RESEARCH ARTICLE



Nitrogen isotope composition of ammonium in PM_{2.5} in the Xiamen, China: impact of non-agricultural ammonia

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Abstract

Since NH₃ is a significant precursor to ammonium in PM_{2.5} and contributes significantly to atmospheric nitrogen deposition but largely remains unregulated in China, the insight into the source of NH₃ emissions by the isotopic investigation is important in controlling NH₃ emissions. In this study, atmospheric concentrations of NH₃ and water-soluble ion composition in PM_{2.5} as well as nitrogen isotope ratios in NH₄⁺ (δ^{15} N-NH₄⁺) in Xiamen, China, were measured. Results showed that average NH₃ concentration for the five sites in Xiamen was 7.9 µg m⁻³ with distinct higher values in the warm season and lower values in the cold season, and PM_{2.5} concentration for the two sites (urban and suburban) was 59.2 µg m⁻³ with lowest values in summer. In the PM_{2.5}, NH₄⁺ concentrations were much lower than NH₃ and showed a stronger positive correlation with NO₃⁻ than that with SO₄²⁻ suggesting the formation of NH₄NO₃ and equilibrium between NH₃ and NH₄⁺. Although the concentrations of NH₃ at the urban site were significantly higher than those at the suburban site, no significant spatial difference in NH₄⁺ and δ^{15} N-NH₄⁺ was obtained. The distinct heavier δ^{15} N-NH₄⁺ values in summer than in other seasons correlated well with the equilibrium isotopic effects between NH₃ and NH₄⁺ which depend on temperature. The initial δ^{15} N-NH₃ values were in the range of waste treatment (-25.42‰) and fossil fuel combustion (-2.5‰) after accounting for the isotope fractionation. The stable isotope mixing model showed that fossil fuel–related NH₃ emissions (fossil fuel combustion and NH₃ slip) contributed more than 70% to aerosol NH₄⁺. This finding suggested that the reduction of NH₃ emissions from urban transportation and coal combustion should be a priority in the abatement of PM_{2.5} pollution in Xiamen.

Keywords Ammonia \cdot Ammonium \cdot PM_{2.5} \cdot Stable nitrogen isotope ratio \cdot Non-agricultural source

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Introduction

Ammonia (NH₃) is the most abundant alkaline gas in the atmosphere and can readily react with available sulfuric and nitric acid to generate ammonium sulfate, ammonium bisulfate, or ammonium nitrate which account for a large fraction of the fine particulate matter (PM_{2.5}) in many Chinese cities (Huang et al. 2014; Wu et al. 2015; Wang et al. 2015). The enhanced PM2.5 has been strongly related to the degradation of visibility, adverse health effects, and climate change. In order to mitigate PM2.5 pollution in China, the central government has established strict industrial emission regulations for SO₂ and NO_x (http://datacenter.mep.gov.cn). Nevertheless, recent studies revealed that the lack of NH₃ emissions control could weaken the effectiveness of PM_{2.5} pollution control based simply on SO2 and NOx emissions reduction (Wang et al. 2013; Zhang and Wu 2013; Backes et al. 2016; Fu et al. 2017). Our previous studies showed that the emission

densities of NH₃ in the coastal urban areas of Fujian province, China, were significantly higher than those in the inland rural areas due to the non-agricultural NH₃ from highly concentrated population and industries (Wu et al. 2017). In addition, NH₃ is estimated to make a great contribution to the total inorganic nitrogen dry deposition due to its rapid dry deposition velocity with respect to that of fine aerosol NH₄⁺ (Russell et al. 2003). The contribution of gaseous NH₃ to inorganic nitrogen deposition in Xiamen Bay, China, was estimated to be 27.4–28.2% which was much higher than that of aerosol NH₄⁺ (17.0–17.7%) (Wu et al. 2018). Therefore, the source apportionment of NH₃ would give the policymakers a great help with the future PM_{2.5} pollution control strategies and ecological safety.

It is well established that fertilizer application and livestock waste are the largest contributors of ammonia emissions on a regional scale while non-agricultural ammonia from humans, waste treatment, and fossil fuel-based emissions may be more important in the urban area (Wang et al. 2015, 2018; Chang et al. 2016; Kang et al. 2016; Wu et al. 2017). However, the low temporal-spatial resolution and high uncertainty of the emission inventory could not fulfill the needs of successful mitigation for NH₃. It has been suggested that the stable N isotope ratios $(^{15}N/^{14}N)$ of anthropogenic atmospheric N depend in part upon their emission sources (Kundu et al. 2010; Felix et al. 2013; Chang et al. 2016; Pan et al. 2016). Therefore, the stable nitrogen isotope ratio in aerosol ammonium (δ^{15} N-NH₄⁺) have been frequently used to investigate sources and processes of atmospheric ammonia (Yeatman et al. 2001; Kundu et al. 2010; Pan et al. 2016, 2018a, b; Wang et al. 2017; Park et al. 2018). Nevertheless, the δ^{15} N-NH₄⁺ values in aerosols may differ very much from those of the initial emitted δ^{15} N-NH₃ values due to the kinetic and equilibrium isotope fractionation takes place in the physicochemical process (e.g., adsorption, evaporation) (Heaton et al. 1997; Hayasaka et al. 2004; Kawashima and Kurahashi 2011; Savard et al. 2017). When NH₃ is in deficit, the δ^{15} N values in NH₄⁺ reflect their NH₃ sources (Jickells et al. 2003). In NH₃-rich atmosphere, however, the ¹⁵N/¹⁴N fractionation under equilibrium reaction of gaseous NH₃ to aerosol NH₄⁺ should generate higher δ^{15} N values in NH₄⁺ than in NH₃ (Heaton et al. 1997; Savard et al. 2017). An experimental $\rm NH_4^+\text{-}NH_3$ isotopic enrichment factor, $\epsilon_{\rm NH4}^{+}\text{-}_{\rm NH3}$ (= $[({}^{15}N/{}^{14}N)_{NH4}{}^{+}/({}^{15}N/{}^{14}N)_{NH3} - 1] \times 1000 \approx \delta^{15}N{}^{-}NH_{4}{}^{+} \delta^{15}$ N-NH₃), was estimated to be + 33% (Heaton et al. 1997), which was in fact close to the calculated difference between δ^{15} N-NH₄⁺ and δ^{15} N-NH₃ (from 19.3 to 48.7% with a mean of $30.4 \pm 8.2\%$) in Niigata Prefecture, Japan (Hayasaki et al. 2004). In addition, the measurements exhibited high mole concentration ratio of gaseous NH_3 to aerosol NH_4^+ (from 1.4 to 22.8 with a mean of 7.5 ± 5.8) confirming that the reaction of gaseous NH₃ to aerosol NH₄⁺ was in equilibrium.

