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**Hydrological Properties of Mediterranean Soils Burned
with Different Fire Intensities**

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25 **ABSTRACT**

26

27 The influence of vegetation cover on soil hydrological properties and its response to
28 the impact of different fire intensities, in a Mediterranean forest environment, has been
29 evaluated. The study was carried out in the Permanent Experimental Field Station of La
30 Concordia (Lliria-Valencia, Spain), on a set of nine erosion plots (4 x 20 m). The Station is
31 located on a calcareous hillside S-SE oriented, with soils of Rendzic Leptosol type and
32 supporting Mediterranean shrubland vegetation. All runoff generated and sediment produced
33 in every rain event was collected from each plot. The set up includes a system of sensors for
34 the continuous monitoring of climatic parameters (air temperature and humidity, rain volume,
35 intensity, etc).

36 In June 1995, a set of experimental fires were carried out on the Station. Three of the
37 plots were burned with high intensity fire, three with moderate intensity and the remainders
38 were left unaltered. Soil water content and water retention capacity (WRC) were measured in
39 the different plots and in two different vegetation covers: under canopy (UC) and in bare soil
40 (BS). The pF curves were also obtained for each fire treatment.

41 During a year after the fires (June 1995-June 1996) great differences in runoff
42 generation between fire treatments and the control plots were observed.

43 No significant differences were detected on water retention capacity between soils UC
44 and BS, in the burned plots. However these differences appeared in the control plots, giving
45 values for UC and BS of $0.130 \text{ cm}^3 \text{ cm}^{-3}$ and $0.180 \text{ cm}^3 \text{ cm}^{-3}$, respectively. Plots
46 corresponding to the high fire intensity treatment showed values of WRC significantly higher
47 than those of the moderate intensity and of the control treatments.

48 The pF curves show that the values of water volume, at the different pressure points
49 studied, were slightly greater on UC soil. Values obtained for BS samples are higher in the

50 fire treatments, showing significant differences in respect to the control plots at pF 1 and 2.
51 These differences were also observed for UC soil, but in this case at pF 2, 2.5 and 4.2.

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54 **Keywords:**

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57 **1. INTRODUCTION**

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59 In Mediterranean areas, shrubland vegetation is often structured in a spotted spatial
60 configuration, playing a significant role in controlling runoff generation and soil loss. The
61 interaction between vegetation development, soil surface properties and water movement
62 strongly influences the structure of Mediterranean ecosystems (Cammeraat *et al.*, 1999).
63 These patterns could change as a result of fire (Moreno, 1999).

64 Forest fires have become a common phenomenon during summer in many European
65 Mediterranean countries. Their effects are more evident on environments like those
66 characteristic of the Mediterranean area (Rubio and Recatala, 2005; Trabaud, 1990). The
67 immediate consequences are the loss of protective vegetation cover and a strong visual impact
68 on the landscape.

69 The soil environment, during and after fire is affected directly by the input of heat and
70 ashes. In the field, the effects of these factors are concomitant, making the identification of
71 individual causes of changes in soil properties, such as degradation of organic matter or
72 changes in aggregate size distribution, between others, difficult (Giovannini and Luchesi,
73 1997).

74 Post-fire conditions on soil surface are of key importance because they determine its
75 response to raindrop splash, overland flow and the development of water repellent soil
76 conditions (De Bano, 1981).

77 One of the most useful ways to study the effects of fire on the soil system is carrying
78 out fires in experimental plots. With this approach, it is possible to know and measure soil
79 conditions before, during and after the fire experiment and to improve knowledge about the
80 hydrology of the zone affected by different intensities of fire (Rubio *et al.*, 1994).

81 The aim of this study is to evaluate post-fire changes in hydrological properties on a
82 typical Mediterranean slope comparing fire affected and not affected soil. The evolution of
83 soil response to runoff processes was also studied in each rainfall event for a year after the fire
84 experiments.

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90 **2. MATERIALS AND METHODS**

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92 **2.1. Study Area**

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94 This work was carried out in the permanent Experimental Station of La Concordia
95 (Lliria-Valencia, Spain), 50 km NW of Valencia city (Figure 1). It is 575 m above sea level,
96 on a forested hillside South-South East facing, with a sclerophyllous shrub cover regenerated
97 after a previous wildfire occurred in 1978. The most abundant species include *Rosmarinus*
98 *officinalis*, *Ulex parviflorus*, *Quercus coccifera*, *Rhamnus lycioides*, *Stipa tenacissima*,
99 *Globularia alypum*, *Cistus clusii* and *Thymus vulgaris*.

