

## **Simultaneous measurement of aqueous redox-sensitive elements and their species across the soil-water interface**

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1 **Simultaneous measurement of aqueous redox-sensitive elements and their species**  
2 **across the soil-water interface**

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23 **Abstract:** The redox-sensitive elements, such as iron, manganese, sulfur, phosphorus,  
24 and arsenic, shift their speciation every millimeter (mm) across the soil-water  
25 interface in the flooded soil environments. Monitoring of element speciation at this  
26 high-resolution (HR) within the SWI is still difficult. The key challenge lies in  
27 obtaining sufficient porewater samples at specific locations along the soil gradient for  
28 downstream analysis. Here with an optimized inductively coupled plasma mass  
29 spectrometry (ICP-MS) method and a HR porewater sampler, we demonstrate  
30 mm-scale element profiles mapping across the SWI in paddy soils.  
31 High-concentrations of iron and manganese ( $> 10 \text{ mg}\cdot\text{L}^{-1}$ ) were measured by ICP-MS  
32 in an extended dynamic range mode to avoid signal overflow. The iron profile along  
33 the SWI generated by the ICP-MS method showed no significant difference ( $p < 0.05$ )  
34 compared to that measured independently using a colorimetric method. Furthermore,  
35 four arsenic (arsenite, arsenate, monomethylarsonic and dimethylarsinic acid), two  
36 phosphorus (phosphite and phosphate) and two sulfur (sulfide and sulfate) species  
37 were separated in 10 min by ion chromatography -ICP-MS with the  $\text{NH}_4\text{HCO}_3$  mobile  
38 phase. We verified the technique using paddy soils collected from the field, and  
39 present the mm-scale profiles of iron, manganese, and arsenic, phosphorus, sulfur  
40 species (relative standard deviation  $< 8\%$ ). The technique developed in this study will  
41 significantly promote the measurement throughput in limited samples (e.g.  $100 \mu\text{L}$ )  
42 collected by HR samplers, which would greatly facilitate redox-sensitive elements  
43 biogeochemical cycling in saturated soils.



45 **Keywords:**

46 porewater

47 soil-water interface

48 arsenic

49 iron

50 manganese

51 sulfur

52 species

53

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57

## 58 **1. Introduction**

59 In flooded soils, the chemical environments of the surface water and saturated  
60 sediment porewater are very different. The surface water is oxidizing due to the high  
61 dissolved O<sub>2</sub>, however, the sediments are generally reducing owing to the lack of O<sub>2</sub>  
62 and the abundance of organic matter (Frenzel et al., 1992). The O<sub>2</sub> can only penetrate  
63 the upper sediment to a depth of a few millimeters (mm) (Ratering and Schnell, 2001).  
64 As a result, the narrow boundary zone between the surface water and sediments, i.e.  
65 soil-water interface (SWI), displays a sharp redox decrease with depth (Huo et al.,  
66 2015; Jones et al., 2018). Iron (Fe), manganese (Mn), and sulfur (S) are the most  
67 important elements in SWI, existing in both solid and dissolved phases through  
68 complex redox reactions (Peng et al., 2019). The redox processes of Fe, Mn, and S  
69 significantly impact the fate of many elements of environmental and agricultural  
70 concern, such as arsenic (As) and phosphorus (P) (Gao et al., 2016; Gao et al., 2006;  
71 Mcadams et al., 2016; Pi et al., 2018). Although it is of great importance to study the  
72 behavior of these elements in SWI, the high-resolution (HR) mm-scale mapping of  
73 those elements and their species has been severely limited to date by the lack of  
74 suitable available methods.

75 Many efforts have been made to measure the mm-scale element profile in  
76 porewater along SWI. The diffusive gradient in thin films technique (DGT) is one of  
77 the best, which can even depict the elements' pattern in  $\mu\text{m}$  scale, however it measure  
78 the flux instead of the equilibrated concentration (Davison and Zhang, 1994; Fang et  
79 al., 2018; Yin et al., 2020). The equilibrated concentrations can be measured by the  
80 diffusive equilibrium in thin films (DET) technique and *in situ* equilibrium dialysis  
81 samplers (peeper) (Arsic et al., 2018; Bottrell et al., 2007; Di et al., 2012; Dočekalová  
82 et al., 2002; Guan et al., 2015; Monbet et al., 2008). The DET probe resolution is 2  
83 mm when using the strip-cutting method (Dočekalová et al., 2002; Gao et al., 2007),  
84 and can reach 1 mm when combining reagent dyeing and computer imaging  
85 densitometry detection (Bennett et al., 2012b; Robertson et al., 2008). Peepers have  
86 relatively low spatial resolution ( $\sim 5$  mm) compared to DET because handling the  
87 water in peeper chambers is not as convenient as the gels in DET probe (Di et al.,  
88 2012; Wen et al., 2019). Recently, we developed a novel porewater sampler, called  
89 *In-situ* Porewater Iterative (IPI) sampler, to monitor the mm-scale heterogeneity of  
90 trace metals in saturated soils (Yuan et al., 2019). The IPI sampler has a comparable  
91 HR ( $\sim 2$  mm) as DET probe. Unlike DET and peeper, the IPI sampler can be used  
92 repeatedly at a certain place without need for removal or destructive sampling.  
93 Another advantage of IPI samplers is to obtain clean liquid porewater sample directly,  
94 which is almost ready for downstream instrumental analysis. Due to these advantages,  
95 the IPI sampler was very suited for mm-scale element profile mapping.



96 Simultaneous measurement of multi-element profiles at HR across SWI presents  
97 significant additional challenges. The sharp and sensitive redox gradient along SWI  
98 requires that the porewater volume sampled should be as small as possible to  
99 minimize the disturbance to the sampling environment (Seeberg - Elverfeldt et al.,  
100 2005), and yet large enough to meet the minimum sample size for sensitivity and  
101 specificity analysis of all the interested parameters (Arsic et al., 2018; Bennett et al.,  
102 2012a; Ding et al., 2016; Motelica-Heino et al., 2003). The HR samplers (e.g. DET,  
103 HR peeper, IPI samplers) designed for element profile mapping generally can only  
104 take less than 0.5 mL solution (Yuan et al., 2019), which is a bare minimum for one  
105 sample injection with most analytical techniques, like inductively coupled plasma  
106 mass spectrometry (ICP-MS) (Xu et al., 2017), ICP- optical emission spectrometry  
107 (Cheng et al., 2012), colorimetric method (Lumbaue et al., 2019). An alternative to  
108 collect more samples is to collect porewater repeatedly at different places or times,  
109 assuming the soil or sediment matrix is homogenous and stable over time. However,  
110 this assumption is severely limiting and constrains the ability to probe the  
111 heterogeneity and dynamics of SWI biogeochemistry as a function of location and in  
112 response to changing environmental conditions (Arsic et al., 2018; Yuan et al., 2019).  
113 Thus, it would be better to solve the issue by optimizing the analytical techniques  
114 used in extracting data from the samples.

115 ICP-MS has been widely applied to understand the element behaviors in various  
116 environments due to its broad spectrum and very low detection limits (Cotta and

117 Enzweiler, 2009). Studies of rhizospheric element profiles have greatly benefited  
118 from developments in ICP-MS technologies. For example, the combination of laser  
119 ablation (LA)-ICP-MS with DGT allows mapping of  $\mu\text{m}$ -scale element fluxes and  
120 provided key information for understanding As behaviors on the root apices  
121 (Williams et al., 2014). More recently, a new approach, called extended dynamic  
122 range (EDR), was introduced to simultaneously measure major and trace metals by  
123 ICP-MS (Hilbig et al., 2017). The EDR mode can attenuate the counts of selected  
124 elements through the spectrometer by tuning the 'rejection parameter a' (Rpa), thus it  
125 enables the detection of major and minor elements in a single run (Hilbig et al., 2017).  
126 EDR mode is potentially ideal to measure multi-element concentrations in  
127 volume-limited samples, such as the porewater sampled by HR samplers, which  
128 contains Fe and Mn over  $10 \text{ mg}\cdot\text{L}^{-1}$ , and other traces at  $\mu\text{g}\cdot\text{L}^{-1}$  concentrations.  
129 However, the combination of HR samplers and ICP-MS under EDR mode has not  
130 been tested to date.

131 The small porewater sample volume also hinders the measurement of element  
132 speciation. Arsenic, P and S species play crucial roles in aquatic biogeochemical  
133 cycling (Chen et al., 2019; Sun et al., 2017). Traditionally, phosphate, P(V), is  
134 measured by colorimetry (molybdate blue) (Rietra et al., 2001), sulfide, S(-II), by  
135 micro-electrode and spectrophotometric method (Laskov et al., 2007), sulfate, S(VI),  
136 by ion chromatography (IC) (Keller-Lehmann et al., 2006), arsenite and arsenate,  
137 As(III,V), by IC-ICP-MS (Gallagher et al., 2001). Summing up, to measure all these

138 analytes in a single sample by standard methodologies, it requires an aliquot volume  
139 of several mL, far more than the porewater volume collected from HR samplers (Xu  
140 et al., 2012; Yuan et al., 2019). Among methods noted here, IC-ICP-MS can measure  
141 all the species, but is limited to the appropriate mobile phase. A review of the  
142 literature (Chen et al., 2019; McDowell et al., 2004; Morton et al., 2005; Reid et al.,  
143 2020; Suzuki et al., 2009) revealed that  $\text{NH}_4\text{HCO}_3$  mobile phase, which is free of As,  
144 P, and S, can separate P, S and As species. Combined with the ICP-MS method  
145 (Hilbig et al., 2017), use of  $\text{NH}_4\text{HCO}_3$  elution IC-ICP-MS thus presents the possibility  
146 of accurate and rapid quantification of key major and minor element concentration  
147 and speciation within SWI profiles from small samples obtained by HR samplers such  
148 as IPI (Yuan et al., 2019), enabling non-destructive, mm-scale and repeated probing of  
149 SWI chemistry over time and therefore overcoming several key limitations of existing  
150 approaches outlined above.

