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Drumbeat LP 'aftershocks' to a failed explosive eruption at Tungurahua Volcano, Ecuador

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Key Points:

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| 7 | • | An episode of accelerating and decelerating long-period (LP) drumbeat earthquakes |
|----|---|---|
| 8 | | is identified at Tungurahua volcano, Ecuador |
| 9 | • | Bayesian gamma point process analysis constrains mirrored sequences of Omori |
| 10 | | Law accelerating and decelerating seismicity |
| 11 | • | Waveforms examined by cross correlation and Q factor reveal a gas driven, repeat- |
| 12 | | ing, single source, which failed to culminate in an eruption |

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13 Abstract

¹⁴ Highly periodic, repetitive long-period (LP) earthquakes, known as 'drumbeats', have

¹⁵ been observed at a range of volcanoes, typically during the ascent of degassed magma.

Accelerating rates of drumbeats have been reported before explosions, and potentially offer forecasts of future activity. However, the broader phenomenology of drumbeats is

offer forecasts of future activity. However, the broader phenomenology of drumbeats poorly understood. Here we describe an episode of over 900 LP earthquakes recorded

¹⁹ in November 2015 at Tungurahua Volcano, Ecuador, that we believe are associated with

a failed explosion. Rates of LP drumbeats accelerated for 10 hours, consistent with an

²¹ Inverse Omori's Law. Before any explosion occurred, seismicity decreased following Omori's

Law, over a further six days. Despite earthquake rates decelerating, amplitudes, spec-

tral peaks, Q values and periodicity remain constant, suggesting there is little change

²⁴ in the source process with time. We argue that the decelerating seismicity is a result of

²⁵ progressive reduction of gas flux, unable to provide sufficient overpressure for explosion.

²⁶ Plain Language Summary

When a volcano is erupting, small earthquakes from the volcano can be used to in-27 fer what internal processes may be occurring. Earthquakes that are very similar to one 28 another and repeat at consistent intervals are known as drumbeat earthquakes. These 29 are of interest in volcanic systems as it implies the earthquakes are generated by a sin-30 gle, repeating source. Previous studies of drumbeat earthquakes at Tungurahua Volcano, 31 32 Ecuador, have described these earthquakes occurring closer together in time and accelerating up to an explosion. In this case, we identify a sequence of drumbeats where the 33 rate accelerates, and without any explosion, decelerates again. We suggest these earth-34 quakes are generated by gas flux which is slowing down. This gas originates beneath a 35 plug at the top of the conduit. We use statistical models to estimate when the volcano 36 may have exploded if the earthquakes had continued to accelerate, and quantify the sub-37 sequent deceleration in earthquake rate. 38

³⁹ 1 Introduction

Active arc volcanoes of andesitic-dacitic composition are often sources of rich seis-40 mic data. Signals at these volcanoes are often dominated by long-period earthquakes (LPs), 41 commonly associated with processes occurring in and around the magma column. Un-42 derstanding these signals could be key to improving our ability to forecast volcanic ac-43 tivity. LPs are characterised by frequencies between 0.5 and 5.0Hz, emergent onsets, and 44 missing clear S wave arrivals (Chouet et al., 1994). They often begin with a mixed fre-45 quency onset, followed by low frequency coda that decays in amplitude with time. This 46 characteristic shape has been modelled as a two part process with an initial excitation 47 trigger and subsequent resonance (Chouet, 1996). These are some of the features that 48 have been used to distinguish different categories of volcano seismic events, attributed 49 to different source processes (Chouet & Matoza, 2013)(fig S1). Swarms of periodic, highly 50 similar, repeating LPs occur in a phenomenon known as drumbeats. Drumbeat seismic-51 ity is commonly associated with degassed magma ascent, however, the broader phenomenol-52 ogy of drumbeats is still poorly established. Locating LPs is generally a very difficult 53 process, however, with one or two stations, careful analysis of the waveforms and their frequency content can tell us about an evolving source mechanism. 55

⁵⁶ Drumbeat earthquakes are best known from the dacite spine extrusion episode at ⁵⁷ Mount St. Helens between 2004 and 2005. Iverson (2008) approximated long term steady-⁵⁸ state behaviour and slowly changing drumbeat rates and amplitudes with frictional stick-⁵⁹ slip at the conduit margins. However, drumbeat seismicity is known to display a vari-⁶⁰ ety of characteristics from many arc volcanoes. Drumbeat seismicity at Soufrière Hills ⁶¹ Volcano appeared in pulses lasting several hours (Green & Neuberg, 2006). These pulses ⁶² were associated with brittle failure of ascending magma at conduit margins (Neuberg

et al., 2006). The behaviour of drumbeats observed at Tungurahua alone is varied. One 63 study examined a six-day episode of steady-state, repeating LPs in 2001 where the Q 64 factors of individual earthquakes were changing through time (Molina et al., 2004). This 65 shift was modelled with repetitive injections of increasingly ash-laden gas. Repeated low 66 frequency (1-3Hz) pulses are recorded in both the seismic and infrasonic record for episodes 67 in 2004 (Ruiz et al., 2006). In July 2013, accelerating drumbeats merged into tremor be-68 fore a large explosions (Bell et al., 2018). A further study identified the incremental break-69 down of an episode of drumbeat LP seismicity during April 2015 (Bell et al., 2017). Build-70 ing on previous models at Soufrière Hills, a more developed plug model argued that LPs 71 were triggered by gas escape and shear failure in the conduit margins with magma as-72 cent. 73

