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Physical and chemical properties of clastic sediments from two caves in the northern karst region of Puerto Rico

Autum Rae Downey

Thesis submitted to the Eberly College of Arts and Sciences at West Virginia University in partial fulfillment of the requirements for the degree of Master of Science in Geology

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> > Morgantown, West Virginia 2020

Keywords: cave sediment, Puerto Rico, karst, geochemistry, X-ray fluorescence, fluorescence spectroscopy, PARAFAC, EEM, organic matter, trace elements Copyright 2020 Autum Downey

ABSTRACT

Physical and chemical properties of clastic sediments from two caves in the northern karst region of Puerto Rico

Autum Rae Downey

Geochemical properties and organic characteristics are essential to understanding the general depositional dynamics of a cave. Most work on clastic cave sediments focuses on physical properties of sediments. Little work has been done on bulk organic matter and geochemical properties.

This project reports physical and geochemical characterization of cave sediments from two caves in the northern karst region of Puerto Rico: El Tallonal Cave and Cueva Clara. Core and grab samples were collected and analyzed for particle size, total organic carbon, total nitrogen, organic matter characteristics using fluorescence spectroscopy, and major and minor elements using X-ray fluorescence (XRF). The sediments were poorly to extremely poorly sorted despite being seemingly homogeneous. Sand- and silt-sized grains composed most of the sediments from both locations. Within El Tallonal Cave, differences in organic chemical properties between saturated and unsaturated sediments were observed. The maximum value of total organic carbon, total nitrogen, and organic C/N ratios were higher, and the range was greater in saturated sediments. A three-component parallel factor analysis (PARAFAC) model was created using collected excitation-emission matrices on organic matter extracted from the sediments. All three modeled components were humic acid-like. Proportions of the modeled components differ between saturated and unsaturated sediments. Elemental concentrations collected using XRF on El Tallonal Cave sediments revealed that Al and Si are negatively correlated, suggesting that clay minerals and quartz occur together resulting in Si dominating in those sediments. Trace elements (Cr, Mn, Co, Ni, As, Rb, and Nb) correlated positively with Ti, implying these elements are sourced from the same parent material. Correlations between Fe, Al, and Mn with P, Co, and As suggest that sorption or co-precipitation with Fe, Al, and Mn oxides is a possible retention mechanism within El Tallonal Cave sediments. No relationship between grain size and geochemical properties was observed. Preliminary comparison with soil and bedrock from the Luqillo Mountains located within the central highlands of Puerto Ricop show similar concentrations of Ti and Zr within El Tallonal Cave sediments suggesting these elements are conservative during transport.

Results of this thesis show that sediment deposition and chemistry in individual caves are complex natural systems. Heterogeneity among cores within single caves was observed. Future studies are needed to better understand the differences observed in saturated and unsaturated sediments, as well as to further explore the mechanisms controlling the geochemical trends observed in cave sediments.

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List of Abbreviations and Acronyms

BIX – biological index

CAM – Cueva Clara (part of the Camuy Cave system)

- cps counts per second
- DI-deionized water
- DOC dissolved organic matter
- EEM excitation-emission matrix
- FI fluorescent index
- FS fluorescence spectroscopy
- ICP OES inductively coupled plasma optical emission spectroscopy
- PARAFAC parallel factor analysis
- PCA principle component analysis
- PCR polymerase chain reaction
- POC particulate organic matter
- SI_C saturation index with respect to calcite
- TAL El Tallonal Cave
- TC-total carbon
- TDS total dissolved solids
- TIC total inorganic carbon
- TN-total nitrogen
- TOC total organic carbon
- XRF X-ray fluorescence

1.0 Introduction

This study provides initial sedimentary descriptions and geochemical and organic properties of cave sediments from the northern karst region of Puerto Rico. Geochemical properties and organic characteristics are essential to understanding the general depositional dynamics of a cave. The range and variability of these parameters in clastic cave deposits are a first step in our ability to use these deposits for paleoclimate reconstruction, specifically meteoric precipitation events such as hurricanes. This study illustrates the variability in clastic sediments (laterally and with depth) from two cave systems. The data presented here demonstrates the complexity in these clastic deposits.

Clastic cave sediments are a promising means to evaluate hydrological models, examine anthropogenic changes, and to provide insight into regional climate change (van Hengstum et al., 2010). Cave sedimentation is a function of local climate variability, and thus many more data are needed to make broader regional - or even global - paleoclimate reconstructions (White, 2007). A detailed examination of the physical and chemical components of clastic cave deposits in individual caves affected recently by a hurricane is an important first step in understanding how storms affect cave sedimentation.

Assessing the physical and chemical properties of clastic cave sediments from the northern karst region of Puerto Rico was done through four major objectives: (1) Explore physical characteristics of clastic cave sediments both laterally through a cave and with depth; (2) Assess the range and variability of total organic carbon, organic C/N ratios, total nitrogen of clastic cave sediment; (3) Evaluate the fluorescence properties of organic matter extracted from clastic cave sediments; (4) Quantify bulk elemental makeup of clastic cave sediments. These four objectives will allow data from the two caves to be compared to each other, as well as to results from caves elsewhere to better document the range of physical and sedimentary processes on caves

2.0 Background

2.1 Cave sediment

Caves contain both chemical and clastic sediments. Chemical sediments defined as materials precipitating in-situ within a cave encompass a broad range of minerals due to the environmental conditions present in a cave. Chemical sediments occur dominantly through supersaturation with respect to a certain mineral. Chemical precipitation of minerals in a cave results in crystals typically composed of carbonates, but may also include sulfates, chlorides, and silica. Clastic sediments include all materials not precipitated in-situ and fall into two broad categories: allogenic and autogenic. Physical and chemical weathering within the cave system releases clasts of the host rock, including abundant carbonate clasts, as well as quartz, chert, clay minerals, and organics and are referred to as autogenic (White, 1988). Sediments sourced from the surface, referred to as allogenic sediments, encompass a variety of material (Figure 1A). Allogenic sediments enter a conduit via two pathways: (1) vertical infiltration via fracture – and – smaller sized pathways; or (2) lateral transport from a direct surface connection (e.g. sinkholes, swallets, vertical shafts) (Figure 1B). The latter dominates volumetrically (Bosch and White, 2007).

Surface derived transport is dependent on the efficiency of surface – to – ground water transfer. For example, surface connections via sinking streams flowing directly into the conduit ceiling will transport a higher sediment load than a surface connection via diffuse infiltration. Once surface water enters the groundwater system, transport of the sediment load must follow groundwater flow. The ability for sediment to remain suspended through the groundwater flow path will control the clastic material entering the conduit. Cave systems commonly contain unknown passages or undiscovered portions that create complex and convoluted groundwater flow. For these reasons, cave sediments are quite complex. Interior cave deposits are subject to remobilization and reworking. The extent to which sediments are remobilized and deposited during high flow events is variable and location-specific. Tracing sediment transport through cave systems based solely on physical sediment characteristics is problematic. Within conduits, drastic variations occur in sedimentary sequences that are only a few meters apart. Attempts to stratigraphically correlate sedimentary sequences laterally through a conduit system has been unsuccessful.



Figure 1: (A) Conceptual illustration of allogenic and autogenic clastic sediment as categorized relative to a conduit. (B) Possible sources of clastic sediment input into a conduit. Sediment sources may contribute both allogenic and autogenic sediment, as indicated by the gradient filled boxes. The conduit system is commonly dominated by allogenic sediment input, indicated by the dominantly allogenic color gradient fill (modified from Bosch and White, 2007).

Cave sediments have been studied and classified based on grain size, shape, and composition to understand the hydrological conditions under which they were deposited. Five facies were first described by Gillieson (1986) and modified by Bosch and White (2007) (Table 1). This classification scheme was described based on Appalachian cave systems. It is important to note that the hydrologic conditions, structural controls, and sediment source areas of Appalachian karst and Puerto Rico vary. Puerto Rico's karst is eogenetic with higher porosity and permeability, lower flow variability, longer lag time in response to an input signal, and lower overall response in discharge than telogenetic karst (Florea and Vacher, 2006). Fornos et al. (2009) introduced a sediment classification based on an eogenetic system (Table 1). These descriptions provide a more accurate reference for describing cave sedimentation in this study.

Cave formation, or speleogenesis, initiates as carbonate rock containing a series of joints, fractures, and bedding planes provide a less resistant pathway for infiltrating groundwater. As groundwater flows through these pathways, a feedback mechanism of dissolution and pathway widening initiates (White, 1988). Widening facilitates the onset of turbulent flow allowing a conduit to carry suspended load. Cave sedimentation and speleogenesis are intimately linked. Sedimentation in a conduit begins as soon as speleogenesis widens a conduit to allow turbulent flow (Farrant and Smart, 2011). This connection is complex and codependent as conduit development is non-linear. A cave may undergo periods of vertical abandonment, clogging, or collapse. In addition to this, a cave system will undergo multiple cycles of sediment accumulation and erosion in response to external forcing due to climatic events or, on a larger scale, glacial and inter-glacial cycles.

2.2 Storm events

The onset of sediment mobility within a conduit is marked by a threshold and strongly non-linear (Herman et al., 2008). This non-linear nature makes deposition an episodic phenomenon. Flooding events drive huge pulses of surface-derived sediment and associated organic carbon into the system. This has recently been illustrated by the effect of Category 5 Hurricane Maria on the northern karst region of Puerto Rico (Figure 2). The extent of surfacederived input is dependent on the intensity of the storm and the flashiness of the system (Herman et al., 2006).

Lithofacies	Sediment characteristics	Hydrologic condition
name Channel faciesª	Well-sorted within a given bed	Deposited in active cave streams; mainly bedload
Thalweg faciesª	Well sorted, gravel, cobble, and boulder size	Cave stream cut through previously deposited channel facies; all finer grained material has been winnowed out
Slackwater faciesª	Well sorted fine-grained clay and silts	Deposited after a flooding event; initially suspended load
Diamicton faciesª	Unsorted, unbedded sediment deposits consisting of all particle sizes	Catastrophic mass wasting event; high relief drainage basin required; not deposited in modern
Waterlain clastic fine- grained ^b	Well-moderately sorted fine-grained deposition; can be finely laminated or unstratified; mostly quartz	Deposited after high flow events
Carbonate ^b	Fine grained carbonate grains to large angular cave breakdown deposits	Breakdown triggered by a permanent change in the water table influencing the stability of the cave system
Mixed ^b	Includes both siliciclastic and carbonate material resulting in a variety of characteristics	No hydrologic restriction
Relict ^b	Older deposits of fine grained red- black silts, mm scale laminations, more clay present than waterlain clastic fine-grained facies	Deposited in ponded areas of the cave system that do not fill with storm water on a regular basis
^a facies describ	bed by Bosch and White (2007)	
~ racies describ	eu by Fornos et al. (2009)	

Table 1: Cave sediment facies and their descriptions.

Doehring and Vierbuchen (1971) described sediment deposition in Cave Springs Cave, Virginia as a result of Hurricane Camille in 1969, reporting a newly deposited terrace up to 1.5 m tall which caused permanent changes in the cave's stream flow and sediment load. Trends between modern cave sediment deposition in response to rainfall events have been found using core scanning μ -XRF on sediment cores from a saturated cave in the Yucatan Peninsula (McNeill-Jewer et al., 2019). Conversely, another saturated cave system in the Yucatan Peninsula did not experience increased sediment deposition in response to Category 1 Hurricane Ingrid, despite a 0.7 m rise in the water table. The difference in sedimentation in response to a major storm demonstrates the extreme variability and heterogeneity observed in karst aquifers.

Attempts to quantify sedimentation rates or sediment budgets with a cave system, although informative, are indirect (Collins et al., 2015b; Hart and Schurger, 2005). During low-flow conditions, cave streams commonly lack suspended load and appear completely clear. It is at this condition when a cave is studied by observers. Due to the inability to directly study cave systems during high flow, the exact mechanism, rate, and sediment distribution during high flow events within a cave system remains unknown (White, 1988).

2.3 Organic carbon

Organic carbon in sediment includes living biomass as well as plant and animal tissue at all stages of decomposition (Figure 3). Decomposition of a sediment's organic matter results from decay of large macromolecules to small monomers and the ultimate transformation back to CO₂. This process takes time. Some molecules decompose more efficiently than others. This leads to recalcitrant pools of organic carbon. Recalcitrant organic matter has essentially shed all liable portions of the material resulting in a stable molecule resistant to further weathering through a process referred to as humification.

Natural organic matter is incredibly complexity. There is currently no way to directly measure the structure and composition of these molecules (Insam, 1996). This makes understanding organic matter cycling complicated. The idea that recalcitrance, stabilization, and carbon age are correlated to one another is debatable (Kleber et al., 2011). The exact structure of humic substances are not known and is thought to vary drastically depending on the source. Due to this complexity, general bulk characterization of organic carbon is used in place of detailed molecular analysis.



Figure 2: Case study example of allogenic input into a conduit system after an extreme storm. Tres Pueblos Sinkhole (A) one week after Hurricane Maria, (B) one month after Hurricane Maria; and (C) one year after Hurricane Maria.



Figure 3: The various components of sediment organic matter. Recalcitrant organic matter are materials that are persistent in nature and resist further decomposition. Liable organic matter will undergo many stages of decomposition before entering the recalcitrant organic matter pool. Time is represented on the left-hand side.

2.3.1 Characterization

Fluorescence spectroscopy is a technique used to characterize DOC. The wavelength at which a molecule fluoresces depends on its structure and size. Molecules capable of fluorescing when excited by lower wavelengths (ultraviolet through visible) are referred to as fluorophores. There are two major categories of fluorophores in natural systems: humic-like (terrestrial derived) and protein-like (microbial derived) (Gabor et al., 2014).

An emission-excitation matrix (EEM) is an assembly of fluorescence data with excitation, emission, and intensity on the X, Y, and Z axis, respectively. EEM's are characterized using several parameters (Table 2). Peaks designated A, C, and M refer to humic-like material while peaks B and T describe tyrosine-like and tryptophan-like material (protein-like), respectively (Figure 4). There are five general regions in an EEM space (Chen et al., 2003). Regions one and two represent aromatic protein, region three is fulvic acid-like, region four is microbial by-product-like, and region five is humic acid-like (Figure 4). Fluorescent indices are also used to characterize an EEM. Humification index quantifies the degree of humification. High values indicate lower H/C ratios and thus a greater degree of humification. Fluorescence index (FI) is used to differentiate between terrestrial-and microbial-derived organic carbon. Fluorescence index values of 1.4 or less indicate DOM of terrestrial origin and values of 1.9 or higher correspond to microbially-derived material. An inverse relationship exists between FI and DOM aromaticity (Gabor et al., 2014). Biological index, like FI, is a measure of terrestrial vs. microbial components (Coble, 1996). Values of BIX between 0.8 and 1.0 correspond to freshly produced DOM of biological or microbial origin, whereas values below 0.6 are considered to contain little autochthonous organic matter (Birdwell and Engel, 2010).

While the peaks and indices described above were developed to quantify the fluorescent characteristics of DOC, they only utilize the intensity of a few discrete points within a data matrix, a matrix containing potentially thousands of data points. Moreover, sediment organic matter exhibits high fluorescence variability that may not be present within the range of these predetermined fluorescence peaks (Shirshova et al., 2009). To get a more complete understanding of an EEM, chemometrics are often used. Parallel factor analysis is a statistical method, like principle component analysis but applied to three dimensions (sample \times emission \times excitation). This method mathematically separates spectra of overlapping fluorescent components. This type of exploratory analysis takes large, complex EEM files and

reduces them into individual components (Bro, 1997). Parallel factor analysis is particularly valuable due to its ability to make direct chemical interpretations from the model.

Parallel factor analysis of a 3-way dataset decomposes the data signal into a set of trilinear terms and a residual array (eq. 1).

$$x_{ijk} = \sum_{f=1}^{F} a_{if} b_{jf} c_{kf} + e_{ijk}$$
 Eq. 1

 x_{ijk} is the intensity of the *i*th sample at the *j*th emission value and at the *k*th excitation value (Figure 5). Each *f* corresponds to a PARAFAC component. Each component has *I a*-values (scores), one for each sample. Each component also has *J b*-values, one for each emission wavelength as well as *K c*-values, one for each excitation wavelength. Parameter a_{if} is directly-proportional to the concentration of the *f*th analyte of sample *I*. b_{jf} is a scaled estimate of the emission spectrum of the *f*th analyte and c_{kf} is linearly proportional to the specific absorption coefficient of the *f*th analyte. Finally, e_{ijk} is the residual representing the variability not accounted for by the model (Murphy et al., 2014).

The PARAFAC model has three assumptions: (1) individual chemical components have distinctly different excitation and emission spectra; (2) fluorescence intensity relates linearly with concentration and spectra are consistent across all wavelengths; and (3) fluorescence intensity is an amalgamation of a fixed number of components (Bro, 1997).

2.3.2 Organic carbon in cave systems

Input of allogenic dissolved organic carbon (DOC) and particulate organic carbon (POC) may arrive via two different pathways: infiltration through the epikarst zone or rapid injection through a direct surface connection. In most karst systems DOC represents a larger flux than POC, and as a result, is studied at greater lengths (Gillieson, 1986). It is thought that negligible organic carbon is generated within the cave itself due to the darkness limiting primary productivity (van Hengstum et al., 2010). While the lack of light obviously eliminates phototropic microbial activity, it is possible that chemoautotrophic metabolism can flourish. The thought that organic matter generation is negligible could be limiting our understanding of organic carbon cycling in these systems.

Measurement	Calculation	Interpretation
BIX (biological index)	Emission intensity at 380 divided by emission intensity at 430 nm at excitation 310 nm	>1 recently produced OM ^a via autotrophic productivity
FI (fluorescence index)	Emission intensity at 470 divided by emission intensity at 520 nm at excitation 370 nm	<1.4 OM ^a of terrestrial origin >1.9 OM ^a of microbial origin
HI (humification index)	Area under the emission spectra from 435–480 divided by the area under the emission spectra from 300–345nm + 435–480 nm, at excitation 254	Higher values indicate increasing degree of humification; low = <5 and increases with decomposition; high values indicate lower H/C ratios
Peak A	Maximum intensity at 260 nm excitation and 380-460 nm emission	Humic acid-like
Peak B	Maximum intensity at 275 nm excitation and 310 nm emission	Tyrosine-like, protein-like
Peak C	Maximum intensity at 350 nm excitation and 420-480 nm emission	Humic acid-like
Peak M	Maximum intensity at 312 nm excitation and 380-420 nm emission	Marine/non-marine humic acid-like
Peak T	Maximum intensity at 275 nm excitation and 340 nm emission	Tryptophan-like, protein-like
Region 1	Excitation <250 nm <335 nm emission	Aromatic protein
Region 2	Excitation <250 nm and emission between 380 and 335 nm	Aromatic protein
Region 3	Excitation <250 nm and emission >350 nm	Fulvic acid-like
Region 4	Excitation between 250-280 nm emission <380 nm	microbial by-product-like
Region 5	Excitation >280 nm and emission >380 nm	Humic acid-like
^a Organic matter		

Table 2: Fluorescent peaks and indices used to characterize an EEM.

Measurement	Calculation	Interpretation
C:A	Peak C intensity: Peak A intensity ratio	High values= humic acid-like Low values = fulvic acid-like
C:T	Peak C intensity: peak T intensity ratio	High values = humic/recalcitrant OMª Low values = fresh/liable OMª
^a Organic matter		

Table 2: Fluorescent peaks and indices used to characterize an EEM (continued).



Figure 4: Peaks and regions indicated spatially within the excitation emission space. Peaks labeled as described by Coble (1996) and regions labeled as described by Chen et al. (2003).



Figure 5: Conceptual model of parallel factor analysis (PARAFAC). PARAFAC decomposes an emission excitation data set, here from i to I into distinct components illustrated on the right-hand side of the figure.

It has been observed that DOC concentrations in the water column and DOC associated with suspended sediment decreases from input to output (Blott and Pye, 2001; Husic et al., 2017; Simon et al., 2007). This suggests that DOC is retained in the system via some metabolic pathway, sorption, or physical entrapment. Concentrations of DOC are variable between input type and hydrologic condition. Sinking streams have been found to contain higher DOC concentrations but had little variation during changing flow conditions. Epikarst waters contained lower DOC concentration but exhibited variability consistent with the onset of changing flow conditions (Pipan and Culver, 2013).

To better understand how organic carbon is transported and retained within a cave, Simon et al. (2010) characterized DOC from major inputs and outputs using PARAFAC analysis. Significant variation in DOC characteristics between input and output were observed and is likely a reflection of the organic carbon source. Component one (humic-like, high molecular weight) was most-strongly represented in the cave stream, component two (humic-like, low molecular weight) was represented in the soil DOC from the karstic surface, and component three (protein-like) was represented in cave drip water. A similar study done by Pipan and Culver (2013) in a different system found that cave drip water contained a lower percent of aromatic compounds than both the sink and resurgence point. Aromatic compounds are highly stable. They are generally more resistant to decomposition and persist in natural systems. DOC with higher aromaticity is considered recalcitrant and thus older than less aromatic DOC. Less aromatic DOC in drip waters indicates the presence of freshly produced DOC, likely a result of microbial activity in the epikarst zone.

