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South Dakota Soils : A Generalized Soils Map of the East River Area of South Dakota

F. C. Westin

G. J. Buntley

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Agricultural Experiment Station

and Cooperative Extension Service

Soils

outh Dakota

F.S. 134A

A Generalized Soils Map of the East River Area of South Dakota

F. C. Westin and G. J. Buntley

Soil Survey Series No. 5, South Dakota Soils, consists of three separate fact sheets, 134A, 134B, and 134C. F.S. 134A. A Generalized Soils Map of the

East River Area of South Dakota.

F.S. 134B. A Generalized Soils Map of the West River Area of South Dakota.

F.S. 134C. The Distribution and Average Yields for Crops Commonly Grown in South Dakota.

South Dakota Soil Forming Factors

The kind of soil that develops in any area is the result of the interaction of five soil forming factors-climate, vegetation or organisms, parent material, relief, and time.

Climate controls the distribution of vegetation. To-gether climate and vegetation often are called the active factors of soil formation. This is because on gently undulating topography within a certain climatic and vegetative zone a characteristic or climax soil will develop unless parent material differences are great. Thus Chernozem soils develop across a great variety of parent materials.

The factor of parent materials exerts its influence on soils principally by determining their texture and to a great extent their mineralogical composition. Whereas climate and vegetation tell what group a soil is in, parent material, to a large, extent, determines its series. For example, Chernozem soils from glacial till are classed in the Vienna series, while Chernozem soils developed in thin loess overlying glacial till are classed in the Kranzburg series.

The factor of relief exerts its greatest influence by determining what drainage a soil will have. Steep slopes have excessively drained, thin soils; flat or depressed areas usually have poorly drained, thick soils. Mature Chernozem or Chestnut soils develop only on undulating relief where climate and vegetation are given full expression.

The factor of time in soil formation can be illustrated by comparing a soil on a flood plain which receives annual increments of alluvium with a soil on a terrace. The former is without horizons although it may have strata of contrasting alluvium, while the latter usually has an ABC horizon sequence. Actually much work s to be done in South Dakota for assessing the im-

Details of these five soil forming factors as they affect soil development in South Dakota will now be considered.

Climate

South Dakota, because of its inland position, has a continental climate with extremes of summer heat, winter cold, and rapid fluctuations of temperature. Annual precipitation ranges from 24 to 25 inches in the southeastern part to about 14 inches in the northwestern part of the state. Most of the precipitation comes in the g and early summer. The fall, winter, and spring ture falls mostly as frontal precipitation and is the result of condensation as warm moist air from the Gulf of Mexico overrides heavier polar air masses. Much of the summer precipitation comes as short hard showers of the convectional thundershower type. In eastern South Dakota, June normally has the most thunder-

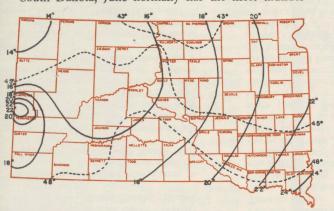


Figure 1. Average annual precipitation and air temperature for South Dakota.

Precipitation (Normal annual precipitation, 1921-50) Temperature Source: U. S. Weather Bureau Records

storms; in western South Dakota most of them normally come in July.

During the cold seasons winds are from the northwest, while they flow from the southeast during the warm season. The annual average surface wind velocity for the state is 10 to 12 miles an hour.

The average number of clear days per year is 120 to 140. There are 100 to 130 partly cloudy days per year on the average and 100 to 120 cloudy days. The normal annual number of hours of sunshine is about 2,850 in the southwestern part of the state to 2,700 in the northeastern part.

Average depth of frost penetration ranges from about

25 inches in the southwestern part to 50 inches in the northeastern part of the state.

The average number of days without killing frost varies from 130 days along the northern part of the state to 160 days in the southeastern part. The Black Hills area generally has shorter growing seasons than the rest of the state, with the average number of frostfree days ranging from 110 to 130 days.

