

1 Alluvial fan records from southeast Arabia reveal multiple
2 windows for human dispersal

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15 **ABSTRACT**

16 The dispersal of human populations out of Africa into Arabia **was** most likely
17 linked to episodes of climatic amelioration, when increased monsoon rainfall led to the
18 activation of drainage systems, improved freshwater availability, and the development of
19 regional vegetation. Here we present the first dated terrestrial record from southeast
20 Arabia that provides evidence for increased rainfall and the expansion of vegetation
21 during both glacial and interglacial periods. Findings from extensive alluvial fan deposits
22 indicate that drainage system activation occurred during Marine Isotope Stage (MIS) 6

23 (ca. 160–150 ka), MIS 5 (ca. 130–75 ka), and during early MIS 3 (ca. 55 ka). The
24 development of active freshwater systems during these periods corresponds with
25 monsoon intensity increases during insolation maxima, suggesting that humid periods in
26 Arabia were not confined to eccentricity-paced deglaciations, and providing
27 paleoenvironmental support for multiple windows of opportunity for dispersal out of
28 Africa during the late Pleistocene.

29 INTRODUCTION

30 Considerable debate surrounds the dispersal of human populations out of Africa.
31 There are two predominant hypotheses concerning the timing of dispersal, each
32 contrasting in their emphasis on the role of the Arabian interior and its changing climate.
33 In one scenario, human populations expanded from an ancestral base in Africa, dispersing
34 rapidly along the coastlines of Arabia to southern Asia ca. 60–50 ka (e.g., Mellars et al.,
35 2013). Another model posits that dispersal into the Arabian interior began much earlier
36 (ca. 130–75 ka), possibly during multiple phases (Groucutt and Petraglia, 2012), when
37 increased rainfall provided sufficient freshwater to support expanding populations. Dated
38 archeological sites in Arabia (Armitage et al., 2011; Petraglia et al., 2011, 2012; Rose et
39 al., 2011) support an earlier model, and correspond with humid phases during Marine
40 Isotope Stages (MIS) 5e (ca. 128–115 ka), 5c (ca. 105–95 ka), and 5a (ca. 85–74 ka).

41 Speleothem and lacustrine records confirm that during interglacial periods such as
42 MIS 5e there were significant increases in humidity in Arabia, following northward
43 incursions of monsoon rainfall (e.g., Burns et al., 2001; Fleitmann et al., 2003; Rosenberg
44 et al., 2011). However, archives such as speleothems require >350 mm of rainfall to
45 initiate growth (e.g., Vaks et al., 2010; Fleitmann et al., 2011), and periods of lesser

46 rainfall that may be sufficient to sustain human populations but not speleothem growth
47 are absent from these records. Consequently, some studies have suggested that climatic
48 conditions during mid-high-latitude glacial periods represent a natural barrier to humans
49 (e.g., Rosenberg et al., 2011), despite the presence of archeological sites dated to early
50 MIS 3 (between ca. 55 and 40 ka) in the United Arab Emirates and Yemen (Armitage et
51 al., 2011; Delagnes et al., 2012). By contrast, alluvial fan sequences have the potential to
52 record a wide range of landscape changes and climatic events. Here we present a unique
53 alluvial fan aggradation record from southeast Arabia spanning the past ca. 160 k.y.
54 Situated along the proposed southern dispersal route, the Al Sibetah record is to date the
55 most comprehensive terrestrial archive from the Arabian Peninsula, and provides
56 evidence for multiple humid episodes during both glacial and interglacial periods.

57 ENVIRONMENTAL SETTING AND METHODS

58 Extensive bajadas emanate from the western flanks of the Hajar Mountains and
59 extend to the Wahiba Sands in Oman (Fig. 1). Situated at the interface of the mountains
60 and the Rub al Khali sand sea, these fans are particularly sensitive to changes in rainfall
61 patterns and thus are excellent indicators of monsoon variability. Modern alluvial fans are
62 largely relict under present climatic conditions; however, previous larger fans extended
63 much further west and are now concealed beneath dunes. Several studies have
64 highlighted the importance of these deposits as potentially important climatic indicators
65 (Maizels, 1987; Burns and Matter, 1995; Preusser et al., 2002; Blechschmidt et al., 2009).
66 The Al Sibetah quarry (N24.33346, E55.76125) exposes a 42 m succession within the
67 partially buried Al Ain fan, comprising interbedded sands, conglomerates, and paleosols
68 in multiple fining-upward cycles.

