the sampling. The sampling points were marked with pins so that we
159 could sample again at the same microsites after burning. Two hours after the fires we
160 took four samples of ash per plot to be analysed and another 36 soil samples.

161

162 *Soil and sediments analysis*

163

164 The soil samples were air-dried, sieved to remove material with diameter >2 mm, and
165 stored in air-tight plastic boxes until analysis. Organic matter content was determined
166 by oxidation with potassium dichromate (Jackson, 1958). Total nitrogen was
167 determined by micro-Kjeldahl automatic analyser using the Bremner method (Black *et*
168 *al.*, 1965). Ammonium and nitrate nitrogen were extracted with 2 M KCl solution and
169 determined by steam distillation by micro-Kjeldahl automatic analyser using the

170 Bremner method (Black *et al.*, 1965). Available phosphorus was determined
171 colorimetrically according to method of Olsen and Dean (Black *et al.*, 1965). The
172 exchangeable cations were extracted with 1 M ammonium acetate solution (Peech,
173 1954), and their contents were determined by flame atomic absorption
174 spectrophotometry. These results are expressed in percentage of exchangeable cations
175 with respect to the cation exchange capacity (CEC), to allow comparisons between
176 different sampling periods. The CEC was determined according to the method of Bower
177 *et al.* (1952).

178 The sediments and runoff generated from each rain event were quantified for all
179 plots. The same methods described for the soil samples were applied to the collected
180 sediments.

181

182 *Statistics*

183

184 Soil data were analysed by two-way analysis of variance (ANOVA) with sampling
185 period (before or after fires) and fire treatment as the main effects. When main effects
186 were significant ($p < 0.05$) we used Tukey's test to compare the treatments.

187 Soil losses and runoff production data were analysed by ANOVA with repeated
188 measures. These variables were measured after each rainfall event on the same
189 experimental units. This design provides a control on the differences among units. Fire
190 treatment was the grouping factor (between- subject effect) and the rainfall events were
191 the repeated measures for each case, which vary within the grouping factor. When
192 significant differences were detected among means, the minimum significant difference
193 for the individual effects and their interaction was calculated using Tukey's test.

194

195

196 **Results and discussion**

197

198 *Changes in soil organic matter and soil nutrients after fire*

199

200 Progression of the fire on the plots was, in general, fairly uniform. The mean duration in
201 the soil of temperatures greater than 100°C was 17.6 minutes for F1 treatment, and 36.3
202 minutes for F2 treatment. Assuming that the temperature measured with one iron rod
203 corresponds to 1 m² of the plot surface, we found that on the F1 plots 50% of their
204 surface had temperatures between 200°C and 400°C (mean value 232°C), whereas on
205 the F2 plots more than 50% of the surface had temperatures between 400°C and 600°C
206 (mean value 439°C). Taking into account these results, we have considered that F2 plots
207 suffered a high fire intensity and the F1 plots suffered a fire of moderate intensity.

208 Table 2 records the values of the soil variables before the fire. Immediately after
209 the fires significant variations were observed for the variables studied, which can be
210 related to fire intensity (Table 2).

211 Fire caused a sharp increase of soil ammonium-N on both fire intensities, but it
212 was more accentuated for the high intensity (25 times greater than before burning). This
213 result accords closely with those reported by other authors (Díaz-Fierros *et al.*, 1990;
214 Giovannini & Lucchesi, 1993). The increase in ammonium-N is attributed to the
215 transformation of organic matter, which occurs at temperatures greater than 210°C. Soil
216 nitrate-N decreased after burning (Table 2). Díaz-Fierros *et al.* (1990) and Wienhold &

217 Klemmedson (1992) similarly found a decrease in nitrate contents between 0.5 and 0.3
218 of the original quantity.

219 Available phosphorus concentrations increased markedly as a result of the fire,
220 especially under the most intense one (Table 2). The increase of 55 mg kg⁻¹ of available
221 P for the F2 treatment and 32 mg kg⁻¹ for F1 was induced mainly by combustion of soil
222 organic matter and mineralization as consequences of the high temperatures.

223 The exchange complex in the calcareous soil of La Concordia is dominated by
224 Ca²⁺ and Mg²⁺. The exchangeable cations Na⁺, K⁺ and Mg²⁺ increased as a result of the
225 fires (Table 2), presumably because of their presence in the ash. However, the
226 exchangeable Ca²⁺ decreased significantly after the fire for both intensities (Table 2),
227 approximately in the same proportion (19% and 17%) as the reduction in the CEC.

228 We did not expect this result, because the temperature at which inorganic
229 calcium volatilises is more than 1400°C (Weast, 1980), although the temperatures
230 required for volatilize cations bound in organic compounds may be significantly less
231 (Raison *et al.*, 1985). Thus, it is possible that a fraction of calcium bonded to soil
232 organic matter had been volatilized where soil surface temperature reached values
233 greater than 677°C, whereas other calcium fractions, in a chemically altered form,
234 remained in the soil.

235

236

237 *Water erosion effects*

238

239 A total of 31 erosive rain events producing runoff were recorded for the complete study
240 period, from August of 1995 to December of 1996 (Figure 1). Seven of them, the ones

241 labelled in Figure 1, produced the greatest runoff and sediment transport, and we
242 selected them for this study. The first five events occurred consecutively after the
243 experimental fires to the end of 1995 (Figure 1). The other two occurred in 1996 and
244 produced the greatest losses of soil. The characteristics of these events are listed on
245 Table 3. The maximum volume of precipitation registered in 30 minutes (I_{30}) of these
246 events ranged from 10.92 mm h⁻¹ to 35.36 mm h⁻¹. Only four other rain events reached
247 an I_{30} between 10 and 15 mm h⁻¹ during the study period, but none of these produced
248 enough sediment to be analysed.

249 Table 4 records the effect of the fire intensity on runoff and sediment
250 production. The unburned plots generated the least runoff and lost the least soil. As an
251 estimation of erosion impact, mean runoff and sediment yield on this plots,
252 corresponding to the seven rainfall events were 2.1 L m⁻² and 0.007 kg m⁻²,
253 respectively. In these plots, runoff generation was on average of 86% less than for the
254 most intense fire and 82% less than for moderate fire. Likewise, there was 98% less
255 sediment lost from the control plots after burning. This effect was accentuated after
256 autumn rain, as observed by other authors working in Mediterranean areas (Giovannini
257 & Lucchesi, 1993; Andreu *et al.*, 1996).

258 Figure 2 shows the mean values of runoff yield and sediment yield for the
259 different fire treatments by each rainfall event. From the second rain event of 18
260 September 1995 to the last event recorded (6 December 1996) there are significant
261 differences in runoff generated and sediment transported between the burned and
262 unburned plots. In the plots subjected to the moderate fire, mean runoff yield was 12.1
263 L m⁻² (6 times greater than for the control plots), and the mean sediment yield was 0.33
264 kg m⁻² (46 times greater than for the control plots). The plots that suffered the high fire

265 intensity produced mean runoff and sediment yield of 15.4 L m⁻² and 0.41 kg m⁻²,
266 respectively. These values are about 1.2 times greater than those measured for the
267 moderate fire intensity plots. This shows the significant differences between plots
268 affected by fire and those unaffected in their response to water erosion processes.

