1 Piperazine Enhancing Sulfuric Acid Based New Particle

2 Formation: Implications for the Atmospheric Fate of Piperazine

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14 **ABSTRACT**: Piperazine (PZ), a cyclic diamine, is one of 160 detected atmospheric 15 amines and an alternative solvent to the widely used monoethanolamine in post-16 combustion CO₂ capture. Participating in H₂SO₄ (SA)-based new particle formation (NPF) could be an important removal pathway for PZ. Here, we employed quantum 17 18 chemical calculations and kinetics modeling to evaluate the enhancing potential of PZ 19 on SA-based NPF by examining the formation of PZ–SA clusters. The results indicate 20 that PZ behaves more like a monoamine in stabilizing SA and can enhance SA-based 21 NPF at the parts per trillion (ppt) level. The enhancing potential of PZ is less than that 22 of the chainlike diamine putrescine, and greater than that of dimethylamine which is 23 one of the strongest enhancing agents confirmed by ambient observations and 24 experiments. After the initial formation of the $(PZ)_1(SA)_1$ cluster, the cluster mainly 25 grows by gradual addition of SA or PZ monomer, followed by addition of $(PZ)_1(SA)_1$ 26 cluster. We find that the ratio of PZ removal by NPF to that by the combination of NPF 27 and oxidations is 0.5-0.97 at 278.15 K. As a result, the participation in the NPF pathway 28 could significantly alter the environmental impact of PZ compared to only considering 29 oxidation pathways.

30 INTRODUCTION

31 Amines are a class of atmospheric nitrogen-containing organic pollutants. Up to 32 now, about 160 different amines have been detected in the atmosphere.¹ Amines are 33 emitted into the atmosphere from both natural and anthropogenic sources including 34 agriculture, biomass burning, animal husbandry, oceans, cooking, smoking and various industrial processes.¹⁻¹² It deserves mentioning that CO₂ capture units will become a 35 36 significant source of amines once the promising amine-based CO₂ capture technology is implemented on a large scale.¹³⁻¹⁵ In recent years, the concern about the fate of amines 37 has been increasing since the transformation of amines could potentially form 38 39 carcinogenic nitrosamines.¹

Several studies have addressed the removal of amines by atmospheric oxidation.¹⁶⁻ ³⁹ The oxidation by hydroxyl radicals (·OH) has been considered to be their main transformation pathway, followed by chlorine radicals (·Cl), at daytime.^{18,32} The reactions with ·OH and ·Cl lead to amines having an atmospheric lifetime on the order of hours.^{1,13,17-19,22-28,32} More importantly, the atmospheric oxidation by ·OH and ·Cl can lead to the formation of N-center radicals, which can further react with NO_x (x = 1, 2) to form hazardous nitrosamines/nitramine (Scheme 1), increasing the environmental

risk of the amines emissions. 17,18,28,30,32 Although the direct reaction of amines with NO_x 47 48 can also lead to the formation of nitrosamines, several studies have shown that it is of little importance under atmospheric condition.^{13,40,41} Besides oxidation reactions, many 49 50 studies have found that amines such as monomethylamine (MA), dimethylamine (DMA), trimethylamine (TMA) and monoethanolamine (MEA) can significantly 51 52 enhance H₂SO₄ (SA)-based new particle formation (NPF) via acid-base reactions, an important process for the formation of atmospheric aerosol particles.⁴²⁻⁷² However, 53 54 there are few studies concerning how the participation of amines in SA-based NPF 55 competes with their oxidation pathways and thereby affects the fate of atmospheric 56 amines. Our recent study indicated that the participation in SA-based NPF is an 57 important sink for MEA, and it could even be comparable to the oxidation reaction pathway initiated by ·OH at 278.15 K.⁷⁰ This implies that the participation in NPF could 58 59 be a significant pathway in determining the fate of other amines, especially those with 60 high enhancing potential for SA-based NPF.



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62 Scheme 1. Major reaction pathways for ·OH/·Cl initiated reactions of amines. Here
63 exemplified using methylamine.

Piperazine (PZ), a cyclic secondary diamine, whose aqueous solution is considered an alternative solvent to MEA, a benchmark and widely used solvent in postcombustion CO₂ capture (PCCC) technology.^{73,74} Once PZ-based PCCC technology is implemented on a large scale, quantities of PZ might be released into the atmosphere from PCCC units due to its relatively high vapor pressure.⁷⁵ PZ has previously been identified as one of the 160 amines detected in the ambient atmosphere.¹ In Zonguldak 70 province, Turkey, the concentration of PZ was found to reach about 4 parts per trillion (ppt), which is higher than that of DMA (about 1 ppt) at the same location.⁷⁶ Our recent 71 72 study showed that the daytime atmospheric oxidation of PZ initiated by ·OH and ·Cl 73 can lead to higher yield of carcinogenic nitrosamines compared to the corresponding 74 oxidation process of MEA, implying higher environmental risk related to PZ than MEA emissions.³² Similar to MEA,⁷⁰ the participation of PZ in SA-based NPF via acid-base 75 reactions could be a significant pathway to compete with its oxidation pathways. 76 77 However, to the best of our knowledge, there are currently no studies that address the 78 participation of PZ in SA-based NPF.

79 The gas phase basicity (GB) of amines has been suggested to be an important parameter in determining their enhancing potential on SA-based NPF.^{53,54} The GB 80 value of PZ (914.7 kJ mol⁻¹) is significantly higher than those of previously studied 81 82 monoamines and ammonia (MA 864.5, DMA 896.5, MEA 896.8, NH₃ 819.0 kJ mol⁻ ¹).⁷⁷ Therefore, based purely on the GB, PZ should have much higher enhancing 83 84 potential on SA-based NPF than NH₃, MA, DMA and MEA under the assumption that 85 they have similar atmospheric concentration. In addition, the steric effect of the -NHgroup in the cyclic PZ should be different from previously studied chainlike 86 amines,^{56,70,78} which could influence its enhancing potential. At last, if the two -NH-87 88 groups of PZ can synergistically interact with SA like chainlike diamine putrescine 89 (PUT), PZ could have much higher enhancing potential than what is expected by its 90 GB. However, due to its rigid cyclic structure, it is not easy to judge whether the two -91 NH- groups of PZ can synergistically interact with SA. All in all, it is difficult to 92 estimate the enhancing potential of PZ on SA-based NPF based on chemistry intuition. 93 In addition, no previous studies have investigated the potential role of cyclic diamines 94 in SA-based NPF. Thus, to comprehensively understand the atmospheric fate of PZ and expand the knowledge of amines enhancing SA-based NPF, it is indispensable toinvestigate the participation of PZ in SA-based NPF.

Here, we investigated the initial step of PZ participating in SA-based NPF by examining the formation of $(PZ)_x(SA)_y$ (x = 0-4, y = 0-4) clusters by a combined method using quantum chemical calculations with the Atmospheric Cluster Dynamics Code $(ACDC)^{79-82}$. The results are compared with previous studies on the DMA–SA, MEA–SA and PUT–SA systems. In addition, the effect of adding water molecules to the clusters was also considered to study the effect of hydration on the cluster formation kinetics of PZ and SA molecules.

104 **COMPUTATIONAL DETAILS**

105 Configurational Sampling and Electronic Structure Calculations. Obtaining a good 106 estimate of the global free energy cluster structures remains a large challenge in 107 modelling atmospheric NPF. Here, we employed a multi-step sampling scheme to search for the global minima of the $(PZ)_x(SA)_y$ (x = 0-4, y = 1-4) clusters. The pure 108 (SA)₁₋₄ clusters were taken from previous work.⁸¹ The multi-step sampling scheme has 109 extensively been applied to study atmospheric cluster formation.^{70,78,83-88} In brief, the 110 111 scheme includes the following six steps: (1) Large number of initial cluster 112 configurations (about 10000 for most of clusters) are randomly generated; (2) All the 113 configurations are initially optimized using the semiempirical PM6 method; (3) Single 114 point energy calculations at the ω B97X-D/6-31+G(d) level of theory is performed on 115 all the optimized configurations; (4) The identified lowest energy configurations within 10-15 kcal mol⁻¹ are fully optimized at the ω B97X-D/6-31++G(d,p) level of theory, 116 followed by a vibrational frequency calculation; (5) For the identified lowest free 117 118 energy configurations (within 1-2 kcal mol⁻¹ of the global minimum), the single point 119 energy was refined with a DLPNO-CCSD(T)/aug-cc-pVTZ calculation; (6) Several of 120 the lowest free energy configurations (about 10 for most of clusters) are subsequently 121 used to build initial configurations for the larger clusters starting over from (1). The 122 initial PZ conformations for building the (PZ)₁(SA)₁ and (PZ)₂ clusters were obtained from our previous AIMD simulation.³² All geometry optimization, vibrational 123 124 frequency calculations and single point energies using the PM6 and ω B97X-D methods were performed in the GAUSSIAN 09 program package.⁸⁹ The DLPNO-CCSD(T)/aug-125 cc-pVTZ calculation was performed in the ORCA 4.0.0 program.⁹⁰ The ω B97X-D/6-126 31++G(d,p) and DLPNO-CCSD(T)/aug-cc-pVTZ methods were selected as the core 127 optimization/frequency and single point energy calculations, respectively, since they 128 129 have shown good performance for studying the formation of atmospheric molecular clusters.^{70,91,92} The Gibbs free energy (G) of the identified clusters were calculated at 130 131 298.15 K with the following formula:

$$G = E + G_{\rm corr} \tag{1}$$

where *E* is the electronic single point energy at the DLPNO-CCSD(T)/aug-cc-pVTZ level of theory and G_{corr} is the Gibbs free energy correction at the ω B97X-D/6- $31^{++}G(d,p)$ level of theory. The formation free energy (ΔG) for each cluster at 298.15 K was calculated by:

137
$$\Delta G = G_{\text{cluster}} - \Sigma G_{\text{monomer}}$$
(2)

138 where G_{cluster} and G_{monomer} is the free energy of the cluster and the constituent molecules, 139 respectively. The ΔG values at other temperatures were calculated under the assumption 140 that enthalpy (ΔH) and entropy (ΔS) change remain constant in the tropospheric 141 temperature range. It is important to note that the ΔG values of the cluster systems used 142 as a comparison (MEA–SA, DMA–SA, PUT–SA), were obtained at the same 143 theoretical level as those in this study. The Cartesian coordinates of the stable PZ–SA 144 clusters were presented in the Supporting Information (SI).