100 Climatically the area belongs to the dry ombroclimate of the lower mesomediterranean
101 belt, according to Thornthwaite's classification. The average annual precipitation is around
102 400 mm with two maximums, autumn and spring, and a dry period from June to September.
103 Mean monthly temperatures range from 13.3°C in January to 25.8°C in August.

104 The soil is a Rendzic Leptosol (FAO-UNESCO, 1988), or Calcic Xerochrept type
105 according to Soil Taxonomy classification (Soil Survey Staff, 1990), developed on Jurassic
106 limestone. This soil has a variable depth, always less than 50 cm, abundant stoniness (\cong 40%)
107 and good drainage.

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110 **2.2. Experimental Set-Up**

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112 The Station consists on a set of nine erosion plots; 4 m wide x 20 m long each, with
113 similar characteristics such as soil morphology, slope gradient, rock outcrops and vegetation
114 cover. The selection of each plot location was made after intensive surveys of the vegetation, soil
115 and morphology patterns, based on across slope transects every two metres.

116 Plots were oriented parallel to the slope and bounded by bricks. At the foot of each
117 plot, a 2 m wide collector ran into a 1500 L tank to record all runoff and sediment produced
118 during each rainfall event. Inside each tank there was a 30 L tank to concentrate the sediments
119 produced, facilitating their collection.

120 A random design of two different fire intensity treatments, with three plots each, was
121 used. These different fire intensities were achieved by the addition of different amounts of
122 fuel load to the plots of each treatment, 40 t ha⁻¹ to reach the high intensity fires and 20 t ha⁻¹
123 for the moderate intensity ones. The fuel necessary to obtain the two fire intensities was taken
124 from the surrounding area using vegetation similar to that present on the plots, and its quantity

125 was calculated using a modification of the method proposed by Etienne and Legrand (1994).
126 The remaining three plots were maintained unburnt to be used as control. The temperatures on
127 soil surface and their duration were measured, on each square metre, by means of
128 thermosensitive paints and thermocouples. The mean soil surface temperatures reached were
129 439 °C for high fire intensity plots and 232 °C for the moderate intensity ones, and the
130 residence time in soil of temperatures greater than 100°C was 36'22" and 17'45" in each fire
131 treatment, respectively (Gimeno-García *et al.*, 2000).

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134 **2.3. Soil Analysis and Measurements**

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136 Soil samples were taken from the first 5 cm of the soil surface before and immediately
137 after fire. These were taken from under canopy (UC) and bare soil (BS). After that, they were
138 air-dried, screened to remove the fraction > 2 mm diameter and stored in plastic boxes for
139 analysis. The mean surface volumetric water content before fire experience was 0.08 cm⁻³ cm⁻³.

140 Soil Water Content (SWC) was calculated for the potentials: 0, -10, -33, -300, and -
141 1500 KPa, or pF 1, 2, 2.5, 3.5, 4.2, using the pressure membrane method (Richards, 1947).

142 Soil Water Retention Capacity (WRC) was calculated for each soil sample using the equations
143 of McLaren and Cameron (1996):

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$$145 \quad \text{(a) } WRC = (\theta_{10} - \theta_{1500}) * \rho_b \quad \text{(b) } WRC = (\theta_{33} - \theta_{1500}) * \rho_b$$

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147 Where WRC is Water Retention Capacity, θ_{10} , θ_{33} , θ_{1500} are gravimetric water
148 volumes at -10, -33, and -1500 KPa, and ρ_b is the bulk density of soil samples. The results
149 were obtained for volumetric units in percentages.

150 Water Retention Capacity with field capacity at -10 KPa and -33 KPa, was calculated
151 for the different plots, and the pF curves were also determined.

152 Climatic parameters and the intrinsic characteristics of the different rainfall events
153 were monitored by a logging system of sensors with GSM transmission of data. Runoff
154 generation dynamics were monitored in each rain event during the studied period. Rainfall
155 intensity was calculated for the maximum volume of precipitation occurring in 30 min (I_{30}).

156 Soil organic matter content was determined by oxidation with potassium dichromate
157 (Jackson, 1958). Electrical conductivity was measured in soil saturation extracts by the method
158 of Richards (1964). Aggregate stability was assessed using a wet-sieving procedure (Primo and
159 Carrasco, 1973). Calcium carbonate content were determined by the Bernard calcimeter
160 method (MAPA, 1986), and pH was determined in saturated paste (Richards, 1954).

161 Standard statistical analyses were applied at 95% of signification level. Analysis of
162 variance and Tukey's test at $\alpha=0.05$ were used to detect differences in WRC according to the
163 different fire treatments and vegetation cover. Climatic parameters were also analyzed with
164 ANOVA's test.

