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LAMONT GEOLOGICAL OBSERVATORY  
PALISADES, NEW YORK

# Technical Report on Seismology No. 38

Crustal Structure of the Arctic Regions  
from the Lg Phase



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Crustal Structure of the Arctic Regions from the

L<sub>g</sub> Phase

by

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## ABSTRACT

The presence of the Lg earthquake surface wave phase is indicative of continental structure between the epicenter and recording station. This fact is used to explore the Arctic regions. Water covered areas of continental structure include the Canadian archipelago, parts of Baffin Bay and Davis Strait, the Barents Sea, Bering Strait and the shallow parts of the Bering Sea and Greenland Sea. Areas of non-continental structure include the Arctic Ocean, Beaufort Sea, Greenland Sea, Norwegian Sea, and the deep parts of the Bering Sea.

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## INTRODUCTION

In a recent paper, Press and Ewing (1952, p. 219) described a seismic surface wave of unusually large amplitudes and characteristic appearance traveling only over paths of a continental nature. Although an adequate theory for this phase is not yet available, discrimination of crustal columns other than that typical of continents against the phase suggested the possibility of using its presence or absence to explore regions of unknown crustal structure. As an example of the method, the Gulf of Mexico was cited (Press and Ewing 1952, p. 228) as an area through which  $L_g$  did not propagate, indicating a non-continental structure for the underlying rocks. This conclusion was subsequently borne out by seismic refraction work. This paper presents the results of a similar study for the Arctic regions.

Press and Ewing assigned the symbol  $L_g$  to the phase,  $L$  because the particle motion was predominantly of the SH or Love type,  $g$  because the waves were thought to be the surface wave counterpart of  $S_g$ , propagated to large distances by a wave guide mechanism in the silicic crust. They presented experimental data for paths across North America, largely using earthquakes along the west coast recorded at Palisades.

Some sample seismograms illustrating the  $L_g$  phase are shown in Figures 1, 2, 3. Its distinctive characteristics are:

1. Sharp commencements with periods usually from 1/2 to 4 seconds, but occasionally as high as 7 seconds. In general, earthquakes slightly offshore or those at very great distances do not produce the shorter periods.

2. Amplitudes far larger than those of any other phase in the same period range for continental paths.

3. Approximately equal amplitudes on all components at the beginning of the phase but greater amplitudes on the transverse horizontal component after a few cycles.

4. A velocity of about 3.5 km/sec. for the beginning. This figure may vary slightly over different paths. It is very close to the shear wave velocity in the upper 15 km of continental crust (Katz 1954).

5. Reverse dispersion, i.e., the shorter periods have the higher group velocities. The periods increase to about 10-15 seconds. The dispersion is not as regular as is the case with Rayleigh waves over oceanic paths.

6. Occasional appearance superposed on the classical long period Love waves.

7. A duration far longer than might be expected for the short period branch of classical Love waves.

8. Occurrence only when the entire path is continental. The insertion of a 100-kilometer segment of oceanic crustal structure is sufficient to prevent the occurrence of  $L_g$  on the seismogram.

This sensitivity of  $L_g$  to horizontal variations of the crustal column forms the basis of this paper. The capacity to distinguish between areas of continental structure and areas of non-continental structure is limited only by the geographical distribution of suitable epicenters and recording stations.

## PROCEDURE

The Arctic was chosen for the first application of the method because it includes several interesting localities of uncertain structure and because of the availability of seismograms with Arctic paths. The collection includes records from the following stations: College, Sitka, Resolute Bay, Palisades, Scoresby Sund, Ivigtut, Copenhagen, Uppsala, Lund and the Berkeley group. At these stations, records from the following instruments were used: College - Wenner EW, Benioff short period Z, McComb-Romberg NS and EW; Sitka - Wenner NS and EW; Resolute Bay - Sprengnether NS, EW, Z, Columbia Z; Palisades - Columbia NS, EW, Z, Benioff long and short period NS, EW, Z; Scoresby Sund - Galitzin-Wilip NS, EW, Z; Ivigtut - Wiechert NS, EW, Grenet Z; Lund - Wiechert NE,-SW, NW-SE; Berkeley - Galitzin NS, EW, Z; Fresno - Sprengnether NS, EW, Z.

Table 1 lists all the paths studied. Each path is designated by a letter which indicates the recording station and a number which indicates the particular shock as listed at that station. With this system the same shock recorded at different stations will have more than one path symbol, e.g. the shock of 12 February 1951 has path symbols P7 to Palisades, C28 to College, and K6 to Sitka.

All records were carefully examined for  $L_g$  and the velocity of the commencement of the phase was calculated. When considerable energy was recorded at the time for  $L_g$  but the distinctive characteristics of  $L_g$ , primarily the high frequencies, were absent, the path was labeled doubtful. No

attempt was made to study variations in velocity, primarily because many of the records do not show good high frequency beginnings. Furthermore, many of the velocities printed here may be approximate, because of the possibility of errors in epicentral location. This is especially true for the shorter paths, e. g. the Alaskan peninsula shocks to College where a small error would be a significant portion of the path.

Some difficulties encountered in application of the method are:

1. The location of many epicenters near shore and offshore, prohibiting their use in exploration of other segments of the path since  $L_g$  excitation is always in doubt. This difficulty may be eliminated by the use of two or more recording stations spanning the segment in question. The station nearer the epicenter then determines the degree of  $L_g$  excitation and the farther one the effect of the intervening segment on the phase.

2. The effect at the source on  $L_g$  excitation, especially the depth of the shock. There are variations in  $L_g$  excitation for shocks listed as normal depth.

3. The discrimination of various types of instruments against parts of the frequency spectrum of  $L_g$ .

4. The effect of mountainous terrain on  $L_g$ . Although this is not thoroughly understood at present, there is meager evidence available which indicates that  $L_g$  will travel a considerable distance through a mountainous region such as the Rocky Mountains (path C-72, Fig. 3).