Under the equilibrium conditions, both the isotopic enrichment factor and the phase distribution of NH_4^+/NH_3

contribute to the difference between δ^{15} N-NH₄⁺ and δ^{15} N-NH₃ (Heaton et al. 1997; Li et al. 2012). Assuming that gaseous NH₃ and aerosol NH₄⁺ are in the state of a well-mixed closed system and not influenced by primary particulate NH₄⁺, the δ^{15} N-NH₃ values can be estimated by using an isotopic mass balance as follows (Heaton et al. 1997):

$$\delta_{\rm NH3} = \delta_{\rm NH4+} - \varepsilon_{\rm NH4}^{+} -_{\rm NH3} \times (1-f) \tag{1}$$

where δ_{NH3} and $\delta_{\text{NH4+}}$ are the isotope ratios of the initial gaseous NH₃ and final aerosol NH₄⁺, respectively, and *f* is the fraction of the initial NH₃ gas converted to the aerosol NH₄⁺. Savard et al. (2017) found that the isotopic difference between δ^{15} N-NH₄⁺ and δ^{15} N-NH₃ was inversely correlated with the ambient temperature in central and southern Alberta, Canada. Based on the relationship between NH₄⁺_(aq)-NH_{3(g)} equilibrium constants and temperature from Urey (1947), Pan et al. (2018b) gave a mathematical regression of $\varepsilon_{\text{NH4}}^+$ -_{NH3} as a function of temperature (Kelvin):

$$\varepsilon_{\rm NH4}^{+}_{-\rm NH3} = 12.4678 \times 10^3 / \text{T} - 7.6694$$
 (2)

The modeling yielded ε_{NH4}^{+} -NH3 of + 34.17% at 25 °C which is in fact very close to the measured isotopic enrichment factor of +33% in Heaton et al. (1997). Based on the IsoSource isotopic mixing model (www.epa.gov), Pan et al. (2016, 2018a) calculated the δ^{15} N value of the initial NH₃ using $\varepsilon_{\rm NH4}^{+}$ -_{NH3} of + 33% and estimated that up to 90% of NH₃ was from fossil fuel emissions during hazy days in urban Beijing. However, the ambient NH₃ concentrations were simulated from WRF-CMAQ which likely led to the uncertainty of the f value. Similarly, the δ^{15} N-NH₃ values were calculated using NH₃ concentrations which were not measured during the PM2.5 sampling period in Seoul (Park et al. 2018). Chang et al. (2016) measured the δ^{15} N values of NH₃ collected by Ogawa passive sampler and estimated that the traffic, waste, livestock, and fertilizer contributed 20.4%, 25.9%, 24.0%, and 29.7%, respectively, to ambient NH₃ during the 2014 APEC summit (3–12 November) in Beijing with stringent air quality control measures. Wang et al. (2017) suggested that the equilibrium between NH₃ and NH₄⁺ was weak according to the low molecular ratios of NH₃ to NO₂ + SO₂ (~0.4) and NH₄⁺ to (NO₃⁻ + SO₄²⁻) (~ 1.1) in urban Beijing in January 2013 and the major contributors to total nitrogen in PM2.5 were estimated from coal combustion (41%), vehicle exhausts (30%), and domestic waste/sewage (14%) by SIAR (stable isotope analysis in R) model. The selection of major emission sources and atmospheric conditions (NH₃-rich or NH₃-poor) played an important role in the source apportionment of ambient NH_4^+/NH_3 (Heaton et al. 1997; Li et al. 2012).

It has been proved that NH_3 generated from the septic tanks and evaporated from the urban green space might influence the ambient NH_3 to a great extent in the urban area in China

(Chang et al. 2015; Teng et al. 2017). In order to meet the emission standards for nitrogen oxides (NO_x) from heavyduty diesel trucks in China, AdBlue (pure urea solution) was widely used in SCR (selective catalytic reduction)-equipped engines to convert NO_x into harmless N₂ in the exhaust stream. At some conditions, over-stoichiometric ammonia availability could result in NH₃ slip in a similar way to that from power plant using SCR denitration technology (Pan et al. 2016). In addition, gasoline and liquid petroleum gasoline engines equipped with a three-way catalytic converter (TWC) can produce NH₃ as a byproduct (Livingston et al. 2009; Behera et al. 2013). Furthermore, the use of TWC with low-sulfur gasoline has been proven to increase NH₃ emissions (Mejía-Centeno et al. 2007). Thus, NH₃ emissions from vehicles in China are expected to increase due to the rapid increase of vehicle numbers and the decrease of sulfur content in gasoline and diesel since 2018. Because the ambient NH₃ is subject to the emission sources, weather conditions, and land use types, the establishment of the localized isotopic signatures (δ^{15} N-NH₃) for the major NH₃ emissions sources is a very important step in the process of ambient NH₃/NH₄⁺ source apportionment.

Xiamen is a coastal city in southeastern China, beside the Taiwan Strait, with an area of 1700 km² and a population of 3.92 million according to the Xiamen Statistic Yearbook of 2017 (http://www.stats-xm.gov.cn/2017/). Besides, the number of tourists and vehicles and the amount of household waste and residential water consumption reached 67.7 million, 1.11 million, 1.74 million tons, and 211.1 million m³ in 2016, respectively. In the meantime, however, the amount of nitrogen fertilizer and fattening pigs in 2016 dropped 60% and 27% with respect to those in 2007, respectively, due to the urban expansion and population shift from a rural area to an urban area. The present population density in Xiamen Island has surpassed the densities in Hong Kong and Singapore, and there will be an estimated 5.0 million residents by 2020 according to the Xiamen City Urban Master Planning (2010-2020). Based on the emission inventory of NH₃ in Xiamen in 2015, humans (septic tanks and latrines), waste treatment (sewage treatment plant and landfill), and traffic contributed 92.4% of the total emission of NH₃ in Siming district in Xiamen Island, while the emission of NH₃ from fertilizer application and livestock waste contributed 81.3% of the total NH₃ emission in Xiang'an district (Wu et al. 2017). These statistics implied that the contribution of urban nonagricultural sources such as vehicle exhausts, septic tank exhausts, and waste treatment to ambient NH₃ might be highly significant at a local scale in the urban area of Xiamen.