269 Generally, the more intense fire resulted in the larger runoff and the greater loss
270 of soil than the moderate fire (Table 4). The latter generated on average, 21.60% less
271 runoff and 19.55% less sediment than the former. However, an inverse trend happened
272 during the three rainfall events immediately after the fire. The cover given by the
273 accumulation of ashes and partly burnt plant material, the decrease in soil porosity and
274 the possible formation of a subsurficial hydrophobic layer in the soil could reduce
275 infiltration, and favour runoff generation on the moderate fire intensity plots, as
276 observed previously by DeBano (1971) and Mallik *et al.* (1984).

277 Application of ANOVA with repeated measures on the runoff and sediment
278 yield data shows that there are statistically significant differences in these variables for
279 fire treatment and for the interaction 'rain x fire' (Tables 5 and 6).

280

281 *Organic matter and nutrient losses by water erosion after fire*

282

283 The total content of organic matter in the sediments for the seven rainfall events was
284 9.72 kg from treatment F2, 8.78 kg for F1, and only 0.072 kg for the control plots.
285 These amounts represent an organic matter loss of 39.91, 36.65 and 0.54 g m⁻² for the
286 three treatments (Table 7). The overall loss of nitrogen was greater for treatment F2
287 than for F1 (Table 7), representing a loss per unit area of 1.59 and 1.46 g N m⁻²,
288 respectively. In the control plots, the total N content in the sediments was significantly

289 less at 0.023 g m^{-2} . The greatest losses of organic matter and total nitrogen occurred
290 during the second rainfall event on 18 September (Figure 3).

291 The losses of both forms of mineral N in the eroded sediments of F1 plots are
292 greater than for F2 (Table 7). This could be due to the differences in N volatilization
293 during combustion. Mean soil temperature of the moderate fires was 240°C , which is
294 enough to volatilize N and to decompose nitrate (Raison, 1979). Nevertheless, the N
295 lost by volatilization in these plots was less than for the intense fires. Chemical analysis
296 of the ashes supports this, as the ashes of the moderate fire are richer in NH_4^+ (average
297 value $51.7 \text{ mg N kg}^{-1}$) than in the ashes from the intense fire ($35.7 \text{ mg N kg}^{-1}$). Thus, the
298 proportion of these ash particles removed with soil by erosion could cause the greatest
299 loss of NH_4^+ after the moderate fires.

300 In both fire treatments, the losses of ammonium N were greater than the losses
301 of nitrate N. In treatment F2, rain on 18 September (II) 1995 and 6 May 1996 produced
302 the greatest losses, between 2.2 and 2.4 mg N m^{-2} (Figure 4a). The greatest losses of
303 nitrate N correspond to the last rainfall event (6 December 1996), with 2.3 mg N m^{-2}
304 (Figure 4b). The losses of both forms of mineral N in the control plots were trivial
305 compared with the losses from the burned plots.

306 The enrichment in mineral N in the sediments from burned plots is related to the
307 increase or decrease of the mineral N concentration in the soil. After the intense rains of
308 autumn (four months after the experimental fires), there was a sharp decrease in
309 ammonium-N in the soil. This decrease was 73% in treatment F2 and 84% in F1. In
310 contrast, increases of 88% and 64% in nitrate-N concentrations in the soil of treatments
311 F2 and F1, respectively, were recorded 4 months after the fire respect to their levels

312 immediately after burning. Thus, the dominance of ammonium or nitrate remaining in
313 the soil results from their losses by erosion and appearing in the sediments.

314 Total loss of available phosphorus in the eroded sediments from the treatment F2
315 was, approximately, twice that lost from F1 (Table 7). Both were much larger than from
316 the control plots. The greatest losses of available P occurred in the rainfall event of 18
317 September (II) (Figure 5). Further differences were maintained until the end of the
318 study.

319 The losses of exchangeable cations showed the same trend as for other variables.
320 Those for K^+ and Mg^{2+} of treatment F2 were approximately, twice the ones for F1
321 (Table 8). The greatest losses on all exchangeable cations correspond to the rainfall
322 event of 18th September (II) (Figures 6 and 7).

323

324 *Net balance*

325

326 A net loss of organic matter, total N, nitrate-N and exchangeable Ca^{2+} occurred under
327 treatment F2, which seems mainly due to the high soil temperatures generated, rather
328 than to the overall losses of these compounds produced by erosion (Table 8). Only
329 losses of nitrate-N and Ca^{2+} , due to combustion, were detected in the plots that had the
330 moderate fire, and they were always less than those observed after the more intense fire
331 (Tables 8 and 9). The most notable difference observed, immediately after the fires, is
332 the increase in 8091 kg ha⁻¹ organic matter and 76.12 kg ha⁻¹ total N for the moderate
333 fire plots, whereas for the high intensity plots there was a decrease of organic matter
334 (942.5 kg ha⁻¹) and total N (58.0 kg ha⁻¹) (Tables 8 and 9). The increase in organic
335 matter and nitrogen after the moderate fire is attributable to the low combustion of the

336 organic matter due to the short duration of high temperatures on the soil surface and to
337 the deposition of partially burnt plant residues, as mention Andreu *et al.* (1996) and
338 Giovannini & Lucchesi (1997).

339 The overall soil losses by erosion during the seven rainfall events were 4.1 t ha⁻¹
340 for the plots in treatment F2, and 3.3 t ha⁻¹ for F1 plots. The losses of organic matter and
341 nitrogen in the sediments are reported in Tables 8 and 9. The greatest losses occurred
342 after the intense fire.

343 Considering the changes in the amounts of organic matter and nutrients,
344 immediately before and after fire, together with the losses produced by erosion, we
345 estimated the inputs and outputs to the soil system, as follows. There was a net loss of
346 1341 kg ha⁻¹ organic matter and 73.97 kg ha⁻¹ of total N for the intense fire treatment
347 (Table 8), and for the moderate fire, an input of 7724 kg ha⁻¹ organic matter and 61.5 kg
348 ha⁻¹ total N (Table 9). In both treatments there was a net loss of nitrate-N, whereas for
349 ammonium-N there was a net gain (Tables 8 and 9). These variations in mineral N
350 contents are strongly related to the intensity of fire.

351 We did not consider nutrient losses as soluble elements in runoff water in the
352 balance because they were much smaller than the losses in the sediments. For example,
353 we measured a maximum loss of water-soluble mineral N of 3.11 mg kg⁻¹ in the high
354 intensity plots for the rainfall of 6 May 1996 compared with 37.2 mg kg⁻¹ in the eroded
355 sediment. Other authors found that the soluble forms of these elements in runoff water
356 represent only 1.5% of the total losses (Soto *et al.*, 1995).

