THE EFFECT OF DRY GRINDING ON ANTIGORITE FROM MULHACEN, SPAIN

A. DRIEF AND F. NIETO

Departamento de Mineralogía y Petrología, Instituto Andaluz de Ciencias de la Tierra, Universidad de Granada-C.S.I.C., Av. Fuentenueva s/n, 18002 Granada, Spain

Abstract—Alteration of the crystal structure of Mulhacén antigorite caused by dry, vibration grinding was investigated by X-ray diffraction (XRD), infrared spectroscopy (IR), thermal analyses (TG), grainsize distribution, and transmission and analytical electron microscopy (TEM, AEM). Grinding for 1 min reduces particles to a size ideal for IR and TG. With prolonged grinding, XRD and electron diffraction patterns showed that the crystal structure was affected mainly along the *c* axis, causing a partial loss of crystallinity. TG analyses revealed that vibration grinding modified mineral dehydration, accelerating the dehydroxylation process and transforming the structural OH to adsorbed water in the resulting matrix. IR spectra and AEM showed that grinding affected the tetrahedral sheet to a lesser extent than the octahedral sheet. Partial release of Mg by preferential destruction of the octahedral sheet after 10 min grinding produced an increase in the Si/Mg ratio in semi-crystalline particles, whereas the amorphous material product after 120 min showed the same composition as the initial antigorite. TEM and grain-size distribution results revealed that grinding led to a general decrease in particle size at the beginning of the experiment followed by the agglomeration of ultrafine particles as grinding proceeded.

Key Words—Antigorite, Crystallinity, Dehydroxylation, Dry Grinding, IR Spectra, Serpentine, Thermal Analysis.

INTRODUCTION

Grinding is a common treatment for phyllosilicates in industrial operations, in research laboratories, and in studies related to soil genesis. Grinding affects the structure and properties of layer-silicates in various ways. The processes involved were widely researched, especially for kaolinite, montmorillonite, and mica (e.g., Halavay et al., 1977; Aglietti et al., 1986). Precise details on the degradation processes of montmorillonite by dry grinding were given by Cícel and Kranz (1981). They suggested that the exchangeable cations are affected first, then the hydroxyl groups and finally, the separation of the tetrahedral and octahedral sheets occurs. In a study on sepiolite, Cornejo and Hermosín (1988) found that the cross-linked ribbon structure of sepiolite is much more resistant to mechanical stress than the layer structure of smectite or kaolinite. Nevertheless, little is known of the behavior of serpentine to grinding. Papirer and Roland (1981), in a study on chrysotile, found that grinding in lowviscosity organic solvents leads to three successive and overlapping processes; rapid defiberization of the asbestos bundles followed by fragmentation of the isolated fibrils and, finally, amorphization and agglomeration resulting in a drastic decrease in the specific area.

The purpose of this work is to clarify the mechanically induced crystal-structure distortion of antigorite to a quasi-amorphous state using a vibration grinder. To our knowledge, this is the first comprehensive investigation on the alteration of the crystal structure of antigorite by vibration grinding.

EXPERIMENTAL

Material

The material used is a serpentinite rock from the Mulhacén group of the Sierra Nevada Complex in the Betic Cordillera (SE Spain). The mineral composition was determined by optical microscopy, X-ray diffraction (XRD), and transmission electron microscopy (TEM), and analytical electron microscopy (AEM). Antigorite was the major serpentine mineral, with small amounts of chrysotile and berthierine (detectable only by TEM). The major impurities were olivine and magnetite (both 5%). Electron microprobe analyses (EMPA) of the Mulhacén antigorite and related chemical formulae, are shown in Table 1.

Grinding procedures

An HSM 100 vibration grinder was used for grinding. It consists of a cylindrical tungsten container (14 cm inner diameter) within which a tungsten ring (10 cm diameter) and cylinder (8 cm diameter) are concentrically placed. The rock was initially ground in amounts of 400 g for 30 s to allow for suitable sample for XRD, TEM, infrared (IR) and thermal gravimetric (TG) analyses. This material was then considered as "starting" material (sample at 0 min grinding). The starting material was ground in amounts of 100 g for 1, 10, 60, and 120 min. After every minute of grinding the process was halted for 30 s to avoid overheating and after 5 min of grinding the process was halted for 90 s for the same purpose. The temperature during grinding never exceeded 60°C. Presented data are total grinding time and do not include the "rest" time.

