

1        **Hysteresis of natural magnetite ensembles: Micromagnetics of**  
2        **silicate-hosted magnetite inclusions based on focused-ion-beam**  
3        **nanotomography**

4        **Even S. Nikolaisen<sup>1</sup>, Richard Harrison<sup>1,2</sup>, Karl Fabian<sup>1</sup>, Suzanne A. McEnroe<sup>1</sup>**

5        <sup>1</sup>Norwegian University of Science and Technology (NTNU), S. P. Andersens veg 15a, 7031 Trondheim, Norway

6        <sup>2</sup>Department of Earth Sciences, University of Cambridge, Downing Street, Cambridge CB2 3EQ, UK

7        **Key Points:**

- 8        • The 3D geometries of more than 400 silicate-hosted magnetite inclusions were mea-  
9        sured using focused ion-beam nanotomography
- 10       • Micromagnetic modelling show that inclusions adopt multiple domain states, span-  
11       ning the entire range of common domain-state diagnostic plots
- 12       • Geometric variability of naturally formed magnetite in silicate, explains scatter of  
13       magnetic parameters in the PSD region of the Day plot

---

Corresponding author: Even Nikolaisen, [even.nikolaisen@ntnu.no](mailto:even.nikolaisen@ntnu.no)

## 14 Abstract

15 Three-dimensional geometries of silicate-hosted magnetic inclusions from the Harcus  
 16 intrusion, South Australia have been determined using focused-ion-beam nanotomography  
 17 (FIB-nt). By developing an effective workflow, the geometries were reconstructed for mag-  
 18 netic particles in a plagioclase (162) and a pyroxene (282), respectively. For each inclusion,  
 19 micromagnetic modelling using MERRILL provided averaged hysteresis loops and backfield  
 20 remanence curves of 20 equidistributed field directions together with average  $M_s$ ,  $M_{rs}$ ,  $H_c$ ,  
 21 and  $H_{cr}$ . The micromagnetic structures within each silicate are single-domain, single-vortex,  
 22 multi-vortex and multi-domain states. They have been analyzed using domain-state diagnos-  
 23 tic plots, such as the Day plot and the Néel plot. SD particles can be subdivided into groups  
 24 with dominant uniaxial anisotropy ( $M_{rs}/M_s \sim 0.5$  and  $10 < H_c < 100$  mT) and mixed  
 25 uniaxial/multiaxial anisotropy ( $M_{rs}/M_s \sim 0.7$  and  $10 < H_c < 30$  mT). Most single-vortex  
 26 particles lie on a trend with  $0 < M_{rs}/M_s < 0.1$  and  $0 < H_c < 10$  mT, while others dis-  
 27 play a broad range of intermediate  $M_{rs}/M_s$  and  $H_c$  values. Single-vortex and multi-vortex  
 28 states do not plot on systematic grain-size trends. Instead, the multi-component mixture of  
 29 domain states within each silicate spans the entire range of natural variability seen in bulk  
 30 samples. This questions the interpretation of bulk average hysteresis parameters in terms  
 31 of grain size alone. FIB-nt combined with large-scale micromagnetic simulations provides  
 32 a more complete characterization of silicate-hosted carriers of stable magnetic remanence.  
 33 This approach will improve the understanding of single-crystal paleomagnetism, and enable  
 34 primary paleomagnetic data to be extracted from ancient rocks.

## 35 1 Introduction

36 A fundamental task in rock magnetism is to identify the magnetic domain states adopted  
 37 by natural remanence carriers, because these control the remanence acquisition process, the  
 38 stability of the remanent magnetization over geological time, and subsequently the reliability  
 39 of the stored paleomagnetic information. Most rocks contain a mixture of magnetic minerals  
 40 covering a broad range of particle sizes and shapes. Rock magnetism broadly classifies these  
 41 particles into superparamagnetic (SP), stable single-domain (SD), pseudo-single-domain  
 42 (PSD), and multi-domain (MD) states, with the boundaries between states originally based  
 43 on experimental results [McNab *et al.*, 1968; Dunlop, 1973; Dunlop and Bina, 1977; Soffel,  
 44 1977]. Based on the magnetic material properties of magnetite, these observed boundaries  
 45 could be related to theoretical calculations for regular particle geometries [Butler and Baner-  
 46 jee, 1975; Argyle and Dunlop, 1984; Enkin and Dunlop, 1987; Enkin and Williams, 1994],  
 47 although the influence of shape has been recognized [Stacey, 1961; Fabian and Hubert,  
 48 1999; Witt *et al.*, 2005]. Theoretical values for the SP-SD and SD-PSD transition sizes agree  
 49 very well with the experimentally determined boundaries of 0.025-0.030  $\mu\text{m}$  and 0.050-  
 50 0.084  $\mu\text{m}$  for crushed and grown magnetite crystals. Experimental studies are limited to syn-  
 51 thetic or natural particle ensembles with high variability in size and shape with the exception  
 52 of a study on synthetic samples generated by lithography, where monocrystallinity could not  
 53 be ensured [Krása *et al.*, 2009]. Magnetite particles directly above the upper SD limit as-  
 54 sume an intermediate state in experimental and theoretical studies. Because they appear to  
 55 be related to a stable and single-domain like experimental behavior, for which Stacey [1961]  
 56 coined the notion of PSD remanence, they are commonly referred to as PSD particles. From  
 57 a theoretical view such particles are intermediate between two characteristic length scales,  
 58 the SD-limit of  $\approx 7 \lambda_{ex}$ , where  $\lambda_{ex} = \sqrt{2A/(\mu_0 M_s^2)}$  is the exchange length, and the domain  
 59 wall width of  $\approx 5 \sqrt{A/K_1}$ . Here  $A$  denotes the exchange constant,  $K_1$  the magnetocrystalline  
 60 anisotropy constant,  $M_s$  the saturation magnetization, and  $\mu_0$  the vacuum permeability. For  
 61 magnetite the intermediate size range lies roughly between 0.070  $\mu\text{m}$  and 0.200-0.300  $\mu\text{m}$ ,  
 62 however, the formation of clearly defined homogeneously magnetized domains does not oc-  
 63 cur below grain sizes of about 2  $\mu\text{m}$ . Nagy *et al.* [2019] modelled grains up to the micron  
 64 range and observed domain wall formation in unconstrained micromagnetic models. What  
 65 is missing, is a comprehensive survey of magnetization states and predicted hysteresis prop-



66 erties for naturally occurring magnetite particles across, and beyond, the SD state. Here we  
 67 develop a method to achieve such a survey based on FIB tomography combined with micro-  
 68 magnetic modeling.

69 To compare our model results to measurements on sized magnetite samples, we con-  
 70 sider several isothermal measurements that have been developed to quantify the character-  
 71 istic domain states of natural remanence carriers. The most common approach is based on  
 72 the measurement of bulk average magnetic hysteresis parameters, such as saturation mag-  
 73 netisation  $M_S$ , saturation remanent magnetisation  $M_{rs}$ , coercivity  $H_c$ , and coercivity of re-  
 74 manence  $H_{cr}$ . Two intensely studied diagrams are the Day plot of  $M_{rs}/M_S$  versus  $H_{cr}/H_c$   
 75 [Day *et al.*, 1977], and the Néel plot of  $M_{rs}/M_S$  versus  $H_c$  [Néel, 1955], which are based on  
 76 bulk hysteresis parameters that can be quickly and easily measured in most rock magnetic  
 77 laboratories [Dunlop, 2002; Tauxe *et al.*, 2002]. It is well known that their interpretation is  
 78 ambiguous without detailed knowledge of the underlying magnetic ensemble [Roberts *et al.*,  
 79 2018]. To relate domain state to geometrical measurements, the theoretical diagram of *But-  
 80 tler and Banerjee* [1975], and later revisions [Fabian *et al.*, 1996; Muxworthy and Williams,  
 81 2006], is commonly used to characterise the domain state of magnetic particles based on par-  
 82 ticle length and axial ratio. This approach assumes that magnetic particles can be effectively  
 83 approximated as prolate ellipsoids.

84 Here we apply our workflow and method to magnetite particles within silicate-hosted  
 85 magnetic inclusions. This setting is abundant and important for studying the Earth's mag-  
 86 netic field, magmatic evolution, geochemical interaction, stability and oxygen fugacity [Fein-  
 87 berg *et al.*, 2005; Tarduno *et al.*, 2020]. Magnetite is the most significant Fe-oxide for paleo-  
 88 magnetic studies, and magnetite is known to form as exsolved or included particles in silicates  
 89 like plagioclase [Davis and Letters, 1981; Feinberg *et al.*, 2005; Usui *et al.*, 2015] and py-  
 90 roxene [Fleet *et al.*, 1980; Frandsen *et al.*, 2004; Renne *et al.*, 2002; Feinberg *et al.*, 2004]  
 91 within a wide range of geologic settings. Davis and Letters [1981] document that magnetite  
 92 exsolves several micrometer long rods in 2-4 different directions in plagioclase from oceanic  
 93 gabbros. The crystallographic relationship between magnetite needles and a plagioclase host  
 94 was determined by [Wenk *et al.*, 2011]. They observed a single set of magnetite needles ori-  
 95 ented parallel to the [001] axis of plagioclase, with the needle elongation direction corre-  
 96 sponding to the [110] direction of magnetite and the (111) planes of magnetite oriented sub-  
 97 parallel to (120) and  $(\bar{1}20)$  planes of plagioclase. Within pyroxene (CPX) iron oxide inclu-  
 98 sions occur parallel to (010), where all inclusions are elongated and with the long axis sub-  
 99 parallel to both [100] and [001] [Renne *et al.*, 2002; Feinberg *et al.*, 2004]. In orthopyroxene  
 100 exsolved oxide lamellae are known from optical and transmission electron microscopy. Il-  
 101 menite lamellae are present as coherent rods and blades where (001) ilmenite is epitaxially  
 102 intergrown on (100) pyroxene, and the a-axis of ilmenite is parallel to the c-axis of pyroxene  
 103 [McEnroe *et al.*, 2004]. There the ilmenite contained hematite lamellae parallel to the (001)  
 104 of the ilmenite. The incorporation of iron into the silicates is not confirmed, but suggested  
 105 to occur at primary subsolidus mineral formation and exsolve magnetite well above its curie  
 106 temperature [Feinberg *et al.*, 2004; Fleet *et al.*, 1980] for pyroxene. The exact process for  
 107 exsolving magnetite in plagioclase is so far not established, but once the orientation relation-  
 108 ship is established the phase boundary between plagioclase and magnetite can be established  
 109 [Wenk *et al.*, 2011].

110 The advent of high-resolution focused-ion beam nanotomography (FIB-nt), combined  
 111 with the development of powerful finite-element micromagnetic (FEM) software optimised  
 112 for rock magnetic applications [Ó Conbhuí *et al.*, 2018; Fabian and Shcherbakov, 2018], pro-  
 113 vides a new opportunity to perform domain-state diagnosis from first principles. FIB-nt is  
 114 performed on a dual-beam electron microscope that combines the high-resolution imaging  
 115 capabilities of a field-emission gun scanning electron microscope (FEG-SEM) with the nm-  
 116 precision slicing capabilities of a FIB [Fagerland, 2014; Brogden, 2015]. This technique is  
 117 widely used to analyse biological samples [Beckwith *et al.*, 2015; Guehrs *et al.*, 2017; Mul-  
 118 ders *et al.*, 2006], alloys [Cao *et al.*, 2009; Ding and Jones, 2011] and geological materials

[Einsle *et al.*, 2016; ter Maat *et al.*, 2020; Warr *et al.*, 2014; Lascu *et al.*, 2015]. The process is destructive, as the slicing destroys the prepared sample area. However, it has higher spatial resolution than comparable non-destructive methods of X-ray computed tomography, enabling the morphology of particles spanning the SD to MD range to be reconstructed in 3D. Rather than attempting to solve the inverse problem of unmixing domain states from a series of magnetic hysteresis measurements, the combined FIB-nt-FEM approach enables forward models of domain states and hysteresis properties to be calculated from direct 3D observations of a representative particle ensemble. Evaluating magnetic signatures in geologic samples will always be challenging. The complexities of particle size, shape, spacing, chemical composition, stoichiometry and stress will generate unique mineralogical settings, and it is a fair question raised by Lascu *et al.* [2015] whether domain-state diagnosis might be an unachievable ideal? Here we show that FIB-nt-FEM is able to resolve domain-state ambiguities associated with the size, shape and spacing of particles. In combination with other techniques, such as energy-dispersive X-ray spectroscopy (EDS) and high-resolution electron backscattered diffraction (HR-EBSD), it has the potential to resolve ambiguities associated with chemical composition, stoichiometry and stress in the future.

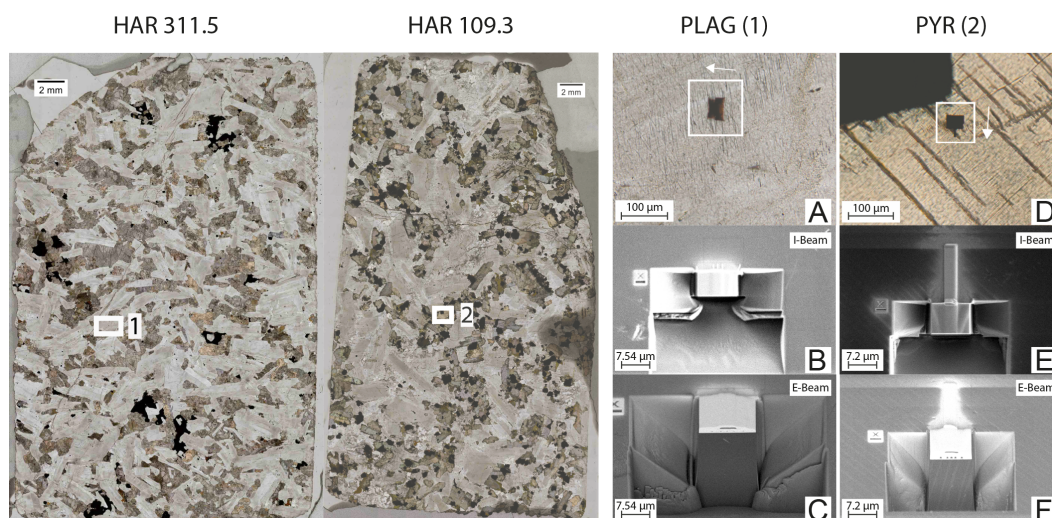
## 2 Materials and Methods

### 2.1 Samples

This study focuses on two gabbroic norite samples obtained from drill cores of the Mount Marcus Intrusion, which is a part of the Giles Complex, South Australia. The sample names 311.5 and 109.3 in Figure 1 correspond to their depth in m. The Giles-Event intrusions were emplaced into the Proterozoic Musgrave Province at about 1090-1040 Ma [Maier *et al.*, 2015]. The Marcus intrusion is of interest because of its striking *remanent* magnetic anomaly and the potential for magmatic nickel sulphide mineralizations [Austin *et al.*, 2014]. The high-coercivity remanent magnetization is interpreted to be carried by single-domain magnetite [Church *et al.*, 2016]. In reflected light the two thin sections in Figure 1 display a wide distribution of magnetite grain sizes, from large, discrete ( $> 100 \mu\text{m}$ ) multi-domain magnetite with oxy-exsolution of ilmenite lamellae and spinel needles and blades. Reduction exsolution of magnetite from ilmenite in the form of blades occurs in isolated ilmenite grains, and in some large ilmenite lamellae, which were first oxy-exsolved from magnetite [Church *et al.*, 2016; Robinson *et al.*, 2016]. Silicate-hosted micrometer- to nanometer-sized inclusions of magnetite are found in feldspars, pyroxenes and amphiboles ( $\ll 0.1 \mu\text{m}$ ). Primary silicates contain the highest concentrations of iron-oxide inclusions and were targeted for 3D imaging using FIB nanotomography. Both thin sections 311.5 and 109.3 are composed of plagioclase, pyroxene, amphibole, intragranular quartz and discrete iron-titanium oxides. Iron-oxide inclusions are present in primary plagioclase and pyroxene and secondary amphibole. Sample 311.5 has a high concentration of inclusions in plagioclase and amphibole. Metamorphic processes have possibly redistributed iron oxides in amphiboles. The intragranular quartz in association with amphibole also suggests the presence of fluids that may have mobilized Fe, and removed it from the system. Sample 109.3 has a lower amount of amphibole and iron-oxide inclusions in plagioclase balanced by an increase of iron-oxide inclusions in the pyroxene. For samples 311.5 and 109.3 the bulk natural remanent magnetization (NRM) values are 34 A/m and 4 A/m, and their bulk magnetic susceptibility values are 0.06 (SI) and 0.10 (SI), respectively.