To investigate the effect of hydration, we studied the $(PZ)_x(SA)_yW_z$ (x = 0-2, y = 0-2, z = 1-5, "W" represents H₂O) clusters. The $(SA)_1W_{1-5}$ and $(SA)_2W_{1-3}$ clusters were taken from previous studies.^{70,93} The global minima of $(SA)_2W_{4-5}$ and $(PZ)_{1-2}(SA)_{1-2}W_{1-5}$ clusters were identified by the same sampling scheme that was used for the unhydrated PZ–SA clusters.

Atmospheric Clusters Dynamic Code (ACDC) Model. The time evolution of 150 151 formation rates, steady-state concentrations and growth paths of clusters were studied 152 using the ACDC code. The detailed theory of the ACDC can be seen in a previous study.⁸¹ Here, the simulation system was treated as a " 4×4 " box for the unhydrated 153 154 PZ-SA system, where 4 is the maximum number of PZ or SA molecules of the clusters. 155 The mobility diameter of the largest cluster is ~ 1.5 nm, which closely resembles the 156 sizes of clusters that can be deemed stable against evaporation in the ambient 157 atmosphere. The $(PZ)_4(SA)_5$ and $(PZ)_5(SA)_5$ clusters were set as the boundary clusters 158 (SI). The ACDC simulations were mainly performed at 278.15 K. In addition, to probe 159 the temperature effect, we also conducted simulations at other temperatures such as 160 258.15, 268.15, 288.15, 298.15 and 313.15 K. The concentration of SA ([SA]) and PZ 161 ([PZ]) were set to be 10^5 , 10^6 , 10^7 and 10^8 cm⁻³ (a range relevant to atmospheric particle formation)^{62,94-97} and 1, 10, 100 ppt (partly higher than that (about 4 ppt) measured in 162 Zonguldak province, Turkey),⁷⁶ respectively. To consider external losses, a constant 163 coagulation sink coefficient of 2.6×10^{-3} s⁻¹ was used.^{98,99} When studying the effect of 164 165 hydration, the simulation system was treated as a " 2×2 " box. The [SA] and [PZ] were set to be 10⁶ cm⁻³ and 10 ppt, respectively, and the simulations were performed at 166 167 278.15 K. The equilibrium hydrate distribution for each cluster was calculated by the equilibrium constant for the formation of the respective hydrate.^{88,93,100} The (PZ)₂(SA)₃ 168

and $(PZ)_3(SA)_3$ clusters were set as the boundary clusters for the hydrated system.

170 RESULTS AND DISCUSSION

171 Cluster Structures. Since many previous studies have already discussed the structures of pure SA clusters,⁸¹ herein, we mainly focus on the clusters $(PZ)_x(SA)_y$ (x = 0-4, y = 172 1-4). The structures of $(PZ)_x(SA)_y$ (x = 0-4, y = 1-4) clusters with the lowest Gibbs 173 174 free energy were presented in Figure 1 and the number of the proton transfer observed 175 in the PZ-SA clusters were shown in Table S2. For the homomolecular PZ clusters, 176 proton transfer is not observed and all the clusters except (PZ)₂ are mainly stabilized by 177 hydrogen bonds (H-bonds), similar to the cases of studied homomolecular NH₃ or amines clusters.^{56,58,59,69,70,78,101,102} However, the (PZ)₂ cluster is stabilized by two 178 179 N…HC interactions. For heteromolecular PZ-SA clusters, proton transfer is observed 180 in all cases and the clusters are stabilized by H-bonds and ionic electrostatic 181 interactions. The proton transfer involves two different patterns. In the first pattern, only one proton of SA is transferred. Therefore, the formation of sulfate ion (SO_4^{2-}) is 182 183 not observed. In this pattern, PZ can accept one proton in two different ways: 1) one 184 PZ molecule only accepts a single proton. The clusters following this way include $(PZ)_1(SA)_{1-3}$, $(PZ)_2(SA)_{1-4}$, $(PZ)_3(SA)_1$ and $(PZ)_3(SA)_3$. Therefore, only one -NH-185 186 group of PZ is protonated for these clusters. 2) one PZ molecule accepts two protons 187 from different SA molecules. Therefore, the two -NH- groups of one PZ are 188 protonated. This phenomenon is only observed in the $(PZ)_1(SA)_4$ cluster.

In the second pattern, one SA molecule transfers two protons to two different PZ molecules, resulting in the formation of a SO_4^{2-} and a single protonated –NH– group in the two PZ molecules, as seen in the $(PZ)_4(SA)_{1-4}$, $(PZ)_3(SA)_2$, and $(PZ)_3(SA)_4$ cluster structures. In the $(PZ)_3(SA)_4$ and $(PZ)_4(SA)_4$ clusters, two and one SA molecules do not transfer any proton, respectively, which makes the number of SA molecules that donate 194 a proton less than the number of protonated PZ. Therefore, one SA in these two clusters has to donate two protons leading to the formation of a SO_4^{2-} . These patterns are vastly 195 different from previously reported amines (MA, MEA, DMA and PUT) and SA clusters 196 with the same composition of acid-base molecules.^{69,70,78} Note that the maximum 197 number of formed SO₄²⁻ is one in all considered clusters, in contrast to the clusters 198 consisting of SA and chainlike diamine PUT, in which several SO_4^{2-} can be formed.⁷⁸ 199 200 The difference results from the fact that it is unfavorable for the rigid structure of PZ to 201 simultaneously accept two protons from one SA as opposed to chainlike diamines such 202 as PUT. Therefore, from a structural point of view, PZ behaves more like a monoamine 203 in stabilizing SA. Another structural feature in all the clusters except (PZ)₁(SA)₄ is that 204 only one of the -NH- groups of PZ interacts with SA or PZ molecule, the remaining -205 NH- group points towards the outside. Therefore, the two -NH- groups of PZ neither behave like chainlike diamine to accept two protons from one SA, nor like MEA to 206 synergistically interact with SA or another amine molecule via two functional 207 groups.70,78 208



Figure 1. Identified lowest free energy structures of the $(PZ)_x(SA)_y$ (x = 0-4, y = 1-4) clusters at the ω B97X-D/6-31++G(d,p) level of theory. The red, blue, gray and white balls represent oxygen, nitrogen, carbon and hydrogen atoms, respectively. Dashed red lines indicate hydrogen bonds.

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214 **Cluster Formation Free Energy.** Both ambient observations and experimental studies 215 have confirmed that DMA is one of the strongest species for stabilizing SA clusters and 216 thus enhancing NPF.^{61,62,71} Therefore, the ΔG values of the DMA–SA system were 217 taken as reference to discuss those of the PZ–SA system. The formation free energy



For the potential use of PZ as an alternative solvent to MEA in PCCC applications, it is interesting to compare the ΔG values of PZ–SA with the MEA–SA systems. It was found that ΔG values of the majority of the PZ–SA clusters are lower than those of the corresponding MEA–SA clusters with the exception of the (PZ)₂₋₄ and (PZ)₁(SA)₃₋₄ clusters. This exception illustrates the important role of the –OH group of MEA in the formation of MEA–SA clusters as presented in our previous study.⁷⁰



Figure 2. Calculated formation free energies (ΔG) (A) for $(PZ)_x(SA)_y$ clusters (x = 0-4, y = 0-4) at the DLPNO-CCSD(T)/aug-cc-pVTZ// ω B97X-D/6-31++G(d,p) level of theory and corresponding total evaporation rates (B) for each cluster at 278.15 K and

the reference pressure of PZ and SA is set to 1 atm.