Oxides	1	2	3	4	5	6	7	8	9	10
Na ₂ O	0.00	0.00	0.01	0.00	0.01	0.00	0.01	0.00	0.00	0.00
MgO	39.12	39.00	39.80	39.30	39.50	39.10	39.52	39.40	39.12	39.68
SiO ₂	43.47	43.30	44.40	43.60	43.90	43.30	43.57	43.60	43.45	44.06
$Al_2 \tilde{O}_3$	2.33	2.44	1.12	2.31	1.93	2.24	1.80	2.26	2.03	1.10
K ₂ O	0.01	0.01	0.01	0.01	0.01	0.00	0.00	0.00	0.01	0.00
TiO ₂	0.01	0.01	0.00	0.02	0.00	0.01	0.01	0.02	0.00	0.00
Ca ₂ O	0.00	0.01	0.00	0.01	0.04	0.00	0.01	0.00	0.00	0.00
Cr_2O_3	0.00	0.00	0.08	0.22	0.05	0.19	0.14	0.13	0.10	0.02
Feo	2.07	2.16	2.27	2.13	2.15	2.18	2.12	2.14	2.07	1.94
MnO	0.05	0.07	0.05	0.04	0.05	0.06	0.05	0.02	0.02	0.05
Total	87.09	87.00	87.80	87.60	87.70	87.10	87.27	87.80	86.85	86.97
		Nur	nber of ior	is on the b	asis of five	O and fou	ur (OH)			
	1	2	3	4	5	6	7	8	9	10
Si	2.01	2.01	2.04	2.00	2.02	2.00	2.01	2.00	2.01	2.04
Mg	2.70	2.69	2.72	2.69	2.70	2.70	2.72	2.70	2.70	2,74
Fe	0.08	0.08	0.08	0.08	0.08	0.08	0.08	0.08	0.08	0.07
Al	0.12	0.13	0.06	0.12	0.10	0.12	0.09	0.12	0.11	0.06
Σ (Mg, Fe, Al)	2.90	2.90	2.86	2.89	2.88	2.90	2.89	2.90	2.89	2.87

Table 1. Electron-microprobe analyses of antigorite and structural formulae on the basis of five O and four (OH).

Examination of the material

An uncovered thin section was prepared for initial optical study and the electron microprobe analysis (EMPA) which followed the procedure as Nieto (1997). Powder XRD patterns were recorded on the starting and ground samples with a Philips PW-1710 diffractometer using CuK α radiation, also following



Figure 1. Infrared spectra of starting and ground antigorite for times indicated.

Nieto (1997). Determination of crystallinity values followed Kisch (1991). The IR spectra were obtained with a Nicolet 20SXB FT-IR spectrometer over the range 4000–400 cm⁻¹ at a scanning speed of 600 cm⁻¹/min from pressed discs using 1.0 mg of ground sample and 800 mg of KBr.

For TEM and AEM analyses, material at 0, 10, and 120 min of grinding, was dispersed in distilled water and deposited on collodion films supported on copper grids and air-dried at room temperature. TEM images, electron diffraction patterns, and semi-quantitative analyses were obtained with a Philips CM20 instrument operated at 200 kV at the "Centro de Instrumentación Científica" (C.I.C.) at Granada University. Experimental conditions were the same as in Sánchez-Navas et al. (1998). The AEM data (not shown) of "starting" antigorite are very close to those obtained by EMPA. Grain-size distribution of the ground samples was determined on a GLAI-CIS-1 apparatus using aqueous solution as a dispersive medium. The TG analyses were performed on 43 mg of sample in a SHIMADZU TGA-50H in air at a heating rate of 10°C/min.

RESULTS

Infrared spectra

Changes in the transmittance band intensities during grinding are illustrated in Figure 1. Table 2 presents the band assignments. Note that all the initial infrared bands increase in value with grinding, and become better defined between 0-1 min (Figure 1). After becoming better defined, however, the values decreased to below the initial values of the starting sample. The transmittance bands at 1066, 622, and 563 cm⁻¹, are characteristic of the Si-O stretching vibrations with perpendicular polarization, of the OH deformation,

Table 2. Transmittance bands and related group-of-atom vibrations according to 1 = Luce (1971), 2 = Heller-Kallai et al. (1975), 3 = Yariv and Heller-Kallai (1975), and 4 = Wilson (1987). W = Wavenumbers.

