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Sensitivity analysis for comparison, validation and physical-legitimacy of

- neural network-based hydrological models
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 SHORT TITLE: Sensitivity analysis for comparison, validation and physical-legitimacy

10

11 ABSTRACT

12 This paper addresses the difficult question of how to perform meaningful comparisons between neural network-based hydrological models and alternative modelling approaches. 13 14 Standard, goodness-of-fit metric approaches are limited since they only assess numerical 15 performance and not physical legitimacy of the means by which output is achieved. 16 Consequently, the potential for general application or catchment transfer of such models is 17 seldom understood. This paper presents a partial derivative, relative sensitivity analysis 18 method as a consistent means by which the physical legitimacy of models can be evaluated. 19 It is used to compare the behaviour and physical rationality of a generalised linear model 20 and two neural network models for predicting median flood magnitude in rural catchments. 21 The different models perform similarly in terms of goodness-of-fit statistics, but behave 22 quite distinctly when the relative sensitivities of their parameters are evaluated. The neural 23 solutions are seen to offer an encouraging degree of physical legitimacy in their behaviour, 24 over that of their generalised linear modelling counterpart, particularly when overfitting is 25 constrained. This indicates that neural solutions are preferable models for transferring into ungauged catchments. Thus, the importance of understanding both model performance and 26 27 physical legitimacy when comparing neural models with alternative modelling approaches is 28 demonstrated.

- **KEYWORDS** | generalised model, index flood, neural network, partial derivative, sensitivity
- 31 analysis, ungauged catchment

33 INTRODUCTION

34 This paper presents an approach for delivering greater meaning from the comparison of artificial neural network (ANN) models with alternative modelling approaches in 35 36 hydrological studies. ANN-based hydrological models are most commonly applied as black-37 box tools and the internal mechanisms by which the model output is generated are not normally explored in hydrological terms. Used in this way, an ANN's primary purpose is the 38 optimisation of complex, non-linear relations between a specific set of hydrological input 39 40 and output data, and standard goodness-of-fit procedures may, therefore, be considered an adequate basis by which to compare its performance to that of other models (Klemes, 1986; 41 Refsgaard and Knusden, 1996). Indeed, assessments of goodness-of-fit have been widely 42 43 used in comparative hydrological modelling studies to argue that ANN models can perform as well as, or better than alternative modelling approaches (e.g. Shrestha and Nestmann, 44 45 2009; Mount and Abrahart, 2011). However, such arguments are informed solely by the 46 degree of optimisation that is achieved by each model. They say nothing about the means 47 by which different models achieve their performance and the relative merits of these 48 alternative means. Indeed, when ANN models are applied solely as black-boxes, their potential relative to other modelling approaches can never be properly understood in a 49 50 generalised or transferrable manner because the extent to which their modelling mechanisms conform to physically-based, hydrological domain knowledge remains untested 51 (Howes and Anderson, 1988; Sargent, 2011). Consequently, critical questions about 52 53 whether ANN modelling mechanisms are more or less reflective of real-world hydrological 54 processes than alternative models are seldom addressed directly (Minns and Hall, 1996; 55 Abrahart *et al.*, 2011), and the relative extent to which they are able to deliver hydrological process insights (i.e. Caswell's (1976) model duality) is not normally evaluated. The purpose
of this paper is to present a method by which these questions may be addressed.

More informative approaches to model comparison are required that explicitly 58 consider the internal behaviours of the different models and assess them according to their 59 60 conformance with the logical, rational and physical expectations of the modeller (c.f. Robinson, 1997). This process is termed model legitimisation and is discussed in a 61 philosophical context by Oreskes et al. (1994) and an applied, hydrological modelling 62 63 context by Mount et al. (in press). Sensitivity analysis (Hamby, 1994) is an important and 64 effective means by which the legitimacy of a hydrological model may be explored. It has been widely applied in conceptual and physically-based modelling over several decades (e.g. 65 McCuen, 1973; Beven and Binley, 1992; Schulz and Huwe, 1999; Radwan et al., 2004; 66 67 Pappenberger et al., 2008; Mishra, 2009; Zhang et al., 2012). A variety of approaches have 68 been used including local (e.g. Turanayi and Rabitz, 2000; Spruill et al., 2000; Holvoet et al., 69 2005; Hill and Tiedeman, 2007), regional (e.g. Spear and Hornberger, 1980) and global-scale methods (Muleta and Nicklow, 2005; Salteli et al., 2008). By contrast, sensitivity analysis 70 has not been widely adopted in ANN modelling studies beyond a few, isolated examples 71 72 (Sudheer, 2005; Nourani and Fard, 2012). This is presumably because the equations that 73 relate inputs and outputs in an ANN are considered complex, inaccessible and difficult to interpret (Aytek et al., 2008; Abrahart et al., 2009), making exploration of model sensitivity 74 75 via direct analysis of the governing equations difficult. Nonetheless, recent progress has 76 been made (Yeung et al., 2010) and relative sensitivity analysis techniques for ANNs have 77 made it possible to assess the internal, mechanistic legitimacy of such models (Abrahart et 78 al., 2012b; Mount et al., in press). However, the focus of these studies has so far been 79 restricted to mechanical considerations. The application of sensitivity analysis to evaluate the physical legitimacy of ANN-based hydrological models, and thus the degree to which
they can be generalised and transferred, remains an outstanding task.