5. Azimuthal effects at the source. No qualitative data are available on this at present.

It is then apparent that although the presence of  $L_g$  is indisputable proof of continental structure\* along the entire path, the absence of  $L_g$ , par-

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\* Continental structure as used here refers to a silicic crust having a thickness of about 35 km as determined by seismic refractions and near earthquake data, although the  $L_g$  wave guide need be no larger than about 15 km, for the short period oscillations.

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ticularly the high frequencies, cannot be used without caution to deduce the existence of a segment of non-continental structure along the path. In short, the presence of  $L_g$  is sufficient proof of continental structure along the entire path. The absence of  $L_g$  is necessary but not sufficient proof of the presence of a segment of non-continental structure.

## RESULTS

Figures 4 and 5 are maps of the Arctic regions with the great circle paths from epicenter to recording station superimposed. Solid lines indicate that no  $L_g$  was recorded, crossed lines that  $L_g$  was recorded, and dashed lines that the presence of  $L_g$  on the record is doubtful.

It is evident from the figures that the water-covered areas of the Canadian archipelago, Baffin Bay and Davis Strait overlie continental structure, although in the case of the latter, no good data are available for the southern part of the strait where it joins with the North Atlantic Ocean, nor is there conclusive evidence for the small deep areas of Baffin Bay. The results are in agreement with geologic interpretations of the area (Eardley, 1951, p. 538). There is a good  $L_g$  path across Greenland extending about 2 degrees off the northeast coast into shallow water. Because of the

nature of the  $L_g$  phase, this is not likely to shed any light on the current theories of the elevation (Nye 1952, p. 529) of the rock floor of Greenland other than to indicate that whatever the elevation of its surface, the crustal column beneath is continental.

Path G1 to Copenhagen traverses the broad continental shelf underlying the Barents Sea clearly indicating continental structure.

In no case is  $L_g$  recorded for paths that traverse the Arctic Ocean basin proper. This evidence is highly favorable to a crustal structure other than continental for the Arctic Ocean basin. To supplement this information, records exhibiting Rayleigh wave dispersion across the Arctic Ocean were examined. In such a study it is highly advantageous to choose a path which has as small a continental segment and as large an oceanic segment as possible. College, Alaska is, at present, the best station from this point of view. However, current instrumentation there is poor for this purpose as the only long period instrument measures the EW or primarily transverse component of ground motion for most Arctic quakes. However, one set of records is available for which NS and EW long period instruments were recording. The dispersion curve of this quake, path C-74, as well as some points from waves tentatively identified as the Rayleigh type on the EW component for paths C-24 and C-10, are shown in Figure 6. Corrections have been made from the curve of Brilliant and Ewing (1954) for the land travel in Alaska and the Alaskan continental shelf. Although there is considerable scatter, note that the curve falls below the oceanic curve of Press and Ewing (1950 p. 275) but above the continental curve of Brilliant and Ewing

(1954 p. 149). This apparently indicates a structure of intermediate nature for the Arctic Ocean basin, but the Rayleigh wave data are so poor that a final decision is not warranted here. Furthermore, recent sounding data in the Arctic by Crary et al. (1952) show that considerable changes need to be made in the map of Emery on which the continental corrections are based. Most certainly, however, the Arctic basin is not underlain by a continental column of rocks. Its identification with the oceanic column of the Atlantic and Pacific or with some intermediate structure awaits evidence which is not readily available at present. Eardley (1951 p. 538) has interpreted a continental structure underlying the Arctic Ocean from geological evidence. Such evidence, he admits, is quite scanty. Gutenberg (1941 p. 27) from PP/P ratios found the Arctic Ocean floor to be similar to that of the Pacific, although Byerly, Mei, and Romney (1949 p. 269) and Mei (1943 p. 149) have pointed out that supplemental studies on directionality of the source are desirable.

No  $L_g$  phases were recorded for any of the paths crossing the Greenland Sea-Norwegian Sea area in the vicinity of Jan Mayen, with the possible exception of a doubtful one at Resolute Bay. Although the Jan Mayen shocks are sometimes thought to be of slightly greater than normal depth, additional evidence for paths G-5 and S-19 favors a non-continental structure for this area.

No  $L_g$ s were recorded from the North Atlantic shocks.

The shocks along the Alaskan peninsula and the Aleutian arc exhibit a very consistent set of data. Beginning with the easternmost shock on

the peninsula and continuing clockwise in azimuth around College to path C-9 it is apparent that most shocks made the  $L_g$  phase. Two did not but they are far enough from shore to be inefficient in  $L_g$  excitation. Path C-9 which is approximately on the margin of the deep part of the Bering Sea made a doubtful  $L_g$ . The remainder of the shocks to the west whose paths cross the deep Bering Sea did not make  $L_g$  at College. The consistency of the results indicates with certainty that the crustal structure of the deep Bering Sea is markedly different from that of the continent.

The large group of Japanese, Kurile and Kamchatka shocks failed to produce an  $L_g$  at College. However, the complicated nature of the path does not permit any specific conclusions to be drawn since most of the paths cross either the Pacific Ocean or the Okhotsk Sea in addition to the Bering Sea.

Most of the shocks whose paths cross the Bering Strait did not make an  $L_g$  at College. However, one shock, C-68 did, although the very high frequencies are missing. In an effort to isolate the reason for the missing  $L_g$ 's, shocks corresponding to paths C-28, L-1 and G-9, and C-32, L-2 and G-10 were examined at Copenhagen and Lund. Neither made an  $L_g$  at College. The former made a very doubtful  $L_g$  at both European stations, the latter made a weak one. This evidence tends to place the reason for the missing  $L_g$  at College on inefficient excitation at the source rather than on a structural anomaly at the Bering Strait. Furthermore, for shock C-68 which did send an  $L_g$  to College, Uppsala also reports  $L_g$  in its station

bulletin. On the other hand, for shock C-67, Uppsala does not report  $L_g$ , and neither is it recorded at College. The conclusion, then, is that continental structure is continuous, at least across the narrow part of the Bering Strait. This is in agreement with the geologic interpretations of the area (Eardley 1951, p. 532).