It is possible to evaluate the climatic factor in South Dakota in relation to the climate of the United States by classifying the climate according to a national scheme. Several schemes may be used, but the system of Thornthwaite (5) classifies the climate of South Dakota as follows: Moist subhumid-southeast South Dakota: Dry subhumid-eastern and south central South Dakota: Semiarid-western South Dakota except for Black Hills. The Black Hills climate ranges from dry subhumid to humid.

The annual precipitation and temperature characteristics of South Dakota are shown in figure 1.

Native Vegetation

Except for the Black Hills, which are timbered, and the river valleys where trees and brush grew, the native vegetation of South Dakota was grassland.

Starting with the eastern border of the state and extending to the eastern edge of the James Valley, the principal association was one of tall grasses. Big bluestem, sand dropseed, and switchgrass were present along with upland and lowland forbs.

Moving westward across the James Valley, the tall grasses gradually dropped out, being found only on sandy soils and on cool northern exposures and the medium and short grasses assumed dominancy. Important species of the midland area were needleandthread. green needlegrass, western wheatgrass, slender wheatgrass, blue grama, prairie junegrass, and buffalograss.

Moving into western South Dakota, due to decreased rainfall, the shorter grasses largely replaced the midgrass species. Here were found blue grama, needleandthread, western wheatgrass, prairie junegrass, and little bluestem.

Certain variations in this general pattern occurred in western South Dakota as a result of extremely sandy or clayey soil texture. For example on the Pierre plain, an area of clay soils, the principal association was one of western wheatgrass, blue grama grass, and buffalograss. In the sand hills of southwestern South Dakota an important association was little bluestem, prairie sandreed, and needleandthread.

Parent Material

The kinds of soil parent material in South Dakota are shown in figure 2. This map was generalized from maps by Flint and Rothrock.

As this map illustrates, there is a large variety of materials from which the soils of the state have developed. They include ancient crystalline rocks in the central Black Hills, sedimentary rocks including shale, sandstone, and limestone in western South Dakota, and glacial materials of several ages in eastern South Dakota. Additional parent materials include loess and alluvial and colluvial materials formed from upland deposits. As soil development is extremely slow on crystalline rocks and as few soils have been recognized to date as having been developed from them in South Dakota, they will not be further discussed.

Sedimentary rocks. Sedimentary rocks have formed by consolidation and cementation of sand, silt, clay and other clastic material, and the precipitation from solution of the carbonates of calicum and magnesium. All of this took place on the floors of ancient seas.

The sands formed sandstone, the silts and clays formed siltstone and shale, and the basic carbonates formed limestone. Few of these rocks in South Dakota are pure -instead they are calcareous sandstones, sandy limestones, and so on. The principal sedimentary rock parent materials include: 1) the Pierre shale of the central part of the West River area; 2) the upper Cretaceous sandstones and sandy shales of the northern part of the West River area; and 3) the Tertiary sandstones and siltstones of the southern part of the West River area.

The Pierre shale area is sometimes called the "gumbo region" because of the plastic clay which weathers from the shale. The strata of the Pierre shale are soft and easily eroded. They generally are not butte-formers, rather they weather into soft rounded hills and ridges with convex tops. The Pierre clay and Lismas clay are soil series developed from Pierre shale.

Upper Cretaceous sandstones and sandy shales of northwestern South Dakota give rise to a great variety of soil textures. The sandstones weather to give rise to sandy soils and the shales which contain admixtures of silts and sands are parent materials for sandy loams, loams clay loams, silty clay loams, silty clays, and clays. The dominant textures are sandy loams and loams.

Morton loam and Vebar sandy loam are soil types mapped in this area.