357 Net inputs of available phosphorus and exchangeable cations, except Ca²⁺ to the
358 soil system are estimated for both fire treatments (Table 8 and 9). Exchangeable Ca²⁺

359 shows a net loss of 598.4 and 497.2 kg ha⁻² for the high and moderate fire intensity
360 plots, respectively.

361 These results accord well with the data reported by several authors (Giovannini
362 *et al.*, 1990; Soto *et al.*, 1995; Andreu *et al.*, 1996), and they confirm that the modifying
363 processes occurring in the soil, and the increase in erosion rates and nutrient losses are
364 related mainly to the intensity of fire.

365

366

367 **Conclusions**

368

369 Fire caused an increase of ammonium N, available phosphorus and Na⁺, K⁺ and Mg²⁺
370 on the soil surface, whereas the nitrate-N content, the CEC and the exchangeable Ca²⁺
371 in the soil decrease after both intense and moderate fire. The combustion caused the
372 losses of organic matter, total N, nitrate N and exchangeable calcium in the most intense
373 fire.

374 Organic matter and nutrients removed with the sediments are closely related to
375 the degree of fire intensity. Soil lost by water erosion took with it organic matter and
376 nutrients. Four months after the fires, a sharp decrease of these constituents in the soil
377 was recorded, which coincides with the highest soil losses produced by the most erosive
378 rainstorms. In a unique rainfall event (18 September, I_{30} 35.36 mm h⁻¹), the greatest
379 losses of organic matter and nutrients were detected, which represent around 50% of the
380 total losses from the burned plots for the whole study period.

381 The calculated net balances show a net loss of organic matter and total nitrogen
382 for plots burned with the most intense fire, whereas an increase in these variables

383 occurred on plots having fire of moderate intensity. Both fire treatments show a net loss
384 of nitrate-N and exchangeable Ca^{2+} , whereas an increase of ammonium-N, available
385 phosphorus and the other cations was found.

386

387

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392

393

394 **References**

- 395 Andreu, V., Rubio, J.L. & Cerni, R. 1994. Long term effects of forest fires on soil
396 erosion and nutrient losses. In: *Soil Erosion and Degradation as a Consequence of*
397 *Forest Fires*, (eds M. Sala and J.L. Rubio), pp. 79-89. Geoforma Ediciones, Logroño.
- 398 Andreu, V., Rubio, J.L., Forteza, J. & Cerni, R. 1996. Postfire effects on soil properties
399 and nutrient losses. *International Journal of Wildland Fire*, **6**, 53-58.
- 400 Black, C.A., Evans, D.D., White, J.L., Ensminger, L.E. & Clark, F.F. 1965. *Methods of*
401 *Soil Analysis: Part 2, Chemical and Microbiological Properties*. American Society
402 of Agronomy, Madison, WI.
- 403 Blank, R.R. & Zamudio, D.C. 1998. The influence of wildfire on aqueous-extractable
404 soil solutes in forested and wet meadow ecosystems along the eastern front of Sierra-
405 Nevada Range, California. *International Journal of Wildland Fire*, **8**, 79-85.

- 406 Bower, C.A., Reitemeier, R.F. & Fireman, M. 1952. Exchangeable cations analysis of
407 saline and alkali soils. *Soil Science*, **73**, 251-261.
- 408 De Bano, L.F. 1971. The effect of hydrophobic substances on water movement in soil
409 during infiltration. *Soil Science Society America Proceedings*, **35**, 340-343.
- 410 Díaz-Fierros, F., Benito, E., Vega, J.A., Castelao, A., Soto, B., Pérez, R. & Taboada, T.
411 1990. Solute loss and soil erosion in burned soil from Galicia (NW Spain). In: *Fire in*
412 *Ecosystem Dynamics: Mediterranean and Northern Perspective*, (eds J.G.
413 Goldammer and M.J. Jenkins), pp. 103-116. SPB Academic Publishing, The Hague.
- 414 Etienne, M. & Legrand, C. 1994. A non-destructive method to estimate shrubland
415 biomass and combustibility. In: *Proceedings 2 International Conference on Forest*
416 *Fire Research*, Vol. I (ed. D.X. Viegas), pp. 425-434. Comissão de Coordenação da
417 Região Centro. Coimbra.
- 418 FAO-UNESCO 1988. *Soil Map of the World. Revised legend. 1: 5 000 000*. FAO,
419 Rome.
- 420 Giovannini, G. & Lucchesi, S. 1993. Effects of fire on soil physico-chemical
421 characteristics and erosion dynamics. In: *Fire in Mediterranean Ecosystems*, (eds L.
422 Trabaud and R. Prodon), pp. 403-412. Commission of the European Communities,
423 Brussels.
- 424 Giovannini, G. & Lucchesi, S. 1997. Modifications induced in soil physico-chemical
425 parameters by experimental fires at different intensities. *Soil Science*, **162**, 479-486.
- 426 Giovannini, G., Lucchesi, S. & Giachetti, M. 1990. Effect of heating on some chemical
427 parameters related to soil fertility and plant growth. *Soil Science*, **149**, 344-350.

- 428 Imeson, A.C., Verstraten, J.M., Van Mulligen, E.J. & Sevink, J. 1992. The effect of fire
429 and water repellency on infiltration and runoff under Mediterranean type forest.
430 *Catena*, **19**, 345-361.
- 431 Jackson, M.L. 1958. *Soil Chemical Analysis*. Prentice Hall, Englewood Cliffs. New
432 Jersey.
- 433 Mallik, A.U., Gimingham, C.H. & Rahman, A.A. 1984. Ecological effects of heather
434 burning. I: Water infiltration, moisture retention and porosity of surface soil. *Journal*
435 *of Ecology*, **72**, 767-776.
- 436 Peech, M. 1954. Determination of exchangeable cations and exchange capacity of soils.
437 Rapid micromethods utilizing centrifuge and spectrophotometer. *Soil Science*, **59**,
438 25-38.
- 439 Raison, R.J. 1979. Modification of the soil environment by vegetation fires, with
440 particular reference to nitrogen transformations: A review. *Plant and Soil*, **51**, 73-
441 108.
- 442 Raison, R.J., Khanna, P.K. & Woods, P.V. 1985. Mechanisms of element transfer to the
443 atmosphere during vegetation fires. *Canadian Journal of Forest Research*, **15**, 132-
444 140.
- 445 Rubio, J.L., Foretza, J., Andreu, V. & Cerni, R. 1996. Effects of forest fires on runoff
446 and soil erosion. In: *Soil Erosion Processes on Steep Lands. Evaluation and*
447 *Modelling* (eds I. Pla Sentís, R. López Falcón and D. Lobo Luján), pp 41-53.
448 CIDIAT. Mérida.
- 449 Soto, B., Basanta, R., Pérez, R. & Díaz-Fierros, F. 1995. An experimental study of the
450 influence of traditional slash-and-burn practices on soil erosion. *Catena*, **24**, 13-23.