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167 **3. MAIN RESULTS AND DISCUSSION**

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169 **3.1. Rainfall characteristics**

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171 Total rainfall collected during the one year period after the experimental fires (June
172 1995-June 1996) was 386.62 mm. The total volume of erosive rainfall, with runoff
173 production, was 321.26 mm distributed in 24 events, showing an average I_{30} of 10.38 mm h⁻¹.
174 The maximum value of I_{30} was 35.36 mm h⁻¹ on 18th September, coinciding with the period

175 when the most aggressive rains in the Mediterranean region usually occurs (Perez Cuevas,
176 1994). The lowest value was 1.4 mm h^{-1} on 15th December (Figure 2A).

177 In winter, the total rain was 172.8 mm (distributed in 15 events), similar to autumn and
178 spring together (148.46 mm in 9 events), but the duration of storms was three times higher
179 than in the rest of year (average of 641 min and 216 min, respectively) (Figure 2B).
180 Moreover, duration was 72% higher and the I_{30} was 70% lower in winter than in the rest of
181 the study period and, as a consequence of this, two periods according rain characteristics can
182 be differentiated during the year of study: the last dates of summer 1995 plus autumn and
183 spring 1996 (period 1), and winter 1995/96 (period 2).

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186 **3.2. Hydrological trends**

187

188 In Table 1 are reported some soil characteristics analyzed after the fire experience,
189 distinguishing between vegetation cover (under canopy and bare soil) and fire treatments.

190 During the first year after the fire experiment, there are marked differences in runoff
191 generation between fire treatments and control plots. Plots affected by high intensity fire give
192 80% more runoff than control plots, meanwhile on plots affected by moderate intensity there
193 is 74% more runoff than on control ones (Figure 3). Benavides-Solorio (2001), shows that
194 runoff rates after the impact of fire on soil produce an increase of even three orders of
195 magnitude in erosion.

196 However, if these data are divided according the two periods defined by rain
197 characteristics, important differences in runoff yield between treatments and periods can be
198 observed. As it was explained above, period 1 is characterized by erosive rains of
199 medium/high I_{30} (average= 18.49 mm h^{-1}) and short duration. Period 2 show rains of low I_{30}

200 (average= 5.51 mm h⁻¹) and long duration, with high soil water content during this period and
201 low runoff values. In this way, runoff in period 2 was 57%, 65%, 18% lower than in period 1
202 for the plots affected by high intensity, moderate intensity and control, respectively (Figure
203 4A). It is possibly explained because the soil profile had not been completely saturated,
204 favouring lower infiltration rates than period 1 but similar for the different fire treatments
205 (Figure 4B), according to the observations made by Cerdà (1996) on similar environments.

206 The data obtained for the burned soils are in agreement with those obtained by
207 different authors (Rubio *et al.*, 1997; Andreu *et al.*, 2002), which indicate that in burned soils
208 the main factors that control runoff are the I₃₀ and the rain distribution during the year, so a
209 direct relation between runoff yield, rainfall volume and intensity was observed. Furthermore,
210 some authors like Lavee *et al.* (1998); Boix Fayos (1998); and Puigdefabregas (1999),
211 conclude that in the Mediterranean environment and in natural conditions, runoff rate is
212 highly dependent on rain regime and antecedent soil moisture conditions. Robichaud (2000)
213 also observed that the rain characteristics after a wildland fire are partially responsible for the
214 runoff rates generated. So, low rainfall intensities could facilitate a gradual wetting of the soil
215 profile, favouring changes or the disappearance of the hydrophobic substances generated on
216 topsoil after the fire, allowing then a normal infiltration rate. This could be the reason for the
217 difference in infiltration rates between the treatments in period 2. The differences in the
218 rainfall distribution through the year of study influence the hydrological trends in runoff yield
219 and in infiltration rate (Figure 4).

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222 **3.3. Water Retention Capacity**

223

224 The soils burned with high intensity fire showed values of WRC significantly higher
225 than those of the moderate and control treatment (Figure 5). After the fire experiment, the
226 burned plots presented homogeneous conditions on WRC for the different vegetation cover,
227 although in BS samples the WRC was slightly lower than in UC samples (Figure 5).