In this study, we conducted a yearlong study of δ^{15} N-NH₄⁺ at two sites (urban and suburban) in Xiamen Bay from December 2016 to November 2017 using the fumigation-precipitation method. We also measured ambient NH₃ concentrations from June 2015 to June 2018 at five sites in

Xiamen Bay using Ogawa passive samplers, which could be used for the estimation of the initial δ^{15} N-NH₃ values. Several local emission sources were selected to collect NH₃ and aerosol NH₄⁺ for stable nitrogen isotope ratio analysis. In the end, we quantified the source contribution to aerosol NH₄⁺ based on the IsoSource isotopic mixing model. The results will provide useful information for the mitigation of non-agricultural sources of NH₃ and PM_{2.5} in urban regions in China.

Methods and materials

Ambient PM_{2.5} and NH₃ sampling

PM_{2.5} were sampled at an urban site located in the Siming campus of Xiamen University (SM, 24.45° N, 118.08° E) in the south of Xiamen Island, and a suburban site located in the Xiang'an campus of Xiamen University (XA, 24.62° N,118.23° E) to the northeast of Xiamen Island (Fig. 1). The urban site is situated between two busy roads about 50 m away from the coast and was selected as a typical coastal urban site. The daily traffic flow rate was around 45,800 on weekends and 18,400 on weekdays based on the field observation in October of 2017. The suburban site is surrounded by villages and scattered plots of cropland about 5 km away from the coast. The distance between the two campuses is around 30 km. The traffic volume near the suburban site is much lower than that near the urban site. PM2.5 samples were collected for 3 days per month (generally the fifth, fifteenth, and twenty-fifth day) using a high-volume sampler (TH1000C, 1.05 m³ min⁻¹, Wuhan Tianhong Instruments, Wuhan, China) adapted with a PM2.5 impactor. Pall quartz filters (25,000 QAT-UP, Pall Co.) for PM2.5 sampling were precombusted to 600 °C for 4 h to remove impurities. The sampling was conducted between December 2016 and November 2017. The concentrations of $PM_{2.5}$ (µg m⁻³) were determined according to the gravimetric difference before and after sampling and the sampling volume. The daily mean concentrations of criteria air pollutants such as SO₂ and NO₂ and meteorological factors including temperature (T), relative humidity (RH), pressure (P), wind speed (WS), wind direction (WD), visibility (Vis), and rainfall (RF) in Xiamen during the sampling periods were obtained from the websites of www.aqistudy.cn and https://rp5.ru, respectively.

Ogawa passive sampler was employed to monitor ambient ammonia concentration at SM, XA, and other three sites (ZZ, IUE, and DD) in the Xiamen Bay from June 2015 to June 2018 (Fig. 1). The Institute of Urban Environment (IUE) is close to two busy expressways, likely influenced by the vehicle exhausts. The Dadeng Middle School (DD) is located in Dadeng Island which is around 1 km to the marine reclamation land at the Xiamen New Airport under construction. The Zhangzhou campus of Xiamen University (ZZ) is



Fig. 1 Location of the $PM_{2.5}$ and NH_3 sampling site (star), power plant (rectangle), waste incineration plant (triangle), and sewage treatment plant (circle) in the Xiamen Bay, China, beside the Taiwan Strait. Approximately 330,000 and 42,000 people live in the inset regions of

SM and XA, respectively. SM, XA, and ZZ denote Siming, Xiang'an, and Zhangzhou campuses of Xiamen University, respectively. DD and IUE represent Dadeng Middle School and the Institute of Urban Environment, respectively

located to the south of Xiamen Island around 9 km away from SM. The Ogawa is a double-sided passive diffusion sampler equipped with two 14-mm citric acid-impregnated quartz filters by the manufacturer (www.ogawausa.com). The samplers were deployed for more than 3 weeks per month covering the sampling period of $PM_{2.5}$. The details of sampling sites, surrounding environment, and calculation procedure are available in Wu et al. (2018).

In addition to the ambient NH₃ sampling, for the collection of NH₃ emissions from waste treatment (sewage treatment station, trash transfer station, and public toilet), fertilizer application (urea and organic fertilizers), and biomass burning (rice straw), we used a two-stage filter pack system to sequentially trap particulate NH₄⁺ on a PTFE filter (2.0- μ m pore size, 30-mm diameter) and gaseous NH₃ on a citrated cellulose filter (2- μ m pore size, 30-mm diameter) at a flow rate of 10 L/min. To minimize the interference of ambient atmospheric NH₃/NH₄⁺ and isotope fractionation during the sampling period, the evaporated NH₃ was trapped in a semi-closed chamber (0.4 × 0.4 m × 1.0 m) with a hole on the roof for setting sampling inlet. The amount of nitrogen applied for urea and organic fertilizers volatilization

experiments was around 100 kg N hm⁻², and the volatilized NH₃ was collected on the first, second, and seventh day. For the biomass burning experiment, the smoke particle and gas NH₃ were both sampled for ¹⁵N analysis. The confined chamber was first cleaned using a vacuum pump for around 10 min. After waiting for another 10 min until the odors fill the entire chamber, the sampler was turned on to collect NH₃/NH₄⁺ for 30 min (Chang et al. 2016).

Chemical and isotope analysis

A punch of quartz filter for PM_{2.5} samples (ϕ 20 mm) and Ogawa NH₃ collection pads were ultrasonically extracted using Milli-Q water (specific resistivity ~ 18.2 M Ω cm). The leachates were analyzed with ion chromatography (PIC-10, Qingdao Puren Instruments Co., China) for NO₃⁻⁻, SO₄²⁻, and NH₄⁺ (Wu et al. 2015). The NH₃ concentration was calculated according to the detected amount of NH₄⁺ in the collection pad, exposed time (min), average ambient temperature (°C), and concentration conversion coefficient (43.8 ppb min/ ng at 20 °C for both pads extracted as one) (https://ogawausa.

com/proctocols/; Wu et al. 2017). A calibration curve was evaluated for each sequence using the standard solutions. A random replicate check was carried out every 10 samples, and the precision fell within 10%. Procedure blanks and recover rates were used to correct the concentrations.

