

1 **Composition of natural phytoplankton community has minor effects**
2 **on autochthonous dissolved organic matter characteristics**

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1 **Abstract**

2 Dissolved organic matter (DOM) is an important component of nutrient cycling,
3 but the role of different organisms controlling the processing of autochthonous
4 DOM remains poorly understood. Aiming to characterize phytoplankton-derived
5 DOM and the effects of complex pelagic communities on its dynamics, we
6 incubated natural plankton communities from a temperate mesohaline estuary
7 under controlled conditions for 18 days. The incubations were carried out in
8 contrasting seasons (spring and autumn) and changes in the planktonic
9 community (phytoplankton, bacteria and microzooplankton), nutrients and DOM
10 were assessed. Our results highlight the complexity of DOM production and fate
11 in natural planktonic communities. Small changes in DOM composition were
12 observed in the experiments relative to the orders-of-magnitude variations
13 experienced in the phytoplankton assembly. We argue that the tight coupling
14 between microbial processing and DOM production by phytoplankton and
15 grazers stabilizes variations in quantity and characteristics of autochthonous
16 DOM, resulting in apparently homogeneous semi-labile DOM pool throughout
17 the experiments. However, seasonal differences in the production and processing
18 of DOM were observed, reflecting differences in the nutrient regimes and initial
19 DOM characteristics in each experiment, but also likely influenced by changes in
20 the successional status of the pelagic community. Acknowledging that
21 characteristics of the DOM derived from phytoplankton growth can vary broadly,
22 heterotrophic processing and successional status of the community are
23 synergistically important factors for shaping those characteristics, and thus
24 affecting the seasonal signature of the semi-labile autochthonous DOM pool.

25 Key words: Pelagic food web; succession; autochthonous DOM

26 **Introduction**

27 Dissolved organic matter (DOM) is an important component of carbon, nitrogen and
28 phosphorus cycling in aquatic systems, operating as storage (Hedges 2002; Jiao et al.
29 2010) and fuelling heterotrophic organisms through the microbial loop (Azam et al.
30 1983; Ferrier-Pages & Rassoulzadegan 1994; Cotner & Biddanda 2002). Primary
31 producers have been proposed as an important DOM source in marine ecosystems,
32 notably in those with limited freshwater influence and inputs of allochthonous DOM
33 (Suksomjit et al. 2009), influencing the DOM composition of surface waters (Biddanda
34 & Benner 1997). Phytoplankton can produce DOM (Thornton 2014), with quantity and
35 characteristics varying with nutrient availability (Mykkestad 1995), phytoplankton
36 species composition (Biddanda & Benner 1997), and bacterial interaction (Ramanan et
37 al. 2016). From studies with axenic phytoplankton cultures, different taxonomic groups
38 have been found to release different types of DOM (Romera-Castillo et al. 2010;
39 Fukuzaki et al. 2014). Additionally, other biological drivers need to be considered to
40 understand environmental DOM dynamics. For example, grazers can also produce
41 DOM (Strom et al. 1997; Ferrier-Pagès et al. 1998), whereas bacteria consume and alter
42 DOM properties (Rochelle-Newall & Fisher 2002; Romera-Castillo et al. 2011; Kinsey
43 et al. 2018). This complex array of biological processes add up to abiotic ones (e.g.
44 photo degradation), resulting in a DOM pool that becomes less and less reactive
45 (Hansell 2013). Thus, labile DOM is fast transformed into recalcitrant DOM, which
46 encompasses different fractions, whose turnover rates becomes increasingly slower: the
47 semi-labile, semi-refractory, refractory and ultra-refractory fractions (Hansell 2013).
48 As a portion of the DOM absorbs light at UV and visible wavelengths (chromophoric
49 DOM – CDOM), and a fraction of CDOM is fluorescent (FDOM), assessment of the
50 optical properties of the CDOM enables a relatively easy and inexpensive way to

51 characterize DOM in the environment (Coble 2007; Fellman et al. 2010). Even if the
52 DOM optical properties lack information on the specific molecular structure, the optical
53 properties constitute a proxy for the composition, sources and DOM processing degree
54 (Stubbins et al. 2014). In coastal zones, where the terrestrial influence overlaps intrinsic
55 processes, characterization of DOM sources and cycling can be even more challenging
56 due to high environmental heterogeneity of the pelagic coastal food web (Yamashita et
57 al. 2008). On top of that, DOM exhibits seasonal dynamics, which is partly linked to
58 biological activity (Markager et al. 2011; Knudsen-Leerbeck et al. 2017). Thus, there is
59 a need to advance from describing autochthonous DOM variability to actually
60 understand the underlying processes and mediators (Markager et al. 2011).

61 This is the second part of a study aimed at characterizing transformations in quantity
62 and characteristics of phytoplankton-derived DOM. Asmala, Haraguchi, Jakobsen, et al.
63 (2018) demonstrated that DOM originating from phytoplankton is rapidly processed and
64 its optical properties are continuously transformed, with dynamics differing strongly
65 over seasons. Phytoplankton composition appeared to have a minor role as driver for the
66 observed DOM dynamics regarding season and inorganic nutrient availability, as the
67 same taxonomical groups (cryptophytes and diatoms) dominated in both seasons
68 (Asmala, Haraguchi, Jakobsen, et al. 2018). Additionally, variations in the DOM
69 quantity were much smaller and were not directly related to variability in phytoplankton
70 biomass (Asmala, Haraguchi, Jakobsen, et al. 2018). The results from Asmala,
71 Haraguchi, Jakobsen, et al. (2018) led us to hypothesize that even if phytoplankton is
72 the primary DOM source, ecological processes, such grazing and succession, are more
73 relevant for the signatures of the semi-labile pool than bulk phytoplankton dynamics.

74 In this study, we examined the role of plankton community characteristics beyond major
75 phytoplankton groups in order to reveal potential community-driven mechanisms in

76 DOM transformation, in contrast to the more numerous studies exploring DOM
77 dynamics in phytoplankton monocultures. Our main objective is to evaluate how
78 temporal changes in plankton community structure influence the DOM transformations
79 observed by Asmala, Haraguchi, Jakobsen, et al. (2018). Specifically, we want to
80 evaluate potential effects of changes in community structure in the DOM dynamics at
81 short-term across two different seasons.

82 **Materials & methods**

83 *Experiment setup*

84 Roskilde Fjord (RF) is a temperate mesohaline estuary with low freshwater input, albeit
85 rich in nutrients, and it is previously reported to have strong signatures of
86 autochthonous DOM (Knudsen-Leerbeck et al. 2017). The experimental setup was
87 described in detail in Asmala, Haraguchi, Jakobsen, et al. (2018). In summary, surface
88 water (about 80 L) was sampled in the inner basin of RF, screened with a 100 μm mesh
89 to remove large zooplankton, and then transferred within one hour from sampling to six
90 glass jars, each containing 10 L. In three of these experimental units, nitrate (NaNO_3)
91 was added daily (days 0-3), totaling $12 \mu\text{mol L}^{-1}$ nitrate addition. No additions were
92 made to the remaining three control units. The bottles were incubated for 18 days at 10
93 $^{\circ}\text{C}$ and with controlled light (16:8 light: dark cycle, at $100\text{--}120 \mu\text{mol m}^{-2} \text{s}^{-1}$) and
94 constantly stirred with a teflon-coated magnetic bar. The experiment was conducted two
95 times representing contrasting seasons in RF: autumn (from 14 September to 2 October
96 2015) and spring (14 March to 1 April 2016).

97 *Sampling strategy:*

98 We employed an adaptive sampling strategy in order to follow daily changes in the
99 plankton community and capture significant alterations in nutrients and DOM over the

100 entire incubation. Samples (100 mL) were taken daily for flow cytometry and FlowCam
101 analysis. Unfortunately, phytoplankton analysis with the flow cytometer was terminated
102 after day 9 in the autumn experiment, due to technical problems with the instrument. *In*
103 *vivo* fluorescence (ex. 470 / em. 685 nm) of the samples was evaluated daily using a
104 Varian Cary Eclipse fluorometer (Agilent). When major changes were detected in the
105 fluorescence, a more comprehensive sampling was conducted: at days 0, 4, 7, 11, and
106 18 (spring) and days 0, 4, 8, 14, and 18 (autumn). Comprehensive sampling involved
107 phytoplankton counts by inverted microscopy, chlorophyll *a* extraction (Chla), bacteria
108 abundances, dissolved organic carbon (DOC), chromophoric dissolved organic matter
109 (CDOM), and nutrient concentrations.

110

111 ***Laboratory analyses***

112 *Chlorophyll a:*

113 Chla was determined according to the method described by Strickland and Parsons
114 (Strickland & Parsons 1972), following the extraction protocol of Holm-Hansen and
115 Riemann (1978). The extracts were kept at -20°C until measured with an AU 10 Turner
116 field fluorometer (Turner Designs, US). Extracted Chla values were used to validate
117 estimated values from *in vivo* fluorescence.

118 *Flow cytometry:*

119 We employed a pulse-shape recording CytoSense flow cytometer (CytoBuoy.com) to
120 analyse phytoplankton. This technique provides phytoplankton counts (cell sizes 1-1000
121 µm) comparable to those obtained with traditional microscopy, although with more
122 reliable counts for cells < 5 µm (Haraguchi et al. 2017). Additionally, it also provides

123 information on cell size and morphology due to its capacity to store the optical profile
124 for each particle, recorded as they travel through the flow cell. The instrument has a 488
125 nm laser, fluorescence sensors (yellow/green – 550 nm, orange - 600-650 nm, red - 650-
126 700 nm) and two scatter sensors, for light scattered parallel (forward scatter) and
127 orthogonal (sideward scatter) to the incident laser beam. Optical particle profiles from
128 live samples (500 μL – 1000 μL , sampled at a flow rate of 8 $\mu\text{L s}^{-1}$) were collected
129 using the software CytoUSB (CytoBuoy.com), with a threshold of 30 mV for the high
130 sensitivity red fluorescence sensor. This trigger was set to include only particles
131 containing Chla (phytoplankton cells). Recorded cells were clustered according to
132 similarities in their optical properties (length, total Forward Scatter (FWS), total red
133 fluorescence (FLR), total orange fluorescence (FLO), and total Sideward Scatter
134 (SWS)), using the software CytoClus3 (CytoBuoy.com). Particles were assigned to one
135 cluster only and the same clustering algorithm was employed for all samples.
136 Taxonomical information was obtained for some of the clusters based on their optical
137 characteristics, pictures taken by the equipment and cross-referenced with microscopy.
138 Carbon biomass was obtained by converting total FWS to volume (Haraguchi et al.
139 2017) and then converting volume to biomass using a generic protist volume-to-carbon
140 conversion formula (Menden-Deuer & Lessard 2000). In order to assess phytoplankton
141 physiological state, we estimated the carbon-to-chlorophyll ratio (C:Chla) of each
142 cluster of the phytoplankton community for all samples separately. For this, total carbon
143 biomass was divided by total Chla, which was estimated by converting the total red
144 fluorescence to Chla concentration using an empirical formula proposed by Haraguchi
145 et al. (2017), for the same location where the inoculum were taken. It needs to be
146 emphasised that the carbon and the Chla estimated by CytoSense differs from the
147 carbon and Chla estimated by microscopy and organic solvent extraction method

148 respectively. Carbon estimates obtained with CytoSense are more precise than
149 microscopy as they are obtained on an individual level (Haraguchi et al. 2017), while
150 CytoSense Chla is based on *in vivo* fluorescence intensity that is more susceptible to
151 physiological changes than Chla (Geider 1987 and references within). Yet, the C:Chla
152 ratio during the experiment is a powerful proxy to follow changes in phytoplankton
153 physiology.