W (cm ⁻¹)	Attributed to	References	
3673	Structural OH	1	
3560	Substitution of trivalent ions for Mg	2	
3436	Adsorbed water	2	
1066	Si-O stretching vibrations		
	with perpendicular polarisation	3	
985	Si-O vibrations in the basal plane	3	
622	OH deformation	3,4	
563	Mg-O out-of-plane	3	
447	Si-O stretching vibrations	1	

and of the Mg-O out-of-plane vibrations, respectively. The bands were well developed in the spectra of the sample ground for 1 min. With prolonged grinding, the intensities of these bands decreased and after 120 min of grinding they disappeared entirely. The OH-stretching vibration band at 3673 cm^{-1} and that of the Si-O vibrations in the basal plane at 985 cm^{-1} show a general decrease by grinding. The former is weakly present in the spectrum corresponding to 120 min grinding, whereas the latter is still well-defined. The band at 3436 cm^{-1} , corresponding to adsorbed water, increases with increasing duration of grinding.

To clarify the relationship between the bands at 3673 and 3436 cm⁻¹, the transmittance ratio T3436/T3673 was measured as a function of time of grinding, as shown in Figure 2. Note that the ratio initially decreased slightly, but then a general increase was observed.

X-ray study

The degree of lattice distortion and amorphization of the samples were studied by powder XRD by comparing the Mulhacén antigorite to the patterns reported by Bailey (1980). When antigorite was ground for 1– 10 min, considerable destruction of the structure occurred, as indicated by a general decrease in the intensity of the reflections (Figure 3). Grinding for 120 min caused complete destruction of the mineral, at least along the c axis, and gives rise to diffraction patterns with only a few weak (*hkl*) reflections.

Figure 4 shows the peak-area evolution for each of the indicated reflections as a function of time of grinding. Significant differences are clear in the rate of decrease in peak area of these diffraction bands. The greatest change was observed for the (001) reflection at 7.22 Å, which decreased in area in the first minute of grinding. The (16,0,1) reflection at 2.52 Å increased in area from 0 to 10 min of grinding and then declined to below the initial value, whereas the combined ($83\overline{2}$) and (16,0,2) reflections at 2.16 Å showed an intermediate behavior between the (001) and (16,0,1) reflections. The *d*(001), the width at half-height, and the



Figure 2. Changes in the transmittance ratio T3436/T3673 during grinding. Note that the scale varies for the time axis.

integral breadth (peak area/peak intensity) values as a function of grinding time are presented in Table 3. The width at half-height and the integral breadth of the basal reflections (001) and (002) increased slowly, with the *d*-values also showing a small increase.



Figure 3. Powder X-ray diffractograms of starting and ground antigorite for times indicated. (CuK α radiation).



Figure 4. Changes in peak area of the basal reflection at 7.22 Å and two (*hkl*) reflections at 2.52 and 2.16 Å during grinding. Note that the scale varies for the time axis.

Electron microscopy

Electron diffraction. The Mulhacén antigorite samples (0, 10, and 120 min grinding) were also examined by electron diffraction. Because of the (001) cleavage, all TEM observations were made approximately perpendicular to the (001) plane. Figure 5 presents electron diffraction patterns showing the effect of grinding on the structure of the mineral. The electron diffraction pattern of the starting antigorite is shown in Figure 5a, and this pattern is very similar to those by Yada (1979), Viti and Mellini (1996), and Viti (1997). After 10 min of grinding, three different materials occurred in the sample: an amorphous material with no significant pattern, a well-crystallized antigorite similar to Figure 5a (neither are shown in Figure 5), and an intermediate transitional material with diffuse rings (Figure 5b). The material ground for 120 min displayed an amorphous material with no significant pattern. Oc-

Table 3. The d(001), the width at half height, and the integral breadth of the (001) and the (002) reflections during grinding. (*) = not measurable.

Grinding		Width at l	nalf height	Integral breadth		
(min)	<i>d</i> (001)	(001)	(002)	(001)	(002)	
0	7.225	0.18	0.17	0.26	0.21	
1	7.225	0.18	0.19	0.26	0.22	
10	7.237	0.20	0.21	0.30	0.25	
60	7.248	0.22	0.22	0.34	0.27	
120	***	***	***	***	***	

casionally, particles show a diffraction ring with a spacing of 2.57 Å.