239 **Evaporation Rates and Cluster Stability.** Comparing the evaporation rate of a cluster 240 to the growth rate due to the collisions with vapor molecules at the given acid and base 241 concentration yields a measure of the stability of the cluster against evaporation. The 242 calculated evaporation rates of all PZ-SA clusters at 278.15 K were presented in Figure 243 2B. According to the condition judging the stability of a cluster (the cluster with evaporation rate lower than 10^{-3} s⁻¹ is stable when the concentration of acid or base 244 245 monomer is around or above ppt level), only $(PZ)_1(SA)_2$, $(PZ)_1(SA)_4$, $(PZ)_2(SA)_2$, 246 (PZ)₂(SA)₄, (PZ)₃(SA)₃, (PZ)₄(SA)₃, and (PZ)₄(SA)₄ clusters can be considered stable enough against evaporation. The three clusters along the diagonal ((PZ)₂(SA)₂, 247 248 $(PZ)_3(SA)_3$, $(PZ)_4(SA)_4$) as well as the $(PZ)_1(SA)_2$ clusters are the most stable with evaporation rates of $10^{-5}-10^{-7}$ s⁻¹. In addition, the clusters (PZ)₂(SA)₃ and (PZ)₃(SA)₄ 249 are relatively stable with evaporation rates in the order of 10^{-2} s⁻¹. By checking the 250 251 evaporation rates for all possible evaporation pathways of each of the PZ-SA clusters 252 (Table S3), it was found that the main decay route for all PZ-SA clusters (except 253 clusters (PZ)₄(SA)₂₋₄) is via evaporation of a PZ or SA monomer. A detailed discussion 254 on the main decay routes for all PZ-SA clusters were presented in SI.

255 It is also interesting to compare cluster evaporation rates for the different amines 256 (MEA, DMA, PZ and PUT) at the same simulation conditions. Generally, most of the 257 PZ-SA clusters have lower evaporation rates than the corresponding 258 DMA/MEA/PUT-SA clusters.^{70,78} For the number of clusters with evaporation rates less than or around 10⁻³ s⁻¹, the PZ–SA system is equal to that of the PUT–SA system 259 and higher than the MEA/DMA-SA systems. The PZ-SA system has more clusters 260 with evaporation rates less than or around 10⁻⁵ s⁻¹ than any other amines-SA systems. 261 This indicates that the PZ-SA system has more stable clusters than the MEA/DMA-SA 262

263 systems and more high stable clusters than the PUT-SA system. In addition, these stable clusters for the PZ-SA system are more even-distributed among the clusters with 264 different size than those of other systems. The above results do not necessarily 265 266 guarantee the faster growth of PZ-SA system compared to the DMA/MEA/PUT-SA systems. However, the higher number of stable clusters will facilitate the growth of the 267 PZ-SA system once the crucial PZ-SA cluster is formed. When the initially formed 268 269 one SA and one base cluster, which are crucial for cluster growth at relevant SA and 270 base concentration for PZ-SA (see Section Growth Pathway), DMA-SA, MEA-SA and PUT-SA systems,^{69,70,78} are compared, the trend in evaporation rate follows 271 272 $(PUT)_1(SA)_1 > (PZ)_1(SA)_1 > (DMA)_1(SA)_1 > (MEA)_1(SA)_1$ at the given acid and base concentrations.70,78 273

274 Steady-State Cluster Concentrations and Formation Rates. The cluster formation rates (J_{PZ}) and steady-state sulfuric acid dimer concentrations ($\sum_{i} [(SA)_2(PZ)_i]$) as a 275 function of the concentration of SA (10⁵-10⁸ cm⁻³) and PZ (1-100 ppt) for the PZ-SA 276 277 system at 278.15 K were presented in Figure 3. The comparison with the DMA-SA $(J_{\text{DMA}}, \sum_{i} [(\text{SA})_2(\text{DMA})_i]), \text{PUT-SA} (J_{\text{PUT}}, \sum_{i} [(\text{SA})_2(\text{PUT})_i]) \text{ and MEA-SA} (J_{\text{MEA}}, J_{\text{MEA}})$ 278 $\sum_{i} [(SA)_2(MEA)_i])$ cluster systems were shown in Figure S3. As can be seen in Figure 279 3, with increasing [SA] and [PZ], J_{PZ} and $\sum_{i} [(SA)_2(PZ)_i]$ gradually increase. As the 280 [PZ] increases, $\sum_{i} [(SA)_2(PZ)_i]$ becomes saturated. More importantly, the J_{PZ} is 1–10³, 281 282 $8-10^3$ and 0.02-0.9 times that of J_{DMA} , J_{MEA} and J_{PUT} , respectively, and $\sum_{i} [(SA)_{2}(PZ)_{i}]$ is 0.7–2, 5–80 and 0.08–0.9 times that of $\sum_{i} [(SA)_{2}(DMA)_{i}]$, 283 $\sum_{i} [(SA)_2(MEA)_i]$ and $\sum_{i} [(SA)_2(PUT)_i]$, respectively, depending on the 284 285 concentration of SA and amines. The enhancing potential of these four amines follows





303 Figure 3. Simulated steady-state SA dimer concentration $\sum_{i} [(SA)_2(PZ)_i]$ (cm⁻³) (A) 304 and the cluster formation rates J_{PZ} (cm⁻³ s⁻¹) out of the simulation systems (B) as a

305 function of [PZ] at 278.15 K.

306 Growth Pathways. The growth pathway and actual Gibbs free energy surface for the PZ-SA clusters at 278.15 K, $[SA] = 10^6$ cm⁻³ and [PZ] = 10 ppt were shown in Figure 307 308 4. As can be seen in Figure 4A, the first step of the PZ-SA system growth is the 309 formation of the (PZ)₁(SA)₁ cluster, similar to the cases of the MEA-SA and DMA-SA systems.⁷⁰ The growth of the formed $(PZ)_1(SA)_1$ cluster mainly proceeds by first adding 310 311 one SA molecule, and then one PZ molecule until the formation of (PZ)₃(SA)₃ cluster. 312 This mechanism is similar to the case of the MEA-SA system.⁷⁰ However, different 313 from MEA-SA system, the (PZ)₃(SA)₃ cluster growth mainly proceeds by first adding 314 one PZ molecule, and then one SA molecule to form (PZ)₄(SA)₄ cluster. Collisions with 315 the (PZ)₁(SA)₁ cluster, instead of PZ or SA molecule, contributes 20-33% to the 316 formation of (PZ)₂(SA)₂, (PZ)₃(SA)₃ and (PZ)₄(SA)₄ clusters. (PZ)₄(SA)₅ (57%) is main 317 cluster leaving the simulation box, followed by (PZ)₅(SA)₅ (24%), (PZ)₆(SA)₅ (10%) 318 and other clusters (9%). Combining the growth pathway with the actual Gibbs free 319 energy surface (Figure 4B), it can be seen that only the cluster $(PZ)_2(SA)_2 \rightarrow$ (PZ)₂(SA)₃ process needs to overcome a small barrier, whereas the remaining processes 320 along the main growth pathway are barrierless after the formation of the $(PZ)_{1}(SA)_{1}$ 321 322 cluster. However, the growth pathway via cluster collisions with (PZ)₁(SA)₁ cluster 323 along the diagonal is barrierless along the entire growth pathway. Combining the 324 growth pathway with the evaporation rate of the PZ-SA system, we conclude that the 325 formation of initial $(PZ)_1(SA)_1$ cluster is the rate determining step for the cluster growth 326 due to its instability compared with other clusters in the growth pathway, similar to cases of clusters containing SA and other amines including MA, MEA and DMA.⁷⁰ 327



Figure 4. Main clustering pathways (A) and actual Gibbs free energy surface (B) for the formation of clusters $(PZ)_x(SA)_y$ (x = 0-4, y = 0-4) at 278.15 K, $[SA] = 10^6$ cm⁻³, and [PZ] = 10 ppt.

331 Hydration Effect. We considered 1–5 H₂O molecules to study the effect of hydration 332 on the formation kinetics of the PZ-SA clusters. It should be noted that only clusters 333 $(PZ)_x(SA)_y$ (x = 0-2, y = 0-2) were considered as a test to investigate the effect of 334 hydration as the computational cost increasing rapidly when studying larger clusters. 335 Details for the discussion on the calculated stepwise hydration free energies and the 336 optimized conformations of the hydrated PZ-SA clusters were presented in the SI. The 337 calculated equilibrium hydrate distributions of the clusters at 278.15 K and relative 338 humidities (RH) 20%, 50% and 80% were presented in Figure S5. Figure S5 shows that 339 the PZ-SA clusters are hydrated by less than three H₂O molecules depending on the 340 RH. The evaporation rates and formation rates compared to dry conditions as a function 341 of RH at 278.15 K were presented in Figure 5. Figure 5A shows that the effect of 342 hydration on the evaporation rates depends on the cluster composition. Hydration has 343 little effect on the $(SA)_2$ cluster and almost no effect on the $(PZ)_2$ and $(PZ)_2(SA)_2$ 344 clusters. However, the evaporation rates of the $(PZ)_1(SA)_2$ and $(PZ)_2(SA)_1$ clusters can 345 be increased up to 50 and 190 times by hydration compared to the dry cases, 346 respectively. However, hydration can greatly decrease (up to 50 times) the evaporation rate of the initially formed $(PZ)_1(SA)_1$ cluster, the rate-determining step for the cluster growth in the system. This is the main reason for the increase in the cluster formation rates (Figure 5B) when hydration is considered compared to the dry case. The cluster formation rates increase up to 2 times compared to the dry case. Therefore, from these small cluster hydration simulations, we can conclude that hydration has a significant effect on the evaporation rates and a minor effect on the formation rates.



Figure 5. Relative evaporation rates (A) and cluster formation rates ([SA] = 10^6 cm⁻³ and [PZ] = 10 ppt) (B) as a function of relative humidities (*RH*) at 278.15 K.