154 *Bacteria:*

155 Bacteria were fixed in glutaraldehyde (1% final concentration) and stored at 4 °C
156 (autumn experiment) and in paraformaldehyde (1% final concentration) and stored at -
157 20 °C (spring) until analysis. For cell counting, bacteria samples were diluted in TE
158 buffer (0.1 M), 10 or 50 times depending on the bacteria concentration. The diluted
159 samples were stained with SYBR green (1:10000 final concentration, Marie et al. 1997)
160 and enumerated by the same equipment employed for the phytoplankton analysis,
161 although with a different set up (200-500 µL, flow rate 8 µL s⁻¹, trigger on 20 mV on
162 the high sensitivity yellow/green 550 nm fluorescence sensor). Heterotrophs were
163 distinguished from phytoplankton by the absence of red fluorescence signal.

164 *FlowCam:*

165 Rotifers and ciliate abundances and body volumes were analysed from live samples
166 using a color FlowCam IV equipped with a FC300 flow cell (Calbet et al. 2014).
167 Samples were kept in dim light at 12 °C and analysed *in vivo* within 4 hours after
168 sampling. The instrument was run in auto image-mode with 4x magnification capturing
169 all particles in the range 15 µm – 1000 µm. The analysis time for each sample was ca.
170 40 min., corresponding to an analysed volume of 20 mL. After processing the sample,
171 recorded images were manually sorted into ciliates and rotifers. Equivalent Spherical

172 Diameter (ESD) and body volume were estimated by the software package VISP 3.17
173 (FluidImagine™) using the area based diameter (ABD) algorithm of VISP 3.17
174 (Jakobsen & Carstensen 2011). Carbon biomass was obtained by converting volume to
175 biomass using a generic protist volume-to-carbon conversion formula (Menden-Deuer
176 & Lessard 2000).

177 *Microscopy:*

178 Microscopy counts in this study were only used to support taxonomical identification of
179 the main phytoplankton groups and were linked to the optical signatures of the most
180 important clusters in each experiment. Fixed samples (acidic Lugol's solution, 2-4%
181 final concentration) were analysed from 10-50 mL Utermöhl chambers (Utermöhl
182 1958), under a size-calibrated inverted microscope (Nikon TI-U, Nikon Instruments
183 Europe B.V.). Sedimented volumes varied depending on the cell concentration. Both
184 phytoplankton and ciliates were enumerated.

185 *Nutrients:*

186 We collected samples for total nutrients, dissolved inorganics, total and total dissolved
187 N and P. Total nutrients (total N; TN and total P; TP) were analysed from unfiltered
188 water samples, whereas dissolved total nutrients (total dissolved N, TDN and total
189 dissolved P, TDP) and inorganic nutrients (nitrite; NO_2^- , nitrate; NO_3^- , ammonium;
190 NH_4^+ , orthophosphate; PO_4^{3-} (DIP) and dissolved inorganic silicate Si) were measured
191 from filtered water samples using combusted GF/F.

192 Dissolved inorganic nutrient samples were stored frozen in 30 mL acid-washed plastic
193 bottles. The samples were analysed on a San ++ Continuous Flow Analyser (Skalar
194 Analytical B.V, Breda, NL) as previously described (Grasshof 1976; Kaas & Markager

195 1998). Detection limits were 0.04, 0.1, 0.3, 0.06 and 0.2 $\mu\text{mol L}^{-1}$ for NO_2^- , NO_3^- , NH_4^+ ,
196 PO_4^{3-} and Si, respectively.

197 Samples for total and total dissolved nitrogen (TN and TDN, respectively) and total and
198 total dissolved phosphorus (TP, TDP, respectively) measurements (20 mL) were
199 collected in 30 mL brown glass bottles filled with Milli-Q water prior to sampling. TN
200 and TP were determined by adding oxidants to the sample followed by autoclaving and
201 were analysed on a San ++ Continuous Flow Analyser (Skalar Analytical B.V, Breda,
202 NL). Detection limits for TN and TP were 1.0 $\mu\text{mol N L}^{-1}$ and 0.1 $\mu\text{mol P L}^{-1}$,
203 respectively.

204 Dissolved inorganic nitrogen (DIN) concentrations were calculated as the sum of the
205 concentrations of NO_2^- , NO_3^- and NH_4^+ . Dissolved organic nutrient concentrations
206 (DON and DOP) were calculated as the difference between total dissolved nutrient
207 (TDN and TDP) and dissolved inorganic nutrient (DIN and DIP).

208 *Dissolved organic carbon (DOC):*

209 DOC was measured with a Shimadzu TOC-VCPH analyser, and the accuracy of
210 measured DOC concentrations was controlled by analysing a seawater reference
211 standard provided by the CRM (consensus reference material) program.

212 *CDOM and FDOM:*

213 CDOM absorption was measured using a Shimadzu 2401PC spectrophotometer with 5
214 cm quartz cuvette over the spectral range from 200 to 800 nm with 1 nm intervals.
215 Milli-Q (Millipore) water was used as the blank for all samples. Excitation-emission
216 matrices (EEMs) of FDOM were measured with a Varian Cary Eclipse fluorometer
217 (Agilent). A blank sample of ultrapure water was removed from the EEMs, as well as
218 the scattering bands. EEMs were corrected for inner filter effects with absorbance

219 spectra (Murphy et al. 2010) and Raman calibrated by normalizing to the area under the
220 Raman scatter peak (excitation wavelength of 350 nm) of an ultrapure water sample run
221 on the same session as the samples. The carbon-specific absorbance wavelengths was
222 calculated by dividing the absorbance of a given wavelength λ by the DOC
223 concentration (Weishaar et al. 2003). Here we employed DOC-specific visible
224 absorbance at 440 nm (SVA₄₄₀), which can be used as an estimator of the proportion of
225 visible-absorbing molecules in the DOM pool. FDOM descriptors often used are peaks
226 at specific excitation/emission wavelengths (Coble's peaks), which can be related, for
227 example, to aromatic amino-acids (peak T) or to humic-like substances (peak C) (Coble
228 1996).

229 *Statistical analyses*

230 Measurements from the experiment were analysed with a linear mixed model that
231 described the effect of treatment over the course of the experiment. Variations in the
232 measurement variable or a log-transform of this (response variable X_{ijk}) were modelled
233 as

$$234 \quad X_{ijk} = t_i + d_j + t_i \times d_j + e_{ijk} \quad \text{Eq. (1)}$$

235 where t_i described the overall difference between control and nitrate addition, d_j
236 described the changes for each day, $t_i \times d_j$ described differences over experiment days
237 between the nitrate addition and control, and e_{ijk} described the residual variation among
238 experimental units. The residual error was modelled as an AR(1) process within each
239 experimental unit to account for potential autocorrelation (repeated measures design).
240 The mixed model was analysed separately for the two seasons. Measurements obtained
241 from the flow cytometer and FlowCam showed scale-dependent variations and were
242 log-transformed, whereas measurements of DIN, DIP, Chla and all DOM variables did

243 not display these tendencies and were assumed approximately normal distributed.
244 Changes in the response variable from one time point to another was calculated as
245 contrasts of parameter estimates of $t_i \times d_j$. The mixed model was analysed using PROC
246 MIXED in SAS version 9.3.
247 A principal component analysis (PCA) was carried out in order to simplify the
248 multidimensional information on community, facilitating dynamics visualization and
249 interpretation. Community data were log transformed prior to PCA analysis.

250 **Results**

251 *Initial conditions*

252 Initial inorganic nutrient concentrations for both spring and autumn experiments were
253 above thresholds considered to limit phytoplankton growth (2.0 μM for DIN and 0.2
254 μM for DIP), with a N:P ratio around 50 in spring and around 1 in autumn (Fig. 1). DIN
255 was mainly composed of $\text{NO}_2^- + \text{NO}_3^-$ in spring and NH_4^+ in autumn (Fig. 1a, b). Initial
256 Chla levels were substantially higher in spring than autumn (~ 13.5 and $1.8 \mu\text{g L}^{-1}$,
257 respectively; Fig. 2a, b). The phytoplankton community was dominated by cryptophytes
258 (*Teleaulax* spp.) in spring and by a mixed community of cryptophytes (*Teleaulax* spp.)
259 and dinoflagellates (*Heterocapsa* cf. *rotundata* and gymnodinioids) in autumn (Fig. 3).
260 The difference in phytoplankton biomass was paralleled by ciliate biomass of around 40
261 $\mu\text{g C L}^{-1}$ in spring and $8 \mu\text{g C L}^{-1}$ in autumn (Fig. 4a, b), with a resulting initial biomass
262 ratio (w/w) between phytoplankton and grazers (ciliates + rotifers) around 60 in spring
263 and 30 in autumn (Fig. 4e, f). Bacterial abundances were similar at the start of the two
264 experiments (Fig. 5), as were the initial planktonic community (Fig. 6). The DOM
265 variables also exhibited distinct characteristics: lower DOC, and peak T, but higher peak
266 C in spring than autumn (Fig. 7 & 8).