82 In this paper, we apply a sensitivity analysis method that can be used to compare the 83 physical legitimacy of ANN-based hydrological models and alternative model counterparts in 84 a direct manner. We exemplify the method by comparing the performance and physical legitimacy of a pair of ANN-based models and an established generalised linear model 85 86 (GLM) for median flood magnitude prediction in ungauged catchments in the UK. First 87 order, partial derivatives of each model's response function are computed, interpreted and used as a consistent means by which the physical legitimacy of each model can be evaluated 88 89 and compared. This focus on response function behaviour is distinctly different to past 90 efforts to assess the physical legitimacy of ANN models, which have traditionally explored 91 internal structural components, such as weights (Abrahart et al., 1999; Olden and Jackson, 92 2002; Anctil et al., 2004; Kingston et al., 2003,2005,2006,2008) and units (Wilby et al., 2003; 93 Jain et al., 2004; Sudheer and Jain, 2004; See et al., 2008; Fernando and Shamseldin, 2009; 94 Jain and Kumar, 2009). However, the uniqueness of ANN structures means that the 95 information derived from them cannot easily be compared directly with that derived from 96 alternative models with different internal structures - thus limiting the comparative value of 97 the information. To overcome this problem, we here assess the physical legitimacy of an 98 ANN's overall response function using a standard relative sensitivity-based method that can be consistently and directly replicated across a range of alternative model types and that is 99 100 widely understood and accepted by hydrologists. Consequently, an evaluation of the 101 physical legitimacy of the means by which each model's performance is obtained 102 accompanies the usual assessments of output validity; enabling the extent to which each 103 model delivers a transferable, general solution to be considered.

105 COMPARING GLM AND ANN-BASED MODELS FOR UNGAUGED CATCHMENT PREDICTION 106 IN THE UK

The modelling of hydrological responses in ungauged catchments remains an important 107 108 focus of research for hydrologists, especially as the majority of the world's river catchments 109 remain ungauged or poorly gauged. In such catchments, the application of distributed physically-based models and statistical approaches is hampered by a lack of input parameter 110 111 knowledge and datasets. Consequently, lumped models which relate broad physiographic, 112 hydrogeologic and climatologic catchment descriptors to flood frequency curves, have long 113 been recognised as offering potential (Rodriguez-Iturbe and Valdes, 1979; Grover et al., 2002). 114

The standard UK method (Natural Environment Research Council, 1975; Vogel and Kroll, 1992; Schrieber and Demuth, 1997) models the relationship between the median of the annual flood series (*QMED*) and a set of regionalised catchment descriptors for rivers in the national, gauged network. The modelled relationship is then applied to ungauged catchments and used to estimate *QMED*, which is subsequently multiplied by a standard, dimensionless growth curve to estimate flood frequency (Institute of Hydrology, 1999).

Four catchment descriptors are used in the standard UK methodology: 1) *AREA* (catchment area in km²); 2) *SAAR* (standard-period average annual rainfall in mm); 3) *FARL* (flood attenuation due to reservoirs and lakes); 4) *BFIHOST* (baseflow index derived from HOST data; Boorman *et al.*, 1995).

125 These catchment descriptors can be thought of as physical controls of *QMED* potential. 126 *SAAR* controls the hydrological inputs to the catchment, *AREA* controls the scaling of the 127 catchment response, whilst *BFIHOST* and *FARL* control the degree of buffering of the input-128 output signal.

Of central importance to the above method is the model that is used to relate *QMED* and the catchment descriptors. These relationships are non-linear and not well represented by standard multiple linear regression. Therefore, the most recent UK method described applies a range of non-linear transformations within a generalised linear modelling (GLM) framework (Kjeldsen *et al.*, 2008; Kjeldsen and Jones, 2009; Kjeldsen and Jones, 2010). The end product is a non-linear regression equation (see Equation 1) from which *QMED* can be estimated directly from the four catchments descriptors.

136 ANN models are also very effective at optimising complex, non-linear relations in 137 hydrological data (American Society of Civil Engineers 2000a,b; Maier and Dandy, 2000; Dawson and Wilby, 2001; Maier et al., 2010; Abrahart et al., 2010; 2012b) and a number of 138 139 studies have highlighted their potential in ungauged catchment prediction (Liong et al., 140 1994; Muttiah et al., 1997; Hall and Minns, 1998; Hall et al., 2000; Dastorani and Wright, 2001; Dawson et al., 2006; Dastorani et al., 2010). Indeed, the UK relationship between 141 142 QMED and catchment descriptors has also been modelled using ANNs and been shown to 143 deliver comparable levels of fit when compared to GLMs (Dawson et al., 2006). However, it 144 remains unclear whether the two modelling approaches are similarly comparable with respect to their physical legitimacy. Models with greater physical legitimacy should be more 145 generally transferrable to new catchment settings. Therefore, determining the physical 146 147 legitimacy of each model is an important element in delivering a physically informed evaluation of how robustly it can be expected to transfer from the gauged catchments upon 148 149 which it is developed, to the ungauged catchments in which it is intended to be applied.

150 In the following sections, the importance of evaluating both model performance and 151 physical legitimacy in ANN model comparisons is exemplified by contrasting the 152 performance and legitimacy of the standard GLM method for QMED prediction with two different ANN-based model counterparts. Its use as an example is particularly appropriate 153 154 because the model inputs and outputs are all physical-based measurements, meaning that 155 patterns observed in inputs and output relations can be interpreted directly in physical 156 terms, also the number of model inputs is relatively small, the first order partial derivatives 157 can be computed for the GLM and directly compared with those of the ANN-based models, 158 and the results of the analysis have real-world relevance and application.

159

160 **Data**

161 A GLM model and two counterpart ANN models for QMED estimation are developed for 162 comparison, with the model inputs conforming to the four used in the standard UK 163 methodology. These inputs were extracted from a pre-filtered set of HiFlows-UK rural catchment data, available at (http://www.environment-agency.gov.uk/hiflows/97503.aspx). 164 165 AREA values are derived from the Centre for Ecology and Hydrology's Integrated 166 Hydrological Digital Terrain Model (based on a 50m grid) and represent surface catchment 167 area projected onto a horizontal plane, draining to the gauging station (Marsh and 168 Hannaford, 2008: 5). SAAR values are derived from UK precipitation records over the 169 standard period 1961-1990. FARL provides a guide to the degree of flood attenuation 170 attributable to reservoirs and lakes above the gauging station. The index ranges from zero (complete attenuation) to one (no attenuation) with values < 0.8 representing a substantial 171 influence on flood response. BFIHOST is derived from the HOST (Hydrology of Soil Types) soil 172 173 data classification and ranges from zero (impermeable) to one (completely permeable). In undisturbed catchments, a strong association exists between Baseflow Index (derived from
archived gauged daily mean flows) and *BFIHOST*. The relationships between *QMED and AREA, SAAR and FARL* are positive, whilst that between *QMED* and *BFIHOST* is negative.