228 These results could indicate changes in physical properties at the soil surface (Andreu,
229 2001; Giovannini, 1994). These changes could be produced in particle-size distribution and
230 aggregation by the re-aggregation of clay-size particles into sand-size (Giovannini and
231 Lucchesi, 1997). When the WRC is calculated on the basis of a matric potential of -33 KPa,
232 the possible effect of the water held by the sand-sized particles is eliminated, and the
233 differences between fire treatments disappear. The fire effect could favour high water holding
234 at low pF values. Between the values -10 KPa and -33 KPa, the water content held is 75% and
235 55% higher for high and moderate intensity treatment, respectively, than control values
236 (Figure 5). Then, there is a significant amount of water held in soil at low pF values for the
237 high intensity treatment. This water is probably retained in the gaps generated by the re-
238 aggregation of clay particles into sand size particles. Guber *et al.* (2003), classifying
239 aggregates by size, using the average water content at -10, -33 and -1500 KPa, found that
240 larger aggregates show the greatest variation of water content and the greatest values for this
241 parameter.

242 Only control plots show significant differences between vegetation cover on WRC
243 (Figure 5). Soil samples taken from bare soil show higher WRC, which is possibly due to the
244 high superficial stoniness that covers the major part of the soil surface (a mean of $59 \pm 3\%$ of
245 surface stoniness). Then, the evapotranspiration rates could be lower because rock fragments
246 block the upward movement of water to the soil surface where evaporation can occur (Nobel,
247 1992). Because of this, the water retained between these pF values is upmost in the bare soil.
248 Bellot *et al.* (1999), found that under canopy soil, and in a Mediterranean environment, not all

249 the rainfall is received by the soil surface since the shrubland canopy interception reduces soil
250 water content due, among others factors, to the major evapotranspiration rates generated by
251 vegetation.

252 Cerdà (1998) and Bellot (1999), found that depending on the shrubland type
253 developed on the same soil, aggregate stability and soil water content can change. So, in
254 natural unburned areas the WRC depends, mainly, on the vegetation type and on soil
255 characteristics.

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258 **3.4. pF curves**

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260 Table 2, shows significant differences between burnt and control treatments for all pF
261 points in UC, except for pF 1, and pF 3.5. In BS the differences between fire treatment values
262 and control ones were only appreciable at low pF values, so the characteristics of the BS
263 samples make easier the physical re-aggregation of clay-size particles to sand-size particles
264 after fire impact than in UC soil (Molina *et al.*, 1998; Llinares, 2001). The soil UC in control
265 plots retains more water than the soil UC and BS in the plots affected by the impact of fire.
266 The pF curves for the high intensity treatment did not show significant differences between
267 UC and BS (Table 2), which could indicate that the pass of fire makes homogeneous soil
268 conditions, as it was observed by Boix Fayos (1997). In the moderate intensity treatment,
269 there are significant differences only for pF 1 and 3.5; meanwhile, for the control plots
270 significant differences were observed for pF 4.2. Those differences between UC and BS on
271 control samples could be due to a reorganization in the microaggregate fraction because at
272 this pF values (pF 4.2), the main factor that possibly determines the matric water retention are

273 the texture and the specific surface of the soil material that join water and soil by adsorption
274 forces (Hillel, 1980).

275 The values of water content for BS samples of fire treatments are higher than those of
276 control samples. In this way, the fire effect on bare soil in relation to its hydrological
277 properties could bring about, initially, an increase in the water content mainly at low pF
278 values (pF 1, 2). This increase shows significant differences for pF 1 and 2 (Table 2).

279 The fact that values of water content in BS samples increase probably depends, among
280 others factors, on structural changes in topsoil after the fire. These structural changes are
281 related to an increase in the macroaggregates fraction favoured by particle cementation
282 processes (Molina *et al.*, 1998; Llinares, 2001; Andreu, 2001).

283 The possibility of a macroaggregates increase in soil surface layers could explain the
284 volume of water retained at low pF values by the soils affected by fire. At these values, the
285 amount of water depends primarily on the capillary effect and the pore-size distribution, and
286 hence, it is strongly affected by soil structure (Hillel, 1980). The rise in the macroaggregate
287 fraction on soil surface accompanied by the decrease of microaggregates, immediately after
288 the fire experiment (Molina *et al.*, 1998; Llinares, 2001), could probably produce an increase
289 on pore volume and water content of soil. These large aggregates present the greatest
290 variations in water content and the greatest values in this parameter (Guber *et al.*, 2003).

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293 **4. CONCLUSIONS**

294

295 The rain distribution during the year after the experimental fires shows clearly two
296 different periods. One is characterized by medium/high intensity and low duration rains in
297 spring and autumn, and the other, in winter, with rains of long duration and low I_{30}
298 characteristics.

299 The impact of fire on soil has important hydrological consequences in the most
300 aggressive rainfall period (spring and autumn). The differences in runoff generation between
301 this period and winter are above 20% for control plots and 60% for the burned plots. These
302 values emphasize the importance of rainfall characteristics in the immediate period after the
303 fire experiments.