451 Weast, R.C. 1980. *Handbook of Chemistry and Physics*, 60 ed. CRC Press, Boca Raton,
452 Florida.

453 Wienhold. B.J. & Klemmedson, J.O. 1992. Effect of prescribed fire on nitrogen and
454 phosphorus in Arizona chaparral soil-plant systems. *Arid Soil Research and*
455 *Rehabilitation*, **6**, 285-296.

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Table 1 Weather at the time of burning

Date	Burned plots	Air temperature / °C	Relative humidity / %	Wind direction	Wind speed / m s ⁻¹
20 June 1995	1	21	71	SE	0.3
	2	22	71	SE	0.3
	4	20	85	SE	0.3
21 June 1995	6 ^a	24	79	SE and SW	0.3 - 1.4
	7	22	82	SE	0.3
	8	22	83	Se	0.3

458

459

^a Plot number 6 suffered repeated changes of speed and wind direction when burning

460

461

Table 2 Mean values of soil variables for the different treatments, before and after fire

Variable	Before fire				After fire ^a			
	High	Moderate	Control	Standard error	High	Moderate	Control	Standard error
Organic matter /%	11.47	9.28	10.15	1.21	11.21 a	11.52 a	10.15 a	1.02
Total N /%	0.43	0.38	0.44	0.04	0.42 a	0.40 a	0.44 a	0.03
NH ₄ ⁺ -N /mg kg ⁻¹	3.54	4.71	2.16	0.09	87.27 c	52.36 b	2.16 a	0.98
NO ₃ ⁻ -N /mg kg ⁻¹	27.39	26.72	19.39	0.75	7.59 b	15.96 b	19.39 a	0.51
P _{available} /mg kg ⁻¹	11.62	6.41	8.08	0.21	66.44 c	38.34 b	8.08 a	1.13
CEC /cmol _c kg ⁻¹	31.25	28.83	27.15	1.81	25.23 b	24.98 b	27.15 a	2.02
Na ⁺ /%	1.51	2.00	1.09	0.57	4.84 b	5.54 b	1.09 a	0.69
K ⁺ /%	4.43	4.75	4.63	0.51	8.34 b	8.139 b	4.63 a	1.01
Mg ²⁺ /%	13.67	12.47	12.79	1.32	18.50 b	16.34 b	12.79 a	1.83
Ca ²⁺ /%	80.39	80.79	81.48	1.72	68.32 b	69.94 b	81.48 a	3.40

462

463 ^a Means not sharing the same letter indicate significant differences for the different fire treatment judged using Tukey's test ($p < 0.05$)

464 **Table 3** Characteristics of the rainfall events

465

Data	Days after fire	Precipitation /mm	Duration /minutes	I_{30} /mm h ⁻¹
23 Aug 1995	63	26.26	90	20.80
31 Aug 1995	71	9.36	285	14.56
18 Sep 1995 (I)	88	7.54	135	10.92
18 Sep 1995 (II)	88	18.72	95	35.36
4 Oct 1995	104	22.62	280	22.40
6 May 1996	316	22.62	165	26.00
6 Dec 1996	526	26.10	210	30.20

466

467

468 **Table 4** Mean values of runoff (l) and sediment (kg) production of the selected rain events for each fire intensity treatment and general
 469 mean of runoff yield (l m⁻²) and sediment yield (kg m⁻²)
 470

Data	Runoff				Sediment			
	High	Moderate	Control	Standard error	High	Moderate	Control	Standard error
23 Aug 1995	16.4	28.2	15.2	5.8	0.27	0.89	0.21	0.3
31 Aug 1995	6.7	10.6	3.6	2.3	0.08	0.19	0.03	0.04
18 Sep 1995 (I) ^a	8.9	12.2	3.8	4	0.04	0.08	0.01	0.03
18 Sep 1995 (II) ^a	312.2	267.6	21.4	38.3	14.9	12.6	0.08	4
4 Oct 1995	249.2	242.8	36.9	34.4	4.81	4.73	0.03	1.2
6 May 1996	289.6	198.8	41.2	37.8	7.43	4.98	0.15	2
6 Dec 1996	350.9	210.2	44.3	64.8	5.05	2.76	0.08	1.2
General mean	176.3	138.6	24.5		4.66	3.75	0.08	
General yield mean	15.4	12.1	2.14		0.41	0.33	0.01	

471 ^a (I) and (II) are two different rain events on the same day

472 **Table 5** Analysis of Variance with repeated measures for the runoff yield

473

474

a) Tests of Between-Subject Effects

Source of variation	df	MS	F	Probability
Fire treatment	2	22.448	18.526	0.003
Error	6	1.212		

475

476

b) Tests involving 'rainfall event' Within-Subject Effect

477

Source of variation	df	MS	F	Probability
Rainfall event	6	11.98	55.015	0.000
Rainfall x fire treatment	12	2.429	11.152	0.000
Error	36	0.218		

478

479

480 **Table 6** Analysis of Variance with repeated measures for the sediment yield

481

482

a) Tests of Between-Subject Effects

Source of variation	df	MS	F	Probability
Fire treatment	2	19262.6	8.172	0.019
Error	6	2357.1		

483

484

b) Tests involving 'rainfall event' Within-Subject Effect

Source of variation	df	MS	F	Probability
Rainfall event	6	14898	28.74	0.000
Rainfall x Fire treatment	12	3936	7.59	0.000
Error	36	518.28		

485

486

487

488 **Table 7** Total losses of organic matter and nutrients per unit area by water erosion

489

	Fire treatment			
	High intensity	Moderate intensity	Control	Standard error
Organic matter /g m ⁻²	39.91	36.65	0.59	0.49
Total N /g m ⁻²	1.59	1.46	0.023	0.02
NH ₄ ⁺ -N /mg m ⁻²	7.13	8.12	0.214	0.14
NO ₃ ⁻ -N /mg m ⁻²	4.82	7.45	0.146	0.23
P _{available} /mg m ⁻²	24.17	11.70	0.137	0.44
Na ⁺ /mg m ⁻²	17.38	12.37	0.067	0.22
K ⁺ /mg m ⁻²	178.50	103.93	0.421	3
Mg ²⁺ /mg m ⁻²	143.75	77.94	0.295	2.17
Ca ²⁺ /g m ⁻²	3.95	3.25	0.018	0.05

490

491

492 **Table 8** High fire intensity plots. Mean contents of organic matter and nutrients in the
 493 soil (kg ha^{-1}) before and after fire; gains or losses due to combustion processes (kg ha^{-1});
 494 losses by transported sediment (kg ha^{-1}) for the seven rainfall events; net balance in the
 495 soil system (kg ha^{-1})
 496