Tertiary sandstones and siltstones of southwestern and south central South Dakota give rise principally to

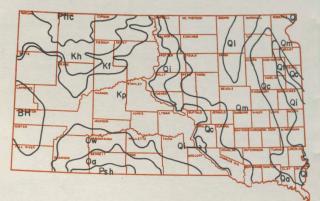


Figure 2. Soil parent materials in South Dakota. West River

Psh Sand Hills

- Oa Oligocene—Arickaree—sandstone and siltstone Ow White River beds—silts and clays Pflc Cannoball, Ludlow, Fort Union undifferentiated
- Kh Hell Creek-sandy shales

Kf Fox Hills sandston

Oa Alluvium

QI Lake Basin—silts, clays, sands QM Glacial till Qc Silty Glacial Materials

Qt Glacial Till and Thin Loess Qi—Thick loess in south, thin loess over loam till in north Source: Flint USGS Prof. Paper 262 and Petsch, B. C., Geologic

sandy and silty soils. The sandy materials on the south are an extension of the Nebraska Sand Hills. Going north the materials progressively have higher silt and clay increments. Some of the strata in this area and also some of those of the Upper Cretaceous area in northwestern South Dakota form benches, plateaus, and buttes because they are more resistant than other associated beds to weathering, mass wasting, and stream cutting.

Pleistocene geology. Pleistocene means "most recent" and this name is given to the events of glaciation. Glacier ice entered the state from the northeast or north and flowed south and west, the western margin of glaciation being the Missouri River. As the ice moved over the preglacial surface it filled valleys, planed off hills, forced the cutting of new valleys, piled up large ridges, and otherwise changed the preglacial topography.

The character of the rocks of the preglacial surface determined to a large extent, the composition of the glacial deposits formed from them. This is because most of the glacial deposits consist of altered rocks of local ori-

Glacial deposits cover South Dakota east of the Missouri River. Geologic evidence consisting of a boulder line of glacial stones shows that an early glacier pushed into western South Dakota 30 to 50 miles west of the present course of the Missouri River. Subsequent geologic erosion has removed this West River glacial drift except for scattered boulders, and a few isolated patches of glacial till, and the soil parent materials in this belt principally are sedimentary in origin.

Glacial deposits are divided physically into four groups: till, outwash, glacial lake deposits, and ice-contact stratified drift. All are present in eastern South Dakota as soil parent materials.

Till, which is the most abundant, is a mixture of all size particles. It is thought to have been deposited from the under part of the flowing ice. Barnes, Houdek, and Vienna are soil series developed from till.

Outwash was deposited by melt water as it flowed away from the ice and consists principally of mixed gravel and sand, usually crossbedded. Ordinarily the outwash material is overlain by alluvium, as in the case of the Fordville soils; or loess, as in the case of the Estelline soils

Glacial lake deposits, called lacustrine materials, consist of parallel-bedded silt and clay with small admixtures of sand. They were formed in depressions or basins temporarily blocked by glaciers and filled with ponded water. The Beotia, Harmony, and Aberdeen are soil series developed in these deposits.

Ice-contact stratified drift accumulated upon or against melting glacier ice. It occurs as knobs or small convex hills usually in rough terrain. In South Dakota the Sioux series, which is associated with the Buse series on the soil association map, is an example.

Although several older glaciers traversed eastern South Dakota, glacial drift left by them has been largely obliterated or covered by drift deposited by more recent ice sheets.

Loess and other wind-deposited sediments. Loess is a nonglacial deposit of wind blown and deposited particles of silt size. The loess in South Dakota came from mixing of silt from nonglacial deposits to the west with silt which was blown out or deflated from outwash bodies and even from the till itself as the glaciers melted The loess may consist of thin veneers to deposits 30 or more feet in thickness.

Strictly speaking, loess refers to particles of silt size. Sandy and silty clay loam materials, also carried and deposited by the wind, are called respectively eolian sand and eolian silty clay loam. They also are important South Dakota soil parent materials. The distribution of these wind-deposited sediments is shown in figure 2 along with the other soil parent materials of South Dakota.

Alluvium. Alluvium consists of stream-laid deposits of gravel, sand, silt, and clay, generally interbedded and almost always mixed. Generally, the alluvium of the West River is clayey in texture while that of the East River is mostly loamy.