151 This study demonstrates simultaneous measurement of multi-element and  
152 multi-species concentration profiles in flooded soil porewater using a HR sampler (IPI,  
153 after Yuan et al. (2019)) combined with the optimized ICP-MS and IC-ICP-MS  
154 method. The typical redox-active elements found in soil porewater, including Fe, Mn,  
155 As, P and S, and the common species of As, P, and S, including phosphite (P(III)),  
156 P(V), S(-II), S(VI), monomethylarsonic acid (MMA), dimethylarsinic acid (DMA),  
157 As(III), As(V), were investigated in flooded paddy soils. This study directly addresses  
158 the challenge of maximizing the chemical information obtainable from increasingly

159 small sample volumes, which would greatly enhance the measurement range,  
160 throughput and application potential of HR samplers.

161

## 162 **2. Materials and methods**

### 163 **2.1 Reagents and materials**

164 All reagents used in this study were of analytical grade or higher, and purchased from  
165 Aladdin Chemical Reagent Co., Ltd. (Shanghai, China), unless stated otherwise.  
166 Element standards for calibration, including As, Fe, Mn, P, S as well as P, S, and As  
167 species, were supplied by Guobiao (Beijing) Testing & Certification Co., Ltd (Beijing,  
168 China). All solutions were prepared with ultrapure water (18.2 MΩ cm, Millipore  
169 Corp., Bedford, USA) deoxygenated by bubbling pure N<sub>2</sub> overnight.

170 Before the soil was sampled from paddy fields in Shaoguan (SG, 25°6'N,  
171 113°38'E), obvious stones and plant debris were mechanically removed by shovels. In  
172 total, ~ 50 kg soils from the top layer (0 - 20 cm) were collected. The soils were  
173 directly transported to the laboratory, and homogenized by passing through a 1.0 mm  
174 diameter wet sieve. The soil characteristics are shown in **Table S1**.

175

### 176 **2.2 Porewater sampler preparation**

177 The IPI sampler used in this study has a similar design as described in detail in our

178 previous study (Yuan et al., 2019), with some minor modifications noted here. A novel  
179 hollow fiber membrane tube (modified polyethersulfone, 20 nm pore size, inner ×  
180 outer diameter × length = 1.0 mm × 1.7 mm × 35 mm, 27.5 μL, Motimo Membrane  
181 Technology Co., Ltd., Tianjin, China) and two pipes (PTFE, inner × outer diameter ×  
182 length = 0.5 mm × 1.0 mm × 180 mm, 35 μL) were used to construct the IPI sampler.  
183 The pore size of the membrane was demonstrated (**Fig. S1**) with scanning electron  
184 microscopy (SEM) JSM-7600 (FJEOL Ltd., Japan). The updated membrane does not  
185 contain fluoride or other potential chelators, which can avoid the complexation of the  
186 membrane for certain analytes (e.g. lead) (Yuan et al., 2019).

187 When the IPI sampler is deployed into solution or saturated soils, solutes around  
188 the hollow fiber membrane tube can diffuse through the membrane (**Fig. S2A**). The  
189 solution inside the tube is pumped out and collected when the diffusion reaches  
190 equilibrium (**Fig. S2B**). During the deployment, silicon caps are applied to seal the IPI  
191 sampler to avoid potential contamination (e.g. gasoline) from the atmosphere. During  
192 each sampling event, 27.5 μL liquid sample in the sampling tube is mixed with 70 μL  
193 ultrapure water in pipes when they are pumped out from the sampler. This indicates ~  
194 100 μL porewater sample can be sampled each time by the IPI sampler, with a  
195 dilution factor of 3.5.

196 Thirty-four IPI samplers were horizontally assembled side by side in a 3D  
197 printed holder (cavity cuboid, length × width × height = 40 mm × 30 mm × 120 mm,  
198 **Fig. S2C-E**). The IPI sampler array, i.e. SWI profiler, can sample the porewater every

199 1.7 mm along SWI (**Fig. S2C**). The SWI profiler has a sampling depth of 60 mm and  
200 was stored in O<sub>2</sub>-free ultrapure water before deploying into flooded soils (**Fig. S2F**),  
201 following the procedure described in Yuan et al. (2019).

202

### 203 **2.3 Analytical method and quality control**

204 Element concentrations were quantified by ICP-MS (NexION 350X, PerkinElmer,  
205 Inc., Shelton, CT USA). The conditions were as follows: EDR mode; **dynamic**  
206 **reaction cell (DRC) mode** (O<sub>2</sub>, gas flow, 1.0 mL·min<sup>-1</sup>); data only analysis; RF power  
207 **1,600W**; plasma gas flow rate 15 L·min<sup>-1</sup>; auxiliary gas flow 1.2 L·min<sup>-1</sup>; nebulized  
208 gas flow 0.94 L·min<sup>-1</sup>; nickel sampling and skimmer cones. The Rpa in EDR mode is  
209 a voltage parameter, which can tune bandpass of *m/z* that has stable trajectories  
210 (Tanner and Baranov, 1999). This parameter functions for the precursors of target  
211 analyte generated from dynamic reaction/collision cell. When Rpa = 0 (default value),  
212 a wide range bandpass of *m/z* is obtained, while non-zero Rpa (0 - 0.24) creates a  
213 narrow bandpass of *m/z*. The sensitivity of *m/z* is proportional to the width of  
214 bandpass, hence the upregulation of Rpa provides an option to suppress the high  
215 sensitivity of *m/z* caused by high abundance analyte or potential interferences. When  
216 using Rpa under different analytical modes, the Rpa value can be manually set for the  
217 interested element.

218 The porewater sample collected by the IPI sampler was introduced into ICP-MS  
219 by a PFA-200 Microflow Nebulizer (0.2 mL·min<sup>-1</sup> uptake rate). Iron, Mn As, P and S  
220 were measured by ICP-MS in EDR and DRC mode. Counts of <sup>91</sup>AsO<sup>+</sup>, <sup>57</sup>Fe<sup>+</sup>, <sup>47</sup>PO<sup>+</sup>,  
221 <sup>48</sup>SO<sup>+</sup> and <sup>55</sup>Mn<sup>+</sup> were recorded.

222 Element species were measured by IC-ICP-MS. The IC (Dionex ICS-1100,  
223 Thermo Scientific, USA) consisted of a standard 25 μL sample loop and an  
224 anion-exchange column (IonPac AS23, 250 mm × 4 mm, Dionex). Mobile phases  
225 used for the separation were 20 mmol·L<sup>-1</sup> NH<sub>4</sub>HCO<sub>3</sub> at pH 10 (Suzuki et al., 2009),  
226 with a flow rate of 1.0 mL·min<sup>-1</sup>. The analytical column was connected to a Type C0.5  
227 Glass Nebulizer of the ICP-MS. The standards were prepared in neutral conditions  
228 (pH 7; 100 μg·L<sup>-1</sup> P(III), P(V), As(III), As(V), MMA and DMA; 1 mg·L<sup>-1</sup> S(-II) and  
229 S(VI)).

230 When developing the calibration curve, a series of standard solutions, containing  
231 1.0/10/100, 2.0/20/200, 5.0/50/500, 10/100/1000, 20/200/2000 μg·L<sup>-1</sup>  
232 As&Mn/P/Fe&S in 2% HNO<sub>3</sub>, were measured (*n* = 3). For As, P and S species, a  
233 series of standard solutions, containing 0/0/0, 1.0/20/500, 2.0/50/1000, 5.0/100/2000,  
234 10/200/5000 μg·L<sup>-1</sup> DMA/P(V)/S(VI) under pH 7, were measured. Peak area was  
235 used to fit the standard curve, and three times standard deviation was used to calculate  
236 the limit of detection (LOD). Data quality was assured by testing a spiked standard  
237 after every 30 samples.

238

## 239 **2.4 The sampling of total elements and element species by IPI samplers**

240 Total elements were prepared in acidic conditions (pH 2; 10  $\mu\text{g}\cdot\text{L}^{-1}$  As and Mn; 10  
241  $\text{mg}\cdot\text{L}^{-1}$  Fe, P, and S), and As species were prepared in neutral conditions (pH 7; 100  
242  $\mu\text{g}\cdot\text{L}^{-1}$  As(III), As(V), MMA and DMA). Those solutions were made by diluting the  
243 relative standards with ultrapure water. To determine the equilibrium time required for  
244 IPI samplers to sample Fe, Mn, As, P, S as well as As species, the time-dependent  
245 response of the sampler to those solutes was investigated in solutions. The samples  
246 inside the samplers were measured after 0, 0.5, 1, 3, 6, and 12 **hr** equilibrium time by  
247 ICP-MS or IC-ICP-MS.

248

## 249 **2.5 Multi-element profile mapping**

250 To detect the element profile, the SWI profilers were inserted into flooded soils in a  
251 pot (diameter  $\times$  height = 12 cm  $\times$  20 cm), with 10 mm above SWI and 50 mm in soils  
252 (**Fig. S2C**). Two replicates were conducted.

253 The paddy pot soils were filled with ultrapure water with  $\sim$  3 cm overlying water,  
254 and the water depth was maintained daily by supplementing ultrapure water during  
255 the experiment. The soils were allowed to stabilize (22  $^{\circ}\text{C}$ , dark conditions) for three  
256 months before the deployment of SWI profilers. Before sampling, the solution inside  
257 IPI samplers was replaced by  $\text{O}_2$ -free ultrapure water driven by an injection pump  
258 (TYD01, Lei Fu, China) (**Fig. S3**), with a velocity of 1.0  $\text{mL}\cdot\text{min}^{-1}$ .