304 The hydrological properties of soil are also affected by the impact of fire. It produced
305 the homogenization of water retention capacity values between vegetation covers (under
306 canopy and bare soil), and the increase in water content of the bare soil in burned plots versus
307 control ones at low pF values. This increase in water content at low pF values could indicate
308 structural changes in the soil surface.

309 In relation with water retention capacity, the pF curves show substantial differences
310 between UC and BS. In some points of the pF curves there are higher values in BS than in UC
311 (statistically not significant). However in the pF range between 3.5 and 4.2, on control plots,
312 the values of soil water retention under canopy soil are slightly higher than on bare soil.

313 In the Mediterranean area, the impacts of fires are magnified by the changing
314 characteristics of the rain regime. This fact and the increase in frequency of forest fires could
315 favour the progressive ecosystem degradation and the increase of desertification risk.

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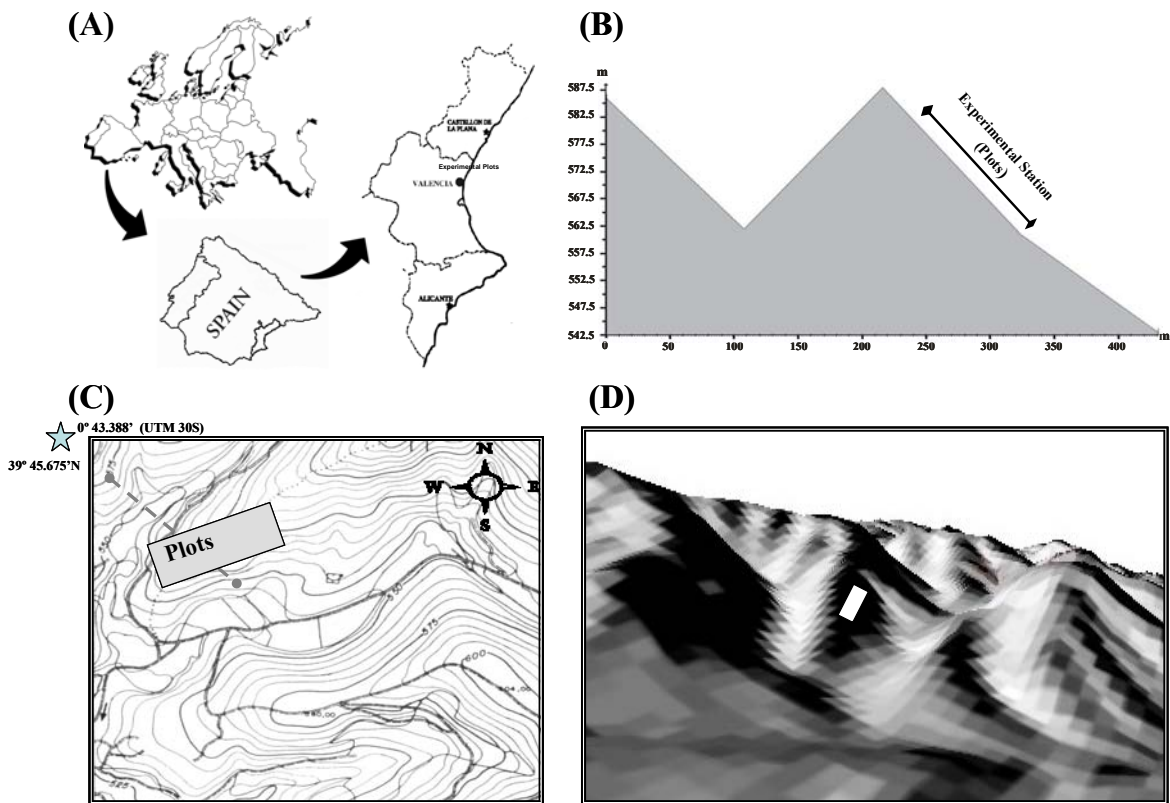
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403 *Figure 1. (A) Geographical location of the Experimental Station of La Concordia (Lliria-Valencia, Spain). Morphological*
 404 *characteristics of the study area: (B) Profile with altitudes and distances. (C) Topographic map with altitudes and*
 405 *coordinates. Grey broken line indicates the profile B. (D) Digital terrain model with the location of the plots (white*
 406 *rectangle).*