The data from which our models are derived are almost identical to those from 177 178 which the GLM that is published in the revitalised UK Flood Estimation Handbook (Kjeldsen et al., 2008) has been developed, and full particulars of the Hi-Flows UK data set can be 179 found in this handbook. A statistical summary of our dataset is provided in Table 1. Some 180 181 minor discrepancies exist between the data used in this study and that used by Kjeldsen et 182 al. (2008) due to our use of the public release version of HiFlows-UK 3.02 rather than the 183 pre-release version originally used. Specifically, our dataset comprises 597 rural catchment 184 records rather than the 602 used previously, and we use an unadjusted flood attenuation 185 variable.

186

187 Model development procedures

188 Three models were developed for comparison.

- QMED_{GLM} a GLM developed on all 597 catchment records, using the methodology
 outlined in Kjeldsen *et al.* (2008).
- ANN_A an optimised ANN, selected from 180 candidate solutions of varying
 complexity and training iterations according to both its goodness-of-fit performance
 and avoidance of evident overfitting.
- ANN_B a purposely over-trained version of ANN_A in which the number of training
 iterations was artificially extended to deliver an overfitted solution. It is included as
 a means of exemplifying the impact of ANN overfitting on the physical legitimacy of a
 network response function.

QMED_{GLM} was developed in accordance with the method of Kjeldsen *et al.* (2008).
Despite the minor differences in the dataset noted above, the resultant regression equation
(Equation 1) remains almost identical to Kjeldsen's original:

201

202
$$QMED_{GLM} = 8.6704AREA^{0.8568} 0.1550^{(\frac{1000}{SAAR})} FARL^{3.3662} 0.0380^{BFIHOST}$$
(1)

203

204 ANN_A and ANN_B comprise a Multi-Layer Perceptron (MLP), with one hidden layer, 205 trained using error back propagation (Rumelhart et al., 1986). The basic structure of these 206 networks is shown schematically in Figure 1. The ANN consists of a number of units or 207 neurons arranged in three layers (although additional hidden layers can be incorporated). 208 The units in the input layer distribute the inputs to the units in the hidden layer, which in 209 turn pass their outputs to the output layer (usually consisting of a single output neuron). Each neuron consists of a weighted set of inputs and an activation function – typically the 210 211 logistic sigmoid function (Equation 2). The output from a single unit is calculated by 212 applying this sigmoid function to the weighted sum of its inputs.

213

- 214 $f(x) = \frac{1}{1+e^{-x}}$ (2)
- 215

Training such networks using back propagation involves presenting the ANN with training data, calculating the error of the network's output with respect to the observed values, propagating this error backwards through the network and adjusting the input weights to the neurons accordingly (to reduce this error). This process must be repeated many times, making minor adjustments to the weights of each cycle (or epoch), until the ANN begins to map input values to the correct output response. The amount by which the weights are adjusted each time can be manipulated by using a learning rate multiplier. Readers that are unfamiliar with ANN concepts, structures and training methods are referred to Kattan *et al.* (2011) or Nelson (2011).

225 The simplicity of this ANN has enabled the development of computational methods 226 for delivering first-order partial derivatives of its response function (Hashem, 1992), which 227 we subsequently use as the basis for our comparative assessment of model legitimacy (see 228 Section 3). This standard ANN has been successfully used in many hydrological studies in 229 the past (Abrahart et al., 2012a) and provides an established non-linear modelling benchmark for ANN studies and a starting point against which more novel approaches can 230 subsequently be compared (Mount et al., 2012). Whilst it is recognised that more advanced 231 232 ANN structures might arguably deliver some additional optimisation advantages, the 233 computational methods required to quantify their response function partial derivatives, and 234 hence deliver directly comparable assessments of their physical legitimacy, are not readily 235 available. Their use is thus avoided in this study.

236 ANN_A was developed using the approach described in Dawson *et al.* (2006) in which 237 a large number of candidate ANNs are trained on a random subset of the data, partitioned 238 according to a 60% calibration to 40% cross-validation ratio. Although there is no agreed 239 standard for splitting the data, this ratio is widely accepted in hydrological modelling 240 (Mount and Abrahart, 2011; See and Openshaw, 2000). 180 candidate models containing 2, 241 3, 4, 5, 6, 7, 8, 9, 10 hidden units were developed with each candidate being trained for up 242 to 20,000 epochs in steps of 1,000, using a learning rate of 0.1 and a momentum value of 243 0.9. Each candidate model was cross-validated using the remaining 40% as a means of 244 preventing overfitting (Giustolisi and Laucelli, 2005; Piotrowski and Napiorkowski, 2013). Overfitting of each candidate solution was evaluated according to its cross-validation scores,
and the candidate solution displaying the best optimisation performance, whilst avoiding
apparent overfitting, was selected as the final model.

ANN_A has nine hidden units, and is trained for 4000 epochs. ANN_B, which we adopt as an example of an overfitted ANN, is structurally identical to ANN_A . However its training epochs have been artificially extended to ten times that of ANN_A (i.e. 40,000 epochs) to promote overfitting. The network unit weights and biases are provided in Table 2 and are used as the inputs to Equation 8, from which relative sensitivity can be computed.

253 It should be noted that the GLM and ANN models utilise the available data records 254 differently during model development. Whilst the GLM uses all 597 records to define the 255 model, each candidate ANN uses only the first 400 records to refine the model, and the 256 remaining 197 records to constrain it via cross-validation. Indeed, the apparent 257 inconsistency with which the GLM and ANN models use the available data could be cited as 258 an argument to negate the fairness of a direct comparison between them. However, this 259 stance fails to credit that both models do use all of the data in the model development 260 process; they just use it in a characteristically different manner that reflects the fundamental differences between each method. In this sense, the models are comparable; 261 262 not because they use the same data in the same way, but rather because each one's use of 263 the data is equally appropriate and justifiable in the context of its own model development 264 method.