267 *Inorganic nutrient dynamics*

268 In spring, DIN and DIP were gradually consumed during the experiment course,
269 although at higher rates at the beginning and end of the experiment (Fig. 1a, c). In
270 autumn, NH_4^+ was the main species of DIN in the control units, and in the treatments,
271 the addition of NO_3^- led to a temporary shift around day 4 in the DIN speciation,
272 although rapidly shifting back to dominance of NH_4^+ during the following days (Fig.
273 1b). During the autumn experiment beginning, changes in the DIN were significant in
274 the controls (Table 1) due to the consumption on NH_4^+ that was the main inorganic N
275 source (Fig. 1b). Whereas, in the treatments, the NO_3^- additions balanced the NH_4^+
276 consumption (Fig. 1b), leading to a non-significant change in the resulting DIN during
277 A1 (Table 1). During both seasons, the addition of nitrate stimulated DIP consumption
278 (Fig. 1c, d).

279 *Successional dynamics*

280 Distinct succession phases, manifested by changes in the phytoplankton (composition,
281 biomass and physiological state) and grazer communities, were used to initiate
282 comprehensive sampling during both experiments. These phases were categorized into
283 four types per season, taking into consideration the succession patterns at short-term and
284 seasonal (spring and autumn) scales. The eight phases are described separately and
285 changes in each are summarized in Table 1 and Table SI.

286 Phase S1 – Expansion (Spring, day 0-4): This phase was characterized by increasing
287 abundance of phytoplankton, dominated by cryptophytes *Teleaulax* spp. (> 80% of the
288 total biomass) and exhibiting a good physiological status, with the lowest C:Chl a ratios
289 of the experiment (Fig. 2c). Ciliate biomass also increased fast during this phase (Fig.
290 4a) with a daily biomass increase of 23 % in the controls and 51% in treatments (Table
291 SI), as did bacterial abundances (Fig. 5a) that almost doubled each day (Table SI).
292 Biomass ratio of phytoplankton to grazers ($r_{P:G}$) decreased from 60 to ~15 (Fig. 4e).

293 Phase S2 – Maturity (Spring, day 5-7): In this phase, phytoplankton biomass peaked
294 and was still dominated by cryptophytes (~70% of the total biomass) (Fig. 3a).
295 However, an increase in the C:Chla ratio was observed (Fig. 2c), with a daily increase
296 of 25% in controls and 19% in treatments (Table 1), indicating a slowing down of the
297 expansive growth. Ciliate biomass remained constant (Fig. 4a), resulting in only minor
298 changes in the $r_{P:G}$ ratio (Fig. 4e). Bacterial abundances declined to levels similar to the
299 initial conditions (Fig. 5a).

300 Phase S3 – Senescence (Spring, day 8-11): Phytoplankton was still dominated by
301 cryptophytes, yet their biomass was declining ~40% per day (Fig. 3a; Table SI) and
302 diatoms (unidentified pennate diatoms and *Skeletonema* sp.) started to increase in
303 treatments (Fig. 3c). No significant changes were observed for C:Chla, which remained
304 at a high level (Fig. 2c). On the other hand, ciliate biomass declined (Fig. 4a, Table SI),
305 and rotifers started to increase in the treatments (Fig. 4a, c); yet, no change in the $r_{P:G}$
306 ratio was observed (Fig. 4e). Bacterial abundance increased again, more than two-fold
307 (Fig. 5a).

308 Phase S4 - Community shift (Spring, day 12-18): Shifts in dominance of phytoplankton
309 (cryptophytes to diatoms, Fig. 3a, c) and grazers (ciliates to rotifers, Fig. 4a, c) were
310 observed. Diatoms increased their biomass until day 14, and then they started to decline
311 (Fig. 3c), reflecting a net increase in C:Chla ratio (Fig. 2c). However, an increase in
312 Chla was observed, which most likely was associated with increasing biomass of
313 miscellaneous nanoflagellates, which resulted in an increase in the $r_{P:G}$ ratio (Fig. 4e).
314 During this phase, experimental units started to behave more erratically, with a large
315 variability in organisms concentrations. Bacterial abundances in the water column
316 continued to increase (Fig. 5a).

317 Phase A1 – Expansion 2 (Autumn, day 0-4 (control) and 0-5 (treatment)):
318 Phytoplankton biomass increased (Fig. 3b) due to growth of cryptophytes (*Teleaulax*
319 spp.) and dinoflagellates (*Heterocapsa* cf. *rotundata* + gymnodinioids). In the controls,
320 phytoplankton grew until day 4 and then declined, while in the treatments growth
321 continued onto day 5. The C:Chla ratio remained relatively constant, but started to
322 increase in the controls from day 3 (Fig. 2d). Ciliates grew well in all units, reaching
323 maximum biomasses at day 4 (controls) or day 5 (treatments), resulting in a decrease in
324 the $r_{P:G}$ ratio (Figs. 4b, 4f). Changes in bacteria could not be assessed due to lack of
325 initial data (Fig. 5b).

326 Phase A2 – Rapid community shift (Autumn, day 4-8 (control) and 5-8 (treatment)): In
327 contrast to spring, the intermediate phases from expansion to community shift were not
328 observed in autumn. Cryptophytes (Fig. 3b) continued the decline from the expansion
329 phase in the controls, as opposed to the increase in diatoms (mainly *Skeletonema* sp. and
330 a small unidentified centric) from day 4 that was observed in both controls and
331 treatments (Fig. 3d). The C:Chla ratio first increased (drastically for control units) and
332 decreased again after day 5, resulting in a net increase of 18% during this phase and
333 reaching similar levels for treatments and controls at day 7 (Fig. 2d, Table 1). Ciliates
334 declined drastically, with one day delay between control and treatment units (Fig. 4b),
335 whereas rotifers increased rapidly (Fig. 4d). The $r_{P:G}$ ratio remained constant below 10
336 at first, but rose towards the end of the phase when grazers collapsed (Fig. 4f). Bacterial
337 abundances increased during this period (Fig. 5b).

338 Phase A3 – Destabilization (Autumn, day 9-14): No high-resolution data from
339 phytoplankton were available after day nine; however, the decline in phytoplankton was
340 confirmed by Chla values and microscopy counts (data not shown). Grazers and
341 bacteria also declined (Fig. 4b, 4d, 5b).

342 Phase A4– Regeneration (Autumn, day 15-18): Phytoplankton and grazers were nearly
343 absent. Bacteria abundances decreased in controls (-24% per day), but not in treatments
344 (Fig. 5b, Table SI).

345 The dynamics of the microbial community were also captured by the PCA, with similar
346 succession patterns being observed in both experiments (Fig. 6). An initial cryptophyte-
347 dominated assembly shifted to diatom-domination. A similar shift was observed for the
348 grazers, with ciliates associated with cryptophytes and rotifers occurring with diatoms
349 (Fig. 6). In spring, the treatments were very similar, while in autumn nitrate additions
350 induced a large change in the biomass (Fig. 2a, b, 3), but smaller changes in the
351 community structure (Fig. 6). Except for the later phase in spring (S4), treatments and
352 seasons exhibited similar trajectories over time. However, differences in the processing
353 velocity were evident from the PCA, with larger distance between days in autumn than
354 in spring, especially during the initial phases (Fig. 6). Unfortunately, data after day 9
355 were not available for the phytoplankton community in autumn, but Chla data indicated
356 that phytoplankton collapsed after day 9.

357 *DOM dynamics*

358 In spring, two distinct periods could be identified for the DOM pool transformation.
359 One was observed at the initial phases (S1 and S2) and marked by the accumulation of
360 DOC and DON (Fig. 7a, c). In phase S1, an increase in protein-like DOM (peak T) was
361 observed at the same time as reduction of humic-like (peak C), followed by
362 accumulation of peak C in phase S2 (Fig. 8c, e). The second period was marked by the
363 accumulation of CDOM (SVA₄₄₀) during phase S3 (treatment units) and S4 (controls)
364 (Fig. 8a). Although DOP was consumed throughout the incubations, differences in the
365 consumption pattern between treatments were observed, with consumption stagnating in
366 phase S2 (controls) or S3 (treatments).

367 In autumn, initial changes in the DOM pool were marked by the decline in peak T
368 during the initial phase A1 (Fig. 8f). Significant changes were noticed later in the
369 experiment: increase in DOC (during A3) and peak C (during A4), along with decrease
370 in peak T (during A4) and CDOM (SVA₄₄₀) (during both A3 and A4) (Fig. 7b, 8b, d, f).

371 **Discussion**

372 Coastal ecosystems display seasonal variations in CDOM and DOC associated with
373 phytoplankton spring blooms, but the dynamics and magnitude of CDOM and DOC
374 vary between years and are hypothesized to depend on bloom composition and
375 interaction with heterotrophs (Minor et al. 2006; Suksomjit et al. 2009). Thus,
376 production of autochthonous DOM can be important in estuarine and coastal areas, but
377 the processes and drivers behind autochthonous DOM seasonal variability remains to be
378 further studied (Markager et al. 2011). One major obstacle impeding our understanding
379 of these processes has been the lack of studies with high-frequency sampling for
380 assessing the role of different communities in the coastal ecosystems (for example
381 plankton, as shown in this study) together with changes in the DOM pool, particularly
382 over different seasons. The daily sampling frequency in this work highlights the fast
383 dynamics and tight coupling between phytoplankton and grazers. The observed
384 dynamics in the plankton community would not have been properly described with
385 lower sampling frequency, suggesting that at least daily sampling is needed in
386 experiments with complex natural plankton communities. DOM and nutrients were
387 sampled only at five occasions, as these constituents were considered less variable.
388 Nevertheless, we recognise that our sampling scheme was inadequate to trace the
389 dynamics of the labile DOM pool, which can be processed by bacteria over time spans
390 of hours (Fuhrman & Ferguson 1986). Thus, the DOM variables analysed in this study
391 most likely represent the dynamics of the semi-labile pool, but considering that most of

392 DOM in marine surface environments can be regarded as semi-labile (Carlson 2002),
393 we believe that our results are representative of natural conditions.