259 Based on the equilibrium test of analytes in this study, the sampling interval was  
260 set as 24 hr. The sampled solution was preserved in the O<sub>2</sub>-free EDTA solution  
261 (Gallagher et al., 2001). The EDTA solution (2 g·L<sup>-1</sup>) was online mixed with the  
262 porewater in a 1:3 volume ratio driven by two injection pumps (**Fig. S3**), with a  
263 velocity of 0.25 and 0.75 mL·min<sup>-1</sup> respectively. The mixed sample was carefully  
264 collected in a clean 0.6 mL centrifuge tube. In total, ~ 100 µL porewater was collected  
265 with an EDTA concentration of 500 mg·L<sup>-1</sup>. Each sample was divided into two parts  
266 (~ 50 µL per part) and measured by ICP-MS and IC-ICP-MS respectively.  
267 Additionally, a microplate reader was used to measure the Fe profile in paddy soil  
268 with the 1, 10-phenanthroline method ( $\lambda = 510$  nm) (Lumbaue et al., 2019). Before  
269 measurement, the O<sub>2</sub>-free colorimetric reagent was online mixed with the porewater  
270 sample in a 1:1 volume ratio driven by two injection pumps (**Fig. S3**), and ~ 200 µL  
271 solution was collected and transferred into 96-well plate (200 µL) for absorbance  
272 determination in a microplate reader (Tecan-Spark, Tecan Trading AG, Zurich,  
273 Switzerland). Finally, mm-scale profiles of multi-element and multi-species were  
274 mapped and evaluated.

275

## 276 **2.6 Data analysis**

277 R software (version 3.5.0) was used to analyze and plot the data in this study. We used  
278 the standard deviation to show the variance of the data. Data of different methods

279 were subjected to one-way analysis of variance (ANOVA) to determine statistical  
280 significance ( $p < 0.05$ ) using SPSS 22 software (SPSS Inc., Chicago, USA).

281

### 282 **3. Results and discussion**

#### 283 **3.1 The time-dependent sampling efficiencies of Fe, Mn, As, P and S by IPI** 284 **samplers**

285 Initially, the trans-membrane diffusion of Fe, Mn, As, P, and S in solution condition  
286 was studied. This was done to determine the applicability of IPI samplers in collecting  
287 solutions with those elements, which is a prerequisite before using the sampler to  
288 collect those elements in soil porewater. Significant peaks of all the five elements  
289 were observed with a 50  $\mu\text{L}$  solution (**Fig. S4**). The equilibration test showed the  
290 concentrations of those elements increased rapidly in the first three **hr**, and then  
291 reached a plateau representing their concentration in solutions (**Fig. S5**). The  
292 time-dependent curves were consistent with our previous report for As, antimony,  
293 cadmium, lead, and nickel (Yuan et al., 2019). Considering the presence of dissolved  
294 organic matters, which could slow down the diffusion of solutes (Dočekalová et al.,  
295 2002; Reynolds et al., 2004; Yuan et al., 2019; Zhai et al., 2018), we deployed the IPI  
296 sampler at 24 **hr** equilibration period in saturated soils.

297

#### 298 **3.2 The effect of Rpa value on Fe and Mn detection**

299 The Fe and Mn signals can be attenuated via tuning the Rpa value in ICP-MS (Hilbig  
300 et al., 2017). The response of Fe and Mn signals to Rpa from 0 to 0.02 is shown in **Fig.**  
301 **S6**. The results indicated that Fe and Mn counts were very sensitive to Rpa value in a  
302 range from 0.0035 to 0.0084, in which Fe counts decreased linearly from 300000 to  
303 25000, and Mn from 85000,000 to 1000,000.

304 Furthermore, the LODs of Fe and Mn are 21.2, 30.9, 210 and 0.504, 1.53, 1.56  
305  $\mu\text{g}\cdot\text{L}^{-1}$  under Rpa 0, 0.005, and 0.01 respectively. Although LODs for Fe and Mn  
306 values decreased when the Rpa increased from 0 to 0.01, however it was sufficient for  
307 the Fe and Mn detection, since their concentrations are often found above  $1\text{ mg}\cdot\text{L}^{-1}$  in  
308 field porewaters (Gustave et al., 2018b; Wang et al., 2019; Xu et al., 2017). When the  
309 ICP-MS was used to measure major elements (e.g. Fe) in low abundance, high Rpa  
310 may lead to the unsuccessful determination due to the relatively low LOD, thus is not  
311 recommended.

312

### 313 **3.3 Comparison of colorimetric and ICP-MS for measuring Fe profiles in field** 314 **samples**

315 The ICP-MS application on Fe and Mn measurement was further investigated with  
316 paddy soil samples. With the method, simultaneous measurement of Fe, Mn, As, P,  
317 and S was achieved in the paddy soil samples. The determination coefficients for all  
318 the five elements were  $> 0.99$  (**Fig. S7**). The LODs for As, P and S were 0.490, 7.76,

319 and  $60.2 \mu\text{g}\cdot\text{L}^{-1}$ , respectively, which agrees well with previous reports where As, P  
320 and S were measured with the ICP-MS (Persson et al., 2009; Yuan et al., 2019). Most  
321 studies on porewater Fe used colorimetric analysis for Fe quantification (Arsic et al.,  
322 2018; Bennett et al., 2012b). To compare our method with the colorimetric method,  
323 two porewater samples were measured using each method.

324 The Fe profile measured by the ICP-MS with different Rpa values is shown in  
325 **Fig. 1**. ICP-MS method with Rpa 0 was unable to measure Fe in soil depth  $> 25$  mm  
326 when  $\text{Fe} > 17 \text{ mg}\cdot\text{L}^{-1}$  in soil porewater because of signal overflow (**Fig. 1A**).  
327 Increasing the Rpa from 0 to 0.005 extended the Fe measurement along soil depth  
328 from  $\sim 25$  mm to  $\sim 36$  mm (**Fig. 1A&B**), however Rpa 0.005 was still unable to avoid  
329 the detector saturation of  $\text{Fe} > 42 \text{ mg}\cdot\text{L}^{-1}$  in deep soil porewaters (**Fig. 1B**). **Fig. 1C**  
330 shows Fe profile could only be measured after the Rpa value was adjusted to 0.01  
331 with  $\text{Fe} \leq 70 \text{ mg}\cdot\text{L}^{-1}$ . Similarly to Fe, Rpa 0.005 or 0.01 allowed Mn measurement  
332 when  $\text{Mn} \leq 3.8 \text{ mg}\cdot\text{L}^{-1}$  in soil porewaters (**Fig. S9**). Although the upper limit was  
333 altered with different Rpa values, the Fe and Mn profiles were identical in top soils (0  
334 - 25mm) (**Fig. 1 A-C, Fig. S9**).

335 The porewater samples collected from the same location were also measured by  
336 the colorimetric method (**Fig. 1D**). Both colorimetric and ICP-MS methods gave  
337 similar results of Fe profiles ( $p > 0.05$ ). However, Fe concentration measured by the  
338 colorimetric method was slightly higher than that obtained by the ICP-MS method.  
339 The higher Fe values obtained by the colorimetric method could be attributed to the

340 interference from other cations (e.g. Mn, calcium, zinc) in the porewater (Hatat-Fraile  
341 and Barbeau, 2019; Miranda et al., 2016). Therefore, Fe concentrations might have  
342 been overestimated as was reported in previous studies (Braunschweig et al., 2012;  
343 Miranda et al., 2016).

344

### 345 **3.4 Profiling of total As, Fe, Mn, P and S across SWI**

346 Using the ICP-MS coupled with IPI samplers, we were able to simultaneously  
347 measure Fe, Mn, As, P, and S at the mm-scale co-distributions of those elements (**Fig.**  
348 **2**, relative standard deviation < 8%). As shown in **Fig. 2**, the Fe remained at low  
349 concentrations in surface water and top-soil porewater, but increased sharply from 9  
350 mm ( $1.5 \text{ mg}\cdot\text{L}^{-1}$ ) below SWI and reached up to  $70 \text{ mg}\cdot\text{L}^{-1}$  in 50 mm deep soils.  
351 Similar trends were observed for Mn, As, and P, which generally increased with depth.  
352 The similar vertical changes of Mn, As, P with Fe agree well with their tightly  
353 coupling in soils induced by dissimilatory Fe reducing bacteria (Arsic et al., 2018; Ma  
354 et al., 2017; Xu et al., 2017). At the 2 - 10 mm zone below SWI, a P pit was observed,  
355 which could be attributed to two reasons. First, unlike Fe, Mn, and As, P in surface  
356 water remained a relatively high concentration ( $\sim 100 \text{ }\mu\text{g}\cdot\text{L}^{-1}$ ), which indicated a  
357 constant P source existing in the surface water. This part of P is believed to have been  
358 released from dead algae degradation (Jarvie et al., 2008). Second, the dissolved P in  
359 surface water was trapped by the Fe oxides formed in SWI, where  $\text{O}_2$  diffused from  
360 surface water reacted with ferrous ions from deep soil (Ajmal et al., 2018; Rietra et al.,

361 2001).

362 The S behavior was distinct from Fe, Mn, P, and As (**Fig. 2**). The concentrations  
363 of S were high at flooded water and dropped with depth. Sulfur is believed to exist as  
364 S(VI) in oxic conditions (Wu et al., 2016), and biotic S(VI) reduction occurs when the  
365 redox potential in soils dropped to a highly reducing condition (after easily used  
366 electron acceptors were consumed, like O<sub>2</sub>, nitrate, Mn and Fe oxides) (Borch et al.,  
367 2010). The subsurface decrease of S is therefore presumably caused by S(VI)  
368 reducing bacteria in anoxic soils (Pester et al., 2012), which transformed mobile S(VI)  
369 to insoluble S(-II) minerals (e.g. FeS, FeS<sub>2</sub>) (Wu et al., 2016).