### 2.2 Focussed-Ion-Beam Slice-and-View Nanotomography

Raw data collection was performed using the Auto Slice & View 4 software on an FEI Helios G4 UX Dual-Beam Focused-Ion-Beam (FIB) microscope at the NTNU NanoLab in Trondheim, Norway. Polished thin sections were coated with 10 nm of Au and the regions of interest were analysed with EDS to confirm that the target iron oxide inclusions were magnetite with minor ( $< 1 \text{ wt}\%$ ) Ti. At the optimal imaging settings for the selected minerals, the



163 **Figure 1.** Thin sections 311.5 and 109.3 with slice-and-view locations 1 (plagioclase) and 2 (pyroxene).  
 164 White arrows in (a) and (d) are aligned with the FIB slice-and-view milling direction (Z-axis). (b) and (e)  
 165 Ion-beam view of the milling volumes outlined by the platinum protective layer. (c) and (f) Electron-beam  
 166 view of the cross section (front face) of the milling volume. The electron beam is oriented at 52° to the ion  
 167 beam.

174 electron beam drift was about 500 nm an hour, which limited the size of the block volume  
 175 that could be successfully imaged. A block volume of material measuring 12x14x12 μm (L-  
 176 W-H), sliced with 30 nm increments, took, on average, 20 hours. With longer sessions, the  
 177 failure rate increased substantially. Block volumes were also limited by hardness of the ma-  
 178 terial and the ability to successfully prepare the trenches and clean the front face of the sliced  
 179 volume. The extracted volume of pyroxene is therefore smaller than for plagioclase since the  
 180 ion-beam has fewer issues with removing material.

181 Sample preparation for slice-and-view was performed by manual instrument opera-  
 182 tion, outside of the automated software. The selected area was first coated with 500 nm of  
 183 Pt. Five 200 nm deep lines that extend the full depth of the milling area were cut into the Pt  
 184 coating and then filled by carbon deposition using the ion beam. The selected area was then  
 185 covered with a further 300 nm of Pt. The central three carbon rods were orientated parallel to  
 186 the slice direction and used as image alignment references. The outside two carbon rods were  
 187 oriented at an angle of  $\pm 60^\circ$  to the slice direction and were used to calibrate the slice thick-  
 188 ness and to enable alignment of images in the slice direction in cases where there are missing  
 189 images, or if a milling session had to be restarted. The surface preparation was finalized with  
 190 3 μm wide and 500 nm thick carbon rectangle, oriented parallel to the slice direction, and  
 191 covered with an additional 300 nm of Pt. This carbon slab is used to keep the E-beam image  
 192 aligned with the front face at all times. Trenches, 15 μm deep, 10-13 μm wide, and 10 μm  
 193 long, were created to the left and right of the target volume, and a trench 20-25 μm wide and  
 194 20 μm long was created in front of the target volume using a standard cross section milling  
 195 routine at 30 keV and 20 nA. The front and sides were then polished with cleaning cross sec-  
 196 tions, first at 1.2 nA and then with a final cleaning step at 0.75 nA. A Pt fiducial mark for ion  
 197 beam alignment was placed on the thin section surface outside of the milled trenches as seen  
 198 in I-beam and E-beam images in Figure 1.

199 For slicing, the Ga-ion beam was kept at 30 keV and 0.75 nA and slice thickness was  
 200 chosen between 10 and 30 nm. Electron-beam imaging applies a mirror detector for the  
 201 Backscattered-Electron (BSE) signal in immersion mode at 5 kV and 0.8 nA. Tilt correc-

tion for the 52° angle between the ion- and electron-beam was performed in the FEI software with dynamic focus by activating the tilt-correction option. Each scan with 6144x4096 pixels at 10  $\mu$ s dwell time, resulted in a 1535x1652 image with pixel size of 7.5 nm in the pyroxene, a 2037x2043 image with pixel size of 7.6 nm in plagioclase. Early trials included the automatic z-axis drift equal to the milling depth and used a fiducial mark outside the milling area for consistent E-beam alignment. However, because this instrument has a constant E-beam drift that is not accounted by the software, it was found to be much more efficient to use the carbon slab as a fiducial mark that remains inside the milling area. This approach reduced the need for additional automated focusing steps after each slice. The high contrast between platinum and carbon, reduces the risk of losing the fiducial mark as the project runs, and increases the success rate of the automated procedure.

### 2.3 Image and object filtering

The raster images of the milled front faces are provided in Tagged Image File Format (TIF) and were stacked in Fiji (ImageJ) and aligned with the stack alignment plugin [Tseng *et al.*, 2011], using the cross-section of the three central carbon trenches as reference pattern. Using their high electron reflectivity and thus brightness, magnetite particles were extracted by gray scale thresholding and exported as binary images. Based on backscatter intensity and cross-referenced with the EDS analyses anything that we could not confirm as magnetite was segmented out. In the case of complex structures and overlapping threshold values, each slice was manually "cleaned" by going through the binary and image stack (side by side), removing any artifacts, or segmented threshold values that were not magnetite. For the pyroxene stack, the thresholding was particularly difficult because the bottom sections of each image had lower overall backscatter intensity than the upper parts. This image stack had to be treated with a non-local means filter, and was segmented using the Dragonfly software from ORS. After the binary segmentation, the stack was passed through 3x3x3 pixel 3D-mean and Gaussian-blur filters to smooth over the pixel boundaries, and to reduce sharp edges in objects at the boundary of the originally cropped volume. Areas with objects in close proximity to each other were manually edited to prevent bridging artifacts when converted to a finite element surface mesh. We excluded incomplete particles at the edge of the milled volume, and particles that are too small to be confidently distinguished from background and noise.

### 2.4 Mesh generation

The stack generated in Fiji was imported to Paraview [Ahrens *et al.*, 2005], then resized based on pixel length, width and slice thickness to regenerate the actual size of the milled volume. A surface mesh of the magnetic particles was generated using a contour filter based on the magnetite brightness threshold. The resulting initial mesh was exported as a stereolithography (STL) file. Both, Meshmixer [Schmidt and Singh, 2010] and Meshlab [Cignoni *et al.*, 2008] were then used to improve the meshes. The quality of each object was evaluated individually. An object was discarded if it did not intersect at least two slices. This imposes a lower-size limit of approximately 30 nm for particles to be modeled in 3D using this method. For geological time-scales, this corresponds approximately to the superparamagnetic transition size for magnetite. The geometric mesh quality for each object was checked, and any holes filled to generate a closed mesh surface. Surface mesh node-density was reduced to generate a more smoothed surface that minimizes the appearance of artificial steps reflecting the finite distance between the measured slices. After eliminating the voxel steps between the slices, mesh node-density was again increased to generate a now smoothed mesh with a target edge length  $\leq 8.9$  nm, the exchange length of magnetite. Target edge length, mesh quality and volume were checked in meshlab. If the mesh still contained artificial sharp edges due to the object's small size, it was remeshed and smoothed in Meshmixer to fit the original size. Iso2mesh and a Matlab script [Fang and Boas, 2009] were used to generate tetrahedral volume meshes from the surface meshes. Also here a target edge length  $\leq 8.9$  nm was used. To keep the node density similar for each object, the edge length was reduced for smaller ob-



jects. The output file was converted to a Patran file using Git Bash (Windows computer) to call a shell script (Convert2Pat) [Ó Conbhuí *et al.*, 2018]. The Patran file is then called by a script to be used by MERRILL for computing the magnetization structures. Aspect ratios of the final meshes were computed from the STL files based on the volume inertia tensor, by computing its eigenvalues and eigensystem by a Python routine. To ensure correct calculations, the proportional relationship between the eigenvalues was tested by comparing it to the measured short and long axis of the same particles in Meshlab.

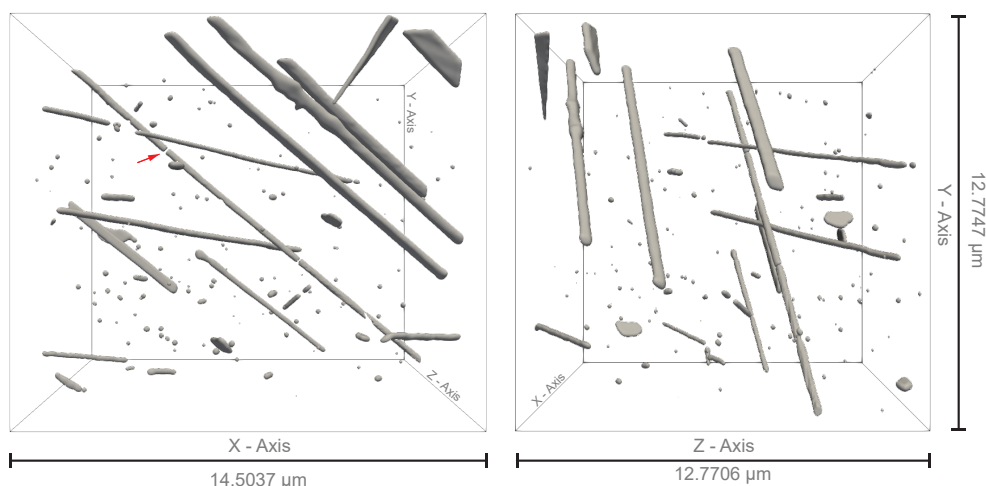
## 2.5 Micromagnetic modelling

Micromagnetic modeling uses the Micromagnetic Earth Related Rapid Interpreted Language Laboratory (MERRILL), a micromagnetics package optimized for rock magnetism [Ó Conbhuí *et al.*, 2018; Fabian and Shcherbakov, 2018]. MERRILL applies finite element tetrahedral meshes to calculate local micromagnetic energies, and a boundary element method to calculate the demagnetizing energy and its gradient [Ó Conbhuí *et al.*, 2018]. For each particle in our study, first an initial remanent domain state was obtained by minimizing the total micromagnetic energy starting from a state of fully randomised spins. Next, for each particle the upper branches of 20 hysteresis loops between the maximal fields of 350 mT and -350 mT was calculated in 5 mT steps. Each of these hysteresis loops was calculated for one of 20 directions chosen from a Fibonacci sphere to achieve an approximately homogenous axis distribution. Backfield remanence curves from 0 to -200 mT in 5 mT steps were calculated for the same 20 field orientations starting from a positive saturation remanence state. While it is possible to run multiple MERRILL operations simultaneously, with the computer power available (Linux server with 20 cores and 128 GB memory) it was still necessary to limit the input file size because of the exponentially increasing time used for minimization. For file sizes < 20 MB an average hysteresis run took 15 min, and an average backfield curve required 35 min. It was possible to complete one hysteresis loop for 250 particles in two field directions within 24 hours. The total dataset presented in this paper equals to approximately 8000 hysteresis loops and 8000 backfield curves. The numerical average of the hysteresis loops and backfield curves over the 20 field directions for each mesh geometry approximately represents a random spherical ensemble of equal particles. Values of  $M_{rs}$ ,  $H_c$  and  $H_{cr}$  were extracted from the average curves. Visual presentation of the magnetization structures of individual models was obtained using the open-source software ParaView ([www.paraview.org](http://www.paraview.org)) [Ayachit, 2015].

## 3 Results

### 3.1 Plagioclase

One cube measuring  $12.77 \mu\text{m} \times 14.50 \mu\text{m} \times 12.77 \mu\text{m}$  (L×W×H) was extracted from plagioclase (PLAG (1) Figure 1 and Figure 2). This volume was reconstructed from 421 slices, each with a thickness of 30 nm. We extracted a total of 162 particles of magnetite and were able to run hysteresis loops for 140 of these. In this volume of plagioclase the magnetite forms two different patterns differentiated by shape and size. The first and most noticeable are magnetite "rods" that are a few to several micrometers long and often extend beyond the milled volume. In Figure 2 we observe these rods oriented in two sets of parallel directions. This observation suggests that these rods are crystallographically oriented, however, we would need to cross reference EBSD and magnetite long axis pole plots to confirm this. Cross sections show that these rods/needles either form flattened or prismatic shapes, with irregular gaps (red arrow in Figure 2) making most of these rods sequences of shorter pieces. The size and volume of these rods are too large to be run by MERRILL in a reasonable time (one smaller  $6.6 \mu\text{m} \times 0.2 \mu\text{m}$  magnetite rod was successfully minimized in zero-field from an initial state of randomised moments and shown in Appendix A.1). Figure A.1 show the general observation for magnetite rods that the magnetization aligns with the length of the rod and generates vortices at each end. The 140 smaller magnetite particles are a mix of



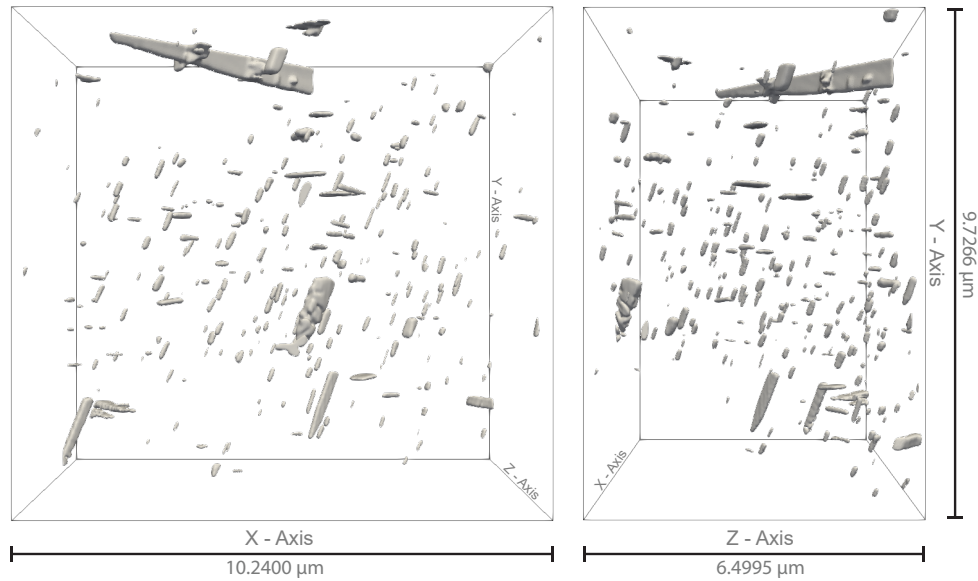
307 **Figure 2.** Magnetite inclusions in a  $12.77 \mu\text{m} \times 14.50 \mu\text{m} \times 12.77 \mu\text{m}$  (L×W×H) volume of plagioclase,  
 308 observed as rods and smaller rather oblate particles. Red arrow highlights the irregular gaps in the magnetite  
 309 rods.