265

266 MODEL PERFORMANCE AND PHYSICAL LEGITIMACY ASSESSMENT

267 Model performance evaluation

268 Each model's performance was evaluated using standard goodness-of-fit metrics to deliver output validation. To ensure a consistent approach the metrics were generated using 269 270 HydroTest (http://www.hydrotest.org.uk), a standardised, open access web site that performs the required numerical calculations (Dawson et al. 2007,2010). Each model's 271 performance is evaluated using RMSE (root mean squared error) and R^2 (R-squared – the 272 273 coefficient of determination) providing an overall measure of model performance; MSRE 274 (mean squared relative error) and MSLE (mean squared logarithmic error) providing two 275 additional measures of performance which place greater emphasis on errors occurring in 276 lower magnitude predictions. These comparative performance statistics are defined as

277
$$RMSE = \sqrt{\frac{\sum_{i=1}^{n} (Q_i - \hat{Q}_i)^2}{n}}$$
 (3)

278
$$R^{2} = \left[\frac{\sum_{i=1}^{n} (Q_{i} - \overline{Q})(\hat{Q} - \widetilde{Q})}{\sqrt{\sum_{i=1}^{n} (Q_{i} - \overline{Q})^{2} \sum_{i=1}^{n} (\hat{Q}_{i} - \widetilde{Q})^{2}}}\right]^{2}$$
(4)

279
$$MSE = \frac{1}{n} \sum_{i=1}^{n} (Q_i - \hat{Q}_i)^2$$
 (5)

280
$$MSLE = \frac{1}{n} \sum_{i=1}^{n} (\ln Q_i - \ln \hat{Q}_i)^2$$
 (6)

where Q_i is observed index flood value *i* (of *n* values), \hat{Q}_i is the modelled value *i*, \bar{Q} is the mean of the observed data, and \tilde{Q} is the mean of the modelled data.

284 Physical legitimacy

285 Following the recent studies of Abrahart et al. (2012b) and Mount et al. (in press), the 286 physical legitimacy of each model was assessed by means of relative, first-order partial 287 derivative sensitivity analysis (see Hamby, 1994 for an overview of sensitivity analysis 288 approaches). Partial derivative sensitivity analysis elucidates the patterns of influence that each model input has on the output (and *vice versa*) across the output range, thus revealing 289 the internal behaviour of the model response function. First order derivatives reveal the 290 291 separate behaviours associated with each model input. When using partial derivatives in 292 model comparison studies, it is necessary to standardise derivative values to rates to avoid the difficulties associated with comparing absolute values derived from different inputs with 293 294 different ranges (Nourani and Fard, 2012). Patterns of relative sensitivity can then be used 295 to directly compare the internal response function behaviour of different models, and 296 legitimacy of these behaviours can then be evaluated according to how well the relative sensitivity patterns conform to the logical, rational and physical expectations of the 297 298 modeller. The relative sensitivity (RS_i) of the output from a model (O) with respect to input (I_i) can be calculated as: 299

300

$$301 \qquad RS_i = \frac{\partial O}{\partial I_i} \cdot \frac{I_i}{O} \tag{7}$$

302

Partial derivatives can be computed for ANNs via the application of a backward chaining partial differentiation rule as outlined in Hashem (1992). Adapted from Hashem's more general rule, for an ANN with sigmoid activation functions (i.e. of standard type, as used in our case study), one hidden layer, *i* input units, *j* hidden units and one output unit 307 (*O*), the partial derivative of a network's output can be calculated with respect to each of its308 inputs as:

309

310
$$\frac{\partial O}{\partial I_i} = \sum_{j=1}^n w_{ij} w_{jO} h_j (1 - h_j) O(1 - O)$$
 (8)

311

where, w_{ij} is the weight from input unit *i* to hidden unit *j*, w_{jO} is the weight from hidden unit *j* to the output unit *O*, h_j is the output of hidden unit *j*, and *O* is the output from the network.

315 One important difference between calculating partial derivatives for multiple input, 316 single output GLMs and ANN models should, however, be noted. When computing partial 317 derivatives of a GLM, there is no need to vary the values of the other inputs to investigate 318 the range of sensitivity responses under different input conditions. This is because GLMs 319 deliver a simple additive response function, such that the relative sensitivity for any one 320 variable will involve only that variable, given that all other parts of the expression will cancel 321 out, during the process of scaling the other variables. Hence, relative sensitivity values for each input to the QMED_{GLM} model (Equation 1) can be computed according to Equations 322 323 (9)–(12). The final relative sensitivities of the QMED_{GLM} model are provided in Equations 324 (13)-(16).

$$326 \qquad \frac{\partial QMED}{\partial AREA} = \frac{0.8568 QMED}{AREA} \tag{9}$$
$$327 \qquad \frac{\partial QMED}{\partial SAAR} = \frac{1864.05 QMED}{SAAR^2} \tag{10}$$

328	$\frac{\partial QMED}{\partial FARL} = \frac{3.3662QMED}{FARL}$	(11)
329	$\frac{\partial QMED}{\partial BFIHOST} = -6.5385QMED.BFIHOST$	(12)
330	$RS_{AREA} = 0.8568$	(13)
331	<i>RS_{SAAR}</i> = 1864.05 / <i>SAAR</i>	(14)
332	$RS_{FARL} = 3.3662$	(15)
333	RS _{BFIHOST} = -6.5385 BFIHOST ²	(16)

The same is not true for ANNs, which are not constrained to produce simple, additive response functions. When computing partial derivatives for an ANN it is therefore necessary to isolate the pattern of relative sensitivity of each input variable in turn by holding the other inputs at fixed values so that the patterns of sensitivity associated with each variable can be interpreted within the context of the other variable states. To this end we adopt a simple three-step methodology.