	High fire intensity					Net balance ^{b, c}
	Soil ^a		Combustion		Sediment	
	Before fire	After fire	Gains	Losses	Losses	
Organic matter	41 597	40 654	-	942	399	(-) 1 341
Total N	1 566	1 508	-	58.0	15.9	(-) 73.9
NH ₄ ⁺ -N	1.28	31.6	30.3	-	0.07	(+) 30.2
NO ₃ ⁻ -N	9.93	2.75	-	7.18	0.04	(-) 7.22
P _{available}	4.21	24.1	19.9	-	0.24	(+) 19.6
Na ⁺	39.2	98.4	59.2	-	0.17	(+) 59
K ⁺	196	294	98	-	1.78	(+) 96.2
Mg ²⁺	188	202	14.2	-	1.44	(+) 12.7
Ca ²⁺	1 821	1 271	-	550	39.5	(-) 589

497 ^a Soil density of the first 5 cm of soil was used for calculation (0.725 g cm^{-3})

498 ^b Inputs (+)

499 ^c Outputs (-)

500 **Table 9** Moderate fire intensity plots. Mean contents of organic matter and nutrients in
 501 the soil (kg ha⁻¹) before and after fire; gains or losses due to combustion processes (kg
 502 ha⁻¹); losses by transported sediment (kg ha⁻¹) for the seven rainfall events; net balance
 503 in the soil system (kg ha⁻¹)
 504

	Moderate fire intensity					Net balance ^{b, c}
	Soil ^a		Combustion		Sediment	
	Before fire	After fire	Gains	Losses	Losses	
Organic matter	33 662	41 752	8 091	-	366	(+) 7 725
Total N	1 377	1 454	76.1	-	14.6	(+) 61.5
NH ₄ ⁺ -N	1.70	19.0	17.3	-	0.08	(+) 17.2
NO ₃ ⁻ -N	9.68	5.78	-	3.90	0.07	(-) 3.97
P _{available}	2.32	13.9	11.6	-	0.12	(+) 11.7
Na ⁺	48.1	109	61	-	0.12	(+) 61.1
K ⁺	194	267	73	-	1.04	(+) 71.9
Mg ²⁺	158	172	14	-	0.78	(+) 13.2
Ca ²⁺	1 689	1 224	-	465	32.5	(-) 497.5

505 ^a Soil density of the first 5 cm of soil was used for calculation (0.725 g cm⁻³)

506 ^b Inputs (+)

507 ^c Outputs (-)

508 **Figure captions**

509

510 **Figure 1** Quantity and intensity (I_{30}) of rain of each erosive rainfall event for the studied
511 period

512

513 **Figure 2** Mean values of (a) runoff yield and (b) sediment yield registered in ‘La
514 Concordia’ plots for the different fire treatments by each studied rainfall event. Mean
515 values not sharing the same letter indicate significant differences for the interaction
516 ‘tratment x rainfall’ using Tukey’s test.

517

518 **Figure 3** Organic matter (a) and total N (b) losses per unit area in each studied rainfall
519 event

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521 **Figure 4** Ammonium-N (a) and nitrate-N (b) losses per unit area in each studied rainfall
522 event

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524 **Figure 5** Available phosphorous losses per unit area in each studied rainfall event

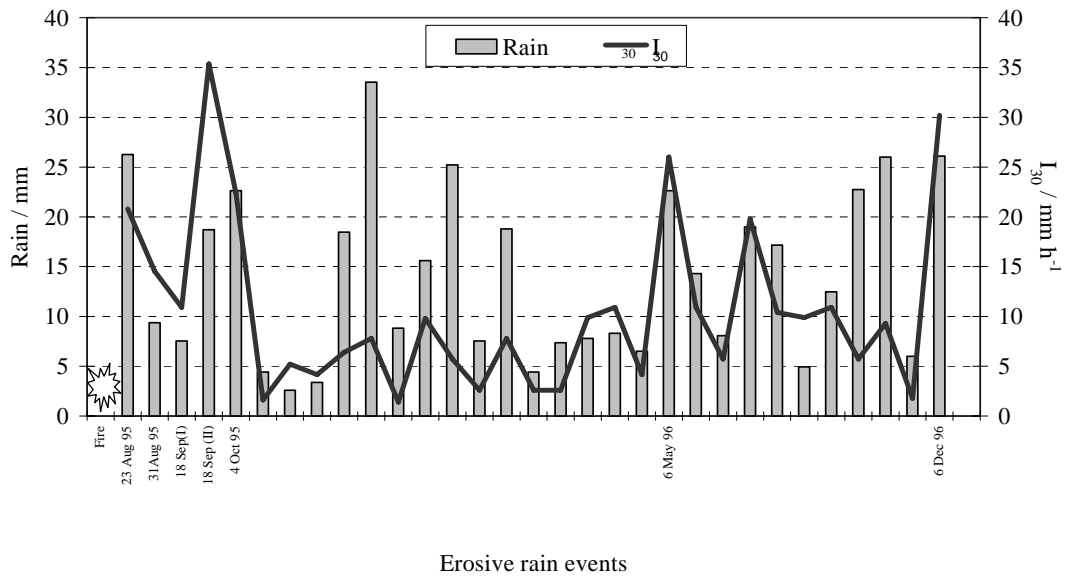
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526 **Figure 6** Exchangeable Na^+ (a) and exchangeable K^+ (b) losses per unit area in each
527 studied rainfall event

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529 **Figure 7** Exchangeable Mg^{2+} (a) and exchangeable Ca^{2+} (b) losses per unit area in each
530 studied rainfall event