303 prolate, oblate and spherical particles with 0.71 median aspect ratio and  $0.131 \mu\text{m}$  median  
 304 equivalent sphere diameter (Figure 4). These are commonly associated with an unknown  
 305 mineral with a lower backscatter intensity that was removed during electron reflectivity seg-  
 306 mentation, as described in section 2.3 (Appendix A.2).

### 310 3.2 Pyroxene

311 The second volume is from pyroxene (PYR (2) in Figure 1) and measures  $6.5 \mu\text{m} \times$   
 312  $10.24 \mu\text{m} \times 9.73 \mu\text{m}$  (L×W×H) (Figure 3). This volume is reconstructed from 647 slices  
 313 with a slice thickness of 10 nm, which contained total of 282 magnetite particles of which  
 314 hysteresis loops were calculated for 273. Magnetite in pyroxene forms elongated particles  
 315 with a median 0.31 aspect ratio (Figure 4b) that are observed in three orientations. This  
 316 silicate-oxide relationship is described by [Fleet *et al.*, 1980; Frandsen *et al.*, 2004; Renne  
 317 *et al.*, 2002]. For the small magnetite particles included in this pyroxene dataset the Ti con-  
 318 tent is expected to be low. The largest object does show sections that contain higher Ti values  
 319 and is excluded from this dataset. EDS overview maps for areas with small particles also  
 320 contain trace amount ( $< 1 \text{ wt}\%$ ) Ti, but there is no visible change in the backscatter inten-  
 321 sity, nor any Ti peak in EDS spot analysis and we can therefore not confirm where the Ti is  
 322 located. Due to the size of the smaller objects and the spot size of the SEM there is a possi-  
 323 bility that the magnetite contain sections that are higher in Ti (ulvöspinel or ilmenite). TEM  
 324 or Microprobe is needed to quantify the Ti content and confirm the mineral identification.  
 325 Based on the cited papers above there is good confidence to identify these smaller inclusions  
 326 as magnetite.

329 Figure 4 and Appendix A.5 summarizes sizes (A) and aspect ratios (B) for the mod-  
 330 elled particles in both plagioclase and pyroxene. The "equivalent sphere diameter" is calcu-  
 331 lated from a sphere with equivalent particle volume. The aspect ratio was calculated using  
 332 the ratio of min over max eigenvalues of the moment of inertia. The complete dataset pre-  
 333 sented in Figure 4 is subdivided according to both the type of silicate host and the zero-field-  
 334 minimized magnetic state (Figure 5 and Table 2). For SD particles we observe a median size  
 335 difference between particles hosted by plagioclase ( $0.062 \mu\text{m}$ ) versus pyroxene ( $0.107 \mu\text{m}$ ).  
 336 The plagioclase-hosted magnetite particles have a difference in shape, with median aspect  
 337 ratios of 0.63 compared to 0.26 for pyroxene-hosted magnetite. The difference in shape is



327 **Figure 3.** Magnetite inclusions in a  $6.5 \mu\text{m} \times 10.24 \mu\text{m} \times 9.73 \mu\text{m}$  (L×W×H) volume of pyroxene, ob-  
 328 served as mainly prolate shapes in three orientations.

338 also observed for SV particles, with median size and aspect ratio of  $0.144 \mu\text{m}$  and  $0.72$ ,  
 339 respectively, for plagioclase and  $0.159 \mu\text{m}$  and  $0.4$ , respectively, for pyroxene. The shape  
 340 difference is not observed for MV particles, however, with median size and aspect ratio of  
 341  $0.247 \mu\text{m}$  and  $0.24$ , respectively, for plagioclase and  $0.214 \mu\text{m}$  and  $0.26$ , respectively, for  
 342 pyroxene. There are three particles characterized as having SW ("swirl") domain structure  
 343 (Appendix A.3). This group refers to a distinct set of large grains with low aspect ratios (me-  
 344 dian size of  $0.370 \mu\text{m}$ , and aspect ratio of  $0.2$  that represents a particle  $0.500 \mu\text{m}$  long and  
 345  $0.100 \mu\text{m}$  wide). These particles were observed only within pyroxene and, because of their  
 346 large size we were only able to complete 20 hysteresis runs for these three particles. We were  
 347 unable to complete backfield runs due to insufficient computer memory.

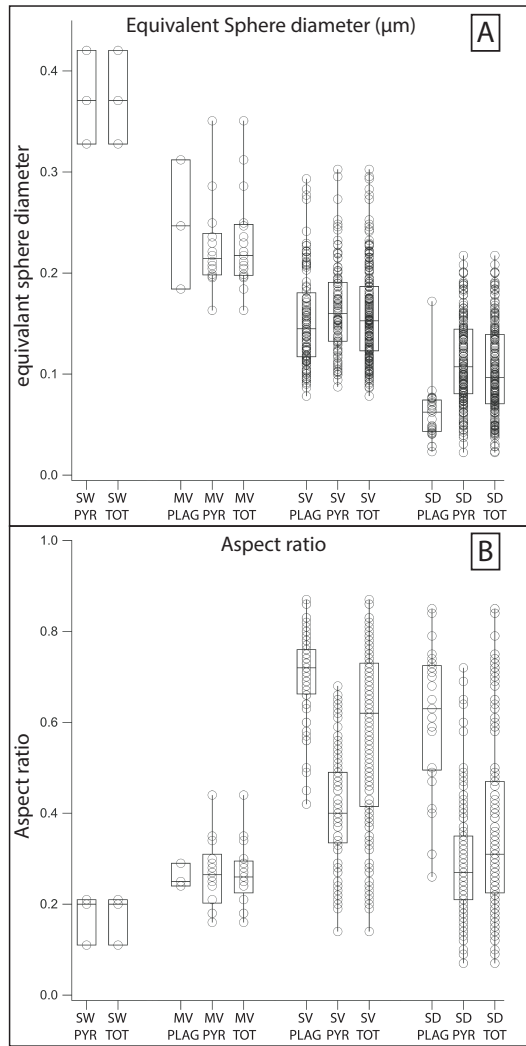
### 351 3.3 Domain characterization

Figure 5 shows the micromagnetic states of representative particles and how they were categorized after minimizing their total magnetic energy from random initial moments. The domain state estimations represent one of many possible local energy minima. The minimization from random initial moments should ideally be repeated multiple times to categorize the full range domain states, and potentially identify the global energy minimum state. This is computationally expensive, but should be encouraged for future studies and as a part of the FIB-nt method in general. Insets show the corresponding average hysteresis loop for each particle. The arrows are color coded by their vertical ( $M_z$ ) component where up = red and blue = down. The vortex core is represented by contour surfaces (in green) enclosing volumes with increased helicity of the magnetization field  $\mathbf{m}$  defined by

$$\text{hel } \mathbf{m} = \mathbf{m} \cdot (\nabla \times \mathbf{m}).$$

352

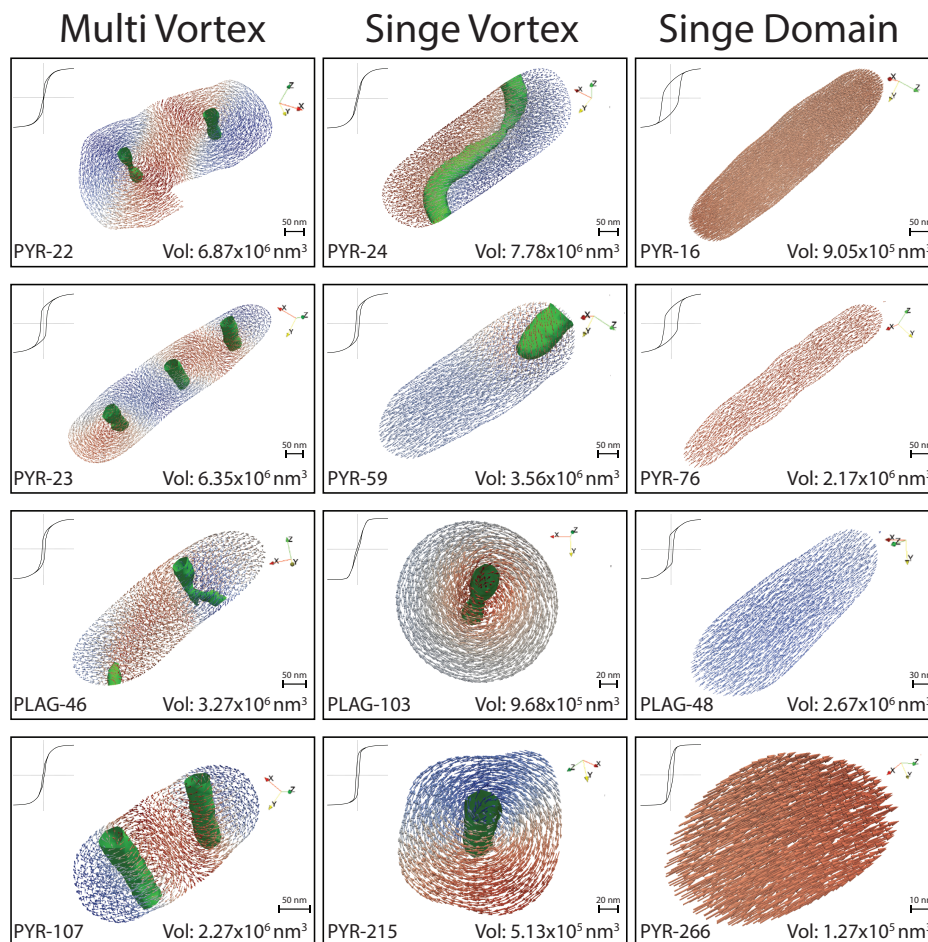
362 Single domain (SD) particles have uniform magnetic moments aligned throughout the  
 363 particle, although the spins may bend along the demagnetizing field lines at edges and  
 364 corners (flower state). The single vortex (SV) is an inhomogeneous magnetization state with low-  
 365 est total micromagnetic energy beyond the single domain limit [Hubert and Schäfer, 1998].



348 **Figure 4.** Box plots that show the relation of particle size (A) and aspect ratio (B), subdivided according to  
 349 the host silicate (plagioclase and pyroxene) and the modelled micromagnetic domain structure (single domain,  
 350 single vortex, multi vortex, swirl).

366 The net energy decrease is obtained by reducing the magnetic stray-field energy through flux  
 367 closure at the expense of increasing exchange and anisotropy energy. Therefore this transi-  
 368 tion can only occur when the decrease in demagnetizing energy overcompensates this in-  
 369 crease. Visually, the single vortex core extends partially or fully through the particle. *Rave*  
 370 *et al.* [1998a]; *Hubert and Schäfer* [1998] find that there is a transition from a simple vortex-  
 371 state to a well-defined multi-domain state, where vortices are embedded features within the  
 372 soft-magnetic Bloch walls. In this paper this transition is described as a multi vortex (MV)  
 373 state from the zero-field minimization. The MV state is characterized by the presence of two  
 374 or more well-defined vortex cores that create a continuously varying directional change in  
 375 magnetization, as opposed to an abrupt angular change that would indicate a domain wall.  
 376 Similar observations have been made in much larger ( $1.4 \mu\text{m}$ ) pyramidal titanomagnetite  
 377 particles [*Khakhalova et al.*, 2018]. Particles with fully developed  $90^\circ$  domain walls were  
 378 in general too large for our server to minimize hysteresis and backfield curves in 20 field  
 379 directions in MERRILL. *Nagy et al.* [2019] did model cuboctahedral magnetite from 0.03-  
 380  $2.7 \mu\text{m}$  equivalent sphere volume diameter and  $0.03\text{-}1.5 \mu\text{m}$  magnetite spheres, studying the





353 **Figure 5.** Magnetic models that express the variability of magnetic signatures within each classification.  
 354 Insets show the simulated average hysteresis loop for each particle, with normalized moments and external  
 355 field from +350 to -350 mT. Arrow colors show the  $M_z$  component, red = up and blue = down. Green contour  
 356 (vortex core) represents areas of increased helicity of the magnetization field. Hysteresis, size and shape pa-  
 357 rameters for each individual particle are summarized in Table 1. Location of these twelve particles are marked  
 358 with bold text in the subsequent figures 7, 8, 9, 10 and 11.

381 evolution of domain structures in magnetite grains with different sizes. They observed the  
 382 formation of primitive bloch-like 71° domain walls that separate uniform regions of magne-  
 383 tization, similar to our observations in the three selected large particles in Appendix A.3. In  
 384 these particles the magnetization pattern mainly consists of vortices or swirls that represent  
 385 seeds of Bloch walls, but due to size constraints, cannot develop into conventional domain  
 386 walls as they occur in well-developed multi-domain patterns [Fabian *et al.*, 1996; Hubert  
 387 and Schäfer, 1998]. Domain walls widths expected for MD magnetite are observed only in  
 388 much larger particles ( $\sim 3 \mu\text{m}$ ) and the SV structures are present until at least  $\sim 1 \mu\text{m}$  [Nagy  
 389 *et al.*, 2019]. Due to the lack of a distinctive terminology, the intermediate sized magnetite  
 390 structures are defined as "swirl" (SW) within this paper.

359 **Table 1.** Hysteresis( $f=f_{emto}$ ), diameter (equivalent sphere diameter) and aspect ratio parameters of the  
 360 average hysteresis loops for the individual particles inset in Figure 5 and highlighted with bold text in Fig-  
 361 ure 7, 8, 9, 10 and 11.