RELIEF AND THE PHYSICAL DIVISIONS

Relief. Relief, as used here, refers to the lay-of-theland. It may be level, undulating, rolling, hilly, rough broken, or mountainous. It may be smooth with a network of small streams, or it may be choppy with many closed basins dotting the landscape. Relief usually varies from acre to acre so it is difficult to show on smallscale soil association maps. However any particular area is usually dominated by certain relief characteristics. For example, most of the Black Hills area is mountainous while most of the glacial Lake Dakota Plain in Spink and Brown Counties is level.

Physical divisions of South Dakota. This section and the map (figure 3) are devoted to a description of the natural land forms of South Dakota as classified by Fenneman and Rothrock and revised by Flint. The st tral Lowland from the Missouri Plateau. As can be noted from the soil association map, this line and the line separating the Chernozem area from the Chestnut area coincide over the northern two-thirds of the state.

The Minnesota River-Red River Lowland (Division 1, figure 3) is a broad, gently undulating, valley-like area with an elevation of 900 to 1,100 feet above sea level. Browns Valley, Minnesota, located midway between Lake Traverse and Big Stone Lake, is the continental divide between drainage to the Artic Ocean and to the Gulf of Mexico. The northeastern slope of the O des Prairies rises sharply nearly 1,000 feet to form the western limit of this lowland.

The Coteau des Prairies (Division 2) is a highland area between the Minnesota-Red River Lowland and



Figure 3. Physical divisions of South Dakota.

- 1. Minnesota River—Red River Lowland 2. Coteau des Prairies 3. James River Lowland
- Lake Dakota Plain
- . James River Highland . Coteau du Missouri
- . Missouri River Trench 8. Northern Plateaus 9. Pierre Hills
- 10. Black Hills
- 11. Southern Plateaus 12. Sand Hills

Source: Flint, R. F. Prof. Paper 262, USGS

the James River Lowland to the west. It slopes gently to the south and west. Its eastern and western slopes are steep at the northern end and taper off on the south. Elevations range from 2,000 feet above sea level on the north to about 1,600 on the south. It is drained to the south by the Big Sioux River, whose tributary streams enter mainly from the east. West of the Big Sioux River. the surface of the Coteau is dotted with lakes, while very few lakes occur east of the river.

Kp Pierre shale BH Black Hills—undifferentiated crystalline metamorphic and sedementary materials East River

Map of South Dakota. See also recent Geological Map of East-ern South Dakota published by State Geological Survey for correlation of the glacial drift sheets.

The James River Lowland (Division 3) is a gently undulating plain lying considerably lower than the Coteau des Prairies on the east and the Coteau du Missouri on the west. The James River drains through the area from north to south and occupies a rather narrow steep sided valley. Elevations range from 1,300 to 1,400 feet above sea level.

The Lake Dakota Plain (Division 4) is the nearly level surface formed by deposition of sediment when glacial Lake Dakota was ponded with water. The area is sandy at the northern end and of a silty clay loam and silty clay texture elsewhere.

The James River Highlands (Division 5) consist of a group of three ridges located at the southern end of the James River Lowland. They are remnants of former stream divides. From east to west, these highlands are Turkey Ridge, James Ridge, and Yankton Ridge.

Turkey Ridge, the largest of the three, is more than 40 miles long, 10 miles wide, and stands more than 300 feet higher than the surrounding country. Below the mantle of glacial drift is bedrock consisting of the Niobrara chalk, overlain in places by the Sharon Springs member of the Pierre shale. These strata are exposed in the Canyon of Turkey Creek; other exposures are rare, as the drift mantle is 30 to 200 feet thick

Yankton Ridge forms the northern bluff of the Missouri Valley from Yankton westward for 16 miles. Below the drift its core is Niobrara chalk, overlain by Pierre shale.

James Ridge, located west of the James River a few miles above its mouth, is only 9 miles long, 11/2 miles wide, and 100 to 260 fett high. Like the other two ridges, it is underlain by shale over chalk.