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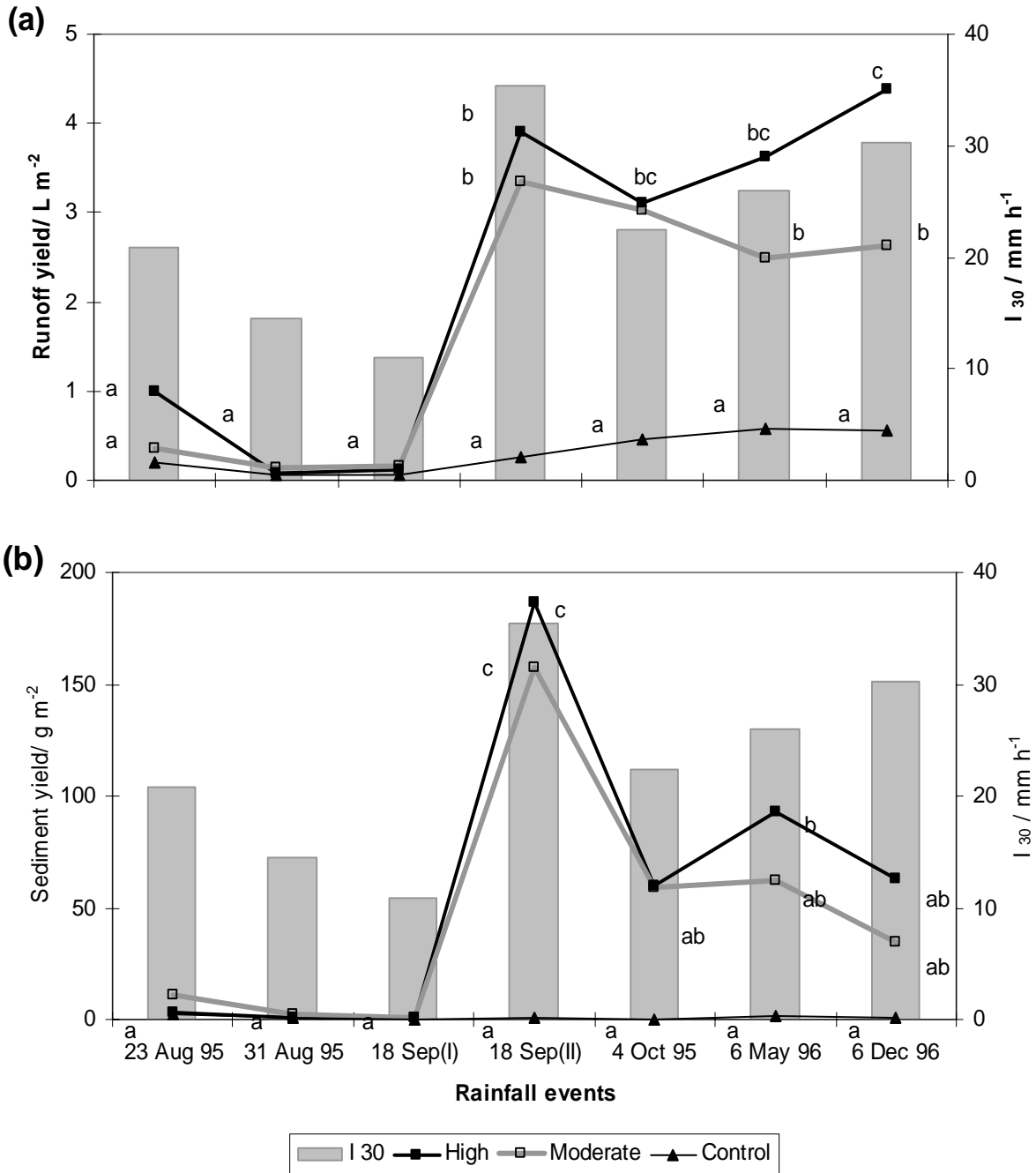
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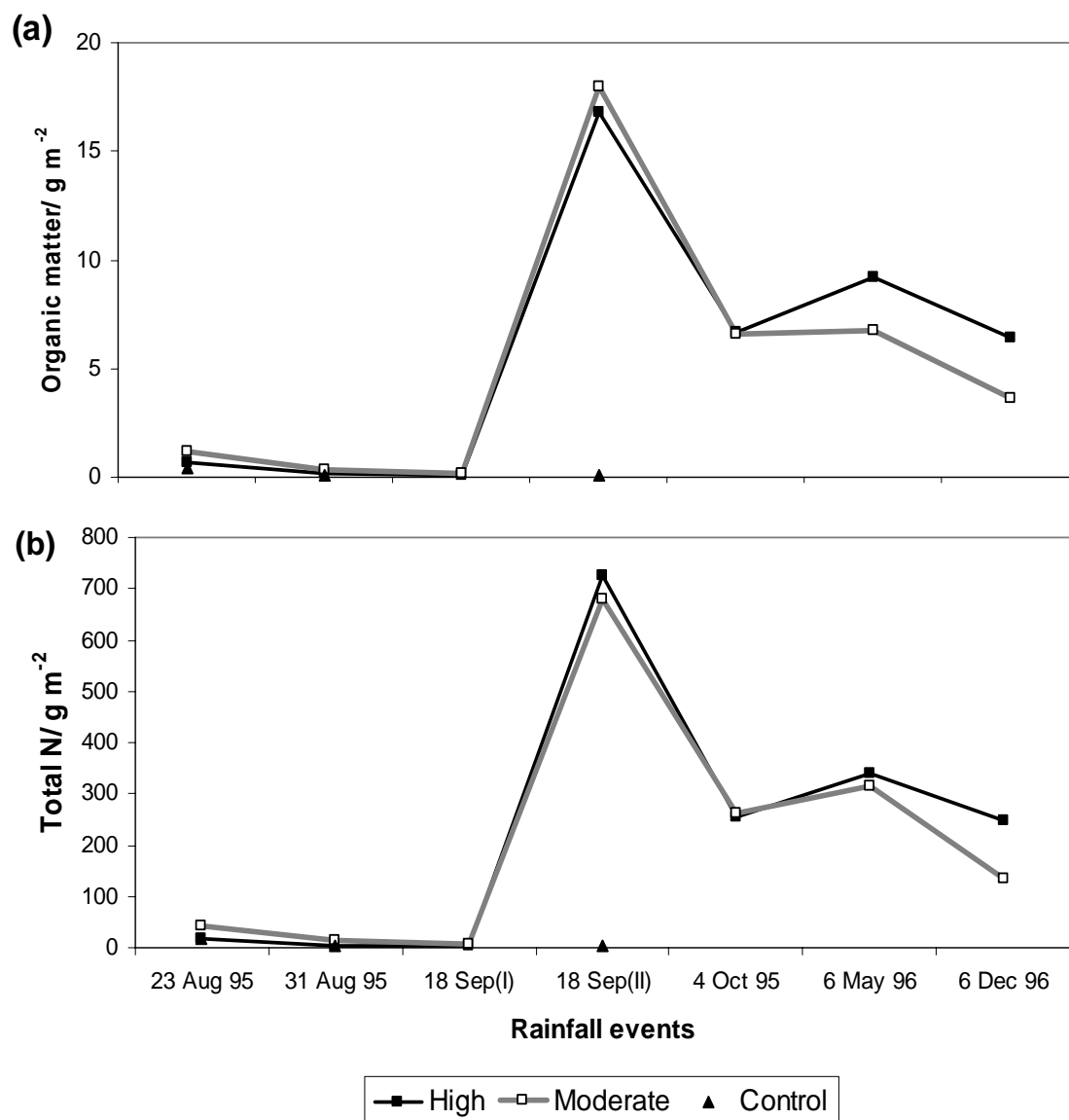
534 Fig. 1