Birdwell and Engel (2010) reported fluorescence index (FI) and humification index (HIX) values for groundwater from a confined portion of the Edwards aquifer to be 1.51-2.65 and 0.17-9.03, respectively. Suggesting that there exists a significant microbial component (protein-like) and an overall lack of terrestrial input into this groundwater system. This signature is consistent during recharge periods, indicating that the influx of nutrient-rich surface material stimulated, rather than inhibited, microbial activity. The dominating microbial signature of the DOC in this study is likely a function of the confined nature of the sampled aquifer. Mudarra et al. (2011) measured the fluorescence properties of multiple spring waters from an unconfined system and found the dominant organic matter was fluvic acid-like.

Within a conduit, sediment organic carbon is retained and stored for an extended period relative to water residence time (Husic et al., 2017). Within the hyporheic zone, total biomass is closely correlated with POC concentration rather than DOC concentration. In marine sediments, molecular weight decreases with depth and age (Lukawska-Matuszewska et al., 2018). These studies support the idea that freshly produced organic matter are large molecules that undergo post depositional humification towards a recalcitrant product (Sobczak et al., 1998).

2.4 Cave sediments as paleo-proxies

Historically paleoclimate studies have focused on speleothems due to their fine resolution and cyclical precipitation banding. Only recently have cave sediments been recognized as a valuable source of information regarding past climatic conditions (White, 2007). The onset of sediment accumulation occurs simultaneously with the formation of the conduit whereas speleothem growth initiates after conduit formation. Because of this, dating clastic sediments reveal a minimum age of the conduit and is likely to be more accurate than dates revealed using speleothems.

There has been a surge of research attempting to reconstruct paleoclimate and paleostorm records through clastic cave sediments in the past two decades. Most reconstruction attempts have focused on organic carbon characterization and dating using both radiocarbon and magnetic polarity techniques (Hearty et al., 2004; Panno et al., 2004; Polk et al., 2013; Polk et al., 2007; van Hengstum et al., 2010). These studies suggest that when a single core is studied carbon dates reveal little to no age reversals and are generally stratigraphically-continuous. Organic matter has been found to decrease with increasing depth and age as a result of progressive microbial degradation under aerobic conditions creating CO₂ and a loss of easily-oxidized organic carbon. Tamalavage et al. (2018) successfully identified sediment organic carbon sources to a sinkhole through time. Periods of increased primary productivity and changes in sediment source areas were quantified using an isotope mixing model. This study most importantly demonstrates that organic carbon variation is present in the depositional record on the karstic surface.

3.0 Location

This study was conducted on two cave systems located in the northern karst region of Puerto Rico, El Tallonal Cave (TAL) and Cueva Clara (CAM). Location TAL was chosen because this cave's water quality and sediment chemistry has been monitored since August 2016.

Datalogger data collected from TAL show that a spike in turbidity in the cave stream coincides with the onset of a rain event (pers. comm. Pantoja-Agred, 2016). The spike in turbidity suggests that this cave system experiences regular allogenic input from the surface. The second location, CAM, was included in this study because sediment accumulation as a result of Hurricane Maria was directly observed (Miller, 2018). Though both locations were flooded as a result of Hurricane Maria, the "normal" conditions are very different between the two systems. El Tallonal cave is a wet cave with a continuously-flowing cave stream whereas CAM is a completely dry cave. Inclusion of TAL and CAM provide a comparison between the physical and chemical sediment characteristics present in these two different systems.

3.1 Geology and landscape

Puerto Rico is part of a volcanic island arc, consisting of mountainous central highlands with steep slopes and large incised valleys, surrounded by a flat lying coastal plain (Figure 6). It is the easternmost island of the Greater Antilles located near the North American-Caribbean plate subduction zone (Cox, 1985; Jolly et al., 2001). Igneous terrain makes up 52 percent of the island surface. These volcanics are concentrated in the central highlands and are basaltic with high compositional variation from silica rich to K-rich calc-alkaline basalts (Jolly et al., 2001). The volcanic central highlands and are surrounded by coastal sedimentary rocks from the late Jurassic to early Eocene (Murphy et al., 2012). Coastal limestone covers about one-fifth of the land area. There are two major karst regions, located on the northern and southern coast of the island (Giusti, 1978) (Figure 7). The sediment studied in this project are transported from two major sources: (1) the volcanic central highlands; and (2) the carbonates.

The Puerto Rican karst is eogenetic, undergoing little to no post-depositional deformation. Carbonates were deposited in shallow, clear water with open circulation, ranging in age from Oligocene to Miocene (Giusti, 1978).



Figure 6: Generalized cross section from the central highlands to the northern coast of Puerto Rico. Surface topography is controlled by the underlying carbonate unit (modified from Monroe 1976).



Figure 7: The island of Puerto Rico with karst regions highlighted. The two cave systems in this project and the path of Hurricane Maria (from southeast to northwest) are indicated. Karst data source: USGS. Hurricane Maria data source: NOAA.

The Puerto Rican karst is eogenetic, undergoing little to no post-depositional deformation. Carbonates were deposited in shallow, clear water with open circulation, ranging in age from Oligocene to Miocene (Giusti, 1978).

The island of PR consists of five major limestone formations (Figure 6). The Aymammon and Aguada limestones have the highest permeability (1-0.02 cm/S) and are the most productive aquifers on the island. Water budget calculations yield a discharge (Q) of 0.11 m³/s and 0.02 m³/s for the two formations, respectively (Giusti, 1978). These formations on average are 85 percent CaCO₃, 4.8 percent MgCO₃, 6 percent SiO₄, and 9 percent insoluble material (Monroe. 1976). Karstification is most intense in the Aymammon limestone and decreases seaward (Giusti, 1978). The tropical climate and eogenetic nature of the Aymammon limestone unit results in a karst topography referred to as magotes (Figure 8). These are rounded mounds with a hardened shell surrounding a softer, eroded interior. The induration and reprecipitation of the uppermost portion of the exposed carbonate produces a "clogging" of the primary porosity and provides important structural integrity (Ford and Williams, 2007). Magotes are asymmetrical, exhibiting a shallow east side in response to the easterly rain typical of the tropical high-pressure systems that pass through the island.

3.2 Climate

Along the northern coast of PR, trade winds trend from the northeast. The annual air temperature mean is 25°C with little yearly variation. Mean annual rainfall for the northern coast of the island is 1550 mm/yr (Morelock et al., 2016)

Major rainfall events are produced during the passage of a cold front or an easterly wave over the region. The frequency of these weather fronts determine whether the island experiences relatively wet or dry conditions (Gomez-Gomez et al., 2014).

3.2.1 Hurricane Maria

On September 20th, 2017 Hurricane Maria made landfall on the southeastern side of PR just under a Category 5 hurricane. The eyewall crossed the island from the southeast to northwest (Figure 7). The storm surge combined with the tides on the northern coast producing a maximum surge of 1.5 m. Rainfall on the island reached upwards of 97 cm in just over 32 hours (Pasch et al., 2017). River discharge in many places reached at or near record levels and in many cases

disabled or destroyed the water level monitoring stations. Severe flooding and mudslides affected much of the island (Pasch et al., 2017). Large volumes of newly-deposited sediment were documented in many caves throughout PR after Hurricane Maria (Miller, 2018).

3.3 The cave systems

3.3.1 El Tallonal Cave (TAL)

El Tallonal Cave is located on private land, protected and preserved by the Citizens of Karst (Cidadanos del Karso) and used for private water supply. It is dammed at the portal and has been instrumented with an Aquatroll® 600 datalogger since 2016 (pers. comm. Padilla, 2019). The cave has one main sinusoid passage with no branching side passages and no known upper or lower levels (Figure 9). Water flows upwards into the sump and discharges from the base of a magote at the cave entrance (Figure 8). The cave is approximately 60 meters long and varies from 1.5 meters to just over 6 meters wide.

Water quality at TAL has been monitored by researchers at the University of Puerto Rico, Mayaguez since August of 2016 as part of a long-term study to better characterize groundwater flow in the region. The groundwater flowpath is relatively short, entering and exiting on either side of a magote, however, dye traces have been unsuccessful (pers. comm. Padilla, 2019) Groundwater response to a rain event is variable depending on the time of year. In the warmer months of the year (August-November) a rain event is marked by an increase in groundwater temperature. In the cooler months (January-May) a rain event is marked by a decrease in groundwater temperature. A spike in turbidity in response to rain event is temporally-consistent indicating that allogenic input coincides with these rain events.

Aqueous systems experiencing the dissolution or precipitation of calcite is determined by the SI_c. An SI value above zero (indicating equilibrium) is supersaturated thus precipitating calcite. A SI_c value less than zero is undersaturated, indicating the water can dissolve calcite. The SIc values were calculated using Ca and alkalinity concentrations and pH from the data provided by the University of Puerto Rico, Mayaguez. All SI_c values from TAL are at or above 0 (maximum SI_c=0.45). Phosphate and TDS (mg/L) experience a sharp increase in October of 2017 and could be a response to the massive flooding that occurred within the system during Hurricane Maria.



Figure 8: (A) Aerial imagery and surface contours of the regional topography surrounding El Tallonal Cave (TAL); Distinct magote peaks are present within the surround area. (B) Aerial imagery and surface contours of the regional topography surrounding Cueva Clara (CAM). Magote karst is absent in this region as the Río Camuy system is contained within a different carbonate rock unit. Sinkholes in this region are karst windows to the Río Camuy River. TAL cave map modified from Lace (2006). CAM cave map modified from Acosta-Colon (2016).



Figure 9: Map of El Tallonal Cave with sampling locations categorized by sampling round. The compartmentalized cave sections are indicated by the dashed boxes. Cave map modified from Lace (2006)
The cave has three main compartments: the front, the middle, and the back (Figure 9). These divisions are based on cave morphology and major sediment bank locations.

The front

The dammed portal inhibits the amount of terrestrial input at the entrance (e.g. windblown plant material, biota, dust). The front is approximately 1.2 meters tall x 3 meters wide. An Aquatroll® 600 datalogger is located approximately two meters into the cave from the dammed entrance. Multiple blue crabs and cave scorpions are in the crevasses of the cave ceiling and clastic sediment banks in this portion of the cave.

Connecting the front and middle of the cave is a narrow passageway, approximately 1meter x 1.5 meters and contains soda straws and stalagmites.

The middle

The middle compartment is approximately seven meters wide by two meters tall and is the largest room in the cave. The sediment banks in this region of the cave are the largest and most accessible. Clastic material within the stream channel contain well-rounded, dominantly pebble and cobble sized grains suggesting that the stream has winnowed out smaller grained material over time. An erosional terrace was observed in the sediment bank on the northern side of the cave stream. This terrace marks the water level when the dam is closed and was present before Hurricane Maria and therefore is not attributed to the flooding caused by the hurricane.

Connecting the middle and back of the cave is a narrow passageway with small (approximately 5 cm long) soda straws on the cave ceiling.

The back

The back compartment of the cave is approximately 1.2 meters wide by one meter tall. It contains a clastic bank that hugs the northern side of the cave wall immediately adjacent to a pool marking the end of the accessible passage.

3.3.2 Cueva Clara (CAM)

Cueva Clara is part of the National Parks Company of the Commonwealth of Puerto Rico natural protected area and is located in the Río Camuy cave system (Figure 8). It is the main

toured portion of the Río Camuy system and has direct vehicle and walkway access to the entrence. The interior of the cave has paved walkways for tourism. The National Parks Company owns and maintains the cave, however, as of 02/16/2019 the park has remained closed to the public as a result of the damage from Hurricane Maria.

This system has been carved by the Río Camuy River, the third largest subterranean river in the world (Nieves-Rivera, 2003). The Río Camuy cave system contains multiple levels above the existing stream level as a result of the Río Camuy River diverting underground as baseflow lowered. Cave sediments from the upper most levels of the system, dated using cosmogenic isotopes (¹⁰Be and ²⁶Al), are up to 4.5 MA (Miller, 2018). Cueva Clara is a dry cave due to the vertical abandonment of the Río Camuy River. It contains two main portals, both at the base of large sinkholes (Figure 10). The Río Camuy River is at the base of the Emplane Sinkhole making the back portion of Cueva Clara a karst window overlooking the bottom level of the Río Camuy system.

The Río Camuy River level rose as result of Hurricane Maria so that it reached the entrance of Cueva Clara via the Emplane Sinkhole, completely flooding it. Once the river levels returned to baseflow after the storm, approximately 0.4 meters of sediment was deposited throughout the cave as determined by the height of sediment accumulation on the paved walkways after the flooding subsided. Cueva Clara also contained a large bat population, all of which were decimated due to the flooding from the hurricane.



Figure 10: Map of Cueva Clara with sample locations labeled. The toured cave entrance and overlook to the Emplane Sinkhole are also labeled. Cave map modified from Nieves-Rivera (2003) and Acosta-Colon (2016).

4.0 Methods

4.1 Sample collection

Three sampling rounds were conducted, and both grab samples and sediment core were collected (Table 3). Core samples were collected using 5 x 30.5 cm polyethylene core sleeves. Cores collected from TAL were sampled using hand force to slide the polyethylene core sleeve into the sediment. Core samples from CAM were collected using an AMS soil corer with a slide hammer attachment.

After collection all samples were refrigerated and kept from direct light until further analysis. A detailed list of all samples and analysis completed on each sample is found in Appendix IV.

Samples collected from the front portion of TAL have been excluded in this study. Sample locations TAL-16 and TAL-17 were collected along a diagonal transect through the sediment bank on either side of the stream in the middle portion of the cave (Figure 11). Sediment collected from the middle portion of TAL are the focus of this study (Figure 12).

Core samples from Cueva Clara were collected from various locations throughout the cave (Figure 10). This sediment, deposited by Hurricane Maria, had since dried and hardened. Desiccation cracks were observed on the sediment surface.

4.1.1 Splitting and describing sediment core

The cores were split in half lengthwise in order to subsample for specific analysis (Table 4). The core sleeve was split using an AMS safety liner splitter and straight edge blade. Once the core was split, both halves photographed.

Once described, one half of the core was subsampled for analysis (Figure 13). Subsamples were collected from depths in which visible change had occurred (e.g. color change, grain size change, and textural change). Each subsample had a thickness of 1 ± 0.5 cm. If a given subsampled section was thinner than one cm the subsample was collected from only the desired section. The depth relative to the top of the core was incorporated into the sample tracking ID (Appendix IV). Samples were later converted to a depth relative to a datum within the cave. The depth relative to the datum of each subsample is found in Appendix IV.

The unsampled half was stored and later analyzed using XRF. After splitting and subsampling the core, it was wrapped in aluminum foil, labeled, and refrigerated.

Table 3: Samples collected organized by sampling round and type of samples collected.

Sampling round	Sampling locations	Method	Number of samples
(1) Feb 2017	TAL ^a	Grab samples	4 Grab samples
(2) July 2018	TAL ^a	Grab samples	11 Grab samples
(3) Feb 2019	TAL ^a	Grab and core samples	20 Grab samples
			18 Core samples
(3) Feb 2019	CAM ^b	Core samples	12 Core samples
^a El Tallonal Cave ^b Cueva Clara			



Figure 11: Generalized cross section of the middle of El Tallonal Cave with sample locations included. Point locations represent grab samples. Sample locations with a length represent sediment cores. Observed water level (dark blue) and dammed water level (light blue) are indicated. Samples were given a depth relative to the datum illustrated in this figure.



Figure 12: The spatial distribution of sample locations TAL-16 and TAL-17. Sample locations TAL-16 and TAL-17 were collected in a diagonal transect along a sediment bank in the middle of the cave. Dammed water level is depicted. A) sample transect location in the middle of the cave. B) location of collected core samples along the transect.

Analysis	Mass needed (g)	Storage container	Sample preparation	
TOC ^ª and TN [♭]	1-2	Plastic petri dish	Air-dried, oven-dried at 60°C and lightly ground	
Particle size	2-3	50 mL centrifuge tube	Sediment, Calgon powder, and DI ^c slurry	
FS ^e	3	Aluminum foil	Alkaline OM ^d extraction	
XRF ^f	N/A	Split plastic core sleeve wrapped in aluminum foil and plastic wrap	Air-dried	
^a Total organic carb	on			
^b Total nitrogen				
^c Delonized water				
° Organic maller ¢ Elucrescence spectroscony				
^f X-ray fluorescence				
 ^e Fluorescence spectroscopy ^f X-ray fluorescence 				

Table 4: Analysis completed on collected subsamples.



Figure 13: Flow chart illustrating subsampling and laboratory analysis.

4.2 Laboratory analysis

4.2.1 Particle size analysis

Subsamples for particle size analysis were prepped using a 1:1.5 (v/v) sediment to five percent Calgon solution ratio. The sediment slurry was placed on a rotary shaker at 70 rpm overnight. Sediment slurries were analyzed using a Beckman Coulter single wavelength LS13-320 particle size analyzer that measures sizes between 0.4 μ m and 2,000 μ m based on Fraunhofer diffraction theory. Particle size is reported as a function of volume percent.

Raw data files from the instrument were organized in R and processed through the GRADISTAT program (Blott and Pye, 2001). This program computes grain size, sorting and skewness, as modified from Folk and Ward (1957) (Figure 14).

4.2.3 Elemental analysis using ICP-OES

Inductively coupled plasma - optical emission spectroscopy (ICP-OES) was used to measure major and trace elemental concentrations of selected reference samples (16 in total) following a lithium borate fusion method. This fusion method involves melting the sample matrix with lithium tetraborate and subsequently dissolving the mixture in nitric acid. Lithium borate fusion facilitates a complete dissolution of the resulting mixture. The fusion and ICP-OES analyses were completed by Activation Laboratories, Inc. of Anacaster Ontario following their lithogeochemsitry method – code 4Litho.

4.2.4 Elemental analysis using XRF

Sediment samples were analyzed for elemental concentrations using XRF. Both the loose sediments and the core samples were measured using a low energy scan (for major elements, plus Ba, V and Cr) and high energy (for trace elements) conditions using a Bruker AXS TRACER III-V ED-XRF equipped with a Rh X-ray tube. Low-energy analysis was completed at 15 kV and 34.4 μ A, with helium, and a count time of 60 s. High-energy analysis was completed at 40 kV and 25 μ A, with an Al–Ti–Cu filter, and a count time of 60 s (Rowe et al., 2012). Core scans were completed in 2.0 mm intervals using an automated core scanner apparatus that houses the handheld XRF described above (Figure 15). All XRF scans were completed at the University of Utah's Records of Environment and Disturbance (RED) lab.



Figure 14: A) Particle size distribution of a sample collected from El Tallonal Cave as a function of volume percent with geometric mean, sorting, and skewness as reported in the GRADISTAT program. B) Grain size, sorting, and skewness divisions in the GRADISTAT program. Modified from Blott and Pye (2001).



Figure 15: A) Photo showing the XRF core scanner located at the University of Utah's RED lab. B) The XRF core scanner with a sediment core loaded and ready for scanning.

Collected XRF data were quantified using a modified version of the mudrock calibration created by Rowe et al. (2012). Selected reference samples were air dried, pulverized to $<74 \mu m$, placed in an X-ray translucent plastic container in order to protect the instrument sensor, and subsequently analyzed using the same X-ray conditions as previously described. These reference spectra were loaded into proprietary Bruker AXS calibration software (*S1CalProcess*). Reference concentrations of each element collected from ICP-OES as described in section 4.2.1 were evaluated against the raw photon counts from each element calculated by inter-element slope and baseline corrected peak heights. Because the software used to create this calibration is proprietary, a full disclosure and discussion of the calibration routines is currently not available. Calibration was created by collaborators from the University of Utah following methods outlined in Kipnis et al. (2019)

Core sample preparation involved storage just above freezing (~0.5 °C) to facilitate drying. For analysis, cores must have water content <40 percent (MacLachlan et al., 2015). Once dried, the core was placed on the automated core scanner and covered in an X-ray translucent plastic wrap in order to protect the instrument sensor and analyzed using the same X-ray conditions as previously described.

After data collection, significant post-processing and data manipulation was performed. Spectra collected from the core were filtered through to pick out any compromised scans. This included scans in which the total photon counts were <3000 cps or where argon photon counts were >1000 cps. Compromised scans occured due to three major reasons; (1) there existed too large of a gap between the instrument sensor and the core; (2) a gap in the core itself; or (3) a gap between the end of one core and the beginning of another. Given that the scan interval was 2.0 mm, excluding compromised scans did not necessarily affect the data trends through the core.

The custom calibration was applied to the filtered raw spectra using proprietary software from Bruker. Once applied, raw photon counts were converted to concentrations and the fully processed data was exported into excel.