Particle	$M_s$ (fAm <sup>2</sup> )	$M_{rs}$ (fAm <sup>2</sup> )	$H_c$ (mT)	$H_{cr}$ (mT)	Diameter ( $\mu$ m)	Aspect ratio
PYR-16	0.43	0.223	92.7	115.3	0.120	0.19
PYR-22	3.30	0.281	3.35	20.6	0.236	0.21
PYR-23	3.05	0.157	32.4	41.8	0.230	0.16
PYR-24	3.74	0.326	8.8	44.4	0.246	0.34
PYR-59	1.71	0.789	18.9	32.7	0.189	0.28
PYR-76	1.04	0.519	76.7	94.3	0.161	0.13
PYR-107	1.09	0.114	9.53	26.6	0.163	0.44
PYR-215	0.246	0.120	15.3	29.0	0.099	0.47
PYR-266	0.061	0.028	26.7	42.8	0.062	0.69
PLAG-46	1.57	0.628	16.9	29.8	0.184	0.29
PLAG-48	1.28	0.644	23.7	34.6	0.172	0.31
PLAG-103	0.465	0.404	9.8	48.2	0.123	0.73

400 **Table 2.** Particles categorization of domain state that are hosted by their respective silicate.

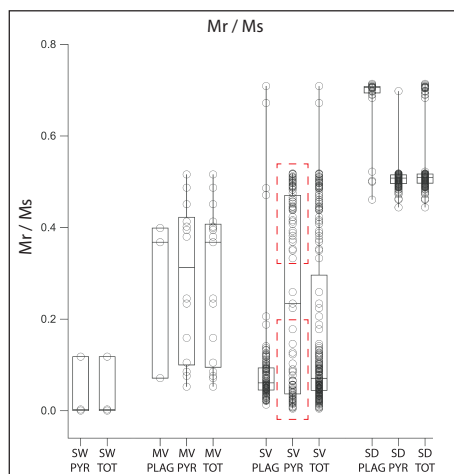
Silicate	SD (green)	SV (red)	MV (blue)	SW (purple)	total
Plagioclase	29	108	3	0	140
Pyroxene	159	97	14	3	273

### 391 3.4 Hysteresis loops

392 Hysteresis results for all simulated inclusions, from both plagioclase (circles) and py-  
 393 roxenes (triangles), are summarised in Figure 6, Figure 7 (Day plot), Figure 9 (Néel plot),  
 394 Figure 11 (Butler-Banerjee plot) and Appendix A.5. Each point represents the average hys-  
 395 teresis data acquired from 20 different field directions and is colour coded according to the  
 396 remanent domain state observed at zero field following energy minimisation from a random  
 397 starting configuration (purple = SD, yellow = SV and turquoise = MV). Each symbol size is  
 398 scaled according to particle size, defined as the diameter of a sphere with equivalent volume.  
 399 The total number of particles adopting each remanent state is represented in Table 2.

401 The two red markers (one for each silicate) in Figures 7 and 9 represent the average  
 402 hysteresis properties for all of the modelled particles. Assuming the two volumes studied  
 403 provide a representative sample of the silicate-hosted component of magnetic mineralogy,  
 404 these two points provide an estimate of the hysteresis properties that would be obtained if  
 405 we measured the bulk properties of plagioclase and pyroxene extracts experimentally. These  
 406 'bulk' hysteresis properties would normally be the only two points available to perform do-  
 407 main state diagnosis. However, when deconstructed to individual particles, we observe the  
 408 data for both volumes extends across the entire range of the common diagnostic diagrams  
 409 and all regions designated to specific domain states. The location of these bulk average points  
 410 also suggest the domain states are unevenly distributed between the silicates. The bulk av-  
 411 erage for pyroxene is dominated by SD particles and for plagioclase is dominated by SV  
 412 particles (Table 2). Figure 6 shows the distribution of particles between the domain states,  
 413 host silicate and  $M_{rs}/M_s$  ranges. For SD there is a distinct separation, with tight clusters of  
 414  $M_{rs}/M_s$  ranges with pyroxene hosted particles at 0.5 and plagioclase hosted particles at 0.7.  
 415 For SV particles there are  $M_{rs}/M_s$  values where particles form clusters. Within plagioclase  
 416 SV particles cluster at  $M_{rs}/M_s$  values  $\approx 6.1 \times 10^{-2}$ , but there are a few points that extend the  
 417 range up to  $M_{rs}/M_s = 0.7$ . SV particles in pyroxene show a bimodal distribution (red boxes)

418 with clusters at  $M_{r,s}/M_s$  values  $> 0.3$  and  $< 0.3$ . MV particles show a similar tendency to  
 419 form two separate clusters at  $M_{r,s}/M_s$  values  $\approx 0.1$  and  $0.4$ . This is most evident for particles  
 420 hosted by pyroxene, but with only three particles extracted from the plagioclase is an insuffi-  
 421 cient sample size to indicate anything about specific  $M_{r,s}/M_s$  ranges. What we have for now  
 422 indicates that these particles cover the same  $M_{r,s}/M_s$  ranges as MV magnetite in pyroxene.  
 423 Three SW particles hosted by pyroxene did successfully complete minimization of the hys-  
 424 teresis loop. The spread in  $M_{r,s}/M_s$  is from  $0.12$  to  $3.44 \times 10^{-4}$ .



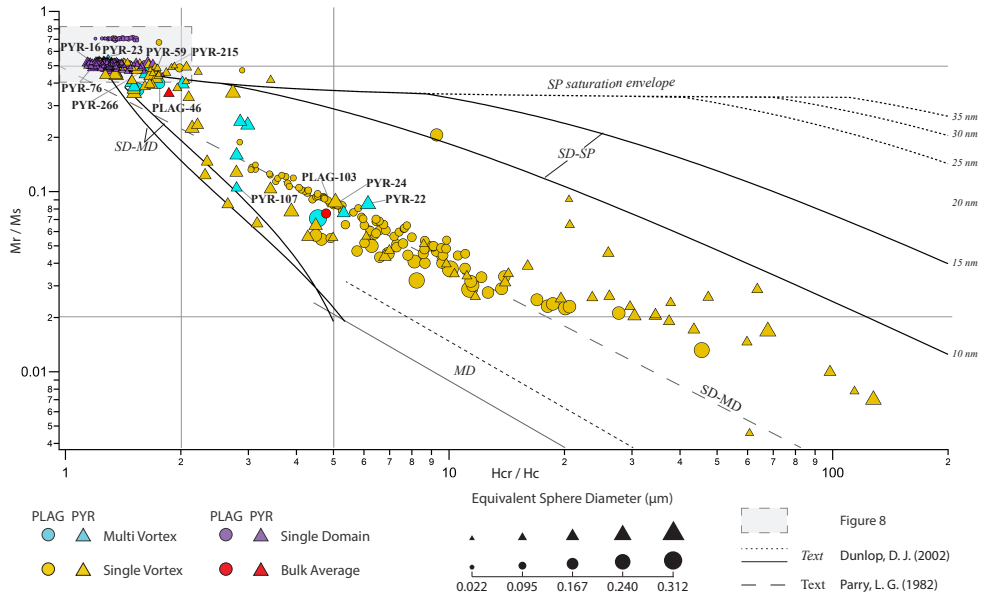
425 **Figure 6.** Box plots for the statistics of individual magnetite  $M_{r,s}/M_s$  in plagioclase, pyroxene and com-  
 426 bined total for the different modelled domain structures.

### 427 3.5 Day plot

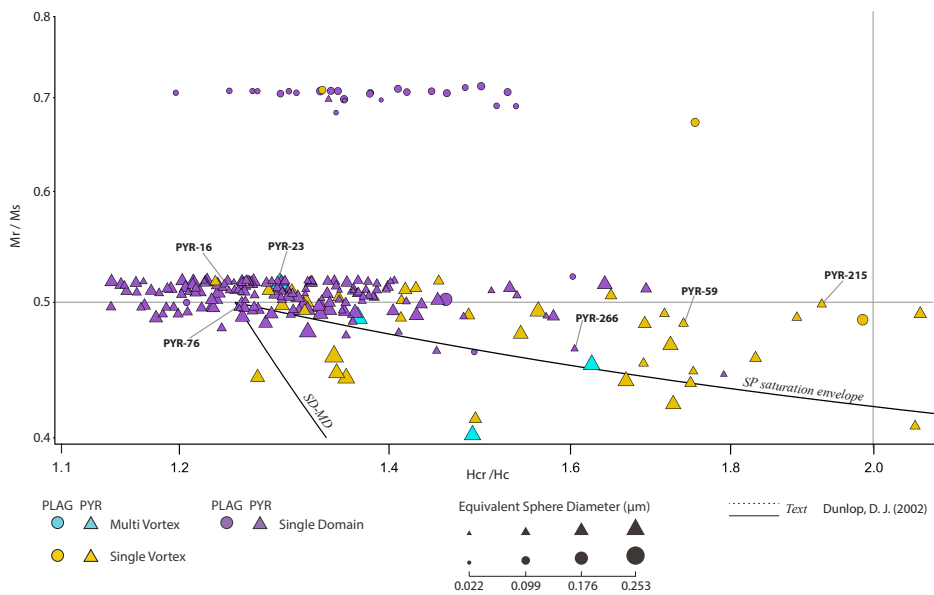
428 In Figure 7 the SD particles plot in the expected region at  $M_{r,s}/M_s \approx 0.5$ . SV particles  
 429 span the whole range from SD to MD. MV particles tend to plot towards the upper left with  
 430  $M_{r,s}/M_s$  values  $> 6.0 \times 10^{-2}$  and  $H_{cr}/H_c < 7$ . The spread of points shows little regard for the  
 431 particle size, highlighted by Figure 8, where particles with varying sizes and domain states  
 432 coexist and even overlap. This observation is also evident throughout Figure 7. At  $M_{r,s}/M_s$   
 433 values  $\approx 0.5$  the cluster of prolate particles is mainly hosted by pyroxene. SD particles hosted  
 434 by plagioclase form a cluster at  $M_{r,s}/M_s$  values  $\approx 0.7$ .

### 439 3.6 Néel plot

440 In a Néel plot (Figure 9) we observe a dominating SD "central ridge" that extends from  
 441  $10 \text{ mT} < H_c < 120 \text{ mT}$ . The SD particles mainly hosted by pyroxene at  $M_{r,s}/M_s \approx 0.5$  align  
 442 with the uniaxial SD (USD) theoretical line. Axis ratios show the majority of these parti-  
 443 cles are prolate in shape (Appendix A.4), and correspond to the median aspect ratio cluster  
 444 of  $0.26$  (Figure 4b). The feldspar-hosted SD particles (higher median aspect ratio of  $0.62$  in  
 445 Figure 4b) form a cluster at  $M_{r,s}/M_s$  values  $\approx 0.7$  and  $10 \text{ mT} < H_c < 30 \text{ mT}$ . These particles  
 446 are oblate spheroids and the increase in  $H_c$  for these particles represents a general increase in  
 447 particle oblateness/flatness (Figure 9A). The low-coercivity ( $< 10 \text{ mT}$ ) region is dominated  
 448 by SV and MV particles. With decreasing coercivities below  $10 \text{ mT}$ , there is also a decrease  
 449 in  $M_{r,s}/M_s$ . Particles with  $M_{r,s}/M_s$  values  $< 0.3$  form a distinct cluster that is dominated by  
 450 SV states (Figure 10). This cluster has a well-defined trend for particles from both plagi-  
 451 olase and pyroxene. While pyroxene-hosted particles are equally sized and evenly distributed  
 452 in Figure 10, there is a decrease in particle size in plagioclase above  $H_c \approx 9 \text{ mT}$ . Particles  
 453 below  $9 \text{ mT}$  and  $M_{r,s}/M_s = 0.07$  align with the theoretical "USD + SP". Those above  $9 \text{ mT}$

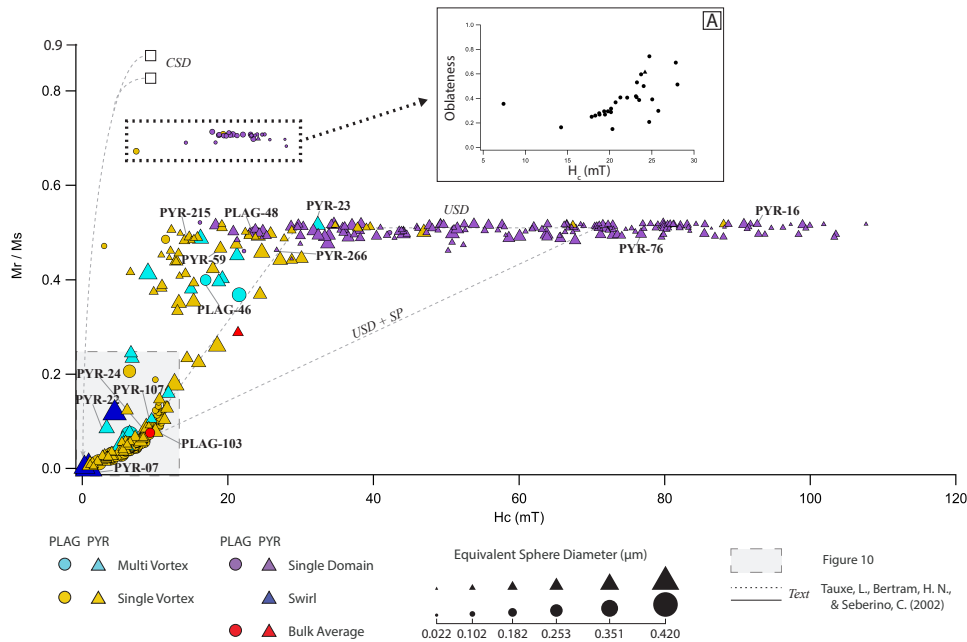


435 **Figure 7.** Day plot with theoretical mixing curves of *Dunlop* [2002] and *Parry* [1982]. Regions for SD,  
 436 PSD and MD from *Dunlop* [2002]. Bold text indicates the positions of particles from Figure 5 and Table 1.



437 **Figure 8.** Expanded area of Figure 7 showing the SD clusters at  $M_{rs}/M_s > 0.43$  and  $H_{cr}/H_c < 2.4$ . Bold  
 438 text indicates the positions of particles from Figure 5 Table 1.

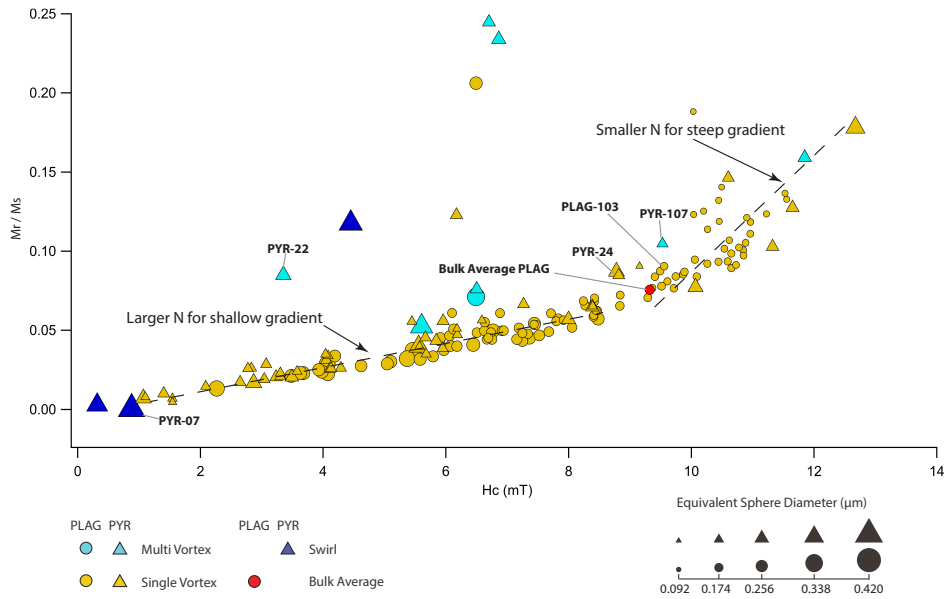
454 demonstrate an increased upwards slope that follows the USD theoretical lines from *Tauxe*  
 455 *et al.* [2002].