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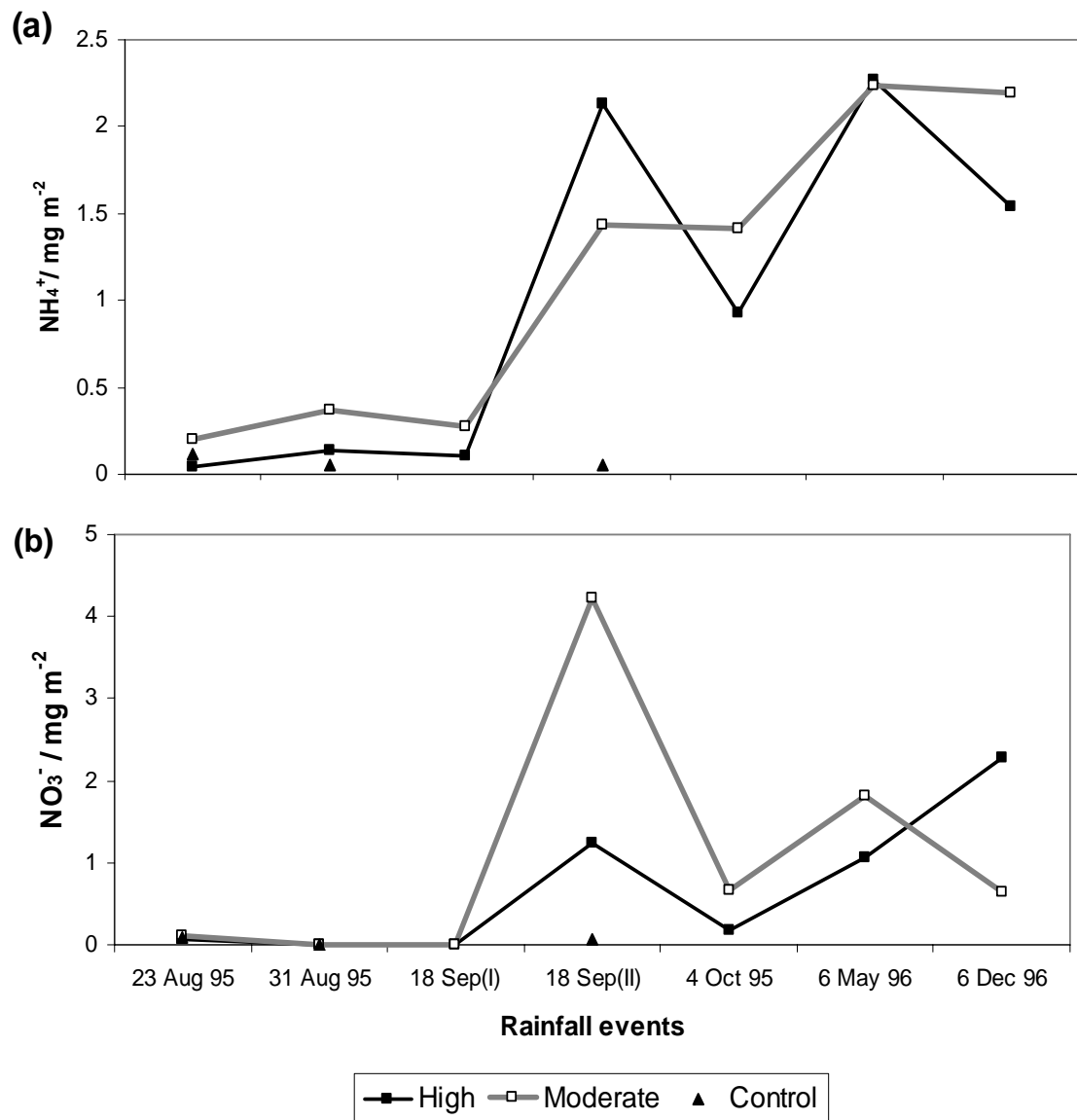
540 Fig. 2
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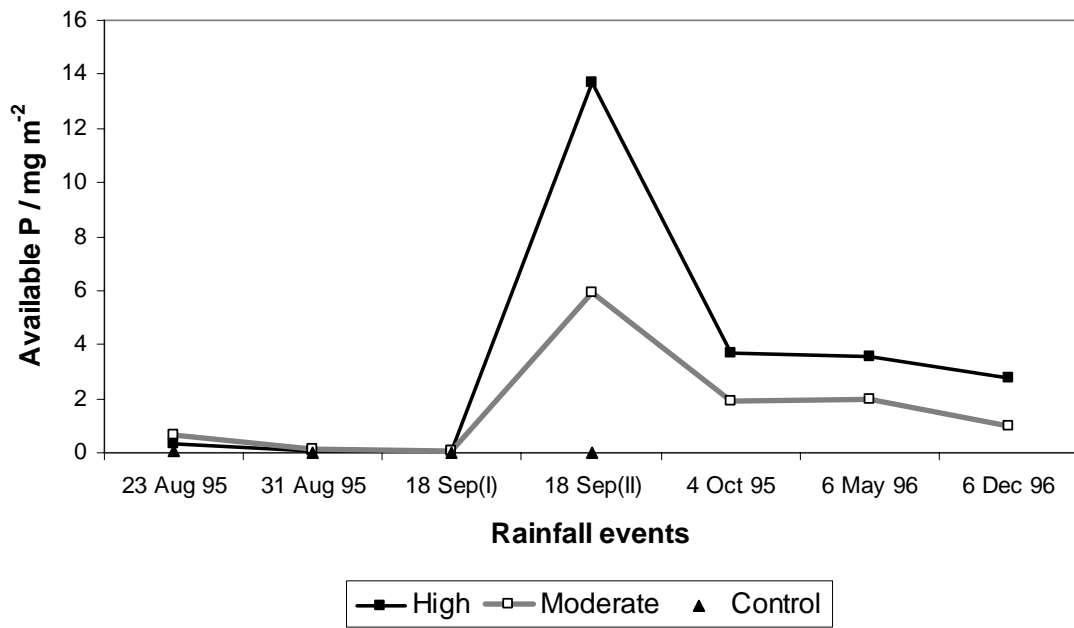
Fig. 3