341

342 Step 1: Compute 25th percentile, median and 75th percentile values for each input variable in
343 the data set.

Step 2: Holding all other variables at either 25th percentile, median or 75th percentile, vary
each input variable in turn from across the range of observed values.

346 Step 3: Plot results and interpret the resultant graphs.

347

Thus, physically speaking, if variable states in our study are held at the 25th percentile (or the 75th percentile in the case of the inverse *BFIHOST* measure), the resultant scenario under test is representative of relatively small, dry catchments with high permeability and high flood attenuation: i.e. low catchment *QMED* potential. Conversely, when variables states are held at the 75th percentile (with *BFIHOST* at the 25th percentile), the resultant scenario under test will be representative of relatively large, wet catchments with low permeability and low attenuation: i.e. high catchment *QMED* potential.

355

356 **RESULTS**

357 Independence

358 Figure 2 and Table 3 present an overview of the data showing the relationships that exist 359 between each of the five variables. AREA is not correlated with any of the other three parameters (correlation coefficient ranging from -0.07 to -0.02). There is a negative 360 correlation between SAAR and BFIHOST (correlation coefficient of -0.42) and a similar 361 362 strength negative relationship between SAAR and FARL (correlation coefficient of -0.39). The 363 only positive correlation is that between BFIHOST and FARL (correlation coefficient of 0.11). 364 These weak relationships indicate a reasonable degree of linear independence between the four variables. The strength of the linear relationship between each of the parameters and 365 QMED ranges from a correlation coefficient score of 0.76 for AREA to -0.07 for FARL. The 366 367 strong linear relationship between QMED and AREA, contrasts with the relative sensitivity 368 scores presented later in this paper for the multiple linear regression model, and in so doing 369 emphasises the additional insights provided by sensitivity analysis over basic statistical 370 measures.

371

372

373 Model skill

Figures 3– 5 present scatter diagrams of observed versus modelled index flood values for the GLM, ANN_A and ANN_B models. The full dataset is depicted in each scatter plot. Figures 3 and 4 reveal comparable amounts of predictive skill for the GLM and ANN_A model. Both plots, indeed, appear to show a reasonable degree of model performance at lower levels, but typically under-estimate the higher magnitude flood events. In contrast the ANN_B model appears to perform well across the range of flood event magnitudes and seems very close to correctly modelling the two largest flood events.

381 Although Figures 3, 4, and 5 provide an interpretive view of the accuracy of the three 382 models, Table 4 provides a more objective, numerical contrast by providing comparative performance statistics for each of the models. It shows that while the ANN_B model is 383 undoubtedly the most accurate overall according to the RMSE and R² measures, the GLM is 384 385 more accurate at modelling low flood indices. Although there appears to be a significant 386 difference between the MSRE statistics of the GLM and the ANN_A model (0.19 and 16.12, respectively) these results need to be treated with caution. A very basic model, that simply 387 predicts the index flood for every catchment as $1 \text{ m}^3 \text{ s}^{-1}$, results in a MSRE statistic of 0.93 – 388 better than both the ANN models and not too dissimilar from the GLM. One would not 389 390 seriously contemplate using such a simple model as a prediction of the index flood in an 391 ungauged catchment so it brings into question the suitability of the MSRE as an appropriate 392 measure of performance. It indicates that a model needs to make only a handful of errors at 393 lower levels (which may not be too far from the observed values) to result in a poor MSRE 394 result. This emphasises the importance of using multiple evaluation criteria and 395 understanding the limitations of individual error measures.

Although the scatter diagrams show reasonably similar performance at lower levels,
one or two over/under predictions have skewed the results. A more appropriate measure of

398 performance at lower levels is perhaps the MSLE used by Pokhrel et al. (2012), the results of 399 which are also presented in Table 4. In this case, although the GLM outperforms the ANN_A and ANN_B models, the results are not too dissimilar. For the simple model (producing 1 $m^3 s^{-1}$ 400 for each case) the MSLE is calculated as 15.36 – significantly higher than the more complex 401 402 models. Given that the ANN_B performs reasonably well for low QMED values and better 403 than the GLM at large QMED values where prediction is normally more problematic, the goodness-of-fit statistics suggest that ANN_B could be considered a reasonable alternative to 404 405 GLM.

406

407 SENSITIVITY ANALYSIS AND PHYSICAL INTERPRETATION OF MODELS

408 GLM

409 Relative sensitivity plots for the GLM are provided in Figure 6 are calculated using Equations 410 (13)–(16). AREA and FARL are both used as simple scaling variables in the model such that 411 the index flood magnitude increases proportionally for larger catchments with lower flood 412 attenuation. The model behaves in a manner that larger catchments produce consistently larger floods, but the overall significance of this behaviour is relatively small. In a simplistic, 413 414 conceptual sense, this is physically legitimate behaviour and one would expect the 415 catchment area to act as a proportionally consistent driver of flood magnitude with a ratio 416 close to unity, as a larger catchment will have proportionally greater hydrological inputs. 417 Importantly, FARL as a driver, is shown to be around four times more important than AREA; 418 a pattern that perhaps highlights the overriding importance of in-channel buffering of flood 419 peaks by lakes and reservoirs in the model.

420 *SAAR* and *BFIHOST* function as more complex drivers of *QMED* and their relative 421 sensitivities vary considerably. Indeed, in certain data ranges each has the potential to

become the most influential driver of index flood magnitude. However, their specific 422 423 patterns of relative sensitivity prove difficult to legitimise in simplified, physical terms. The proportionally greater sensitivity of index flood magnitude to increases in wetness in low 424 425 rainfall catchments, as opposed to ones possessing high rainfall, does not correspond well 426 with broad hydrological notions. The expectation would be to find low antecedent moisture 427 in low rainfall catchments to result in enhanced infiltration, reduced propensity for Hortonian overland flow and correspondingly lower index flood sensitivity compared to 428 429 higher rainfall catchments. This suggests that there is a substantive runoff buffering 430 mechanism in wet catchments that is not present in dry ones. Whilst one may postulate that 431 factors such as different vegetation types in dry and wet catchments may buffer flood 432 responses differently, it is difficult to envisage their impact being sufficient to produce the 433 magnitude of difference observed in the relative sensitivity plot. Moreover, the pattern 434 appears counter to notions of antecedent moisture which would be expected to be lower in 435 dry catchments and, therefore, would act to proportionally reduce catchment runoff and 436 index flood magnitude.