456 **Figure 9.** Néel plot with the location of particles from Figure 5 in bold and Table 1. Theoretical lines of  
 457 USD, USD+SP and CSD from *Tauxe et al.* [2002]. Inset (A) shows the increasing oblateness for the SD  
 458 particles at  $M_{r_s}/M_s = 0.7$  with increasing  $H_c$ .

### 461 3.7 Butler-Banerjee plot

462 The Butler-Banerjee plot in Figure 11 delineates a two-dimensional size-shape range  
 463 for SD particles of magnetite and titanomagnetite [*Butler and Banerjee, 1975*]. Contrary to  
 464 the Day and Néel plots, which interpret domain state from hysteresis parameters, the Butler-  
 465 Banerjee plot predicts domain states based on the length of a particle's long axis compared  
 466 to its axis ratio. One is therefore able to characterize a particle's domain state purely based  
 467 on shape and size. The theoretical boundaries in the background were calculated either by a  
 468 domain model [*Butler and Banerjee, 1975*] (dashed line) or by 3D micromagnetic modeling  
 469 [*Muxworthy and Williams, 2008*] (solid lines). In *Butler and Banerjee [1975]* the particles  
 470 have cuboidal shapes with  $180^\circ$  domain walls, while in *Muxworthy and Williams [2008]* the  
 471 magnetic particles are SD or SV parallelepipeds arranged in chains to assess magnetostatic  
 472 stabilization in chains of magnetosomes from magnetotactic bacteria. Our dataset defines  
 473 a clear transition between SD and SV states that is in good agreement with the 'no interac-  
 474 tion' boundary identified by *Muxworthy and Williams [2008]* for isolated parallelepipeds.  
 475 Our data in Figure 11 clearly reflect the influence of the silicate host mineral on the inclu-  
 476 sion shape and subsequently its magnetic properties. While plagioclase-hosted magnetites  
 477 (circles) plot predominantly at aspect ratios (AR) above 0.6, pyroxene-hosted magnetites (tri-  
 478 angles) have AR less than 0.6. Particles that are able to adopt MV remanence states typically  
 479 plot in the region with  $AR < 0.4$  and long axis  $> 0.300 \mu\text{m}$ . The phase boundary between  
 480 SD and SV/MV states in this region deviates from that of *Muxworthy and Williams [2008]*,  
 481 most likely reflecting the difference between the real particle geometries used here versus the  
 482 ideal parallelepipeds used by *Muxworthy and Williams [2008]*. Figure 11 demonstrates that



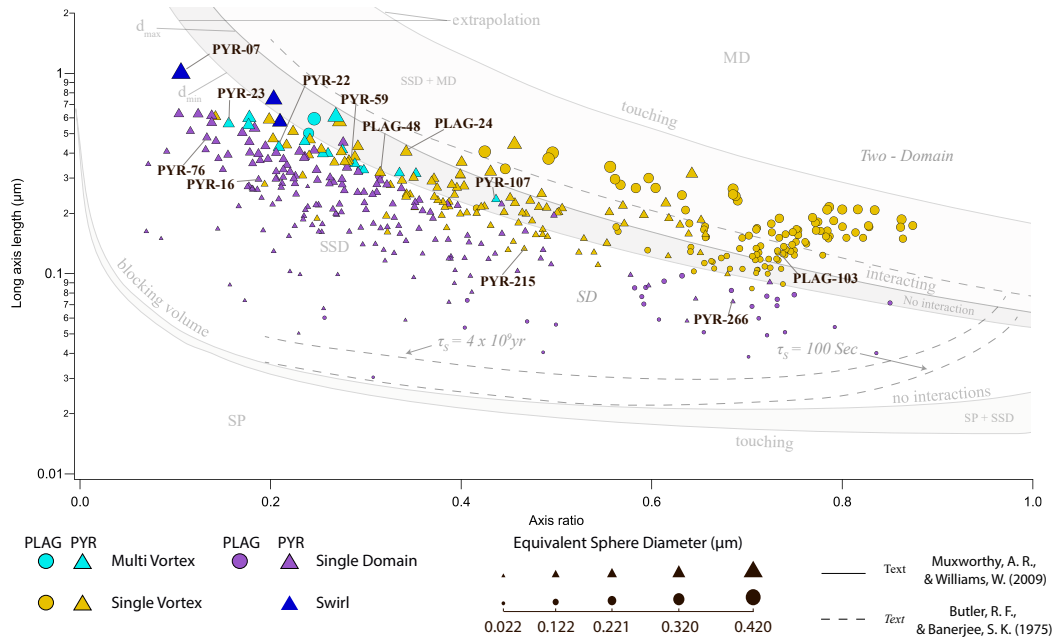
459 **Figure 10.** Expanded area of  $M_{r,s}/M_s < 0.25$  and  $H_c < 14\text{mT}$  from Figure 9. Bold text indicates particles  
 460 from Figure 5, Table 1 and the bulk average for plagioclase. Particle PYR-07 is displayed in Appendix A.3.

483 MV and SW particles start to dominate over SV states with decreasing AR and increasing  
 484 long axis length.

#### 489 4 Discussion

490 As micromagnetics takes us beyond the SD theories that have been the foundation of  
 491 rock magnetism since the pioneering work of Néel, the ability to obtain an accurate break-  
 492 down of the different classes of magnetic behaviour that exist within a sample becomes in-  
 493 creasingly important. One of the most striking results of this study is the shear range of mag-  
 494 netic domain states and hysteresis properties that are observed within just two small  $\sim 10\mu\text{m}^3$   
 495 volumes of silicate. The observed behaviour spans the entire range of popular domain-state  
 496 diagnostic plots (Figures 7, 9 and 11). In contrast, bulk hysteresis measurements reduce this  
 497 complexity to a single average data point (e.g. the red points in Figures 7, 9), discarding  
 498 valuable information about the true nature of the underlying magnetic ensemble. Classify-  
 499 ing a bulk sample as either 'SD', 'PSD', or 'MD' on the basis of the Day diagram is clearly  
 500 meaningless when particles in the SV state plot across the entire 'SD/PSD/MD' range, and  
 501 the largest MV particles plot much closer to the SD region than the MD region (Figure 7).  
 502 Of the three diagnostic plots explored here, the most successful is the Butler-Banerjee plot  
 503 (Figure 11). This diagram works well because it recognises the fundamental importance of  
 504 particle size and shape as the dominant factors controlling the domain state of magnetite.  
 505 However, our results also demonstrate that size and shape alone are often poor predictors  
 506 of magnetic hysteresis properties. Generalised geometric information, such as equivalent  
 507 spherical volume, is not sufficient to predict the magnetic properties – the specifics of indi-  
 508 vidual particle geometries really do matter. Although time consuming and computationally  
 509 intensive, the FIB-nt-FEM workflow presented here provides a route to rock magnetic char-  
 510 acterisation that not only yields a quantitative breakdown of the domain states present, but  
 511 a realistic estimate of the range of magnetic properties associated with the particle ensemble.  
 512 As the database of magnetic properties linked to specific particle geometries expands,  
 513 it will eventually become possible to use look-up tables and/or machine-learning approaches  
 514 to translate 3D particle geometries into useful magnetic parameters (including, for example,





485 **Figure 11.** Butler-Banerjee diagram that shows the relationship between a particle's long axis and its axis ratio.  
 486 axis ratio = 0 references an infinitely long needle and axis ratio = 1 references an equidimensional particl  
 487 (e.g. cube or sphere). Bold text indicates the position of particles visible in Figure 5 and Table 1, except  
 488 for particle PYR-07 which is displayed in Appendix A.3.

515 energy barriers for thermal switching) without the need to perform additional micromagnetic  
 516 simulations. With increased automation, the experimental methods outlined here could in-  
 517 deed become a routine feature of future rock magnetic characterisation workflows.

518 The Néel plot (Figure 9) shows a distinct "central ridge" of SD particles along the USD  
 519 line at  $M_{r_s}/M_s \approx 0.5$ . *Tauxe et al.* [2002] argue that the  $H_c$  increase for SD particles is pro-  
 520 portional to an increase in the length to width (L/W) ratio. Figure 4a in *Tauxe et al.* [2002]  
 521 shows that the increase from  $H_c = 30$  mT to 70 mT for USD particles is caused by an in-  
 522 crease in L/W from 1.3:1 to 2:1 While we do observe that particles with  $H_c > 100$  mT have  
 523 high L/W ratio ( $> 10 : 1$ ), for coercivities  $< 100$  mT we do not find a clear correlation be-  
 524 tween coercivity and aspect ratio, which may be due to the large spread of USD particle sizes  
 525 (0.022-0.253  $\mu\text{m}$ ). Our data imply that the size of a USD particle is at least as important as  
 526 its aspect ratio in determining the switching field, highlighting again the need for detailed  
 527 geometrical information to produce an accurate prediction of the magnetic properties of an  
 528 ensemble. Another interesting modeling result is that the characteristic coercivity range for  
 529 SD magnetite particles predominantly covers the range  $10 \text{ mT} < H_c < 100 \text{ mT}$ , which is in  
 530 line with experimental results [*Church et al.*, 2016]. Without external stress, higher  $H_c$   
 531 values appear to require very unusual particle geometry, while SD structures with  $H_c < 10$  mT  
 532 can occur only in a very narrow grain size and shape region.

533 A cluster of feldspar-hosted SD particles with  $M_{r_s}/M_s$  values of 0.7 (Figure 9 and 8)  
 534 falls in between the theoretical values for cubic SD ( $M_{r_s}/M_s = 0.83-0.87$ ) [*Joffe and Heuberger,*  
 535 1974; *Tauxe et al.*, 2002] and uniaxial SD ( $M_{r_s}/M_s = 0.5$ ). This remanence ratio is similar to  
 536 the value of  $M_{r_s}/M_s = 0.75$  calculated by *Harrison et al.* [2019] for particles with a combi-  
 537 nation of uniaxial anisotropy (restricting moments to a basal plane) and hexagonal anisotropy  
 538 (defining multiple easy axes within that basal plane). This cluster contains mainly oblate par-  
 539 ticles (Appendix A.4), and therefore the intermediate value of  $M_{r_s}/M_s$  is entirely consistent

544 **Table 3.** Individual particles with different magnetization structure that are all within the designated SD  
 545 area of the Day plot

Object		$M_{rs}/M_s$	$H_c$ (mT)	$H_{cr}$ (mT)	Aspect Ratio	Sphere Diameter ( $\mu\text{m}$ )
PYR-137	SD	0.516	34.2178	46.4489	0.31	0.142
PLAG-146	SD	0.500	42.2442	50.9584	0.41	0.062
PYR-197	SV	0.519	19.1958	27.875	0.49	0.111
PLAG-135	SV	0.709	19.3504	25.7931	0.71	0.078
PYR-23	MV	0.516	32.3692	41.8226	0.16	0.229
PYR-21	MV	0.490	16.2453	22.2288	0.18	0.249

540 with these particles having a mixed uniaxial/multi-axial anisotropy. A positive correlation  
 541 between coercivity and oblateness is observed (inset to Figure 9). This improved correla-  
 542 tion may be a result of the smaller size variations within this oblate cluster (0.023-0.090  $\mu\text{m}$ )  
 543 compared to that of the pyroxene-hosted USD prolate cluster.

546 The SV domain state is observed across a wide range of particle sizes and shapes, and  
 547 produces correspondingly large ranges of  $H_c$  (0-88 mT),  $M_{rs}/M_s$  (0-0.7) (Figure 9) and  
 548  $H_{cr}/H_c$  (1.5-107) (Figure 7). The SV particles in Figure 6 belong to different shape clus-  
 549 ters, where PLAG contains more oblate and PYR more prolate particles with respect to the  
 550 shape classification in Appendix A.4. For a constant demagnetizing factor (N), one would  
 551 expect a linear increase of  $M_{rs}/M_s$  with  $H_c$  with slope inversely proportional to N (Figure  
 552 4.5 in *Stacey and Banerjee* [1974] and Figure 5.16 in *Dunlop and Özdemir* [1997]). The in-  
 553 creased slope of the  $M_{rs}/M_s$  vs  $H_c$  plot for  $H_c > 9$  mT (Figure 10) indicates, therefore, a  
 554 smaller average demagnetizing factor for the smaller particles compared to the larger particles  
 555 with  $H_c < 9$  mT (Appendix A.6).

556 Another key result of this study is the observation of SV and MV particles that display  
 557 hysteresis properties similar to those expected for much smaller SD particles (Table 3). The  
 558 MV state is stable in large grains (median 0.225  $\mu\text{m}$ ) with low aspect ratios  $< 0.5$  (Figure 4a  
 559 and b). Despite their larger sizes, these particles plot more towards the upper left of a Day  
 560 diagram (SD and PSD regions), rather than towards the MD region. Large, slightly flattened  
 561 needle ( $a \gg b > c$ ) particles with low aspect ratio readily adopt MV states during zero-field  
 562 relaxation from a random starting configuration (e.g. PYR-23 in Figure 5), but are forced  
 563 into a uniformly magnetised remanence state after exposure to a saturating field. This pro-  
 564 duces an apparent disconnect between the domain-state classification used in this study and  
 565 the observed hysteresis properties. This disconnect becomes significant if, as intended, the  
 566 zero-field relaxation state is more representative of a weak-field thermoremanent magnetisa-  
 567 tion (TRM) state than the corresponding saturation remanence state. Such disconnects have  
 568 been observed in dusty olivine by *Lappe et al.* [2013], where stable SV states were observed  
 569 after acquisition of weak-field TRM but were converted to metastable SD states after appli-  
 570 cation of a room-temperature saturating field. Our study raises the question of whether high-  
 571 field hysteresis measurements are always a useful means to classify the potential weak-field  
 572 remanence states adopted by a given particle geometry, especially for SV and MV states,  
 573 which may well be the dominant carriers of remanent moment. Such ambiguities are avoided  
 574 using the FIB-nt-FEM approach, ensuring that an appropriate combination of physical mod-  
 575 els is always used for a given ensemble. A full theory of the remanence state adopted by such  
 576 particles requires us to calculate the multitude of energy barriers between all possible reman-  
 577 ent states. Although this is an active area of research, it is outside the scope of the current  
 578 study. As discussed above, the state adopted will be highly dependent on the external field  
 579 history. By introducing the energy barrier calculations, we would expect to see a better sep-  
 580 aration line between the remanence states of all the particles, resulting in adjustments to the  
 581 theoretical lines in Figure 11 from *Butler and Banerjee* [1975]; *Muxworthy and Williams*  
 582 [2008].