The method for the isolation of ammonium in PM_{2.5} was derived from the HCl fumigation method in Kundu et al. (2010) and Howa et al. (2014). The punch of quartz filter (75-mm diameter) was exposed to HCl fuming in a desiccator overnight to remove NO₃⁻. The effectiveness of the approach was evaluated in terms of the remained NH₄⁺ and NO₃⁻ on the filters after HCl fumigation. The extracts of the remained NH₄⁺ were filtered to a borosilicate test tube. For the gaseous NH₃ samples, the filter was extracted without HCl fumigation. Excess powered sodium tetraphenylborate (the solubility at 25 °C is 470 g/L) was added to the ammonium salt solutions to ensure complete precipitation of the ammonium ions (Howa et al. 2014). The low solubility of ammonium tetraphenylborate in water (the solubility at 25 °C is 9.7 mg/L) means that very little ammonium remains dissolved form. The precipitates were filtered using a GF/F membrane filter (25-mm diameter, Whatman) and then washed with 5 mL Milli-Q water. The precipitates on the filter papers were allowed to dry for 4.5 h at 55 °C in a vacuum desiccator and then placed in a tin cup for elemental analyzer/isotope ratio mass spectrometer (EA/IRMS) (vario EL cube EA + IsoPrime100, Elementar) to determine nitrogen isotope ratio $({}^{15}N/{}^{14}N)$. The tin cups were precleaned with acetone to remove organic contaminants prior to use. It should be noted that the preparation method used in this study for isolating the NH_4^+ fraction of the mixture extract solution might not be interfered by organic nitrogen. The tetraphenylborate isolation method was proved to improve the discriminatory power over bulk analysis (Howa et al. 2014). The solid ammonium tetraphenylborate was oxidized in the combustion column packed with copper oxide at 960 °C, and then the derived NO2 was introduced into the reduction column to reduce to N2 at 600 °C. The resultant N2 gas was purified and transferred to IRMS for isotope measurement. The isotopic composition of N was reported in parts per thousand relative to atmospheric N_2 as follows:

$$\delta^{15}N(\%) = \left[\left({^{15}N/^{14}N} \right)_{\text{sample}} / \left({^{15}N/^{14}N} \right)_{N2} - 1 \right] \times 1000(3)$$

where the absolute $^{15}\text{N}/^{14}\text{N}$ value of N_2 is $3.6765\times 10^{-3}.$ A standard with a $\delta^{15}\text{N}$ value of 1.18% from Indiana University was used as a standard to calculate $\delta^{15}\text{N}$ ratios. The average analytical error for $\delta^{15}\text{N}$ measurements based on repeated duplicate analysis of the acetanilide standard was 0.25%.

IsoSource mixing model for source identification

Based on "bottom-up" methods, previous study has indicated that non-agricultural emission sources (humans, waste treatment, and vehicle traffic) contribute around 92% of the total NH₃ emissions in the urban district while agricultural emission sources (fertilizer application and livestock waste) contribute around 81% of the total NH₃ emissions in the suburban district in 2015 (Wu et al. 2017). Considering the surrounding environment at the sampling sites and the prior investigation, the ambient NH₃ near the two sites in Xiamen during the study period have been shown to originate from four major sources: fossil fuel combustion (coal combustion and vehicle exhausts), NH₃ slip (derived from urea that is used in SCR-equipped diesel engines and power plants), N fertilizer application, and human excreta.

The stable nitrogen isotopic composition of a mixture can be used to determine the contribution of different sources using mathematical mixing models. In this study, we used the IsoSource model (version 1.3.1, www.epa.gov) to calculate ranges of source proportional contributions to the initial NH₃. The model solves iteratively for feasible mixing solutions and has been recently used for atmospheric nitrogen source apportionment (Chang et al. 2016; Pan et al. 2016; Wang et al. 2017; Ti et al. 2018). This model addresses every possible combination of source proportions (summing to 100%) incrementally (e.g., 1%), then calculates the predicted isotope value for each combination using linear mass balance equations. These predicted values are then examined to determine which ones fall within some tolerance range (typically 0. 1%) of the observed receptor isotope value, and all of these feasible solutions are recorded.

In this study, the measured δ^{15} N-NH₃ value from volatilized organic fertilizer (-21.34%) was used as end member δ^{15} N-NH₃ value for fertilization because this kind of organic fertilizer was widely used for campus lawn maintenance and vegetables grown near the sampling sites. This value was in the range of reported δ^{15} N-NH₃ values from rice and wheat field in Taihu lake region, China (Ti et al. 2018), but higher that of the NH₃ volatilization from cornfield in the USA (Felix et al. 2013) and urea in laboratory (Change et al. 2016) (Table S1). Because the sewage treatment plants and trash transfer stations were far away from the two sampling sites, the measured δ^{15} N-NH₃ values from campus public toilet (– 25.42‰) were used as endmember δ^{15} N-NH₃ value for human excreta. The δ^{15} N-NH₃ values of NH₃ slip (-12.95‰) were obtained from the measurements in SCR-equipped power plant emissions in Felix et al. (2013) due to the lack of data in China (Pan et al. 2016; Ti et al. 2018). As seen from the Table S1, the δ^{15} N-NH₃ values in vehicle emissions (-4.6 to -2.2% inside tunnel) obtained by Felix et al. (2013) were much heavier than those of -17.8 to -9.6% in Chang et al. (2016) but were within the range of coal combustion (-7 to)2%; Freyer 1978). Thus, the median value of the range of fossil fuel combustion (-2.5%) from the US-based study was used as the endmember δ^{15} N-NH₃ value, which can be distinguished from other emissions. The initial δ^{15} N-NH₃ was

calculated from δ^{15} N-NH₄⁺ according to Eq. (1) because of the isotope exchange equilibrium between NH₃ and NH₄⁺ in air (Pan et al. 2016, 2018a).

HYSPLIT backward trajectories

The Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model (https://ready.arl.noaa.gov/) has been widely used to trace the origin of air masses and establish the relationships between source and receptor (Kundu et al. 2010; Beyn et al. 2015). In this study, the backward trajectories were calculated every 6 h with a travel time of 72 h and at 100 m above sea level for each day during the sampling period. Meteorological data were used from the NOAA GDAS (Global Data Assimilation System) database (Su et al. 2015). All the trajectories in each season were separately subjected to cluster analysis using the cluster algorithm of HYSPLIT model. Previous studies have showed that the backward trajectories could be influenced by the starting height (Koracin et al. 2011; Stein et al. 2015). Therefore, the HYSPLIT model was run for starting height at 100 m, 500 m, and 1000 m to investigate the surface influence (Fig. S1). No significant difference was found for the backward trajectory cluster at different starting height. One possible reason for the similar trajectory clusters was that the air mass mainly traveled along the coastal region or from the ocean. The air masses originated from Yangtze River Delta urban agglomerations could carry the coastal urban air pollutants to the sampling sites, while the air masses originated from the South China Sea were clean and could dilute the air pollutants to a great extent.