370 To the best of our knowledge, there is no reported analytical method that can  
371 simultaneously measure major and trace elements in the redox gradient zone across  
372 SWI. Traditionally, a large volume of subsamples are required to measure Fe, Mn and  
373 P by colorimetric methods or ICP-OES (Arsic et al., 2018; Rietra et al., 2001; Serrat,  
374 1998; Wang et al., 2019; Yi et al., 2019), S by IC (Keller-Lehmann et al., 2006), and  
375 most traces (e.g. As and antimony) by ICP-MS (Gustave et al., 2018a; Gustave et al.,  
376 2019). It is obviously beneficial for studies on element biogeochemical cycles when  
377 all the elements can be measured in one injection. The ICP-MS based method  
378 dramatically reduces the time consumption for multi-element, thus significantly  
379 increases the measurement throughput. The method can also be coupled with other  
380 HR samplers, for example, DET and HR peeper.

381

### 382 3.5 As, P, and S species measurement with IC-ICP-MS

383 The As, P, and S species vary along SWI and determine their environmental fates. In  
384 this study, simultaneous detection of four As, two P, and two S species in solution and  
385 anoxic soil porewater was achieved with the  $\text{NH}_4\text{HCO}_3$  as the mobile phase (**Fig. 3**).  
386 The retention times were 6.6, 7.5, 3.2, 3.9, 6.0, 10, 3.2, and 8.8 min for P(V), P(III),  
387 DMA, As(III), MMA, As(V), S(-II), and S(VI) respectively. The separation of four As  
388 species agrees well with previous work using the same chromatographic conditions  
389 (Suzuki et al., 2009). The results also demonstrated  $\text{NH}_4\text{HCO}_3$  mobile phase can be  
390 extended to measure P and S species. The determination coefficients for all the three  
391 elements are  $> 0.96$  (**Fig. S8**).

392 When applied to soil porewater, 2 As (As(III,V)), 1 P (P(V)), and 2 S (S(-II, VI))  
393 species were detected (**Fig. 3**). Arsenite, P(V), and S(-II) ( $> 70\%$ ) represent the  
394 dominant As, P and S in soil porewater, respectively. These results agree well with  
395 previous reports from multiple soils when the HPLC-ICP-MS, IC or colorimetric  
396 methods were used (Chen et al., 2019; Han et al., 2018; Xu et al., 2017). Besides, a  
397 susceptible P peak was detected in both solution and porewater with a retention time  
398 of 200 s (**Fig. 3**). The retention time is very close to the column dead time for the  
399 IC-ICP-MS. which indicates this compound may be cations or small molecules.  
400 Among the P species found in environmental samples, phosphine was the only one  
401 with neutral in charge, but it is only found in highly reducing environments (Han et al.,

402 2002), thus can be excluded. Therefore, we extrapolate the P peak at 200 s is likely an  
403 interference, such as  $^{47}\text{Ti}^+$ , which has the same mass-to-charge ratio of  $^{47}\text{PO}^+$ .

404 Many types of mobile phases were designed to measure the P, S, or As species  
405 (**Table 1**). Among the typical mobile phases, the carbonate-based mobile phase  
406 ( $\text{NH}_4\text{HCO}_3$ ) is better than the others due to its capability to separate the common  
407 species of As, P, and S in 10 min. Other chemicals were not suited for various reasons.  
408 The P or S containing compounds, like  $\text{K}_2\text{SO}_4$ , 1-butanesulfonic acid,  
409 1-hexanesulfonic acid and  $\text{NH}_4\text{H}_2\text{PO}_4$ , (Branch et al., 1989; Hirata and Toshimitsu,  
410 2007; Paik et al., 2010) were first excluded because of their interferences with P and S  
411 measurement. Nitrate-based mobile phase, such as  $\text{HNO}_3$  and  $\text{NH}_4\text{NO}_3$ , (Jackson and  
412 Bertsch, 2001; Paik et al., 2010) can be used to separate P and S species, but its  
413 elution strength is weaker than that of  $\text{CO}_3^{2-}$ , which must be used in strong acid  
414 condition to shorten the retention time (Jackson and Bertsch, 2001), or in neutral  
415 condition with very long running time (Vriens et al., 2014). The acidic condition is  
416 not ideal for S(-II) detection, because it encourages gaseous  $\text{H}_2\text{S}$  formation. The alkali  
417 hydroxyl-based mobile phase is widely used to As and S species measurement by  
418 IC-ICP-MS (Divjak and Goessler, 1999; Jackson and Bertsch, 2001). However, the  
419 hydroxyl-based mobile phase often contains alkali metal cations, like sodium, which  
420 can clog the plasma torch, sampling com, or skimmer com by the inorganic salts  
421 deposition. Thus,  $\text{NH}_4\text{HCO}_3$  is the best choice to develop a mobile phase for  
422 simultaneously separating P, S, and As species.



423 The  $\text{NH}_4\text{HCO}_3$  mobile phase pH (10) is appropriate to form P, S, and As species  
424 of different charges (Divjak and Goessler, 1999; McDowell et al., 2004; Reid et al.,  
425 2020). However, the alkali condition may cause Fe oxides precipitation when the  
426 mobile phase mixed with the high Fe porewater. The precipitates could bind with the  
427 anions, and potentially interfere with the testing and cause clog in the analytical  
428 column. Thus, EDTA was pre-added into samples before sample injection to the IC.  
429 EDTA could mitigate the precipitation of metal ions within a wide range of pH  
430 (Almkvist et al., 2013; Gallagher et al., 2001; Samanta and Clifford, 2006). With  
431 EDTA addition, no metal oxides precipitation was observed during IC-ICP-MS  
432 measurement.

433

### 434 **3.6 Profiling of P, S and As species across SWI**

435 The optimized IC-ICP-MS method was further verified with soil porewater collected  
436 by SWI profilers. The vertical changes of As(III), As(V), P(V), S(-II) and S(IV) are  
437 depicted in **Fig. 4**. Among the species, As(III), As(V), P(V) and S(-II) were low in top  
438 layers, and increased gradually from 7.9, 0.70, 60  $\mu\text{g}\cdot\text{L}^{-1}$  and 1.6  $\text{mg}\cdot\text{L}^{-1}$  at  $\sim 13$  mm  
439 below SWI to over 200, 50, 200  $\mu\text{g}\cdot\text{L}^{-1}$  and 15  $\text{mg}\cdot\text{L}^{-1}$  in deep soils respectively. The  
440 S(VI) change was consistent with the total S, which decreased along the soil depth  
441 (**Fig. 4**). The As, P and S trends agree with their fates in flooded soils, and the results  
442 revealed by other methods (Arsic et al., 2018; Han et al., 2018; Mcadams et al., 2016;  
443 Robertson et al., 2008).

444 According to the element profiles, we noticed that there was an overlap between  
445 dissolved Fe and S(-II) in pretty high concentrations at >15 mm below SWI (**Fig.**  
446 **2&4**). Ferrous ions are the main form of dissolved Fe in reducing condition, and not  
447 supposed to co-exist with S(-II) in aqueous phase in theory because they form  
448 insoluble FeS (Rickard, 2006). However, the co-existence of **ferrous ions** and S(-II)  
449 were frequently reported in many field samples (Pagès et al., 2011; Robertson et al.,  
450 2008; Robertson et al., 2009) and we speculate that the **ferrous ions** may be chelated  
451 with dissolved organic matters and stabilized in aqueous phase.

452 Although the redox-sensitive elements speciation has received much attention,  
453 most analytical methods for those species were developed for single element only  
454 (Divjak and Goessler, 1999; Han et al., 2018; Jackson and Bertsch, 2001). In  
455 comparison with those methods, the optimized method in the current study has higher  
456 measurement throughput, which can simultaneously measure As, P and S species  
457 within 10 min in a single run. The method is well suited for coupling with HR  
458 samplers to map multi-species profiles across SWI.

459

#### 460 **4. Conclusions**

461 Simultaneous measurement of multi-element (Fe, Mn, As, P, S) and multi-species  
462 (As(III), As(V), MMA, DMA, P(III), P(V), S(-II) and S(VI)), in volume-limited  
463 samples (~ 100 µL level), was achieved with the ICP-MS and IC-ICP-MS analysis.

464 The method demonstrated both high sensitivity ( $\mu\text{g}\cdot\text{L}^{-1}$  level) and high throughput.  
465 Combining the optimized analytical methods and a HR porewater sampler (i.e. IPI),  
466 we successfully measured the mm-scale co-distributions of multi-element (Fe, Mn, As,  
467 P and S) and multi-species (As(III), As(V), P(V), S(-II) and S(VI)) along SWI. With  
468 rapid measurement of multiple parameters from limited samples, the optimized  
469 analytical methods enable researchers to measure aqueous chemistry (including pure  
470 solution, surface water, porewater and groundwater) at high throughput. Moreover,  
471 the optimized analytical methods are very well suited for limited samples analysis in  
472 HR samplers (e.g. IPI, DET and peeper). The combination of the method and HR  
473 samplers can provide as much as possible aqueous parameters, thus facilitate studies  
474 of elements cycling in micro interfaces (e.g. SWI) of soils, sediments and other  
475 aqueous environments.

476

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483

484 **Appendix A. Supplementary data**

485 Supplementary data associated with this article can be found in the online version at

486 xxxxxx.

487

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701 precipitation of cadmium and arsenate at the Brushite– Fluid interface. *Environ. Sci. Technol.*  
702 53:194-202.