583 The blue particles seen in Figure 9, 10 and 11 are large ( $0.370\ \mu\text{m}$  median equivalent  
584 sphere diameter) oblate to prolate shapes (Figure 4a, b and Appendix A.4) which show a ten-  
585 dency to form domain walls when minimized in zero-field (Appendix A.3). The backfield  
586 curves were not able to be calculated on a reasonable timescale, and therefore are not plotted  
587 on a Day diagram (Figure 7). On a Néel diagram (Figure 9), two SW particles plot close to  
588 the origin, as would be expected for MD behaviour. One, however, displays similar proper-  
589 ties to MV particles, suggesting that they share many characteristics, and that the transition  
590 from MV to SW to MD is likely gradual one for particles with this shape.

591 In Figure 7 the theoretical Day plot SD-MD and SD-SP mixing lines show the mod-  
592 ified calculated binary curves *Dunlop* [2002] and the SD-MD mixing curve derived from  
593 natural magnetite mixtures in *Parry* [1982]. Because effects like thermal variations, stress,  
594 interactions, impurities or inclusions are not included in the models of this study, the hystere-  
595 sis properties of our modelled particles represent ideal magnetite crystals. When the entire  
596 data set is compared to theoretical mixing lines, the closest agreement occurs for the SD-MD  
597 mixing line from *Parry* [1982], that represents a grain-size trend of measured mixtures of  
598 SD and MD particles. The samples used by *Parry* [1982] were generated by crushing natu-  
599 ral MD magnetite and sieving for various grain sizes. Because crushing introduces internal  
600 stress anisotropy that may be larger than the magnetocrystalline anisotropy, the correlation  
601 with our dataset appears to be spurious, but based on the fact that our data consistently lie  
602 above the binary mixing lines from *Day et al.* [1977]; *Dunlop* [2002].

603 It is unavoidable that the magnetic models presented here contain several uncertainties  
604 related to reconstruction artifacts, model assumptions, and unknown physical influence fac-  
605 tors. Shape uncertainty is directly connected to the extraction of selected mineralogy from  
606 the binary stack and the accuracy of the resulting finite-element meshes. The alignment of  
607 fiducial marks in the image stack to reconstruct the volume contributes to the shape uncer-  
608 tainty. The carbon rods are deposited from carbon gas reacting with the ion beam. As ob-  
609 served in the reassembled stack, there are irregular shifts in image placement that also im-  
610 pact the reconstructed shape, especially of the smallest particles. Evidence for this is seen in  
611 the meshes as irregular pinching in elongated particles, where images are shifted. Thresh-  
612 olding in FIJI is another uncertainty in reconstructing the true shape of particles. With the  
613 variation in size, the extraction of the smaller particles will become suppressed when se-  
614 lecting the threshold based on the median size. A simple calculation of the volume change  
615 when increasing the particle volume by adding a pixel in each direction leads to an estimate  
616 of up to 10-25% for the larger particles and 50-130% for the smaller particles (based on an  
617 increase in the representing sphere diameter of 7 nm pixel width). The true change in vol-  
618 ume is dependent on more complicated measures such as the particle orientation and the  
619 milling thickness. We were also concerned that variable amount of mesh smoothing would  
620 possibly change the hysteresis parameters and therefore the domain state. To test this we cal-  
621 culated average hysteresis properties for 3 particles of different sizes at 3 different stages of  
622 smoothness. First stage is where the particle shape is strongly voxelated (directly from the bi-  
623 nary image), second is at intermediate smoothing and the last where all sides are completely  
624 rounded (axis ratio is preserved in all stages). From the resulting hysteresis properties we  
625 did not observe any significant changes beyond the deviation expected for models starting  
626 with randomized moments. This is in agreement with a study by *Rave et al.* [1998b] that at  
627 resolutions below the exchange length, particle surface discretization does not influence mi-  
628 cromagnetic model results. Model assumptions that contribute to the overall uncertainty are  
629 the material parameters for magnetite that in the natural minerals can vary for example due  
630 to impurities. Physical unknowns are lattice orientations, related to the direction of the easy  
631 anisotropy axes, or internal stress [*Hodych, 1990; ter Maat et al., 2020*].

632 This study presents a workflow for extracting the 3D geometry of a nanoscale mag-  
633 netic particle ensemble and calculating its micromagnetic properties. To date it is the most  
634 comprehensive micromagnetic data set of natural remanence carriers, and acts as a proof-of-  
635 concept study for obtaining statistical data on the magnetic domain state and hysteresis prop-

636 erities of silicate-hosted magnetic inclusions. This work is of particular relevance to single-  
 637 crystal paleomagnetic studies, which offer the best possibility of extracting reliable primary  
 638 paleomagnetic remanence data from ancient rocks with complex geological histories. The  
 639 key challenge in such studies is targeting primary remanence carriers that are most likely to  
 640 have escaped thermal, chemical or viscous remagnetization. In the case of the Marcus in-  
 641 trusion, stable remanence carriers hosted in primary feldspar and pyroxene grains could be  
 642 targeted directly, avoiding both remanence carriers hosted by secondary amphibole as well  
 643 as interstitial MD magnetite that could carry a viscous overprint. The combination of FIB-  
 644 nt and micromagnetic simulations provides a full statistical breakdown of the domain states  
 645 present in a representative single crystal. Further studies will enable the range of local en-  
 646 ergy minimum states adopted by the SD, SV, MV, SW and MD states, as well as the energy  
 647 barriers that separate them, to be calculated, and the corresponding blocking/unblocking be-  
 648 haviour during natural and laboratory cooling/heating cycles to be modelled. With that we  
 649 would be one step closer to a comprehensive model of remanence acquisition that takes us  
 650 beyond Néel's SD theory. Ultimately this will increase both the reliability of single-crystal  
 651 paleomagnetic studies and our confidence to interpret them, paving the way to unlocking  
 652 hitherto inaccessible parts of the geomagnetic record.

## 653 5 Conclusions

- 654 1. FIB-nt protocols have been developed for 3D imaging of silicate-hosted magnetite  
 655 inclusions down to the stable SD limit of 30 nm.
- 656 2. Silicate-hosted magnetite inclusions in plagioclase and pyroxene span a wide range  
 657 of domain state behaviours, including SD, SV, MV, SW and MD.
- 658 3. In plagioclase, a dominant cluster of oblate SD particles is found with mixed uniaxial/multi-  
 659 axial anisotropy and  $M_{rs}/M_s = 0.7$ . In pyroxene, SD particles are predominantly prolate  
 660 with uniaxial anisotropy and  $M_{rs}/M_s = 0.5$ .
- 661 4. The range of hysteresis properties observed in just two  $\sim 650 - 2300 \mu\text{m}^3$  volumes of  
 662 pyroxene and plagioclase, span almost the entire range of common domain-state diagnostic  
 663 plots. This indicates that bulk average hysteresis parameters cannot predict a unique domain  
 664 state, even for the particle ensemble inside a single exsolved silicate crystal.
- 665 5. Because shape and size equally determine the hysteresis properties of natural mag-  
 666 netite inclusions, the Butler-Banerjee plot performs best in terms of domain-state diagnosis,  
 667 despite failing to account for MV states.
- 668 6. The morphology of the magnetite inclusions highlights the need to extend Néel's  
 669 SD theory in the context of single-crystal paleomagnetism towards SV and MV magnetiza-  
 670 tion states, as these may be dominant carriers of stable paleomagnetic remanence in silicates.

## 671 References

- 672 Ahrens, J., B. Geveci, and C. Law (2005), *Paraview: An end-user tool for large data visual-*  
 673 *ization. The visualization handbook*, 717 p. pp., Kiteware.
- 674 Argyle, K., and D. Dunlop (1984), Theoretical domain structure in multidomain magnetite  
 675 particles, *Geophys. Res. Lett.*, 11.
- 676 Austin, J., D. Hillan, P. Schmidt, and C. Foss (2014), Understanding magnetism in the Giles  
 677 Complex, Musgrave Block, SA, *Preview*, 171, 41–44.
- 678 Ayachit, U. (2015), *The ParaView Guide: A Parallel Visualization Application*, Kitware.
- 679 Beckwith, M. S., K. S. Beckwith, P. Sikorski, N. T. Skogaker, T. H. Flo, and Ø. Halaas  
 680 (2015), Seeing a mycobacterium-infected cell in nanoscale 3d: correlative imaging by  
 681 light microscopy and fib/sem tomography, *PLoS One*, 10(9).

- 682 Brogden, V. (2015), Method and apparatus for slice and view sample imaging, uS Patent  
683 9,218,940.
- 684 Butler, R. F., and S. K. Banerjee (1975), Theoretical single-domain grain size range in mag-  
685 netite and titanomagnetite, *J. Geophys. Res.*, *80*, 4049 – 4058.
- 686 Cao, S., W. Tirry, W. Van Den Broek, and D. Schryvers (2009), Optimization of a fib/sem  
687 slice-and-view study of the 3d distribution of ni<sub>4</sub>ti<sub>3</sub> precipitates in ni-ti, *Journal of mi-  
688 croscopy*, *233*(1), 61–68.
- 689 Church, N. S., K. Fabian, and S. A. McEnroe (2016), Nonlinear Preisach maps: Detecting  
690 and characterizing separate remanent magnetic fractions in complex natural samples, *Jour-  
691 nal of Geophysical Research-Solid Earth*, *121*(12), 8373–8395.
- 692 Cignoni, P., M. Callieri, M. Corsini, M. Dellepiane, F. Ganovelli, and G. Ranzuglia (2008),  
693 Meshlab: an open-source mesh processing tool., in *Eurographics Italian chapter confer-  
694 ence*, vol. 2008, pp. 129–136.
- 695 Davis, K. E. J. E., and P. S. Letters (1981), Magnetite rods in plagioclase as the primary car-  
696 rier of stable NRM in ocean floor gabbros., *Earth and Planetary Science Letters*, *55*(1),  
697 190–198.
- 698 Day, R., M. Fuller, and V. A. Schmidt (1977), Hysteresis properties of titanomagnetites:  
699 grain-size and compositional dependence, *Phys.Earth Planet. Inter.*, *13*, 260–267.
- 700 Ding, R., and I. P. Jones (2011), An fib–sem slice-and-view study of three-dimensional beta  
701 phase distribution in ti–6al–4v, *Journal of electron microscopy*, *60*(2), 149–155.
- 702 Dunlop, D. (1973), Superparamagnetic and single-domain threshold sizes in magnetite, *J.  
703 Geophys. Res.*, *78*, 1780–1793.
- 704 Dunlop, D. (2002), Theory an application of the day plot ( $m_{rs}/m_s$  versus  $h_{cr}/h_c$ ) 1. theoreti-  
705 cal curves and tests using titanomagnetite data, *J. Geophys. Res.*, *107*, 1–22.
- 706 Dunlop, D. J., and M.-M. Bina (1977), The coercive force spectrum of magnetite at high  
707 temperatures: evidence for thermal activation below the blocking temperature, *Geophys. J.  
708 R. astr. Soc.*, *51*, 121–147.
- 709 Dunlop, D. J., and Ö. Özdemir (1997), *Rock magnetism: fundamentals and frontiers*, vol. 3,  
710 Cambridge university press.
- 711 Einsle, J. F., R. J. Harrison, T. Kasama, P. Ó. Conbhuí, K. Fabian, W. Williams, L. Wood-  
712 land, R. R. Fu, B. P. Weiss, and P. A. Midgley (2016), Multi-scale three-dimensional char-  
713 acterization of iron particles in dusty olivine: Implications for paleomagnetism of chon-  
714 dritic meteorites, *American Mineralogist*, *101*(9), 2070–2084.
- 715 Enkin, R. J., and D. Dunlop (1987), A micromagnetic study of pseudo-single-domain rema-  
716 nence in magnetite, *J. Geophys. Res.*, *92*, 12,726–12,740.
- 717 Enkin, R. J., and W. Williams (1994), Three-dimensional micromagnetic analysis of stability  
718 in fine magnetic grains, *J. Geophys. Res.*, *99*, 611–618.
- 719 Fabian, K., and A. Hubert (1999), Shape-induced pseudo-single-domain remanence, *Geo-  
720 physical Journal International*, *138*(3), 717–726.
- 721 Fabian, K., and V. P. Shcherbakov (2018), Energy barriers in three-dimensional micromag-  
722 netic models and the physics of thermoviscous magnetization, *Geophysical Journal Inter-  
723 national*, *215*(1), 314–324, doi:10.1093/gji/ggy285.
- 724 Fabian, K., A. Kirchner, W. Williams, F. Heider, T. Leibl, and A. Hubert (1996), Three-  
725 dimensional micromagnetic calculations for magnetite using FFT, *Geophys. J. Int.*, *124*,  
726 89–104.
- 727 Fagerland, S. K. (2014), Investigation of focused ion beam/scanning electron microscope  
728 parameters for slice and view and energy dispersive x-ray spectroscopy of embedded brain  
729 tissue, Master’s thesis, Institut for fysikk.
- 730 Fang, Q., and D. A. Boas (2009), Tetrahedral mesh generation from volumetric binary and  
731 grayscale images, in *2009 IEEE International Symposium on Biomedical Imaging: From  
732 Nano to Macro*, IEEE, doi:10.1109/isbi.2009.5193259.
- 733 Feinberg, J. M., H.-R. Wenk, P. R. Renne, and G. R. Scott (2004), Epitaxial relationships of  
734 clinopyroxene-hosted magnetite determined using electron backscatter diffraction (ebsd)  
735 technique, *American Mineralogist*, *89*(2-3), 462–466.