There is an excellent correlation between water depth and areas through which  $L_g$  will not propagate. In no case is there  $L_g$  propagation beneath water deeper than 1000 fathoms for any great distance. The possibility that the absence of  $L_g$  over these paths might be due to the water column rather than the underlying rocks is ruled out because, although it is present on all three components,  $L_g$  is primarily a shear wave polarized horizontally. This component of motion could not be affected by the water. Thus it appears that water depths actually reflect changes of the crustal column. Indeed, geophysical methods of any sort have yet to detect an area where continental structure underlies oceanic depths.

#### SUMMARY

The  $L_g$  surface wave phase was used to distinguish areas of continental crustal structure from areas of non-continental structure in the Arctic regions. Water-covered areas of continental structure include the Canadian archipelago, parts of Baffin Bay and Davis Strait, the shallow parts of the Bering Sea and of the Greenland Sea, the Barents Sea and Bering Strait. Areas of non-continental structure include the Arctic Ocean, Beaufort Sea, Greenland Sea, Norwegian Sea and the deep Bering Sea. Nowhere is continental structure detected beneath oceanic depths.

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TABLE I

## MASTER LIST OF SHOCKS USED IN ARCTIC STUDY

## BERKELEY GROUP (B)

No.	Date D-M-Y	Hour	Lat.	Long.	Geograph- ical loc.	Depth	Mag.	in km	Lg	sec	Comp	Remarks
B1	22-4-51	12 36 16	76°N	73°W	Baffin Bay		4500	Yes	3. 5	Z		Mineral - poor data for Lg
							4810	Yes	3. 45	Z		Mt. Hamilton - poor data for Lg
B2	6-6-51	16 10 52	71 1/2N	8°W	Jan Mayen	60	7	No				Mineral Fresno

TABLE I

## MASTER LIST OF SHOCKS USED IN ARCTIC STUDY

## (COLLEGE - C)

No.	Date D-M-Y	Hour 23 33 38	Lat. 22N	Long. 100 1/2E	Geographical Location S. Yunnan Prov. China	Depth 6 3/4	Mag. in km Lg sec	Lg Vel in km/ sec	Comp.	Remarks
C1	2-2-50	23 33 38							No	
C2	3-2-50	16 45 29*	54N	162W	Aleutians		1440	Yes 3.23	E	Excellent Lg - Long Period Preceding
C3	4-2-50	02 07 53*	54N	162W	Aleutians		1440	Yes 3.27	E	Same as above
C4	7-2-50	10 37 22	46N	152E	Kurile Isl.			No		
C5	18-2-50	14 39 30	54N	164W	Aleutians		1500	Yes 3.26	E	
C6	27-3-50	13 04 04	53 1/2N	173E	Aleutians			No		
C7	4-4-50	18 44 10	52N	101E	Border USSR & Outer Mongolia		6 1/2	No		S. P. Benioff only
C8	5-4-50	01 17 15	52N	177W	Aleutians			No		S. P. Benioff only
C9	26-4-50	12 18 28	53N	170W	Aleutians			Possibly	Z, E	No H. F.
C10	5-6-50	11 16 12	87N	45E	No. Polar Region			No		Good dispersed train of surface waves
C11	27-6-50	15 41 54	45 1/2N	140E	Off Hokkaido		6 1/2			No

TABLE I (Cont'd)  
(COLLEGE - 2)

No.	Date D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag.	in km sec	Lg Comp.	Lg Vel. in km/ No	Remarks
C12	9-7-50	16 09 55	36N	72E	Pakistan						
C13	12-7-50	11 09 15	53N	166W	Aleutians		6 1/4	1400	Yes 3.48	E	
C14	14-7-50	12 06 45	52N	171W	Aleutians				No		
C15	8-8-50	05 12 00	55N	134 1/2W	S. E. Alaska		1250	Possibly	E	Not a good Lg	
C16	15-8-50	14 09 30	28 1/2N	97E	CBI Border		8 1/2		No		
		Aftershocks									
C17	22-8-50	07 40 07	53N	160E	Off Kamchatka				No		
C18	2-9-50	02 47 23	52 1/2N	169W	Aleutians		100	6 3/4	1820	Yes 3.52	E
C19	10-9-50	03 21 20	35N	140 E	Near Honshu			6 3/4		No	
C20	16-9-50	21 58 15	52 1/2N	178E	Aleutians		100	6.6	2180	No	
C21	28-9-50	21 47 01	54 1/2N	134 1/2W	Qu. Charlotte Isl.			5900	Possibly	E	
C22	5-11-50	17 37 25	33N	134 1/2E	Off Shikoku, Japan				No		
C23	22-11-50	10 16 26*	51N	176W	Aleutians				No		
C24	9-1-51	16 00 24	81N	122E	Arctic Ocean			6 3/4-6 1/2	No		Good dispersed train of surface waves

TABLE I (Cont'd)

(COLLEGE - 3)

No.	Date	D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag.	Lg Vel. in km/sec.			Remarks
									No			
C 25	15-1-51	01 50 36		51N	178 1/2E	Aleutians						
		02 29 46										
C 26	18-1-51	21 15 50		52N	177W	Aleutians						No
C 27	24-1-51	07 17 01		33N	115 3/4	California	6					No
C 28	12-2-51	17 22 02		66N	136E	Siberia						No
C 29	5-3-51	14 46 00		53N	163W	S. of Alaska Pen.						No
C 30	5-3-51	20 11 45		29N	128E	Ryukyu Isl.	150	7				No
C 31	19-3-51	20 28 55		57N	160E	Kamchatka						No
C 32	14-4-51	13 32 59		61N	136E	Siberia						No
C 33	22-4-51	12 36 16		76N	73W	Baffin Bay	2790	Doubtful 3.27	E	Weak		
C 34	29-4-51	07 35 46		80 1/2N	121E	Arctic Ocean						No
C 35	25-4-51	21 59 35		81N	121E	Arctic Ocean						No
C 37	10-3-51	19 44 52		51N	180W	Aleutians	60					
		19 47 14										
C 38	5-6-51	16 57 47		30N	132E	Kyushu	100	6 3/4-7				No
C 39	6-6-51	16 10 52		71 1/2N	8W	Jan Mayen	60	7	4640	Doubtful 3.34	E	No H. F.

