

1 **Resilience of SAR11 bacteria to rapid acidification in the high latitude open**  
2 **ocean**

3 Manuela Hartmann<sup>1</sup>, Polly G Hill<sup>1</sup>, Eithne Tynan<sup>2</sup>, Eric P Achterberg<sup>2,3</sup>, Raymond J G  
4 Leakey<sup>4</sup> and Mikhail V. Zubkov<sup>1\*</sup>

5 **Author affiliations:**

6 <sup>1</sup>National Oceanography Centre, Southampton, European Way, Southampton SO14  
7 3ZH, UK

8 <sup>2</sup>Ocean and Earth Science, National Oceanography Centre Southampton, University  
9 of Southampton, Southampton, SO14 3ZH, United Kingdom

10 <sup>3</sup>GEOMAR Helmholtz Centre for Ocean Research, 24148 Kiel, Germany

11 <sup>4</sup>Scottish Association for Marine Science, Scottish Marine Institute, Oban, Argyll  
12 PA37 1QA, UK

13 **Authors' contribution:** M.H. and P.G.H. contributed equally to this work.

14

15 **Running title:** Effect of acidification on SAR11

---

\* **Correspondence:**

Mikhail V. Zubkov, Ocean Biogeochemistry & Ecosystems Research Group, National  
Oceanography Centre, Southampton, European Way, Southampton, SO14 3ZH,  
United Kingdom, Tel: +44 (0)23 8059 6335, Fax: +44 (0)23 8059 6247,  
[mvz@noc.soton.ac.uk](mailto:mvz@noc.soton.ac.uk)

## 16 **Abstract**

17 Ubiquitous SAR11 *Alphaproteobacteria* numerically dominate marine planktonic  
18 communities. Because they are excruciatingly difficult to cultivate, there is  
19 comparatively little known about their physiology and metabolic responses to long-  
20 and short- term environmental changes. As surface oceans take up anthropogenic,  
21 atmospheric CO<sub>2</sub>, the consequential process of ocean acidification could affect the  
22 global biogeochemical significance of SAR11. Shipping accidents or inadvertent  
23 release of chemicals from industrial plants can have strong short-term local effects  
24 on oceanic SAR11. This study investigated the effect of 2.5 fold acidification of  
25 seawater on the metabolism of SAR11 and other heterotrophic bacterioplankton  
26 along a natural temperature gradient crossing the North Atlantic Ocean, Norwegian  
27 and Greenland Seas. Uptake rates of the amino acid leucine by SAR11 cells as well  
28 as other bacterioplankton remained similar to controls despite an instant ~50%  
29 increase in leucine bioavailability upon acidification. This high physiological  
30 resilience to acidification even without acclimation, suggests that open ocean  
31 dominant bacterioplankton are able to cope even with sudden and therefore more  
32 likely with long-term acidification effects.

33

## 34 **Key words**

35 CARD-FISH / flow cytometric cell sorting / isotopic tracer labelling / pCO<sub>2</sub>  
36 perturbation

37

## 38 **Introduction**

39 Heterotrophic bacteria, among them the ubiquitous alphaproteobacterial SAR11  
40 clade (hereafter SAR11, (Morris *et al.*, 2002)), form an abundant and important  
41 component of pelagic marine microbial communities, dominate the remineralisation  
42 of phytoplankton-derived organic matter (Martin *et al.*, 1987, Boyd *et al.*, 1999) and  
43 constitute a nutrient-rich food source for the smallest eukaryotes (Sherr & Sherr,  
44 2002, Hartmann *et al.*, 2013). SAR11 are present in surface waters throughout the  
45 world's oceans (Brown *et al.*, 2012) but experience unique environmental conditions  
46 in the polar oceans, including a highly seasonal radiation regime, low water  
47 temperature and perennial or continuous sea ice-cover.

48 Human activities were shown to have a large, cumulative impact on the majority of  
49 the global ocean (~66%), ocean acidification and sea surface temperature being the  
50 major factors followed by ocean-based pollution, shipping and destructive fishing  
51 (Halpern *et al.*, 2015). Bacteria of polar surface waters may be subject to more rapid  
52 and intense ocean acidification due to the higher solubility of carbon dioxide (CO<sub>2</sub>) at  
53 low water temperatures leading to reduced pH values and a decreased buffering  
54 capacity (Sabine *et al.*, 2004, Fabry *et al.*, 2009). Apart from the slow process of  
55 ocean acidification, shipping activities can locally decrease pH due to NO<sub>x</sub> and SO<sub>x</sub>  
56 emissions (Hassellöv *et al.*, 2013) or by accidental release of acidifying agents into  
57 the water (Mamaca *et al.*, 2009). Acidification may have indirect effects on the  
58 speciation of trace metals, e.g. increasing the availability of iron and copper because  
59 of their higher solubility at lower pH (Morel *et al.*, 2003). However, culture studies  
60 have shown that despite the higher solubility the bioavailability of iron could  
61 decrease as a result of acidification (Shi *et al.*, 2010). The effect of lowered pH on  
62 inorganic nutrient speciation is less well known. Using theoretical speciation  
63 diagrams Zeebe and Wolf-Gladrow (2001) showed significant differences for

64 phosphate, silicate, iron and ammonia speciation if pH drops from 8.1 to 7.8. Bulk  
65 dissolved organic matter concentrations have been shown to be uninfluenced by  
66 acidification in short-term (MacGilchrist *et al.*, 2014) and long-term experiments (Zark  
67 *et al.*, 2015). However, dissolved organic matter is highly complex and differences  
68 due to acidification could be masked in bulk analyses resulting in changes of  
69 metabolic responses and community compositions.

70 Abundance of the SAR11 clade has been shown to co-vary negatively with  
71 particulate organic carbon but positively with turnover rates of dissolved free amino  
72 acids (Giebel *et al.*, 2011). Transcriptional responses of SAR11 from the North  
73 Pacific subtropical gyre indicate that these organisms acquire nitrogen from  
74 dissolved organic matter in addition to its usage as an energy source (Sharma *et al.*,  
75 2014). Besides organic matter, temperature and salinity have been proposed to drive  
76 distribution and abundance of different SAR11 groups (Brown *et al.*, 2012,  
77 Herlemann *et al.*, 2014).

78 Studies of acidification are currently mainly focused on ocean acidification where  
79 small or large volumes of samples are incubated for comparatively short (days) or  
80 long (up to one year) under increased pCO<sub>2</sub> usually mimicking predicted end of the  
81 century, i.e. 1000ppm, levels of pCO<sub>2</sub> (Riebesell, 2004, Engel *et al.*, 2005, Motegi *et*  
82 *al.*, 2013, Gattuso *et al.*, 2014, Richier *et al.*, 2014, Zark *et al.*, 2015). Alas, microbial  
83 communities can react rapidly (within hours) to elevated CO<sub>2</sub> levels. For example,  
84 the nitrogen fixation rates of *Trichodesmium* can increase by up to 41% within 4-6  
85 hours after exposure to high CO<sub>2</sub> conditions (Hutchins *et al.*, 2009).

86 It is often assumed that the effect of (rapid) acidification on marine bacterioplankton  
87 will be marginal due to the natural variability of surface ocean pH (Joint *et al.*, 2011)

88 and the potential of bacterioplankton to grow at a wide range of pH, e.g. an isolates  
89 of the Roseobacter group exhibit optimal growth rates between pH 7.2 And 7.8, but  
90 stopped growing at pH 8.2 (Giebel *et al.*, 2011). Indeed, in estuarine or upwelling  
91 regions pH can change significantly in a short period of time (up to 1 unit) (Hofmann  
92 *et al.*, 2011). However, measurements of sites in the Pacific Ocean and Antarctica  
93 were considerably more stable, e.g. showing only 0.024-0.096 units variability  
94 (Hofmann *et al.*, 2011). Therefore, they could be more sensitive to abrupt pH  
95 changes.