394 The study site Roskilde Fjord (RF) is a shallow and microtidal temperate estuary, with
395 long freshwater residence time, high nutrient inputs and low freshwater discharge
396 (Kamp-Nielsen 1992; Staehr et al. 2017). Because of those characteristics, RF is an
397 ideal environment for investigating the biogeochemical processing of DOM, since
398 disturbances from physical transport are low and nutrients and carbon are primarily
399 processed within the system (Asmala, Haraguchi, Markager, et al. 2018). RF has
400 contrasting characteristics between spring and autumn, changing from net autotrophic,
401 DIN-rich and DIP-poor in spring to net heterotrophic, DIN-poor and DIP-rich in autumn
402 (Staehr et al. 2017). Concentrations and composition of organic C, N and P pools in RF
403 change substantially over the seasonal scale, displaying a gradual decrease in
404 bioavailable DOC from spring to autumn, whereas the proportions of bioavailable DON
405 and DOP are more variable and possibly related to occurrences of higher phytoplankton
406 biomass (Knudsen-Leerbeck et al. 2017). This highlights the importance of
407 phytoplankton in regulating the composition of DOM associated with different forms of
408 organic N and P. In our experiments, the initial DOM pool had contrasting
409 characteristics between seasons: spring DOM had more allochthonous characteristics
410 and autumn DOM was more autochthonous, most likely related to the higher freshwater
411 inputs during early spring and further processing of DOM over summer (Asmala,
412 Haraguchi, Jakobsen, et al. 2018). The spring bloom in RF is fuelled by inorganic
413 nutrients, mainly from land, whereas summer and autumn production is mainly
414 sustained by nutrient inputs from sediments and occasional intrusions of deeper waters
415 from the Kattegat (Knudsen-Leerbeck et al. 2017; Staehr et al. 2017).

416 ***Phytoplankton derived DOM signature: physiological status or species composition?***

417 Phytoplankton grown under replete light and nutrient conditions with constant
418 temperature maintain a constant optimum Chla quota per cell, whereas changing light,
419 temperature and nutrient limitation can lead to substantial variations in cellular Chla
420 (Geider 1987 and references within). Cellular decreasing Chla, or the change in the
421 specific cell ratio in C:Chla, is therefore a proxy for the physiological status of the
422 phytoplankton community and is reflecting the balance between carbon fixation and
423 other growth processes (Fig 2c, d). Under nutrient deplete conditions, the Chla per cell
424 is reduced in order to lower the accumulation of solar energy within the cell to avoid
425 photochemical damage. Yet the cell does not entirely stop photosynthesis, and the cell
426 needs to store the solar energy as photosynthetic products; either as fatty acids or release
427 them to the water as DOM. The excess DOM is either passively diffused or excreted
428 actively, with the dominant process depending on the physiological status of the
429 organism and species (Thornton 2014). The release of DOM by phytoplankton can be
430 enhanced by excess levels of light (Cherrier et al. 2015) and nutrient limitation, with
431 enhanced release of simple carbohydrates under N or P limitation (Mykkestad 1995;
432 Biddanda & Benner 1997). Nutrient status also influences the bioavailability of released
433 DOM compounds, with DOM originating from phytoplankton grown under nutrient-
434 deplete conditions being less bioavailable for bacteria (Obernosterer & Herndl 1995;
435 Puddu et al. 2003).

436 It has been demonstrated that different phytoplankton species, growing under axenic
437 conditions, produce DOM with distinct quantity and quality, and that DOM production
438 is correlated to phytoplankton abundance (Romera-Castillo et al. 2010; Fukuzaki et al.
439 2014). In our experiments, similarity between seasons was observed for phytoplankton,
440 with dominance of cryptophytes (*Teleaulax* spp.) in the initial phase, followed by a
441 transition phase to a final diatom-dominated community after nutrient depletion set in

442 (P in spring and N in autumn). Interestingly, no difference in the bulk DOM and CDOM
443 was observed between phases dominated by different taxonomical groups in any of our
444 experiments. Furthermore, variability in the DOC concentrations was modest in
445 comparison with biomass changes experienced by phytoplankton, and DOC and
446 phytoplankton biomass dynamics were decoupled. Hence, phytoplankton composition
447 and biomass appear to have a secondary role in structuring the bulk and optical DOM
448 characteristics of the semi-labile pool in a complex community.

449 In spring, during phase S1, cryptophytes dominated the phytoplankton biomass, while
450 an increase in DON and protein-like FDOM (peak T) was observed. During this phase,
451 cryptophytes notably increased their biomass and exhibited good physiological state,
452 indicated by the low C:Chl_a ratios. Peak T can be used as a proxy for fresh
453 autochthonous DOM (Coble 2007; Hansen et al. 2016), and the exudation of
454 proteinaceous compounds by healthy phytoplankton associated with peak T signal has
455 been described by Romera-Castillo et al. (2010). Bacterial peak T production was also
456 reported, with the largest fraction occurring within the cells rather than in the
457 extracellular phase (Fox et al. 2017). Considering that our FDOM measurements were
458 derived from gently filtered samples, it is reasonable to assume that most of peak T in
459 the experiments was from the dissolved pool. Production rates of FDOM are reported to
460 be higher for bacteria than phytoplankton (Romera-Castillo et al. 2011); yet, in our
461 experiments the increase in peak T was only observed during S1, in contrast to bacteria
462 increases that were observed in other phases (S3, S4 and A2). Thus, the increase in peak
463 T during the spring expansion (S1) phase is most likely linked to phytoplankton growth,
464 either directly or mediated through bacteria. The high initial peak C values indicated
465 allochthonous/terrestrial DOM, which was more labile and readily consumable within
466 the first four days of the experiment (Asmala, Haraguchi, Jakobsen, et al. 2018), as

467 evidenced by the sharp decline of this variable. This was followed by an increase in
468 peak C of almost similar magnitude between day 4 and 7 (phase S2), which in our
469 closed experimental units could only be associated with autochthonous production.
470 Peaks with optical signatures similar to terrestrial sources (peak A and C) are found to
471 be produced by marine phytoplankton grown under axenic conditions (Romera-Castillo
472 et al. 2010; Fukuzaki et al. 2014) and, notably, by bacteria growing on phytoplankton
473 exudates (Romera-Castillo et al. 2011; Kinsey et al. 2018).

474 We suggest that this accumulation of humic-like substances was more likely derived
475 from bacterial processing of phytoplankton exudates (peak T). Thus, the balance
476 between peak T production and consumption, and autochthonous peak C accumulation
477 appears to be related to the physiological state of the phytoplankton cells and not only
478 the biomass of net autotrophic communities.

479 In the autumn experiment, even though cryptophytes were growing and accounted for
480 ~50% of the phytoplankton biomass in the first phase, a decrease in peak T was
481 observed. We attribute the opposite behaviour of peak T between seasons to differences
482 in the overall metabolic state of the community, which is net autotrophic in spring and
483 net heterotrophic in autumn (Staehr et al. 2017). Asmala, Haraguchi, Jakobsen, et al.
484 (2018) pointed out that differences in the limitation patterns between seasons (P-limited
485 spring and N-limited autumn) also might influence the bacterial efficiency to degrade
486 DOM, with P-surplus in autumn boosting the bacterial capacity to utilize DOM in
487 comparison to spring. Thus, while in spring DIN:DIP ratios were high and production
488 exceeded consumption, a net accumulation of DON and peak T was observed; in
489 autumn, when DIN:DIP was low and heterotrophy predominant, peak T was readily
490 consumed, even if produced, resulting in a net decrease. Those findings suggest that
491 phytoplankton physiology might be more relevant than species composition for the

492 specific DOM production, whereas the role of phytoplankton for governing the DOM
493 pool is small when other community components are present. This means that the DOM
494 pool in the environment is regulated in synergy by the entire food web / community
495 structure, and not just by the primary producers.

496 ***Role of heterotrophs***

497 Although DOM production might be controlled by phytoplankton, the bacterial
498 community is essential in regulating DOM quantities and characteristics in the
499 environment (Guillemette & del Giorgio 2012). Previous studies have demonstrated that
500 bacterial composition and activity are tightly coupled to primary production and that
501 bacterial communities quickly adapt to efficiently use phytoplankton-derived DOM,
502 resulting in modest changes in the DOM pool (Sarmiento & Gasol 2012; Landa et al.
503 2016; Hoikkala et al. 2016; Luria et al. 2017). In both experiments the proportion of
504 bioavailable DOC was reported to be low, indicating tight food-web coupling and fast
505 bacterial processing of freshly produced DOM in both seasons (Asmala, Haraguchi,
506 Jakobsen, et al. 2018). Thus, we argue that our results display the net effect of
507 planktonic communities on autochthonous DOM dynamics, despite the fact we lack
508 detailed information on bacterial community composition and activity.

509 In our experiments, changes in the bacterial abundances followed the Chla dynamic in
510 spring, whereas such correspondence was less evident in autumn, indicating a tighter
511 coupling between bacteria and photosynthesis in spring and the importance of other
512 DOM sources in autumn. The biomass proportion between autotrophs and heterotrophs
513 can be used as a proxy for community structure, varying according to resources
514 availability and relative turnover rates of autotrophs and heterotrophs (Gasol et al.
515 1997). The proportion of autotrophs to grazers ($r_{P:G}$) (Fig. 4e, f) reflected differences in
516 the initial community structure of each experiment, and here we used it as a proxy for

517 the biomass between auto- and heterotrophs. Differences in $r_{P:G}$ indicate that spring was
518 more autotrophic with slower turnover rates than autumn, further supported by the fast
519 successional transition in the later (Fig. 6). This was also evidenced by the dynamic of
520 the main inorganic nitrogen species in each experiment, with a large contribution of
521 NH_4^+ to DIN indicating the importance of heterotrophic regeneration processes,
522 especially in autumn.

523 Although it was not possible to distinguish the effects of grazing from other processes
524 in our experiments, we argue that grazing was essential to fuel the microbial loop,
525 especially during autumn, and likely contributed to the DOM pool dynamic. In addition,
526 the intense ciliate grazing on cryptophytes, probably opened a niche for diatom growth,
527 as ciliates preferentially graze on nanoflagellates rather than large and colonial diatoms
528 (Kivi & Setälä 1995; Granéli & Turner 2002; Sommer et al. 2005). It is not possible to
529 conclude whether the shift observed in phytoplankton community in both experiments
530 was driven by nutrient limitation or grazing, and most likely it was due to a combination
531 of both. Interestingly, diatoms grew after inorganic nutrient depletion (phase S4 and
532 A2), which could be due to association with bacteria, remineralising nutrients in the
533 surrounding phycosphere around the diatom cells (Amin et al. 2012). Although no
534 significant effects of phytoplankton composition was observed in our experiments, the
535 increased contribution of diatoms likely marked the shift from water column processes
536 to processes related to surfaces (i.e. suspended particles, detritus, marine snow). In
537 mature communities the substrates become more important and particle-attached life
538 strategies become dominant, as the external nutrient inputs are limited and recycled
539 nutrients drive the system production (Wetzel 1995). Thus, the faster development of
540 the autumn experiment and the occurrence of the later phases (destabilisation and
541 regeneration, phase A3 and A4, respectively) probably reflect the higher dependency of

542 the mature communities on nutrient recycling, and therefore on heterotrophic and
543 microbial nutrient processing. Those mature stages were not reached during the spring
544 experiments, due to the slower turnover rates.