703 **List of tables**704 **Table 1** Separation of As, P and S species with IC-ICP-MS under different conditions

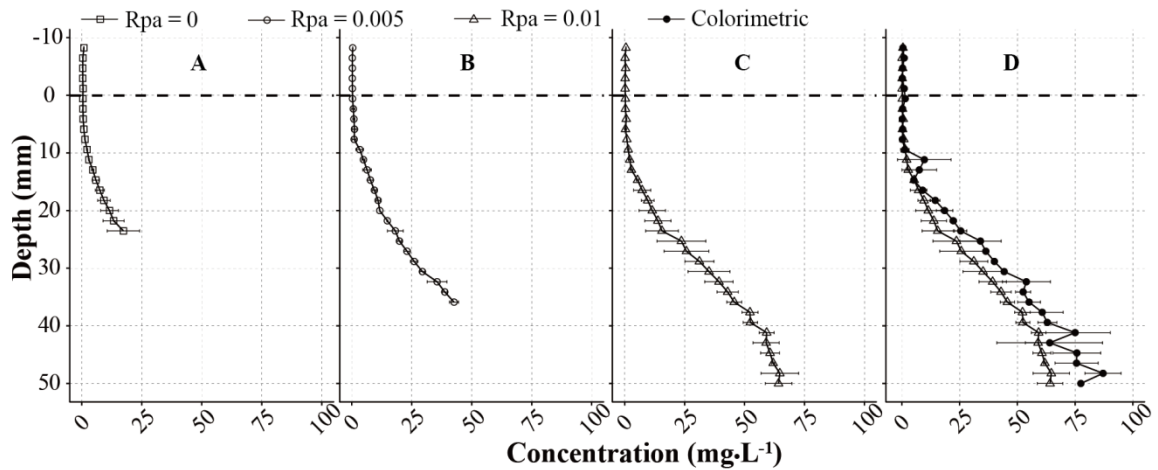
Mobile phase	pH	Species measured	Tr.§	Ref.
HNO <sub>3</sub>	1.3 - 2.6	As(III, V), MMA, DMA	5	(Jackson and Bertsch, 2001)
C <sub>4</sub> H <sub>10</sub> O <sub>3</sub> S <sup>†</sup>	3.0	As(V), DMA	5	(Hirata and Toshimitsu, 2007)
C <sub>6</sub> H <sub>14</sub> O <sub>3</sub> S <sup>‡</sup>	5.1	P(V)	5	(Chen et al., 2009)
NH <sub>4</sub> NO <sub>3</sub>	7.0	S(IV, VI), thiosulfate	5	(Lin and Jiang, 2009)
	7.5	As(III, V), MMA, DMA, S(VI)	15	(Vriens et al., 2014)
NH <sub>4</sub> NO <sub>3</sub> <sup>+</sup>	7.2	As(III, V), MMA, DMA	15	(Paik et al., 2010)
NH <sub>4</sub> H <sub>2</sub> PO <sub>4</sub>	+8.2			
	10	As(III, V), MMA, DMA	9	(Suzuki et al., 2009)
NH <sub>4</sub> HCO <sub>3</sub>	10	P(III, V), S(-II, VI)	10	This study
K <sub>2</sub> SO <sub>4</sub>	10.5	As(III, V), MMA, DMA	7	(Branchet et al., 1989)
(NH <sub>4</sub> ) <sub>2</sub> CO <sub>3</sub>	11.2	As(V), P(V), S(VI)	4	(Divjak et al., 1999)
		As(III, V), MMA, DMA	8	(Jackson and Bertsch, 2001)
NaOH	> 12	S(-II, IV, VI)	8	(Divjak and Goessler, 1999)

705 § measurement throughput (min per sample); † 1-butanefulfonic acid; ‡ 1-hexanesulfonic acid; arsenite,  
706 As(III); arsenate, As(V); monomethylarsonic, MMA; dimethylarsinic, DMA; hypophosphite, P(I);  
707 phosphite, P(III); phosphate, P(V); sulfide, S(-II); sulfite, S(IV); sulfate, S(VI).

708



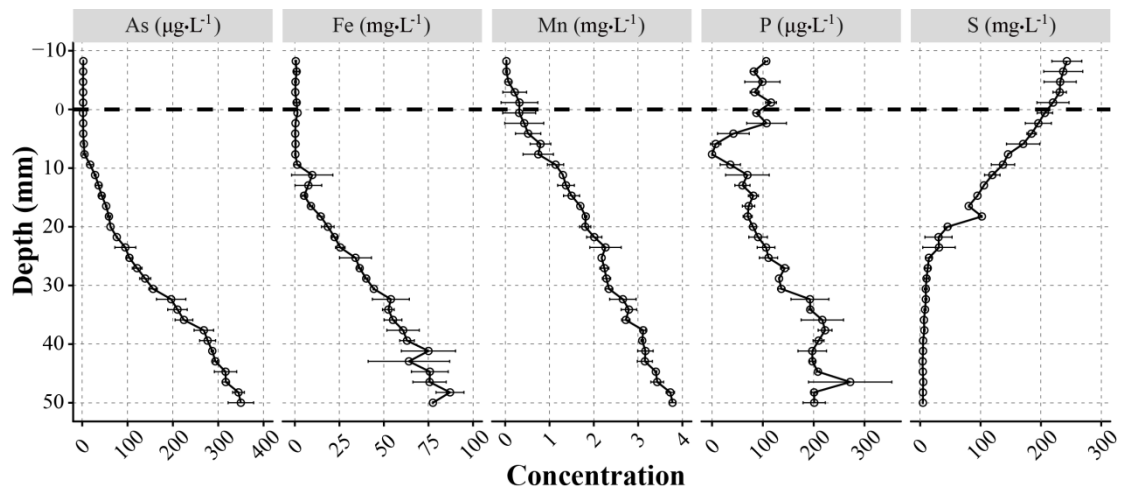
709 **List of figures**



710

711 **Figure 1** Iron profiles measured by ICP-MS under extended dynamic range mode or  
712 colorimetric method (phenanthroline) in Shaoguan (SG) paddy. A-C): Fe profile  
713 measurement with ICP-MS method when ‘rejection parameter a’ (Rpa) = 0, 0.005 and  
714 0.01 respectively; D) Fe profile measured by colorimetric and ICP-MS method (Rpa =  
715 0.01). **The triangle symbols in C) and D) share the same dataset.** The error bar is the  
716 standard deviation ( $n = 2$ ). The black dashed line at depth zero is the soil-water  
717 interface.

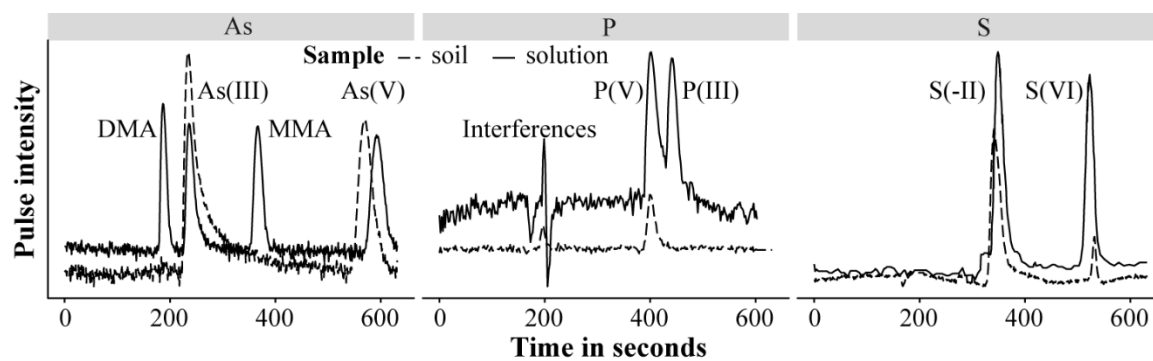
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720 **Figure 2** Mapping of As, Fe, Mn, P, and S profiles in Shaoguan (SG) paddy with  
 721 ICP-MS.

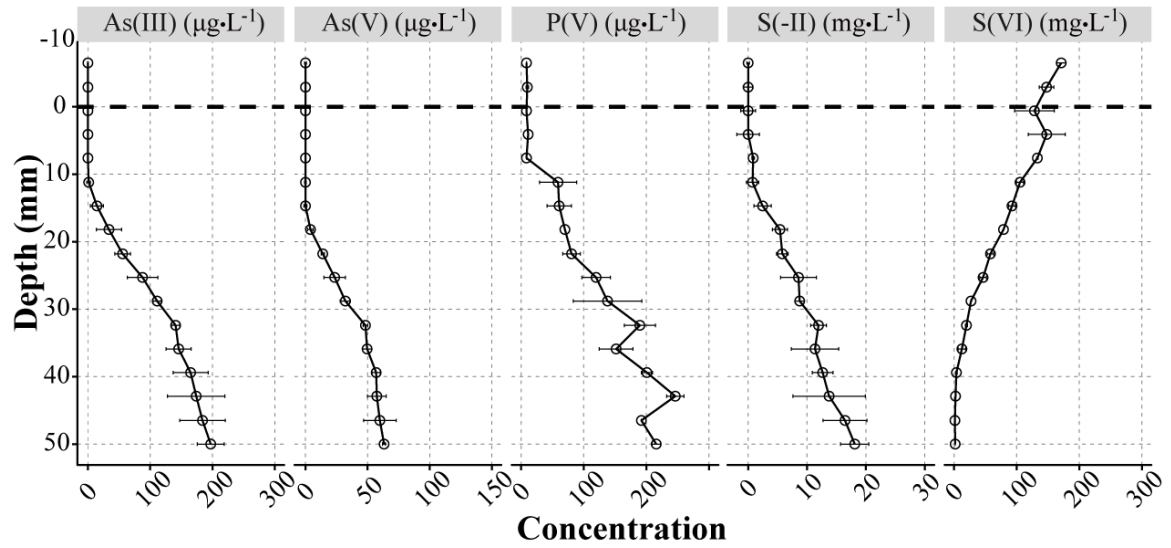
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724 **Figure 3** Arsenic (As), phosphorus (P), and sulfur (S) species detected by IC-ICP-MS.  
 725 The samples include soil porewater and solution. The solution was prepared in neutral  
 726 conditions (pH 7; 100  $\mu\text{g}\cdot\text{L}^{-1}$  phosphite (P(III)), phosphate (P(V)), arsenite (As(III)),  
 727 arsenate (As(V)), monomethylarsonic (MMA) and dimethylarsinic (DMA); 1  $\text{mg}\cdot\text{L}^{-1}$   
 728 sulfide (S(-II)) and sulfate (S(VI))).