355 Implications. We have revealed that PZ at ppt level can significantly enhance SA-356 based NPF. The enhancing potential of PZ is higher than that of DMA and MEA, and lower than that of PUT.^{70,78} The order of the enhancing potential of the amines is 357 358 consistent with that of their GB, further indicating the important role of GB of amines 359 involved in SA-based NPF. In addition, we showed that the two -NH- groups of PZ 360 cannot synergistically interact with SA, making PZ behave similarly to monoamines as 361 opposed to chainlike diamines in stabilizing SA. In this way, one of the -NH- groups points outwards from the clusters. The existence of the exterior -NH- group on the 362 363 surface of the clusters would make the subsequent growth mechanism of the PZ-SA 364 nucleus different from other cases of amines-SA nucleation. It could be interesting to 365 probe the further growth mechanism of PZ-SA nucleation in future by studying larger 366 cluster structures.

367 Obviously, the participation in SA-based NPF is one removal pathway for the emitted PZ, similar to the case of MEA. The removal rate constants (k_{SA}) of PZ by 368 participating in SA-based NPF are estimated to be 3.2×10^{-10} and 4.0×10^{-10} cm⁻³ 369 370 molecule⁻¹ s⁻¹ at 278.15 K at dry (RH = 0) and 50% RH conditions, respectively (SI). 371 Previous studies have shown that the reactions with ·OH and ·Cl are important removal pathways for PZ at daytime, due to their high reaction rate constants ($k_{OH} = 2.7 \times 10^{-10}$ 372 cm⁻³ molecule⁻¹ s⁻¹, $k_{Cl} = 4.7 \times 10^{-10}$ cm⁻³ molecule⁻¹ s⁻¹) at 278.15 K.^{28,32} The three 373 374 reactive agents (·OH, ·Cl and SA) toward PZ can coexist in the atmosphere and the 375 concentration of \cdot Cl ([\cdot Cl]) and [SA] is estimated to be around 0.01–0.1 and 1–19 times of that of \cdot OH ([\cdot OH]) during daytime, respectively.¹⁰⁵⁻¹¹⁰ Based on k_{SA} , k_{OH} , k_{Cl} , [SA], 376 377 [·OH] and [·Cl], we estimated the contribution of the participation in SA-based NPF 378 (Con_{SA}) to the removal of PZ by $k_{SA}[SA]/(k_{OH}[\cdot OH] + k_{CI}[\cdot CI] + k_{SA}[SA])$ at 278.15 K 379 and RH=0 or 50% (Table S6). As can be seen in Table S6, Con_{SA} is 50%-97%, 380 indicating the participation in SA-based NPF is a dominant removal pathway for PZ at 381 278.15 K, especially at high [SA]. It was found that the Con_{SA} to the removal of PZ has a negative temperature dependence in temperature range of 258.15–313.15 K (Figure 382 383 S6). When the temperature effect is considered, the participation in SA-based NPF still 384 play an important role in removing PZ at all atmospheric conditions, except a combined 385 condition of low [SA], low RH and high temperature. Therefore, if ignoring the 386 participation of PZ in SA-based NPF, the contribution of atmospheric oxidation by ·OH 387 and ·Cl to the removal of PZ will be highly overestimated. In addition, the contribution 388 of the participation in SA-based NPF on PZ removal is higher than that to the MEA 389 removal at both 0 and 50% RH conditions, and 278.15 K (SI). More importantly, the 390 high contribution of the NPF pathway to the removal of PZ decreases the overall nitrosamine yield compared to only considering the atmospheric ·OH and ·Cl oxidation pathways. The higher contribution of the NPF pathway to the removal of PZ than that to the removal of MEA decreases their relative risk of nitrosamine formation. A detailed discussion on the reevaluation of the overall nitrosamine yield of PZ and MEA was presented in SI. Therefore, this study further stresses that the participation in SA-based NPF should be considered for other atmospheric amines to evaluate the environmental risk, especially for the formation of carcinogenic nitrosamine.

398 ASSOCIATED CONTENT

399 Supporting Information. Details of the chosen boundary conditions, comparison for 400 the formation free energies at low and high theoretical level, observed number of proton 401 transfers in the PZ-SA system, stepwise hydration free energies, removal rate constants 402 of PZ by participating in SA-based NPF, main decay route for all PZ-SA clusters, thermochemical information of PZ-SA clusters, ΔG values of PZ-SA clusters at 403 404 298.15 K, evaporation rates for all evaporation pathways of PZ-SA clusters, the cluster 405 formation rates and steady-state SA dimer concentrations as a function of temperature, 406 conformations of the hydrated PZ-SA clusters, hydrate distribution of studied clusters, 407 Con_{SA} to the removal of PZ and MEA, detailed discussion on the reevaluation of the overall nitrosamine yield of PZ and MEA and coordinates of all optimized clusters. 408

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421 **REFERENCES**

- 422 (1) Ge, X. L.; Wexler, A. S.; Clegg, S. L. Atmospheric amines Part I. A review. *Atmos.* 423 *Environ.* 2011, 45, 524-546.
- 424 (2) Yao, L.; Wang, M. Y.; Wang, X. K.; Liu, Y. J.; Chen, H. F.; Zheng, J.; Nie, W.;
- 425 Ding, A. J.; Geng, F.-H.; Wang, D. F.; Chen, J. M.; Worsnop, D. R.; Wang, L. Detection
- 426 of atmospheric gaseous amines and amides by a high-resolution time-of-flight chemical
- ionization mass spectrometer with protonated ethanol reagent ions. *Atmos. Chem. Phys.* **2016**, 16, 14527-14543.
- 429 (3) Zheng, J.; Ma, Y.; Chen, M.; Zhang, Q.; Wang, L.; Khalizov, A. F.; Yao, L.; Wang,
- 430 Z.; Wang, X.; Chen, L. Measurement of atmospheric amines and ammonia using the
- 431 high resolution time-of-flight chemical ionization mass spectrometry. *Atmos. Environ.*432 2015, 102, 249-259.
- 433 (4) You, Y.; Kanawade, V. P.; de Gouw, J. A.; Guenther, A. B.; Madronich, S.; Sierra-
- Hernández, M. R.; Lawler, M.; Smith, J. N.; Takahama, S.; Ruggeri, G.; Koss, A.;
 Olson, K.; Baumann, K.; Weber, R. J.; Nenes, A.; Guo, H.; Edgerton, E. S.; Porcelli,
- Olson, K.; Baumann, K.; Weber, R. J.; Nenes, A.; Guo, H.; Edgerton, E. S.; Porcelli,
 L.; Brune, W. H.; Goldstein, A. H.; Lee, S. H. Atmospheric amines and ammonia
- 430 L., Brune, W. H., Goldstein, A. H., Lee, S. H. Atmospheric annues and annuolia 437 measured with a chemical ionization mass spectrometer (CIMS). *Atmos. Chem. Phys.*
- 437 inteastied with a chemical folization mass spectrometer (CHVIS). Atmos. Chem. F1 438 **2014**, 14, 12181-12194.
- 439 (5) Hemmilä, M.; Hellén, H.; Virkkula, A.; Makkonen, U.; Praplan, A. P.; Kontkanen,
 440 J.; Ahonen, L.; Kulmala, M.; Hakola, H. Amines in boreal forest air at SMEAR II
 441 station in Finland. *Atmos. Chem. Phys.* 2018, 18, 6367-6380.
- (6) Hong, Y.; Liu, Y. R.; Wen, H.; Miao, S. K.; Huang, T.; Peng, X. Q.; Jiang, S.; Feng,
 Y. J.; Huang, W. Interaction of oxalic acid with methylamine and its atmospheric
 implications. *RSC Adv.* 2018, 8, 7225-7234.
- 445 (7) Mao, J. B.; Yu, F. Q.; Zhang, Y.; An, J. Y.; Wang, L.; Zheng, J.; Yao, L.; Luo, G.;
- Ma, W. C.; Yu, Q.; Huang, C.; Li, L.; Chen, L. M. High-resolution modeling of gaseous
 methylamines over a polluted region in China: source-dependent emissions and
 implications of spatial variations. *Atmos. Chem. Phys.* 2018, 18, 7933-7950.
- 449 (8) van Pinxteren, M.; Fomba, K. W.; van Pinxteren, D.; Triesch, N.; Hoffmann, E. H.;
- 450 Cree, C. H. L.; Fitzsimons, M. F.; von Tümpling, W.; Herrmann, H. Aliphatic amines
- 451 at the Cape Verde Atmospheric Observatory: Abundance, origins and sea-air fluxes.
 452 *Atmos. Environ.* 2019, 203, 183-195.
- 453 (9) Dawson, M. L.; Varner, M. E.; Perraud, V.; Ezell, M. J.; Wilson, J.; Zelenyuk, A.; 454 Gerber, R. B.; Finlayson-Pitts, B. J. Amine-amine exchange in aminium-
- 455 methanesulfonate aerosols. J. Phys. Chem. C 2014, 118, 29431-29440.
- (10) Chen, H.; Wang, M.; Yao, L.; Chen, J.; Wang, L. Uptake of gaseous alkylamides
 by suspended sulfuric acid particles: Formation of ammonium/aminium salts. *Environ. Sci. Technol.* 2017, 51, 11710-11717..
- 459 (11) Marrero-Ortiz, W.; Hu, M.; Du, Z.; Ji, Y.; Wang, Y.; Guo, S.; Lin, Y.; Gomez-
- 460 Hermandez, M.; Peng, J.; Li, Y.; Secrest, J.; Levy Zamora, M.; Wang, Y.; An, T.;
- 461 Zhang, R. Formation and optical properties of brown carbon from small alpha-462 dicarbonyls and amines. *Environ. Sci. Technol.* **2018**, 53, 117-126.
- 463 (12) Kumar, M.; Li, H.; Zhang, X.; Zeng, X. C.; Francisco, J. S. Nitric acid-amine
- 464 chemistry in the gas phase and at the air-water interface. J. Am. Chem. Soc. 2018, 140,
 465 6456-6466.
- 466 (13) Nielsen, C. J.; Herrmann, H.; Weller, C. Atmospheric chemistry and
- 467 environmental impact of the use of amines in carbon capture and storage (CCS). *Chem.*
- 468 Soc. Rev. 2012, 41, 6684-6704.