TABLE I (Cont'd)  
(COLLEGE - 4)

TABLE I (Cont'd)

(COLLEGE - 5)

No.	Date	D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag.	Lg	Comp.	Remarks
				54 1 / 2N	160 W	Off Alaska Pen.		6 1 / 2			
									in km	sec	
C55	8-11-51	13 45 39									No
C56	12-11-51	08 09 26	47N	154E	Kuriles			6 1 / 2			No
C57	15-11-51	19 42 12	52 1 / 2N	160 1 / 2E	Kamchatka	60	6 1 / 4				No
C58	18-11-51	09 35 43	31N	90 1 / 2E	E. Tibet		7 1 / 2				No
C59	24-11-51	04 30 24	57 1 / 2N	176 1 / 2W	Aleutians	60					No
C60	21-12-51	08 37 28	26 1 / 2N	100E	Yunnan, China						No
C61	25-12-51	15 58 28	49N	155 1 / 2E	Kuriles	60					No
C62	12-3-52	12 13 10	64N	22W	W. Coast of Iceland						No
C63	19-3-52	09 04 18	41N	125E	Korea						No
C64	1-4-52	00 37 41. 5	48. 0N	113. 8W	N. W. Montana		2750	Yes 3. 50	E	Weak	
C65	19-6-52	12 12 56	23N	100E	S. Yunnan Prov. China		6 1 / 2		No		
C66	17-8-52	16 02 05	30 1 / 2N	91 1 / 2E	Tibet		7 1 / 4				No
C67	14-9-52	09 34 10	93 1 / 2E	34N	Shanghai, China						No
C68	30-9-52	12 52 00	28 1 / 2N	102E	China		6 1 / 2	8030	Yes 3. 66	E	No H. F. Also recorded at Uppsala

TABLE I (Cont'd)

## (COLLEGE - 6)

No.	Date D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag.	in km Lg. sec.	Lg. Vel. in km /	Comp.	Remarks
C69	10-10-52	18 47 37	30 1/2N	69E	Cent. Pakistan			No			
C70	8-12-52	15 09 30	23N	99 1/2E	CBI Border			No			
C71	10-12-52	05 58 06	71N	7W	Jan Mayen Isl.			No			Looks slightly deep, perhaps 100 km
C72	23-11-47	09 46 05. 5	47°47'N	112°02'W	S. W. Montana			3120	Yes 3. 59	N, E	Excellent Lg
C73	28-1-48	15 51. 3	38N	68E	Turkistan			No			
C74	18-2-48	20 29. 8	82N	43E	Arctic Region			6 3/4		No	
C75	25-5-48	07 11. 3	30N	99 1/2E	Siking Prov. China			7 1/4		No	
C76	28-9-48	21 36. 6	23N	94E	Burma			No			

TABLE I (Cont'd)

## (COPENHAGEN - G)

No.	Date	D-M-Y	Hour	Lat.	Long.	Mag.	in km sec.	Lg 3.68	Lg Vel. in km/	Remarks
G1	18-2-48	20 29.8	82N	43E	Arctic Region	3080	Yes	3.68	Excellent Lg	
G2	26-4-48	09 32 20	51N	34W	N. Atlantic	3070	No		Good Rayleigh wave dispersion	
G3	8-7-48	12 34.6	71N	6W	Near Jan Mayen Land	1920	No			
G4	28-8-48	02 27.8	57N	161E	Kamchatka	7170	No			
G5	2-3-49	06 54.6	72N	3W	Arctic Ocean	2290	No			
G6	19-4-49	15 19.2	48N	154E	Kuriles	7950	No			
G7	27-9-49	15 30 43	60N	149W	S. Alaska	4930	No			
G8	6-6-51	16 10 52	71 1/2N	8W	Jan Mayen	60	7	2600	No	
G9	12-2-51	17 22 02	66N	136E	Siberia	5675	Doubtful 3.27	N	Weak if present	
G10	14-4-51	13 32 39	61N	136E	Siberia	6120	Yes	3.38	N, E, Z No H. F.	

TABLE I (Cont'd)

(SITKA - K)

No.	Date	D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag:	in km	Lg sec	Comp.	Remarks
K1	2-2-50	23 33 38	22N	100 1/2E	S. Yunnan Prov.	6 3/4	No					
K2	4-4-50	18 44 10	52N	101E	Border USSR & China	6 1/2	No					
K3	5-6-50	11 16 12	87N	45E	N. Polar Region	No						
K4	9-7-50	16 09 53	36N	72E	Pakistan	No						
K5	26-8-50	04 39 27	65N	162W	150 mi. NW of Nome	6 1/2	No					
K6	12-2-51	17 22 02	66N	136E	Siberia	No						
K7	14-4-51	13 32 59	61N	136E	Siberia	6 3/4	3650	No				
K8	22-4-51	12 36 16	76N	73W	Baffin Bay	3200	Yes 3. 30	N, E				
K9	6-6-51	16 10 52	71 1/2N	8W	Jan Mayen	60	7	No				
K10	13-8-51	18 33 40	43N	32 1/2E	Black Sea	6 1/2	No					
K11	18-11-51	09 26 33	31N	90 1/2E	E. Tibet	7 1/2	No					
K12	21-12-51	08 37 28	26 1/2N	100E	Yunnan, China	No						
K13	31-10-55	18 37 49	46 1/2N	112W	Helena, Mont.	1790	Yes 3. 55					
												Periods less than 4 sec

TABLE I (Cont'd)

(LUND - L)

No.	Date D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag. in km	Lg sec.	Comp	Remarks
L1	12-2-51	17 22 02	66N	136E	Siberia		5650	Doubtful		
L2	14-4-51	13 32 59	61N	136E	Siberia		6100	Yes 3.44	NW, NE	No H. F.