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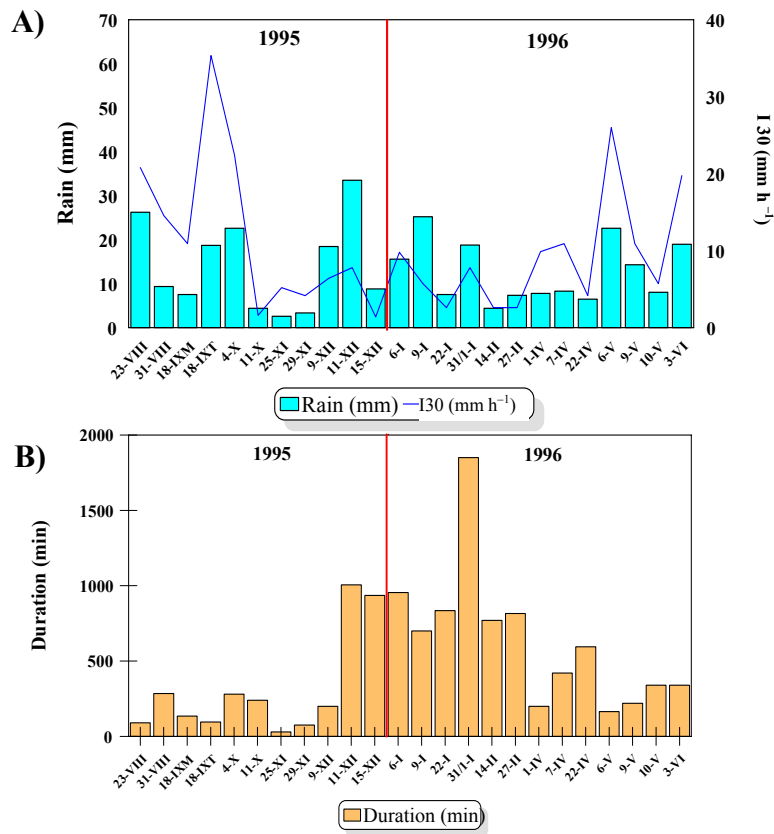
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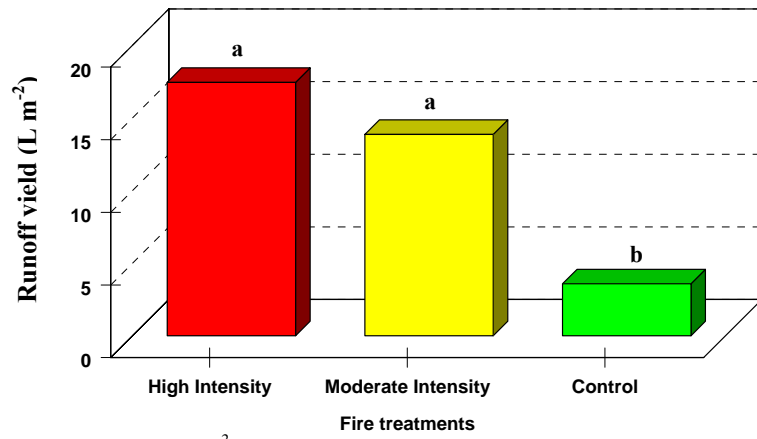
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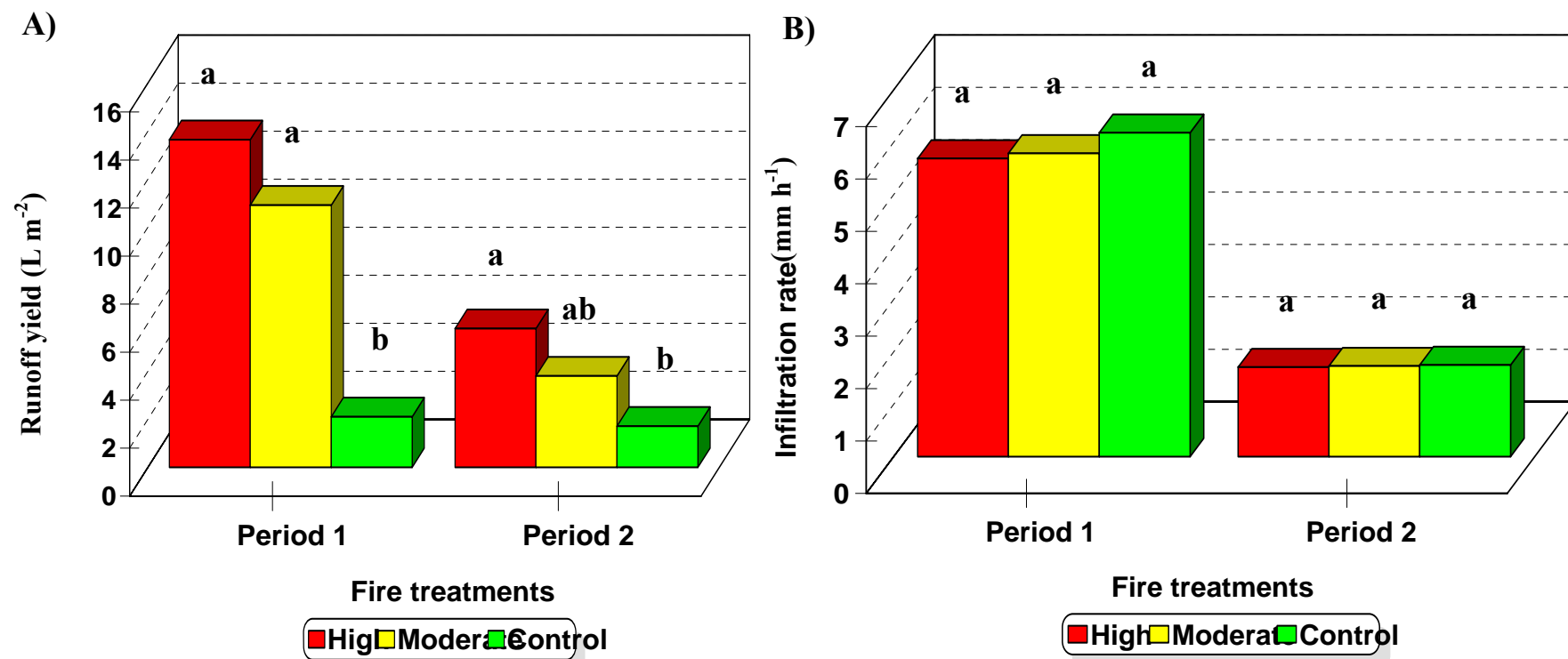
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Figure 2. Volume (A), intensity (A), and duration (B), of erosive rains occurring during the studied period.