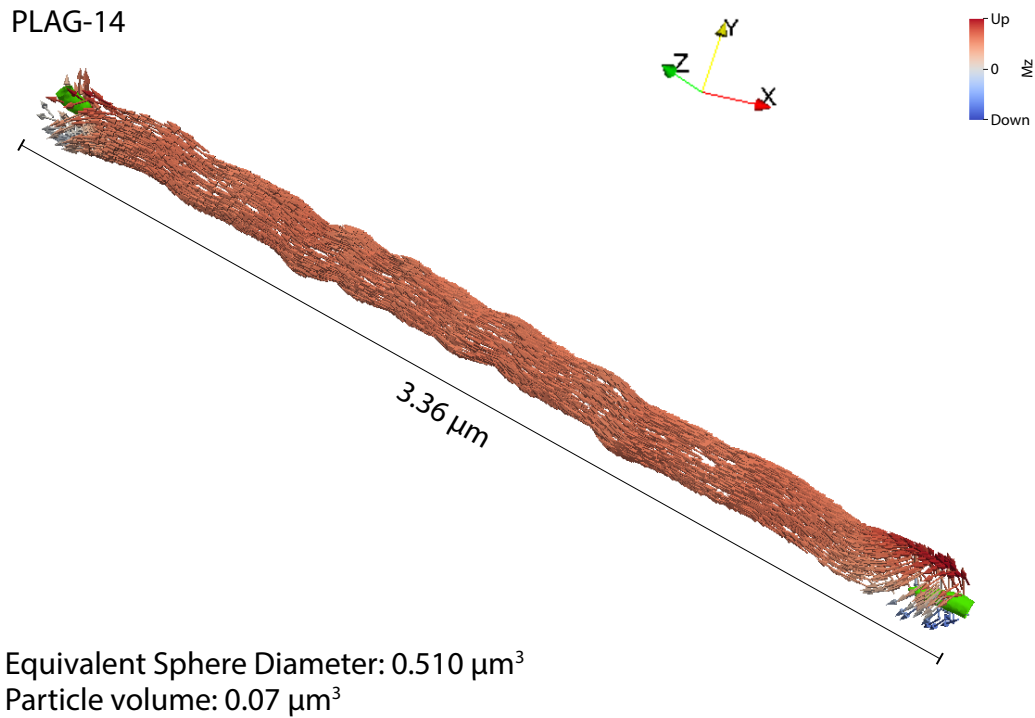
- 736 Feinberg, J. M., G. R. Scott, P. R. Renne, and H.-R. Wenk (2005), Exsolved magnetite inclu-  
737 sions in silicates: Features determining their remanence behavior, *Geology*, *33*(6), 513,  
738 doi:10.1130/g21290.1.
- 739 Fleet, M., G. A. Bilcox, and R. L. Barnett (1980), Oriented magnetite inclusions in pyrox-  
740 enes from the Grenville Province, *The Canadian Mineralogist*, *18*(1), 89–99.
- 741 Frandsen, C., S. Stipp, S. McEnroe, M. Madsen, and J. Knudsen (2004), Magnetic domain  
742 structures and stray fields of individual elongated magnetite grains revealed by magnetic  
743 force microscopy (MFM), *Physics of the Earth and Planetary Interiors*, *141*(2), 121–129,  
744 doi:10.1016/j.pepi.2003.12.001.
- 745 Guehrs, E., M. Schneider, C. M. Günther, P. Hessing, K. Heitz, D. Wittke, A. L.-S. Oliver,  
746 N. Jakubowski, J. Plendl, S. Eisebitt, et al. (2017), Quantification of silver nanoparticle  
747 uptake and distribution within individual human macrophages by fib/sem slice and view,  
748 *Journal of nanobiotechnology*, *15*(1), 21.
- 749 Harrison, R. J., X. Zhao, P. Hu, T. Sato, D. Heslop, A. R. Muxworthy, H. Oda, V. S. Kuppili,  
750 and A. P. Roberts (2019), Simulation of remanent, transient, and induced forc diagrams for  
751 interacting particles with uniaxial, cubic, and hexagonal anisotropy, *Journal of Geophys-  
752 ical Research: Solid Earth*, *124*(12), 12,404–12,429.
- 753 Hodych, J. (1990), Magnetic hysteresis as a function of low temperature in rocks: evidence  
754 for internal stress control of remanence in multi-domain and pseudo-single-domain mag-  
755 netite, *Phys. Earth Planet. Inter.*, *64*(1), 21–36.
- 756 Hubert, A., and R. Schäfer (1998), *Magnetic Domains*, Springer, Berlin, Heidelberg, New  
757 York.
- 758 Joffe, I., and R. Heuberger (1974), Hysteresis properties of distributions of cubic single-  
759 domain ferromagnetic particles, *Phil. Mag.*, *29*, 1051–1059.
- 760 Khakhalova, E., B. M. Moskowitz, W. Williams, A. R. Biedermann, and P. Solheid (2018),  
761 Magnetic vortex states in small octahedral particles of intermediate titanomagnetite, *Geo-  
762 chemistry, Geophysics, Geosystems*, *19*(9), 3071–3083, doi:10.1029/2018gc007723.
- 763 Krása, D., C. D. W. Wilkinson, N. Gadegaard, X. Kong, H. Zhou, A. P. Roberts, A. R. Mux-  
764 worthy, and W. Williams (2009), Nanofabrication of two-dimensional arrays of mag-  
765 netite particles for fundamental rock magnetic studies, *Journal of Geophysical Research*,  
766 *114*(B2), doi:10.1029/2008jb006017.
- 767 Lappe, S.-C. L., J. M. Feinberg, A. Muxworthy, and R. J. Harrison (2013), Comparison and  
768 calibration of nonheating paleointensity methods: A case study using dusty olivine, *Geo-  
769 chemistry, Geophysics, Geosystems*, *14*(7), 2143–2158.
- 770 Lascu, I., R. J. Harrison, Y. Li, J. R. Muraszko, J. E. T. Channell, A. M. Piotrowski, and  
771 D. A. Hodell (2015), Magnetic unmixing of first-order reversal curve diagrams using prin-  
772 cipal component analysis, *Geochemistry, Geophysics, Geosystems*, *16*(9), 2900–2915, doi:  
773 10.1002/2015gc005909.
- 774 Maier, W., H. Howard, R. Smithies, S. Yang, S.-J. Barnes, H. O'Brien, H. Huhma,  
775 and S. Gardoll (2015), Magmatic ore deposits in mafic–ultramafic intrusions  
776 of the giles event, western australia, *Ore Geology Reviews*, *71*, 405–436, doi:  
777 10.1016/j.oregeorev.2015.06.010.
- 778 McEnroe, S., J. Skilbrei, P. Robinson, F. Heidelbach, F. Langenhorst, and L. Brown (2004),  
779 Magnetic anomalies, layered intrusions and mars, *Geophysical Research Letters*, *31*(19).
- 780 McNab, T. K., R. A. Fox, and A. J. F. Boyle (1968), Some magnetic properties of magnetite  
781 ( $\text{Fe}_3\text{O}_4$ ) microcrystals, *J. Appl. Phys.*, *39*, 5703 – 5711.
- 782 Mulders, J. L., G. Knott, and B. Lich (2006), Dualbeam slice & view: Practical aspects for  
783 collecting 3d cortex image data, *Microscopy and Microanalysis*, *12*(S02), 1324–1325.
- 784 Muxworthy, A. R., and W. Williams (2006), Critical single-domain/multidomain grain sizes  
785 in noninteracting and interacting elongated magnetite particles: Implications for magneto-  
786 somes, *Journal of Geophysical Research: Solid Earth*, *111*(B12).
- 787 Muxworthy, A. R., and W. Williams (2008), Critical superparamagnetic/single-domain grain  
788 sizes in interacting magnetite particles: implications for magnetosome crystals, *Journal of  
789 The Royal Society Interface*, *6*(41), 1207–1212, doi:10.1098/rsif.2008.0462.

- 790 Nagy, L., W. Williams, L. Tauxe, and A. R. Muxworthy (2019), From nano to micro: Evolution  
791 of magnetic domain structures in multidomain magnetite, *Geochemistry, Geophysics,*  
792 *Geosystems*, 20(6), 2907–2918.
- 793 Néel, L. (1955), Some theoretical aspects of rock-magnetism, *Adv. Phys.*, 4, 191–243.
- 794 Ó Conbhuí, P., W. Williams, K. Fabian, P. Ridley, L. Nagy, and A. R. Muxworthy (2018),  
795 MERRILL: Micromagnetic Earth related robust interpreted language laboratory, *Geo-*  
796 *chemistry, Geophysics, Geosystems*, 19, doi:10.1002/2017gc007279.
- 797 Parry, L. (1982), Magnetization of immobilized particle dispersions with two distinct particle  
798 sizes, *Physics of the Earth and Planetary interiors*, 28(3), 230–241.
- 799 Rave, W., K. Fabian, and A. Hubert (1998a), Magnetic states of small cubic particles with  
800 uniaxial anisotropy, *J. Magn. Magn. Mater.*, 190, 332–348.
- 801 Rave, W., K. Ramstöck, and A. Hubert (1998b), Corners and nucleation in micromagnetics,  
802 *J. Magn. Magn. Mater.*, 183, 329–333.
- 803 Renne, P. R., G. R. Scott, J. M. G. Glen, and J. M. Feinberg (2002), Oriented inclusions of  
804 magnetite in clinopyroxene: Source of stable remanent magnetization in gabbros of the  
805 messum complex, namibia, *Geochemistry, Geophysics, Geosystems*, 3(12), 1–11, doi:  
806 10.1029/2002gc000319.
- 807 Roberts, A. P., L. Tauxe, D. Heslop, X. Zhao, and Z. Jiang (2018), A critical appraisal of the  
808 “Day” diagram, *Journal of Geophysical Research: Solid Earth*, 123(4), 2618–2644, doi:  
809 10.1002/2017jb015247.
- 810 Robinson, P., S. A. McEnroe, N. Miyajima, K. Fabian, and N. Church (2016), Remanent  
811 magnetization, magnetic coupling, and interface ionic configurations of intergrown rhom-  
812bohedral and cubic Fe-Ti oxides: A short survey, *American Mineralogist*, 101(3-4), 518–  
813530.
- 814 Schmidt, R., and K. Singh (2010), Meshmixer: an interface for rapid mesh composition, in  
815 *ACM SIGGRAPH 2010 Talks*, pp. 1–1.
- 816 Soffel, H. C. (1977), Domain structure of titanomagnetites and its variation with temperature,  
817 *J. Geomag. Geoelectr.*, 29, 277–284.
- 818 Stacey, F. D. (1961), Theory of magnetic properties of igneous rocks in alternating fields,  
819 *Phil. Mag.*, 6, 1241–1260.
- 820 Stacey, F. D., and S. K. Banerjee (1974), *The physical principles of rock magnetism*, Elsevier,  
821 Amsterdam.
- 822 Tarduno, J. A., R. D. Cottrell, R. K. Bono, H. Oda, W. J. Davis, M. Fayek, O. van’t Erve,  
823 F. Nimmo, W. Huang, E. R. Thern, et al. (2020), Paleomagnetism indicates that primary  
824 magnetite in zircon records a strong hadean geodynamo, *Proceedings of the National*  
825 *Academy of Sciences*, 117(5), 2309–2318.
- 826 Tauxe, L., H. N. Bertram, and C. Seberino (2002), Physical interpretation of hysteresis loops:  
827 Micromagnetic modeling of fine particle magnetite, *Geochemistry Geophysics Geosys-*  
828 *tems*, 3, art. no.–1055.
- 829 ter Maat, G. W., K. Fabian, N. S. Church, and S. A. McEnroe (2020), Separating geometry-  
830 from stress-induced remanent magnetization in magnetite with ilmenite lamellae  
831 from the Stardalur basalts, Iceland, *Geochemistry, Geophysics, Geosystems*, 21, doi:  
832 10.1029/2019gc008761.
- 833 Tseng, Q., I. Wang, E. Duchemin-Pelletier, A. Azioune, N. Carpi, J. Gao, O. Filhol, M. Piel,  
834 M. Théry, and M. Balland (2011), A new micropatterning method of soft substrates re-  
835veals that different tumorigenic signals can promote or reduce cell contraction levels, *Lab*  
836 *Chip*, 11, 2231–2240, doi:10.1039/C0LC00641F.
- 837 Usui, Y., T. Shibuya, Y. Sawaki, and T. Komiya (2015), Rock magnetism of tiny exsolved  
838 magnetite in plagioclase from a Paleoproterozoic granitoid in the Pilbara craton, *Geochem-*  
839 *istry, Geophysics, Geosystems*, 16(1), 112–125, doi:10.1002/2014gc005508.
- 840 Warr, L. N., J. Wojatschke, B. M. Carpenter, C. Marone, A. M. Schleicher, and B. A. van der  
841 Pluijm (2014), A “slice-and-view”(fib–sem) study of clay gouge from the safod creeping  
842 section of the san andreas fault at 2.7 km depth, *Journal of Structural Geology*, 69, 234–  
843244.



- 844 Wenk, H.-R., K. Chen, and R. Smith (2011), Morphology and microstructure of magnetite  
 845 and ilmenite inclusions in plagioclase from adirondack anorthositic gneiss, *American Min-  
 846 eralogist*, 96(8-9), 1316–1324.
- 847 Witt, A., K. Fabian, and U. Bleil (2005), Three-dimensional micromagnetic calculations for  
 848 naturally shaped magnetite: Octahedra and magnetosomes, *Earth and Planetary Science  
 849 Letters*, 233(3-4), 311–324.