96 The aim of the present study was to investigate the effect of acidification on protein  
97 synthesis of natural bacterioplankton from ocean waters along a latitudinal gradient  
98 from the North Atlantic to the Arctic. Importantly the study examined, for the first  
99 time, the immediate ( $\leq 2$  h) and therefore direct, group-specific response of SAR11,  
100 to pH perturbation, thereby avoiding the indirect effects resulting from pH-induced  
101 changes to other components of the natural community. In addition, comparison of  
102 bacterioplankton response along the natural temperature gradient offered insights  
103 into the relationship between acidification and temperature; both of which are  
104 predicted to increase in future years. Seawater samples were acidified using a  
105 combined approach of acid and sodium bicarbonate addition in order to maintain  
106 alkalinity levels (Riebesell *et al.*, 2010). Bacterial activity was assessed by measuring  
107 the uptake and turnover of tritiated leucine, the assimilation of which is routinely used  
108 to study bacterial biomass production (Kirchman *et al.*, 1985). Leucine assimilation is  
109 a ubiquitous and obligate metabolic pathway which is most likely to react  
110 immediately to artificial acidification. This approach therefore offered a highly  
111 sensitive method to detect positive or negative effects of acidification on SAR11 and  
112 other heterotrophic bacterioplankton.

113

## 114 **Materials and Methods**

### 115 **Seawater collection and acidification**

116 The study was conducted between the 6<sup>th</sup> and 17<sup>th</sup> June 2012 as part of the UK  
117 Ocean Acidification research programme's Arctic research cruise on the RRS *James*  
118 *Clark Ross* (cruise JR271). Eleven open sea stations were sampled along a  
119 latitudinal transect crossing the North Atlantic Ocean, Norwegian and Greenland  
120 Seas, including ice-covered waters of the Fram Strait (Table S1; Fig. 1). At each  
121 station a 10 L seawater sample was collected from 10-20 m depth using a Niskin  
122 bottle mounted on a stainless steel frame with a Sea-Bird 9/11 plus conductivity,  
123 temperature and density (CTD) profiler. The sample was decanted into an acid  
124 cleaned polycarbonate carboy and subsequently divided into an acidified and a  
125 control treatment.

126 Acidification (i.e. the manipulation of seawater to increase the ambient pCO<sub>2</sub> to a  
127 target level of 1000 ppm, resulting in a pH of 7.79±0.07) was achieved by the  
128 addition of ultra-clean hydrochloric acid and sodium bicarbonate; an approach which  
129 increases CO<sub>2</sub> and decreases pH whilst maintaining a balanced carbonate system  
130 (Riebesell *et al.*, 2010). 50ml samples were removed from control and acidified  
131 bottles at the beginning and end of incubations to measure the resultant dissolved  
132 inorganic carbon (DIC) and total alkalinity (TA), and to verify this remained stable  
133 throughout the incubation. DIC was analysed with an Apollo SciTech DIC analyzer  
134 (AS-C3) using a LI-COR (7000) CO<sub>2</sub> infrared detector. TA was determined using an  
135 open-cell titration (Dickson *et al.*, 2007) with the Apollo SciTech's AS-ALK2 Alkalinity  
136 Titrator. Analyses of Certified Reference Materials (CRMs) (A.G. Dickson, Scripps,

137 batch 117) were used at the beginning, middle and end of each analytical run to  
138 ensure accuracy of the measurements. Precision was taken as the standard  
139 deviation of repeated analysis of CRMs, which in this case was  $\pm 1.5 \mu\text{mol kg}^{-1}$  for TA  
140 and  $\pm 3 \mu\text{mol kg}^{-1}$  for DIC. The remaining variables of the carbonate system, including  
141  $\text{pCO}_2$ , were calculated using CO2SYS (MATLAB version 1.1) (Lewis & Wallace,  
142 1998, van Heuven *et al.*, 2011) using the constants of Mehrbach *et al.* (1973) refitted  
143 by Dickson and Millero (1987).

144 Additional triplicate samples (1.6 mL) were taken at the beginning of incubations  
145 from the control and acidified seawater and fixed with 20% paraformaldehyde (PFA,  
146 1% final concentration) for determination of bacterial cell abundance by flow  
147 cytometry (Fig. 2a and Fig. S1), as outlined below.

#### 148 **Microbial uptake of leucine by radioisotope dilution bioassay**

149 The effect of acidification on leucine uptake, an indicator of bacterial protein  
150 synthesis, was determined using the radioisotope dilution bioassay approach  
151 developed for freshwater systems (Wright & Hobbie, 1966) and subsequently  
152 employed extensively in the marine environment (Hill *et al.*, 2011). [4,5- $^3\text{H}$ ] L-Leucine  
153 ( $140 \text{ Ci}\cdot\text{mmol}^{-1}$ , Hartmann Analytic, Germany) was added to the control and acidified  
154 seawater samples at a range of concentrations (0.1, 0.2, 0.4, 0.6, 0.8, 1.0 nM) and  
155 incubated in parallel at *in situ* temperature and under low-intensity, indirect artificial  
156 light. Three samples (1.6 mL) were incubated for each concentration and fixed  
157 sequentially after 10, 20 and 30 minutes by the addition of 20% (w/v) PFA (1% final  
158 concentration) for 1 hour at room temperature (RT) in the dark.

159 Each fixed bioassay sample was filtered through a  $0.2 \mu\text{m}$  pore-size polycarbonate  
160 filter (Nuclepore, Whatman, UK) to collect cells, washed twice with 3 mL ultrapure

161 water (Purelab, Elga Process Water, Marlow, UK) and the filters placed into plastic  
162 vials to which 3 mL Goldstar scintillation cocktail (Meridian, UK) were added. The  
163 radioactivity retained by particulate material was then measured on ship by liquid  
164 scintillation counting (Tri-Carb 3100, Perkin Elmer, UK) to assess microbial leucine  
165 uptake.  $^3\text{H}$ -leucine uptake rates were determined for each concentration of  $^3\text{H}$ -  
166 leucine addition from the regression of uptake versus incubation time (Fig. S2a). The  
167 ambient concentration of the leucine pool and its turnover time were determined from  
168 the x-axis and y-axis intercepts of the regression of leucine uptake rate versus  
169 concentration, respectively (Fig. S2b).

#### 170 **Leucine uptake rates of different bacterial clusters by flow cytometry**

171 Additional control and acidified seawater samples were incubated with 0.4 nM  $^3\text{H}$ -  
172 leucine (final concentration) for 2 hours at ambient temperature and fixed with PFA  
173 (1% final concentration) for subsequent flow cytometric sorting of different bacterial  
174 clusters for stations 6 to 15 (Table S1). Samples were stained with SYBR Green I  
175 nucleic acid stain (Sigma-Aldrich, UK) in the presence of tri-potassium citrate  
176 (Sigma-Aldrich, UK) at RT (Marie *et al.*, 1997). Stained samples were analysed, and  
177 bacterial cells enumerated and sorted by flow cytometry (FACSort, Becton  
178 Dickinson, UK) according to the light-scattering properties ( $90^\circ$  or side light scatter)  
179 and relative concentration of SYBR Green I stain per particle (green fluorescence;  
180 FL1,  $530 \pm 30$  nm). The total bacterioplankton population could be divided into two  
181 distinct clusters differentiated by their nucleic acid content: low nucleic acid and high  
182 nucleic acid containing bacteria (hereafter LNA and HNA respectively, Fig. S1). The  
183 HNA and LNA cells were enumerated by the addition of an internal standard  
184 comprising a mixture of 0.5  $\mu\text{m}$  and 1.0  $\mu\text{m}$  multifluorescent latex beads  
185 (Polysciences, USA) of known concentration (Zubkov & Burkill, 2006).