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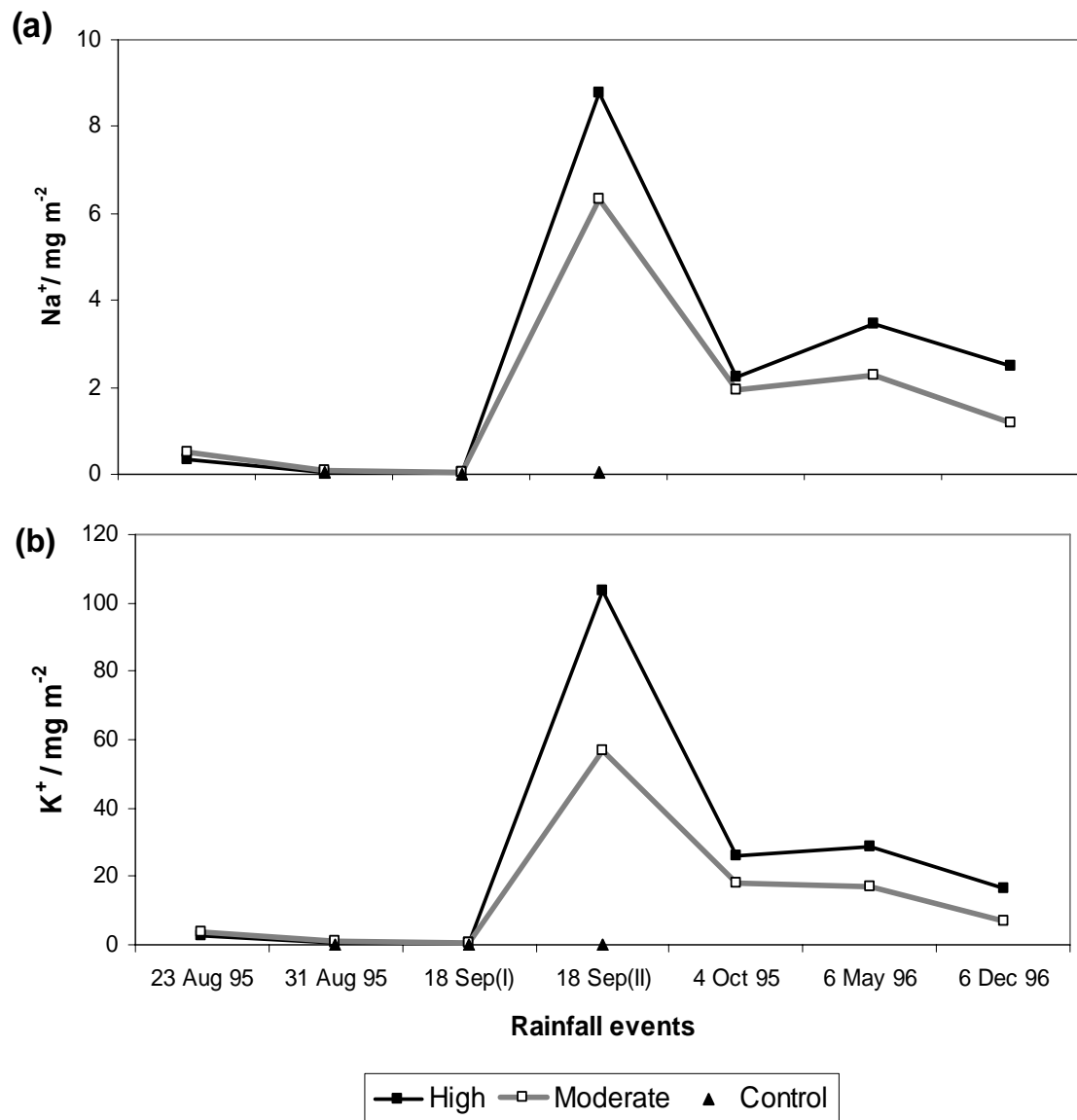
Fig. 4



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Fig. 5

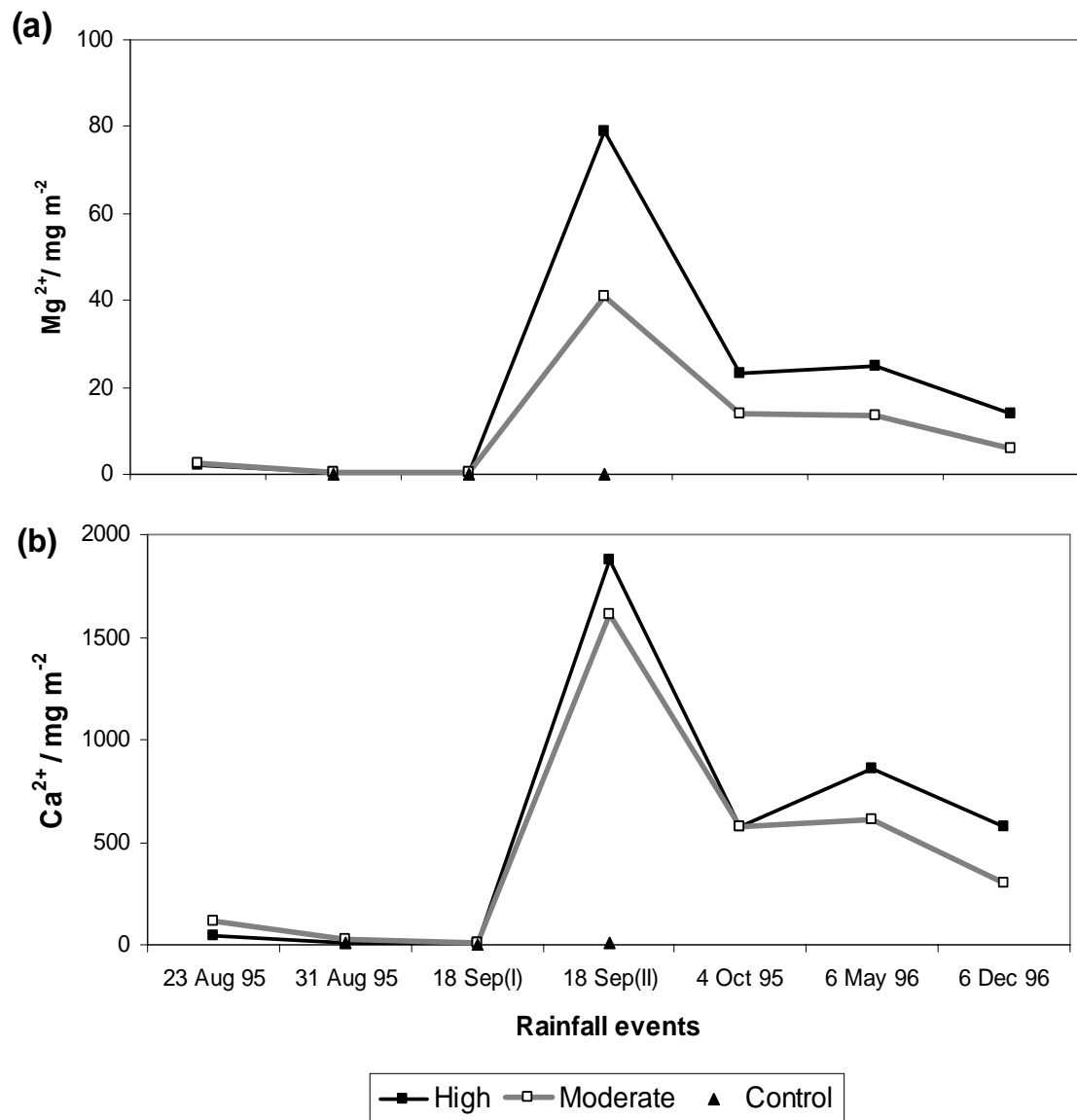
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Fig. 6

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Fig. 7

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