TABLE I (Cont'd)  
(PALISADES - P)

No.	Date D-M-Y	Hour 12-12-56*	Lat. 23N	Long. 100E	Geographical Location S. Yunnan Prov. China	Mag. 6 1/2	in km 12950	Lg sec	Comp.	Remarks	Lg Vel. in km/ No
P1	19-6-52	12-12-56*	23N	100E	Eastern Tibet	11950	No				
P2	17-8-52	16 02 05*	30 1/2N	91 1/2E	Shanghai Prov.	11600	No				
P3	14-9-52	09 34 10*	34N	93 1/2E	Szechwan Prov. China	6 1/2	12300	No			
P4	30-9-52	12 52 00*	28 1/2N	102E	N. Afghanistan	10600	No				
P5	27-11-52	07 20 13*	37N	70E	Weak at Pali- sades						
P6	8-12-52	15 09 30*	23N	99 1/2E	China-Burma Border	12800	No				
P7	12-2-51	17 22 02	66N	136E	Near Verkhoyanski Mts., Siberia	6 1/2	No				
P8	14-4-51	13 32 59*	61N	136E	" " "	6 3/4	8350	No			
P9	22-4-51	12 36 16*	76N	73W	Baffin Bay	3900	Yes 3. 58	Z			
P10	29-4-51	07 35 46*	80 1/2N	121E	Arctic Ocean	6430	No				
P11	6-6-51	16 10 52	71 1/2N	8W	Jan Mayen	60	7				
P12	12-2-53	04 31 16*	65N	133W	Yukon, Canada	4550	Yes 3. 57	N	Excellent Lg		

TABLE I (Cont'd)

## (RESOLUTE BAY - R)

No.	Date	Geographical Location	Depth	Mag.	in km	Lg sec	Comp.	Remarks
	D-M-Y				610	Yes 3.76		Velocity doubtful
	Hour	Lat.	Long.					in km/
R1	22-4-51	12 36 16	76N	73W	Baffin Bay			
R2	6-6-51	16 10 52	71 1/2N	8W	Jan Mayen	60	7	Doubtful 3.60
R3	27-9-51	19 24 12	49N	129W	Off Vancouver, B.C.	3200	Yes 3.54	Z
R4	20-9-51	12 38 40	65N	154W	Central Alaska	2350	Yes 3.46	E
R5	12-2-51	17 22 02	66N	136E	Siberia	6 1/2	No	

TABLE I (Cont'd)  
 (SCORESBY SUND - S)

No.	Date D-M-Y	Hour	Lat.	Long.	Mag.	in km	Lg sec	Comp.	Remarks
	1-3-29	07 31 0	50. 2N	130. 7W	5440	Doubtful	E	No H. F. - not a good Lg	
S1								3. 63	
S2	4-7-29	04 28 28	63. 2N	147. 3W	Central Alaska	4550	No		
S3	4-7-29	07 14 18	55. 7N	35. 5W	N. Atlantic Ocean	1770	No		Good train of dispersed sur- face waves
S4	4-7-29	07 31 10	55. 7N	35. 5W	N. Atlantic Ocean	1770	No		Good train of dispersed sur- face waves
S5	4-7-29	05 56 32	55. 7N	35. 5W	N. Atlantic Ocean	1770	No		Good train of dispersed sur- face waves
S6	25-3-32	23 58 39	62. 5N	153. 3W	Near Anchorage, Alaska	4760	Doubtful	E	Not a sharp Lg
S7	28-7-34	21 37 04	55N	156. 8W	Off Alaska Peninsula	5490	No		
S8	24-9-35	22 12 25	49. 3N	129. 2W	Off Vancouver Island	5490	Yes 3. 32		S. P. missing - L. P. insts. may have missed be- ginning
S9	22-3-38	15 22 11	57. 6N	132. 1W	Off. Qu. Charlotte Islands	5270	Yes 3. 55	E	S. P. missing

TABLE I (Cont'd)

## (SCORESBY SUND - 2)

No.	Date	D-M-Y	Hour	Lat.	Long.	Geographical Location	Depth	Mag	in km	Lg sec	Comp.	Remarks
S10	29-5-40	01 57 46		66.6N	135.4W	Near MacKenzie River		3990	Yes	3.42	Z	Excellent - very H.F. some L.P.
S11	30-7-41	01 51.5		60.9N	149.2W	Near Anchorage		4870	No			
S12	28-2-48	01 58 05		53.5N	133W	Qu. Charlotte Islands		5450	No			
S13	27-6-48	21 39.2		56N	158W	S. of Alaska Peninsula		5520	No			
S14	23-8-49	20 24 32		53N	132W	Off Brit. Columbia		5210	Yes	3.52	N	Good Lg
S15	31-10-49	01 39 32		56N	135W	Sitka, Alaska		5000	Yes	3.49	E	Good Lg but not very H.F.
S16	27-3-50	13 04 04		53.5N	173E	Aleutians		6290	No			
S17	6-4-50	21 48 02		49N	129W	Vancouver, B.C.		5530	No			
S18	5-6-50	11 16 16		87N	50E	North Polar Region		2080	No			
S19	20-6-50	14 11 45		74.5N	80E	Jan Mayen		1080	No			Good train of dispersed waves
S20	12-7-50	11 09 15		53N	166W	Aleutians		5990	No			
S21	12-2-51	17 22 02		66N	136E	Siberia			No			
S22	14-4-51	13 32 59		61N	136E	Siberia			No			

TABLE I (Cont'd)  
 (UPPSALA - U)

No.	Date D-M-Y	Hour	Lat. 71 1/2N	Long. 8W	Geographical Location	Depth	Mag.	in km	Lg	Remarks		
										U1	6-6-51	16 10 52
(IVIGTUT - V)												
No.	Date D-M-Y	Hour	Lat. 81N	Long. 56N	Geographical Location	Depth	Mag.	in km	Lg	Lg Vel in km/ sec	Comp.	Remarks
V1	15-4-45	02 35.2		164E	Off Kamchatka			5840	No			
V2	8-11-45	09 05 41	81N	7W	Greenland Sea			2560	Yes 3.73	N		Good H. F.
V3	8-11-45	10 62.5	81N	7W	Greenland Sea			2560	Yes 3.66	N		
V4	12-1-46	20 25.7	59N	147.5W	Gulf of Alaska			4970	No			
V5	1-4-46	12 28.9	54N	164W	Aleutians			6040	No			
V6	16-10-47	02 09 45	64.5N	148.8W	40 mi. S. W. of Fairbanks			4660	Yes 3.68	Z		Good H. F.