Transmission electron microscopy. The TEM image in Figure 6 shows an example of the particle microcrystal morphology at the final stage of grinding. This micrograph reveals that serpentinite rock alteration occurs in two steps. First, the disintegration of the primary coarser particles produces ultrafine particles with diameters of 125–200 Å (arrows in Figure 6). Second, these particles adhere to each other to form large agglomerations (large particles in Figure 6). Mechanochemical reactions, such as the transformation of structural OH groups to adsorbed water, increased surface energy, and the formation of the amorphous material during grinding probably play an important role.

Analytical electron microscopy. The chemical analyses of starting antigorite and that ground for 10 and 120 min were obtained by AEM. A change of composition (Figure 7) affected the two major elements (Si and Mg). The change is related to both the extent of grinding and the particle crystallinity (amorphous vs. semicrystalline). Particles analysed after 10 min of grinding (dark circles) which exhibited an electron diffraction pattern similar to Figure 5b were affected by a drastic increase in the Si/Mg ratio. In fact, there is depletion of Mg by preferential destruction of the octahedral sheet. Normalizing of this composition on the



Figure 5. Changes in the electron diffraction patterns of: (a) starting antigorite in the (001) orientation, (b) semi-crystalline antigorite after 10 min of grinding.



Figure 6. Transmission electron micrograph showing the external morphology of the particles at the end of grinding (120 min).

basis of five O atoms and four (OH) produces Si > 2, indicating a lower negative charge, presumably related to (OH), would be necessary to maintain the maximum possible value (Si = 2). However, amorphous material at 120 min of grinding (open triangles) with no apparent diffraction by electrons nearly reproduced the composition of the initial antigorite, whereas that at 10 min of treatment (open circles) represents an intermediate.

Thermal analyses

The TG curve of the sample ground for 1 min shows two water losses at different temperature ranges



Figure 7. Representation in a ternary diagram [Al-Si-(Mg + Fe)] of the chemical data of antigorite at different stages of degradation (SM-0 = starting material, SC-10 = semicrystalline particles after 10 min of grinding, AM-10 and AM-120 are amorphous materials after 10 and 120 min of grinding, respectively).

(Figure 8). The low-temperature water loss, not detected at 0 min of grinding, occurs between 0-200°C and represents adsorbed water (the maximum loss is at 80°C). Structural water is lost between 600-800°C, with a maximum loss occurring at 748°C. The temperature corresponding to the maximum loss of structural OH decreased from 748°C (after 1 min of grinding) to 662°C (after 120 min of grinding). The TG curve of a sample after 120 min of grinding showed a completely different behavior, with a less marked difference between the two end-members (i.e., adsorbed water vs. structural water loss) and water loss becomes more evenly distributed. Weight loss of structural water also decreased significantly (Figure 9). In contrast, the weight loss of the adsorbed water increased (Figure 9) and the temperature corresponding to where maximum loss occurs increased from 80 to 155°C (Figure 8). The overall water loss increased slightly and the value corresponding to 120 min (11.85%) was very close to the theoretical value for pure antigorite (Caillère et al., 1982; Viti and Mellini, 1996).

Grain-size distribution

The variation in the particle-size fractions as a function of grinding time (Table 4) shows three stages: 1) grinding for 1 min increases the number of $\leq 2-\mu m$ particles, which reaches a maximum, and produces a relative decrease in particle sizes with diameters of 2–



Temperature °C

Figure 8. TG curves of starting and ground antigorite for times indicated.

10 μ m, as well as the elimination of 10–20- μ m particles; 2) grinding for 60 min results in a general decrease in the $\leq 2-\mu$ m fraction and a relative increase in the 2–10- μ m fraction and a small proportion (1.43%) of the 10–20- μ m size particles occurred again after 10 min of grinding; 3) at 120 min the $\leq 2-\mu$ m size particles increased and the 2–10- μ m fraction decreased once again.