Similarly, the sensitivity of the index flood to catchment permeability is counter to basic physical principles with index floods seen to be an order of magnitude more sensitive to a unit change in permeability in a highly permeable catchment when compared with the same proportional change in an impermeable one. Whilst the overall negative relative sensitivity of *QMED* to *BFIHOST* is conceptually legitimate, the specific pattern is difficult to legitimise physically as is the magnitude of the relative sensitivity observed relative to that of the other variables. The sensitivity analysis thus indicates only partial physical legitimacy of the GLM, with the pattern of sensitivity of *QMED* to *SAAR* and *BFIHOST* being particularly difficult to rationalise.

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449 ANN_A

450 Relative sensitivity plots for the ANN_A model are provided in Figure 7. Importantly, none of 451 the plots exhibit the extreme, localised sensitivity variability that one would expect from an 452 over-fitted model (see ANN_B below), which in the context of the model skill statistics 453 reported above, suggests ANN_A offers a reasonable solution. ANN_A is characterised by 454 generally lower relative sensitivity values in comparison to those observed for the GLM, 455 coupled with enhanced complexity in the sensitivity responses across each variable's data 456 range, the form of which is strongly influenced by the values of the other variables.

457 The relatively high sensitivity of QMED to AREA highlights the central importance of 458 catchment size as a determinant of index flood magnitude in this model. This pattern of 459 behaviour is an approximate counterpart of the GLM plot. Relative sensitivity remains 460 roughly consistent at a value close to 1 and AREA is seen to act as a scaling variable in a 461 physically-legitimate manner. However, the same degree of legitimacy is not observed in either the low or high *QMED* potential plots. Here opposing trends in the relative sensitivity 462 are observed. When all other inputs are set to high QMED potential, proportional changes in 463 464 catchment area of small catchments is seen to have almost 10 times the impact on QMED 465 than the same proportional change in large catchments. The pattern reverses when inputs 466 are set to low QMED potential. This model behaviour is very difficult to legitimise in physical 467 terms.

Low values associated with BFIHOST highlight the general insensitivity of QMED 468 469 to catchment permeability in this model. As expected, BFIHOST has a generally negative 470 influence on QMED such that as permeability increases, QMED reduces. A general increase in QMED's sensitivity to BFIHOST is observed as the other inputs are set to increasing levels 471 472 of QMED potential. This indicates an increased importance of permeability as a constraint 473 on index flood magnitude in catchments with high potential for generating large index floods. However, the very low magnitude of the sensitivities observed makes it difficult to 474 475 draw any clear conclusions about the physical legitimacy of the patterns observed beyond 476 the fact that BFIHOST is clearly not a particularly important driver of QMED.

In contrast to the GLM, FARL acts as a relatively modest driver of QMED, 477 indicating that the ANN_A model is less heavily influenced by in-channel controls of peak 478 479 discharge magnitude than the GLM. In simplistic physical terms, one would expect a 480 reduction in flood attenuation to drive a proportional increase in QMED, and the positive 481 relative sensitivity plots confirm this basic assumption. However, the precise form of the 482 sensitivity relationship between QMED and FARL is more difficult to legitimise. The GLM represents the relationship as one of simple scaling and this same basic pattern exists for 483 484 low and median QMED potential plots across medium to high FARL data ranges (i.e. medium 485 to low levels of attenuation) where relative sensitivity is consistently about 0.5. However, at 486 lower FARL data ranges the proportional response of QMED to change in FARL reduces substantially to 0.1. When other inputs are set to high QMED potential, the decreasing trend 487 488 is consistent across all FARL ranges. This is less easily rationalised and is most likely attributable to the scarcity of catchments with low FARL values in the data resulting in a lack 489 490 of data constraint on the form of the ANN model covering this data range, irrespective of 491 the values of the other inputs.

492 The pattern of sensitivities observed for SAAR can only be partially legitimised in 493 generalised physical terms. At a very simplistic level, the scaling behaviour of SAAR observed 494 in the low QMED potential plot is perhaps reasonable given that proportionally wetter 495 catchments should indeed result in proportionally greater floods. However, the patterns 496 observed in the median and high QMED potential plots possess elements that are both 497 physically rational and irrational. The increasing sensitivity to SAAR at low and mid data ranges could feasibly be explained in terms of antecedent moisture. Indeed, the on-average 498 499 lower antecedent moisture in dry catchments could be expected to result in a smaller 500 proportion of the rainfall contributing to runoff; leading to reduced hydrograph flashiness and proportionally lower QMED sensitivity to SAAR in dryer catchments. Similarly, the 501 502 decline in sensitivity in the upper data ranges could be argued to be due to the fact that the 503 catchment is already so wet that any additional rainfall makes relatively little difference to 504 the index flood. However, this explanation ignores the role of overland, Hortonian flow in 505 saturated, wet catchments which one would expect to drive an increase in the relative 506 sensitivity in the upper data ranges. Finally, the negative relative sensitivity observed in the 507 extreme upper ranges of the high QMED potential plot is physically-irrational as it suggests 508 that proportionally increasing the catchment wetness will reduce the proportional response 509 in *QMED*; in extreme cases even resulting in a reduction in *QMED*.

510 For each of the model inputs the behaviour of the ANN_A model is seen to be 511 particularly influenced by the states of the input variables. When these are set to their 512 median values (i.e. indicative of median *QMED* potential), the majority of the relative 513 sensitivity plots indicate that the response function produces a model behaviour that can be 514 physically-legitimised. However, this legitimacy is less certain when other variables are set 515 at their 25th percentile values (i.e. indicative of low *QMED* potential) and completely breaks down when set at their 75th percentile value (i.e. indicative of high *QMED* potential). Indeed, under the latter condition, *AREA*, *FARL* and *SAAR* drive *QMED* in a manner that is particularly difficult to explain in hydrological terms. Crucially then, a link can be made between the lack of physical legitimacy in the model's behaviour in the upper and lower quartiles of the solution space and a lack of coincident data points which exist there to constrain the form of the ANN model.