431 *Figure 3. Total values of runoff yield (L m⁻²) corresponding to the different fire treatments during the studied period. Values*
 432 *not sharing the same letter indicate significant differences between fire treatment using Tukey's test (P<0.05).*

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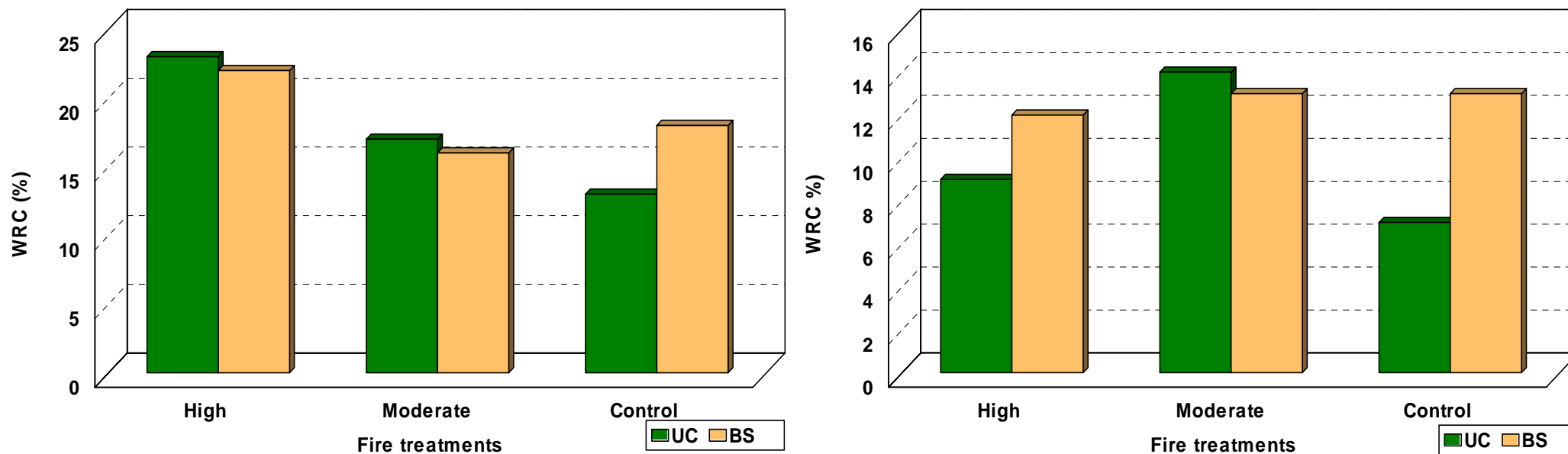


450 Figure 4. Total values of runoff yield (L m⁻²) (A) and infiltration rate (mm h⁻¹) (B), for the different treatments and periods described during 1995-96. Values not sharing the same letter indicate
 451 significant differences between fire treatment using Tukey's test ($P < 0.05$).