DISCUSSION

The degradation of the Mulhacén antigorite by a vibration grinder is described as a series of parallel changes in the crystal structure and chemistry of the mineral. The increase in band intensities evident in the IR spectra between 0-1 min of grinding is explained by a decrease in particle size, producing an ideal IR spectrum and, therefore, increased transmittance-band intensities. Changes in the structural O-H and the Si-O vibrations may be used to monitor the relative resistance of the octahedral and tetrahedral sheets to grinding. The general decrease in intensity of the band at 3673 cm⁻¹, the disappearance of bands at 622 and 563 cm⁻¹, and the persistence of the transmittance bands at 985 and 447 cm⁻¹ all suggest that grinding affects the tetrahedral sheet to a lesser extent than the octahedral sheet.



Figure 9. Loss of water during heat treatment. Note that the scale varies for the time axis.

The chemical data support the results of the IR spectra in that grinding affects the tetrahedral sheet less than the octahedral sheet; Mg decreases after 10 min of grinding, Si increases, and Al remains constant. At the end of grinding, the composition of the resulting amorphous material is quite similar to that of the fresh antigorite. This suggests that early grinding causes the decrease of Mg by preferential destruction of the octahedral sheet, giving rise to a relative increase of Si in the particles. At the end of grinding, the resulting matrix consists of Mg Si-rich material similar in composition to the starting antigorite and corresponding to residual tetrahedral and octahedral sheets weakly bound within the system.

In Figure 2, the T3436/T3673 transmittance ratio decreased slightly at first due to the more ideal grain size and thus to the improved band of the structural O-H stretching vibration. The change in band characteristics at 3673 cm⁻¹ which is related to the structural O-H stretching vibration, and the broad band at 3436 cm⁻¹ which is related to the randomly-attached OH ions and adsorbed water, is clearly related to a change in the status of the hydroxyl groups. These groups are irreversibly displaced from their structural positions during grinding and remain in the system bound either as water molecules adsorbed on the degraded matrix or as OH groups linked by unsaturated broken bonds.

Table 4. Grain-size evolution during grinding.

Particle size (µm)	Grinding time (min)						
	0	1	10	60	120		
0–2	31.62	53.78	40.76	33.80	41.5		
2 - 10	68.00	46.24	57.84	64.87	58.50		
10 - 20	0.39	0.00	1.43	0.00	0.00		

From Figure 8, a close relationship appears to exist between the adsorbed and the structural water, which is consistent with the conclusion of the IR results. After the dehydroxylation of the mineral by dry grinding, the OH groups remained in the system bound to the degraded matrix either in the form of OH groups linked by unsaturated broken bonds or as adsorbedwater molecules. The weight loss corresponding to samples after 120 min of grinding (11.85%) is very close to the theoretical value of pure antigorite. Grinding seems to be an important process for water extraction, including structural water, thus accelerating the dehydroxylation process.

From the XRD results grinding in a vibration grinder disrupts the structure along the c axis more than within the (001) plane, since all basal reflections do not occur after 120 min, whereas the (16,0,1) reflection is well represented even after 120 min of grinding. This is confirmed by both electron diffraction and IR; some particles of the material ground for 120 min showed a diffuse ring at 2.57 Å, whereas the band at 1066 cm⁻¹, assigned to the Si-O stretching vibration with perpendicular polarization, did not occur. The general increase of the width at half-height and the integral breadth of the (001) and (002) reflections and the d(001)-values throughout grinding suggest that the vibration grinding destroys the crystal-structure of the mineral before complete amorphization. The structure of the semicrystalline particles differs from the starting material by forming "islands" in the octahedral sheet as a result of Mg and OH loss. Mg loss is obvious from the AEM results whereas the OH loss is expected to compensate the loss of positive charge from the octahedral sheet. Moreover, OH loss is also apparent from the IR (progressive increase in the transmittance ratio T3436/T3673) and TG (decrease in percentage of the structural-water loss) results. Although the octahedral "islands" produce faults in the periodicity of the layer stacking (c axis), the (001) plane would be maintained by the strong bonds existing between the Si tetrahedra.

In studies on kaolins, Juhász (1980) and Gonzales *et al.* (1991) found the specific surface area to increase initially upon grinding and to decrease as grinding continues. Papirer and Roland (1981), Henmi and Yoshinaga (1981), Cornejo and Hermosín (1988), and Pérez Rodriguez *et al.* (1988) reported similar behavior in chrysotile, imogolite, sepiolite, and pyrophyllite, respectively. Juhász (1980) referred to the increase in specific surface area as the "degree of dispersivity" and its decrease the "formation of aggregates by the compaction of the fine particles."