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- 523
- 524 ANN_B

Relative sensitivity plots for the ANN_B model are provided in Figure 8. This ANN model is 525 526 intentionally over-fitted and the impact of this over-fitting is clearly seen in the relative 527 sensitivity plots. The degree of local variability in relative sensitivity is highly exaggerated 528 when compared to ANN_A with variables switching between both negative and positive 529 responses in QMED at different data ranges. QMED responds to AREA and SAAR (the most 530 influential drivers in the model) in an irrational manner with high magnitude, localised 531 variation in relative sensitivity being particularly characteristic of the patterns observed. The relative sensitivity plots of QMED to AREA and SAAR are characterised by complex 532 533 polynomial forms with no consistent trends in the relationship. The patterns observed are 534 indicative of data over-fitting and lack any physical legitimacy.

535 Relative sensitivity of *QMED* to *FARL* behaves in a more constrained manner 536 than *AREA* or *SAAR*, ranging from +0.8 to -0.3 indicating the relative lack of sensitivity to this 537 variable in ANN_B. However, the sensitivity plots for low and median *QMED* potential show 538 both positive and negative responses at different data ranges. Indeed, these plots suggest 539 that in certain data ranges, a proportional decrease in flood attenuation will see a 540 proportional reduction in flood magnitude: a result that lacks physical legitimacy. The high 541 *QMED* potential plot is very similar to that of ANN_A

Relative sensitivity of *BFIHOST* to *QMED* is very muted with this variable being an almost irrelevant driver of index flood magnitude when other variables are set to low and median *QMED* potential. Localised complexity in the relative sensitivity is observed, particularly across low *BFIHOST* values where low and median *QMED* potential plots switch between positive and negative relative sensitivity values in a physically-irrational manner. The high *QMED* potential plot is perhaps more rational as it displays a flatter, negative response which indicates a negative scaling behaviour.

In contrast with ANN_A, local variation in relative sensitivity for *AREA* and *SAAR* becomes highly exaggerated when other variables are held at their low *QMED* potential values. This again highlights difficulties of fitting a 'bottom heavy' physically-legitimate ANN model, through upper regions of a solution space that lack sufficient coincident higher magnitude data points to constrain the form of the model.

554

555 Physical legitimacy

556 The broad physical legitimacy of the different model sensitivity plots are compared in Table 557 5. It is clear that none of the models behave in a manner that can be physically rationalised for all input variables. The GLM displays a basic level of physical legitimacy in the behaviour 558 559 of AREA and FARL but this is lacking for SAAR and BFIHOST drivers. ANN_A displays varying 560 degrees of physical legitimacy in the sensitivity between QMED and each of the input 561 variables, with the least rational responses occurring when other variables are set to the 562 high QMED potential values. However, in all cases, when other variables are set to their 563 median values, the relative sensitivities of the ANN are physically legitimate at least in part.

Indeed, in this sense ANN_A arguably performs better than its GLM counterpart albeit 564 565 delivering slightly less favourable goodness-of-fit. ANN_B is over-fitted and the patterns 566 observed in its relative sensitivity plots cannot be legitimised in a physical sense. However, this lack of model legitimacy is in contrast to the goodness-of-fit statistics which indicate 567 568 ANN_B to be the best model. Thus, developing techniques that can deliver a clear physical or 569 mechanistic interpretation of input relative sensitivity analysis patterns in ANN modelling 570 scenarios represents an important consideration for future research. Indeed, the presented 571 results serve as a clear demonstration of the dangers associated with evaluating models on 572 the basis of statistical performance validation approaches alone.

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574 SUMMARY AND CONCLUSIONS

575 This paper has addressed the difficult question of how to make meaningful comparisons 576 between artificial neural network-based hydrological models and alternative modelling 577 approaches. Comparisons which are based solely on goodness-of-fit metrics (i.e. the 578 standard black-box approach presented in much of the literature) are very limited because 579 they only consider model performance and not the means by which the performance is obtained. The commonly encountered limitation of metric equifinality, in which metric 580 581 scores for the models being compared are insufficiently different to enable conclusive 582 differentiation of the best or preferred model, is evident in our results. Our example of 583 median flood modelling provides a clear demonstration of this with the fit scores obtained 584 by the ANN and GLM models delivering inconclusive evidence about relative overall model 585 performance.

586 However, the limitations of goodness-of-fit metrics are arguably more fundamental 587 if there is a requirement to compare the transferability of each model from one hydrological

588 context to another. In such cases, the physical legitimacy of each model must also be 589 evaluated and compared in a direct manner. Models used in ungauged catchment 590 prediction are a good example of those that must ultimately be transferred, and that This study has presented a 591 therefore require evaluation of their physical legitimacy. 592 consistent means by which the physical legitimacy of ANN models can be evaluated and 593 compared with alternative modelling approaches. The application of relative sensitivity analysis in our median flood modelling example has enabled the physical legitimacy of two 594 595 ANN-based models to be compared directly with the GLM counterpart used as standard in 596 the UK. Tables 4 and 5 provide clear evidence that a general ANN modelling approach can deliver models as good as the GLM approach currently used in the UK Flood Estimation 597 598 Handbook, both in terms of their performance and their legitimacy. Whilst the paper does 599 not purport to be a competition between ANNs and GLMs, in this isolated case the evidence 600 does lend some support to the view that ANN-based models may have some advantages 601 over their GLM counterparts. However, one can only build good physically-legitimate ANN 602 models if ample data of sufficient quality exist, and if the model development process is 603 sound. It is also evident from this evaluation that ANN solutions can only deliver physical 604 legitimacy if issues such as overfitting are avoided.