- 469 (14) Zhu, L.; Schade, G. W.; Nielsen, C. J. Real-time monitoring of emissions from
 470 monoethanolamine-based industrial scale carbon capture facilities. *Environ. Sci.*
- 471 *Technol.* **2013**, 47, 14306-14314.
- 472 (15) da Silva, G. Formation of nitrosamines and alkyldiazohydroxides in the gas phase:
- 473 the CH₃NH + NO reaction revisited. *Environ. Sci. Technol.* **2013**, 47, 7766-7772.
- 474 (16) Borduas, N.; Murphy, J. G.; Wang, C.; da Silva, G.; Abbatt, J. P. D. Gas phase
 475 oxidation of nicotine by OH radicals: kinetics, mechanisms, and formation of HNCO.
 476 *Environ. Sci. Technol. Lett.* 2016, 3, 327-331.
- 477 (17) Xie, H. B.; Li, C.; He, N.; Wang, C.; Zhang, S. W.; Chen, J. W. Atmospheric
 478 chemical reactions of monoethanolamine initiated by OH radical: mechanistic and
 479 kinetic study. *Environ. Sci. Technol.* 2014, 48, 1700-1706.
- 479 Kinetic study. *Environ. Sci. Technol.* 2014, 48, 1700-1706.
 480 (18) Xie, H. B.; Ma, F. F.; Wang, Y. F.; He, N.; Yu, Q.; Chen, J. W. Quantum chemical
- 480 (18) Ale, II. B., Ma, P. P., Walg, T. P., He, N., Fu, Q., Chen, J. W. Quantum chemical
 481 study on ·Cl-initiated atmospheric degradation of monoethanolamine. *Environ. Sci.*482 *Technol.* 2015, 49, 13246-13255.
- 483 (19) Xie, H. B.; Ma, F. F.; Yu, Q.; He, N.; Chen, J. W. Computational study of the 484 reactions of chlorine radicals with atmospheric organic compounds featuring NH_x -pi-485 bond (x = 1, 2) structures. J. Phys. Chem. A **2017**, 121, 1657-1665.
- 486 (20) Onel, L.; Thonger, L.; Blitz, M. A.; Seakins, P. W.; Bunkan, A. J.; Solimannejad,
- M.; Nielsen, C. J. Gas-phase reactions of OH with methyl amines in the presence or
 absence of molecular oxygen. an experimental and theoretical study. *J. Phys. Chem. A* **2013**, 117, 10736-10745.
- 490 (21) Tan, W.; Zhu, L.; Mikoviny, T.; Nielsen, C. J.; Wisthaler, A.; Eichler, P.; Muller,
- 491 M.; D'Anna, B.; Farren, N. J.; Hamilton, J. F.; Pettersson, J. B. C.; Hallquist, M.;
- 492 Antonsen, S.; Stenstrom, Y. Theoretical and experimental study on the reaction of tert-493 Butylamine with OH radicals in the atmosphere. *J. Phys. Chem. A* **2018**, 122, 4470-
- 494 4480.
- 495 (22) Barnes, I.; Wiesen, P.; Gallus, M. Rate coefficients for the reactions of OH radicals 496 with a series of alkyl-substituted amines. *J. Phys. Chem. A* **2016**, 120, 8823-8829.
- 497 (23) Onel, L.; Blitz, M. A.; Breen, J.; Rickard, A. R.; Seakins, P. W. Branching ratios
 498 for the reactions of OH with ethanol amines used in carbon capture and the potential
 499 impact on carcinogen formation in the emission plume from a carbon capture plant.
- 500 Phys. Chem. Chem. Phys. **2015**, 17, 25342-25353.
- 501 (24) Onel, L.; Blitz, M.; Dryden, M.; Thonger, L.; Seakins, P. Branching ratios in
 502 reactions of OH radicals with methylamine, dimethylamine, and ethylamine. *Environ.*503 Sci. Technol. 2014, 48, 9935-9942.
- 504 (25) Nicovich, J. M.; Mazumder, S.; Laine, P. L.; Wine, P. H.; Tang, Y.; Bunkan, A.
- 505 J.; Nielsen, C. J. An experimental and theoretical study of the gas phase kinetics of
- 506 atomic chlorine reactions with CH₃NH₂, (CH₃)₂NH, and (CH₃)₃N. *Phys. Chem. Chem.*
- 507 *Phys.* **2015**, 17, 911-917.
- 508 (26) Borduas, N.; Abbatt, J. P.; Murphy, J. G.; So, S.; da Silva, G. Gas-phase
 509 mechanisms of the reactions of reduced organic nitrogen compounds with OH radicals.
 510 *Environ. Sci. Technol.* 2016, 50, 11723-11734.
- 511 (27) Borduas, N.; da Silva, G.; Murphy, J. G.; Abbatt, J. P. D. Experimental and 512 theoretical understanding of the gas phase oxidation of atmospheric amides with OH
- 513 radicals: kinetics, products, and mechanisms. J. Phys. Chem. A **2015**, 119, 4298-4308.
- 514 (28) Onel, L.; Dryden, M.; Blitz, M. A.; Seakins, P. W. Atmospheric oxidation of
- 515 piperazine by OH has a low potential to form carcinogenic compounds. *Environ. Sci.*
- 516 Technol. Lett. 2014, 1, 367-371.

- 517 (29) Sarma, P. J.; Gour, N. K.; Bhattacharjee, D.; Mishra, B. K.; Deka, R. C. Hydrogen
- 518 atom abstraction from piperazine by hydroxyl radical: a theoretical investigation. *Mol.*
- 519 *Phys.* **2017**, 115, 962-970.
- 520 (30) Borduas, N.; Abbatt, J. P.; Murphy, J. G. Gas phase oxidation of 521 monoethanolamine (MEA) with OH radical and ozone: kinetics, products, and 522 particles. *Environ. Sci. Technol.* **2013**, 47, 6377-6383.
- 523 (31) Onel, L.; Blitz, M. A.; Seakins, P. W. Direct determination of the rate coefficient
- 524 for the reaction of OH radicals with monoethanol amine (MEA) from 296 to 510 K. J. 525 *Phys. Chem. Lett.* **2012**, 3, 853-856.
- 526 (32) Ma, F. F.; Ding, Z. Z.; Elm, J.; Xie, H. B.; Yu, Q.; Liu, C.; Li, C.; Fu, Z. Q.; Zhang,
- 527 L. L.; Chen, J. W. Atmospheric oxidation of piperazine initiated by Cl: unexpected 528 high nitrosamine yield. *Environ. Sci. Technol.* **2018**, 52, 9801-9809.
- 529 (33) Nielsen, C. J.; D'Anna, B.; Bossi, R.; Bunkan, A. J. C.; Dithmer, L.; Glasius, M.;
- 530 Hallquist, M.; Hansen, A. M. K.; Lutz, A.; Salo, K.; Maguta, M. M.; Nguyen, Q.;
- 531 Mikoviny, T.; Müller, M.; Skov, H.; Sarrasin, E.; Stenstrøm, Y.; Tang, Y.; Westerlund,
- 532 J.; Wisthaler, A. Atmospheric degradation of amines (ADA). *ISBN 978-82-992954-7-*
- 533 5, http://urn.nb.no/URN:NBN:no 30510, University of Oslo, Oslo, 2012.
- (34) Olszyna, K.; Heicklen, J. The inhibition of photochemical smog-Vl. The reaction
 of O₃ with diethyihydroxylamin. *Sci. Total Environ.* 1976, 5, 223-230.
- (35) Atkinson, R.; Perry, R. A.; Pitts, J. N. Rate constants for the reaction of the OH
 radical with CH₃SH and CH₃NH₂ over the temperature range 299-426 K. J. Chem. *Phys.* 1977, 66, 1578-1581.
- 539 (36) Atkinson, R.; Perry, R. A.; Pitts, J. N. Rate constants for the reactions of the OH
- radical with $(CH_3)_2NH$, $(CH_3)_3N$, and $C_2H_5NH_2$ over the temperature range 298-426 K. *J. Chem. Phys.* **1978**, 68, 1850-1853.
- 542 (37) Harris, G. W.; Atkinson, R.; Pitts, J. N. Kinetics of the reactions of the hydroxyl 543 radical with hydrazine and methylhydrazine. *J. Chem. Phys.* **1979**, 83, 2557-2559.
- 544 (38) Harris, G. W.; Pitts, J. N. Notes. Rates of reaction of hydroxyl radicals with 2-
- 545 (dimethylamino)ethanol and 2-amino-2-methyl-1-propanol in the gas phase at 300 ± 2 546 K. *Environ. Sci. Technol.* **1983**, 17, 50-51.
- 547 (39) Gleim, J.; Heicklen, J. The oxidation of diethylhydroxylamine by NO₂. *Int. J.* 548 *Chem. Kinet.* **1982**, 14, 699-710.
- 549 (40) Lv, C. L.; Liu, Y. D.; Zhong, R.; Wang, Y. Theoretical studies on the formation of 550 N-nitrosodimethylamine. *J. Mol. Struct.: THEOCHEM* **2007**, 802, 1-6.
- 550 N-Introsodimentifianme. *J. Mol. Struct.*. *THEOCHEM* **2007**, 802, 1-0. 551 (41) Hanst, P. L.; Spence, J. W.; Miller, M. Atmospheric chemistry of N-nitroso
- dimethylamine. *Environ. Sci. Technol.* **1977**, 11, 403-405.
- 553 (42) Chen, H.; Ezell, M. J.; Arquero, K. D.; Varner, M. E.; Dawson, M. L.; Gerber, R.
- 554 B.; Finlayson-Pitts, B. J. New particle formation and growth from methanesulfonic
- acid, trimethylamine and water. *Phys. Chem. Chem. Phys.* **2015**, 17, 13699-13709.
- 556 (43) Chen, H.; Varner, M. E.; Gerber, R. B.; Finlayson-Pitts, B. J. Reactions of
- 557 methanesulfonic acid with amines and ammonia as a source of new particles in air. *J.* 558 *Phys. Chem. B* **2016**, 120, 1526-1536.
- 559 (44) Xu, J.; Finlayson-Pitts, B. J.; Gerber, R. B. Proton transfer in mixed clusters of 560 methanesulfonic acid, methylamine, and oxalic acid: implications for atmospheric 561 particle formation. *J. Phys. Chem. A* **2017**, 121, 2377-2385.
- 562 (45) Arquero, K. D.; Xu, J.; Gerber, R. B.; Finlayson-Pitts, B. J. Particle formation and
- 563 growth from oxalic acid, methanesulfonic acid, trimethylamine and water: a combined
- september 2017, 19, 28286-28301.