545 Differences in the autotrophs, heterotrophs and $r_{P:G}$ at the beginning and during each
546 experiment were not reflected to the same extent in the DOM pool. This can be
547 attributed to the intricate dynamic between nutrients and the pelagic community,
548 resulting in DOM that is constantly produced, but differing by the dominant origin
549 process (Kujawinski 2011). This can be exemplified for the mesohaline Chesapeake
550 Bay, where similar rates of DON release over different seasons were observed, although
551 DON was mainly released from autotrophs in May switching to be released by
552 heterotrophs in October (Bronk et al. 1998). Similar to this study, our experiments
553 showed that nitrate-rich and -poor conditions likely favoured DOM production directly
554 from phytoplankton and mediated by grazers, respectively. Hence, our results suggest
555 that nutrient conditions are the main controlling factor for the balance between
556 autotrophic and heterotrophic processes of DOM production; however, these processes
557 tend to produce DOM with similar bulk and optical characteristics, giving the
558 impression of DOM resilience. This suggests that the semi-labile DOM pool measured
559 with our methods is relatively resilient to changes in phytoplankton composition,
560 biomass and physiological state. As parts of the DOM pool are very reactive and can
561 change over time scales of hours (Kirchman et al. 1991; Obernosterer et al. 2008), the
562 fast responses of the microbial and phytoplankton communities seem to stabilise DOM
563 optical characteristics. This implies that tightly coupled microbial processing
564 counterbalances large changes in the phytoplankton community and, most likely, the
565 associated production of DOM. On the other hand, the resilient DOM pool might

566 change slowly over the season resulting in a seasonal signature that reflects the
567 gradually changing composition and status of the community.

568 ***Plankton succession and the reactivity of autochthonous DOM***

569 Characteristics of marine DOM are a result from its origin and processing history,
570 comprising a mixed pool of molecules, with different degree of availability to bacteria
571 and other organisms. As labile autochthonous DOM is rapidly consumed, the resulting
572 DOM pool is increasingly recalcitrant as the heterotrophic degradation continues
573 (Benner 2002). Although most of the DOM components in our experiments are
574 considered somehow resistant to biological transformation, the experiments exhibited
575 changes in DOM characteristics with divergent trajectories over the experiments,
576 indicating that DOM was continuously processed (Asmala, Haraguchi, Jakobsen, et al.
577 2018). In spring, most changes occurred during the first half of the experiment, while in
578 autumn most of the changes were observed at the end of the experiment (Asmala,
579 Haraguchi, Jakobsen, et al. 2018). In this study, we show that those changes in the
580 DOM pool were related to phytoplankton production (spring) and remineralisation
581 (autumn). The role of successional changes in our experiments was further supported by
582 the net increase of DOC concentrations in both experiments, accompanied by
583 contrasting patterns between seasons for CDOM, with net accumulation in spring and
584 consumption in autumn (Asmala, Haraguchi, Jakobsen, et al. 2018). In this study, we
585 highlight the CDOM dynamic by including DOC-specific visible absorbance (SVA₄₄₀),
586 which also showed net increase in spring and decrease in autumn. The SVA₄₄₀
587 dynamics showed that changes in the CDOM pool were not gradually occurring over
588 the course of the experiments, but at specific successional stages of the community,
589 primarily associated with phytoplankton decay. In spring, a drastic CDOM
590 accumulation was observed in phases S3 (treatment units) and S4 (controls), following

591 the collapse of the cryptophytes and ciliates populations. In autumn, a CDOM decrease
592 was only observed after the planktonic community collapsed (phases A3 and A4) and
593 when a large amount of DOC was released. This indicates that in autumn, the
594 planktonic community was likely producing DOM that was quickly processed and after
595 the community collapse, bacteria had only more recalcitrant CDOM available, resulting
596 in the accumulation of recalcitrant DOC. This could also explain the magnitude of the
597 observed changes in both experiments, where terrestrial DOM associated with net
598 primary productivity and slower turnover rates in spring would promote larger changes
599 in the DOM pool as it would be more labile. Conversely, the background DOM in
600 autumn is likely to be more recalcitrant, as it has been exposed to degradation over
601 summer, and as the turnover rates of the community are also faster, any fresh DOM
602 produced would be quickly processed. However, when the pelagic community
603 collapsed, the microbial community had to process the more recalcitrant DOM, leading
604 to delayed and smaller changes in the pool in comparison to spring.

605 Those results suggest that over the annual cycle, along with the succession of planktonic
606 communities, the DOM pool follows the reactivity continuum, driven by the patterns in
607 nutrient availability. Our results are aligned with the concept that new nutrients
608 stimulate net community production leading to accumulation of non-labile DOM that
609 has been previously described for Roskilde Fjord (Asmala, Haraguchi, Jakobsen, et al.
610 2018) and for larger scales (Hansell & Carlson 1998; Romera-Castillo et al. 2016).

611 However, in this study we emphasize the complexity of DOM processing by natural
612 communities, showing that species composition (for both autotrophs and heterotrophs)
613 play a secondary role for determining signatures of environmental DOM, as species
614 composition in nature result from community development and maturity (in response to
615 nutrient limitation), and dilution of water masses. Analysing changes in DOM in light of

616 the community successional stages might provide a better framework for interpreting
617 autochthonous DOM processing in a broad range of environmental conditions, not only
618 over the seasonal but also at spatial scales.

619 **Conclusions**

620 Phytoplankton composition and biomass appear to have a minor effect on changes in the
621 bulk and optical characteristics of DOM, when heterotrophic community components
622 (grazers and bacteria) are present. We provide a comprehensive view on the
623 summarised complex processing of the autochthonous DOM produced by autotrophs in
624 the environment in contrast to experiments with axenic cultures or monocultures. Our
625 results suggest that the successional stages of diverse biological communities have an
626 effect on DOM cycling in coastal areas and that this community effect should be
627 considered in studies of environmental DOM. Our results indicate that even if
628 phytoplankton photosynthesis is the primary initial source of DOM, the extent of
629 subsequent heterotrophic processing and the ecological status of the community, which
630 are related to patterns in nutrient limitation, ultimately governs the flow of DOM in
631 natural environments and the seasonal signature of the semi-labile and recalcitrant
632 autochthonous DOM pool.

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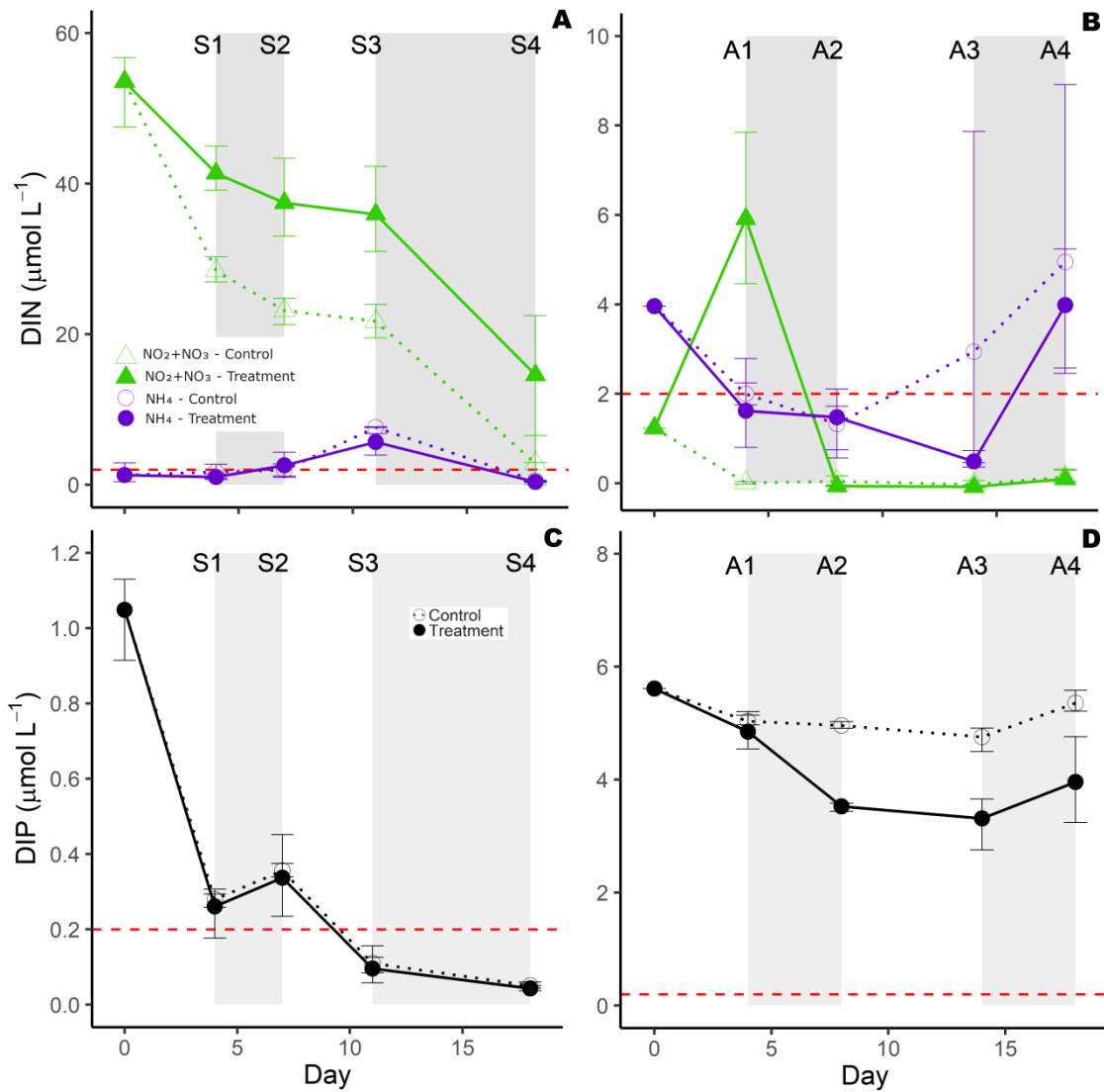
865

866 Table 1: Significant rates of change ($P < 0.05$) derived from the repeated measures mixed
 867 model for the different phases of the two seasonal experiments. Rates were estimated as
 868 the difference between start and end day divided by number of days. Non-significant
 869 rates ($P > 0.05$) are not shown. Phases with a significant change in rates between control
 870 (C) and treatment (T) are shaded in grey.