186 Radioisotopically labelled HNA and LNA cells were sorted at low flow rate ( $\sim 12 \mu\text{L min}^{-1}$ ) for 1, 2, 3 and 4 minutes, filtered directly onto  $0.2 \mu\text{m}$  pore-size polycarbonate  
187 filters (Nuclepore, Whatman, UK), washed three times with 3 mL ultrapure water and  
188 transferred into plastic vials to which 8 mL of Goldstar scintillation cocktail were  
189 added. Radioactivity within the sorted cells was assayed using an ultra-low level  
190 liquid scintillation counter (1220 Quantulus, Wallac).  
191

## 192 **Identification of dominant bacterial phyla and clades by CARD-FISH**

193 CAlysed Reporter Deposition Fluorescence *In Situ* Hybridisation (CARD-FISH) on  
194 flow cytometrically sorted LNA and HNA cells was performed to determine the  
195 taxonomic composition of the bacterial populations at all stations. Before sorting for  
196 CARD-FISH, the flow cytometer was extensively cleaned with Decon (5%, Fisher,  
197 UK), commercially available thin bleach and ultra-pure water (Millipore, UK).  
198 Subsequently, a new sheath fluid filter with a pore-size of  $0.1 \mu\text{m}$  was inserted, and  
199 the flow cytometer run for 1 hour at high flow rate ( $\sim 200 \mu\text{L sec}^{-1}$ ). In order to validate  
200 cleanliness of the flow cytometer a 60 ml water sample was taken off the sorting line,  
201 filtered onto a  $0.2 \mu\text{m}$  pore-sized polycarbonate filter (Nuclepore, Whatman, UK),  
202 dried and counterstained with Vectashield<sup>®</sup> DAPI (4',6-diamidino-2-phenylindole,  
203 Vector Laboratories Ltd., UK) for 10 min at RT. The filtered volume of 60 ml equalled  
204 approximately the volume filtered during cell sorts. At least two transects of the filter  
205 were inspected to detect potential contaminants using an epifluorescence  
206 microscope (Axioscope, Zeiss, Germany) equipped with a LED light source pE-300  
207 (CoolLED, UK). If more than 20 cells were counted per transect, the cleaning  
208 process was repeated. The cleanliness of the instrument was checked repeatedly  
209 during the cell sorting for CARD-FISH.

210 Flow cytometric cell sorting was carried out as described above and at least 50 000-  
211 100 000 cells were collected onto each filter. At four stations samples had to be pre-  
212 concentrated in order to achieve these cell numbers. Therefore, 60 ml of PFA-fixed  
213 (1% v/v) sample was concentrated at a flow rate of 2.5 ml min<sup>-1</sup> onto a 0.1 µm pore-  
214 size Teflon filter (M-Tech Diagnostics, UK) using a syringe pump (KD Scientific,  
215 USA).

216 Overall hybridisation efficiency was determined using the Eub338I-III probe mix  
217 (Amann *et al.*, 1990) targeting Eubacteria. To identify LNA, which have estimated  
218 low ribosomal content due to their small cell size (Kemp *et al.*, 1993, Morris *et al.*,  
219 2002), a mix of six different probes and one unlabelled helper probe targeting SAR11  
220 *Alphaproteobacteria* (Morris *et al.*, 2002, Gomez-Pereira *et al.*, 2013) were used in  
221 order to enhance hybridisation signal. Additionally, a probe targeting the SAR86  
222 *Gammaproteobacteria* core cluster (Zubkov *et al.*, 2001) was applied. The majority of  
223 HNA were identified by probes targeting the *Bacteroidetes* (Manz *et al.*, 1996),  
224 *Alphaproteobacteria* (Neef, 1997), and *Gammaproteobacteria* (Manz *et al.*, 1992)  
225 phyla. Information on each probe used, including target bacterial groups, nucleic acid  
226 sequence and percentage of formamide used for hybridisation, can be found in the  
227 supplementary material (Table S2).

228 CARD-FISH was performed as described in Gomez-Pereira *et al.* (2013). Briefly,  
229 cells were permeabilised with lysozyme (10 mg·ml<sup>-1</sup>, Sigma Aldrich, UK) and  
230 achromopeptidase (60 U·ml<sup>-1</sup>, Sigma Aldrich, UK) at 37°C for 1 h and 30 min,  
231 respectively. Filters were hybridised overnight (SAR11 probe mix) or for 3 h  
232 (remaining probes) at 46°C at varying formamide concentrations (Table S2).  
233 Hybridisation buffer and probes (50 ng·ml<sup>-1</sup>) were mixed at a 300:1 ratio. Positive  
234 hybridisation of the probe was detected using fluorescently labelled tyramide

235 Alexa488 (Life Technologies, Carlsbad, CA, USA) at a ratio of 1:1000 in amplification  
236 buffer. Finally cells were counterstained with DAPI (4',6-diamidino-2-phenylindole),  
237 mounted in Vectashield<sup>®</sup> antifading reagent and enumerated microscopically.

## 238 **Data analysis**

239 All statistical analyses were performed using SigmaPlot. For normally distributed  
240 data Student's t-tests were used to compare treatments, while for non-normally  
241 distributed data Wilcoxon signed rank tests were applied. P values below 0.05 were  
242 considered significant. Errors were calculated according to standard error  
243 propagation procedures.

244

## 245 **Results and Discussion**

### 246 **Bacterial abundance and identification of dominant populations**

247 Bacterial abundances ranged from  $2 \cdot 10^5$  to  $2 \cdot 10^6$  cells  $\text{ml}^{-1}$ . Although LNA  
248 represented occasionally a significant proportion of the bacterial population (up to  
249 45%, station 8), HNA were dominant at all stations (on average  $72 \pm 10\%$ , Fig. 2a).  
250 CARD-FISH analyses on flow cytometrically sorted LNA using a mix of seven  
251 different SAR11 probes to enhance the signal (Gomez-Pereira *et al.*, 2013) revealed  
252 that the LNA were almost exclusively composed of SAR11 *Alphaproteobacteria*  
253 ( $90 \pm 3\%$ ,  $n=10$ , Fig. 2b). Therefore, we will refer to LNA as SAR11 from this point  
254 onward. If single probes are used on LNA, the hybridisation efficiency can be very  
255 low, e.g. only  $18 \pm 8\%$  of the sorted population positively hybridised with the general  
256 eubacterial probe (Eub338I-III, Fig. 2b) and the resulting signal was very weak.  
257 Although slightly higher, perhaps attributable to the use of a probe mix, the here  
258 measured total abundance of SAR11 ( $25 \pm 9\%$ ) is comparable to a study carried out

259 in the Western Arctic Ocean in May-June, where FISH using a single probe revealed  
260 that SAR11 comprised up to  $15\pm 11\%$  of the bacterioplankton community (Malmstrom  
261 *et al.*, 2007). In addition to the SAR11 probe mix, a SAR86 targeted probe was  
262 applied to the sorted LNA cells (SAR86\_1245, Fig. 2b) in order to identify the minute  
263 remaining part of the population. However, SAR86 target cells were detected in only  
264 4 out of 9 samples and then just a minute fraction exhibited positive hybridisation  
265 signals (0.2-1% of the cells). Similar to SAR11, SAR86 *Gammaproteobacteria* have  
266 a small genome size and could hence be expected within the LNA (Dupont *et al.*,  
267 2012). Potentially these values underestimate the presence of SAR86 in the samples  
268 due to the above-mentioned limitations of using a single probe (Fig. 2c).

269 The hybridisation signal for the HNA was very strong and 84-97% of cells were  
270 positively hybridised with probe EUB338 I-III (Fig. 2c). In contrast to LNA, HNA  
271 represented a very heterogeneous mix of bacteria. The selected probes in this study  
272 allowed identification of the majority of the cells ( $61\pm 18\%$ ). *Bacteroidetes* were  
273 present at all studied stations where, with the exception of one of the southernmost  
274 stations (Station 6, 10%), they represented a major fraction of the total sorted HNA  
275 cells (on average  $42\pm 12\%$ ) (Fig. 2c). At one of the northernmost stations (Station 15)  
276 *Bacteroidetes* could not be determined due to a lack of sample material. *Alpha*- and  
277 *Gammaproteobacteria* abundances were more variable and significantly lower (4-  
278 26% and 1-17% respectively) (Fig. 2c).

279 The homogeneity and methodological accessibility of SAR11 as revealed by our  
280 above analyses (Fig. 2b) makes it an ideal candidate to directly and quantitatively  
281 study the effect of acidification on a particular group of heterotrophic bacteria instead  
282 of bulk bacterial matter, where potential signals could be masked by the presence of  
283 other bacterial populations.

284 **Effect of acidification on bioavailable concentrations, microbial uptake rates**  
285 **and turnover time of leucine**

286 Flow cytometric analyses of relevant microbial populations (bacterioplankton, LNA,  
287 HNA and small eukaryotes) exhibited no significant fluctuations in cell abundance in  
288 relation to acidification (Fig. S3). Turnover times of mRNA in bacteria suggest that  
289 they are able to react in a matter of minutes to environmental change (Selinger *et al.*,  
290 2003, Steglich *et al.*, 2010). Hence, the experimental set-up of  $\leq 2$  hours allowed  
291 sufficient time to determine a direct effect of acidification.