The Coteau du Missouri (Division 6) is part of the Missouri Plateau of the Great Plains province, separated from the main body of the Missouri Plateau by the Missouri River. This highland area is covered with glacial deposits and underlain by Pierre shale and older formations. Several broad shallow sags traverse the coteau, which marks the positions of former stream val-For eastern continuations of the Grand, Moreau, Jenne, Bad, and White Rivers. These sags are shown in plate 7 of USGS Professional Paper 262.

The Missouri River Trench (Division 7) averages a little over a mile in width with the valley floor 300 to 600 feet below the tops of the steep dissected bluffs. The river flows south-southeast with a gradient of about a foot per mile. Erosion and deposition are believed to be in equilibrium. As early travelers to the region reported the water to be turbid, rapid erosion apparently was in progress before the advent of agriculture, although cultivation in the tributary region certainly has d to the sediment load.

The Northern Plateaus (Division 8) is a series of plateaus and isolated buttes underlain by the Fox Hills sandstone and younger Cretaceous strata. It ranges in elevation from 2,000 to over 3,000 feet above sea level.

Pierre Hills (Division 9) consist of a series of smooth hills and ridges with rounded tops. The region is underlain by the Pierre shale formation and has lower elevations (1,800 to 2,800 feet) than the plateau country which rims it to the north and south.

The Black Hills (Division 10) is a region of mountainous terrain consisting of a series of turned up sedimentary strata, called hogbacks, arranged concentrically around a core of ancient crystalline rocks. Elevations range from 3,200 feet to about 7,000 feet.

The Southern Plateaus (Division 11) are divided into two regions. The large area to the southwest consists of a series of benches and buttes, underlain by Tertiary sandstones, siltstones, and shale. Elevations range from 2,800 to 3,600 feet. The Badlands comprise the northwestern part. The second area occurs in southeastern South Dakota principally in Union County. This is a stream-dissected highland underlain by a thick mantle of loess. Elevations range from 1,200 to 1,500 feet.

The Sand Hills (Division 12) is an extension of the Sand Hills region of Nebraska. It consists of a series of rounded hills interspersed with low, swampy areas, the whole region being underlain by eolian sand. Elevations range from 3,00 to 3,600 feet.

Time

Time is important in the formation of a soil. If the materials are easily eroded by wind and water, as in the case of the Pierre shale, the soil of steep slopes is destroyed almost as fast as it is formed. On undulating topography, soil formation on these materials and erosion go on at about the same pace. On flat slopes, due to the grass root mat which retards destructive processes. deeper soils develop which relatively are older from the standpoint of soil formation, than are the undulating and rolling soils. Thus the time factor is relative and varies across materials of the same geological age. Much research needs to be done in South Dakota evaluating the precise role of time as a soil forming factor.

Black silty clay loams and loams of the subhumid grasslands 34. Deep, friable, well to excessively drained silty clay loams from loess on nearly level to rolling upland (mostly Moody, Crof-ton, Nora and Trent series) mp Crook 35. Deep, friable, well and moderately welldrained silty clay loams from loess over glacial till on sloping upland (mostly Kranzburg and Brookings series)

36. Deep, friable, well and moderately well-

37. Deep, friable well-drained silty soils from

Hidewood and Brookings series)39. Deep, friable, well-drained, poorly drained and excessively drained silty soils on

rolling uplands consisting of hills inter-

spersed with depressions, sloughs and

lakes of all sizes (Most probable series-

Poinsett, Parnell, Sinai and Kranzburg)

loamy soils from loess over eolian sand on gently rolling uplands (Mostly Flan-dreau, Egeland and Maddock series)