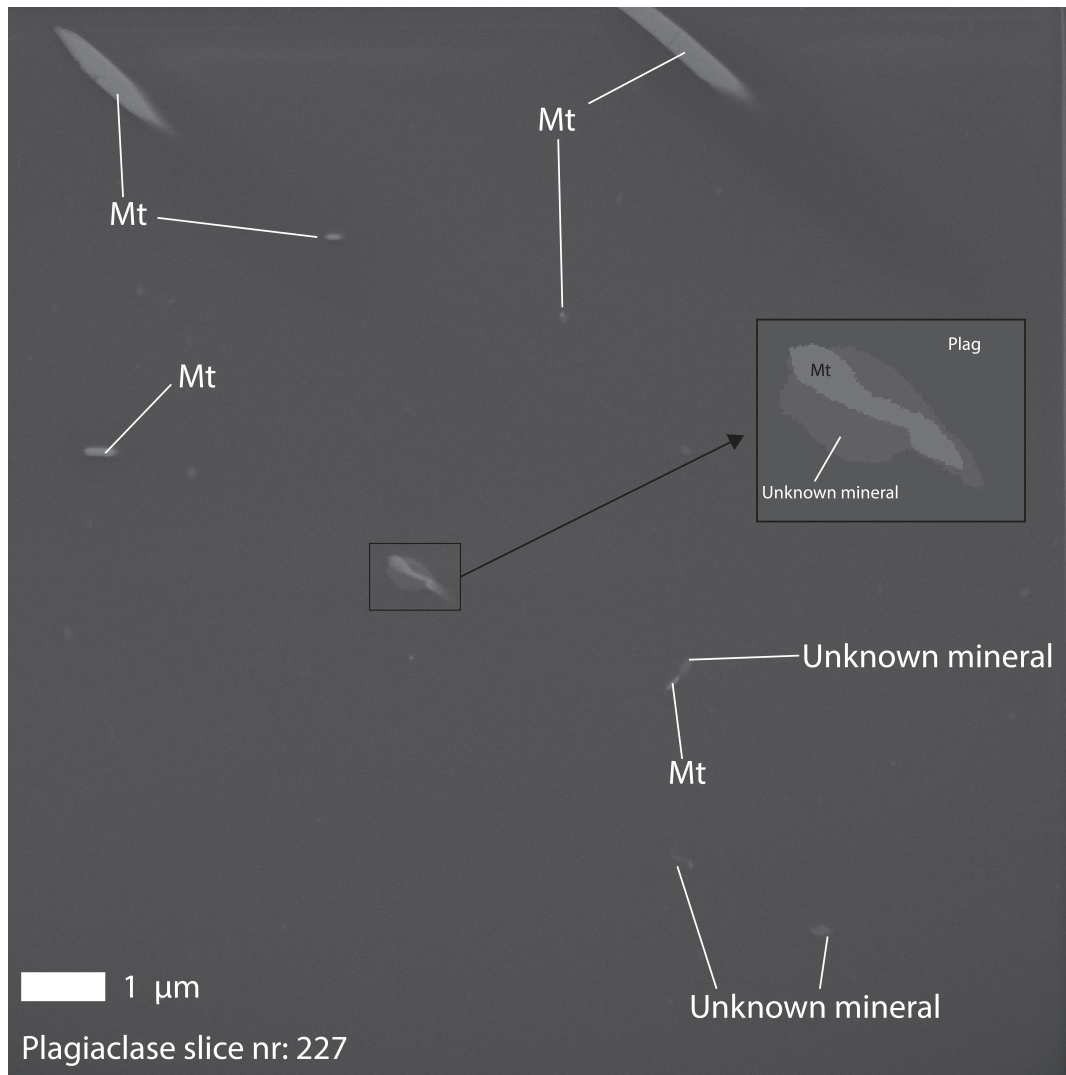
850 **A:**



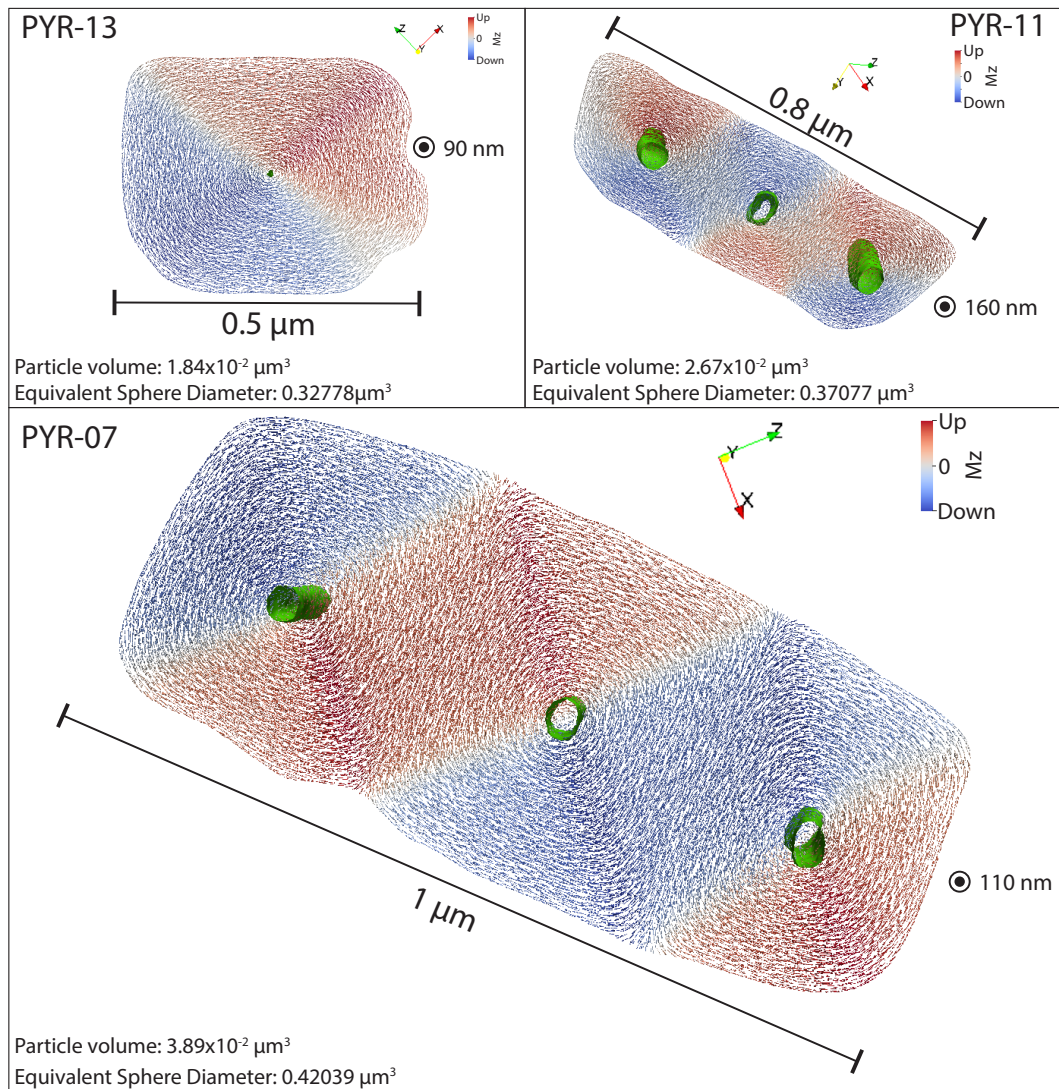
851 **Figure A.1.** Magnetic model of a  $3.36 \mu\text{m}$  long magnetite needle in plagioclase, minimized in zero-field.

### 862 Acknowledgments

863 The FIB-SEM and Slice-and-View work was performed at the NTNU Nanolab and the Nor-  
 864 wegian infrastructure for micro- and nanofabrication as part of a PhD-project funded by  
 865 NTNU. The Research Council of Norway is acknowledged for the support to the Norwegian  
 866 Micro- and Nano-Fabrication Facility, NorFab, Project 245963/F50. The calculations and  
 867 usage of a Linux server was possible with granted access by IGP and a designated worksta-  
 868 tion with remote access. We thank Phil Schmidt for providing these very interesting samples.  
 869 We also wish to thank Lesleis Nagy and Thomas Berndt for their detailed and constructive  
 870 reviews. Particle meshes for replication of results and appendix figures can be found at the  
 871 online repository dataverse.no (<https://doi.org/10.18710/Q0F8PA>).



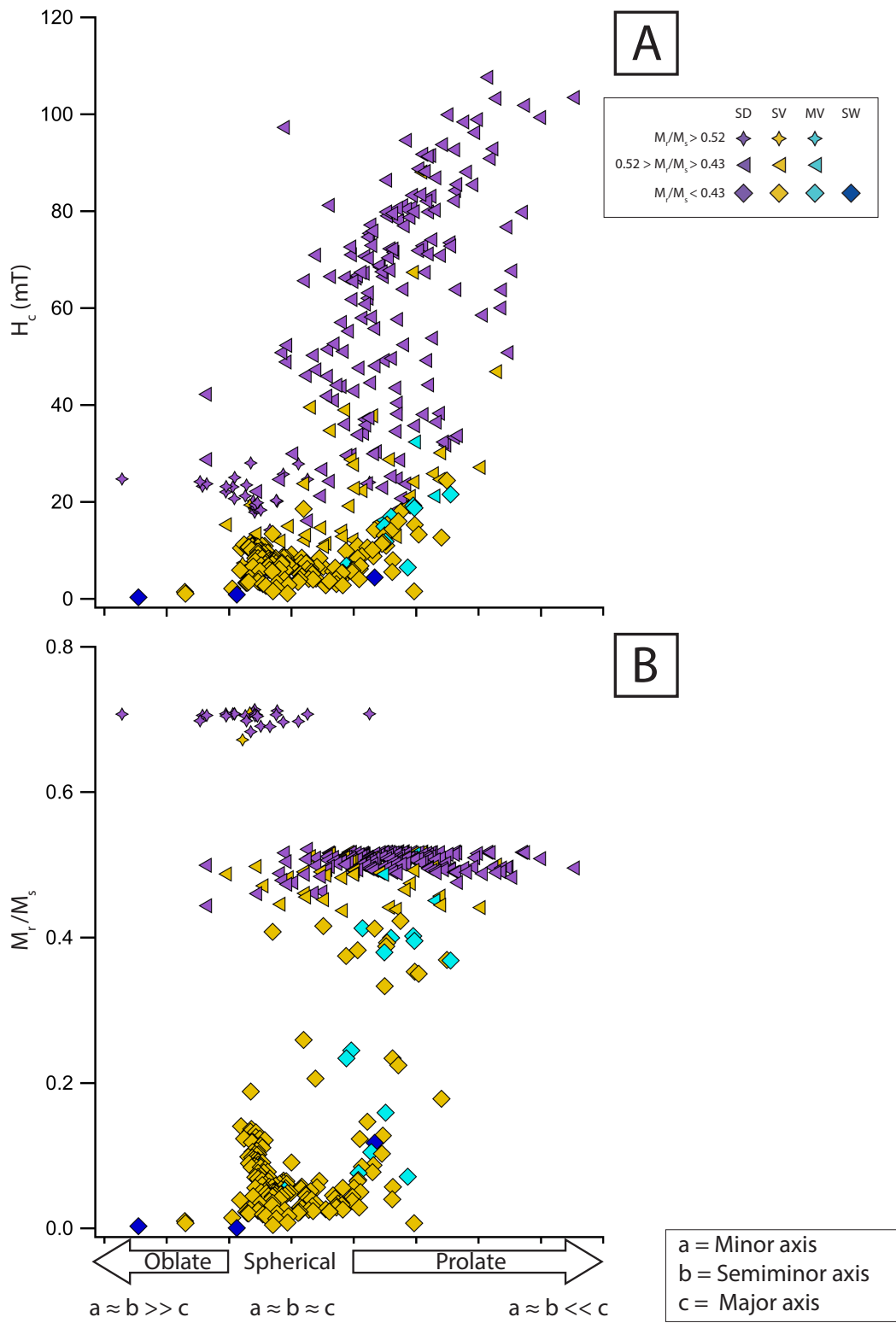
852 **Figure A.2.** Image of slice 227 in plagioclase showing mainly magnetite associated with an unknown  
853 mineral that often occurs at an edge or coated around the magnetite.



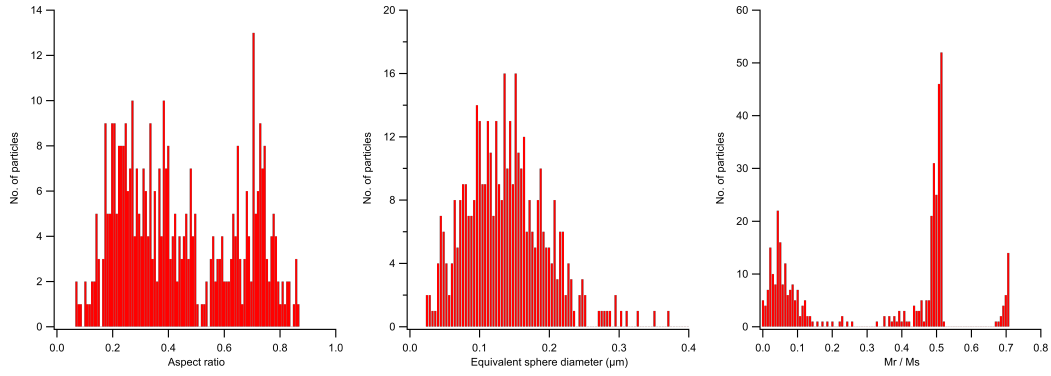
854

**Figure A.3.** Rendering of the zero-field minimization of large particles labeled as SW.

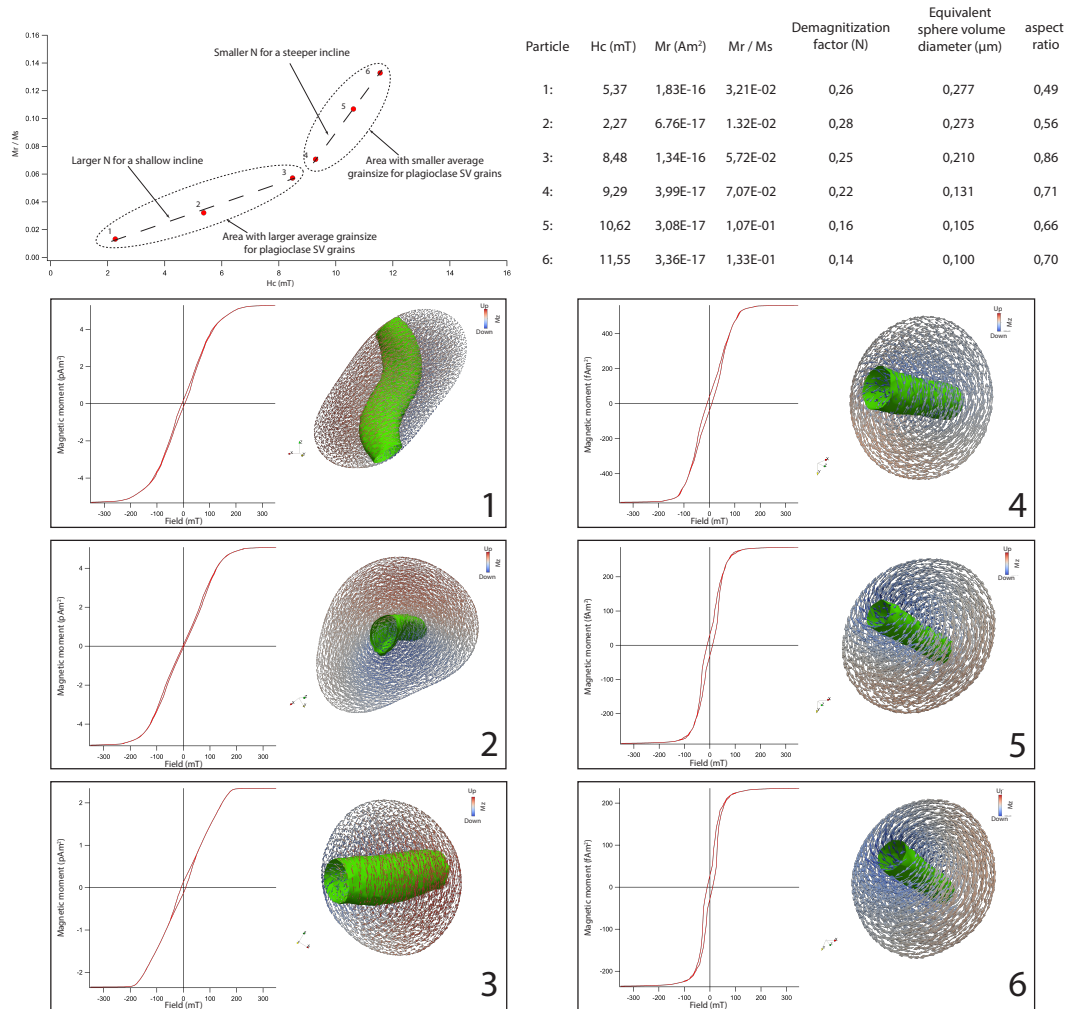




855 **Figure A.4.** a) Coercivity,  $H_c$  and b) remanence ratio,  $M_{rs}/M_s$  as a function of particle prolate-  
 856 ness/oblateness.



857 **Figure A.5.** Histograms showing the particle distribution of  $M_r/M_{r,S}$ , Aspect ratio and Equivalent sphere  
 858 diameter.



859 **Figure A.6.** Showing a selection of six particles that represents the two different slopes observed among  
 860 SV plagioclase particles. N was calculated and volume corrected from the slope between  $M_r$  and  $H_c$ . The  
 861 dotted ellipses showing the clusters of similar particle sizes that differs in the two trends.