69 Paleoenvironmental (geochemical, carbonate isotope, magnetic susceptibility,
70 phytolith, laser granulometry) analyses of the fine-grained sediment component (<2 mm)
71 were used to detect the response of fan aggradation processes to climatic changes. The
72 chronology is constrained by 12 single aliquot regenerative optically stimulated
73 luminescence dates. Full methods and additional paleoenvironmental and chronological
74 information are presented in the GSA Data Repository¹.

75 PALEOENVIRONMENTAL RECONSTRUCTIONS AND CHRONOLOGY

76 The Al Sibetah quarry comprises medial-distal alluvial fan sediments visible in
77 multiple exposures 20–100 km downfan (Fig. 1). The fan is very low gradient (0.2°–
78 0.5°), featuring imbricated gravel beds ~20 km from the fan apex, indicative of a fluvial
79 (climatic) rather than a gravity-driven (base-level change) regime. The mixed clast and
80 fine-grained sedimentary composition of the sequence is typical of mixed-load streams
81 with braided reaches. Stream gravels, typified by high susceptibility values, fine upward
82 into calcareous silts and sands, paleosols, or eolian material over 13 aggradation
84 phases (Fig. 2), each indicating a gradual waning of stream flow. Thick
85 homogeneous paleosols with extensive rootlets and bioturbation represent gradual
86 sedimentation, landscape stabilization, and grassland development, which typically
87 occurs over many years. Six periods of stream channel aggradation followed by soil
88 and grassland development occurred during late MIS 6, with a mean depositional age
89 of 158.7 ± 12.9 ka. This is in keeping with a peak in summer monsoon intensity and
90 insolation between ca. 160 and 150 ka (e.g., Clemens and Prell, 2003). Phytolith data
91 indicate a mix of C₃ (more humid) and C₄ (more arid) grassland types, with C₃ pooid

92 phytolith morphotypes accounting for between 27% and 60%, and panocoid and
93 chloridoid values ranges being 4.7%–10% and 3%–10%, respectively. The $\delta^{13}\text{C}$ values
94 from pedogenic carbonates indicate an overall increase in the proportion of C_4 grasses
95 across the landscape throughout late MIS 6 (–4.89‰ to +0.99‰). Soil and grassland
96 formation and increased regional humidity are generally accompanied by a depletion of
97 $\delta^{18}\text{O}$ values, increased hydrolysis, and decreased salinization, and decreased particle size
98 and magnetic susceptibility values following the cessation of detrital and clastic influx
99 (Fig. 2).

100 Three phases of stream channel activation and grassland development occurred
101 between ca. 130 ka and ca. 88 ka, representing wet phases during MIS 5e (ca. 128–115
102 ka), MIS 5c (ca. 105–95 ka), and MIS 5a (ca. 85–74 ka). MIS 5e channel gravels are
103 asymmetric in cross section with greater imbrication and depth, reflecting a localized
104 increase in gravel accumulation and point-bar development indicative of a larger
105 meandering channel. Soil formation and vegetative processes were restricted during MIS
106 5 (likely due to a more mobile channel with strong currents), and the zone of grassland
107 formation was farther downfan. The $\delta^{13}\text{C}$ values during MIS 5 are between –3.45‰ and
108 +0.58‰ and indicate a dominance of C_4 grasses throughout the landscape; however, a
109 higher proportion of detrital and clastic material makes climatic interpretations from
110 isotope data problematic. Panicoid values are highest (13%) during MIS 5e, with an
111 overall higher proportion of woody taxa in the region during MIS 5 than MIS 6, and the
112 highest values during MIS 5e. More mesic conditions are observed during MIS 5e and
113 MIS 5a, and more xeric conditions during MIS 5c. This is supported by generally lower
114 hydrolysis and greater salinization during MIS 5c.

115 MIS 5 sediments are overlain by a 2-m-thick sequence of eolian material dated to
116 ca. 73 ka, reflecting dune mobilization and increased aridity during MIS 4. These are
117 overlain by stream-channel gravels, which represent the redevelopment of a braided
118 channel network during a subsequent humid phase. This stratigraphic unit was identified
119 in distal exposures of the Al Sibetah fan at Remah (Fig. 1) ~40 km to the
120 southwest and dated to ca. 55 ka (Farrant et al., 2012), and is indicative of increased
121 rainfall during early MIS 3. No Holocene-age fan aggradation was recorded at Al
122 Sibetah; however, an age of ca. 5 ka was recorded in overlying dune sands, representing a
123 phase of dune formation during the mid-Holocene that broadly corresponds with the end
124 of increased Holocene rainfall in Arabia (Parker, 2009).