4.2.5 Total carbon, total organic carbon, total inorganic carbon, and total nitrogen

Approximately three grams of sediment were placed into a plastic petri dish and left to air dry for two days. Once the sample was air dried it was homogenized and oven dried at 60°C for 24 hours. After oven drying, the sample was homogenized a second time. After preparation, the

samples were sent to the University of Florida's Stable Isotope Mass Spectrometry Lab for analysis.

4.2.5.1 TC and TN

Samples were analyzed using Carlo Erba NA1500 CNHS elemental analyzer. After flash combustion the sample gas was passed through a hot reduction column (650°C) consisting of reduced elemental copper to remove oxygen and a chemical (magnesium perchlorate) trap to remove water, finally reaching a gas chromatography column at 125°C that separates the N₂ and CO₂ gases.

4.2.5.2 TIC and TOC

Total inorganic carbon (TIC) was measured coulometrically using a UIC (Coulometrics) 5017 CO_2 coulometer. The sediment samples were placed into vials topped with septa. A double needle assembly was used to purge the sample vial of atmospheric gas using CO₂-free nitrogen carrier gas. Acid was then injected into the sample vial and evolved CO₂ was carried to the coulometer where total carbon as CO₂ was measured.

Total organic carbon (TOC) was obtained by taking the difference between total carbon (TC) and TIC.

4.2.6 Fluorescence Spectroscopy (FS)

Characterization of sediment organic carbon was performed using a Horiba Scientific FluoroLog-3 spectrofluorometer system. This system is configured with 450-Watt light source, single grated emission monochromator, and a single grated excitation monochromator. Excitation-emission matrices (EEM) were collected from λ ex: 240-400 nm and λ em: 280-540 nm and a slit width of one nm for both the excitation source and emission detector. 1st and 2nd order Rayleigh scattering were removed before and after each scan.

Aqueous tryptophan, tyrosine, and humic acid were used as a standard reference for the EEM scans (Figure 16). Tryptophan and tyrosine are both amino acids and are used as a reference for the characteristic peak T (λ ex: 275 nm; λ em: 340 nm) and peak B (λ ex: 275 nm; λ em: 310 nm) in Coble (1996), respectively. Humic acid is a more complicated material and therefore does not have a distinct fluorescence peak but rather a region (λ ex: 250-400 nm; λ em:

375-550). Specifications of the fluorometer used in this study restricted the possibility of measuring fluorescence response of raw sediments (Appendix III)

Organic matter was extracted into an aqueous form. There many ways to extract organic matter from a solid medium with results dependent on the method. Comparisons between two extraction methods were performed: (1) 0.1 M sodium hydroxide and; (2) 0.1 M sodium pyrophosphate (Na₄P₂O₇) (Chen and Hur, 2015; Greenfield et al., 2018) (Appendix II). Results from this comparison concluded that using 0.1 NaOH as the extraction solution yielded a higher florescence response and thus all extractions were performed using NaOH as the extraction solution. All extractions were completed by following a modified version of the method outlined in Hur et al (2014).

Air-dried sediment measured between three and four grams was mixed with 0.1 M NaOH solution in a 1:10 sediment:solution (m/m) ratio and placed overnight on a reciprocating shaker. The sediment-alkaline slurry was then centrifuged at 7500 rpm for 40 minutes. The supernatant liquid was then filtered through a Fisher brand P4 cellulose filter and collected for analysis on the fluorometer following the parameters previously outlined. All organic extractions were diluted by 50 percent when ran for analysis on the fluorometer.

Once an EEM was collected, it was exported to a CSV file. The data were then manipulated in R (Appendix XIII). Raw data collected from the fluorometer was manipulated in R using the "*eemR*" package (Massicotte, 2019). This package removed first and second order Rayleigh scattering from a 15 nm range and interpolated intensity over the Rayleigh scattering range (Figure 29). All EEMs were normalized to a percent of the maximum intensity for each scan. This was done to normalize the data across all samples for PARAFAC modeling Indices BIX and FI were calculated using formulas contained in the "*eemR*" package. Parallel factor analysis was performed using the "*staRdom*" R package (Pucher et al., 2020). Model results from the PARAFAC analysis are reported in section 5.2. Details on the PARAFAC modeling code are found in Appendix XIII.

4.3 Data analysis

Stratigraphic columns and core photographs were created and modified in Adobe Photoshop and Illustrator. Figures were created using a combination of Grapher by Golden Software, R, and Microsoft Excel and edited using Adobe Illustrator and Microsoft Excel. Post processing was completed in R. All R scripts used in this project can be found in the git repository dedicated to this project (<u>https://github.com/ardowney/Masters_thesis</u>).



Figure 16: Aqueous phase excitation emission matrix of tryptophan, tyrosine, and humic acid. Regions one through five are indicated. Region classification from Chen et al. (2003).

5.0 Results

5.1 Sedimentary descriptions

5.1.1 Bulk grain size characteristics

Particle size data were collected on subsamples from TAL (n=156) and CAM (n=58) (Figure 17). Sand- and silt-sized grains composed most of the sediments from both locations. The samples from TAL contained a slightly higher proportion of clay than those from CAM. Proportion of sand and silt ranged from 0 percent to 95 percent. Clay proportions are less variable, ranging from 0 percent to 20 percent.

Medium sand was the dominant sand-sized fraction in all subsamples collected from TAL. Coarse silt was the dominant grain size in the mud-sized fraction (silt and clay) in all subsamples collected from TAL. Fine sand was the dominant sand sized fraction in all CAM subsamples. Coarse and medium silt was the dominant mud sized fraction in all CAM subsamples.

All sediments collected from TAL and CAM are categorized as poor to extremely poorly sorted (>2.0). Inspection in the lab with a magnifying glass showed that many of the more "clayy" layers contained small amounts of fine to medium sand. It is likely that the presence of these sand grains resulted in the poorly sorted classification. Another factor potentially controlling the poorly sorted classification are the cohesiveness of clay particles resulting in clumping of smaller clay particles into larger grains and skewing the results collected on the particle analyzer. Though precautions were taken to avoid clumping (section 4.2.1) completely eliminating clumping of clay particles is effectively impossible.

Due to the cohesiveness present in the TAL sediments, a high proportion of clay was expected from the particle size analysis. This was not the case, suggesting that only a small proportion of clay is needed in order create a cohesive sediment. Additional cohesiveness could be a result of microcrystalline Fe, Al, and Mn oxide coatings on the sediment grains.

5.1.2 TAL surface sediments

Surface sediments were collected from one, two and three meters away from the stream throughout the cave (Table 5). Mean grain size was largest (140 μ m) in the back of the cave, 2 meters from the stream (TAL-E2-01-00). This sample contained 68 percent sand and had mean grain size larger than the surrounding sediment (Figure 18). This sample also had the poorest

sorting (7.1; extremely poorly sorted). Sediment with the lowest mean grain size (12 μ m) is also located in the back of the cave, one meter away from the stream. This sample contained 99 percent fine-grained materials (silt and clay). In the middle of the cave (TAL-C), mean grain size one, two, and three meters away from the stream had little variation but sediment furthest away from the stream were the least sorted (Figure 18).

Sorting values on all surface sediments are >2.0, categorizing them as very poorextremely poorly sorted (Appendix VI). Lower sorting values indicate better sorting. There is a negative correlation between percent silt and sorting (Figure 19). The correlation between sorting and percent clay is not statistically-significant ($R^2=0.12 p$ value > 0.05). Within the fine-grained material (silt and clay) course silt is the dominant size. Within the sand sized fraction, medium sand is the dominant grain size. There is no apparent trend in grain size or sorting spatially from the back to the front of the cave (Figure 18).

Sediments collected from one, two and three meters from the stream also had no apparent trend between location with either grain size or sorting. This is likely a function of the complex geomorphology of the cave passage. Groundwater flow before the stream enters the conduit in the back portion of the cave (labeled sump in figure 9) limits the material able to enter the cave. This controls grain size transport. Fluvial flood deposits include larger grains than those within TAL. While mostly composed of sub-sand sized grains, the poorly sorted nature of surface sediments in TAL indicates allogeneic input and rapid deposition during flood waters. This is further evidenced by the onset of turbidity accompanied by a rain event in the cave stream (pers. comm. Pantoja-Agred, 2016). Very coarse to coarse sand in the front of the cave suggests that high flow events supply enough energy to transport larger grains from the sump to the dammed portal.

It is important to note that well-rounded, pebble-sized clasts were present in the stream channel throughout the cave. This suggests that the cave stream has winnowed out the sediment, presumed to have previously filled TAL, leaving behind only the larger clasts in the stream channel.

Relative location within cave	Sample location	Distance from cave stream (m)	Tracking ID
Front	TAL-A	1	TAL-A1-01-00
		2	TAL-A2-01-00
		3	TAL-A3-01-00
Passage	TAL-B	1	TAL-B1-01-00
between front and middle		2	TAL-B2-01-00
		3	TAL-B3-01-00
Middle	TAL-C	1	TAL-C1-01-00
		2	TAL-C2-01-00
		3	TAL-C3-01-00
Passage	TAL-D	1	TAL-D1-01-00
between middle and		2	TAL-D2-01-00
back		3	TAL-D3-01-00
Back	TAL-E	1	TAL-E1-01-00
		2	TAL-E2-01-00
		3	TAL-E3-01-00

Table 5: Surface sediment sample locations and associated tracking IDs.



Figure 17: Grain size distribution in El Tallonal Cave (TAL) and Cueva Clara (CAM). Sand, silt, and clay are reported as a fractional proportion.



Figure 18: A) Sediment from the sediment bank surface were collected one, two, and three meters away from the cave stream at locations TAL-A through TAL-E. Spatial variation in B) grain size and C) sorting of sediment surface samples from El Tallonal Cave (TAL). Dashed lines represent poorly sorted (sorting value 2-4) and extremely poorly sorted (sorting value >4). Sorting values are unitless. Spatial variation in grain size and sorting of sediment surface samples from El Tallonal Cave (TAL). Sediment from the sediment bank surface were collected one, two, and three meters away from the cave stream at locations TAL-A through TAL-E. Dashed lines represent poorly sorted (sorting value 2-4) and extremely poorly sorting (sorting value >4). Sorting values are unitless.



Figure 19: Relationship between percent silt and sorting from surface sediments collected one, two, and three meters away from the stream at locations TAL-A through TAL-E.

5.1.3 TAL core sediments

The dammed portal raises the water level in the cave and lowers the natural flow velocity. Since the dam must be released for researchers to enter the cave, the exact water level when the dam is operational is not exactly known but can be inferred based on the geomorphology of the sediment banks in the middle of the cave. The cave stream flows through the middle room of the cave, resulting in deposition of sediment banks on either side of the stream. The northern sediment bank contains an erosional "terrace" (Figure 12). The sediment bank on the southern side of the cave stream is lower in height and inferred to be the erosional terrace on that side of the cave stream based on the meander bend at that location (Figure 9). Water level when the dam is closed submerges most of the sediment bank on the southern side of the stream and reaches the bank terrace level on the northern side of the stream. Sediments from TAL are deemed saturated or unsaturated based on the location relative to the erosional terrace. The erosional terrace was present prior to the flooding resulting from Hurricane Maria and therefore not caused by the hurricane.

Course silt is the dominant fine-grained material and medium sand is the dominant sand sized material, consistent with trends observed in the surface sediments. Contrary to the surface sediments there is no significant correlation between sorting and percent silt, or sand was observed ($R^2=0.09 p$ value >0.05).

Redox zonation results from variation in oxygen availability in a natural environment. Speciation, precipitation, or complexation may occur based on the redox conditions of the surrounding environment. The type of redox zonation is controlled by the thermodynamic availability of electron acceptors in an environment (Pinti, 2011). Redox zonation within a sedimentary environment is often observed as color changes resulting from speciation or mineral precipitations. Cores collected from TAL exhibit little "true" stratification. Most layering in the cores are due to variation in cohesiveness, color change, or redox zonation. Cohesive layering below a loose sediment layer separated by a sharp boundary is common.

5.1.3.1 TAL-14

Three cores were collected at TAL-14 sample location. These cores were collected consecutively with depth. Gaps between core segments are due to incomplete recovery. Grain size in these three cores are 50±10 percent sand sized (Figure 20). Proportion of sand decreases

to between 9 and 6 percent from 55 to 60 cm depth. There is a pebble sized clast of limestone was observed at 44 cm depth at a boundary between sand and silt layer (Figure 20). A dominantly sand layer from 43 to 52 cm depth contained small (1 ± 0.5 cm) clay "chunks" presumed to be rip up clasts.

5.1.3.2 TAL-15

Two consecutive cores with depth were collected at TAL-15 sample location (Figure 20). Gaps between core segments are a result of incomplete recovery. A cohesive layer from 6-12 cm contained a broken piece of a speleothem (soda straw; 2 cm in length). Below this cohesive layer is black material that continues through the rest of the core. This black material is angular and brittle. The top of the second core at TAL-15 sample location marks a drastic change in composition and sedimentary characteristics. Layers of the black material are separated by thin layers (0.5 cm) of cohesive finer grained material. Multiple larger clasts (pebble sized) of this black grained material were observed. These larger clasts were sand-sized grains cemented together. The cause of this cementation is unknown. These cemented grains were individually stratified with sequences of brown/reddish brown and black banding within the larger cemented grains. In total, three sequences of this black material followed by a thin clay layer are present. A sandy layer containing well-rounded pebble sized quartz grains. These grains are likely the same material as those observed in the active stream channel. A total of five pebble sized quartz grains were observed in a "pocket" within a thin (1 cm) sandy layer. At approximately 45 cm depth a pristine crystal of calcite is present.

Manganese oxides precipitate preferentially on quartz (Ford and Williams, 2007). While no quantified data were collected on the elemental composition of any of the black material in this core, unquantified XRF scans were. Raw XRF spectra collected from this black material shows higher cps in the energy range associated with Mn florescence. Based on this it likely that the black material is indeed manganese oxide.

Core locations TAL-14 and TAL-15 are located 1.5 m apart. Despite being collected proximal from the same bank, these two core locations have little similarity. No sediment layering or stratification similar enough to be correlated between the two core locations are present.



Figure 20: Core photographs and proportion of sand from location TAL-14 and TAL-15.

Cores TAL-14 and TAL-15 sampled locations were the only core location in which consecutive cores were collected with depth. The lack of stratigraphic correlation between these two locations suggests that correlation of the remaining core sample locations is implausible. There is too much spatial variation between the cores to make correlative interpretations. This is also illustrated by the high variability observed in the sediment characteristics of the previously described surface sediments (Figure 18).

5.1.3.3 TAL-16

Sample location TAL-16 consists of eight sediment cores collected along a diagonal transect along the northern sediment bank in the middle of the cave (Figure 12). Each core is treated separately. This is because the cores are not consecutive in depth and do not contain continuous, correlatable layers.

Four cores were collected from the bank above the erosional terrace and four were collected below the terrace (Table 6; Figure 12). Core TAL-16-01 is dominantly sand (60 percent) and contains no stratification. At 16 cm depth is a thin layer (1 cm) of approximately 100 percent fine-grained material (Figure 21). The transition from this fine-grained layer to the sandier layer above is marked by a sharp boundary in sedimentary texture and no color change. Notable sedimentary textures include a small piece (0.5 cm) of black material and a larger (1 cm) chunk of clay, likely a rip-up clast.

Core TAL-16-02 contained little to no variation in grain size and lacked any stratigraphic layering or color change. The entire 24 cm of core is 50 ± 5 percent sand (Figure 21). Core TAL-16-03 is 50 percent sandy material. The first 13 cm of core is more cohesive and less sandy than the bottom 7 cm, however, based on the particle size data the proportion of sand similar in both layers. A small piece (0.2 cm) of black material was observed at nine cm depth (Figure 21).

Core TAL-16-04 is the last core on the upper terrace. It is closest to the water table when the dam is closed. This sample is extremely cohesive. Hand force was not enough to slide the core sleeve into the sediment. A "chunk" of this approximately 20 cm section of the bank was sampled instead of collecting a true core. It is dominantly fine-grained (60-90 percent mud-sized) and has little variation in sedimentary texture. There are "hair-like" strands of black material that transitions into distinct banding (Figure 21). It is likely that the banding observed in this core is redox zoning resulting from changes in oxygen availability due to fluctuations in water level.

Core	Saturated/Unsaturated
TAL-16-01	Unsaturated
TAL-16-02	Unsaturated
TAL-16-03	Unsaturated
TAL-16-04	Unsaturated/saturated depending on water level
TAL-16-05	Saturated
TAL-16-06	Saturated
TAL-16-07 TAL-16-08	Saturated Saturated

Table 6: Cores collected from sample locationTAL-16 and position relative to the dammedwater level.



Figure 21: Core photographs and proportion of sand from TAL-16-01, TAL-16-02, TAL-16-03, and TAL-16-04. Cores are not displayed in stratigraphic order.

The next four core are subaqueous when the dam is closed. Core TAL-16-05 varies from 40-50 percent sand and contains no sedimentary layering. Proportion of sand decreases with depth (Figure 22). Color changes were observed as horizontal redox banding. The top two and bottom two cm of core TAL-16-06 were dominantly fine-grained (80 percent) (Figure 22). No sedimentary layering is present. Core TAL-16-07 is dominantly sand (50-60 percent). Redox zones rather than sharp banding were observed in core TAL16-07 (Figure 22). A small piece (0.5 cm) of black material was observed at 5 cm depth, possibly a manganese oxide coating. Core TAL-16-08 is 50±10 percent sandy material. Just as the three other subaqueous core, no sedimentary layering was observed. Redox banding separated by a gradational boundary was observed as well (Figure 22).

Redox banding present in the four subaqueous core varies between two distinct colors: greenish (Munsell color: 5Y 5/3) and black (Munsell color: 2.5Y 4/1). The boundary between the redox bands are variable in shape and clarity, some boundaries are sharp while others are gradational.

5.1.4 CAM core sediments

Cores were collected from various locations through CAM (Figure 10). In general, all CAM sediments contained slightly-higher proportions of sand than sediment from TAL (Figure 17). Core from CAM contained more distinct sedimentary layering than those collected from TAL. The redox banding observed in TAL sediments are absent in CAM sediments likely because the sediments in CAM were recently deposited in a dry cave and have not undergone repeated wet/dry cycles like the TAL sediments.

Large desiccation cracks covered the sediment surface throughout the entire CAM location. Again, the top 30.5 cm of sediment were cored at various locations throughout the cave. This sediment was likely deposited by back flooding of the Río Camuy River as a result of the flooding caused by Hurricane Maria. The flooding entered the cave from the Emplane sinkhole in the northeast portion of the cave (Figure 10). The water remained in the cave until the water levels in the Río Camuy River returned after the floodwaters receded. The core collected throughout the cave is variable spatially and no obvious stratigraphic correlation was observed between the cores.



Figure 22: Core photographs and proportion of sand from TAL-16-05, TAL-16-06, TAL-16-07, and TAL-16-08. Cores are not displayed in stratigraphic order.

Core CAM-05-01 is composed almost entirely of sand while CAM-07-01 collected 50 cm away is dominantly mud (Figure 23). Core CAM-07-01 contains blocky clay chunks that are likely rip up clasts. This same texture was observed in core CAM-01-01 to a lesser degree. Core CAM-01-01 contained the most obvious stratigraphic layering. Angular pebble-sized carbonate clasts were observed in the top 10 cm of CAM-10-01 (Figure 23). It is unknown where these clasts originated but are likely sourced from the lower levels of the Río Camuy cave system or the region surrounding the Emplame sinkhole. Carbonate clasts were not observed in any of the other core collected from this cave.

Black material was observed within core CAM-01-01 and CAM-07-01 (Figure 23). The black material, possibly organic, is contained in the bottom 10 cm of both cores. This material was inferred to be organic rather than a manganese oxide due to the charcoal like texture and brittle nature of the material. Manganese oxide in TAL sediments was amorphous and microcrystalline, contrary to the black material in this core. Remaining core from this location lack this black material.

All particle size data collected as part of this project are found in Appendix VI.

5.2 Elemental data

5.2.1 Carbon and nitrogen

All surface sediments from TAL (Table 6) and core samples from location TAL-16 were analyzed for TOC, TIC, and TN (Appendix VII). Concentrations of TOC range from 0.01 to 0.2 mol/kg. These concentrations converted to percent of organic carbon range from 0.13 percent to 2.4 percent. Concentrations of TN range from 2.9×10^{-3} to 1.7×10^{-2} mol/kg. Concentrations of TIC range from 0 to 4.25×10^{-2} mol/kg, or 0 to 0.5 percent. Molar ratios of organic C: N range from 3.90 to 12. 2 (Table 7). There exists a strong positive correlation between TC (mol/kg) and TOC (mol/kg) (R²=0.97 *p* value < 0.05) (Figure 24). The R² value of the TC vs TIC relationship is lower but still significant (R²=0.52 *p* value < 0.05) (Figure 24). On a log-linear scale, percent organic carbon and organic C/N are positively correlated (R²= 0.88 *p* value <0.05 (Figure 25). There is a strong positive relationship between percent organic carbon and TN (R²=0.94, *p* value <0.05). Location TAL-C1-01-00 collected from the middle of the cave contained the highest percent organic carbon value (1.58 percent) out of all surface sediment samples (Figure 26).