40. Deep, friable, well-drained silty and

41. Deep, friable, well-drained silty soils

42. Deep, friable to firm, well to moderately

43. Moderately deep to shallow; firm to hard, imperfectly drained, silty clay loams and

Bend and Eckman series)

from loess over outwash gravel on nearly

level terraces and plains (Mostly Estel-

well-drained silty soils from lacustrine deposits on nearly level glacial lake

silty clays with claypans developed from

lacustrine deposits on level glacial lake plains (Mostly Aberdeen, Harmony and

plains (Mostly Beotia, Harmony, Grea

loess over glacial till and clay loams from glacial till on undulating to rolling upland (Mostly Kranzburg and Beadle

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drained silty clay loams from loess over

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- 38. Deep, friable, well, moderately well, and somewhat poorly drained silty soils on nearly level uplands (Mostly Kranzburg,

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- Exline series) 44. Deep friable well-drained silty soils of sloping plains associated with thin fri-able loamy soils occuring in drainage BLACK valleys (Mostly Beotia and Buse series) 45. Deep, friable, well-drained silty soils asso-ciated with deep, friable, poorly drained
 - highly calcareous silty soils; both on nearly level lacustrine plains (Mostly Overly and Bearden series) 46. Deep, friable, well-drained silty clay loams associated with moderately deep imperfectly drained claypan soils on an undulating to rolling plain (Most prob
 - able series-Reliance and DeGrey) 47. Deep, friable, well-drained silty clay loams associated with poorly drained soils on an undulating plain (Most prob-able series-Reliance and Tetonka)
 - Black and very dark grayish brown loams and clay loams of the subhumid grasslands 48. Deep, friable, well-drained loams and
 - light clay loams, some with deep-lying claypans, on a gently undulating glacial plain (Most probable series-Bonilla,
 - Houdek, Vienna and Cresbard) 49. Deep, friable, well-drained loams and

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White Butte

light clay loams associated with claypan soils and poorly drained soils on an undulating glacial plain (Mostly Houdek Cresbard, Cavour, Tetonka and Miranda

- 50. Deep friable, well-drained loams and light clay loams associated with mostly shallow claypan soils on an undulating glacial plain (Mostly Houdek, Miranda Cavour, Tetonka and Cresbard series)
- 51. Deep, friable, well and moderately welldrained loams on a gently undulating glacial plain (Mostly Barnes, Svea and Buse series)
- 52. Deep, firm, well-drained clay loams and deep, friable well-drained loams associated with claypan soils on an undulat ing glacial plain (Mostly Beadle, Houdek and Cavour series) 53. Deep, firm, well-drained clay loams in
- association with claypan soils on a nearly level glacial plain (Mostly Beadle and
- Cavour series) 54. Deep friable well-drained loams and silt loams developed from local alluvium and
- Joans developed from local antivitit and glacial till on a sloping plain below hilly areas (Mostly Lane and Houdek series)
 55. Shallow firm claypan soils associated with shallow friable saline soils on a nearly level plain (Mostly Exline and Maple region) series)
- 56. Deep, friable, well-drained light loams developed from stratified materials on a nearly level plain (Mostly Hand soils)
- 57. Deep, friable, well-drained silty clay loams associated with moderately deep clay soils and poorly drained depressional soils on the rolling plain adjacent to Missouri River (Mostly Reliance, Pierre and Tetonka series)
- 58. Moder by shallow and shallow some-what sively drained loams over out-wash sand and gravel on nearly level to rolling plains (Mostly Renshaw and
- Sioux series) 59. Deep, friable, well and moderately welldrained clay loams on gently undulating glacial plains (Mostly Forman and Aas-60. Deep, friable, imperfectly drained highly
- calcareous silt loams of glacial lake plains (Mostly Bearden series)
- 61. Deep, friable, well and moderately welldrained bick surfaced, (due to worm activity bams associated with depres-sional soils of the undulating glacial plain (Mostly Singsaas, Oak Lake and Parnell series)
- 62 .Deep, friable to firm well-drained, clay loams containing shale chips and developed from glacial till having a high content of shale chips, of the strongly undulating glacial plain (Mostly Edgeley and Cresbard series)