125 **DISCUSSION**

126 The sequence at Al Sibetah provides a unique record of climate-driven landscape
127 changes in Arabia between MIS 6 and MIS 3. This period coincides with key processes in
128 human evolution: the origin of *Homo sapiens* in Africa (ca. 200–150 ka), the early
129 dispersal of humans out of Africa (ca. 130–80 ka), and the post-MIS 5 dispersal into Asia
130 during MIS 3 (ca. 60–50 ka). The availability of freshwater in the southern dispersal zone
131 during these periods would have been critical to the demographic expansion of early
132 human populations. Increased vegetation and grassland development within range of the
133 mountains would have provided an attractive location for hunting, while the development
134 of shallow aquifers along the bajada would have extended potable water resources for
135 many months of the year by feeding springs and water holes.

136 Increased freshwater availability ca. 160–150 ka was accompanied by the
137 development of savannah-type grasslands with a woodland component. While increased

138 humidity at this time is documented in the Levant, central Negev, and southern Jordan
139 (Bar-Matthews et al., 2003; Petit-Maire et al., 2010; Vaks et al., 2010), speleothem data
140 from southern regions of Arabia suggest a continually hyperarid climate, with monsoon
141 rainfall displaced substantially south due to increased glacial boundary conditions
142 (Fleitmann et al., 2011). This study presents significant evidence for increased rainfall
143 during MIS 6 in southern Arabia, demonstrating that the inland convection of monsoon
144 rainfall is not prevented during glacial periods. This corresponds with marine evidence
145 for monsoon intensification ca. 165–150 ka (e.g., Malaizé et al., 2006; Caley et al., 2011),
146 and incipient soil formation reported in the Wahiba between ca. 160 and 140 ka (Preusser
147 et al., 2002; Radies et al., 2004).

148 Fan activation throughout MIS 5 is consistent with speleothem and lake records
149 from Oman, Yemen, and Saudi Arabia (i.e., Fleitmann et al., 2011; Rosenberg et al.,
150 2011), and with an early expansion of human populations out of Africa (Groucutt and
151 Petraglia, 2012). MIS 5e channel flow was typified by a larger fluvial system, which at
152 times may have extended to the Persian Gulf coast (Farrant et al., 2012),
153 facilitating demographic connectivity between mountainous and coastal regions.
154 Drainage activation in the southernmost reaches of the bajada during MIS 5 led to the
155 formation of the ~1400 km² paleolake Saiwan (Rosenberg et al., 2012), while to the
156 north, extensive braided stream development occurred at Wadi Dhaid (Atkinson et al.,
157 2013). Findings from Al Sibetah also indicate that MIS 5c was the weaker of the three
158 MIS 5 humid periods, with decreased woody taxa and more xeric conditions than MIS 5e
159 or 5a. It is important that the timing of freshwater availability and grassland development
160 at Al Sibetah coincides with the occupation of Jebel Faya ca. 130–90 ka (Figs. 1 and 3),

161 suggesting that fan drainage processes may have facilitated population movements
162 through the region during **that** time.

163 Other studies have stated that ca. 75–10 ka, intense aridity prevailed in Arabia
164 (e.g., Fleitmann et al., 2011; Rosenberg et al., 2011), preventing population movements
165 through the interior. The posited major expansion of human populations out of Africa ca.
166 60–50 ka is in keeping with this notion, suggesting an exclusively coastal route (Mellars
167 et al., 2013). However, paleoenvironmental evidence from Arabia indicates that sufficient
168 freshwater resources existed in the interior to support expanding populations. Drainage
169 activation between ca. 60 and 50 ka **is** recorded in several regions (Krbetschek, 2008;
170 McLaren et al., 2009; Farrant et al., 2012; Parton et al., 2013), congruent with an increase
171 in monsoon intensity recorded in marine records (e.g., Clemens and Prell, 2003; Govil
172 and Naidu, 2010), speleothem data from Socotra (Burns et al., 2003), and a Middle
173 Paleolithic site in Yemen dated to ca. 55 ka (Delagnes et al., 2012). Therefore, the notion
174 that arid conditions during MIS 3 would have prevented dispersals through the Arabian
175 interior is no longer tenable and thus challenges an exclusively coastal route of migration.
176 **It is** significant that the periodicity of increased rainfall presented by the Al Sibetah and
177 surrounding regional records **suggests** that incursions of monsoon rainfall, and in turn,
178 periods of demographic expansion through the interior of Arabia, may have been driven
179 by insolation maxima (Fig. 3), rather than mid-high-latitude deglaciations.