Figure 23: Core photographs and proportion of sand from CAM-01-01, CAM-05-01, CAM-07-01, and CAM-10-01. Cores are not displayed in stratigraphic order.

At sample locations TAL-A, TAL-B, and TAL-C percent organic carbon was highest one meter away from the stream and decreased with increasing distance from the stream (Figure 26). At sample location TAL-E, the trend is opposite, organic carbon values increase with increasing distance from the stream. Sample location TAL-D contained the highest percent organic carbon value two meters away from the stream. Surface sediments from the front of the cave contains the highest TIC concentrations relative to the other collected surface sediments. Surface sediments collected from the back of the cave contained the lowest TIC values relative to the other collected surface sediments (Figure 26). Surface sediments contained higher TN values than either the core samples.

Sample location TAL-16 consisted of eight cores collected in a diagonal transect across the sediment bank (Figure 12). TAL-16-01 through TAL-16-04 are not saturated when the dam is closed. The bottom four core samples are saturated (Table 6). The top 1 cm of sediment collected from the first four core have a consistent decrease in organic carbon with depth. Sediment collected from the top 1 cm of each saturated core contain higher organic carbon values than the unsaturated sediment core (Figure 27).

Sediments collected from the portion of the bank that are saturated contain higher concentrations and a larger range of organic carbon and TN (Figure 28). Organic C/N values are generally higher in saturated samples except for two unsaturated samples with higher organic C/N values (Figure 28). The ranges of organic C/N in both the saturated and unsaturated sediments is comparable, although the saturated samples contained slightly-higher organic C/N values.

5.2.2 Fluorescence spectroscopy

Fluorescence data were collected on organic matter extracted from solid sediment using an alkaline extraction solution of the subsamples collected from TAL-16-01 through TAL-16-05 (Figure 29). Cores TAL-16-01 through TAL-16-05 includes all unsaturated sediment cores and one saturated sediment core (Table 6). The number of saturated samples analyzed by fluorescence spectroscopy FS (n=5) differs from the saturated samples analyzed for TOC, TC, TIC, and TN (n=38). Due to the limited number of saturated sediment extractions analyzed, the FS data collected provides a preliminary comparison.

		TOC ^a	TIC ^b	TC℃	TN ^d	Organic
		(mol/kg)	(mol/kg)	(mol/kg)	(mol/kg)	C/N ^e
Surface (n=15)	Minimum	4.91x10 ⁻²	1.67x10 ⁻³	5.99x10 ⁻²	6.43x10 ⁻³	7.63
	Maximum	1.32x10 ⁻¹	3.33x10 ⁻²	1.48x10 ⁻¹	1.36x10 ⁻²	1.22
	Average	8.42x10 ⁻²	1.35x10 ⁻²	9.77x10 ⁻²	9.09x10 ⁻³	9.20
	Standard deviation	2.44x10 ⁻²	9.42x10 ⁻³	2.50x10 ⁻²	2.06x10 ⁻³	1.35
Un- saturated (n=24)	Minimum	1.08x10 ⁻²	< MDL ^f	1.08x10 ⁻²	2.86x10 ⁻³	3.50
	Maximum	6.08x10 ⁻²	2.00x10 ⁻²	7.99x10 ⁻²	8.57x10 ⁻³	9.10
	Average	2.71x10 ⁻²	3.68x10 ⁻³	3.07x10 ⁻²	5.32x10 ⁻³	5.00
	Standard deviation	1.11x10 ⁻²	5.37x10 ⁻³	1.59x10 ⁻²	1.17x10 ⁻³	1.36
Saturated (n=38)	Minimum	1.63x10 ⁻²	< MDL ^f	1.08x10 ⁻²	2.86x10 ⁻³	3.50
	Maximum	1.96x10 ⁻¹	4.25x10 ⁻²	2.29x10 ⁻¹	1.71x10 ⁻²	1.20
	Average	6.49x10 ⁻²	1.41x10 ⁻²	7.90x10 ⁻²	7.96x10 ⁻³	7.29
	Standard deviation	4.75x10 ⁻²	1.21x10 ⁻²	5.62x10 ⁻²	3.42x10 ⁻³	2.45

Table 7: Summary table of all collected carbon and nitrogen concentrations from sample location TAL-16. Surface samples are not included in the un-saturated sample set.

^a Total organic carbon

^b Total inorganic carbon

^cTotal carbon

^d Total nitrogen

^e Organic carbon:nitrogen ratio ^f Less than method detection limit



Figure 24: Relationship between total carbon (TC), total organic carbon (TOC), total inorganic carbon (TIC), and total organic carbon from El Tallonal Cave sediments (TAL).


Figure 25: Relationship between total nitrogen (TN), organic C/N ratio, and total organic carbon (TOC; %) from El Tallonal Cave sediments (TAL).



Figure 26: A) Sediment from the sediment bank surface were collected one, two, and three meters away from the cave stream at locations TAL-A through TAL-E. Spatial distribution of B) fraction of organic carbon (FOC); C) total inorganic carbon (TIC); and D) organic C/N ratios of all surface sediments collected from El Tallonal Cave (TAL).



Figure 27: Total organic carbon (TOC; %) collected from the top of each core collected from TAL-16 ordered by depth relative to the datum. Dammed water level is 118 cm below datum.



Figure 28: Variability in organic C/N ratios, total nitrogen (TN), and total organic carbon (TOC; %) between saturated and unsaturated sediments. Dammed water level is 118 cm below datum.



Figure 29: Excitation-emission matrix (EEM) data collected on the alkaline extraction from TAL-16 sample location subsamples. 1^{st} and 2^{nd} order Rayleigh scattering has been removed. The EEMs have been normalized to the max intensity of each scan. EEMs were collected from λex : 240-400 nm and λem : 280-540 nm.

The FS data collected from these sediments had BIX values from 0.68 to 1.87 and FI values from 0.89 to 1.54 (Figure 30). The highest BIX value was obtained from 12 cm below the top of the sediment bank surface. Lower BIX and FI values were found in samples with the highest organic carbon values (Figure 30). Samples with a BIX value >1 are composed of fresh autotrophic organic matter. All saturated samples have BIX values >1. Unsaturated samples with BIX values >1 are also present. There is no clear connection between proximity to the cave stream and BIX values >1. Samples with organic matter containing FI values >1.90 are considered to originate from a microbial origin. There were no sediments in this sample set containing FI values >1.90. The maximum FI value (1.50) was found in core TAL-16-04. This is the closest unsaturated core to the cave stream and is likely to experiences partial submergence with fluctuating water level. Unsaturated samples generally had lower organic C/N and a large range of BIX and FI values.

Corrected EEM's with 1st and 2nd order Rayleigh scattering removed were used to create a PARAFAC model. The model was created using a combination of the "eemR" and the "staRdom" packages in R (Massicotte, 2019; Pucher et al., 2020). The model creates an output of six possible models ranging from two to seven components. There is no clear rule to choosing the appropriate number of components (Murphy et al., 2013). Models without distinct peaks in the residual terms are desired however, it is unlikely that a given model will completely describe an entire dataset. The chosen model must minimize the number of leverage outliers and the maximum absolute value of the intensity in the residuals. The absolute values of the maximum intensity are used as a PARAFAC model can result in negative residual intensities as well.

A PARAFAC model with three components was chosen as the model with the lowest number of outliers and a relatively low maximum absolute value intensity (Table 8). All three modeled components are contained in region five of the EEM space (humic acid-like) (Figure 31). The proportion of each component was normalized to the total contribution of each sample (Figure 32). Component two dominates and component one is the most variable between the samples (Figure 33). The proportion of component one is highest on saturated samples, independent of organic carbon concentration (Figure 34). The opposite is true for component three. It is lowest in the saturated samples. The sample with the largest proportion of component one also contains the highest organic C/N ratio (Figure 34). Unsaturated samples generally contain a lower proportion of component one and lower org C/N ratios.



Figure 30: Relationship between fluorescence index (FI) and biological index (BIX) and total organic carbon (TOC; %). Saturated and unsaturated subsamples are indicated. FI values <1.40 indicate organic matter of terrestrial origin (dashed line). BIX values >1 indicate fresh autotrophic organic matter (dashed line).

Table 8: Results of the six parallel factor analysis (PARAFAC) models. The model with the minimum number of emission, excitation, and sample out lairs and minimum absolute value of intensity in the residuals is desired. The PARAFAC model with three components was chosen.

Model ^a	Number of emission out liars	Number of excitation out liars	Number of sample out liars	Max residual intensity (absolute value)
2 Components	5	5	4	0.44
3 Components ^b	5 ^b	5 ^b	4 ^b	0.25 ^b
4 Components	5	6	4	0.33
5 Components	5	6	4	0.20
6 Components	5	6	4	0.25
7 Components	5	6	4	0.25
^a n=35 samples ^b Chosen model				



Figure 31: Excitation-emission matrices (EEM) of the three-component parallel factor analysis (PARAFAC) model. Regions one through five are labeled. All modeled components fall within region five (humic acid-like).



Figure 32: Proportion of the modeled parallel factor analysis (PARAFAC) components one, two, and three in the sample set categorized by core. Dammed water level is indicated.



Figure 33: Ternary diagram with the three modeled parallel factor analysis (PARAFAC) components. Components are reported as a fractional proportion.



Figure 34: Total organic carbon (TOC; %) as a function of the contribution of each component. Saturated and unsaturated samples are indicated.

Saturated sediments contained lower proportions of PARAFAC component three. Proportion of component three varied with depth in unsaturated sediments as well. The top of each unsaturated core had the highest proportion of component three and decreased with depth. This pattern was consistent in all unsaturated cores at location TAL-16 (Figure 32).

5.2.2 Elemental concentrations

Data collected from XRF analysis were quantified using a custom calibration for the four unsaturated core samples from location TAL-16 (Table 6) and the core collected from location TAL-14. Major elements (Na, Al, Si, K, P, Ca, Ti, and Fe) are reported in weight percent and minor elements (Ba, Cr, Cu, Co, Mn, Ni, Zn, Rb, Nb, Mo, and Y) are reported as ppm (Appendix VIII). Concentrations reported as 0 weight percent are present at low (below calibration detection) concentrations. Quantified data were then converted to mol/kg.

The most abundant element by an order-of-magnitude is Si (Figure 35). Aluminum is the second most abundant element and Fe is next. Principle component analysis (PCA) was performed using major and minor elements (Figure 36). Principle component analysis is a mathematical algorithm that reduces the dimensionality while simultaneously maintaining most of the variation in a dataset (Ringnér, 2008). In most cases, a dataset is reduced to two or three principle components. The selected principle components are uncorrelated and therefore, explain most of the variability in a dataset. Variables plotting close together, far from the origin, exhibit similar variation in a dataset (Bro and Smilde, 2014). Elemental concentrations plot in all four quadrants of the PCA biplot (Figure 36). Silicon and Al plot almost directly across from one another. Aluminum and Si have a negative relationship. These elements, along with the associated elemental concentrations plotting near them in the PCA space, compose group one and group two elements within this dataset (Figure 36). Groups one and two contain elements composing most mineral assemblages (i.e. silica tetrahedra and aluminum octahedra) and therefore likely represent the overall lithology. Group three, containing all other elemental concentrations plot between quadrant one and three of the PCA biplot. These elements are clustered around an x value of -1 (absolute minimum) and a y-value of zero, representing a heavy weight on component one. This reveals that most of the variation in group three elements is described by principle component one. The opposite is true for groups one and two.

A Pearson Correlation Matrix was created using the "Hmisc" (Harrell and Dupont, 2020) package in R to explore correlations between pairs of elements (Appendix XIII). These two major elements, Al and Si, were found to be strongly negatively-correlated (R=-0.86 p value <0.05) (Figure 37). As such, these relationships provide information about the lithology of these sediments. Composed of alternating layers of aluminum octahedra and silica tetrahedra, phyllosilicate minerals result in some positive correlation in Al and Si concentrations (Essington, 2015). The strong negative correlation between Al and Si suggests that quartz and clay minerals occur together, resulting in high Si concentrations and low Al concentrations. Molar concentrations of Na and K are associated with Al and Si, respectively (Figure 37). Correlation between Si and K (R=0.71 p value <0.05) is not as strong as Al and Na (R=0.99 p value <0.05). This lower R value is due to the presence of outliers with unexpectedly-high Si concentrations. Potassium is the only elements strongly correlated with Si. Aluminum is positively correlated with P and Cu (Figure 37). Copper is the only trace element associated with Al (R = 0.66 p value <0.05). Organic phosphorus accounts for 20-70 percent of the total P pool in a sedimentary environment (Essington, 2015). Within these sediments, phosphorus is negatively correlated with TOC (R= -0.71 p value <0.05). Sorption of P is known to have a strong and almost irreversible affinity for sorption to Al and Fe oxyhydroxides (Von Wandruszka, 2006). The positive correlation between Al and P suggests that most of the P pool in these sediments could be sorbed to aluminum oxide. Within a cave environment, P can be supplied by bat guano. While CAM was known to have a large bat population (before Hurricane Maria), it is unknown if TAL contains a bat population. No bats were observed during field work in TAL.

Dissolution of the carbonate rock matrix supplies most of the Ca to these cave sediments and therefore is not well correlated with other elemental concentrations (Figure 37). While Ca can be supplied by allogenic materials (e.g. silicate and phyllosilicates), Ca supplied by carbonate dissolution can be expected to outweigh Ca supplied by other sources (Essington, 2015). Water samples collected by the University of Puerto Rico, Mayaguez reveal that the water in TAL is supersaturated with respect to calcite (average SIc=0.29, n=9) (pers. comm. Pantoja-Agred, 2016). Kinetic thresholds on the precipitation of calcite suggest that at low supersaturation values, calcite (SI_c <1) will not actually precipitate the predicted value at equilibrium (White, 1988). This kinetic threshold results in more Ca in solution than is expected of which can be retained in the clastic sediments within the cave.



Figure 35: Elemental concentrations plotted in molar units from highest to lowest average value of A) the entire dataset; B) major elements; and C) trace elements. Symbols are averages; error bars are a total of one standard deviation.



Figure 36: Eigenvalues of principle components 1-7 and scatter plot of the principle component analysis (PCA). Dashed line represents the three groups of elemental concentrations present within the PCA space.

Trace elements in this project refer to any element with concentrations at or below 0.01 mol/kg (Figure 35C). Titanium is likely present in nature as the stable oxide, rutile. Due to the stability of rutile, Ti tends to accumulate in sediments and is often used as an indication to the degree of weathering. Titanium is positively correlated with several trace elements, most notably Cr, As, and Mn (Figure 37). Chromium and Ti are very strongly correlated (R=0.99 *p* value <0.05). In addition to Ti, As is positively correlated with Rb (R=0.85 *p* value <0.05), Nb (R=0.64 *p* value <0.05), Ni (R=0.75 *p* value <0.05), and Co (R=0.72 *p* value <0.05) in these sediments (Figure 37). The only strong positive correlation with Fe is Co (R=0.79 *p* value <0.05).

Manganese rapidly precipitates as microcrystalline oxides when environments transition from suboxic to oxic (Dixon et al., 1990). As a result, Mn oxide precipitation is a good indication of changing water level. Precipitation of microcrystalline Mn oxide was observed in core TAL-16-04 resulting from fluctuating water level (Figure 21). Manganese is positively correlated with Ti and Cr (R=0.69 *p* value <0.05) in these sediments.

Retention of trace elements Pb, As, Cr, Co, and Ni in a sedimentary environment occur through sorption to Al, Fe, and Mn oxides (Cerqueira et al., 2011; Dixon et al., 1990; Ma and Hooda, 2010; McKenzie, 1980). The strong positive correlation between Mn and Pb could indicate Pb sorption to Mn oxide.

Sorption and co-precipitation are a dominant retention mechanism for trace elements in a sedimentary environment. However, correlation does not imply sorption as a definite retention mechanism. This project did not focus on understanding why these correlations are present. As such, implying sorption or co-precipitation are meant as a potential cause for these correlative relationships.

The correlation matrix reveals several cases where the correlation coefficient is exceedingly-low, but still significant (e.g. Fe vs. Si) (Figure 37). Low p values (below α =0.05) implies that a change in the predictor variable is associated with a change in the response values. This is possible even with noisy and highly-variable data. While p values are a widely used statistical parameter, caution when using this parameter in large datasets is advised (Lin et al., 2013). A p value measures the distance between the data and the null hypothesis, typically using the standard deviation of the predictor values (i.e. the standard error) (Biau et al., 2010).



Figure 37: Pearson correlation matrix of elemental concentrations quantified using XRF. Matrix was created using the "Hmisc" package in R. Grey shaded values represent p-values < 0.05. Boxed values indicate an R value > 0.6 (blue shaded) and <-0.6 (orange shaded).

The standard error between any two parameters shrinks as sample size increases. Problems when interpreting p values arise when the standard error between the parameters becomes so small that even small distances from the estimated values and the null hypothesis become statistically significant. This results in p values that approach zero, not as a rejection of the null hypothesis but as an artifact of the large sample size (Lin et al., 2013). The computed correlation matrix contained many data points (n=669). Low R values and p values may be an artifact of this large sample size. That is not to say that variables with high correlation coefficients are void. There still exists significant statistical power in those highly-correlated variables.

When Ti, Cr, and As are plotted against Zr two distinct groups are observed (Figure 38). Pb and Ni also exhibit this trend. When separated out by core, this bifurcation is due to variation in core locations (Figure 38). Core collected from location TAL-14 show higher Zr concentrations than core collected from location TAL-16. An additional population of sediment with higher Ti, Cr, and As values was observed and isolated to the TAL-16-04 core sediments (Figure 38). Sediment collected in this core was between the saturated and unsaturated zones. The saturated condition of the core varies depending on the water level and as a result, Mn oxides have precipitated throughout this core. Higher concentrations of Ti, Cr, and As may be a result of the fluctuating saturated and unsaturated conditions. Given that sample location TAL-14 is further from the inferred dammed water level and that sampling was collected consecutively with depth, it is unknown if these sediments were saturated (Figure 11).

Subsamples were collected from a range of 1±0.5 cm for particles size analysis. XRF data (collected every 2 mm) was averaged from a one cm range corresponding to the collected subsample range for particle size analysis. No statistically-significant correlations between particle size and elemental concentrations were observed in the collected dataset from TAL. Subsamples collected for particle size and the XRF averaged data did not contain much overlap. The lack of trends between grain size and elemental concentrations may be due to the limited sample size. Additionally, discrepancies in exact sample depth between the averaged XRF data and the subsamples collected for particle size analysis are a potential issue effecting the lack of correlation.



Figure 38: Zirconium vs Ti, Pb, and As separated by core. Two distinct groups are observed due to variation in concentrations from different core locations. Solid blue ellipse represent TAL-16 cores and orange dashed ellipse represents TAL-14 cores.

6.0 Discussion 6.1 Sediment facies

Sediment facies classifications by Bosch and White (2007) are described through observations made in teleogenetic cave systems and based on physical sediment characteristics (Table 1). Grain sizes and sediment sorting are categorized relative to the system. Bosch and White (2007) state that the content in a cave system is dependent on the available source material and that most sedimentary deposits will not contain distinct facies. The purpose of this project was not to fully characterize the facies present in these caves. Additionally, applying these facies to sediments from TAL is somewhat difficult as the collected core showed little true "stratification" in grain size. This classification provides a loose framework for describing cave sediments. Considering bulk sedimentary characteristics, a minimum of two facies were observed in TAL: thalweg facies present in the active stream channel and channel facies present in the bank sediments (Table 1).

The thalweg facies is created when a secondary (younger) stream channel cuts through previously deposited channel facies, winnowing out smaller materials and leaving larger clasts behind (Table 1). Well-rounded cobbles and pebbles were observed in the active stream channel and similar well-rounded pebbles were found in core TAL-15-02 leading to this thalweg classification (Figure 20). Additionally, the cave stream is active in contrast to the sediment bank which have been in place for some time based on the presence of small (1-2 cm) stalactites on the top of the sediment bank in some locations. Bank sediments from TAL were poorly sorted and lacked gravel and pebble sized clasts (except for TAL-15-02). These sediments could be considered a channel facies though it is not possible to distinguish channel facies from slackwater facies here (Table 1). It is likely that these sediments were deposited through sequences of channel and slackwater facies conditions.

Sediments collected from CAM are likely to have been deposited by backflooding of the Río Camuy River resulting from flooding in response to Hurricane Maria. This interpretation is based on the flooding observed in Tres Pueblos Sinkhole located approximately 0.6 km from CAM (Figure 2). Additional evidence of this backflooding was the "bath ring" observed in the Emplane Sinkhole during sediment collection at CAM (Figure 39). For water to reach CAM, based on elevation difference, the Río Camuy River experienced a 20-meter rise in water level. This depositional condition is characteristic of the slackwater facies, however, CAM sediments

are poorly sorted and contained what are interpreted as rip-up clasts (Table 1). Slackwater facies are reported as the first centimeters of clay capping the very top of a sedimentary deposit. Sediments from CAM do not contain any such clay cap. Sediment core from CAM contained no graded bedding or sorting indicating gradual suspended load settling via waning transport energy.