292 Acidification did not affect the bulk community leucine uptake rates in offshore  
293 waters of the North Atlantic Ocean or Norwegian and Greenland Seas (t-test,  
294  $p=0.73$ ; Fig. 3a). Effects of acidification on different parts of the bacterial community  
295 could be potentially masked in the bulk community analyses. However, average  
296 cellular leucine uptake rates of SAR11 and HNA separated by flow cytometrical cell  
297 sorting across eight stations (Mann-Whitney,  $p=0.97$  and  $0.66$  respectively, Fig. 4)  
298 were not affected by acidification either. These two bacterial groups are the major  
299 consumers of dissolved leucine in these waters, as shown by comparing the sum of  
300 the uptake rates of SAR11 and HNA versus the total uptake measured (Fig. S4).  
301 Presumably due to their larger size, HNA usually consumed significantly more  
302 leucine  $\text{cell}^{-1} \text{h}^{-1}$  than SAR11 (Mann-Whitney,  $p<0.001$ , Fig. S5) and accounted for  
303  $83 (\pm 16\%)$  of the total microbial leucine uptake, therefore dominating bacterial  
304 production in these waters whilst, owing to their strong latitudinal variation in  
305 abundance (Fig. 2a), SAR11 contributed 3-45% of total bacterial production.

306 In contrast to leucine uptake, acidification did significantly increase leucine  
307 bioavailability (Wilcoxon signed rank,  $p=0.005$ ; Fig. 3b) which, in turn, increased the

308 turnover time (Wilcoxon signed rank,  $p < 0.001$ ; Fig. 3c). Because cell numbers in  
309 the acidified and control sample were comparable (Fig. S3), the increased leucine  
310 bioavailability could not be attributed to organic matter release by cell death due to  
311 the acidification of the sample. It is more likely that pH-induced dissolution of  
312 colloidal material resulted in higher concentrations of bio-available amino acids in the  
313 water (Hansell & Carlson, 2014). On average, concentration of leucine in acidified  
314 samples was 1.8 times higher with up to three times higher concentrations at two  
315 stations. Increase organic matter is often considered to have a negative impact on  
316 members of the SAR11. However, the stable leucine uptake rates suggest that  
317 acidification and the measured release of amino acids does not have a detrimental  
318 effect on this group. Despite the comparatively more stable pH in the surface open  
319 ocean (Hofmann *et al.*, 2011), the results are in accordance with the ability of  
320 bacterioplankton in estuaries, upwelling zones and other aquatic systems, such  
321 freshwater lakes, to cope with very large (order of magnitude) and rapid (within  
322 hours) pH changes. The ubiquitous distribution of SAR11 (Brown *et al.* 2012), already  
323 suggested that they are adapted over a wide pH range and our study confirmed that  
324 they can persist under rapidly changing conditions as well.

325 The above results suggest that SAR11 and other heterotrophic bacterioplankton  
326 encountered in this study are tolerant to pH changes, thereby supporting the  
327 conclusions of Joint *et al.* (2011) who deduced theoretically that microorganisms will  
328 adapt to ocean acidification. However, acidification is just one parameter in the  
329 currently changing ocean, so cumulative effects of for example increases in sea  
330 surface temperature and UV radiation will have to be taken into consideration.  
331 Moreover, other organisms might be not able to adapt to pH change, therefore  
332 community structure and thus biogeochemistry will perhaps change dramatically to

333 which bacteria cannot adapt that easily. Indeed, some bacterial groups, e.g.  
334 Gammaproteobacteria and Flavobacteriaceae, seem to be more susceptible to pH  
335 changes as seen by subtle shifts in bacterial community composition in long-term  
336 microcosm studies (Krause *et al.*, 2012).

### 337 **Bacterioplankton activity along the latitudinal transect and temperature** 338 **gradient**

339 A positive linear correlation was observed between total microbial community leucine  
340 uptake rate and temperature ( $R^2=0.53$ ,  $p=0.001$ , Fig. 5a) with a single high outlier at  
341  $3.45^\circ$  C corresponding to the northernmost station in relatively warm Atlantic-  
342 influenced water (Station 11, Fig. 1). In contrast, the bioavailable leucine  
343 concentration was independent of temperature ( $R^2=0.006$ ,  $p=0.73$ , Fig. 5b).

344 Average cellular leucine uptake rates of SAR11 followed the same trend of  
345 increasing uptake rate with increasing temperature ( $R^2=0.24$ ,  $p=0.05$ , Fig. 5c) as  
346 observed for the total microbial community. The cellular uptake trend observed for  
347 SAR11 was enhanced at the population level ( $R^2=0.86$ ,  $p<0.001$ , Fig. 5d) reflecting  
348 their increased abundance along the temperature gradient towards elevated  
349 temperatures (Fig. 2a). The positive correlation of both abundance and leucine  
350 uptake rates (Fig. 5) of SAR11 bacteria with temperature suggests that, owing to  
351 their tolerance to acidification, SAR11 could thrive in polar regions following the  
352 predicted rise in sea surface temperature (Rhein *et al.*, 2013). In addition, SAR11  
353 could profit from the immediate increase in bio-available leucine following  
354 acidification (Fig. 3) since it has been shown that SAR11-specific transporter gene  
355 expression increases significantly after the addition of dissolved organic matter  
356 (Poretsky *et al.*, 2010). Indeed, different temperature-adapted phylotypes can co-

357 occur within the SAR11 population indicating potential ecological niche differentiation  
358 between SAR11 subgroups (Brown *et al.*, 2012).

359 In contrast to LNA, the cellular and population uptake rates of the heterogeneous  
360 HNA were not significantly correlated with temperature ( $R^2=0.12$ ,  $p=0.16$ , Fig. 5e,  
361 and  $R^2=0.11$ ,  $p=0.18$ , Fig. 5f, respectively). Compositional similarity of the HNA  
362 population, according to our FISH analyses (Fig. 2c), across stations, could  
363 potentially indicate that temperature has limited effect on the leucine uptake rates of  
364 the HNA. No relationships were observed between cellular uptake rates and the  
365 abundance of each of the three identified HNA groups (*Bacteroidetes*,  
366 *Alphaproteobacteria* and *Gammaproteobacteria*) (Fig. 2c); however  $38\pm 18\%$  of the  
367 HNA remained unidentified.

368 Considerable latitudinal variation was observed in the HNA:SAR11 ratio of average  
369 cellular leucine uptake rates (Fig S6). At lower latitudes (60-75°N) the difference in  
370 cellular uptake rate was much less pronounced with similar values recorded for both  
371 SAR11 and HNA at several stations. Interestingly, significantly higher cellular uptake  
372 rates were observed for HNA in the Fram Strait region compared to lower latitudes  
373 (Mann-Whitney,  $p=0.007$ , 77-79°N; Fig 4), despite a weak negative correlation  
374 between latitude and total microbial uptake rates ( $R^2=0.3$ ,  $p=0.009$ ) This could  
375 perhaps be due to a shift in community composition owing to the fact that three of  
376 these four highest latitude stations were associated with ice cover.

## 377 **Conclusions**

378 Acidification of seawater mimicking end of the century  $pCO_2$  levels or accidental  
379 release of acidic compounds in shipping resulted in an instantaneous increase (1.8  
380 fold on average) of bioavailability of the amino acid leucine. In contrast, it had no



381 effect on amino acid uptake rates of SAR11 and other heterotrophic bacteria along a  
382 natural temperature gradient suggesting a high physiological tolerance of  
383 bacterioplankton to acidification in the North Atlantic Ocean, Norwegian and  
384 Greenland Seas. A positive correlation between temperature and cell abundance as  
385 well as amino acid uptake rates of SAR11 bacteria indicated that SAR11 might  
386 benefit from anticipated future climate-induced increases in sea surface temperature  
387 by increasing their population abundance and productivity.