Results and discussion

Seasonal variation of NH₃ concentration

The annual mean NH₃ concentration for all sites was $7.90 \pm$ 3.79 μ g m⁻³ with a range of 1.80 to 22.44 μ g m⁻³ (*n* = 185) (Table S2). There were significant positive correlations for the NH₃ concentrations between the five sites (r = 0.358 - 0.691, p < 0.05) (Table S3). Generally, the concentrations of NH₃ exhibited higher values in warm season and lower values in cold season (Fig. 2). Significant correlations between NH₃ levels and monthly average site temperature (T) were also observed (r = 0.345 - 0.535, p < 0.05, n = 37), likely due to the increased NH₃ emissions from soil, fertilization, waste treatment, and green space as well as the dissociation of NH₄NO₃ aerosol brought about by increased temperature (Yeatman et al. 2001; Phan et al. 2013; Hu et al. 2014). However, NH₃ concentration had no correlation with rainfall (RF) and wind speed (WS). In urban Shanghai, there was a highly significant relationship between hourly average WS and NH₃ concentration (Chang et al. 2016). In addition, there were clear morning and evening peaks of NH₃ levels in urban area which were consistent with morning and evening rush hour and mixing layer variations (Wang et al. 2015, 2016). In this study, the triweekly time resolution might cloud the influence of WS and RF on NH₃ concentrations. Therefore, industrial and traffic emissions which are insensitive to temperature could not be the major factor influencing the seasonal pattern of ambient NH₃. Similar temperature dependent patterns have also been reported in urban Shanghai, Beijing, Zhengzhou and Taihu Lake region, China (Wang et al. 2015, 2018; Xu et al. 2016; Ti et al. 2018), and downtown Toronto, Canada (Hu et al. 2014). Notable decrease of NH_4^+ and $NO_3^$ in PM2.5 was observed in the warm season from June to October in Fig. 3 further supporting the effect of temperature on ambient NH₃. Of the five sites, the IUE had significant higher NH₃ concentrations than those at others based on paired-samples t tests (p < 0.001, 2-tailed) (Table S4). Because the reclaimed water from the sewage treatment station was applied to the landscape in the office park of IUE, the high levels of NH₃ at this site was probably influenced by NH₃ from both the landscape water body and the vehicular exhausts. Sewage treatment plant has been reported to be an important source of ambient NH₃ in an urban area (Reche et al. 2012; Chang et al. 2016). According to the t test analysis, the ambient NH₃ concentrations at SM were significantly higher than those at XA (p < 0.001, 2-tailed) even though N fertilizer application could be an additional emission source at XA. The busy traffic in particular emissions from passenger cars equipped with TWC could make a significant contribution to ambient NH₃ at SM site (Phan et al. 2013; Wang et al. 2016, 2018). Strong positive correlations between NH₃ and CO (mainly emitted as vehicle exhaust) concentrations were observed at the traffic station in Rome (Italy) and a mixed-use urban area (office, commercial, residential, and traffic) east of downtown Shanghai (China) (Perrino et al. 2002; Chang et al. 2016). Similar significant correlations were also observed in urban Seoul in spring, fall, and winter except in summer (Phan et al. 2013). The contribution of vehicle emissions to ambient NH₃ in urban Shanghai was estimated to be 12.6-24.6% based on the difference of the NH₃ concentration above the open section of a traffic tunnel and the ambient concentration at the urban site (Wang et al. 2018). In addition, approximately 330,000 and 42,000 people live in the region around 2 km to SM and XA sites, respectively. Thus, the emissions of NH₃ from residential building toilets and septic tanks were expected to contribute more to the ambient NH₃ at SM than at XA (Chang et al. 2015). The spatial distribution of NH_3 around Xiamen Bay indicated that the urban area is the potential NH₃ emission region in Xiamen.

Table S2 summarizes the reported ambient NH₃ concentration in different cities in the world. The measurements in this study were lower than those determined in Beijing and Xi'an in China, Taihu lake region in China, Santiago in Chile, Rome **Fig. 2** Monthly variation in NH₃ concentration (IUE, DD, XA, SM, and ZZ) and average temperature (T) in Xiamen Bay, China, from June 2015 to Jun 2018



in Italy, and Okhla and Delhi in India, while higher than those reported in Shanghai and Nanjing in China, Ho Chi Minh City in Vietnam, Madrid in Spain, Niigata in Japan, and some cities in Canada and USA, and were comparable to those measured in urban Seoul (Korea), Barcelona (Spain), and Taichung (China). These comparisons suggested that urban nonagricultural sources could potentially influence ambient NH₃. Generally, the seasonal patterns of NH₃ were characterized by higher values in summer and lower values in winter (Table S2). However, the concentrations of NH₃ in Delhi urban-agro area in India was lower in summer than those in winter and fall due to less agricultural activities in summer (Sharma et al. 2010). Higher ambient NH₃ concentrations were also reported over the cornfield when the fertilizer application rate increased from 40 to 135 kg N ha⁻¹ (Felix et al. 2014).

Seasonality of PM_{2.5} and its water-soluble ions

Table 1 summarizes the statistics of $PM_{2.5}$ and its watersoluble ion concentrations. The concentration for $PM_{2.5}$ at SM ranged from 14.0 to 154.8 µg m⁻³ with a mean of 61.8 ± 34.9 µg m⁻³, which was slightly higher than that at XA (ranged from 13.6 to 123.8 µg m⁻³ with a mean of 56.4 ± 30.9 µg m⁻³). The PM_{2.5} values in this study were lower than the Chinese National Ambient Air Quality Standard of 75 $\mu g~m^{-3}$ for PM_{2.5} (Grade II, GB3095–2012) but much higher than the annual air quality guideline levels of 25 μ g m⁻³ for PM_{2.5} recommended by the World Health Organization (WHO 2006). Although the difference for PM_{2.5} concentration between the two sites was not statistically significant based on the paired t test (p = 0.404, 2-tailed) (Table S3), the $PM_{2.5}$ concentrations at the two sites were strongly correlated (r = 0.592, p < 0.001, n = 36). In addition, the average PM_{2.5} concentration at the urban site was slightly higher than the mean value of 51.5 μ g m⁻³ in 2011/2013 in the urban area of Xiamen (Wu et al. 2015). Seasonal variations in cations (Na⁺, NH₄⁺, K⁺, Mg²⁺, and Ca²⁺), anions (Cl⁻, NO₃⁻, and SO_4^{2-}), and $PM_{2.5}$ are depicted in Fig. 3. In this study, the trend for PM_{2.5} was higher in the winter, spring, and fall, and lower in the summer. The lower summer values for PM_{2.5} were apparently due to clean air masses from the South China Sea and Taiwan Strait (Figs. S1 and S2), lower gaseous precursors, and higher temperature (Fig. S3) unfavorable for the gas-to-aerosol conversion (Seinfeld and Pandis 2006; Wu et al. 2015). In addition, higher rainfall in summer promotes the scavenging of particles by wet deposition. The sum of the total water-soluble inorganic ions correlated well with PM25 (r = 0.331 - 0.673, p < 0.05) and accounted for $33 \pm 16\%$ and $35 \pm 20\%$ of PM_{2.5} at SM and XA, respectively. The mass fraction of water-soluble ions in PM2.5 were lower than