Facies developed by Fornos et al. (2009) describe sediments from an eogenetic cave system. Sediments described in this study were collected from a submerged cave system. Facies descriptions include both physical sediment characteristics and organic matter concentrations. Considering the differences in saturated and unsaturated sediments observed in this project, describing both organic and physical properties are necessary when studying sediments from a submerged cave system. Facies are grouped into waterflow coarse-grained deposits, waterlain clastic fine-grained deposits, carbonate deposits, and mixed facies (Table 1). Slackwater facies from Bosch and White (2007) roughly correlates to waterlain clastic fine-grained facies described by Fornos et al. (2009). Waterlain clastic fine-grained facies include thinly-laminated brown clays, silts, and fine sands. Mineralogy of these sediments is dominantly siliceous, and composition corresponds to soil material transported by rainwater. Thinly-laminated clays and silts were also described by Bull (1981) in cave sediments within a vadose system and concluded that neither grain size nor laminae thickness were related to current direction. Waterlain clastic fine-grained facies are deposited through surface derived input and periodic accumulation or organic matter as plant material is present. This waterlain clastic fine-grained facies can be applied to sediments from TAL, though differences in sediment characteristics are present. Sediments from TAL do not contain thinly-laminated clays and silts and lack obvious organic matter laminations.

Sediment data collected from both caves are inherently biased. Neither cave was fully mapped for the presence and types of sediments and therefore it is almost certain that other facies are present but were not captured in this study. The difficulty in applying facies classification to sediments collected in this project is due to the lack of a full map, variability within caves, the generalization of facies descriptions, and the geomorphological and groundwater controls on cave sedimentation.



Figure 39: Photos of the Emplane Sinkhole during the February 2019 sampling. The Río Camuy River is approximately 20 meters below the "bath ring" A) The sinkhole photographed from the back of CAM. The overlook photographed is a popular attraction during the cave tour through CAM. B) Photograph of Emplane Sinkhole from another perspective emphasizing the "bath ring" created by the Río Camuy River backflooding resulting from the Hurricane Maria flooding.

6.2 Significance of sediment saturation

The dynamic nature of the sediment solution makes it a complex environment. Heterogeneity is a characteristic property of sediment environments. This includes the sediment solution as well. This heterogeneity facilitates the presence of microenvironments. Microenvironments can occur as a result of the physical heterogeneity of the sediment or by a limiting nutrient (e.g. O₂). There exist many microenvironments within a sedimentary environment. For example, O₂ availability around a single sediment particle can vary from 21 percent to 0 percent in just two mm (Madigan et al., 2018).

A distinct difference in organic chemical properties between saturated and unsaturated sediments was observed. The range of TOC, TN and organic C/N ratios were higher in saturated sediments. Saturated sediments also contained higher concentrations than unsaturated sediments. These saturated sediments, as the name suggests, contain more water within the pore space. This water and the elemental components within it are hereinafter referred to as the sediment solution. Chemical reactions such as hydration, hydrolysis, acid-base, oxidation-reduction and complexation occur largely within the sediment solution (Essington, 2015). These reactions control the fate and behavior of several geochemical species as well as retention via precipitation- dissolution and adsorption-desorption. In addition to facilitating chemical fate and behavior, the sediment solution is the principle mode of transport for materials into and out of these sediments. The presence of redox reactions and electromotive potential are correlated to pore water volume (Wanzek et al., 2018). The presence of extensive redox zonation in the saturated core collected from location TAL-16 illustrates this point.

Chess et al. (2010) described grain size trends and sediment facies from Butler Cave, VA. Sample selections were based on a "low and wet" and "high and dry" classification on the basis that sediment sourced from the lower portions of the cave were recently deposited in relation to sediments from higher portions of the cave. Color variation between these two sediments were quantified by collecting optical absorption spectra through diffuse reflectance spectroscopy and calculating color ratios. A distinct difference in color ratios were found between low and high sediment. This color variation is attributed to the varying degrees of hydration of iron oxides. The cause of color variation in the sediments collected from this project was not explored but may be due to variable degrees of hydration of iron oxides like those described in Chess et al. (2010)

The physical differences between core TAL-14 and TAL-15 (1.5 m apart) could be attributed to the presence of microenvironments (Figure 20). The section of core TAL-15 containing Mn oxide contained larger sand grains, creating more pore space and potentially a larger volume of sediment solution facilitating the precipitation of Mn oxide. Differences between saturated and unsaturated sediments were less evident in the collected FS data and PARAFAC model. This is partially due to the low number of saturated samples analyzed as previously discussed in section 5.2.2. However, the saturated samples included in the PARAFAC model contained a lower proportion of component three (Figure 32). A PARAFAC model including additional saturated samples is recommended in order to more accurately evaluate these data. The large range of organic C/N ratio in saturated sediments suggest that C/N ratios may not be indicative of changing carbon sources. All saturated samples had a FI value <1.40 indicating DOM of terrestrial origin opposed to microbial origin. While this might be the case, it is important to remember that the indices calculated in this study were created for non-cave environments. The environmental conditions within a cave system may lead to incorrect interpretation of these data. While the calculated indices are still a valuable tool for evaluating organic matter properties, a re-evaluation may be necessary when characterizing organic matter in a cave environment.

Cave sediments contain an archive of potentially critically important paleoclimatic information. The data presented in this project suggest that paleoclimate reconstructions should be focused on unsaturated sediments or those that do not experience repeated wet-dry cycles in order to minimize the amount of post depositional diagenesis. Considering the dynamic nature of cave systems and complex connection between cave sedimentation, cave morphology, and speleogenesis, it remains difficult to eliminate the possibility of sediments being saturated in-situ at some point in a caves evolution.

6.3 Input of clastic sediments

Sediments from TAL and CAM indicate that there is likely a mixture of carbonate and igneous source materials. The onset of turbidity coinciding with a rain event at TAL is evidence that sedimentation is dominantly surface derived in TAL The source area for most sediments in this region of PR is the central volcanic highlands (section 3.1). The Luquillo Critical Zone Observatory located in the northeastern portion of the central highlands. Quartz diorite and

volcaniclastics ranging from basaltic to andesitic are the two dominant lithologies in that region (Figure 40). Bedrock weathering rates in the Luquillo Mountain region of PR were found to experience the fastest chemical weathering rates of any granitoid rocks on earth's surface (Schulz and White, 1999). Weathering rates of the volcaniclastic lithologies were found to be several orders of magnitude higher than the quartz diorite (Buss et al., 2017). Some amount of clay minerals was present in soils above both lithologies though the quartz diorite soil contained a higher proportion (Table 9). Soils above the Luquillo Mountain lithologies provide information on the chemical weathering processes resulting in the alteration of primary igneous minerals into secondary clay minerals. Physical weathering processes also play an important role during transport from the central highlands into the karst system. The Luquillo Mountain region is not the exact source area of the sediments in TAL but still provide a useful preliminary comparison.

Sediments from TAL are depleted in Si, Al, Fe, K, Sr, Y, and Nb when compared to the Luquillo Mountain soils from both lithologies (Figure 40). Depletion in TAL sediments may suggest that these elements are not transported and sequenced within the cave system. While this could be the case it is important to consider how complexation and speciation may affect the mobility of these sediments both in the conduit, after deposition, and during transport from the source area. Similar concentrations of Ti and Zr in TAL sediments compared to the Luquillo Mountain soils could be an indication these elements are conservative during transport from volcanic highlands into the cave system (Figure 40). The increase in Ca concentrations between TAL sediments and the volcanic soils is possibly due to the additional supply of Ca into the system from carbonate dissolution. However other cave sediments within the northern karst of PR have not been analyzed and therefore no data exists to confirm the Ca source.

Major elements Al and Si are strongly negatively correlated in these sediments. The reason for this is unknown. The negative correlation may suggest a lack of clay minerals. However, chemical weathering of the central highlands has been found to produce secondary clay minerals quite rapidly (Table 9). It is more likely that clay minerals and quartz occur together resulting in Si dominating in those sediments. This is evidenced by the poorly sorted nature of TAL sediments. Sections of core with low Si concentrations and high Al concentrations could indicate a lack of quartz or a high proportion of Al oxides. X-ray diffraction was not included in this project; therefore, these interpretations cannot be validated.

Quartz diorite ^a					Volcaniclastics ^b					
Bedrock			Soil		Bedrock		Soil			
%	Mineral	%	Mineral	%	Mineral	%	Mineral			
56.4	Plagioclase	53.7	Quartz	36.0	Plagioclase	56.0	Kaolinite			
24.9	Quartz	27.1	Kaolinite	24.0	Chlorite	26.3	Quartz			
9.50	Biotite	15.7	Biotite	10.0	Quartz	11.3	Illite			
6.30	Hornblend	3.40	Goethite	9.40	Pyroxene	3.90	Goethite			
1.90	Geothite			8.00	Epidote	1.70	Hematite			
1.80	K-feldspar			5.80	K-feldspar					
				3.60	Amphibole					
				0.40	Kaolinite					
^a Bedrock and soil data from White et al. (1998)										
^b Bedrock and soil data from Buss et al. (2017)										

Table 9: Mineral assemblages and proportions of the quartz diorite bedrock and soil and the volcaniclastic bedrock and soil from the Luquillo Mountains.



Figure 40: Elemental composition of A) bedrock and B) overlying 0.2 m of soil collected from the two dominant lithologies in the Luquillo Critical Zone Observatory. C) Average elemental concentrations of El Tallonal Cave (TAL) sediments plotted against the soil elemental concentrations of the two lithologies. Error bars represent one standard deviation total. The Luquillo Critical Zone Observatory data are from White et al. (1998) and Buss et al. (2017).

Group three elements observed in the PCA analysis include almost all the trace elements in the dataset (Figure 36). Trace elements are rarely found in discrete mineral phases in a sedimentary environment. Rather, these elements exist as inclusions, incorporated into a mineral structure through isomorphic substitution, associated with a mineral structure (as well as organic matter), through sorption or co-precipitation (Essington, 2015). Because trace elements are not inherently connected to discrete mineral phases, concentrations tend to be highly variable depending on the environment. This property makes geochemical trends in trace metal concentrations potentially important in understanding cave sediment sources and transport.

Several trace elements correlate positively with Ti in these sediments (Figure 37). The exact mechanism controlling this is not known. One hypothesis is that Ti and those trace elements are supplied from the same volcanic source. Titanium is generally present in sediment as a highly stable oxide rutile and has historically been used as an indication of weathering. The concentration of Cr in lake sediments has been found to increase with decreasing grain size (Mwamburi, 2016). Chromium concentration in the central highlands as of this project has not been evaluated. Though Cr concentrations tend to be elevated in sediments sourced from maficultra mafic parent material (Essington, 2015). The strong correlation between Cr and Ti suggests that Cr could be used as an additional indication to the degree of weathering.

Organic matter from epikarst, terrestrial plants, and autotrophic metabolism has been successfully traced to a sinkhole through time using physical and chemical sediment properties, namely ¹³C isotope analysis (Tamalavage et al., 2018). This work demonstrates that it is possible for sediment characteristics to vary depending on the source. An important distinction between sediments deposited at the base of a sinkhole and sediments within a cave system is that sediments deposited within a sinkhole have not yet entered the subsurface. This fact minimizes some complexity when tracing a sediment source. A groundwater flow path in a karst aquifer almost always contain some unknown or unexplored portion and therefore adds uncertainty exactly where allogenic sediments enter an aquifer. The lag time between a rain event and the onset of temperature change within TAL is effectively zero (pers. comm. Pantoja-Agred, 2016). This suggests that surface water is entering the cave system rapidly and potentially not far from the main input location in the back compartment of the cave (Figure 9). However, tracing groundwater flow at TAL has been unsuccessful (pers. comm. Padilla, 2019). Identifying sediment source areas to a sinkhole is a promising conclusion that suggests sediment sources to a

cave system can more confidently be traced when the groundwater flow path has been well established.

6.4 Karst water and karst sediment organic matter

Dissolved organic carbon in karst aquifers and cave systems has been well studied (Birdwell and Engel, 2010; Brown and McKay, 2009; Hartland et al., 2012; Husic et al., 2017; Simon et al., 2007; Simon et al., 2010). Organic carbon within suspended sediment from a cave stream has been found to be retained and/or degraded through a caves flow path (Husic et al., 2017). Post-depositional sediment organic matter has not been explored to the same extent. Results from this project suggest significant post depositional diagenesis and organic matter transformations are important processes controlling the fate and behavior of chemical species. Organic matter characterization via FS has been a prominent area of research within soil, agriculture, and waste management studies and provides valuable information about the type of organic matter present in karst sediments.

Parallel factor analysis was done on cave water DOC by Simon et al. (2010). A threecomponent model best fit the variation in fluorescent characteristics. Component two from their model closely matched component two from the PARAFAC model in this project. Component two in Simon et al. (2010) was similar to humic-like compounds and was represented in soil above the epikarst zone. The Simon et al. (2010) sample set was small (n=18) and as a result lowers the ability of the PARAFAC model to produce statistically significant results and potentially skews the effect size of the sample set. Additionally, DOC from the soil samples were extracted using deionized water (DI) and not an alkaline solution and therefore only water extractable organic matter was analyzed. These factors make it difficult to perform any quantifiable comparison to the EEM data collected in this project.

A dominantly-microbial signature was observed (FI >1.9) in groundwater DOC from a confined portion of the Edwards aquifer leading to the conclusion that surface input stimulates microbial activity (Birdwell and Engel, 2010). No FI values were higher than 1.9 in the analyzed organic matter extracts in this project. The extent of microbially-generated organic matter is unknown in TAL. Given the increase in total organic carbon in saturated sediments, it is likely that if organic matter production is occurring, it is generated within the saturated sediments. Though highest FI and BIX values were observed in unsaturated sediments. The lower values of

FI in this project when compared to Birdwell and Engel (2010) suggest that applying these predefined indices to environmental conditions differing from those used to define them eliminate portions of EEM data that are potentially informative in understanding organic matter dynamics.

Panno et al. (2004) measured organic carbon and collected radiocarbon age dates of clastic cave sediments from Illinois and found that organic carbon concentrations decreased with increasing depth from the bank surface and were stratigraphically consistent with no age reversals. This finding is consistent with organic carbon concentrations measured in unsaturated sediments (Figure 27). Age dating was not done as part of the PR project and remains unknown if this decreasing organic carbon concentrations in unsaturated sediments are correlated to age.

6.5 Relevance to storm events

Fine-grained cohesive sediments present in a cave system has important implications for contaminant fate and transport in a karst aquifer. Clastic input into cave or conduit must first enter the subsurface, most often through groundwater flow. The presence of these cohesive sediments in a groundwater flow path with known organic contaminants may have a high potential for sorption onto the sediment particles. Cave sediments associated with contamination will cause retention of a contaminant in an aquifer. Sediments within a flowing water increase transport. Large storm events such as Hurricane Maria will mobilize previously deposited sediments and the associated sorbed contaminants, potentially entering public or private water supplies.

Sediment mobility is within a conduit is marked by a threshold (Herman et al., 2008). The installed datalogger in TAL shows that turbidity coincides with even small rain events (pers. comm. Pantoja-Agred, 2016). In contrast, sediment input within CAM occurs only once the Río Camuy river reaches the cave level. Cueva Clara cave does not experience continuous groundwater input as it is an upper (older) level of the Río Camuy system and has since been abandoned by the Río Camuy river (Miller et al., 2017). The differences between these two systems show that a threshold marking the onset of sediment transport and surface derived input is highly variable and specific to a given system. Collins et al. (2015a) documented cave sedimentation from a submerged cave in Yucatan with Hurricane Ingrid (Category 1) and found no deposition associated with the 0.7 m rise in the water table. Considering the differences between the two caves included in this project puts this into context. The cave systems included

in this project are drastically different when it comes to sediment deposition in response to a rain event. Deposition depends both on cave morphology, groundwater flow, and the magnitude of the rain event. These factors must first be considered when investigating sediment deposition in response to major storm events.

Identifying paleostorms within the sedimentary record is an area of increased research as frequency of global catastrophic weather events has increased as a result of climate change (McCloskey and Keller, 2009). It is important to uncover how past climate change has impacted the frequency and severity of extreme weather events to further understand how it will affect the occurrence of extreme weather events now and in the future. Donnelly (2005) reconstructed a record of storm surge deposits from a backbarrier salt-pond located in PR going back 2000 years. This reconstruction was done based on the presence of courser grained material within the sedimentary record of this salt-pond. Cave morphology can restrict the flow in such a way that grain size changes may not be indicative of floodwaters. The sediment collected from CAM were known to be deposited through floodwaters but contained only sub sand sized grains. Properties other than grain size (i.e. trace element, organic matter, isotopes) will provide more information when reconstructing paleostorms using cave sediments.

Caves act as a sediment sink. Clastic input is injected into a subsurface depositional environment, protecting it from the erosive processes occurring within subaerial systems (Gale, 1984). This makes cave sediments important archives of paleostorm data. Despite the complexity of cave sedimentation, it remains worthwhile to continue exploring these deposits as paleoclimate archives. While the storm record reconstructed by Donnelly (2005) is still significant finding, coastal sediments are continuously reworked and remobilized. This coupled with the direct influence of the ocean and inability to distinguish between tropical storms, hurricanes, or tsunamis places some uncertainty in the paleostorm record. Cave systems located inland eliminates the influence of tsunamis or other localized coastal flooding events and distinction often difficult to make for coastal deposits (Gabriel et al., 2008; van Hengstum et al., 2008; White, 2007). Additionally, cave sediments record deposition going back potentially much further than 2000 years depending on the system. The sedimentary records preserved within cave systems have potential to not only evaluate and confirm the pre-existing paleostorm record but extend it further into the past.

7.0 Recommendations for future work

The data collected during this projected revealed that sediment input into a cave system is exceptionally complex. This is not implying that storm signatures are unidentifiable, rather this project suggests that many more data are necessary in order to identify source areas and transport mechanisms before a confident assessment of storm deposits can be made.

7.1 Organic matter

As previously addressed, dynamic geochemical cycling was observed in the saturated sediments from TAL This observation only hints at the complexity of the geochemical processes occurring in these sediments.

A time series study focused on organic matter transformations and degradations in the cave system would be particularly valuable. Time series studies over a long enough time period will useful for paleoclimate and paleostorm reconstructions. Interactions between sediment and water should also be studied at a smaller, more detailed resolution. In order to use organic matter as a marker of allogenic input, a complete understanding of in-situ organic matter transformations is necessary. A comparison between the organic matter transformations in saturated sediments could aid in a better understanding of organic matter degradation and recalcitrance. The lack of photosynthetic metabolism in the cave environment provides a useful comparison to similar studies performed in subaerial environments.

From the data collected in this project, distinction between in-situ microbial organic matter and allogenic organic matter is not possible. Further work on FS analysis must be done in order to confidently characterize organic matter sources (Appendix III). This project performed organic matter extractions using an alkaline solution. Improvements in instrument capabilities in recent years have allowed researchers to collect fluorescence data from solid samples. Performing FS on solid samples eliminates the inherent bias associated with alkaline organic matter exactions and decreases sample preparation times. Solid phase FS on these sediments will provide a true look at the bulk organic matter characteristics and not just those alkalineextractable organic matter characteristics. Additional analysis using nuclear magnetic resonance may provide a useful comparison to the collected fluorescence data.

More data on bulk organic matter concentrations of these clastic cave sediments will also be useful moving forward.

7.2 Post-depositional diagenesis and metal cycling

Each element has a distinct chemical behavior within the sedimentary environment. This behavior is dependent on several factors. The fate and behavior of an element is determined by the oxidation state of that element, the oxidation state of the surrounding solution and solid phases, acid-base controls, and thermodynamic properties. Data collected in this study did not focus on understanding these factors. As such, predications about elemental fate and transport are gathered from bulk elemental concentrations. Many more data are needed in order to provide a more detailed understanding of geochemical behavior in this sedimentary environment.

The fate and behavior of elements are certain to differ between saturated and unsaturated sediments. Exploring the differences between these two sedimentary environments is needed to better understand post-depositional diagenesis. This could be done by focusing on solid-solution interfaces at a smaller scale. Exploring the microbial ecosystem in the cave is also necessary. Metabolic pathways can control most of the chemical fate and behavior by facilitating speciation, complexation, etc. General characterization of the metabolic pathways present in TAL through DNA sequencing or polymerase chain reaction analysis will be informative.

In order to understand the TAL system, flux calculations from input to output should also be performed. Changes in groundwater chemistry from input to output will provide insight into retention and mobility of certain chemical species in these sedimentary deposits.