Accelerating seismicity, has been related to material failure in the Failure Forecast 74 Method (FFM) (Main, 1999; Voight, 1988). New statistical methods allow analysis of 75 point process data, revealing properties of precursory sequences. Improved methods help 76 to quantify data, identify changes and understand underlying processes. We can exam-77 ine seismicity rates with relationships such as the Modified and Inverse Omori's Laws. 78 By contrasting accelerating and decelerating seismicity with models and examining this 79 'mirrored' effect we can investigate the significance of failed explosions, better understand 80 the physics of the process and develop forecasting statistics. 81

Here we describe a six day sequence of accelerating and decelerating drumbeat LP 82 earthquakes at Tungurahua during November 2015, associated with a 'failed' explosive 83 eruption. We use a Bayesian gamma point process model (Bell et al., 2017) to examine 84 the acceleration of seismicity rate, and the subsequent decelerating rate of seismicity. We 85 find that the drumbeats both accelerate and decelerate according to a power law with 86 an exponent value, $p = 0.96 \quad 0.51$ and $p = 0.97 \quad 0.12$ respectively. Despite prolonged 87 decaying temporal rates of seismicity, the earthquakes show strong similarity with fam-88 ilies persisting across the six day sequence and amplitudes unchanging. This suggests 89 a slowing rather than a breakdown of the driving source mechanism following a failed 90 eruption. 91

First we introduce the activity and data recorded at Tungurahua during November 2015. We then present the seismic data, along with the statistical methods for analysis. We model the data using a Bayesian point process methodology, testing different rate models and estimating parameter posterior distributions. We analyse earthquake properties including waveform similarity, families and Q factor values. We finally present a model for accelerating and decelerating drumbeats, and discuss the implications this has for magma ascent dynamics at Tungurahua.

⁹⁹ 2 Data & Methods

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2.1 Tungurahua

Tungurahua is a 5,032m high and esitic stratovolcano in the Central Cordillera of 101 the Ecuadorian Andes (Hall et al., 1999). The most recent phase of activity occurred be-102 tween 1999 and 2016 with notable sub-Plinian activity in 2006 (Mothes et al., 2015). Un-103 rest at Tungurahua was typically associated with high rates of LP seismicity. Between 104 the major explosive episodes of 2014 and 2016, heightened seismicity accompanied de-105 formation and repeating tilt cycles (Bell et al., 2017; Neuberg et al., 2018; Marsden et 106 al., 2019). This study focuses on an episode of drumbeats during one such cycle in Novem-107 ber 2015. The drumbeats persist for six days and did not culminate in any explosion. 108 There was then a repose period of 3 months before the final explosions in February 2016. 109 Drumbeat seismicity in persisted for several weeks in April 2015 and was accompanied 110 by small explosions and ash emissions (Bell et al., 2017). Whilst in October and early 111

November 2015, small pulses of drumbeat seismicity emerged and ceased over just a few
hours or days and are as yet unstudied (fig 1).

114 2.2 Monitoring data

The Instituto Geofísico de la Escuela Politécnica Nacional (IGEPN) maintain a vol-115 cano monitoring network on Tungurahua. The network includes short period and broad-116 band seismometers, DOAS gas flux stations, infrasound stations, tiltmeters, GPS, cam-117 eras and acoustic flow monitors. From a seismic network of 11 stations, IGEPN main-118 tain a catalogue of detected, classified, and where possible, located events. Over 90% of 119 events were recorded at RETU, a short period seismometer at elevation over 4000m, ap-120 proximately 2000m from the crater rim. This proximity means the signal to noise ratio 121 (SNR) is high and many small, shallow events are recorded. We manually picked 932 events 122 from 25 - 30 November 2015 for this study, representing all detectable events at RETU. 123 These events were only visible at the one station and with emergent onsets and no clear 124 S-phases, locating the events was not possible. As the seismicity is only recorded at this 125 uppermost station, we believe these LPs are associated with shallow processes in the top 126 2000m of the conduit (Bell et al., 2018). The similarity of the waveforms indicates that 127 they are all closely co-located within a small depth range. Given this co-location, we use 128 the maximum amplitudes of individual events as a relative comparison for magnitude. 129 615 of the manually picked events appear in the IGEPN catalogue. However, there are 130 only 20 events which are located and have estimated magnitudes, all of which are less 131 than magnitude 1.5, and carry large uncertainties. The seismicity on 25 November is the 132 first clearly identifiable sequence of LP events as in the preceding days, the signal at RETU 133 is dominated by emission tremor. 134

Details of surface observations and ash column heights are extracted from daily reports produced by the Observatorio del Volcán Tungurahua (OVT) (https://www.igepn . edu. ec/), and used in conjunction with the seismic data for temporal analysis. Explosion counts and radial tilt measurements at station RETU are also collected from IGEPN catalogues.