Combining TEM and grain-size results shows that dispersion and the breaking of coarser particles produces smaller but more numerous particles that produce high transmittance-band intensities in IR spectra. With continued grinding, the ultrafine particles tend to adhere to each other. Due to the increased surface energy, the action of the amorphous material, and probably the water adsorbed in the matrix formed by grinding (as indicated by the IR spectra and TG curves), larger agglomerates are formed.

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REFERENCES

- Aglietti, E.F., Porto Lopez, J.M., and Pereira, E. (1986) Mechanochemical effects in kaolinite grinding. *International Journal of Mineral Processing*, **16**, 125–146.
- Bailey, S.W. (1980) Structures of layer silicates. In Crystal Structure of Clay Minerals and Their X-ray Identification, G.W. Brindley and G. Brown, eds., Mineralogical Society, London, 1–123.
- Caillère, S., Hénin, S., and Rautureau, M. (1982) Minéralogie des argiles: II. Classification et nomenclature. Masson, Paris, 189 pp.
- Cícel, B. and Kranz, G. (1981) Mechanism of montmorillonite structure degradation by percussive grinding. *Clay Minerals*, 16, 151–162.
- Cornejo, J. and Hermosín, M.C. (1988) Structural alteration of sepiolite by dry grinding. *Clay Minerals*, **2**, 391–398.
- González, G.F., Ruiz Abrio, M.T., and González, R.M. (1991) Effects of dry grinding on two kaolins of different degrees of crystallinity. *Clay Minerals*, 26, 549–565.
- Halavay, J., Jónás, K., Elek, S., and Inczédy, J. (1977) Characterisation of the particle size and the crystallinity of certain minerals by infrared spectrophotometry and other instrumental methods: I. Investigation of clay minerals. *Clays* and *Clay Minerals*, 25, 451–456.
- Heller-Kallai, L., Yariv, S., and Gras, S. (1975). Hydroxyl stretching frequencies of serpentine minerals. *Mineralogi*cal Magazine, 40, 197–200.
- Henmi, T. and Yoshinaga, N. (1981) Alteration of imogolite by dry grinding. *Clay Minerals*, 16, 139–149.
- Juhász, A.Z. (1980) Mechano-chemical activation of kaolin minerals. Acta Mineralogica-Petrographica, 24, 121–145.
- Kisch, H.J. (1991) Illite crystallinity: Recommendations on sample preparation, X-ray diffraction settings and interlaboratory samples. *Journal of Metamorphic Geology*, 9, 665–670.
- Luce, R.W. (1971) Identification of serpentine varieties by infrared absorption. US Geological Survey Professional Paper, **750B**, 199–201.
- Nieto, F. (1997) Chemical composition of metapelitic chlorites: X-ray diffraction and optical property approach. *European Journal of Mineralogy*, **9**, 829–841.

- Papirer, E. and Roland, P. (1981). Grinding of chrysotile in hydrocarbons, alcohol and water. *Clays and Clay Minerals*, 29, 161–170.
- Pérez-Rodríguez, J.L., Madrid, L., and Sánchez-Soto, P.J. (1988) Effects of dry grinding on pyrophyllite. *Clay Minerals*, 23, 399–410.
- Sánchez-Navas, A., Martín-Algarra, A., and Nieto, F. (1998) Bacterially-mediated authigenesis of clays in phosphate stromatolites. *Sedimentology*, **45**, 519–533.
- Viti, C. (1997) Electron diffraction patterns of natural antigorites: A still unknown modulated crystal structure. In *Electron Crystallography*, D.L. Dorset, S. Hovmöller, and X.D. Zoo, eds., Kluwer Academic Publishers, The Netherlands, 419–422.
- Viti, C. and Mellini, M. (1996) Vein antigorites from Elba Island, Italy. European Journal of Mineralogy, 8, 423–434.
- Wilson, M.J. (1987) A Hand Book of Determinative Methods in Clay Mineralogy. Blackie, London, 308 pp.
- Yada, K. (1979) Microstructures of chrysotile and antigorite by high resolution electron microscopy. *Canadian Mineralogist*, **17**, 679–691.
- Yariv, S. and Heller-Kallai, L. (1975) The relationship between the IR spectra of serpentines and their structures. *Clays and Clay Minerals*, 23, 145–152.

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