A more detailed petrographic investigation coupled with X-ray diffraction analysis is recommended. It is speculated that Fe and Al oxide coatings precipitate in-situ. This may contribute to the cohesiveness of these sediments despite the low proportions of clay sized particles. The presence of Fe and Al oxides will have a strong influence on the sorption and complexation of many chemical species and organic matter.

Once more XRF data are obtained and quantified, re-evaluating the relationship between grain size and elemental concentrations is needed.

7.3 Source area investigation

Identifying the exact source area of the TAL sediments will be exceptionally informative. This question is two-fold. The groundwater flow path through the karst aquifer must first be identified. Attempts to do this, as of this project, have been unsuccessful. Once identified, surface waters entering the sink connected to TAL must be traced back to a headwater source in

the central highlands. The approach to this answer is not exactly known though isotopic investigations may prove to be useful.

8.0 Conclusion

This project concludes that there exists a complex relationship between surface derived input and post-depositional in-situ chemical transformations. Chemical differences between saturated and unsaturated sediments in TAL was observed. Interactions between solids and the sediment solution must be considered when using geochemical parameters as a tool for paleoclimate or paleo-storm reconstructions.

Fluorescent spectroscopy and PARAFAC modeling on alkaline organic matter exactions from these sediments conclude that bulk organic matter is humic acid-like (region 5; Figure 4). FI values from saturated sediments were <1.4 indicating organic matter of terrestrial origin. Proportion of PARAFAC component three increased with depth in core TAL-16-01, TAL-16-02, and TAL-16-03. Saturated sediments had lower proportions of component three. No obvious trend in PARAFAC components and FOC, organic C/N, of TN concentration.

X-ray fluorescence data concludes that Si and Al are strongly negatively correlated, suggesting that clay minerals and quartz occur together resulting in Si to dominant in those sediments, though X-ray diffraction is needed to confirm this. Titanium is correlated with several trace elements. More data are needed to better understand the mechanism controlling those trends. Correlation between certain trace elements and Fe, Al, or Mn was observed suggesting that sorption and complexation to oxides facilitates retention of these elements in the sediment. Again, more data are needed to confirm this. Preliminary comparison with bedrock and soil chemistry from the central highlands show enrichment and depletion of various elements, hinting at the mobility and retention during sediment transport.

Sediments collected from both TAL and CAM demonstrate extreme heterogeneity present spatially in the cave. Geochemical parameters collected from TAL also demonstrate the importance of microenvironments within sedimentary environments.

Information regarding the chemical makeup of clastic cave sediment is limited. A frame of reference for the elemental data reported here does not exist. The creation of a matrix-specific XRF calibration and high-resolution elemental (inorganic ad organic) data collected for this project provides such a reference. As this project illustrates, cave sedimentation is a function of local climate, geomorphology, and groundwater flow dynamics. Further work on physical and chemical characteristic of cave sediments will provide a broader frame of reference for
paleoclimate reconstructions. Exploring clastic cave deposits as paleoclimate archives, as of this project, is in its infancy.

9.0 References

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Appendix I: Naming convention for samples collected from El Tallonal Cave and Cueva Clara, PR

All samples collected from El Tallonal Cave or Cueva Clara begin with TAL or CAM, respectively, followed by a hyphen and number (or letter) referring to sample location followed by a number referring to the sample number with depth at that location followed by a hyphen and number indicating the depth within that sample at which it was subsampled. (TAL-sample location-sample number at that location-depth (cm) from the top of that sample number)

- Samples taken with depth and core samples are given a sample location number followed by a hyphen and number corresponding to the sample number taken at that location. This is followed by an additional hyphen and two digits indicating the depth at which the sub-sample was collected from that sample number.
 - Example 1: TAL-15-02-05 refers to a sample taken from location 15 at 5 cm below the top of the second sample collected at location 15.
 - Example 2: TAL-17-07-18 refers to sediment collected from sample location
 17 at 18 cm below the top of the 7th sample taken at location 17.
 - Example 3: TAL-11-01-00 refers to the sediment bank surface at location 11.
 - \circ Example 4: TAL-12-04-00 refers to the 4th sample with taken at location 12.
 - Surface sediment samples are given letters referring to sample location followed by a number referring to meters away from active cave stream
 - Example 1: TAL-A1-01-00 refers to a grab sample collected at location A at 1 meter away from the stream
 - Example 2: TAL-D3-01-00 refers to a grab sample collected at location D at 3 meters away from the stream

Appendix II: Organic matter extraction method comparison

There many ways to extract organic matter from a solid medium with results dependent on the method. Two extraction solutions were compared: (1) 0.1 M sodium hydroxide and; (2) 0.1 M sodium pyrophosphate (Na₄P₂O₇) (Chen and Hur, 2015; Greenfield et al., 2018). The extraction solutions were mixed from a 1 N NaOH stock solution and solid Na₄P₂O₇. All extractions were completed by following a modified version of the method outlined in Hur et al. (2014) and are further described in section 4.2.6. The resulting solutions were different in color (Figure II-1). The Na₄P₂O₇ extraction solutions were consistently dark, almost black, in color while the NaOH solutions tended to be a light brown but varied from dark brown to clear. Extractions using the two solutions were then analyzed for non-purgeable organic carbon (NPOC) on a Shimadzu TOC-V carbon analyzer (Figure II-2). Extraction completed with Na₄P₂O₇ extracted more total non-purgeable organic carbon than when NaOH was used. However, the Na₄P₂O₇ extraction solution yield almost no fluorescence response (Figure II-3). Extractions completed with NaOH responded with a higher intensity signal. Due to this lack of fluorescence signal when Na₄P₂O₇ was used, extractions were completed using NaOH via the method modified from Hur et al. (2014).



*Figure II-1: Photographs comparing the NaOH and Na*₄*P*₂*O*₇ *organic matter extraction solutions.*



Figure II-2: Non-purgeable organic carbon values reported for the two alkaline extraction methods. Error bars represent one standard deviation total.



Figure II-3: Emission-excitation matrices (EEM) collected from the two organic matter extraction solutions. LEFT: NaOH organic matter extraction solution; RIGHT: Na4P4O7 organic matter extraction solution.

Appendix III: Solid phase FS summary

The fluorometer is traditionally used to analyze aqueous samples; however, it was initially attempted to collect fluorescence data directly from the solid phase. Collecting fluorescence spectroscopy (FS) data from a solid sample eliminates the need to perform tedious organic matter extractions and potential bias. Solid phase fluorescence measurements were attempted on tryptophan and humic acid standards (Figure III-1). These solids were placed into the solid sample holder specifically designed for the make and model of fluorometer used. The solid sample holder was placed at an angle 30°-40° relative to the photon detector. The solid phase scans were completed (\lambda ex: 240-400 nm; \lambda em: 280-540 nm) with 1st and 2nd Rayleigh scattering removed. A consistent fluorescence signal was collected from solid phase tryptophan at λex: 270-325 nm; λem: 320-360 nm. However, solid phase humic acid did not yield a fluorescent response (Figure III-2). The concentration of photons supplied to the solid sample by the xenon bulb in the model of fluorometer used is not high enough to excite the medium. This issue does not exist when measuring FS on aqueous mediums as the excitation source can easily penetrate the solution. A study done by Muller et al. (2011) cites a similar issue occurring when attempting to detect a fluorescence signal from solid organic waste. Humic acid and other similarly dark colored sample mediums contain highly light absorbing chromophores that attenuate the fluorescence signal before it reaches the detector. The study successfully detected a solid phase fluorescence response when the xenon bulb was replaced with a laser source as it can supply higher concentrations of photons to the solid sample. Given that the Fluorolog-3 does not has a laser excitation source, organic carbon characterization was completed by extracting organic matter from the sediments into an aqueous form.

With necessary intrument upgrades solid phase FS may be possible. Collecting FS on solid phase is valuble for understanding bulk organic matter characteristics in-situ. This is not possible when extracdting organic matter into an aqueous form. Many more data and experiments are needed in order to determine the necessary intrument upgrades and method accuracy.



Figure III -1: EEM collected from solid phase tryptophan and humic acid from a Horiba® Fluoro-log 3 configured with 450-Watt light source, single grated emission monochromator, and a single grated excitation monochromator

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CAM-01-01-11	11	x	x			x		CAM-09-01-09	09	x	x			x	
CAM-01-01-12	12	x	x			x		CAM-09-01-15	15	x	x			x	
CAM-01-01-13	13	х	х			х		CAM-09-01-17	17	х	х			х	
CAM-01-01-15	15	х	х			х		CAM-09-01-20	20	х	х			х	
CAM-01-01-16	16	х	х			х		CAM-09-02-00	26	х	х			х	
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CAM-02-01-02	02	х	х			х		CAM-09-02-08	34	х	х			х	
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CAM-02-01-11	11	х	х			х		CAM-09-02-15	41	х	х			х	
CAM-02-01-13	13	х	х			х		CAM-09-02-18	44	х	х			х	
CAM-02-01-15	15	х	х			х		CAM-10-01-00	0	х	х			х	
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CAM-04-01-05	05	X	x			X		TAL-17-02-02	122	X	x			X	
CAM 04 01 12	12	X	x			X		TAL-17-02-03	120	X	x			x	
CAM 04 01 16	15	x	x			X		TAL-17-02-06	120	X	x			X	
CAM 04 01 20	20	x	x			x		TAL-17-02-11	13/	x	x			x	
CAM-05-01-00	20	×	×			v		TAL-17-02-14	134	v	×			~	
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CAM-05-01-07	07	x	x			x		TAL-17-02-25	145	x	x			x	
CAM-05-01-11	11	x	x			x		TAL-17-02-27	147	x	x			x	
CAM-05-01-14	14	x	x			x		TAL-17-03-00	135	x	x			x	
CAM-05-01-17	17	x	x			x		TAL-17-03-04	139	x	x			x	
CAM-06-01-00	00	х	х			х		TAL-17-03-08	143	х	х			х	
CAM-06-01-05	05	х	х			х		TAL-17-03-11	146	х	х			х	
CAM-06-01-07	07	х	х			х		TAL-17-03-13	148	х	х			х	
CAM-06-01-12	12	х	х			х		TAL-17-03-15	150	х	х			х	
CAM-06-01-19	19	х	х			х		TAL-17-03-18	153	х	х			х	
CAM-06-02-00	20	х	х			х		TAL-17-03-22	157	х	х			х	
CAM-06-02-08	28	х	х			х		TAL-17-04-00	147	х	х			х	
CAM-06-02-14	34	х	х			х		TAL-17-04-04	151	х	х			х	
CAM-06-02-17	37	х	х			х		TAL-17-04-07	154	х	х			х	
CAM-07-01-00	00	х	х			х		TAL-17-04-11	158	х	х			х	
CAM-07-01-04	04	х	х			х		TAL-17-04-16	163	х	х			х	
CAM-07-01-07	07	х	х			х		TAL-17-05-00	166	х	х			х	
CAM-07-01-11	11	х	х			х		TAL-17-05-03	169	х	х			х	
CAM-07-01-14	14	х	х			х		TAL-17-05-08	174	х	х			х	
CAM-07-01-17	17	х	х			х		TAL-17-05-11	177	х	х			х	
CAM-07-01-19	19	х	х			х		TAL-16-01-00	40	х	х	х		х	х
CAM-08-01-03	03	х	х			х		TAL-16-01-07	47	х	х	х		х	х

Appendix IV: All collected subsamples with tracking ID and analysis completed.

tracking ID ^a	depth below datum (cm)	particle size	TC/TOC /TN	FS	ICP ^b	XRF^{c}	XRF data quantified	tracking ID ^a	depth below datum (cm)	particle size	TC/TOC /TN	FS	ICP ^b	XRF^{c}	XRF data quantified
TAL-16-01-11	51	х	х	х		х	х	TAL-16-08-17	184	х	x			х	
TAL-16-01-14	58	x	x	x		x	x	TAL-16-08-20	187	x	x			x	
TAL-16-01-17	57	х	x	х		х	x	TAL-16-08-23	190	х	х			х	
TAL-16-01-19	59	х	x	х		х	х	TAL-16-08-26	193	х	х			х	
TAL-16-01-20	60	х	х	х		х	х	TAL-A1-01-00	177	х	х			х	
TAL-16-01-23	63	х	х	х		х	х	TAL-A2-01-00	162	х	х			х	
TAL-16-02-00	49	х	х	х		х	х	TAL-A3-01-00	159	х	х			х	
TAL-16-02-05	54	х	х	х		х	х	TAL-B1-01-00	177	х	х			х	
TAL-16-02-10	59	х	х	х		х	х	TAL-B2-01-00	151	х	х			х	
I AL-16-02-15	64	х	x	x		x	x	I AL-B3-01-00	152	x	x			х	
TAL 16-02-20	69 64	X	x	X		X	x	TAL-C1-01-00	1/2	X	X		X	X	
TAL 16 02 05	61	X	x	X		X	x	TAL-02-01-00	122	X	X		X	X	
TAL-10-03-05	00 70	x	x	x		x	x	TAL-03-01-00	44	x	x		х	x	
TAL-10-03-09	73	×	×	× ×		~	×	TAL-D1-01-00		×	× ×			×	
TAL-16-03-14	75	Ŷ	Ŷ	x		Ŷ	x	TAL-D3-01-00		Ŷ	Ŷ			x	
TAL-16-03-18	79	x	x	x		x	x	TAL-E1-01-00	178	x	x		x	x	
TAL-16-04-00	95	x	x	x		x	x	TAL-E2-01-00	160	x	x		x	x	
TAL-16-04-03	98	x	x	x		x	x	TAL-E3-01-00	158	x	x		x	x	
TAL-16-04-06	101	х	х	х		х	x	TAL-14-01-00	105	х	х			х	х
TAL-16-04-08	103	х	x	х		х	x	TAL-14-01-03	123	х	х			х	х
TAL-16-04-12	110	х	x	х		х		TAL-14-01-08	113	х	х			х	х
TAL-16-04-18	116	х	х	х		х		TAL-14-01-10	115	х	х			х	х
TAL-16-05-00	118	х	х	х		х		TAL-14-01-12	117	х	х			х	х
TAL-16-05-02	120	х	х	х		х		TAL-14-01-15	120	х	х			х	х
TAL-16-05-05	123	х	х	х		х		TAL-14-02-00	137	х	х			х	х
TAL-16-05-07	125	х	х	х		х		TAL-14-02-06	143	х	х			х	х
TAL-16-05-09	127	х	х	х		х		TAL-14-02-08	145	х	х			х	х
TAL-16-05-12	130	х	х	х		х		TAL-14-02-14	151	х	х			х	х
TAL-16-05-15	133	х	x	х		х		I AL-14-02-18	155	х	х			х	х
I AL-16-05-17	135	X	x	X		X		TAL-14-02-23	160	X	X			X	x
TAL-16-05-23	141	X	x	X		X		TAL 14 02 02	109	X	X			X	x
TAL-10-00-27	140	x	x	х		x		TAL-14-03-03	17Z	x	x			x	x
TAL-10-00-00	131 133	x	x			x		TAL-14-03-05	1/4	x	x			x	x
TAL-10-00-02	134	×	×			x y		TAL-14-03-00	181	x	x y			x Y	×
TAL-16-06-05	136	Ŷ	Ŷ			Ŷ		TAL-15-01-00	104	x	Ŷ			x	^
TAL-16-06-07	138	x	x			x		TAL-15-01-04	108	x	x			x	
TAL-16-06-09	140	x	x			x		TAL-15-01-07	111	x	x			x	
TAL-16-06-12	143	x	x			x		TAL-15-01-13	117	x	x			x	
TAL-16-06-15	146	х	x			х		TAL-15-01-17	121	х	х			х	
TAL-16-06-16	147	х	x			х		TAL-15-01-20	124	х	х			х	
TAL-16-06-17	148	х	х			х		TAL-15-01-22	130	х	х			х	
TAL-16-06-18	149	х	х			х		TAL-15-01-26	126	х	х			х	
TAL-16-06-20	151	х	х			х		TAL-15-02-00	148	х	х			х	
TAL-16-07-00	147	х	х			х		TAL-15-02-04	152	х	х			х	
TAL-16-07-02	149	х	х			х		TAL-15-02-08	156	х	х			х	
FAL-16-07-04	151	х	х			х		I AL-15-02-10	158	х	х			х	
I AL-16-07-08	155	X	x			x		I AL-15-02-13	161	x	x			х	
TAL-16-07-12	159	X	x			x		TAL 12 01 00	100	X	x			x	
TAL-10-08-00	107	x	x			x		TAL-13-01-00		x	x			x	
TAL-10-00-01	100	x	x			x		TAL-13-01-03		x	x			x	
TAL-10-00-03	173	×	×			x v		TAL-13-01-05		x	x y			x v	
TAL-16-08-08	175	×	×			×		TAL-13-01-10		×	×			×	
TAL-16-08-11	178	x	x			x		TAL-13-01-19		x	x			x	
TAL-16-08-14	181	x	x			x		TAL-13-01-24		x	x			x	

tracking ID ^a	depth below datum (cm)	particle size	TC/TOC /TN	FS	ICP ^b	XRF^{c}	XRF data quantified
TAL-13-02-04		х	х			Х	
TAL-13-02-08		х	х			х	
TAL-13-02-09		х	х			х	
TAL-13-02-12		х	х			х	
TAL-13-02-15		х	х			х	
TAL-13-02-21		х	х			х	
TAL-13-02-23		х	х			х	
TAL-13-02-26		х	х			х	
TAL-01-01-00	165	х	х		х	х	
TAL-02-01-00	120	х	х		х	х	
TAL-04-01-00	158	х	х		х	х	
TAL-05-01-00	168	х	х			х	
TAL-11-01-00	50	х	х		х	х	
TAL-11-02-00	79	х	х		х	х	
TAL-11-03-00	98	х	х			х	
TAL-11-04-00	108	х	х			х	
TAL-11-05-00	120	х	х		х	х	
TAL-11-06-00	129	х	х		х	х	
TAL-11-07-00	136	х	х		х	х	
TAL-12-01-00	158	х	х			х	
TAL-12-02-00	176	х	х		х	х	
TAL-12-03-00	182	х	х		х	х	
TAL-12-04-00	185	х	х			х	

Appendix V: Core photographs and particle size data collected from the remaining core not explicitly discussed in the text

Five cores were collected from the southern bank in the middle of the cave. Like TAL-16 sample locations, cores were collected in a diagonal transect along the bank (Figure 12). Each core is treated separately as they were not collected consecutively with depth.

Core TAL-17-01 is proportionally 50±10 percent sandy material (Figure V-1). The remaining length of the core is a consistent sedimentary texture, the proportion of sandy material decreases with depth. Some subtle redox banding is present.

Within core TAL-17-02 the proportion of sandy material transitions from 50 percent to 20 percent (Figure V-1). Subtle redox is present. Banding boundaries are consistently sub horizontal.

Redox banding is seemingly absent in core TAL-17-03 (Figure V-1). Grain size transitions from dominantly sandy (85 percent sand) to dominantly fine-grained (60 percent mud). No obvious sedimentary layering is present.

Core TAL-17-04 contains the highest proportion of sand at this sample location (Figure V-1). No obvious sedimentary or redox layering is present. Core TAL-17-05 contains an obvious redox zone in the bottom half of the core (4-12 cm) (Figure V-2). This redox zone contains small (0.1 cm) pieces of black material. The top of the core is 80 percent sandy material and deceases with depth.



Figure V-1: Cores photographs and proportion of sand from TAL-17-01, TAL-17-02, TAL-17-03, and TAL-17-04. Core are not displayed in stratigraphic order.



Figure V-2:Core photograph and proportion of sand from TAL-17-05.