729



730

731 **Figure 4** Profiles of P, S and As species in SG paddy measured by SWI profiler and  
 732  $\text{NH}_4\text{HCO}_3$  eluent.

733

734 **Simultaneous measurement of aqueous redox-sensitive elements and their species**  
735 **across the soil-water interface**

736 Zhao-Feng Yuan <sup>1, 2, 3</sup>, Williamson Gustave <sup>1, 2, 4</sup>, Raju Sekar <sup>5</sup>, Jonathan Bridge <sup>6</sup>, Jia-Yue Wang <sup>1</sup>,  
737 Wei-Jia Feng <sup>1</sup>, Bin Guo <sup>7\*</sup> and Zheng Chen <sup>1\*</sup>

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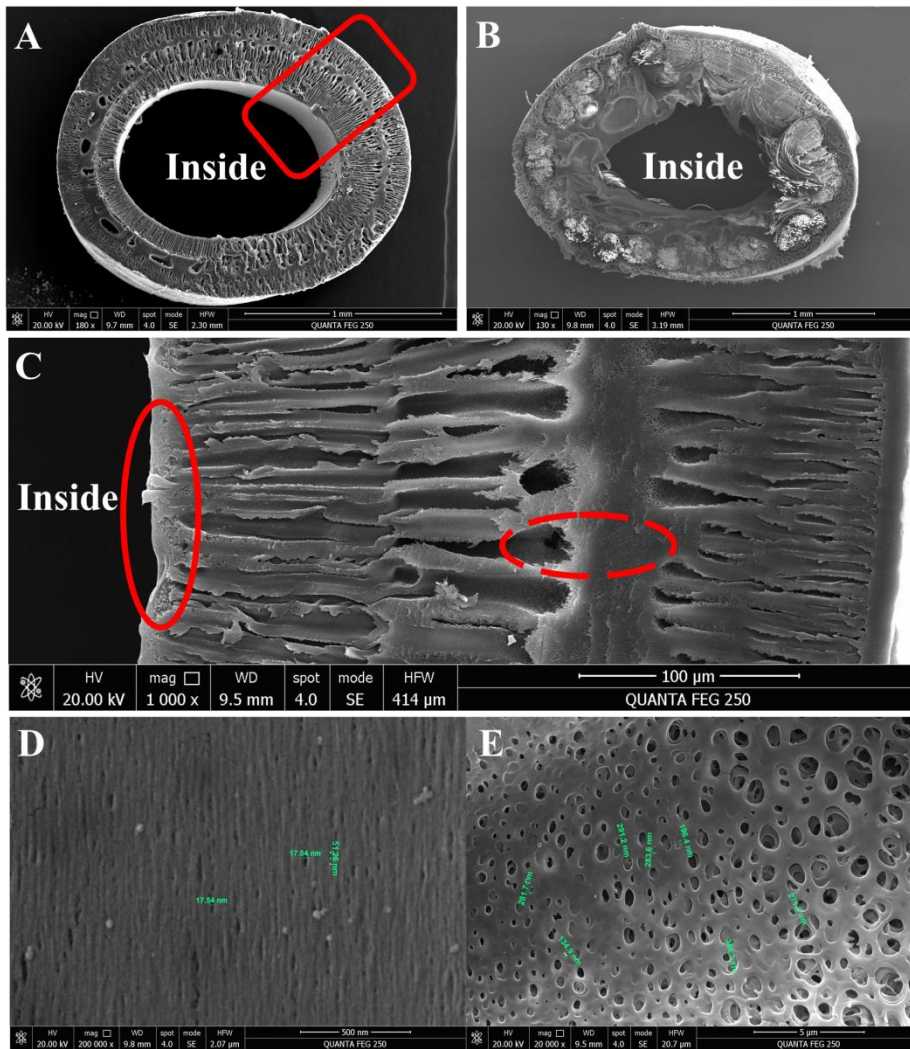
757 **Table S1** The selected physicochemical characteristics of Shaoguan (SG) paddy soils  
758 used in this study

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<b>Location</b>	<b>Soil type</b>	<b>As (mg kg<sup>-1</sup>)</b>	<b>Fe (g kg<sup>-1</sup>)</b>	<b>P (mg kg<sup>-1</sup>)</b>	<b>S (%)</b>	<b>Mn (g kg<sup>-1</sup>)</b>
<b>SG</b>	loam	146	227	40.0	0.40	1.50

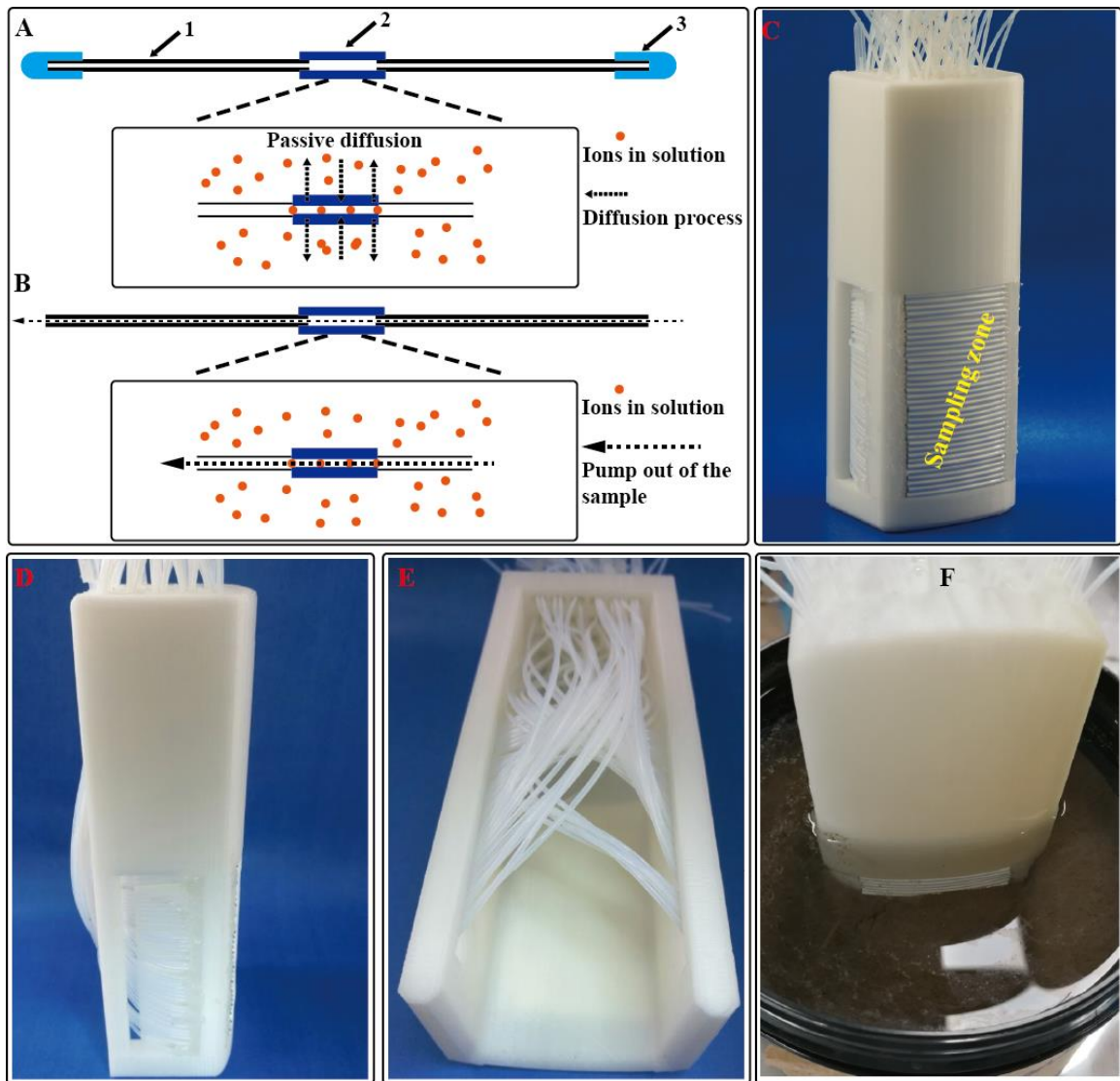
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760

761 **Figure S1** Pore size of the membrane determined by SEM. A-B) vertical section of  
 762 the novel modified polyethersulfone membrane and the old polyvinylidene fluoride  
 763 membrane respectively, the red frame is the selected region for further magnification  
 764 in C); and solid and dashed eclipses in C) are the selected region for pore size  
 765 determination in D) and E) respectively. The pore size of the novel membrane for  
 766 analytes is determined by its inside-layer with a pore size of ~20 nm.