To conclude it is clear that comparing ANN models to alternative approaches on the basis of goodness-of-fit is insufficient, and that sensitivity analysis offers an important means by which the physical legitimacy of ANN models can be compared with that of counterpart models. Indeed, hydrological modellers using ANNs can and should be striving to evaluate the physical legitimacy of their models as well as their performance. By applying sensitivity analysis to ANN models a sense of trust is introduced that goes part of the way to addressing one of the key issues in the international ANN river forecasting research agenda 612 of Abrahart et al. (2012a), specifically the need for advanced diagnostic techniques that can 613 help counter criticisms of the black-box nature of such models (e.g. Babovic, 2005). It is, 614 therefore, surprising that it remains almost entirely absent from ANN studies and highlights 615 the importance of a broader research agenda to develop robust, computational sensitivity 616 analysis methods across the range of data-driven techniques currently being used in 617 hydrological modelling. Such an agenda should include additional investigations that more fully explore the impact of different architectural structures in ANN models especially the 618 619 potential bearing that internal complexity might have on the relative sensitivity of solutions 620 to particular types of hydrological modelling problem.

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849 TABLE CAPTIONS

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	Median	Minimum	Maximum	25 th Percentile	75 th Percentile
AREA (km ²)	148.70	1.63	4586.97	68.00	327.81
BFIHOST	0.47	0.20	0.97	0.40	0.57
FARL	0.99	0.65	1.00	0.96	1.00
SAAR (mm)	1096	558	2848	830	1375
QMED	43.54	0.14	992.85	12.92	117.71

Table 2. Network weights and biases. Input neurons I1 - I4 (AREA, BFIHOST, FARL, SAAR,

respectively); Hidden neurons H1 – H9; Output neuron O (QMED)

ANNa

		Weight			Weight			Weight			Weight			Weight
11	H1	2.112	12	H1	1.287	13	H1	-1.858	14	H1	-4.078	H1	0	-2.004
11	H2	-0.211	12	H2	-0.392	13	H2	-1.591	14	H2	-0.154	H2	0	-0.797
11	H3	2.907	12	H3	-6.502	13	H3	2.196	14	H3	4.048	H3	0	4.901
11	H4	-1.170	12	H4	2.792	13	H4	-0.347	14	H4	-3.403	H4	0	-1.904
11	H5	0.245	12	H5	-0.337	13	H5	-2.473	14	H5	0.521	H5	0	-1.001
11	H6	0.009	12	H6	-1.236	13	H6	-1.627	14	H6	0.087	H6	0	-0.533
11	H7	-13.412	12	H7	-4.484	13	H7	1.478	14	H7	2.806	H7	0	-7.586
11	H8	-1.236	12	H8	0.008	13	H8	-0.782	14	H8	-0.284	H8	0	-0.921
11	H9	-6.588	12	H9	-2.458	13	H9	0.998	14	H9	1.157	H9	0	-3.972

ANNb

		Weight			Weight			Weight			Weight			Weight
11	H1	-1.877	12	H1	20.295	13	H1	0.185	14	H1	-14.475	H1	0	-2.575
11	H2	-16.987	12	H2	-3.354	13	H2	1.693	14	H2	2.498	H2	0	-13.556
11	H3	-3.798	12	H3	-0.008	13	H3	-2.085	14	H3	-7.115	H3	0	4.112
11	H4	5.559	12	H4	-0.845	13	H4	1.849	14	H4	-18.273	H4	0	-4.311
11	H5	-2.996	12	H5	4.687	13	H5	-6.742	14	H5	6.914	H5	0	-1.337
11	H6	8.318	12	H6	-8.377	13	H6	2.917	14	H6	8.574	H6	0	4.750
11	H7	8.324	12	H7	-3.983	13	H7	-3.674	14	H7	10.392	H7	0	3.969
11	H8	11.702	12	H8	-19.838	13	H8	-2.518	14	H8	16.069	H8	0	-2.763
11	H9	1.210	12	H9	-3.488	13	H9	-3.777	14	H9	6.853	H9	0	-3.085

Biases

Neuron	Bias ANNa	Bias ANNb
H1	-0.596	-0.708
H2	-0.175	-1.927
H3	-3.240	0.049
H4	-0.315	-1.594
H5	0.413	2.982
H6	-0.098	-7.794
H7	-1.459	-0.996
H8	-0.508	0.627
H9	-0.720	0.278
0	0.282	1.707

875 Table 3. Correlation matrix for model variables

	AREA	BFIHOST	FARL	SAAR	QMED
AREA	1.00	-0.02	-0.07	-0.05	0.76
BFIHOST		1.00	0.11	-0.42	-0.27
FARL			1.00	-0.39	-0.07
SAAR				1.00	0.24

Table 4. Numerical accuracy of different models under test

	GLM	ANN _A	ANN _B	
RMSE (m ³ s ⁻¹)	43.09	47.49	33.18	
R ²	0.89	0.88	0.94	
MSRE	0.19	16.12	1.91	
MSLE	0.13	0.51	0.33	

885Table 5. Physical legitimacy of GLM and ANN models

Input Variable	QMED potential of other	Does the pattern of sensitivity response conform to conceptual notions of physically-rationality?				
	catchment variables	GLM	ANNA	ANN _B		
	Low		No	No		
AREA	Median	} Yes	Yes	No		
	High		No	No		
	Low	1	Yes	No		
SAAR	Median	} No	In Part	No		
	High		No	No		
	Low		In Part	No		
FARL	Median	} Yes	In Part	No		
	High		No	No		
	Low		No	No		
BFIHOST	Median	} No	In Part	No		
	High		In Part	In Part		





Figure 2. Scatter plot matrix of model variable with linear regression lines fitted







Figure 3. GLM versus QMED









Figure 5. ANN_B model versus QMED



Figure 6. Relative sensitivity of QMED to model inputs: GLM



Figure 7. Relative sensitivity of QMED to model inputs: ANNA



Figure 8. Relative sensitivity of QMED to model inputs: ANN_B