Spring experiment	Phases								
	S1		S2		S3		S4		
	C	T	C	T	C	T	C	T	
DIN ($\mu\text{M d}^{-1}$)	-6.17	-3.11						-3.71	-3.81
DIP ($\mu\text{M d}^{-1}$)	-0.19	-0.20			-0.06	-0.06			
Chla ($\mu\text{g L}^{-1} \text{d}^{-1}$)	2.82	2.16	-2.72	-2.43	-2.74		2.91	2.26	
C:Chla ($\% \text{d}^{-1}$)			25%	19%	-13%		-11%		
DOC ($\mu\text{M d}^{-1}$)									
DON ($\mu\text{M d}^{-1}$)			4.05						
DOP ($\mu\text{M d}^{-1}$)				-0.05	-0.04				
SVA ₄₄₀ ($\text{m}^2 \text{g}^{-1} \text{C d}^{-1}$)	-0.004		0.008			0.035	0.017	-0.003	
Autumn experiment	A1		A2		A3		A4		
	C	T	C	T	C	T	C	T	
DIN ($\mu\text{M d}^{-1}$)	-0.80			-1.53				0.92	
DIP ($\mu\text{M d}^{-1}$)	-0.14	-0.19		-0.33			0.15	0.16	
Chla ($\mu\text{g L}^{-1} \text{d}^{-1}$)		1.47		1.25	-0.48	-1.78			
C:Chla ($\% \text{d}^{-1}$)		23%	18%		-28%				
DOC ($\mu\text{M d}^{-1}$)					22.21	38.91			
DON ($\mu\text{M d}^{-1}$)	2.31		-3.19						
DOP ($\mu\text{M d}^{-1}$)									
SVA ₄₄₀ ($\text{m}^2 \text{g}^{-1} \text{C d}^{-1}$)			-0.005			-0.005	-0.005	-0.004	

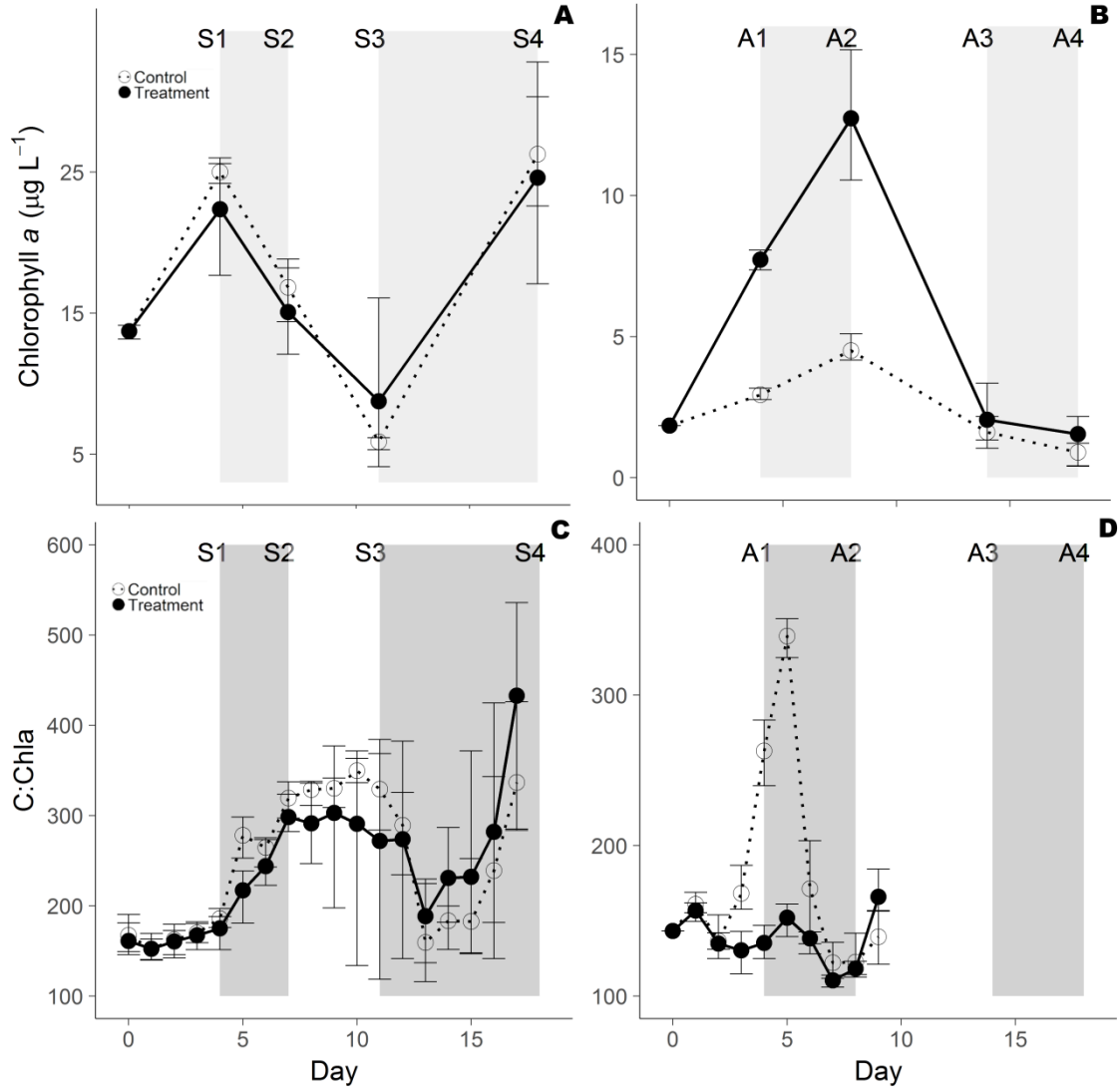
871

872 Figure 1. Dissolved inorganic nutrients concentrations during the course of the
 873 experiment in spring (left column) and autumn (right column) experiments. Dashed
 874 horizontal lines indicate nutrient thresholds potentially limiting phytoplankton growth
 875 ($2.0 \mu\text{mol L}^{-1}$ for DIN and $0.2 \mu\text{mol L}^{-1}$ for DIP). Note the difference in scaling between
 876 spring and autumn experiments. Error bars show maximum and minimum of replicated
 877 units. The different phases observed in the experiments are indicated with shaded
 878 background.



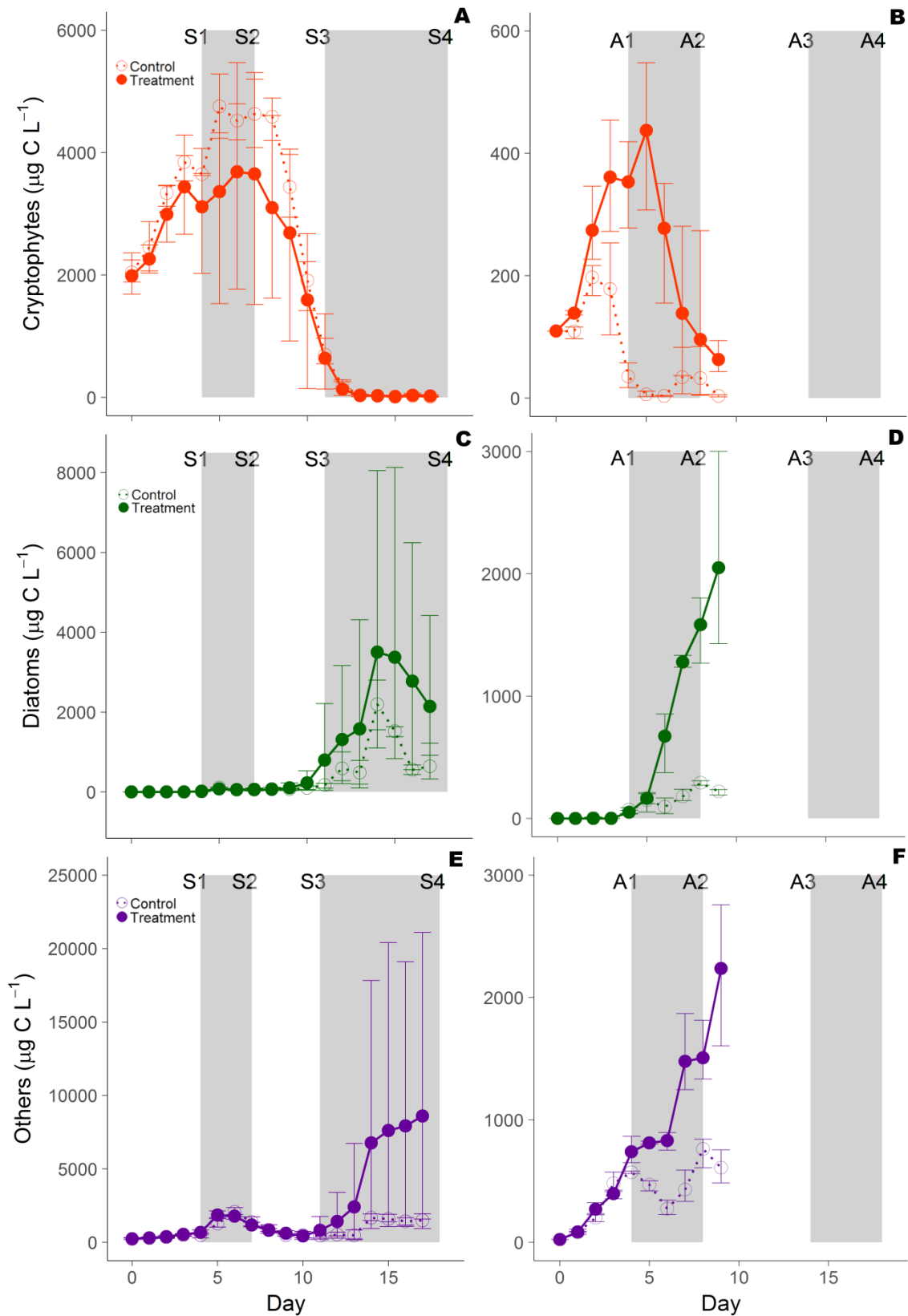
879

880 Figure 2. Phytoplankton biomass (chlorophyll *a*) and physiological state (C:Chla)
 881 temporal dynamics in spring (left column) and autumn (right column) experiments.
 882 Note the difference in scaling between spring and autumn experiments. Error bars show
 883 maximum and minimum of replicated units. The different phases observed in the
 884 experiments are indicated with shaded background.