Sandy soils of the Subhumid grasslands

- 63. Deep, friable to loose, excessively drained and somewhat poorly drained sandy loams and loamy sands of glacial outwash plains (Mostly Hecla, Hamar and Maddock series)
- 64. Wet sandy soils (No suitable series name presently available) Dark Grayish Brown, predominantly silty
- soils of the semiarid grasslands 65. Deep, friable, well-drained silt loams and

silty clay loams associated with deep, well-drained clay loams and poorly drained depressional soils on a gently undulating glacial plain (Mostly Agar, Raber and Tetonka series)

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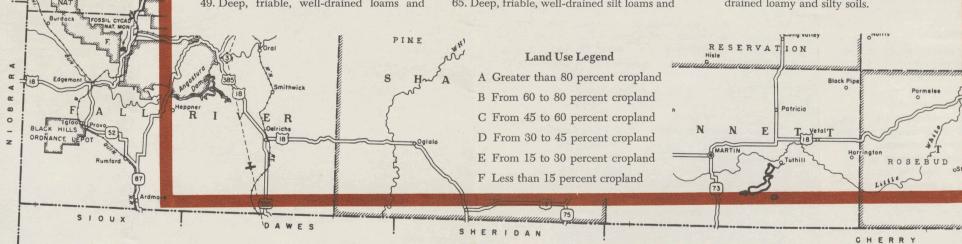
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66. Deep, friable, well-drained silty clav loams some moderately deep over glacial till, and clay loams associated with claypan soils on an undulating glacial plain (Mostly Eakin, Raber, Agar and Miranda series)

Dark Gravish Brown loams and clay loams of the semiarid glassland 67. Deep, friable, well-drained loams associ-

- ated with claypan soils and poorly drained soils on an undulating glacial plain. (Mostly Williams, Cavour, Cresbard and Tetonka series) 68. Deep, friable, well-drained loams associ-
- ated with poorly drained soils and claypan soils on a rolling glacial plain (Mostly Williams, Tetonka and Cavour series)
- 69. Deep, firm, well-drained clay loams asso-ciated with poorly drained soils and clay-pan soils on a strongly undulating glacial plain (Mostly Raber, Tetonka and Cavour series) 70. Moderately deep and shallow, friable,
- well and excessively drained loams, sandy loams and silt loams over sand and gravel on gently undulating glacial outwash plains (Probable series–Oahe, Akaska, Maddock)
- Clayey soils of the Semiarid grasslands 71. Deep and moderately deep, firm to hard, well and excessively drained clays on an undulating to rolling plain (Probable series-Pierre, Promise and Lismas) Hilly soils
- 72. Thin, friable, excessively drained loams associated with moderately thick, well-drained loams and poorly drained depressional soils, on hilly uplands in the Sub-humid Grasslands (Probable series-Buse, Barnes and Parnell)
- 73. Thin, friable, excessively drained loams associated with moderately thick, welldrained loams, poorly drained depressional soils and elaypan soils on hilly topog-raphy of the semiarid grasslands (Prob-able series—Zahl, Williams, Tetonka and Cavour)
- 74. Shallow, friable loams from sandstone and shale bedrock on hilly plains of the semiarid grassland (Probable series-Bainville)
- 75. Thin, firm to hard clays from shale on hilly slopes along the Missouri River Valley (Probable series-Lismas, Pierre)
- Soils from Alluvium 76. Soils from Missouri River Alluvium-mixed sandy, loamy, and clayey soils with the sands occurring mostly along the water course and the clays along the bluffs with the loamy soils in between
- 77. Soils from James River Alluvium mostly somewhat poorly drained loams and silt loams of bottomlands and low terraces 78. Soil from Big Sioux River Alluvium. Mix-
- ed soils grading in texture from sand to clay and in drainage from excessive to poor. Some low terraces on which occur loams over gravel.
- 79. Soils from Local Alluvium. Mostly deep, moderately well to somewhat poorly drained loamy and silty soils.



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