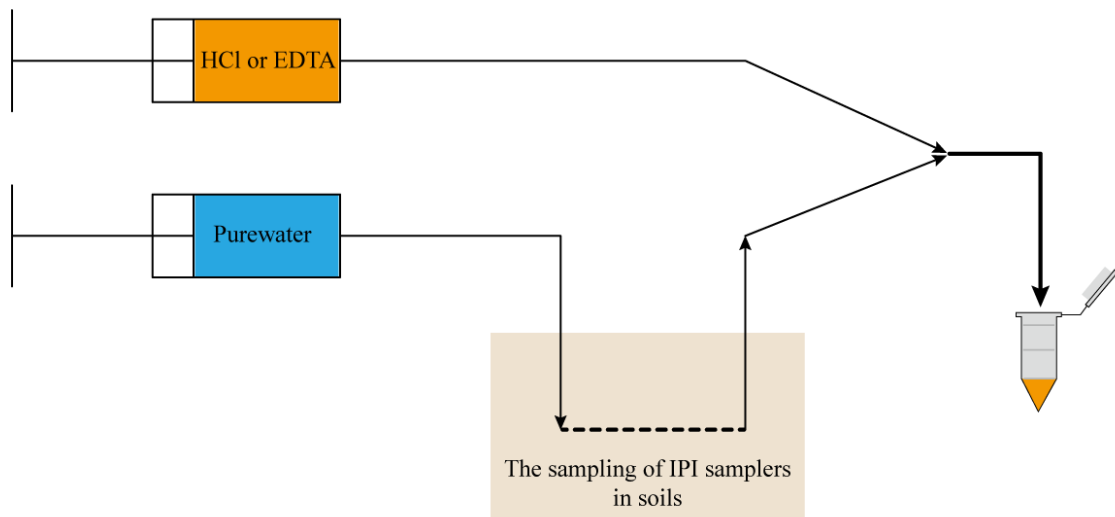


767

768 **Figure S2** Schematic diagram of *In-situ* Porewater Iterative (IPI) sampler. A)  
 769 elements ions diffuse through the hollow fiber membrane; B) the solution inside the  
 770 hollow fiber membrane tube is pumped out from the tube and collected; C) photo of  
 771 IPI sampler array (SWI profiler), SWI profiler consists of 34 IPI samplers, which are  
 772 assembled horizontally in the plastic holder; D) photo of IPI sampler (side); E) photo  
 773 of IPI sampler (back and bottom); F) photo of SWI profiler deployed in soils. Note  
 774 (A): 1. pipe; 2. hollow fiber membrane tube; 3. silicon cap.

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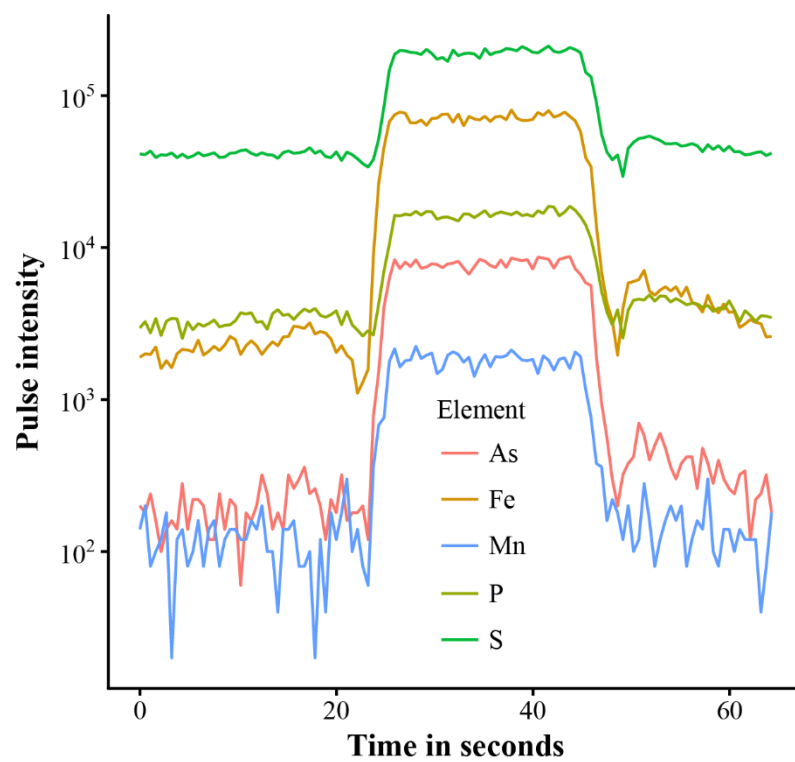




776

777 **Figure S3** The sampling process of IPI samplers in saturated soils. After the sampling  
 778 of the IPI sampler reaches equilibrium, the sample is collected powered by an  
 779 injection pump with pure water, and online mixed with the HCl or EDTA solution to  
 780 preserve the sample. The arrows indicate the flow direction.

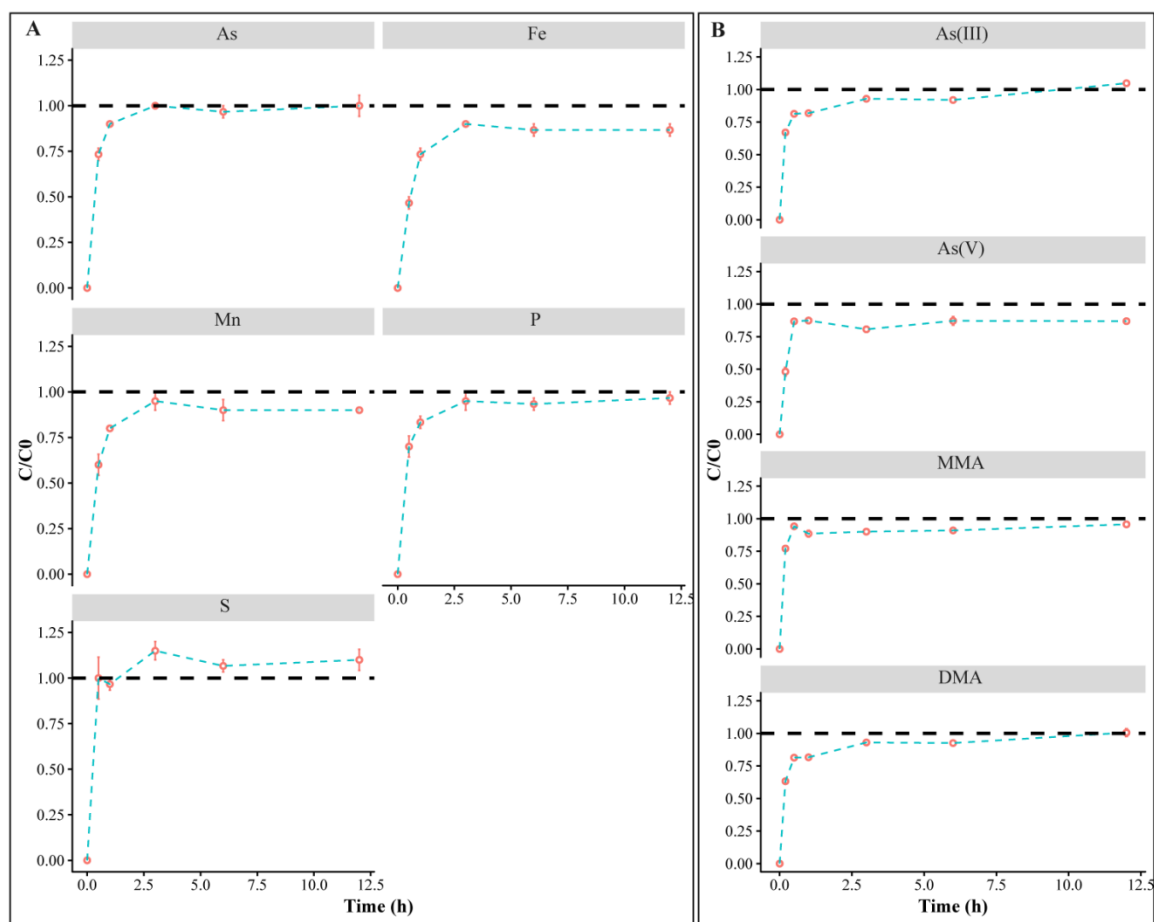
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782

783 **Figure S4** Measurement of As, Fe, Mn, P, and S in 50  $\mu\text{L}$  solution by data only  
784 analysis in ICP-MS. The solution was prepared in 2%  $\text{HNO}_3$ , with 10  $\mu\text{g}\cdot\text{L}^{-1}$  As and  
785 Mn, 100  $\mu\text{g}\cdot\text{L}^{-1}$  P, and 1.0  $\text{mg}\cdot\text{L}^{-1}$  Fe and S.