Fig. 3 Variation of PM_{2.5} and its chemical species from December 2016 to November 2017 at SM and XA

	SM (urban, $n = 36$)		XA (suburban, $n = 36$)	
	Range	Average \pm SD	Range	Average \pm SD
sPM _{2.5} (µg m ⁻³)	14.0-154.8	61.8 ± 34.9	13.6–123.8	56.4 ± 30.9
$Na^{+} (\mu g m^{-3})$	0.08-1.94	0.54 ± 0.45	0.09-1.33	0.49 ± 0.30
$NH_4^+ (\mu g m^{-3})$	0.18-9.49	2.60 ± 2.21	0.22-7.41	2.46 ± 1.80
$K^{+} (\mu g m^{-3})$	0.03-0.69	0.22 ± 0.14	0.04-0.49	0.18 ± 0.10
Mg^{2+} (µg m ⁻³)	0.03-0.40	0.10 ± 0.08	0.04-0.18	0.08 ± 0.03
$Ca^{2+} (\mu g m^{-3})$	0.16-3.28	0.68 ± 0.66	0.19-1.16	0.49 ± 0.20
$Cl^{-} (\mu g m^{-3})$	0.01-2.10	0.48 ± 0.48	0.09-2.25	0.44 ± 0.45
$NO_3^{-}(\mu g m^{-3})$	0.51-27.41	5.26 ± 6.91	0.77-28.31	4.31 ± 5.65
$SO_4^{2-} (\mu g m^{-3})$	1.06-15.13	6.34 ± 3.30	1.32-14.70	6.30 ± 3.37
$SO_2 (\mu g m^{-3})$	7.58-16.94	11.64 ± 4.01	7.81-19.09	12.92 ± 4.10
$NO_2 (\mu g m^{-3})$	18.43-49.89	30.94 ± 12.39	22.20-42.27	31.72 ± 7.15
$NH_3 (\mu g m^{-3})$	2.57-13.18	7.54 ± 3.31	1.80-9.53	5.36 ± 2.20
$n-NH_4^+/n-NO_3^-$	0.43-13.46	3.83 ± 3.24	0.57-9.59	3.42 ± 2.14
$n-NH_4^+/n-SO_4^{2-}$	0.75-6.12	2.29 ± 1.20	0.73-5.50	2.41 ± 1.24
$n-NH_4^+/n-(NO_3^- + SO_4^{-2})$	0.27-2.49	1.16 ± 0.52	0.32-2.21	1.18 ± 0.40
n-NH ₃ /n-NH ₄ ⁺	0.58-12.69	3.88 ± 3.29	0.34-7.67	2.94 ± 2.28
n-NH ₃ /n-NO ₂	0.14-1.91	0.80 ± 0.49	0.12-1.16	0.49 ± 0.27
n-NH ₃ /n-SO ₂	0.63-6.22	2.78 ± 1.70	0.38-3.41	1.76 ± 0.90
$n-NH_3/n-(NO_2 + SO_2)$	0.12-1.46	0.61 ± 0.37	0.09–0.87	0.38 ± 0.20

Table 1 Mass concentrations of water-soluble ions in $PM_{2.5}$, NH_3 , SO_2 , NO_2 , and molecular ratios of NH_4^+ to NO_3^- , NH_4 to SO_4^{2-} , NH_4^+ to $NO_3^- + SO_4^{2-}$, NH_3 to NH_4^+ , NH_3 to NO_2 , NH_3 to SO_2 , and NH_3 to $NO_2 + SO_2$ at SM and XA sites in Xiamen Bay from December 2016 to November 2017

41.1-49.5% in Beijing-Tianjin-Hebei region (Gao et al. 2018), 43% at Oimu Island in Bohai Sea (Zong et al. 2016), and comparable to 35% in urban Guangzhou and 34% in suburban Zhuhai in China (Tao et al. 2017), indicating that the water-soluble inorganic ions might not play a decisive role in ambient PM2 5 in Xiamen. Secondary inorganic ions including SO_4^{2-} , NO_3^{-} , and NH_4^{+} (SNA) constituted the majority of the total water-soluble inorganic ions (more than 80%). The proportion was similar to the values in the North China Plain (average 86%) (Xu et al. 2016). It was clear that the increase of PM2.5 was associated with the formation of SNA suggesting that the secondary formation of inorganic aerosol played an important role in PM2.5 pollution. Furthermore, the NO₃⁻ concentrations exhibited strong correlation with its precursor NO₂ at SM (r = 0.81, p < 0.001, n = 36) and XA (r =0.723, p < 0.001, n = 36) while the SO₄²⁻ concentrations correlated with SO₂ only at SM (r = 0.38, p = 0.01, n = 36) and there was no significant correlation between NH₄⁺ and NH₃ at both sites (Fig. S4). The NO_3^- in $PM_{2.5}$ is generally in the form of NH₄NO₃ that is formed by the oxidation of NO_x to the gas phase of HNO₃ and then reaction with NH₃. The low temperature in winter and spring can beneficially promote the gas phase of HNO₃ to transform into aerosol NH₄NO₃ (Seinfeld and Pandis 2006).

According to the paired-samples *t* test results for NH_4^+ and NH_3 between SM and XA in Table S4, the spatial difference of NH_4^+ was more influenced by the post-emission