180 **CONCLUSIONS**

181 The Al Sibetah alluvial fan sequence provides a unique and sensitive record of
182 landscape change in southeast Arabia between MIS 6 and MIS 3. Phases of monsoon-
183 driven alluvial fan aggradation **occurred** during both glacial and interglacial periods, in

184 line with insolation maxima. The occurrence of humid periods not previously identified
185 in lacustrine or speleothem records highlights the complexity and heterogeneity of the
186 Arabian paleoclimate, and suggests that interior migration pathways through the Arabian
187 Peninsula may have been viable approximately every 23 k.y. since at least MIS 6.

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318 **FIGURE CAPTIONS**

319 Figure 1. Map showing location of the study site and extent of bajada system in southeast
320 Arabia, including other identified sections of the Al Ain fan (UAE—United Arab
321 Emirates). Inset is photo of the exposed quarry site (person circled for scale). Jebel Faya
322 (Armitage et al., 2011) and paleolake Saiwan (Rosenberg et al., 2012) are also shown.

324

325 Figure 2. Stratigraphic record and photograph of the sequence at Al Sibetah (southeast
326 Arabia), showing aggradation phases 1–13,
327 optically stimulated luminescence (OSL) ages, geochemical data
328 (hydrolysis, salinization), carbonate isotope values ($\delta^{18}\text{O}_{\text{carb}}$, $\delta^{13}\text{C}_{\text{carb}}$), magnetic
329 susceptibility values, and particle size analysis of <2 mm fine sediment fraction.

333

334 Figure 3. Comparison of paleoenvironmental records and dated Middle Paleolithic
335 archeological sites from Arabia. Gray bars indicate humid periods relevant to this study.

336 MIS—Marine Isotope Stage; VPDB—Vienna Peedee belemnite; XRF—X-ray
337 fluorescence; IOSM—Indian Ocean Summer Monsoon

338 Archeological sites: A—Shi’bat Dihya, Yemen (Delagnes et al., 2012); B—

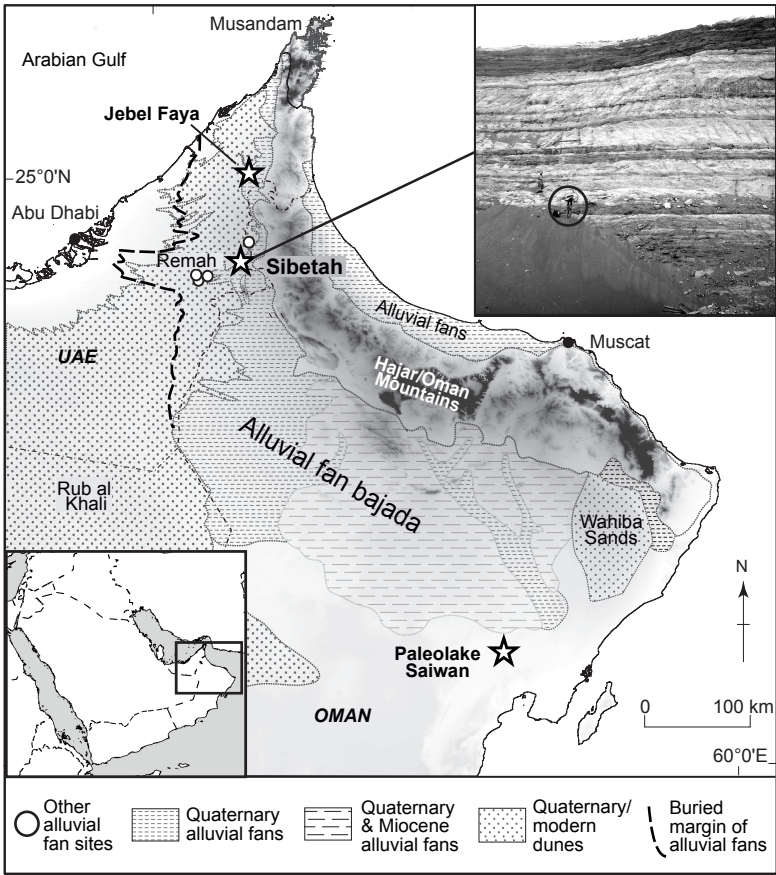
339 JQ1, northern Saudi Arabia (Petraglia et al., 2011); C—JKF, northern Saudi Arabia

340 (Petraglia et al., 2012); D—Aybut Auwal, southern Oman (Rose et al., 2011); E—Jebel