140 **2.3** Methods

The seismic data is initially processed using the ObsPy toolkit (Krischer et al., 2015). 30 second duration waveforms are sliced and bandpass filtered between 1 and 40Hz. The maximum amplitude of each event is extracted. Fast Fourier Transform (FFT) of each signal is calculated to generate a periodogram. We find the power spectral density (PSD) for frequencies sampled at an interval of 0.01Hz and extract the maximum value as the fundamental peak frequency.

The Q factor for each event is calculated using an auto-regressive moving average 147 (ARMA) technique, adapted from *Seismo-Volcanalysis* software (Lesage, 2007). The Q 148 factor is a non-dimensional number that describes how quickly or slowly wave energy dis-149 sipates and is often strongly linked to the fundamental peak frequency. Auto-regressive 150 methods have been successfully used to analyse changing LP frequency contents (Kumagai 151 & Chouet, 1999; Lokmer et al., 2008). The approach is similar to the commonly used 152 Sompi method (Hori et al., 1989). A signal is composed of a number of individual har-153 monic decaying oscillations. Each component can be represented in complex frequency 154 space and quantified by a peak frequency (f, Hz) and growth rate (g, s^{-1}) . (Kumazawa 155 et al., 1990). We generated cumulative f-q diagrams for all filters between 2 and 30 and 156 points that cluster around a pole at the spectral peak are used to calculate Q (Eq 1) (Cusano 157 et al., 2008). In this automated adaptation of the ARMA method, a hierarchical clus-158 tering method is used to automatically select points in the complex frequency space (Eads, 159 2008).160



Figure 1. *a)* Seismicity at Tungurahua 2015-2016. Red bars show daily event counts, blue dots mark daily explosions, black line shows cumulative seismicity, grey shading marks period of interest for this study. Top panel marks surface observations - blue shows known episodes of drumbeats, grey shows ash ejection and orange are sightings of incandescent glow in crater. **b)** Accelerating drumbeats, 25 November. *c)* Decelerating drumbeats, 25 November. *d)* Penultimate day of drumbeats, 29 November. All 6hr extracts from RETU.

$$Q = \frac{f}{2g} \tag{1}$$

We determine the maximum cross correlation coefficient between 0 and 1, for all pairs of events in our catalogue and use a threshold value to group events into families (Waite et al., 2008; Yukutake et al., 2017; Park et al., 2019). Following previous studies, including that of LP drumbeat seismicity in April 2015 at Tungurahua, the threshold is set at 0.7 (Petersen, 2007; Thelen et al., 2011; Bell et al., 2017).

We also calculate earthquake inter-event times (IETs) and their periodicity to highlight times of pronounced drumbeat activity. (Bell et al., 2017) defines periodicity as the ratio between the mean, , and standard deviation, , of IETs. Events randomly distributed in time, with an average rate, , will have a probability density function such that - = 1. Clustered events have periodicities less than 1, whereas periodic events have periodicities greater than 1.

Finally we considered models for the accelerating and decelerating components the drumbeat episode. Previous studies of accelerating seismicity have modelled rates using power, exponential and hyperbolic relationships (Ignatieva et al., 2018; Bell et al., 2018). In accelerating and decelerating components we opted to model the event rates using an exponential relation (Eq 2) and a power law. For the decelerating event rates this is the Modified Omori's Law (Eq 3) and for the accelerating rates, an Inverse Omori's Law (Eq 4).

$$n(t) = ke^{-t} \tag{2}$$

$$n(t) = k_1(t \quad t_f) \quad p_1 \tag{3}$$

$$n(t) = k_2(t_f \quad t)^{-p_2} \tag{4}$$

¹⁷⁹ We model the drumbeat sequence as an inhomogeneous Gamma process (Bell et ¹⁸⁰ al., 2018). We define the point of maximum seismicity as, t_0 , separating the accelerat-¹⁸¹ ing and decelerating components. We use a Bayesian approach with PyMC3 implemen-¹⁸² tation (Salvatier et al., 2016). We use Markov Chain Monte Carlo (MCMC) to sample ¹⁸³ the posterior distributions of model parameters. We run 5000 iterations. We provide ini-¹⁸⁴ tial estimates for parameters p, t_f, k and \cdot . The prior distribution and rate parameters ¹⁸⁵ used are detailed in table S1.

186 **3 Results**

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Across the six day period from 25-30 November we see an initial increase in the rate of seismicity before a rapid deceleration (fig 3). The peak in event rate occurs at 10:00 on 25 November (fig 2). During the drumbeat episode the radial tilt increases and decreases in a range of 10 rad.

3.1 Drumbeat Onset

The first 10 hours of the drumbeat sequence is markedly different from the activity observed thereafter. In the first 10 hours, the event rate increases, the individual event amplitudes increase slightly and the seismicity becomes increasingly periodic (fig 2). The point process modelling shows the accelerating rates of seismicity can be defined by a power law (fig 3a). The best fitting exponent, p = 0.96 0.51. The best fit value for