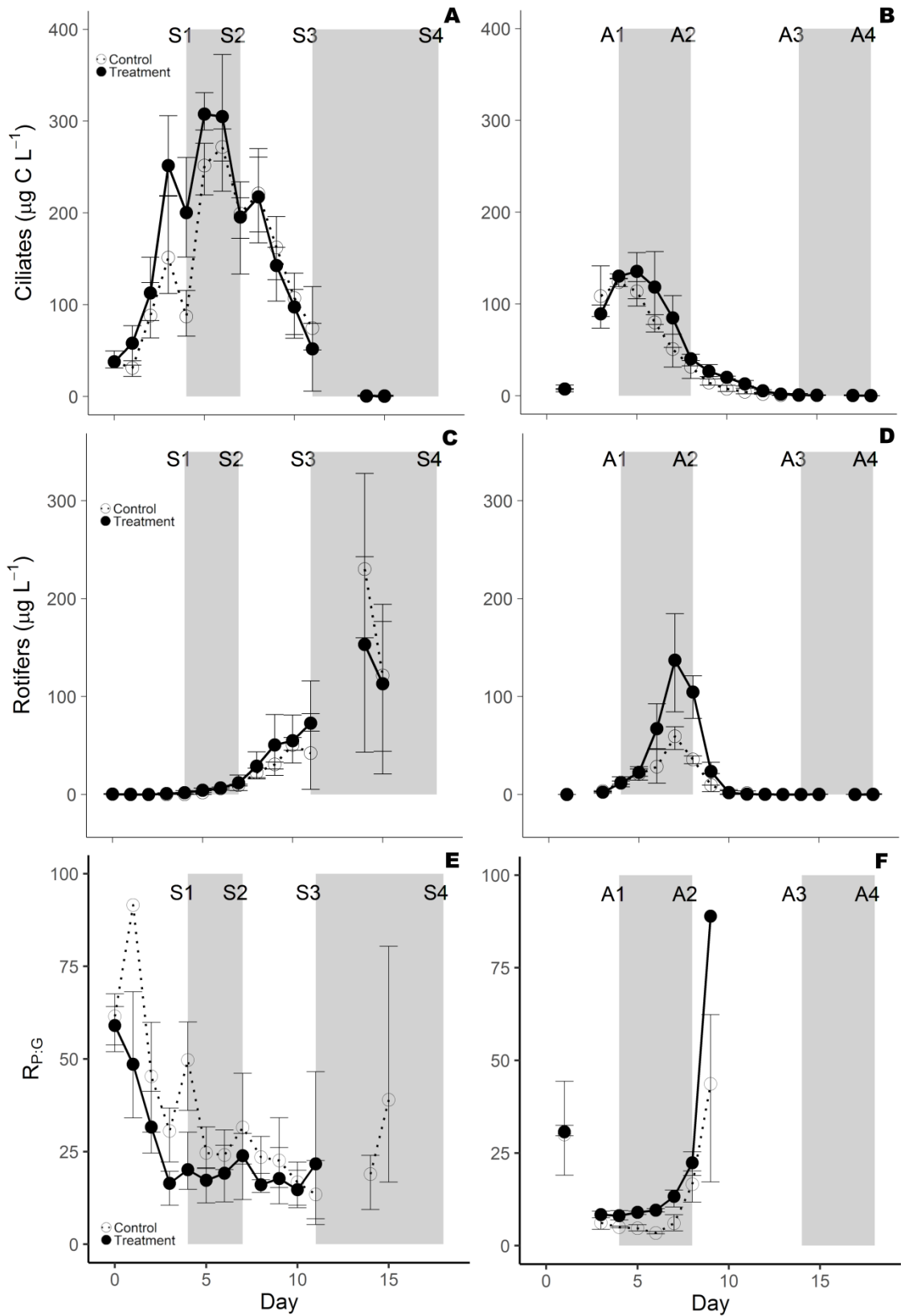
885

886 Figure 3. Temporal dynamics of the main phytoplankton groups in spring (left column)
 887 and autumn (right column) experiments. Note the difference in scaling between spring
 888 and autumn experiments. Error bars show maximum and minimum of replicated units.
 889 The different phases observed in the experiments are indicated with shaded background.



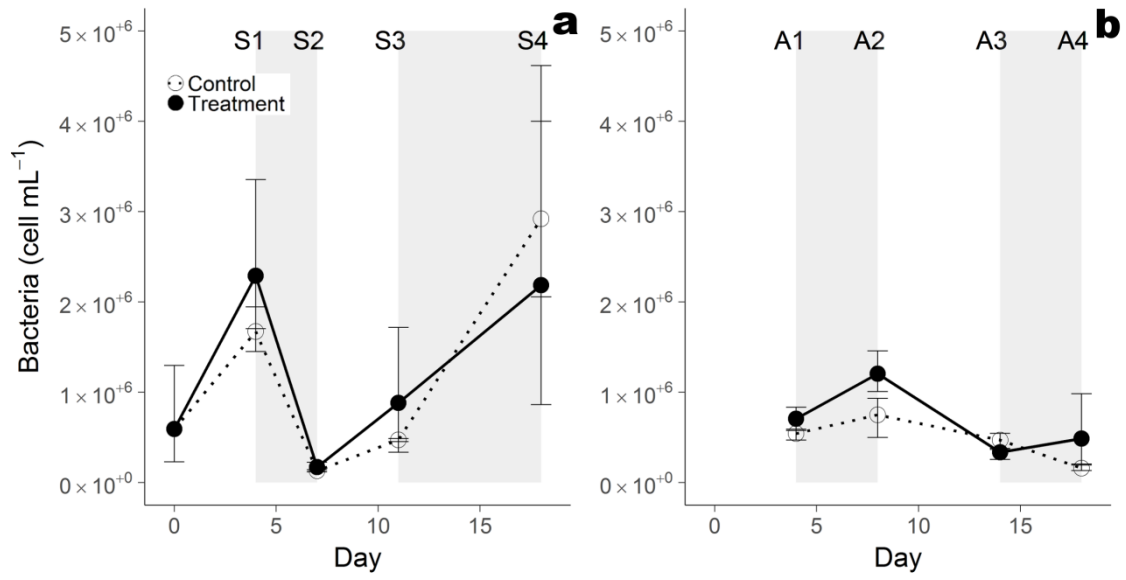
890

891 Figure 4. Ciliates and rotifer biomasses and the biomass ratio of phytoplankton to
 892 grazers ($R_{P:G}$) in spring (left column) and autumn (right column) experiments. Error
 893 bars show maximum and minimum of replicated units. The different phases observed in
 894 the experiments are indicated with shaded background.



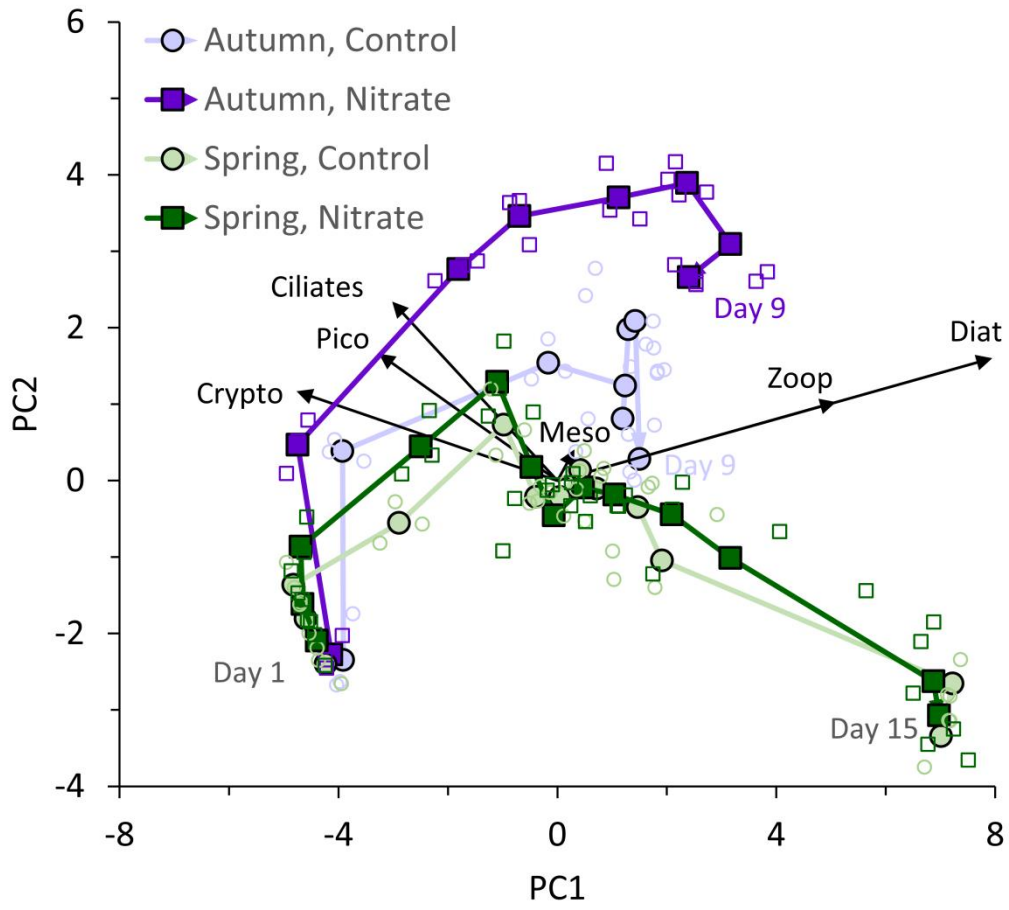
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896 Figure 5. Free-living bacteria abundances in spring (a) and autumn (b). Error bars show
897 maximum and minimum of replicated units. The different phases of the experiments are
898 indicated with shaded background.