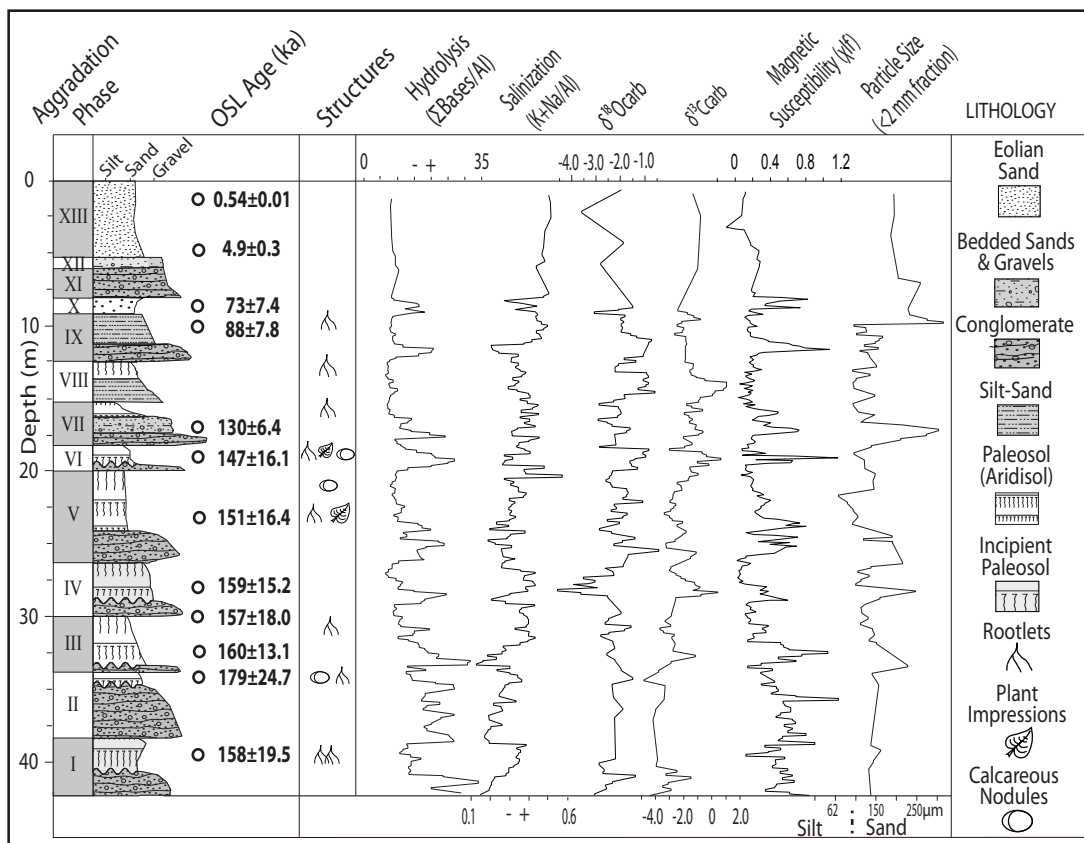
341 Faya (assemblages A–C), United Arab Emirates (UAE) (Armitage et al., 2011).
342 Paleoclimatic records: F—fluvial deposits, central Saudi Arabia (McLaren et al., 2009);
343 G—Aqabah paleolake, UAE (Parton et al., 2013); H—Al Ain alluvial fan (UAE) at
344 Remah (i), (Farrant et al., 2012) and Al Sibetah (ii); I—paleosols, Wahiba Sands, Oman
345 (Preusser et al., 2002), J—paleolake Saiwan, Oman (Rosenberg et al., 2012); K—
346 paleolakes Mundafan and Khujaimah, southern Saudi Arabia (Rosenberg et al., 2012);
347 L—Moomi Cave speleothems, Socotra, Yemen (Burns et al., 2003); M—speleothems
348 from Hoti Cave (Oman) and Mukalla Cave (Yemen) (Fleitmann et al., 2011); N—
349 Arabian Sea productivity records from foraminiferal assemblages (solid line) (Caley et
350 al., 2011) and bromine X-ray fluorescence counts (dashed line) (Ziegler et al., 2010);
351 O—summer monsoon stack (dashed line), summer monsoon factor (black line) (Clemens
352 and Prell, 2003); P—summer insolation at 30°N (Wm^{-2}) (Berger and Loutre, 1991).
Dashed boxes represent either age uncertainties or inferred ages.

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357 ¹GSA Data Repository item 2015xxx, details on methodologies, optically stimulated
358 luminescence geochronology, and phytolith analysis, is available online at
359 www.geosociety.org/pubs/ft2015.htm, or on request from editing@geosociety.org or
360 Documents Secretary, GSA, P.O. Box 9140, Boulder, CO 80301, USA.



Ash_Parton_Figure1.pdf



Ash_Parton_Figure2.pdf