- (46) Arquero, K. D.; Gerber, R. B.; Finlayson-Pitts, B. J. The role of oxalic acid in new
 particle formation from methanesulfonic acid, methylamine, and water. *Environ. Sci. Technol.* 2017, 51, 2124-2130.
- 568 (47) Xu, J.; Perraud, V.; Finlayson-Pitts, B. J.; Gerber, R. B. Uptake of water by an 569 acid-base nanoparticle: theoretical and experimental studies of the methanesulfonic
- 570 acid-methylamine system. Phys. Chem. Chem. Phys. 2018, 20, 22249-22259.
- 571 (48) Chen, H.; Finlayson-Pitts, B. J. New particle formation from methanesulfonic acid
 572 and amines/ammonia as a function of temperature. *Environ. Sci. Technol.* 2017, 51,
 573 243-252.
- 574 (49) Lin, Y.; Ji, Y.; Li, Y.; Secrest, J.; Xu, W.; Xu, F.; Wang, Y.; An, T.; Zhang, R.
- 575 Interaction between dicarboxylic acid and sulfuric acid-base clusters enhances new
- 576 particle formation. *Atmos. Chem. Phys. Discuss.* **2018**, 1-46. DOI: 10.5194/acp-2018-577 975
- 578 (50) Zhao, H.; Jiang, X.; Du, L. Contribution of methane sulfonic acid to new particle 579 formation in the atmosphere. *Chemosphere* **2017**, 174, 689-699.
- (51) Liu, L.; Li, H.; Zhang, H.; Zhong, J.; Bai, Y.; Ge, M.; Li, Z.; Chen, Y.; Zhang, X.
 The role of nitric acid in atmospheric new particle formation. *Phys. Chem. Chem. Phys.*
- **2018**, 20, 17406-17414.
- 583 (52) Chu, B. W.; Kerminen, V. M.; Bianchi, F.; Yan, C.; Petaja, T.; Kulmala, M. 584 Atmospheric new particle formation in China. *Atmos. Chem. Phys.* **2019**, 19, 115-138.
- 585 (53) Glasoe, W. A.; Volz, K.; Panta, B.; Freshour, N.; Bachman, R.; Hanson, D. R.;
- McMurry, P. H.; Jen, C. Sulfuric acid nucleation: An experimental study of the effect of seven bases. J. Geophys. Res-Atmos. 2015, 120, 1933-1950.
- 588 (54) Olenius, T.; Halonen, R.; Kurtén, T.; Henschel, H.; Kupiainen-Määttä, O.; Ortega,
- 589 I. K.; Jen, C. N.; Vehkamäki, H.; Riipinen, I. New particle formation from sulfuric acid 590 and amines: Comparison of monomethylamine, dimethylamine, and trimethylamine. *J.*
- 591 Geophys. Res-Atmos. 2017, 122, 7103-7118.
- 592 (55) Yu, H.; McGraw, R.; Lee, S. H. Effects of amines on formation of sub-3 nm 593 particles and their subsequent growth. *Geophys. Res. Lett.* **2012**, 39, L02807.
- 594 (56) Kurtén, T.; Loukonen, V.; Vehkamäki, H.; Kulmala, M. Amines are likely to 595 enhance neutral and ion-induced sulfuric acid-water nucleation in the atmosphere more 596 effectively than ammonia. *Atmos. Chem. Phys.* **2008**, 8, 4095-4103.
- 597 (57) Erupe, M. E.; Viggiano, A. A.; Lee, S. H. The effect of trimethylamine on 598 atmospheric nucleation involving H₂SO₄. *Atmos. Chem. Phys.* **2011**, 11, 4767-4775.
- 599 (58) Nadykto, A. B.; Herb, J.; Yu, F. Q.; Xu, Y. S. Enhancement in the production of
- 600 nucleating clusters due to dimethylamine and large uncertainties in the
- thermochemistry of amine-enhanced nucleation. *Chem. Phys. Lett.* **2014**, 609, 42-49.
- Nadykto, A. B.; Herb, J.; Yu, F. Q.; Xu, Y. S.; Nazarenko, E. S. Estimating the lower
 limit of the impact of amines on nucleation in the earth's atmosphere. *Entropy* 2015,
 17, 2764-2780.
- 605 (60) Paasonen, P.; Olenius, T.; Kupiainen, O.; Kurtén, T.; Petäjä, T.; Birmili, W.;
- Hamed, A.; Hu, M.; Huey, L. G.; Plass-Duelmer, C.; Smith, J. N.; Wiedensohler, A.;
- 607 Loukonen, V.; McGrath, M. J.; Ortega, I. K.; Laaksonen, A.; Vehkamäki, H.;
- 608 Kerminen, V. M.; Kulmala, M. On the formation of sulphuric acid -amine clusters in 609 varying atmospheric conditions and its influence on atmospheric new particle
- 610 formation. *Atmos. Chem. Phys.* **2012**, 12, 9113-9133.
- 611 (61) Jen, C. N.; McMurry, P. H.; Hanson, D. R. Stabilization of sulfuric acid dimers by
- 612 ammonia, methylamine, dimethylamine, and trimethylamine. J. Geophys. Res-Atmos.
- 613 **2014**, 119, 7502-7514.