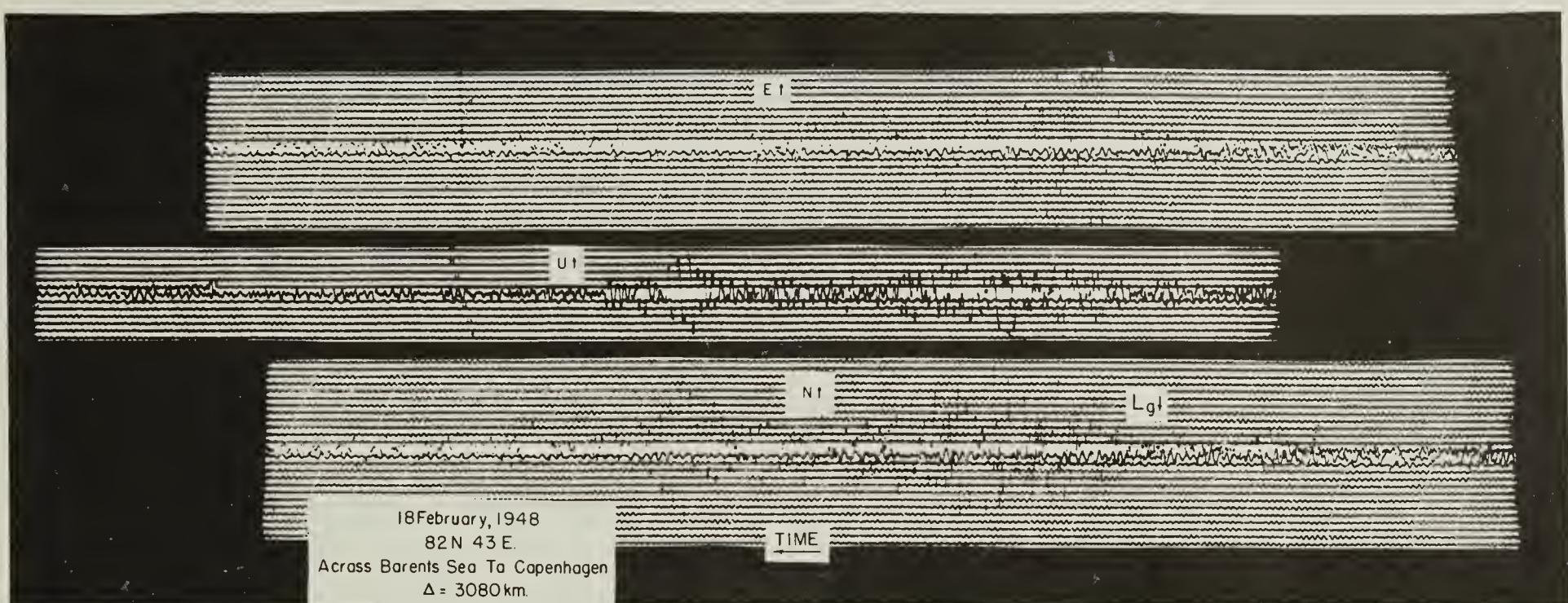


Figure 1. Seismograms:  $L_g$  across Barents Sea to Copenhagen

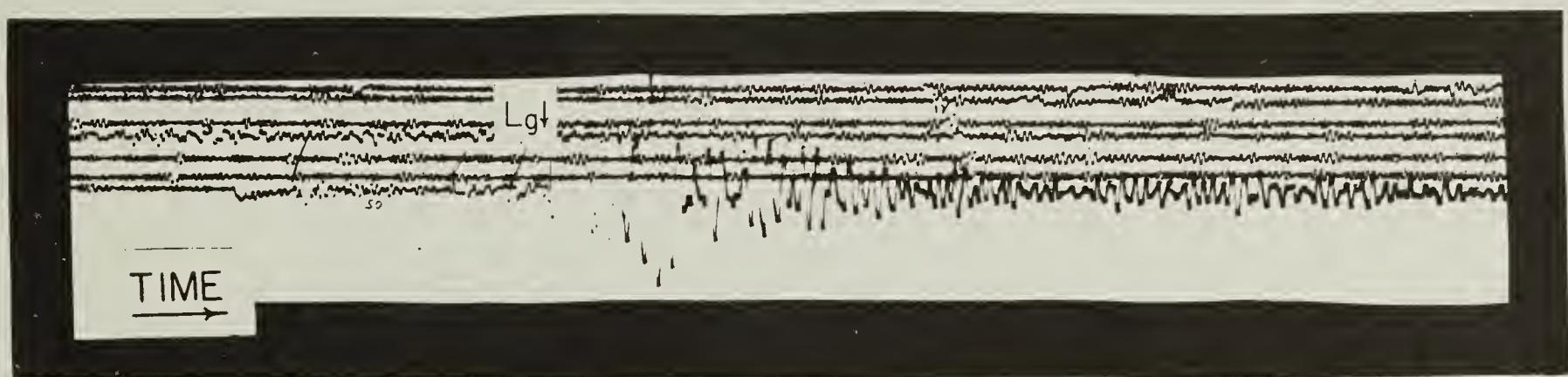


Figure 2. Seismograms:  $L_g$ , Aleutians to College, Alaska