### 388 **Funding**

389 This work was supported by the UK Natural Environment Research Council, the UK  
390 Department of Environment, Food and Rural Affairs (Defra), and the UK Department  
391 of Energy and Climate Change (DECC) via the UK Ocean Acidification research  
392 programme (NERC Grants NE/H017097/1, NE/H017348/1 and NE/H016988/1)

### 393 **Acknowledgements**

394 We thank the captain, officers and the crew of the RRS James Clark Ross on JR271  
395 for their help and support at sea, and Elaine Mitchell for provision of additional flow  
396 cytometry data. We would also like to thank Dr Bernhard M. Fuchs for critical  
397 comments on earlier drafts of the manuscript. We are also grateful to the UK Natural  
398 Environment Research Council, the UK Department of Environment, Food and Rural  
399 Affairs (Defra), and the UK Department of Energy and Climate Change (DECC) for  
400 funding the research cruise via the UK Ocean Acidification research programme, and  
401 to the Danish, Icelandic and Norwegian diplomatic authorities for granting permission  
402 to travel and work in Greenland, Iceland and Svalbard coastal and offshore waters.  
403 The bathymetry in Fig. 1 is reproduced from the GEBCO Digital Atlas published by  
404 the British Oceanographic Data Centre on behalf of IOC and IHO, 2003.

405

406 The authors declare that there are no conflicts of interest.

407 **References**

408 Amann RI, Binder BJ, Olson RJ, Chisholm SW, Devereux R & Stahl DA (1990)

409 Combination of 16S rRNA-targeted oligonucleotide probes with flow cytometry for

410 analyzing mixed microbial populations. *Appl Environ Microbiol* **56**: 1919-1925.

411 Boyd PW, Sherry ND, Berges JA, *et al.* (1999) Transformations of biogenic

412 particulates from the pelagic to the deep ocean realm. *Deep-Sea Res Pt II* **46**: 2761-

413 2792.

414 Brown MV, Lauro FM, DeMaere MZ, *et al.* (2012) Global biogeography of SAR11

415 marine bacteria. *Mol Syst Biol* **8**.

416 Dickson AG & Millero FJ (1987) A comparison of the equilibrium constants for the

417 dissociation of carbonic acid in seawater media. *Deep-Sea Res Pt I* **34**: 1733-1743.

418 Dickson AG, Sabine C & Christian JR (2007) *Guide to best practices for ocean CO<sub>2</sub>*

419 *measurements*. N. Pac. Mar. Sci. Org., Sidney, BC, Canada.

420 Dupont CL, Rusch DB, Yooseph S, *et al.* (2012) Genomic insights to SAR86, an

421 abundant and uncultivated marine bacterial lineage. *ISME J* **6**: 1186-1199.

422 Engel A, Zondervan I, Aerts K, *et al.* (2005) Testing the direct effect of CO<sub>2</sub>

423 concentration on a bloom of the coccolithophorid *Emiliana huxleyi* in mesocosm

424 experiments. *Limnol Oceanogr* **50**: 493-507.

425 Fabry VJ, McClintock JB, Mathis JT & Grebmeier JM (2009) Ocean acidification at

426 high latitudes: the bellweather. *Oceanography* **22**: 160-171.

427 Gattuso J-P, Brewer PG, Hoegh-Guldberg O, Kleypas JA, Pörtner H-O & Schmidt

428 DN (2014) Cross-chapter box on ocean acidification. *Climate Change 2014: Impacts,*

429 *Adaptation, and Vulnerability Part A: Global and Sectoral Aspects Contribution of*  
430 *Working Group II to the Fifth Assessment Report of the Intergovernmental Panel on*  
431 *Climate Change*, (Field CB, Barros VR, Dokken DJ, et al., eds.), pp. 129-131.  
432 Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.

433 Giebel HA, Kalhoefer D, Lemke A, Thole S, Gahl-Janssen R, Simon M & Brinkhoff T  
434 (2011) Distribution of Roseobacter RCA and SAR11 lineages in the North Sea and  
435 characteristics of an abundant RCA isolate. *ISME J* **5**: 8-19.

436 Gomez-Pereira PR, Hartmann M, Grob C, Tarran GA, Martin AP, Fuchs BM,  
437 Scanlan DJ & Zubkov MV (2013) Comparable light stimulation of organic nutrient  
438 uptake by SAR11 and *Prochlorococcus* in the North Atlantic subtropical gyre. *ISME J*  
439 **7**: 603-614.

440 Halpern BS, Frazier M, Potapenko J, et al. (2015) Spatial and temporal changes in  
441 cumulative human impacts on the world's ocean. *Nature Communications* **6**: 7.

442 Hansell DA & Carlson CA (2014) Biogeochemistry of Marine Dissolved Organic  
443 Matter. pp. 712ff. Academic Press, London, UK.

444 Hartmann M, Zubkov MV, Scanlan DJ & Lepère C (2013) In situ interactions  
445 between photosynthetic picoeukaryotes and bacterioplankton in the Atlantic Ocean:  
446 evidence for mixotrophy. *Environ Microbiol Rep* **5**: 835-840.

447 Hassellöv I-M, Turner DR, Lauer A & Corbett JJ (2013) Shipping contributes to  
448 ocean acidification. *Geophys Res Lett* **40**: 2731-2736.

449 Herlemann DPR, Woelk J, Labrenz M & Jurgens K (2014) Diversity and abundance  
450 of "Pelagibacterales" (SAR11) in the Baltic Sea salinity gradient. *Syst Appl Microbiol*  
451 **37**: 601-604.

452 Hill PG, Mary I, Purdie DA & Zubkov MV (2011) Similarity in microbial amino acid  
453 uptake in surface waters of the North and South Atlantic (sub-)tropical gyres. *Prog*  
454 *Oceanogr* **91**: 437-446.

455 Hofmann GE, Smith JE, Johnson KS, *et al.* (2011) High-frequency dynamics of  
456 ocean pH: a multi-ecosystem comparison. *Plos One* **6**.

457 Hutchins DA, Mulholland MR & Fu FX (2009) Nutrient cycles and marine microbes in  
458 a CO<sub>2</sub>-enriched ocean. *Oceanography* **22**: 128-145.

459 Joint I, Doney SC & Karl DM (2011) Will ocean acidification affect marine microbes?  
460 *ISME J* **5**: 1-7.

461 Kemp PF, Lee S & Laroche J (1993) Estimating the growth rate of slowly growing  
462 marine bacteria from RNA content. *Appl Environ Microbiol* **59**: 2594-2601.

463 Kirchman D, Knees E & Hodson R (1985) Leucine incorporation and its potential as  
464 a measure of protein synthesis by bacteria in natural aquatic systems. *Appl Environ*  
465 *Microbiol* **49**: 599-607.

466 Krause E, Wichels A, Gimenez L, Lunau M, Schilhabel MB & Gerdtz G (2012) Small  
467 changes in pH have direct effects on marine bacterial community composition: a  
468 microcosm approach. *Plos One* **7**.

469 Lewis E & Wallace DWR (1998) Program developed for CO<sub>2</sub> system calculations.  
470 *ORNL/CDIAC-105 Carbon Dioxide Information Analysis Center, Oak Ridge National*  
471 *Laboratory, US Department of Energy, Oak Ridge, Tennessee.*

472 MacGilchrist GA, Shi T, Tyrrell T, Richier S, Moore CM, Dumousseaud C &  
473 Achterberg EP (2014) Effect of enhanced pCO<sub>2</sub> levels on the production of dissolved  
474 organic carbon and transparent exopolymer particles in short-term bioassay  
475 experiments. *Biogeosciences* **11**: 3695-3706.

476 Malmstrom RR, Straza TRA, Cottrell MT & Kirchman DL (2007) Diversity,  
477 abundance, and biomass production of bacterial groups in the western Arctic Ocean.  
478 *Aquat Microb Ecol* **47**: 45-55.

479 Mamaca E, Girin M, le Floch S & el Zir R (2009) Review of chemical spills at sea and  
480 lessons learnt.

481 Manz W, Amann R, Ludwig W, Wagner M & Schleifer KH (1992) Phylogenetic  
482 oligodeoxynucleotide probes for the major subclasses of proteobacteria - problems  
483 and solutions. *Syst Appl Microbiol* **15**: 593-600.

484 Manz W, Amann R, Ludwig W, Vancanneyt M & Schleifer KH (1996) Application of a  
485 suite of 16S rRNA-specific oligonucleotide probes designed to investigate bacteria of  
486 the phylum cytophaga-flavobacter-bacteroides in the natural environment.  
487 *Microbiology* **142**: 1097-1106.