atmospheric chemistry than its gaseous precursor in Xiamen. Strong correlations were found between NH_4^+ and SO_4^{2-} and NO_3^- at both sites (r = 0.59 - 0.90, p < 0.001, n = 36) (Fig. S4). The molecular ratios of NH_4^+ to $NO_3^- + SO_4^{2-}$ and NH_3 to $\mathrm{NH_4}^+$ were on average of 1.16 ± 0.52 and 3.88 ± 3.29 at SM and 1.18 ± 0.40 and 2.94 ± 2.28 at XA, respectively (Table 1). From this point, NH_4^+ was fully neutralized with SO_4^{2-} and NO₃⁻, and there might be sufficient equilibrium reactions between gas NH_3 and aerosol NH_4^+ in the atmosphere when gas NH₃ concentration exceeds that of NH₄⁺, especially during the warmer season from May to October. In the atmosphere at Maki Monitoring Station in Niigata Prefecture, Japan, the molecular ratios of NH₃ to NH₄⁺ ranged from 1.4 to 22.8 with a mean of 7.5 ± 5.8 , which were in accordance with the isotopic fractionation under equilibrium conversion of gas NH₃ to aerosol NH4⁺ (Hayasaki et al. 2004). In contrast, the low molar ratios of ambient NH₃ to $(NO_2 + SO_2)$ (< 1) appeared to show ammonia-poor atmospheric conditions during the sampling period at the two sites. Based on the obvious monthly variations of NH_4^+ and NO_3^- concentrations in Fig. 3, it can be assumed that part of NH4⁺ existed as relatively less stable ammonium salts (e.g., NH₄NO₃). It is expected that the diffusion of NH₃ back to the atmosphere during the reversible reaction and equilibrium between NH₃ and NH₄⁺ induced significant isotopic fractionation, thus exhibiting much higher δ^{15} N-NH₄⁺ values in PM_{2.5} than its initial ammonia sources (Table S1) (Kawashima and Kurahashi 2011). Therefore, a net isotope effect between NH₄⁺-NH₃ ($\epsilon_{\rm NH4}^+$ -_{NH3}) should be added to Eq. (1) to calculate the δ^{15} N values of initial NH₃ before putting into IsoSource model.

Isotopic signatures of NH₄-N in PM_{2.5}

The δ^{15} N-NH₄⁺ values in PM_{2 5} over the study period showed a wide variability among the samples (Fig. 4). The δ^{15} N-NH₄⁺ ranged from 5.29 to 26.13% with a mean of $13.96 \pm 4.56\%$ at SM site and from 1.52 to 27.03% with a mean of $14.34\pm$ 5.52% at XA site (Table S5); however, they were not statistically different (p = 0.678) (Table S4). Our measurements were heavier than those reported in urban Beijing by Pan et al. (2016, 2018a, b), Baengnyeong Island in Korea by Park et al. (2018), urban Colorado in the USA by Moore (1977), rural and coastal sites in Ireland and England by Heaton et al. (1997), and Yeatman et al. (2001) in aerosol, and similar to the measurements in Taihu Lake region in China by Ti et al. (2018), urban Seoul in Korea by Park et al. (2018), rural Niigata in Japan by Hayasaka et al. (2004), Jeju Island in Korea by Kundu et al. (2010), rural Yurihonjo in Japan by Kawashima and Kurahashi (2011), and urban and industrial areas in Alberta, Canada (Savard et al. 2017) in aerosol (Table S5). As summarized in Table S1, the δ^{15} N-NH₄⁺ from ocean and inside tunnel were in the range of -8 to -5%(Jiekells et al. 2003) and 29.6 to 40.0% (Park et al. 2018), respectively, which were beyond the scope of the measured δ^{15} N-NH₄⁺ values in PM_{2.5} at the coastal sites in this study. So the particulate NH_4^+ from the two emission sources was not likely to be important contributors.

Although the different nitrogen isotope analysis techniques might influence the measurements, the significant differences for the δ^{15} N-NH₄⁺ values at different locations over the world as summarized in Table S5 were more affected by the emission sources and/or isotopic fractionation. For the different surroundings between the two sites in this study, the comparison results indicated that the post-emission processes (e.g., equilibrium fractionation between gaseous NH₃ and aerosol NH₄⁺) might play a more important role than that of NH₃ emission sources in determining δ^{15} N-NH₄⁺ values. At both sites, the seasonal trend for δ^{15} N-NH₄⁺ was slightly heavier in



Fig. 4 Variations of δ^{15} N-NH₄⁺ (%) in PM_{2.5} at SM and XA in Xiamen from December 2016 to November 2017

the summer $(16.06 \pm 5.09\%$ at SM and $18.91 \pm 4.58\%$ at XA; June to September) than in other seasons. In addition, a significant positive correlation between ambient temperature and δ^{15} N-NH₄⁺ values was observed at both sites (r = 0.33, p < 0.05 at SM; r = 0.67, p < 0.001 at XA) (Fig. S4). Similar good correlation between δ^{15} N ratios and ambient temperature was also observed for PM₁₀ collected from Chennai, India, during late winter (Pavuluri et al. 2010). The seasonal trends agreed with the theoretical concepts that temperature effects on isotopic fractionation may also affect the δ^{15} N values of atmospheric reaction products.

Besides, the seasonality of the δ^{15} N-NH₄⁺ values has been often attributed to the changing of the initial δ^{15} N-NH₃ (Kundu et al. 2010; Kawashima and Kurahashi 2011; Park et al. 2018; Ti et al. 2018). Ti et al. (2018) recently reported higher δ^{15} N-NH₃ values from fertilizer application and livestock waste in winter but lower values in summer and fall in Taihu Lake region, China. The peak in spring agricultural activity was accompanied by a decrease in $\delta^{15}\text{N-NH}_3$ values at some US ammonia monitoring network sites (Felix et al. 2017). Chang et al. (2016) measured the δ^{15} N-NH₃ values of different volatilized ammonia and found no clear difference between summer and winter. The higher δ^{15} N-NH₄⁺ values in summer and fall in Jeju Island and Baengnyeong Island in Korea were more explained by the air mass from China (including enhanced biomass burning) than by isotope fractionation (Kundu et al. 2010; Park et al. 2018). Although the measured δ^{15} N-NH₄⁺ values were similar to the δ^{15} N-NH₄⁺ values of rice straw burning smoke (Table S1), the average contribution of biomass burning to NH3 in Xiamen was estimated to be only 0.4% in our previous study (Wu et al. 2017). The facts that the volatilized NH₃ emissions such as animal waste, fertilizer application, and waste treatment with depleted ¹⁵N (Table S1) reach maximum due to the highest temperature in summer while the fossil fuel-related NH3 emissions which are isotopically more enriched in ¹⁵N reach maximum in winter due to the energy demand for heating are in conflict with the observed seasonal trend of δ^{15} N-NH₄⁺ in this study.

Furthermore, the higher δ^{15} N-NH₄⁺ values associated with the stagnant air masses observed in Beijing and Jeju Island (Pan et al. 2018b; Kundu et al. 2010) were consistent with a more pronounced isotope fractionation due to the longer atmospheric lifetime of aerosol NH₄⁺ during the transport. It is reasonable to assume that atmospheric chemical reactions caused the lower values of δ^{15} N-NH₄⁺ in winter in Korea. In this study, the 3-day backward trajectories from the NOAA's HYSPLIT model in summer appeared to exhibit shorter distance with respect to those in other seasons (Fig. S1) suggesting a possible effect of stagnant air masses on δ^{15} N-NH₄⁺ values in summer. In the summer, the high temperature enhances the NH₃ volatilization rates and the NH₃ pool in the atmosphere seems to be filled with NH₃. The enhanced isotope fractionation followed by the intensified gas-to-particle conversion (NH₃ \leftrightarrow NH₄⁺) at high NH₃ concentration favors the production of heavier particulate ¹⁵NH₄⁺ in the atmosphere (Kawashima and Kurahishi 2011). Therefore, the equilibrium isotope effect between the gaseous NH3 and aerosol NH₄⁺ and its temperature dependence might be the most important factor influencing the seasonality of δ^{15} N-NH₄⁺ (Heaton et al. 1997; Kawashima and Kurahashi 2011; Li et al. 2012: Savard et al. 2017).