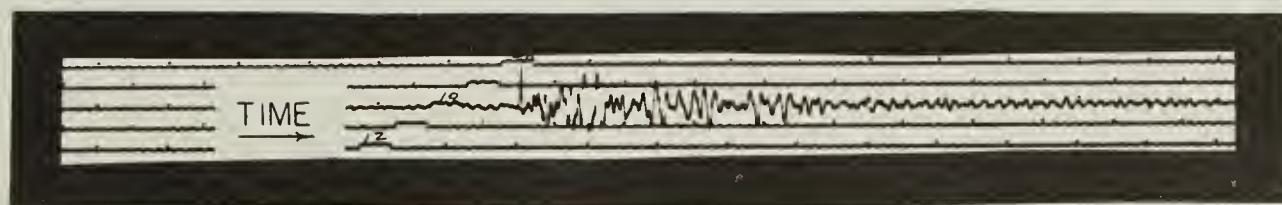


Figure 3. Seismograms:  $L_g$ , S. W. Montana to College, Alaska

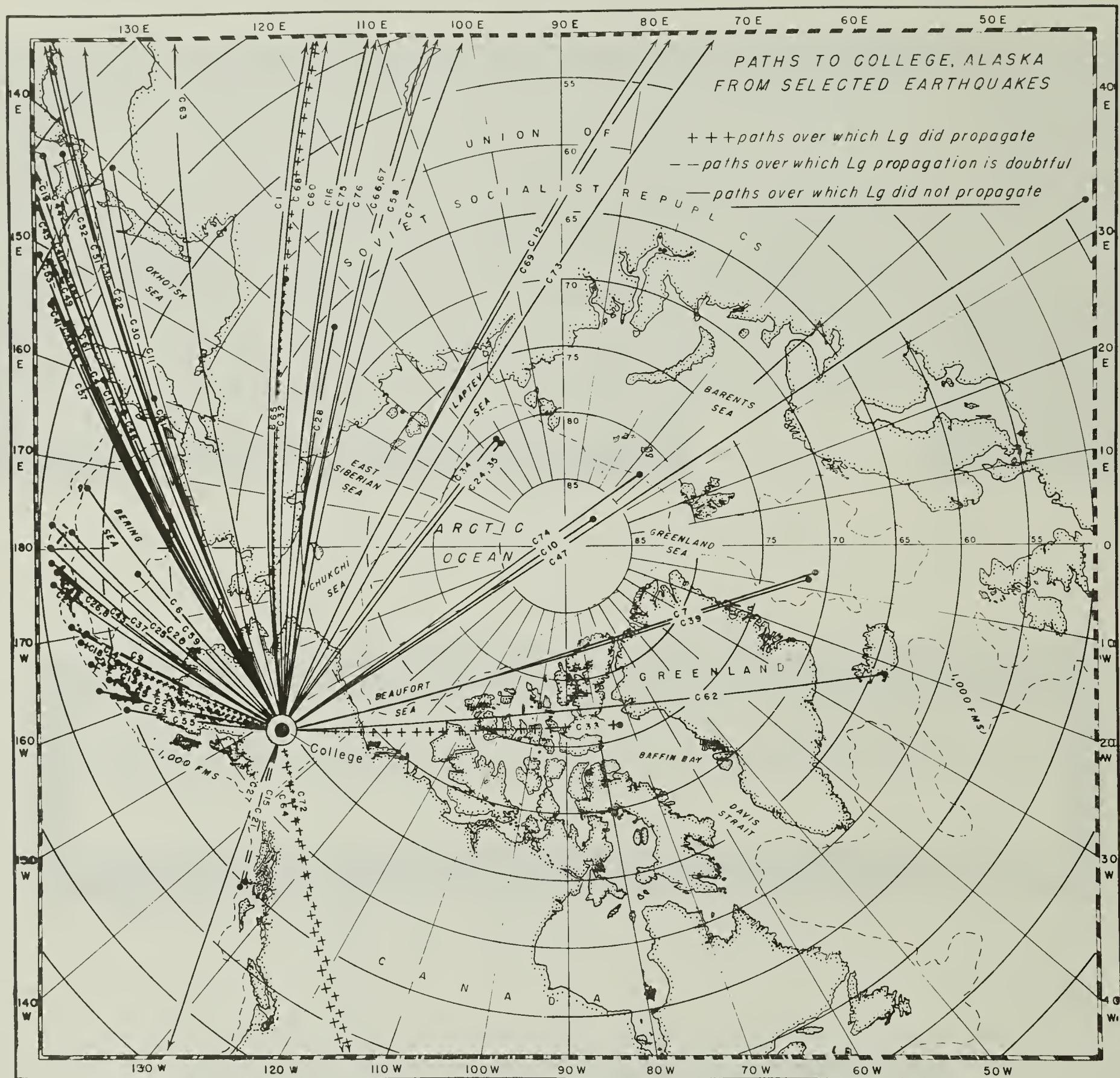


Figure 4. Paths across Arctic regions to College, Alaska

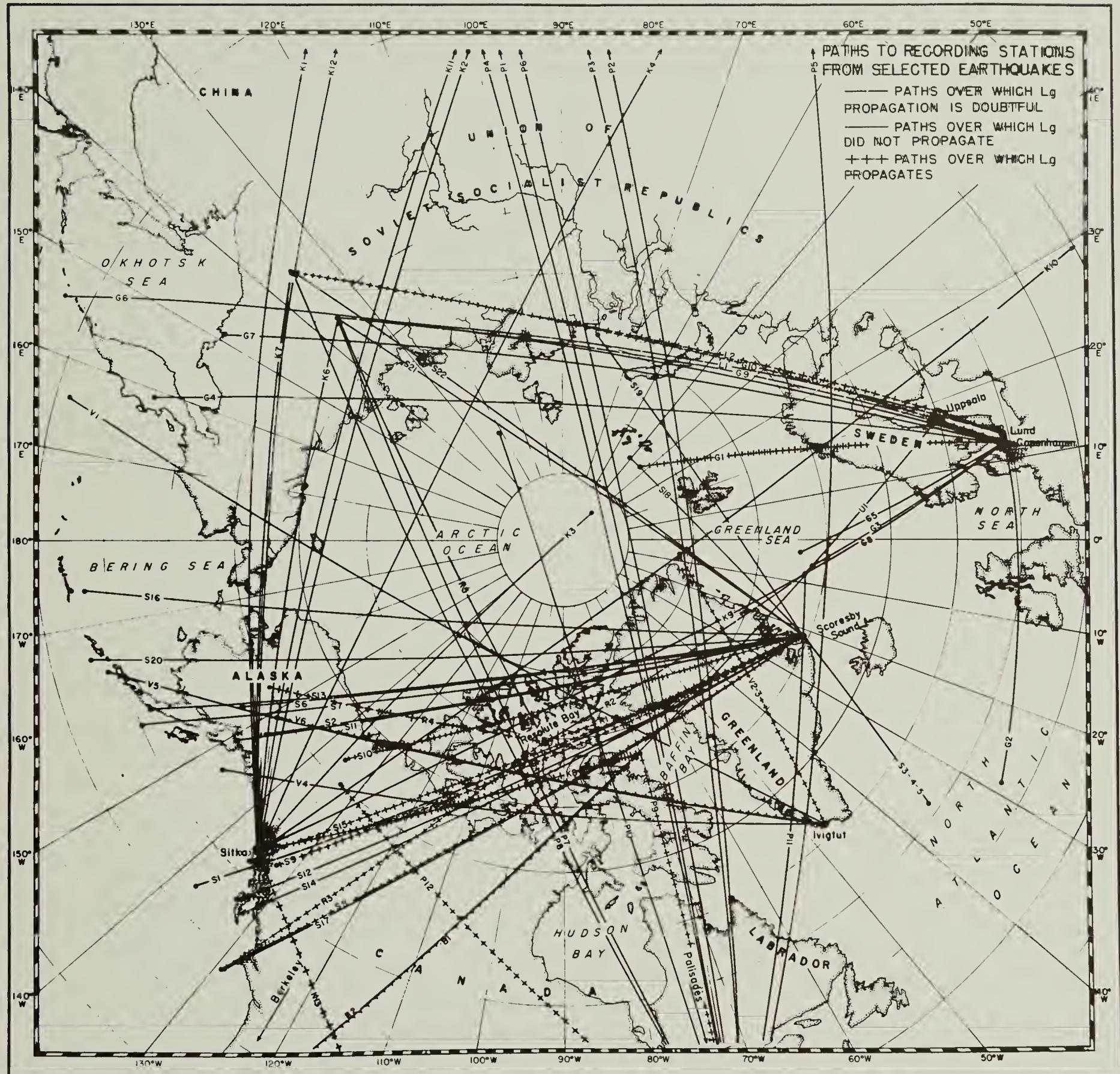