488 Marie D, Partensky F, Jacquet S & Vaulot D (1997) Enumeration and cell cycle  
489 analysis of natural populations of marine picoplankton by flow cytometry using the  
490 nucleic acid stain SYBR Green I. *Appl Environ Microbiol* **63**: 186-193.

491 Martin JH, Knauer GA, Karl DM & Broenkow WW (1987) VERTEX - Carbon cycling  
492 in the Northeast Pacific. *Deep-Sea Res Pt I* **34**: 267-285.

493 Mehrbach C, Culberson CH, Hawley JH & Pytkowicz RM (1973) Measurement of the  
494 apparent dissociation constants of carbonic acid in seawater at atmospheric  
495 pressure. *Limnol Oceanogr* **18**: 897-907.

496 Morel FMM, Milligan AJ & Saito MA (2003) Marine bioinorganic chemistry: the role of  
497 trace metals in the oceanic cycles. . *The Oceans and Marine Geochemistry -*  
498 *Treatise on Geochemistry*,(Holland HD & Turekian KK, eds.), Elsevier-Pergamon,  
499 Oxford UK.

500 Morris RM, Rappe MS, Connon SA, Vergin KL, Siebold WA, Carlson CA &  
501 Giovannoni SJ (2002) SAR11 clade dominates ocean surface bacterioplankton  
502 communities. *Nature* **420**: 806-810.

503 Motegi C, Tanaka T, Piontek J, Brussaard CPD, Gattuso JP & Weinbauer MG (2013)  
504 Effect of CO<sub>2</sub> enrichment on bacterial metabolism in an Arctic fjord. *Biogeosciences*  
505 **10**: 3285-3296.

506 Neef A (1997) Anwendung der in situ Einzelzell-Identifizierung von Bakterien zur  
507 Populationsanalyse in komplexen mikrobiellen Biozönosen. Doctoral Thesis Thesis,  
508 Technische Universität München.

509 Poretsky RS, Sun SL, Mou XZ & Moran MA (2010) Transporter genes expressed by  
510 coastal bacterioplankton in response to dissolved organic carbon. *Environ Microbiol*  
511 **12**: 616-627.

512 Rhein M, Rintoul SR, Aoki S, *et al.* (2013) Observations: Ocean. *Climate Change*  
513 *2013: The Physical Science Basis Contribution of Working Group I to the Fifth*  
514 *Assessment Report of the Intergovernmental Panel on Climate Change*, (Stocker TF,  
515 Qin D, Plattner G-K, Tignor M, Allen SK, Boschung J, Nauels A, Xia Y, Bex V &  
516 Midgley PM, eds.), Cambridge University Press,, Cambridge, United Kingdom and  
517 New York, NY, USA.

518 Richier S, Achterberg EP, Dumousseaud C, Poulton AJ, Suggett DJ, Tyrrell T,  
519 Zubkov MV & Moore CM (2014) Phytoplankton responses and associated carbon  
520 cycling during shipboard carbonate chemistry manipulation experiments conducted  
521 around Northwest European shelf seas. *Biogeosciences* **11**: 4733-4752.

522 Riebesell U (2004) Effects of CO<sub>2</sub> enrichment on marine phytoplankton. *J Oceanogr*  
523 **60**: 719-729.

524 Riebesell U, Fabry VJ, Hansson L & Gattuso J-P (2010) Guide to best practices for  
525 ocean acidification research and data reporting. Publications Office of the European  
526 Union, Luxembourg.

527 Sabine CL, Feely RA, Gruber N, *et al.* (2004) The oceanic sink for anthropogenic  
528 CO<sub>2</sub>. *Science* **305**: 367-371.

529 Selinger DW, Saxena RM, Cheung KJ, Church GM & Rosenow C (2003) Global  
530 RNA half-life analysis in *Escherichia coli* reveals positional patterns of transcript  
531 degradation. *Genome Res* **13**: 216-223.

532 Sharma AK, Becker JW, Ottesen EA, Bryant JA, Duhamel S, Karl DM, Cordero OX,  
533 Repeta DJ & DeLong EF (2014) Distinct dissolved organic matter sources induce  
534 rapid transcriptional responses in coexisting populations of *Prochlorococcus*,  
535 *Pelagibacter* and the OM60 clade. *Environ Microbiol* **16**: 2815-2830.

536 Sherr EB & Sherr BF (2002) Significance of predation by protists in aquatic microbial  
537 food webs. *Anton Leeuw Int J G* **81**: 293-308.

538 Shi DL, Xu Y, Hopkinson BM & Morel FMM (2010) Effect of ocean acidification on  
539 iron availability to marine phytoplankton. *Science* **327**: 676-679.

540 Steglich C, Lindell D, Futschik M, Rector T, Steen R & Chisholm SW (2010) Short  
541 RNA half-lives in the slow-growing marine cyanobacterium *Prochlorococcus*.  
542 *Genome Biol* **11**: R54.

543 van Heuven S, Pierrot D, Rae JWB, Lewis E & Wallace DWR (2011) MATLAB  
544 program developed for CO<sub>2</sub> system calculations. *ORNL/CDIAC-105b Carbon*  
545 *Dioxide Information Analysis Center, Oak Ridge National Laboratory, US*  
546 *Department of Energy, Oak Ridge, Tennessee.*

547 Wright RT & Hobbie JE (1966) Use of glucose and acetate by bacteria and algae in  
548 aquatic ecosystems. *Ecology* **47**: 447-464.

549 Zark M, Riebesell U & Dittmar T (2015) Effects of ocean acidification on marine  
550 dissolved organic matter are not detectable over the succession of phytoplankton  
551 blooms. *Science Advances* **1**, DOI: 10.1126/sciadv.1500531.

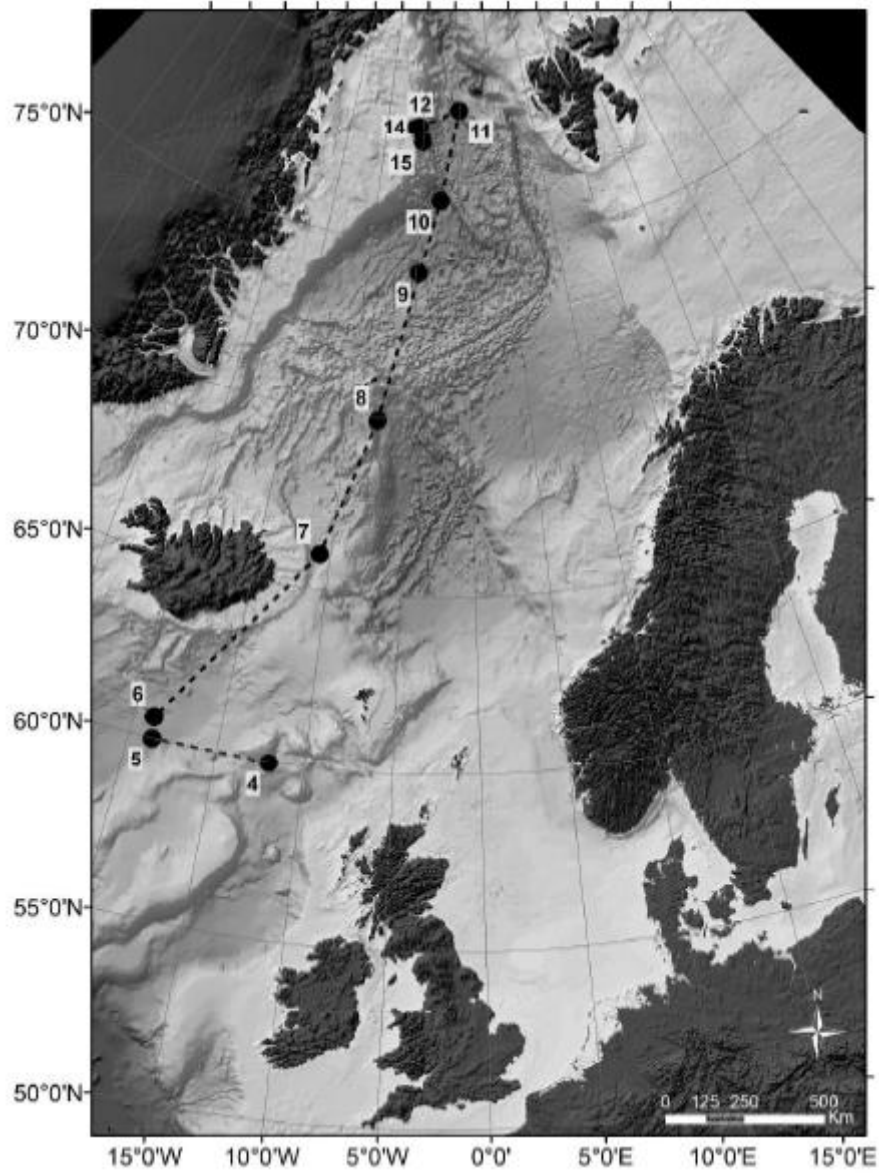
552 Zeebe RE & Wolf-Gladrow D (2001) *CO<sub>2</sub> in seawater: equilibrium, kinetics, isotopes*.  
553 Elsevier, Amsterdam.