Contributions of NH₃ to particulate NH₄⁺

The initial δ^{15} N-NH₃ values before the start of gas-to-particle conversion were calculated based on the assumption that there existed nitrogen isotope equilibrium exchange between gas NH_3 and aerosol NH_4^+ (Pan et al. 2016; Park et al. 2018). The calculated δ^{15} N-NH₃ values ranged from -17.48 to -4.85% at SM and from -20.81 to -4.11% at XA, which were similar to the measured δ^{15} N-NH₃ values of – 11.5 to -8.5% at the highway interchange in urban Seoul (Park et al. 2018), -14.5 to -1.0% in Niigata Prefecture, Japan (Hayasaki et al. 2004), -24.2 to -3.0% in Taihu Lake region, China (Ti et al. 2018), and $-10 \pm 2.6\%$ in Colorado, USA (Moore 1977), and were in the range (-42.4 to +7.1%) of those reported for 9 AMoN sites across the USA (Felix et al. 2017) but were much higher than those reported in urban Beijing (-43.1 to -27.1%), Chang et al. 2016), near coalfired power plant (-38.0 to -21.1%) and fertilizer plant (-31.3%) in Alberta, Canada (Savard et al. 2017). The calculated δ^{15} N-NH₃ values were generally heavier than those in NH₃ from fertilizer application, livestock excreta, and waste treatment, but lighter than coal combustion and traffic exhausts (Table S1). It should be noted that the measured δ^{15} N-NH₃ values increased from -32.9% on the first day to -19.06%on the second day and -6.39% 7 days after urea application. The trend that volatilized NH₃ after fertilization becomes isotopically heavier over time was also observed in Ti et al. (2018) and Frank et al. (2004). However, the δ^{15} N-NH₃ values changed little for organic fertilizer application experiments. Thus, strong nitrogen fractionation may mask NH₃ source signatures when ambient NH₃ in rural agricultural area is mainly originated from urea fertilizer application.

The IsoSource isotopic mixing model was used to estimate the relative contributions of the four different emission sources on the basis of the calculated δ^{15} N-NH₃ (Fig. 5). The results indicated that fossil fuel combustion-merged with coal combustion and vehicle emissions-dominated the ambient NH₃, with a mean contribution of 44% (in the range of 15-85%) at SM and 59% (in the range of 11-90%) at XA, which were much higher than the results in Taihu Lake region (26%, 19%, and 20% at urban, suburban, and rural sites, respectively) (Ti et al. 2018) and urban Beijing (8.8–29.4%) (Chang et al. 2016). Based on bulk δ^{15} N signatures and a Bayesian isotope mixing model, Wang et al. (2017) estimated



Fig. 5 Relative contributions of individual emissions to atmospheric NH₃ in urban and suburban Xiamen estimated from the IsoSource model

that fossil-derived NH₃ (coal combustion and traffic sources) contributed about 30% of bulk N in PM2.5 in urban Beijing. In comparison, fertilizer applications, NH₃ slip, and human excreta only accounted for 13-17%, 16-26%, and 12-14% on average of the ambient NH₃, respectively. The annual contribution ratios of NH₃ from fertilizer application to fossil fuel combustion (0.22-0.37) at the two sites were much lower than the ratio (0.74) at the county level using emission factors (Wu et al. 2017). With a total of 1.11 million civil vehicles in 2016 and a stable growth trend of vehicle numbers and urbanization in Xiamen, regulatory control of urban transportation and coal combustion in Xiamen and surrounding areas will become increasingly important to air quality improvement.

As depicted in previous studies, the uncertainties in the determination for NH₃ end members and isotopic effect taking place during the air-land exchange and gas-aerosol exchange dominate the overall uncertainties in the source apportionment of ambient NH₃ (Chang et al. 2016; Felix et al. 2017; Pan et al. 2018a). As summarized in Table S1, results published for end member data for a single source are variable. In addition, performing near-source sampling to determine site-specific δ^{15} N-NH₃ of individual sources instead of calculating δ^{15} N-NH₃ from δ^{15} N-NH₄⁺ values can be made in order to improve and provide statistical confident data (Felix et al. 2017; Savard et al. 2017).

Conclusions

Previous studies have indicated that agricultural emission sources were the largest contributing emission sources to ambient NH₃ at a county or larger scale level, while nonagricultural NH₃ in the urban district could contribute more to the local NH₃ due to its short lifetime. In this study, attempts have been made to measure the ambient NH₃, NH₄⁺, and δ^{15} N-NH₄⁺ values in PM_{2.5} in Xiamen to separate NH₄⁺ into four different source components based on the isotope mixing model. The temperature was identified as the mainspring leading to the seasonal variations of NH₃, NH₄⁺, and δ^{15} N-NH₄⁺

values. We confirmed that the atmosphere in Xiamen was NH₃-rich based on the high molecular ratios of NH₃ to NH_4^+ and NH_4^+ to $NO_3^- + SO_4^{2-}$ (>1). A distinct seasonal trend of δ^{15} N-NH₄⁺ values could be detected: lighter in winter and heavier in summer. The initial δ^{15} N-NH₃ values were calculated under the conditions of nitrogen isotope equilibrium fractionation between gas NH₃ and aerosol NH₄⁺. Source apportionment by the IsoSource isotopic mixing model indicated that fossil fuel-related emissions (fossil fuel combustion and NH₃ slip) dominated aerosol NH₄⁺ in urban Xiamen, which were basically consistent with the emission inventory in urban Xiamen. These findings suggested that the $\delta^{15}N$ of NH₃ and NH₄⁺ can be used as a useful tool to effectively perform source apportionment of atmospheric NH₃ in an urban area and understand how the nitrogen isotope exchange reaction between NH₃ and NH₄⁺ is influenced by the temperature. Further studies on the relationship between δ^{15} N-NH₃ from the emission sources and δ^{15} N- NH₄⁺ at a receptor site can help to develop effective PM2.5 control strategies.

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