Appendix VI: Summary table of all collected particle size data from a Beckman Coulter single wavelength LS13-320 particle size analyzer

tracking ID	mean (um)	sorting	skewness	% sand	% silt	% clay	% very coarse sand	% coarse sand	% medium sand	% fine sand	% very fine sand	% very coarse silt	% coarse silt	% medium silt	% fine silt	% very fine silt
TAL-02-01-00	273	3.37	-1.99	89.5	10.5	0	4.77	23.7	44.2	15.3	15.6	1.93	1.73	3.07	3.53	0.23
TAL-04-01-00	13.0	5.82	-0.50	21.2	54.7	24.1	0	0	0	4.43	16.8	20.7	13.9	9.07	9.43	1.67
TAL-A1-01-00	62.5	8.41	-0.61	50.4	58.7	10.7	1.97	16.3	21.4	6.27	4.77	13.9	11.1	6.90	6.17	0.80
TAL-A2-01-00	58.2	5.71	-1.14	59.4 46.0	64.6 51.0	8.37	U	0.33	15.3	34.8	8.97 10.0	10.7	9.53	6.17 6.97	5.20	0.70
TAL-A3-01-00	38.8 87.5	0.00	-0.84	40.Z	51.9 66.7	3.30	0	1.17	9.27	22.9	12.8 6.37	17.2	10.5	0.07	0.00	0.93
TAL-B1-01-00	18.6	3.55	-0.57	16.6	196	740	n n	2.00	20.1	3.87	12.6	20.7	21.9	18.07	14.63	0.15
TAL-B3-01-00	25.2	2.57	-0.57	15.0	17.2	2.30	0 0	0.00	0.00	1 43	13.6	30.5	26.6	15.93	943	0.00
TAL-C1-01-00	27.4	4.61	-0.74	32.6	36.8	9.20	Ő	0.33	0.57	17.3	14.4	18.0	18.2	12.77	8.97	0.23
TAL-C2-01-00	37.9	5.66	-0.84	48.4	53.7	9.93	0	0.50	4.33	30.7	12.9	11.2	12.1	9.53	8.27	0.53
TAL-C3-01-00	35.2	6.65	-0.62	43.5	50.0	12.7	0	1.70	13.8	15.9	12.2	14.9	12.0	8.07	7.57	1.13
TAL-D1-01-00	115.0	5.43	-1.04	68.2	73.2	3.50	0.97	1.87	36.4	16.8	3.07	7.40	9.47	6.53	4.77	0.20
TAL-D2-01-00	19.7	4.74	-0.68	22.6	26.9	13.1	0	0.47	0.47	8.90	12.8	21.0	19.6	13.0	10.2	0.43
TAL-D3-01-00	22.1	5.30	-0.71	30.3	35.2	14.2	0	0.00	0.33	15.5	14.4	19.1	15.7	10.5	9.2	0.90
TAL-E1-01-00	11.8	2.37	-0.60	0.10	1.88	4.50	0	0.00	0.00	0.00	0.10	12.8	25.5	32.2	24.1	0.90
TAL-E2-01-00	140.7	7.11	-0.91	68.7	75.6	3.17	5.03	29.3	26.1	6.10	2.20	4.13	6.03	8.90	8.70	0.33
TAL-E3-01-00	17.8	3.39	-1.44	4.13	0.13	9.97	0	1 22	460	0 60	4.13	37.0	20.0	13.8	9.10	0.30
TAL-11-07-00	32.2 227	3.69 4.56	-0.42	20.4 79.2	30.1 83.1	0.90	3.84	35.5	4.0Z	9.00 7.17	11.1 ∡∩q	25.0	21.0 530	13.Z 3.47	2.86	0.00
TAL-11-03-00	20.53	4 13	-1.20	20.0	23.1	12.4	0.04	00.0	20.0 0	1.00	19.0	31.1	177	10.47, 10.8	7.67	0.14
TAL-11-04-00	237	4.15	-1.12	77.5	81.3	0.00	5.33	34.2	29.9	5.82	2.23	7.22	8.27	5.02	2.02	0.00
TAL-11-05-00	14.7	4.57	-0.90	13.5	17.3	17.5	0	0	0	0.37	13.1	27.8	18.9	12.2	9.74	0.38
TAL-11-06-00	20.5	4.39	-0.76	24.0	27.8	10.5	0	0	0.05	9.68	14.2	19.4	19.2	15.4	11.47	0.18
TAL-11-07-00	86.2	5.12	-0.78	62.9	67.9	2.61	0	5.49	30.2	21.2	6.04	7.34	10.5	9.37	7.03	0.22
TAL-12-01-00	30.3	4.64	-1.12	36.8	40.7	10.6	0	0.30	0.53	13.4	22.5	25.1	13.4	7.97	5.83	0.33
TAL-12-02-00	144	4.31	-0.50	65.8	70.3	0.00	2.57	2.30	29.2	6.67	7.07	13.5	11.5	7.03	2.13	0.00
TAL-12-03-00	203	6.64	-1.91	83.8	89.3	6.53	4.50	29.5	39.0	8.97	1.83	1.90	1.50	1.57	3.27	1.47
TAL-12-04-00	168	7.13	-1.41	75.6	82.1	6.27	7.13	29.7	26.4	9.00	3.97	5.37	4.80	3.67	3.70	0.50
TAL-14-01-00	62.8	3.56	-0.41	52.9	56.5	0.50	0	1.67	11.9	25.1	14.7	16.2	14.0	10.2	6.17	0.07
TAL-14-01-03	/0./	3.55	-0.26	54.2	58.0	0.00	0	2.50	16.0	22.0	13.7	16.9	14.3	9.33	5.27	0.03
TAL-14-01-08	43.0	4.21	-1.31	49.2	52.6 73.7	0.77	U	0.00	10.93	23.3	25.1	20.5	11.0	5.83	5.40	0.27
TAL-14-01-10	09.1 88.2	3.04	-1.27	70.0 64.0	675	0.00	0	3.70	12.0	40.7 31.6	10.4	9.27	13.0	0.07 8.07	4.4/	0.23
TAL-14-01-12	48 0	3 36	-0.34	45 N	48.0	3.00	0	0.70	1.23	23.8	20.0	25.0	14.4	793	4.57	0.00
TAL-14-02-00	49.4	5.36	-1.23	58.9	63.6	9.27	Ő	047	3.93	36.1	18.4	12.1	807	5.63	5.30	0.80
TAL-14-02-06	42.0	4.13	-1.14	44.2	47.6	6.60	Ö	1.10	2.20	17.1	23.8	23.1	12.9	7.67	5.33	0.17
TAL-14-02-08	118	4.71	-0.61	66.4	71.2	0.00	2.03	12.3	32.0	16.3	3.73	5.67	11.5	10.4	6.00	0.00
TAL-14-02-14	126	4.86	-0.67	67.2	72.0	0.00	2.63	13.0	36.6	11.5	3.47	6.03	9.93	10.4	6.47	0.00
TAL-14-02-18	178	5.80	-1.24	75.0	80.3	1.30	1.50	31.7	39.3	2.93	0.27	1.80	5.10	8.40	8.37	0.00
TAL-14-02-23	90.2	5.31	-0.63	62.0	67.3	2.50	1.27	11.3	23.4	17.7	8.30	8.27	11.2	9.37	6.47	0.10
TAL-14-03-00	28.9	1.99	-0.63	9.8	11.3	0	0	0	0	0.50	9.30	43.3	26.3	14.0	6.57	0.00
TAL-14-03-06	20.1	2.16	0.32	6.1	8.67	0	0	0	1.27	0.83	4.03	21.1	33.5	26.9	12.3	0.03
TAL-14-03-09	77.3	5.68	-0.05	49.1	55.3	0	0.40	18.7	24.6	3.63	2.33	8.53	16.4	16.8	9.10	0.00
TAL-14-03-11	69.5	4.79	-0.14	51.5	56.6	0	0	7.27	24.8	15.0	4.33	8.47	15.9	15.2	8.87	0.03
TAL-14-03-13	82.U 45.2	3.20	-0.80	03.0	47.2	3.70	0.37	7.40	22.0	21.2 17.9	12.0	0.27 24.0	9.77	0.00	5.97 5.04	0.10
TAL-15-01-00	40.Z 32.0	430	-0.97	37.1	47.2	7.50	0	0	0.20	17.0	20.0 18.0	24.9	16.5	10.04	7.52	0.15
TAL-15-01-04	49.5	2.67	-0.50	46.2	48.8	0.00	0	0	0.21	20.9	25.1	20.0	15.2	9.24	5.60	0.24
TAL-15-01-13	67.8	3.82	-1.02	61.3	64.8	2.91	Ő	1 39	6 15	37.6	16.3	12.1	12.3	7.51	3 78	0.04
TAL-15-01-17	24.8	4.89	-0.80	30.0	34.3	11.6	Õ	0	0.88	16.0	13.1	19.7	19.4	11.7	7.47	0.22
TAL-15-01-20	69.7	4.87	-1.41	67.1	71.3	6.37	Õ	1.96	9.98	38.4	16.8	9.75	7.32	4.87	3.93	0.65
TAL-15-01-22	45.5	5.01	-1.19	54.9	59.3	8.61	0	0.82	2.13	31.1	20.9	13.8	10.4	6.50	5.22	0.57
TAL-15-01-26	55.7	3.73	-0.20	52.2	56.3	0.00	0	3.38	5.42	27.1	16.4	11.8	13.6	13.2	9.08	0.10
TAL-15-02-00	54.8	4.36	-0.62	48.7	53.0	3.65	0	0.92	12.0	25.6	10.3	17.7	14.8	9.01	5.96	0.16
TAL-15-02-04	174	4.57	-0.92	79.7	84.1	0.00	12.4	14.2	19.9	23.0	10.2	7.25	5.11	4.54	3.36	0.07
TAL-15-02-08	126	4.21	-0.67	72.8	77.1	0.00	4.38	1.35	22.5	24.8	10.7	7.45	6.68	8.05	5.02	0.01

tracking ID	mean (um)	sorting	skewness	% sand	% silt	% clay	% very coarse sand	% coarse sand	% medium sand	% fine sand	% very fine sand	% very coarse silt	% coarse silt	% medium silt	% fine silt	% very fine silt
TAL-15-02-10	114	4.16	-1.89	80.8	83.9	3.27	0	3.22	22.2	44.3	11.1	5.00	2.99	2.78	3.61	1.55
TAL-15-02-13	32.4	8.23	-0.29	43.7	52.1	12.31	0	3.73	21.1	15.8	3.16	6.25	6.56	13.1	17.2	0.88
TAL-15-02-18	84.0	4.71	-0.38	61.3	66.3	1.77	4.16	9.98	7.59	24.0	15.6	10.8	12.3	8.59	5.11	0.04
TAL-16-01-00	229	3.90	-1.35	82.7	86.0	0	4.60	25.3	36.8	13.1	2.87	3.30	5.40	5.37	3.20	0.03
TAL-16-01-07	39.3	4.09	0.71	28.0	33.0	0	0.00	6.27	12.1	6.30	3.23	14.9	27.6	21.3	8.23	0.00
TAL-16-01-11	191	3.91	-1.07	78.4	82.0	0	3.17	2.90	34.7	15.5	4.17	5.70	7.77	5.40	2.77	0.00
TAL-16-01-14	54.0	4.03	0.47	33.8	38.7	0	0.13	6.73	18.1	6.07	2.80	21.0	26.4	14.9	4.00	0.00
TAL-16-01-17	11.5	2.58	-0.74	0.0 70 c	1.9	7.13	0.00	0.00	0.00	0.00	0.03	12.9	26.7	31.3	21.4	0.47
TAL-16-01-18	196	4.50	-1.26	79.5	83.5	0.00	3.95	21.7	39.9	12.3	1.65	2.45	4.65	7.25	6.10 10.2	0.05
TAL-10-01-19	208	1.42	-1.07	70.4 81.6	11.4 85.3	0.00	2.07	22.1	JZ.4	14.0	2.95	3.07 2.20	3.20	6.20	6 10	0.43
TAL-16-01-23	152	4 26	-0.84	71.8	76.0	n n	1 73	164	36.6	14.0	3.07	6.17	10.6	8.03	3 37	0.20
TAL-16-02-00	64.0	5.31	-0.08	50.6	56.3	Ő	0.33	8.47	24.5	13.0	4.37	6.90	12.0	17.3	13.10	0.10
TAL-16-02-05	37.6	3.94	0.66	28.6	33.4	õ	0.13	4.63	10.7	7.8	5.33	14.9	26.4	20.2	9.87	0.03
TAL-16-02-10	37.0	8.39	-0.29	44.9	53.5	11.20	0.00	8.37	20.2	12.1	4.20	6.10	7.53	14.4	15.5	0.40
TAL-16-02-15	68.3	5.53	-0.11	48.7	54.6	1.53	2.40	9.47	22.0	11.1	3.80	9.67	17.8	14.2	8.03	0.03
TAL-16-02-20	72.8	5.40	0.17	45.4	51.4	0	3.90	15.0	15.7	7.07	3.70	11.2	21.3	15.0	7.03	0.00
TAL-16-03-00	129	5.50	-0.48	62.9	68.5	0	4.40	23.7	24.6	6.83	3.30	7.53	12.2	11.0	6.30	0.03
TAL-16-03-05	97.4	5.11	-0.53	62.1	67.3	1.27	2.23	11.0	26.1	16.5	6.30	7.63	12.2	10.5	6.23	0.03
TAL-16-03-09	160	4.85	-0.96	74.0	78.7	0.00	3.57	19.1	36.0	12.7	2.57	4.40	5.90	8.43	7.17	0.07
TAL-16-03-12	85.7	8.26	-0.99	65.9	73.8	9.23	1.93	13.5	31.6	16.2	2.67	3.73	5.43	7.97	7.63	0.17
TAL-16-03-14	88.9	4.82	-0.18	53.6	58.7	0	1.40	11.2	27.9	10.8	2.20	11.1	17.1	13.0	5.27	0.00
TAL-16-03-18	/3.5	4.48	-0.02	47.7	52.6	0	0.73	7.63	24.7	11.4	3.23	14.7	19.5	12.6	5.63	0.00
TAL-16-04-00	40.7	4.11	-0.65	40.9	44.0	0.17 14.6	0	0.95	0.40	10.0	0.20	24.0 20.1	10.0 21.9	0.90 14 g	0.40 10.37	0.07
TAL-10-04-05	21.2	334	-1.00	3.30 12.5	144	8.07	0	0	0	0.07	9.20 12.4	25.1	21.0	13.6	8.07	0.20
TAL-16-04-00	33.5	293	-1.48	30.1	31.8	3.80	0	0	0	3 03	27.0	33.5	17.1	10.0	543	0.13
TAL-16-04-12	24.6	3.22	-1.40	19.0	21.0	6.17	Ő	ŏ	ŏ	0.83	18.2	33.3	20.1	12.9	8.33	0.00
TAL-16-04-18	37.5	5.31	-0.49	35.7	40.8	7.87	0.20	6.17	6.17	10.3	12.8	22.6	16.2	10.2	7.30	0.20
TAL-16-05-00	28.7	4.29	-0.77	34.0	37.8	7.97	0	0.47	0.67	13.8	19.1	18.9	15.7	12.6	10.40	0.43
TAL-16-05-02	40.3	5.74	-0.98	53.7	59.0	10.6	0	0.43	3.27	32.6	17.4	10.4	9.00	7.77	7.67	0.90
TAL-16-05-05	28.9	4.92	-0.62	36.1	40.7	8.70	0	0.90	2.10	17.7	15.3	16.6	13.5	12.1	12.33	0.77
TAL-16-05-07	36.2	5.91	-0.90	51.4	56.9	11.6	0	0.13	2.73	31.1	17.4	10.6	8.83	8.07	8.53	1.00
TAL-16-05-09	28.1	6.20	-0.78	46.3	52.2	14.9	0	0.00	0.80	26.1	19.4	10.8	9.73	8.50	8.63	1.13
TAL-16-05-12	40.4	6.68	-0.93	56.1	62.4	13.1	0	0.27	5.67	37.8	12.4	8.03	7.83	6.53	6.80	1.57
TAL-16-05-15	39.6	4.80	-1.14	46.7	50.9	8.73	0	0.00	1.37	27.4	17.9	18.8	12.3	7.53	5.73	0.23
TAL-16-05-17	32.0	5.85	-0.88	46.4	51.9	12.7	U	0.00	1.27	28.5	16.6	13.1	11.2	8.00	7.50	1.03
TAL-10-00-20	୦୦. । ୨୮. ହ	5.42 5.40	-0.95	44.4 34 3	49.4	11.4	0	0.00	1.07	27.0	13.9	10.1	12.5 14.1	0.00	7.20	0.77
TAL-10-00-27	20.0	0.49 4 75	-0.89	18.2	39.Z	14.9	0	0.00	0.77	19.9	13.Z 8.13	18.3	14.1	9.50	0.57	1.47
TAL-16-06-02	29.1	6.01	-0.78	43.5	22.J	14.0	0	0.43	1 37	25.0	16.8	13.4	11.5	847	7 77	1.13
TAL-16-06-03	37.2	5.30	-1.05	49.3	54.1	10.6	Ő	0.23	1.23	28.2	19.7	14.5	10.9	7.67	6.57	0.50
TAL-16-06-05	33.7	4.56	-1.04	40.0	43.9	8.40	0	0.00	0.80	21.0	18.1	20.6	15.0	9.10	6.50	0.30
TAL-16-06-07	34.4	4.56	-0.95	40.3	44.3	8.47	0	0.77	1.00	19.0	19.5	19.4	14.7	9.70	7.23	0.23
TAL-16-06-09	41.4	4.91	-1.12	48.6	52.9	9.03	0	1.00	2.30	24.4	21.0	18.1	11.3	7.07	5.57	0.33
TAL-16-06-12	48.1	3.82	0.27	38.9	43.4	0.00	0	5.33	9.10	13.8	11.0	17.8	20.1	14.8	8.20	0.07
TAL-16-06-15	41.4	4.31	-0.82	49.5	53.5	5.63	0	0.00	0.70	27.8	21.0	13.4	12.5	10.6	8.10	0.27
TAL-16-06-16	41.4	4.74	-0.97	49.5	53.8	7.33	0	0.77	2.10	26.2	20.4	14.0	11.6	9.43	7.80	0.30
TAL-16-06-17	48.6	3.74	-0.54	45.7	49.4	2.73	0	2.20	3.73	20.1	19.7	20.6	14.3	9.47	6.97	0.27
TAL-16-06-20	61.5	5.08	-0.59	55.3	60.4	4.33	0	5.67	13.7	23.1	12.8	10.5	12.9	10.1	6.73	0.13
TAL-16-06-22	31.5	2.97	-0.11	26.1	29.2	0.97	U	0.00	0.17	14.7	11.2	24.2	22.9	15./	10.00	0.13
TAL-16-07-00	09.9 2022	0.20 5.02	-1.10	6∠.4 ⊿1 ว	00.1 45 0	0.5U	U	5.93 0	21.5 ∩	24.3 10.2	10.6	14.6	0.33	5.17 10.2	4.53 9.77	0.60
TAL-10-07-02	20.3 66.2	0.00 1 07	-0.92	41.3 5/ 0	40.0 50.0	202	0	430	227	19.0	22.U 0.27	14.0 11.0	10.4 1/ 2	10.2	0.77 6.20	0.00
TAL-16-07-04	86 0	4 69	-0.00	68.1	724	2.90	0 N	373	186	36.4	943	777	7 67	670	5 70	0.40
TAL-16-07-12	133	3.12	-1.46	80.1	82.6	0.00	0	1.23	29.4	43.3	6.20	4.87	5.77	5.47	3.60	0.20
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tracking ID	mean (um)	sorting	skewness	% sand	% silt	% clay	% very coarse sand	% coarse sand	% medium sand	% fine sand	% very fine sand	% very coarse silt	% coarse silt	% medium silt	% fine silt	% very fine silt
TAL-16-08-00	75.6	4.16	-0.12	53.3	57.8	0	0.40	8.27	18.2	15.9	10.53	14.50	15.50	10.83	5.90	0.03
TAL-16-08-01	142	4.15	-0.92	74.3	78.3	0	1.83	11.5	35.2	19.4	6.33	5.47	7.47	7.70	5.00	0.03
TAL-16-08-03	74.5	4.74	-0.81	62.1	66.7	3.80	0	6.20	14.5	26.4	15.0	10.9	10.4	7.40	5.30	0.20
TAL-16-08-06	24.3	4.37	-0.85	32.4	36.3	9.63	0	0.00	0.00	9.87	22.6	17.5	15.6	13.1	11.1	0.67
TAL-16-08-08	66.8	3.61	-0.52	56.3	60.0	0.47	0	0.97	10.90	33.3	11.2	12.1	14.8	10.2	5.93	0.10
TAL-16-08-11	31.8	5.47	-0.87	47.0	52.0	11.20	0	0.00	0.67	25.9	20.4	11.9	10.1	9.30	9.47	1.03
TAL-16-08-14	27.3	4.96	-0.75	37.7	42.3	10.40	0	0.57	0.63	16.3	20.2	15.8	13.2	11.2	10.9	0.83
TAL-16-08-17	31.5	5.48	-0.67	42.0	47.2	10.63	0.37	2.00	2.33	17.1	20.2	14.9	10.9	9.87	10.7	1.03
TAL-16-08-20	31.5	5.49	-0.82	46.3	51.4	11.17	0	0.60	1.13	22.9	21.7	12.7	9.53	8.90	10.1	1.30
TAL-16-08-23	54.8	2.51	-0.37	47.6	50.2	0	0	0.83	1.77	14.5	30.4	28.4	12.8	7.40	3.80	0.00
TAL-16-08-26	89.5	4.04	-0.50	63.3	67.5	0	0	5.50	23.7	25.1	9.07	7.67	12.5	11.0	5.50	0.00
TAL-17-01-00	47.7	2.57	-0.43	44.1	46.6	0	0	0	0.13	17.6	26.3	24.1	16.7	10.4	4.75	0.01
TAL-17-01-03	62.9	2.88	-0.94	59.3	61.8	0.70	0	0	1.06	31.0	27.2	17.0	10.7	7.51	4.78	0.04
TAL-17-01-07	43.8	3.26	0.10	37.6	41.3	0	0.32	2.58	3.32	14.5	16.9	20.3	17.9	14.6	9.44	0.08
TAL-17-01-12	84.4	2.86	-0.42	62.8	65.9	0	U	1.79	13.1	25.2	22.7	19.6	9.90	5.40	2.28	0.00
TAL-17-01-16	38.5	7.59	-0.73	51.7	59.1	14.69	U	4.16	10.1	25.3	12.3	10.9	7.10	7.14	8.10	0.30
TAL-17-01-24	42.7	2.13	-0.59	34.3	30.2	0.00	0	1 72	752	0.3U	29.0	30.4 10.9	17.0	9.30	3.00	0.00
TAL-17-02-00	33.0 72.1	0.00 0.04	0.39	50.7	00.4 62.4	0.00	0	0.02	7.00	20.0	7.90 20 E	10.0	10.0	24.7 7.41	2 40	0.04
TAL-17-02-02	25.2	2.84	-0.49	25.0	28.4	0.00	0	0.92	0.21	29.0	20.0	21.0	12.4 24.1	13.6	5.42	0.01
TAL-17-02-03	52.2	2.00	-0.12	20.9 45.6	20. 4 48.5	0.00	0	0.55	2.04	7. 4 2 21.0	21.8	2/3	2 4 .1 15.1	0 32	5.65	0.01
TAL-17-02-00	64.8	2.02	-0.52		-0.J 55.7	0.00	0	2 21	2.2J 8.57	21.0	17.0	18.6	153	0.02 0.42	4.65	0.04
TAL-17-02-14	40 9	3.03	-0.14	39.1	42.2	0.00	0 0	0.65	1 17	16.3	20.9	20.4	17.3	13.5	9.26	0.02
TAL-17-02-17	38.0	2 57	-0.21	32.8	35.5	0.00	Õ	0.00	0.04	12.1	20.0	26.4	21.0	12.8	6.89	0.05
TAL-17-02-21	74.4	2.91	-0.99	66.6	69.2	0.00	Ő	Õ	2.18	42.8	21.7	11.6	9.26	7.41	5.04	0.11
TAL-17-02-25	49.4	4.36	-1.12	57.9	61.7	5.45	0	0.24	0.99	33.6	23.0	11.1	9.68	8.36	7.14	0.36
TAL-17-02-27	28.6	2.31	-0.18	18.2	20.5	0.00	Ō	0.00	0.00	3.2	15.0	31.0	25.4	16.1	9.17	0.08
TAL-17-03-00	54.0	2.91	-0.31	47.3	50.4	0.00	0	0.68	3.52	22.0	21.1	20.2	15.9	10.7	5.79	0.02
TAL-17-03-04	134	3.10	-1.65	83.0	85.3	0.00	0	1.56	26.2	47.4	7.88	3.91	3.07	5.11	4.68	0.24
TAL-17-03-08	35.0	2.65	-0.13	31.2	34.1	0.00	0	0	0.13	11.1	20.0	23.8	20.6	15.6	8.68	0.04
TAL-17-03-11	87.5	3.30	-0.78	66.6	69.8	0.00	0	1.36	13.9	38.4	12.9	11.1	9.50	7.98	4.76	0.04
TAL-17-03-13	73.5	3.44	-0.82	62.6	65.8	0.56	0	1.24	7.28	38.6	15.5	12.0	9.89	8.30	6.61	0.14
TAL-17-03-15	68.4	3.35	-0.68	60.7	64.0	0.00	0	0	5.37	42.7	12.7	10.9	10.7	11.0	6.63	0.02
TAL-17-03-18	76.3	3.13	-0.82	64.7	67.6	0.00	0	0.46	5.47	43.3	15.4	11.1	10.8	8.14	5.17	0.10
TAL-17-03-22	37.7	3.51	-0.57	39.6	42.9	2.76	0	0.34	0.68	18.0	20.5	19.2	15.9	12.46	9.82	0.32
TAL-17-04-00	77.7	2.72	-0.76	63.1	65.7	0.00	0	0.34	4.79	37.0	21.0	18.1	9.91	5.78	3.05	0.01
TAL-17-04-04	122	3.26	-1.38	79.0	81.7	0.00	0	1.19	24.9	46.5	6.38	3.76	5.52	7.24	4.40	0.08
TAL-17-04-07	105	3.19	-1.10	74.3	77.1	0.00	0	1.52	16.7	46.3	9.81	6.50	8.32	7.02	3.82	0.08
TAL-17-04-11	239	2.16	-2.96	95.3	95.4	0.00	0	7.51	47.8	37.8	2.15	1.09	0.72	1.21	1.65	0.07
TAL-17-04-16	65.2	4.46	-0.91	61.4	65.6	3.25	0	1.45	9.51	38.1	12.4	8.46	9.28	9.75	7.63	0.19
TAL-17-05-00	164	2.75	-2.69	90.4 70.0	91.4	0.99	U	1.38	29.7	5Z.Z	7.12	2.52	1.57	1.80	2.58	0.11
TAL-17-00-03	112	2.97	-1.00	13.5	70.2	0.00	0	1.10	ZZ.1	37.7	12.5 14.4	10.9	1.69	0.31	2.43	0.01
TAL-17-00-06	0.00	3.40	-0.10	49.0	20.4	0.00	0	2.99	0C.1	20.0	14.4	10.0	10.1	11.1	0.70	0.00
CAM 01 01 00	42.4	2.00	-0.02	37.9	41.Z	1 43	0	0.24	1.11	19.4	17.Z	20.0	21.0 16.0	13.0	0.90	0.01
CAM-01-01-00	34.6	2.80	-0.30	32.0	34.5	1.43	0	0	0.00	6.71	25.7	22.0	17.2	12.4	8 94	0.34
CAM-01-01-02	24.8	3.74	-0.72	26.6	30.1	5.62	0 0	0 0	0.01	10.0	16.3	20.2	18.2	15.6	13.2	0.22
CAM-01-01-09	38.8	2 91	-0.55	36.3	39.0	1 18	0 0	ñ	0.20	14.2	21.8	20.2 24 0	17.6	12.5	8.35	0.00
CAM-01-01-11	37.0	3.28	-0.94	38.3	41 ∩	3.68	Ň	0.62	0.02	12.6	25.6	23.1	15.9	11 1	7 72	0.11
CAM-01-01-12	50.9	3.07	-0.70	46.0	48.9	1.55	õ	0	2.36	20.0	23.1	21.4	14.6	9.67	6.65	0.13
CAM-01-01-15	16.1	4.69	-0.64	23.1	27.4	14.5	Ō	1.56	0.00	3.16	20.0	17.7	15.0	14.5	13.9	1.34
CAM-01-01-19	25.8	4.64	-0.73	32.4	36.6	9.48	Ō	0.00	0.49	10.12	20.2	19.7	14.7	12.8	10.5	0.42
CAM-02-01-02	69.5	3.16	-0.77	62.3	37.1	0.61	0	2.09	4.53	29.4	26.3	15.3	9.34	7.16	5.11	0.14
CAM-02-01-08	41.1	2.43	-0.41	35.5	64.5	0.00	0	0	0.04	9.83	25.7	30.0	17.8	11.0	5.55	0.05