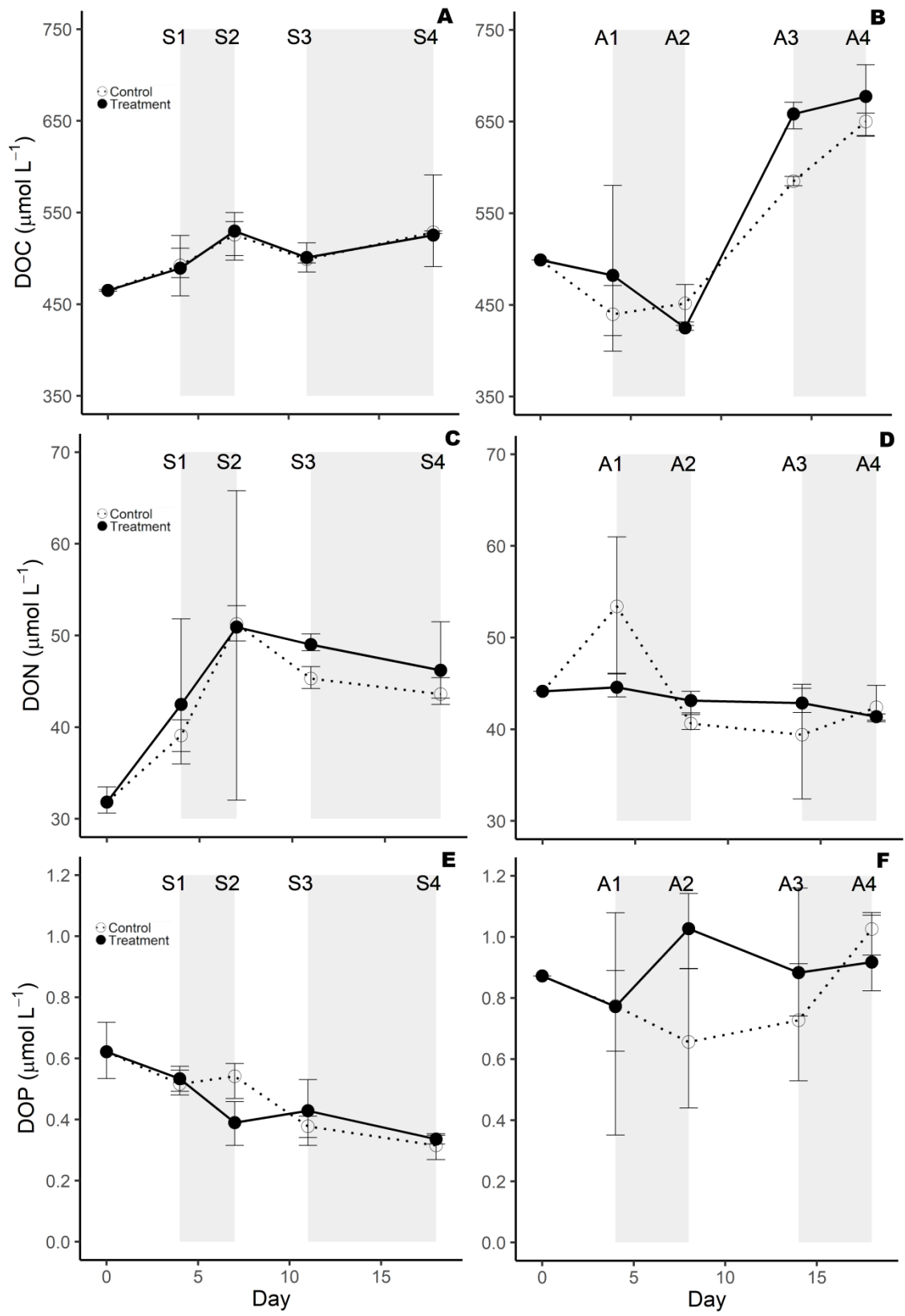
899

900 Figure 6. Trajectories of planktonic components over the course of the experiments
 901 obtained from Principal Components Analysis (PCA). Empty circles represent the
 902 individual observations and filled symbols show daily averages across the triplicate
 903 experimental units. Loadings of PC1 and PC2 are shown with arrows: Crypto =
 904 cryptophytes, Pico = pico-eukaryotes, Ciliates = ciliates, Meso = *Mesodinium rubrum*,
 905 Zoop = rotifers, and Diat = diatoms.



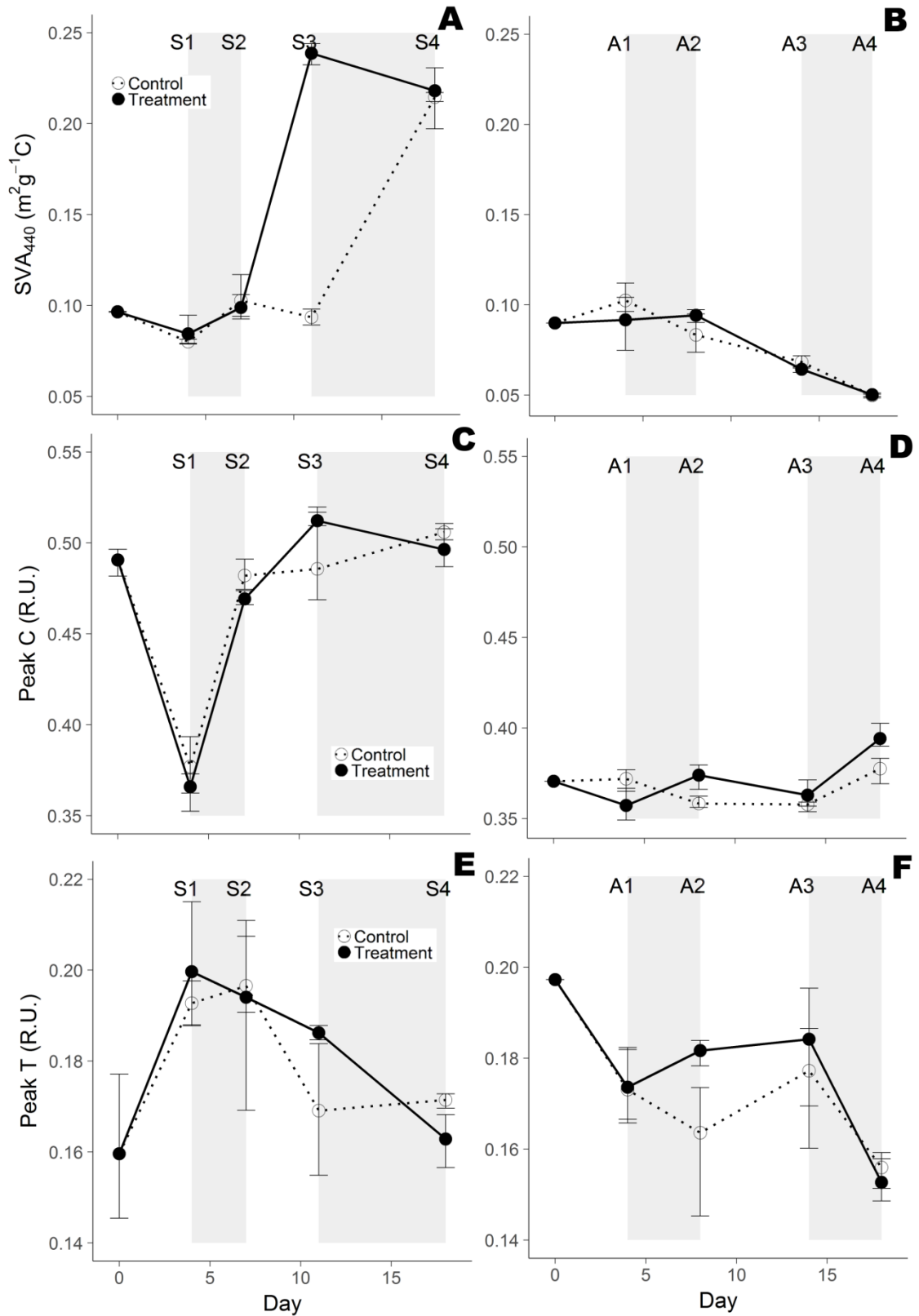
906

907 Figure 7. Dissolved organic carbon, nitrogen and phosphorus concentrations in spring
 908 (left column) and autumn (right column) experiments. Error bars show maximum and
 909 minimum of replicated units. The different phases observed in the experiments are
 910 indicated with shaded background.



911

912 Figure 8. CDOM (SVA₄₄₀) and FDOM (peak C and peak T) components in spring (left
 913 column) and autumn (right column) experiments. Error bars show maximum and
 914 minimum of replicated units. The different phases observed in the experiments are
 915 indicated by the shaded background colours.



916

917

918 **Table SI:** Significant rates of change ($P < 0.05$) derived from the repeated measures
919 mixed model for the different phases of the two seasonal experiments, for the plankton
920 community components and FDOM variables. Rates were estimated as the difference
921 between start and end day divided by number of days. For phytoplankton,
922 microzooplankton and bacteria response variables, rates are expressed as relative
923 changes per day. Non-significant rates ($P > 0.05$) are not shown. Phases with a
924 significant change in rates between control (C) and treatment (T) are shaded in grey.
925

Spring experiment	Phases							
	S1		S2		S3		S4	
	C	T	C	T	C	T	C	T
$C_{\text{cryptophytes}}$ (% d^{-1})					-38%	-40%	-52%	-40%
C_{diatoms} (% d^{-1})	85%	108%	81%			48%	22%	39%
C_{others} (% d^{-1})		37%	53%		-43%	-32%		
C_{ciliates} (% d^{-1})	23%	51%	31%		-23%	-35%	-63%	-54%
C_{rotifers} (% d^{-1})			107%	69%		61%	58%	
A_{bacteria} (% d^{-1})	41%	51%	-57%	-58%	38%	43%	29%	
Peak-C (R.U. d^{-1})	-0.028	-0.031	0.035	0.034		0.011		
Peak-T (R.U. d^{-1})	0.008	0.010			-0.007			-0.003
Autumn experiment	A1		A2		A3		A4	
	C	T	C	T	C	T	C	T
$C_{\text{cryptophytes}}$ (% d^{-1})	-26%			-48%	-76%			
C_{diatoms} (% d^{-1})	188%	167%	43%	136%				
C_{others} (% d^{-1})	43%	96%		-14%		-47%		
C_{ciliates} (% d^{-1})	152%	151%	-29%	-25%	-39%	-41%		
C_{rotifers} (% d^{-1})	136%	128%	29%	72%	-45%	-54%		
A_{bacteria} (% d^{-1})						-19%	-24%	
Peak-C (R.U. d^{-1})		-0.003	-0.003	0.004		-0.002	0.005	0.008
Peak-T (R.U. d^{-1})	-0.006	-0.006					-0.005	-0.008

926