- 614 (62) Almeida, J.; Schobesberger, S.; Kurten, A.; Ortega, I. K.; Kupiainen-Maatta, O.;
- 615 Praplan, A. P.; Adamov, A.; Amorim, A.; Bianchi, F.; Breitenlechner, M.; David, A.;
- 616 Dommen, J.; Donahue, N. M.; Downard, A.; Dunne, E.; Duplissy, J.; Ehrhart, S.;
- Flagan, R. C.; Franchin, A.; Guida, R.; Hakala, J.; Hansel, A.; Heinritzi, M.; Henschel,
 H.; Jokinen, T.; Junninen, H.; Kajos, M.; Kangasluoma, J.; Keskinen, H.; Kupc, A.;
- H.; Jokinen, T.; Junninen, H.; Kajos, M.; Kangasluoma, J.; Keskinen, H.; Kupc, A.;
 Kurten, T.; Kvashin, A. N.; Laaksonen, A.; Lehtipalo, K.; Leiminger, M.; Leppa, J.;
- Loukonen, V.; Makhmutov, V.; Mathot, S.; McGrath, M. J.; Nieminen, T.; Olenius, T.;
- 621 Onnela, A.; Petaja, T.; Riccobono, F.; Riipinen, I.; Rissanen, M.; Rondo, L.;
- 622 Ruuskanen, T.; Santos, F. D.; Sarnela, N.; Schallhart, S.; Schnitzhofer, R.; Seinfeld, J.
- 623 H.; Simon, M.; Sipila, M.; Stozhkov, Y.; Stratmann, F.; Tome, A.; Trostl, J.;
- 624 Tsagkogeorgas, G.; Vaattovaara, P.; Viisanen, Y.; Virtanen, A.; Vrtala, A.; Wagner, P.
- 625 E.; Weingartner, E.; Wex, H.; Williamson, C.; Wimmer, D.; Ye, P.; Yli-Juuti, T.; 626 Carslaw, K. S.; Kulmala, M.; Curtius, J.; Baltensperger, U.; Worsnop, D. R.;
- 627 Vehkamäki, H.; Kirkby, J. Molecular understanding of sulphuric acid-amine particle
 628 nucleation in the atmosphere. *Nature* 2013, 502, 359-363.
- 629 (63) Jen, C. N.; Bachman, R.; Zhao, J.; McMurry, P. H.; Hanson, D. R. Diamine-
- sulfuric acid reactions are a potent source of new particle formation. *Geophys. Res. Lett.* **2016**, 43, 867-873.
- 632 (64) Wang, C. Y.; Jiang, S.; Liu, Y. R.; Wen, H.; Wang, Z. Q.; Han, Y. J.; Huang, T.;
 633 Huang, W. Synergistic effect of ammonia and methylamine on nucleation in the earth's
- atmosphere. A Theoretical Study. J. Phys. Chem. A **2018**, 122, 3470-3479.
- 635 (65) Lv, S. S.; Miao, S. K.; Ma, Y.; Zhang, M. M.; Wen, Y.; Wang, C. Y.; Zhu, Y. P.;
- Huang, W. Properties and atmospheric implication of methylamine-sulfuric acid-water
 clusters. J. Phys. Chem. A 2015, 119, 8657-8666.
- 638 (66) Ma, Y.; Chen, J.; Jiang, S.; Liu, Y.-R.; Huang, T.; Miao, S.-K.; Wang, C.-Y.;
- 639 Huang, W. Characterization of the nucleation precursor $(H_2SO_4-(CH_3)_2NH)$ complex:
- 640 intra-cluster interactions and atmospheric relevance. *RSC Adv.* **2016**, 6, 5824-5836.
- 641 (67) Qiu, C.; Zhang, R. Multiphase chemistry of atmospheric amines. *Phys. Chem.*642 *Chem. Phys.* 2013, 15, 5738-5752.
- 643 (68) Zhang, R.; Wang, G.; Guo, S.; Zamora, M. L.; Ying, Q.; Lin, Y.; Wang, W.; Hu,
- M.; Wang, Y. Formation of urban fine particulate matter. *Chem. Rev.* 2015, 115, 3803-3855.
- 646 (69) Elm, J. Elucidating the Limiting Steps in Sulfuric Acid-Base New Particle 647 Formation. J. Phys. Chem. A **2017**, 121, 8288-8295.
- 648 (70) Xie, H. B.; Elm, J.; Halonen, R.; Myllys, N.; Kurtén, T.; Kulmala, M.; Vehkamäki,
- 649 H. Atmospheric Fate of Monoethanolamine: Enhancing new particle formation of
- (71) Yao, L.; Garmash, O.; Bianchi, F.; Zheng, J.; Yan, C.; Kontkanen, J.; Junninen,
 H.; Mazon, S. B.; Ehn, M.; Paasonen, P.; Sipila, M.; Wang, M. Y.; Wang, X. K.; Xiao,
- 654 S.; Chen, H. F.; Lu, Y. Q.; Zhang, B. W.; Wang, D. F.; Fu, Q. Y.; Geng, F. H.; Li, L.;
- 655 Wang, H. L.; Qiao, L. P.; Yang, X.; Chen, J. M.; Kerminen, V. M.; Petaja, T.; Worsnop,
- 656 D. R.; Kulmala, M.; Wang, L. Atmospheric new particle formation from sulfuric acid
- and amines in a Chinese megacity. *Science* **2018**, 361, 278.
- 658 (72) Ma, X.; Sun, Y.; Huang, Z.; Zhang, Q.; Wang, W. A density functional theory
- study of the molecular interactions between a series of amides and sulfuric acid. *Chemosphere* 2019, 214, 781-790.
- 661 (73) Rochelle, G.; Chen, E.; Freeman, S.; Van Wagener, D.; Xu, Q.; Voice, A. Aqueous
- piperazine as the new standard for CO₂ capture technology. *Chem. Eng. J.* **2011**, 171,
- *663* 725-733.

- 664 (74) Fine, N. A.; Nielsen, P. T.; Rochelle, G. T. Decomposition of nitrosamines in CO₂
 665 capture by aqueous piperazine or monoethanolamine. *Environ. Sci. Technol.* 2014, 48,
 666 5996-6002.
- 667 (75) *Kirk-Othmer Encyclopedia of chemical technology*, 3rd ed., Vols 1-26. John Wiley 668 and Sons: New York, NY, 1978-1984; 1978; Vol. V2, p296.
- 669 (76) Akyüz, M. Simultaneous determination of aliphatic and aromatic amines in
 670 ambient air and airborne particulate matters by gas chromatography-mass spectrometry.
 671 *Atmos. Environ.* 2008, 42, 3809-3819.
- (77) Hunter, E. P. L.; Lias, S. G. Evaluated gas phase basicities and proton affinities of
 molecules: An update. *J. Phys. Chem. Ref. Data* 1998, 27, 413-656.
- 674 (78) Elm, J.; Passananti, M.; Kurtén, T.; Vehkamäki, H. Diamines can initiate new particle formation in the atmosphere. *J. Phys. Chem. A* **2017**, 121, 6155-6164.
- 676 (79) Olenius, T.; Kupiainen-Maatta, O.; Ortega, I. K.; Kurtén, T.; Vehkamaki, H. Free
- 677 energy barrier in the growth of sulfuric acid-ammonia and sulfuric acid-dimethylamine
 678 clusters. J. Chem. Phys. 2013, 139, 084312.
- 679 (80) McGrath, M. J.; Olenius, T.; Ortega, I. K.; Loukonen, V.; Paasonen, P.; Kurtén,
- 680 T.; Kulmala, M.; Vehkamäki, H. Atmospheric cluster dynamics code: a flexible method
- for solution of the birth-death equations. *Atmos. Chem. Phys.* **2012**, 12, 2345-2355.
- 682 (81) Ortega, I. K.; Kupiainen, O.; Kurtén, T.; Olenius, T.; Wilkman, O.; McGrath, M.
- 583 J.; Loukonen, V.; Vehkamäki, H. From quantum chemical formation free energies to 684 evaporation rates. *Atmos. Chem. Phys.* **2012**, 12, 225-235.
- 685 (82) Shi, X.; Zhang, R.; Sun, Y.; Xu, F.; Zhang, Q.; Wang, W. A density functional
- theory study of aldehydes and their atmospheric products participating in nucleation. *Phys. Chem. Chem. Phys.* 2018, 20, 1005-1011.
- (83) Elm, J.; Jen, C. N.; Kurtén, T.; Vehkamäki, H. Strong hydrogen bonded molecular
 interactions between atmospheric diamines and sulfuric acid. *J. Phys. Chem. A* 2016,
 120, 3693-3700.
- 691 (84) Myllys, N.; Ponkkonen, T.; Passananti, M.; Elm, J.; Vehkamäki, H.; Olenius, T.
- Guanidine: A highly efficient stabilizer in atmospheric new-particle formation. J. Phys.
 Chem. A 2018, 122, 4717-4729.
- 694 (85) Myllys, N.; Olenius, T.; Kurtén, T.; Vehkamäki, H.; Riipinen, I.; Elm, J. Effect of
 695 bisulfate, ammonia, and ammonium on the clustering of organic acids and sulfuric
 696 Acid. J. Phys. Chem. A 2017, 121, 4812-4824.
- 697 (86) Elm, J.; Myllys, N.; Hyttinen, N.; Kurtén, T. Computational study of the clustering
- 698 of a cyclohexene autoxidation product $C_6H_8O_7$ with itself and sulfuric acid. J. Phys. 699 Chem. A **2015**, 119, 8414-8421.
- 700 (87) Li, H.; Zhong, J.; Vehkamäki, H.; Kurtén, T.; Wang, W.; Ge, M.; Zhang, S.; Li,
- 701 Z.; Zhang, X.; Francisco, J. S.; Zeng, X. C. Self-catalytic reaction of SO₃ and NH₃ to 702 produce sulfamic acid and its implication to atmospheric particle formation. *J. Am.*
- 702 produce suffamile acid and its implication to atmospheric particle formation. *J. Am.* 703 *Chem. Soc.* **2018**, 140, 11020-11028.
- 704 (88) Liu, L.; Kupiainen-Maatta, O.; Zhang, H.; Li, H.; Zhong, J.; Kurten, T.;
- Vehkamäki, H.; Zhang, S.; Zhang, Y.; Ge, M.; Zhang, X.; Li, Z. Clustering mechanism
 of oxocarboxylic acids involving hydration reaction: implications for the atmospheric
 models. J. Chem. Phys. 2018, 148, 214303.
- 708 (89) F Frisch, M. J.; Trucks, G. W.; Schlegel, H. B.; Scuseria, G. E.; Robb, M. A.;
- 709 Cheeseman, J. R.; Scalmani, G.; Barone, V.; Mennucci, B.; Petersson, G. A.; Nakatsuji,
- 710 H.; Caricato, M.; Li, X.; Hratchian, H. P.; Izmaylov, A. F.; Bloino, J.; Zheng, G.;
- 711 Sonnenberg, J. L.; Hada, M.; Ehara, M.; Toyota, K.; Fukuda, R.; Hasegawa, J.; Ishida,
- 712 M.; Nakajima, T.; Honda, Y.; Kitao, O.; Nakai, H.; Vreven, T.; Montgomery, J. A., Jr.;
- 713 Peralta, J. E.; Ogliaro, F.; Bearpark, M.; Heyd, J. J.; Brothers, E.; Kudin, K. N.;