Figure 5. Paths across Arctic regions to all stations except College

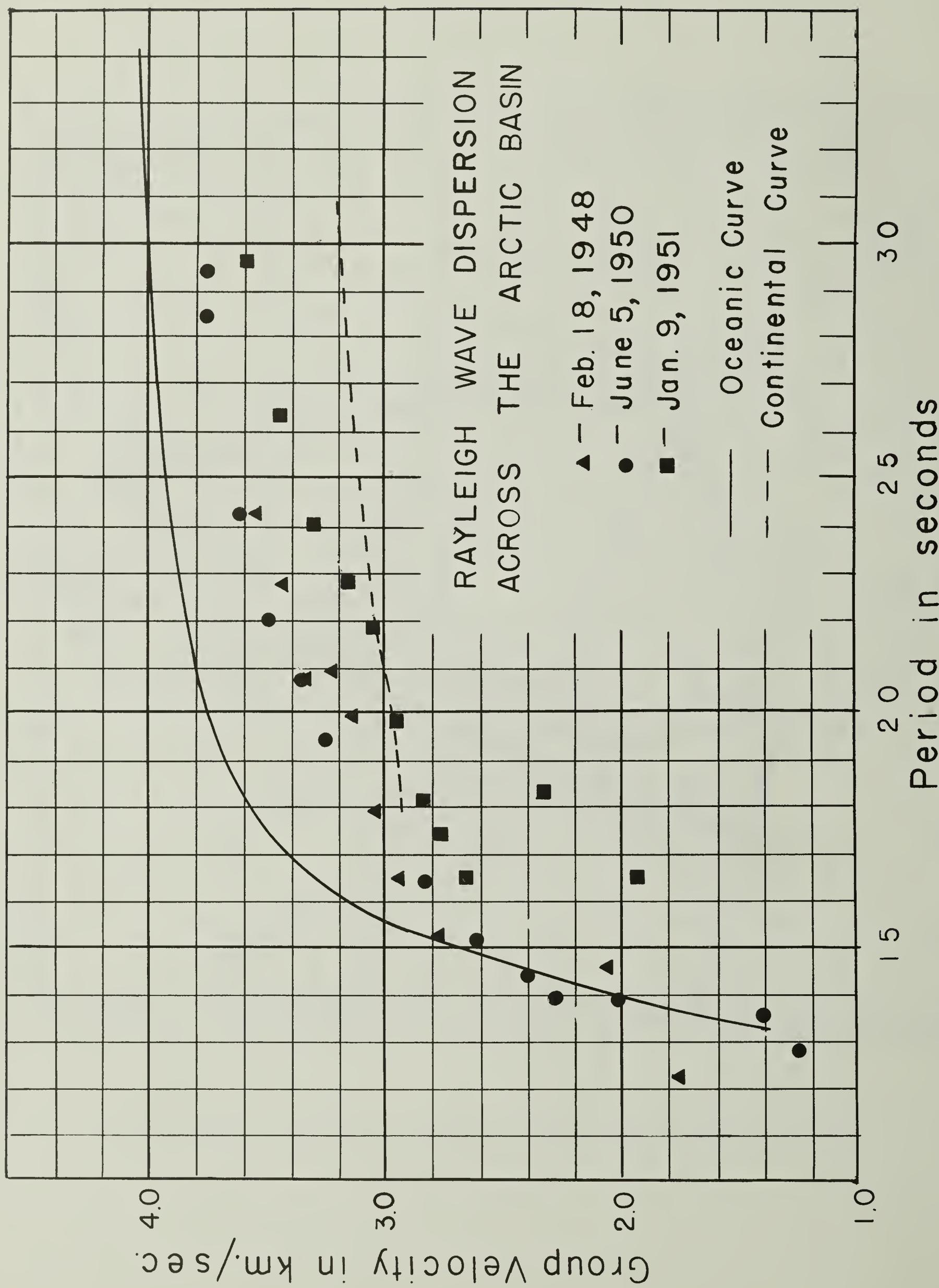


Figure 6. Rayleigh wave dispersion across the Arctic Ocean Basin

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