554 Zubkov MV & Burkill PH (2006) Syringe pumped high speed flow cytometry of  
555 oceanic phytoplankton. *Cytometry Part A* **69A**: 1010-1019.

556 Zubkov MV, Fuchs BM, Burkill PH & Amann R (2001) Comparison of cellular and  
557 biomass specific activities of dominant bacterioplankton groups in stratified waters of  
558 the Celtic Sea. *Appl Environ Microbiol* **67**: 5210-5218.

559





560

561

Fig. 1. Location of stations sampled during this study in the North Atlantic Ocean, Norwegian and Greenland Seas.

562

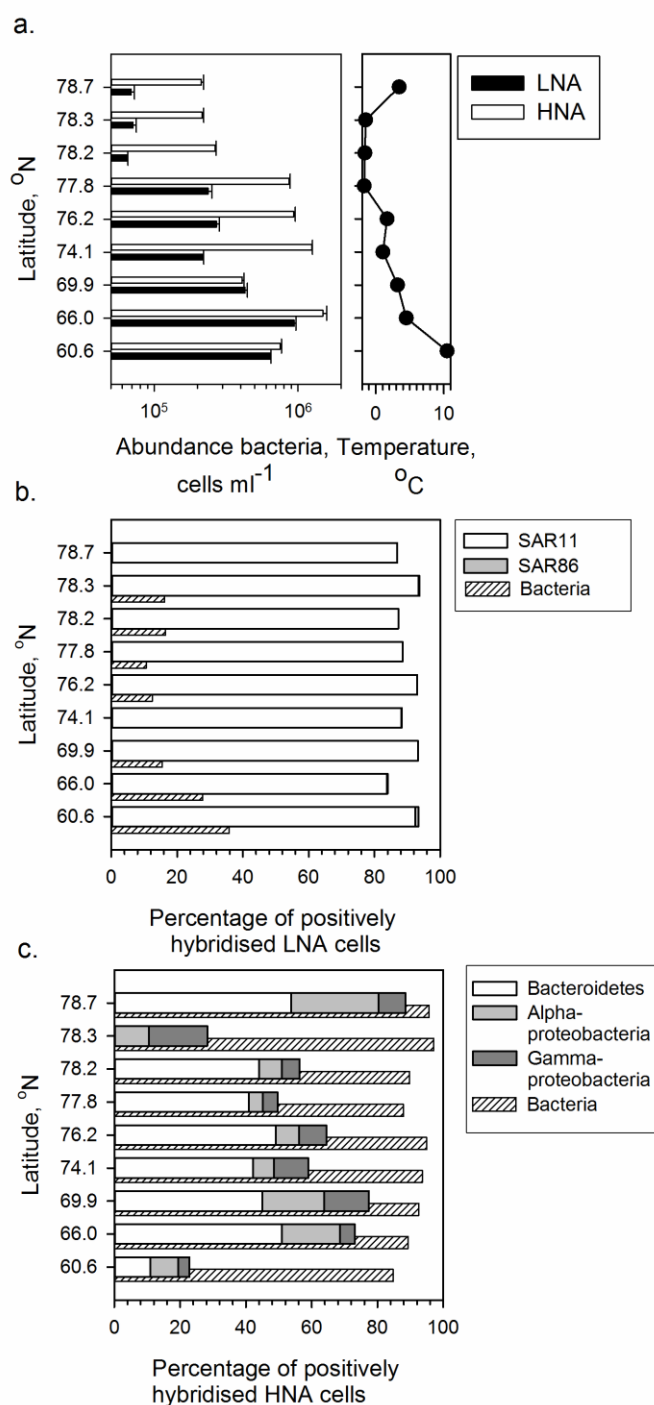
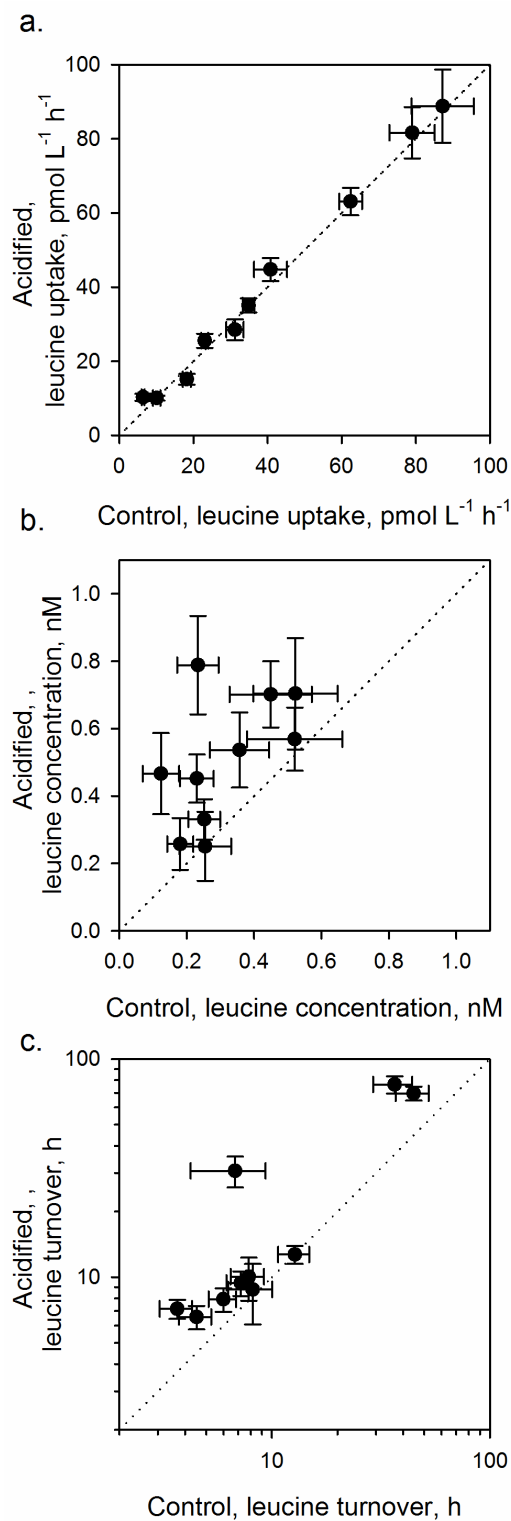


Fig. 2. Abundance (a) and community composition (b, c) of low nucleic acid and high nucleic acid bacteria (LNA and HNA, respectively) in offshore waters along a latitudinal transect crossing the North Atlantic Ocean, Norwegian and Greenland Seas. To simplify temperature and latitudinal correlation, the temperature measured along the latitudinal transect is shown in (a) next to the abundance. CARD-FISH, using probes targeting the main bacterial phyla, was performed on flow cytometrically sorted LNA and HNA cells (b, c) in order to establish community composition of the bacterial populations. Abbreviations in Fig2b: SAR11 *Alphaproteobacteria* (SAR11), SAR86 *Gammaproteobacteria* (SAR86). The error bars in (a) represent standard difference of mean ( $n=2$ ). Note that due to technical problems data were unavailable for the Bacteria (stations 9 and 11, b) and the Bacteroidetes (station 12, c) as assessed by CARD-FISH.