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456 *Figure 5. Water Retention Capacity (WRC) calculated at matric potentials of -10 KPa and -1500 KPa (left side), and calculated at matric potentials of -33 KPa and -1500 KPa (right side), for*
457 *the different fire treatments and vegetation cover, immediately after the fire experiment (1995). UC (under canopy), BS (bare soil). Values not sharing the same letter indicate*
458 *significant differences for the different vegetation cover (lower case) according to the Tukey's test. Differences between treatments are also showed (capital letter)*
459 *Tukey's test ($P < 0.05$).*

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462 *Table 1. Summary statistics for temperatura data (C°) measured with the thermosensitive paints*

Fire treatment	High (4Kg m ⁻²)			Moderate (2Kg m ⁻²)		
	1	4	8	2	6	7
Plots						
N	80	80	80	80	80	80
Mean (C°) ^a	417.78 a	448.09 a	434.91 a	239.90 b	239.46 b	217.54 b
Median	420	454	420	226	226	198
SD	118.78	132.63	147.32	90.71	91.58	81.61

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464 ^a Different lower case letter among High and Moderate treatments indicates statistically significant difference at

465 *P*<0.05

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484 Table 2. Mean values of some soil properties for each fire treatment and vegetation cover in 1995, after fire
 485 experiment.

<u>Before fire experience</u>						
Vegetation cover	A.S. (%)	pH (water)	E.C.(dS/m)	CO₃Ca (%)	O.M.(%)	
UC	-	7.29a	1.14a	45.42a	12.11a	
BS	-	7.5b	0.61b	50.15b	8.49b	

<u>After fire experience</u>						
Vegetation cover	Treatment	A.S. (%)	pH (water)	E.C.(dS/m)	CO₃Ca (%)	O.M.(%)
UC	High	33.68a	7.20a	3.59a	45.85a	12.85a
UC	Moderate	31.59a	7.34a	2.51ab	46.94a	11.85a
UC	Control	35.32a	7.33a	1.04b	45.65a	12.33a
BS	High	28.13a	7.21a	2.71a	48.35a	9.57ab
BS	Moderate	23.46a	7.38ab	1.70b	49.01a	11.18a
BS	Control	23.53a	7.52b	0.68c	49.00a	7.98b

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487 UC, Under Canopy; BS, Bare Soil. A.S., aggregate stability; E.C., electrical conductivity; CO₃Ca, calcium carbonate;
 488 O.M., organic matter. Values not sharing the same letter in columns indicate significant differences between fire
 489 treatment for the different vegetation cover, using Tukey's test (P<0.05).

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503 *Table 3. Water Content (%), for the different vegetation cover and treatments for the soil samples taken after the fire*
 504 *experiment.*

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506	Levels of comparison	pF 1	pF 2	pF 2.5	pF 3.5	pF 4.2
507	High I	47.79a	38.56a	24.49a	22.44a	15.67a
508	UC Moderate I	48.76a	32.96b	30.03b	22.48a	15.64a
509	Control	47.52a	34.35b	27.46ab	23.81a	20.87b
510	High I	45.37a	37.09a	27.30a	22.95a	14.95a
511	BS Moderate I	52.05b	31.78b	28.71a	24.28a	15.72a
512	Control	47.47ab	33.19b	28.24a	22.65a	14.99a

514	Levels of comparison	pF 1	pF 2	pF 2.5	pF 3.5	pF 4.2
515	High UC	47.79a	38.56a	24.49a	22.44a	15.67a
516	Intensity BS	45.37a	37.09a	27.30a	22.95a	14.95a
517	Moderate UC	48.76a	32.96a	30.03a	22.48a	16.64a
518	Intensity BS	52.05a	31.78a	28.71a	24.28b	15.72a
519	Control UC	47.52a	34.35a	27.46a	23.81a	20.87a
520	Control BS	47.47a	33.19a	28.24a	22.65a	14.99b

521 (UC) Under canopy; (BS) Bare soil. A) significant differences in water content between fire treatments depending on
 522 vegetation cover, and B) significant differences in water content between vegetation cover depending on fire intensity
 523 treatments, for different pF values according to ANOVA's test. Values not sharing the same letter in a column
 524 indicate significant differences for the different treatment and vegetation cover, judged using Tukey's test (P<0.05).
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