- 714 Staroverov, V. N.; Kobayashi, R.; Normand, J.; Raghavachari, K.; Rendell, A.; Burant,
- 715 J. C.; Iyengar, S. S.; Tomasi, J.; Cossi, M.; Rega, N.; Millam, J. M.; Klene, M.; Knox,
- 716 J. E.; Cross, J. B.; Bakken, V.; Adamo, C.; Jaramillo, J.; Gomperts, R.; Stratmann, R.
- E.; Yazyev, O.; Austin, A. J.; Cammi, R.; Pomelli, C.; Ochterski, J. W.; Martin, R. L.;
- 718 Morokuma, K.; Zakrzewski, V. G.; Voth, G. A.; Salvador, P.; Dannenberg, J. J.;
- Dapprich, S.; Daniels, A. D.; Farkas, O.; Foresman, J. B.; Ortiz, J. V.; Cioslowski, J.;
 Fox, D. J. *Gaussian 09*; Gaussian, Inc.: Wallingford, CT, 2009.
- (90) Neese, F. The ORCA program system. *Wiley Interdiscip. Rev. Comput. Mol. Sci.*2012, 2, 73-78.
- 723 (91) Elm, J.; Kristensen, K. Basis set convergence of the binding energies of strongly
- hydrogen-bonded atmospheric clusters. *Phys. Chem. Chem. Phys.* **2017**, 19, 1122-1133.
- (92) Elm, J.; Bilde, M.; Mikkelsen, K. V. Assessment of binding energies of
 atmospherically relevant clusters. *Phys. Chem. Chem. Phys.* 2013, 15, 16442-16445.
- (93) Loukonen, V.; Kurtén, T.; Ortega, I. K.; Vehkamäki, H.; Padua, A. A. H.; Sellegri,
 K.; Kulmala, M. Enhancing effect of dimethylamine in sulfuric acid nucleation in the
- presence of water a computational study. *Atmos. Chem. Phys.* **2010**, 10, 4961-4974.
- 730 (94) Kerminen, V. M.; Petäjä, T.; Manninen, H. E.; Paasonen, P.; Nieminen, T.; Sipilä,
- 731 M.; Junninen, H.; Ehn, M.; Gagné, S.; Laakso, L.; Riipinen, I.; Vehkamäki, H.; Kurten,
- 732 T.; Ortega, I. K.; Dal Maso, M.; Brus, D.; Hyvärinen, A.; Lihavainen, H.; Leppä, J.;
- Lehtinen, K. E. J.; Mirme, A.; Mirme, S.; Hõrrak, U.; Berndt, T.; Stratmann, F.; Birmili,
 W.; Wiedensohler, A.; Metzger, A.; Dommen, J.; Baltensperger, U.; Kiendler-Scharr,
- A.; Mentel, T. F.; Wildt, J.; Winkler, P. M.; Wagner, P. E.; Petzold, A.; Minikin, A.;
- 736 Plass-Dülmer, C.; Pöschl, U.; Laaksonen, A.; Kulmala, M. Atmospheric nucleation:
- highlights of the EUCAARI project and future directions. *Atmos. Chem. Phys.* 2010, 10, 10829-10848.
- 739 (95) Kürten, A.; Li, C.; Bianchi, F.; Curtius, J.; Dias, A.; Donahue, N. M.; Duplissy, J.;
- 740 Flagan, R. C.; Hakala, J.; Jokinen, T.; Kirkby, J.; Kulmala, M.; Laaksonen, A.;
- 741 Lehtipalo, K.; Makhmutov, V.; Onnela, A.; Rissanen, M. P.; Simon, M.; Sipilä, M.;
- 742 Stozhkov, Y.; Tröstl, J.; Ye, P.; McMurry, P. H. New particle formation in the sulfuric
- acid–dimethylamine–water system: reevaluation of CLOUD chamber measurements
 and comparison to an aerosol nucleation and growth model. *Atmos. Chem. Phys.* 2018,
- 745 18, 845-863.
- 746 (96) Tomicic, M.; Bødker Enghoff, M.; Svensmark, H. Experimental study of H₂SO₄
- aerosol nucleation at high ionization levels. *Atmos. Chem. Phys.* **2018**, 18, 5921-5930.
- 748 (97) Zhang, R. Y. Getting to the critical nucleus of aerosol formation. *Science* 2010,
 749 328, 1366-1367.
- 750 (98) Kontkanen, J.; Lehtipalo, K.; Ahonen, L.; Kangasluoma, J.; Manninen, H. E.;
- 751 Hakala, J.; Rose, C.; Sellegri, K.; Xiao, S.; Wang, L.; Qi, X.; Nie, W.; Ding, A.; Yu,
- H.; Lee, S.; Kerminen, V.-M.; Petäjä, T.; Kulmala, M. Measurements of sub-3 nm
 particles using a particle size magnifier in different environments: from clean mountain
- top to polluted megacities. *Atmos. Chem. Phys.* **2017**, 17, 2163-2187.
- 755 (99) Maso, M. D.; Hyvärinen, A.; Komppula, M.; Tunved, P.; Kerminen, V.-M.;
- Lihavainen, H.; Öviisanen, Y.; Hansson, H.-C.; Kulmala, M. Annual and interannual variation in boreal forest aerosol particle number and volume concentration and their connection to particle formation. *Tellus B* **2017**, 60, 495-508.
- 759 (100) Henschel, H.; Kurtén, T.; Vehkamaki, H. Computational study on the effect of 760 hydration on new particle formation in the sulfuric acid/ammonia and sulfuric
- 761 acid/dimethylamine systems. J. Phys. Chem. A **2016**, 120, 1886-1896.
- (101) Li, S.; Qu, K.; Zhao, H.; Ding, L.; Du, L. Clustering of amines and hydrazines in
 atmospheric nucleation. *Chem. Phys.* 2016, 472, 198-207.

- (102) Ling, J.; Ding, X.; Li, Z.; Yang, J. First-principles study of molecular clusters
 formed by nitric acid and ammonia. J. Phys. Chem. A 2017, 121, 661-668.
- (103) Temelso, B.; Morrison, E. F.; Speer, D. L.; Cao, B. C.; Appiah-Padi, N.; Kim,
 G.; Shields, G. C. Effect of mixing ammonia and alkylamines on sulfate aerosol
- 768 formation. J. Phys. Chem. A 2018, 122, 1612-1622.
- 769 (104) Yu, F.; Luo, G. Modeling of gaseous methylamines in the global atmosphere:
- impacts of oxidation and aerosol uptake. *Atmos. Chem. Phys.* 2014, 14, 12455-12464.
- (105) Wang, D. Y. S.; Hildebrandt Ruiz, L. Secondary organic aerosol from chlorine initiated oxidation of isoprene. *Atmos. Chem. Phys.* 2017, 17, 13491-13508.
- (106) Dhulipala, S. V.; Bhandari, S.; Hildebrandt Ruiz, L. Formation of oxidized organic compounds from Cl-initiated oxidation of toluene. *Atmos. Environ.* 2019, 199,
- 775 265-273.
- 776 (107) Le Breton, M.; Hallquist, Å. M.; Pathak, R. K.; Simpson, D.; Wang, Y.;
- Johansson, J.; Zheng, J.; Yang, Y.; Shang, D.; Wang, H.; Liu, Q.; Chan, C.; Wang, T.;
- 778 Bannan, T. J.; Priestley, M.; Percival, C. J.; Shallcross, D. E.; Lu, K.; Guo, S.; Hu, M.;
- Hallquist, M. Chlorine oxidation of VOCs at a semi-rural site in Beijing: significant chlorine liberation from $CINO_2$ and subsequent gas- and particle-phase CI-VOC
- chlorine liberation from ClNO₂ and subsequent gas- and particl
 production. *Atmos. Chem. Phys.* 2018, 18, 13013-13030.
- 782 (108) Zheng, J.; Hu, M.; Zhang, R.; Yue, D.; Wang, Z.; Guo, S.; Li, X.; Bohn, B.; Shao,
- 783 M.; He, L.; Huang, X.; Wiedensohler, A.; Zhu, T. Measurements of gaseous H₂SO₄ by
- AP-ID-CIMS during CAREBeijing 2008 Campaign. Atmos. Chem. Phys. 2011, 11,
 7755-7765.
- 786 (109) Berresheim, H. Gas-aerosol relationships of H₂SO₄, MSA, and OH: Observations
- in the coastal marine boundary layer at Mace Head, Ireland. J. Geophys. Res. 2002,
 107, 5-1–5-12.
- 789 (110) Wingenter, O. W.; Sive, B. C.; Blake, N. J.; Blake, D. R.; Rowland, F. S. Atomic
- 790 chlorine concentrations derived from ethane and hydroxyl measurements over the
- 791 equatorial Pacific Ocean: Implication for dimethyl sulfide and bromine monoxide. J.
- 792 Geophys. Res.-Atmos. 2005, 110, 10.