563

564



565

Fig. 3. Microbial uptake rate (a), bioavailable concentration (b) and turnover time (c) of leucine in control versus acidified experimental treatments from offshore waters along a latitudinal transect crossing the North Atlantic Ocean, Norwegian and Greenland Seas. The dotted line represents a 1:1 relationship with the distribution of points on or near this line indicating no impact of acidification relative to the control. Error bars indicate propagated error of regression slope, x-axis intercept and y-axis intercept in (a), (b) and (c) respectively.

566

567

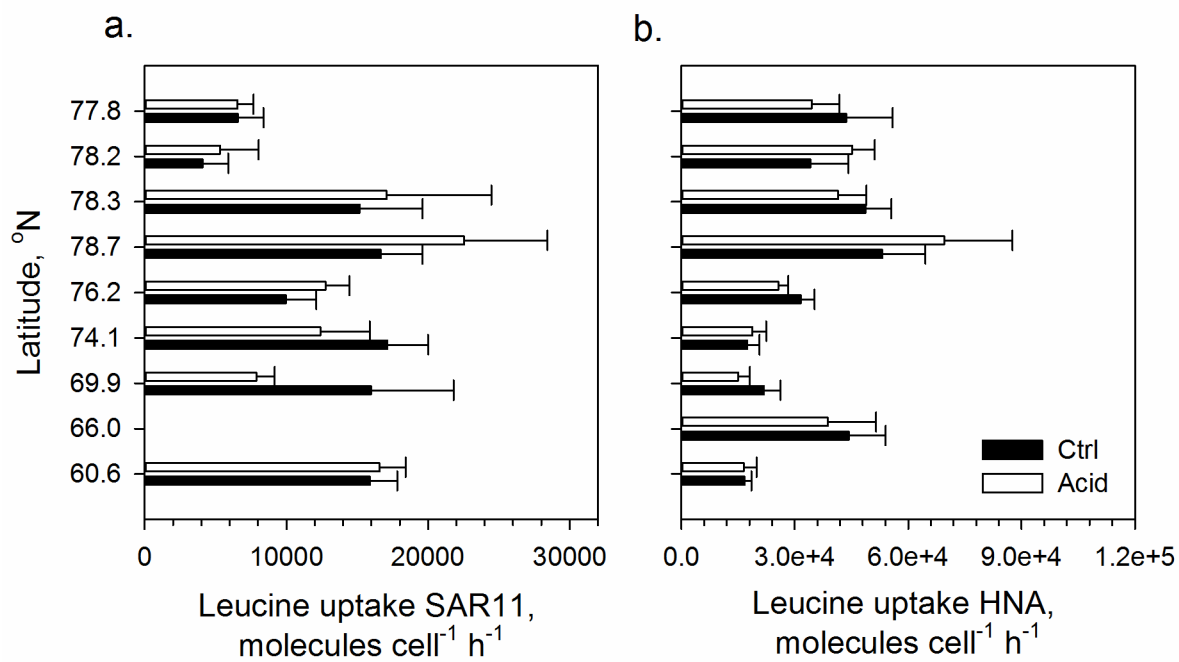
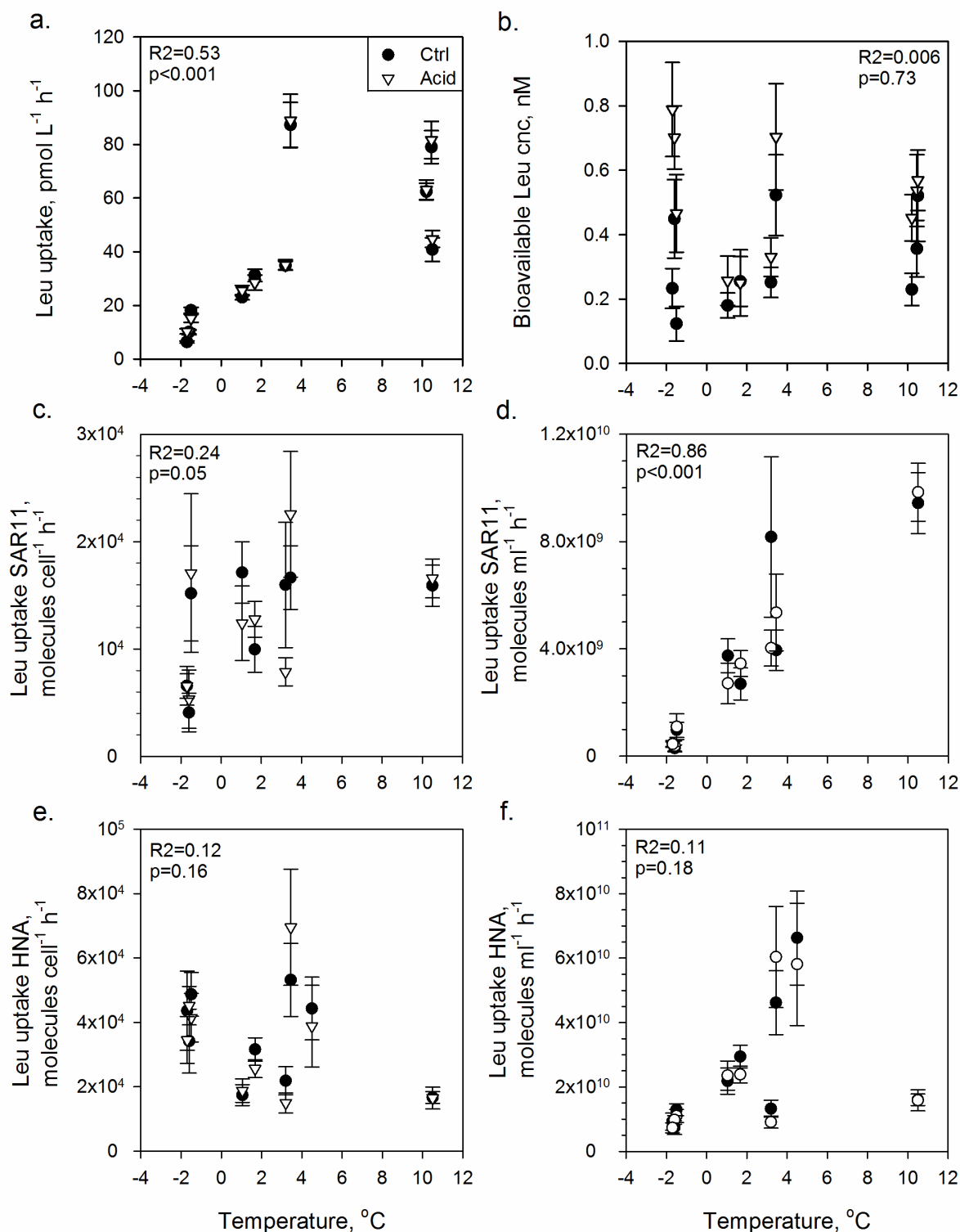


Fig. 4. Average ( $n=4$ ) cellular leucine uptake rates of SAR11 Alphaproteobacteria (SAR11, a) and high nucleic acid (HNA, b) bacteria in control (Ctrl) versus acidified (Acid) experimental treatments from offshore waters along a latitudinal transect crossing the North Atlantic Ocean, Norwegian and Greenland Seas. Error bars indicate propagated error of bioassay regression slope and standard deviation of different flow cytometric sorts ( $n=4$ ). Arrows indicate stations under ice cover. Note that data for leucine uptake by SAR11 were unavailable from station 7.

568

569

570



571

Fig. 5. Microbial uptake rates and bioavailable concentration of leucine versus temperature in control (Ctrl) and acidified (Acid) experimental treatments from offshore waters along a latitudinal transect crossing the North Atlantic Ocean, Norwegian and Greenland Seas. (a) Total microbial community uptake rate; (b) Bioavailable leucine concentration; (c) SAR11 *Alphaproteobacteria* (SAR11) average cellular uptake rate; (d) SAR11 population uptake rate; (e) High nucleic acid bacteria (HNA) average cellular uptake rate; (f) HNA population uptake rate.  $R^2$  and  $p$  values of the linear regression are given for each panel. Please note the different scales on the y-axis of the panels. Error bars in (a) and (b) indicate propagated error of regression slope and intercept. In (c-f) error bars show the propagated error of regression slope, standard deviation of different flow cytometric sorts ( $n=4$ ) and standard difference of counts ( $n=2$ ). Note that data on sorted bacterial cells were unavailable from stations 4 and 5.

572