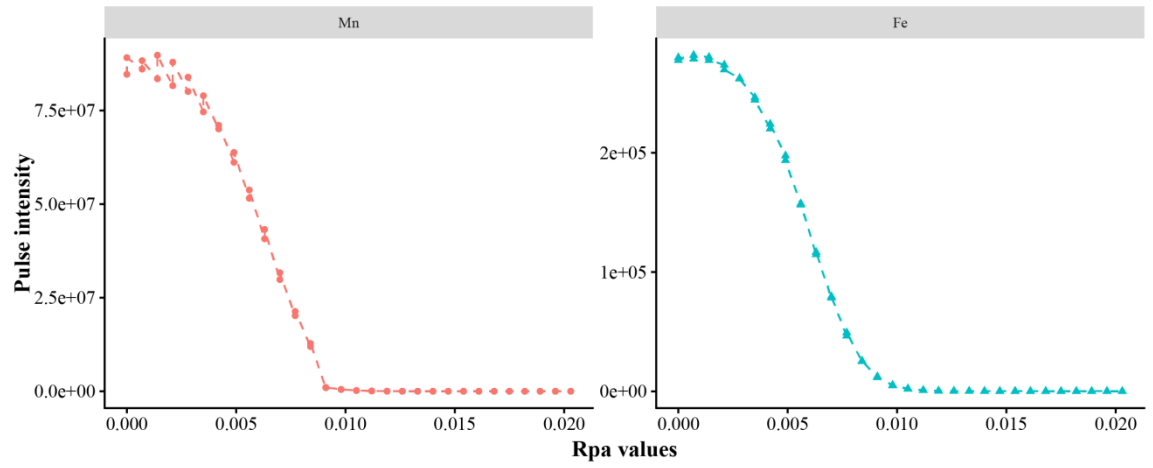
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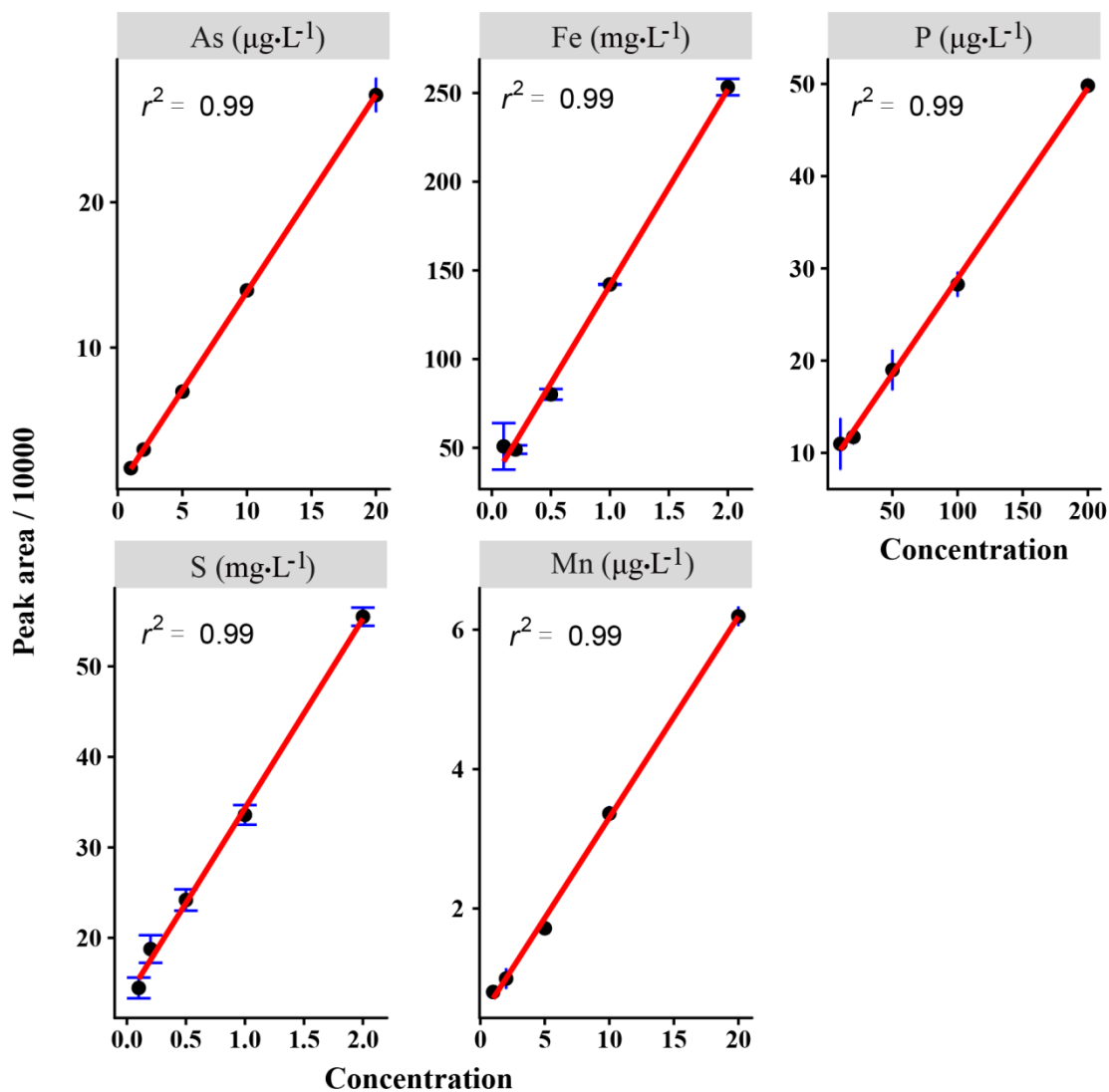
788 **Figure S5** The dynamic response of total elements (A) and redox species (B) in IPI  
 789 samplers (normalized element concentration in solutions). A) total elements include  
 790 arsenic (As), iron (Fe), manganese (Mn), phosphorus (P) and sulfur (S) measured by  
 791 ICP-MS; B) four As species, including arsenite [As(III)], arsenate [As(V)],  
 792 monomethylarsonic (MMA) and dimethylarsinic (DMA), were tested by IC-ICP-MS.  
 793 The error bar is the standard deviation ( $n = 3$ ).

794



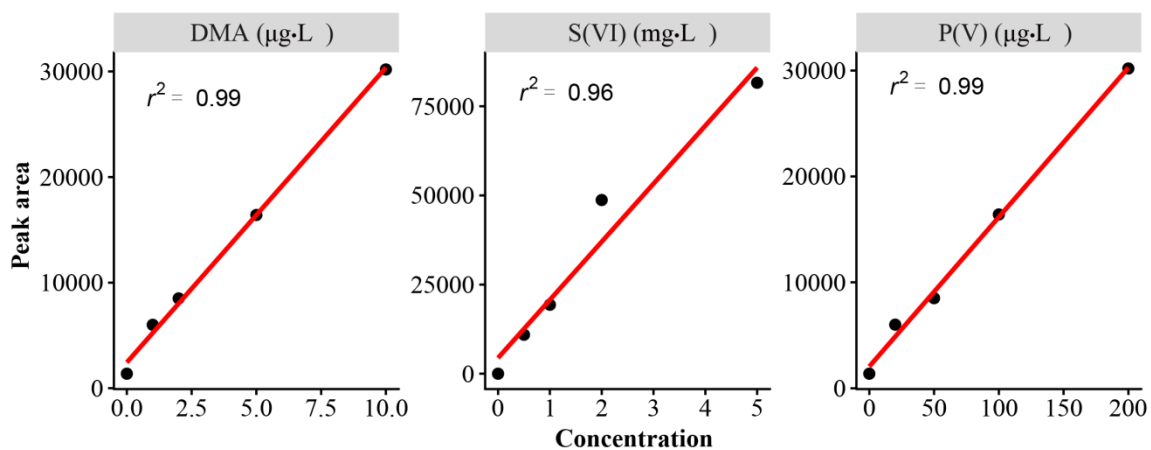
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796 **Figure S6** Response of Fe and Mn to different Rpa values.



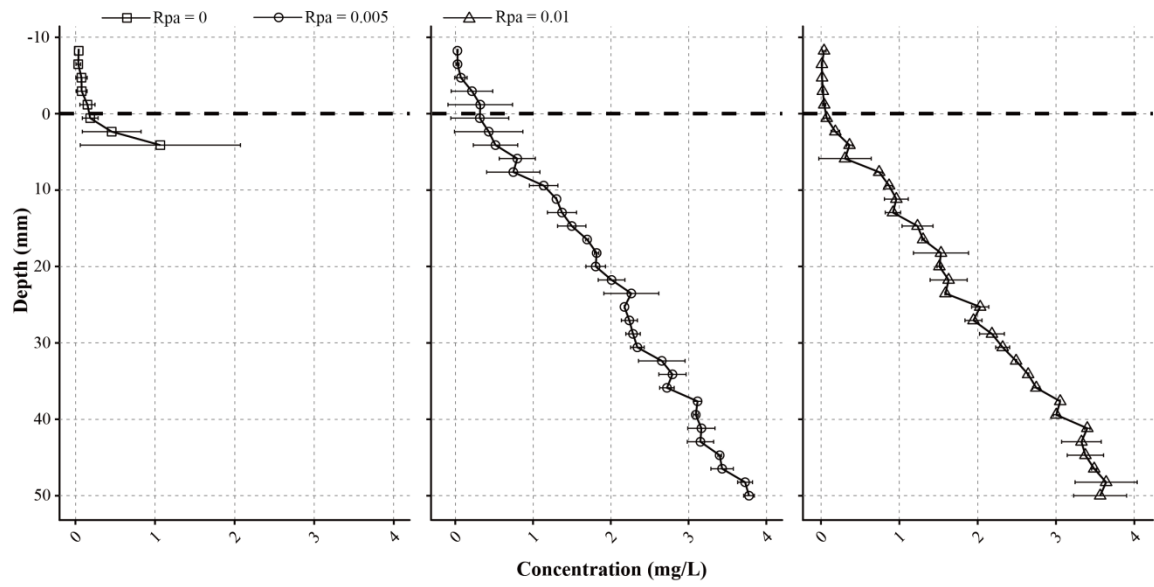
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798 **Figure S7** The calibration curve of As, Fe, Mn, P, and S ( $n = 3$ ). The standards,  
 799 containing 1.0/10/100, 2.0/20/200, 5.0/50/500, 10/100/1000, 20/200/2000 µg·L<sup>-1</sup>  
 800 As&Mn/P/Fe&S, were prepared in 2% HNO<sub>3</sub>. The error bar is the standard deviation.



801

802 **Figure S8** The calibration curve of dimethylarsinic (DMA), sulfate (S(VI)) and  
 803 phosphate (P(V)). The standards, containing 0/0/0, 1.0/20/500, 2.0/50/1000,  
 804 5.0/100/2000, 10/200/5000  $\mu\text{g}\cdot\text{L}^{-1}$  DMA/P(V)/S(VI), were prepared under neutral  
 805 conditions.



806

807 **Figure S9** Manganese (Mn) profile measured by ICP-MS with different Rpa values in  
 808 Shaoguan (SG) paddy. Rpa values were set at 0, 0.005, and 0.01 respectively. The  
 809 error bar is the standard deviation ( $n = 2$ ).