tracking ID	mean (um)	sorting	skewness	% sand	% silt	% clay	% very coarse sand	% coarse sand	% medium sand	% fine sand	% very fine sand	% very coarse silt	% coarse silt	% medium silt	% fine silt	% very fine silt
CAM-02-01-13	33.0	3.95	-1.14	40.3	53.1	6.60	0	0	0.04	11.7	28.6	22.7	11.7	9.82	8.44	0.43
CAM-02-01-17	55.8	3.22	-0.51	50.0	49.0	1.01	0	1.22	4.29	22.3	22.3	20.0	13.9	9.09	5.88	0.09
CAM-03-01-00	52.7	2.67	-0.55	48.4	51.6	0.00	0	0.17	0.85	20.4	27.0	24.3	13.4	8.35	5.49	0.11
CAM-03-01-04	43.5	2.85	-0.73	41.5	57.3	1.19	0	0	0.24	15.7	25.6	23.8	15.6	10.7	7.07	0.10
CAM-03-01-12	43.7	3.01	-0.59	41.9	56.8	1.26	0	0.22	0.61	17.9	23.2	21.8	16.4	11.3	7.33	0.11
CAM-03-01-17	35.9	2.63	-0.07	31.1	68.9	0	0	0	0.21	11.5	19.4	24.3	22.0	15.3	7.34	0.05
CAM-05-01-00	149	5.84	-1.77	80.8	85.7	4.48	0	3.23	33.5	21.3	8.07	4.13	2.32	2.77	4.59	0.88
CAM-05-01-05	120	3.37	-0.99	73.7	76.9	0	0	1.62	31.1	27.2	12.2	10.3	6.40	5.43	4.04	0.14
CAM-05-01-07	64.8	3.66	-0.57	57.0	60.7	0.65	0	4.51	11.1	25.0	19.2	15.4	10.0	8.91	7.68	0.24
CAM-05-01-09	80.5	4.77	-0.78	62.6	67.2	2.14	0	3.34	24.8	21.9	11.3	10.1	8.26	8.32	8.20	0.34
CAM-05-01-11	84.9	3.27	-0.51	63.7	67.1	0	0	2.46	15.9	22.4	22.0	16.9	8.47	6.31	4.52	0.08
CAM-05-01-14	78.0	4.49	-0.92	64.2	68.4	2.46	0	8.07	20.7	27.1	13.9	10.5	7.83	7.21	7.36	0.46
CAM-05-01-17	160	3.27	-1.41	82.9	85.6	0	0	0	37.3	27.7	9.82	5.00	3.76	4.49	3.63	0.19
CAM-06-01-00	90.7	3.34	-0.80	68.7	31.3	0	0	2.16	16.1	33.4	17.1	11.0	8.0	6.9	5.27	0.15
CAM-06-01-05	80.7	3.54	-0.57	63.2	36.8	0	0	2.66	15.5	28.6	16.4	12.3	10.0	8.5	5.95	0.10
CAM-06-01-07	69.4	3.45	-0.39	58.1	41.9	0	0	1.98	11.5	27.2	17.5	13.6	12.4	10.1	5.83	0.04
CAM-06-01-12	44.7	3.35	0.29	36.5	63.5	0	0	2.86	6.30	13.7	13.6	20.7	20.9	14.9	6.83	0.02
CAM-06-01-19	28.4	2.21	-0.16	17.9	82.1	0	0	0	0	1.24	16.7	29.8	27.9	17.4	6.97	0.02
CAM-06-02-00	81.5	3.64	-0.58	63.7	36.3	0	0	3.12	16.5	27.7	16.4	11.9	9.26	8.36	6.64	0.14
CAM-06-02-08	30.5	2.34	0.03	20.3	79.7	0	0	0.81	0.08	2.26	17.2	30.4	26.2	16.2	6.89	0.02
CAM-06-02-14	31.3	2.30	-0.24	21.2	78.8	0	0	0	0	3.83	17.4	32.5	24.5	14.5	7.21	0.07
CAM-06-02-17	37.2	2.48	-0.21	29.8	70.2	0	0	0	0.09	9.90	19.8	30.4	21.0	12.3	6.42	0.08
CAM-07-01-07	29.2	2.24	-0.34	18.2	20.3	0	0	0	0	1.61	16.6	33.3	24.2	15.7	8.49	0.09
CAM-07-01-11	32.6	2.34	-0.36	24.9	27.1	0	0	0.68	0	3.51	21.3	31.1	22.2	14.0	7.70	0.10
CAM-07-01-14	36.0	2.57	-0.03	29.2	32.0	0.45	0	0	0.36	8.25	19.9	28.6	21.3	13.8	7.10	0.06
CAM-07-01-17	20.0	2.20	-0.00	12.0	14.0	0.45	0	16.2	0	0.58	12.2	33.0 25.6	20.0	17.0	9.90	0.10
CAM 02 01 02	20.0	2.10	-0.40	16.0	17.0	1 16	0	10.3	0.01	2.27	10.2	30.0	24.0 25.2	10.3	0.20 11.2	0.10
CAM-08-01-03	20.2	2.00	-0.33	10.4	02.0 81.0	0	0	0	0.01	0.03	10.1	27.9	20.0	17.0	8.20	0.21
CAM-08-01-11	23.0 24 Q	2.22	-0.41	13.0	85.6	1 26	0	0	0	1.02	12.1	32.1	20.0	16.7	10.20	0.00
CAM-08-01-15	263	2.40	-0.30	13.1	87.0	0	0	0	0	1.02	11.8	32.1	20.5	17.4	9.25	0.23
CAM-08-01-19	28.5	2.10	-0.24	14.6	854	ñ	ñ	ñ	ñ	0.70	13.9	35.7	26.9	15.7	7 14	0.07
CAM-09-01-05	66.7	4.01	0.34	42.5	57.5	ŏ	2.34	8.19	12.9	8.52	10.6	23.6	19.3	10.5	4.06	0.01
CAM-09-01-09	80.2	4.67	0.04	51.3	48.7	Ő	3.31	12.1	14.8	11.4	9.61	15.5	16.5	10.6	6.02	0.08
CAM-09-01-15	14.3	3.75	-0.84	9.21	78.7	12	0	0	0	1.77	7.44	21.2	24.4	18.2	14.1	0.71
CAM-09-01-17	46.4	2.82	0.60	32.9	67.1	0	Ō	4.29	1.93	8.11	18.2	31.4	21.5	11.0	3.16	0.00
CAM-09-01-20	30.5	4.67	0.12	30.1	66.2	4	0	5.24	5.92	8.32	10.6	14.7	18.1	19.2	13.8	0.29
CAM-09-02-00	56.7	3.27	0.17	43.3	56.7	0	0	3.26	9.86	14.26	15.9	23.4	18.7	10.5	4.11	0.01
CAM-09-02-05	64.7	4.20	0.45	40.5	59.5	0	2.60	11.0	9.49	7.06	10.3	23.6	20.5	11.2	4.19	0.01
CAM-09-02-08	41.7	2.70	0.12	33.2	66.8	0	0	1.09	1.64	12.6	17.8	26.6	22.7	12.7	4.80	0.01
CAM-09-02-11	59.2	3.41	0.15	45.5	54.5	0	0.43	4.98	7.82	14.7	17.5	23.0	16.7	9.7	5.09	0.05
CAM-09-02-15	45.4	3.76	0.71	29.8	70.2	0	0.67	8.89	5.38	4.81	10.0	25.4	24.1	14.2	6.56	0.04
CAM-09-02-18	45.1	3.08	0.64	30.5	69.5	0	0.55	4.19	4.27	8.46	13.0	29.3	23.9	12.1	4.15	0.00
CAM-10-01-00	53.5	4.24	0.38	38.0	43.0	0	0	6.51	15.3	7.03	7.55	16.9	22.4	16.3	6.34	0.01
CAM-10-01-10	60.6	3.92	0.32	41.0	45.6	0	0	0.95	15.0	10.0	8.56	20.6	21.8	12.4	4.20	0.00
CAM-10-01-15	52.7	3.17	-0.10	45.5	49.0	0	0	1.11	6.12	21.6	16.8	19.4	18.2	11.0	5.85	0.05
CAM-10-01-21	50.7	3.33	-0.11	46.1	49.8	0	0	0.24	5.92	22.0	17.1	16.8	17.0	12.8	7.22	0.06
CAM-10-01-24	49.6	2.71	-0.28	44.0	46.9	0	0	1.34	1.08	20.4	22.3	23.7	17.5	10.0	4.72	0.04

Appendix VII: Summary table of all collected TC, TOC, and TN data

racking ID	TC (mol/kg)	TIC (mol/kg)	TOC (mol/kg)	foc (%)	organic C/N	racking ID	TC (mol/kg)	TIC (mol/kg)	TOC (mol/kg)	foc (%)	organic C/N
TAL -16-01-00	0.08	0.02	0.06	0.73	8 51	TAL-16-07-12	0.05	0.00	0.05	0.55	9 16
TAL-16-01-07	0.03	0.01	0.03	0.31	4 52	TAL-16-08-00	0.09	0.02	0.08	0.00	8 84
TAL-16-01-11	0.03	0.00	0.03	0.36	5.25	TAL-16-08-01	0.07	0.02	0.06	0.69	10.1
TAL-16-01-14	0.02	0.00	0.02	0.28	4.66	TAL-16-08-03	0.09	0.01	0.07	0.9	8.07
TAL-16-01-17	0.03	0.00	0.03	0.33	4.81	TAL-16-08-06	0.17	0.01	0.16	1.95	12.0
TAL-16-01-19	0.03	0.00	0.03	0.34	5.66	TAL-16-08-08	0.05	0.01	0.05	0.55	8.02
TAL-16-01-23	0.07	0.02	0.05	0.64	6.22	TAL-16-08-11	0.15	0.01	0.13	1.59	10.9
TAL-16-02-00	0.04	0.00	0.04	0.48	6.22	TAL-16-08-14	0.21	0.03	0.18	2.17	11.5
TAL-16-02-05	0.03	0.00	0.03	0.32	4.66	TAL-16-08-17	0.22	0.02	0.20	2.35	11.4
TAL-16-02-10	0.03	0.00	0.02	0.29	4.83	TAL-16-08-20	0.23	0.03	0.20	2.36	11.5
TAL-16-02-15	0.03	0.00	0.02	0.29	4.23	TAL-16-08-23	0.03	0.00	0.03	0.34	4.96
TAL-16-02-20	0.03	0.01	0.02	0.30	5.00	TAL-16-08-26	0.04	0.01	0.04	0.47	6.85
TAL-16-03-00	0.04	0.01	0.03	0.39	9.10	TAL-A1-01-00	0.11	0.01	0.10	1.24	10.3
TAL-16-03-05	0.03	0.00	0.03	0.33	4.81	TAL-A2-01-00	0.11	0.03	0.08	0.95	7.91
TAL-16-03-09	0.02	0.00	0.02	0.26	4.33	TAL-A3-01-00	0.08	0.02	0.06	0.76	8.86
TAL-16-03-12	0.02	0.00	0.02	0.20	3.89	TAL-B1-01-00	0.12	0.00	0.11	1.37	10.7
TAL-16-03-14	0.01	0.00	0.01	0.13	3.79	TAL-B2-01-00	0.10	0.00	0.10	1.15	12.2
TAL-16-03-18	0.02	0.00	0.02	0.21	4.08	TAL-B3-01-00	0.07	0.01	0.06	0.68	8.81
TAL-16-04-00	0.01	0.00	0.01	0.16	3.73	TAL-C1-01-00	0.15	0.02	0.13	1.58	9.70
TAL-16-04-03	0.02	0.00	0.02	0.29	4.23	TAL-C2-01-00	0.12	0.01	0.10	1.23	9.56
TAL-16-04-06	0.02	0.00	0.02	0.30	4.37	TAL-C3-01-00	0.12	0.02	0.10	1.22	8.89
TAL-16-04-08	0.03	0.00	0.03	0.32	4.66	TAL-D1-01-00	0.06	0.01	0.05	0.59	7.64
TAL-16-04-12	0.03	0.00	0.03	0.33	4.81	TAL-D2-01-00	0.10	0.03	0.08	0.92	7.66
TAL-16-04-18	0.02	0.00	0.02	0.21	3.50	TAL-D3-01-00	0.07	0.02	0.06	0.68	7.93
TAL-16-05-00	0.13	0.03	0.10	1.18	9.17	TAL-E1-01-00	0.07	0.01	0.06	0.72	7.63
TAL-16-05-02	0.13	0.03	0.10	1.15	9.58	TAL-E2-01-00	0.08	0.00	0.08	0.92	9.75
TAL-16-05-05	0.16	0.02	0.14	1.71	11.1	TAL-E3-01-00	0.10	0.00	0.10	1.16	10.4
TAL-16-05-07	0.14	0.04	0.10	1.26	9.80	TAL-01-01-00	0.11	0.02	0.09	1.12	8.71
TAL-16-05-09	0.16	0.04	0.12	1.44	10.5	TAL-02-01-00	0.05	0.01	0.03	0.42	5.44
TAL-16-05-12	0.08	0.03	0.05	0.62	7.23	TAL-04-01-00	0.06	0.02	0.04	0.48	5.60
TAL-16-05-15	0.08	0.02	0.05	0.63	8.16	TAL-05-01-00	0.02	0.00	0.02	0.21	4.08
TAL-16-05-17	0.07	0.02	0.05	0.55	6.41						
TAL-16-05-23	0.02	0.00	0.02	0.20	3.89						
TAL-16-05-27	0.07	0.02	0.05	0.63	6.12						
TAL-16-06-00	0.11	0.02	0.08	0.99	8.25						
TAL-16-06-02	0.08	0.03	0.05	0.64	7.46						
TAL-16-06-03	0.12	0.03	0.08	1.00	8.33						
TAL-16-06-05	0.10	0.02	0.08	0.94	7.83						
TAL-16-06-07	0.13	0.03	0.10	1.18	9.17						
TAL-16-06-09	0.09	0.02	0.06	0.76	8.06						
TAL-16-06-12	0.11	0.03	0.09	1.04	8.66						
TAL-16-06-15	U.16	0.02	U.14	1.67	10.25						
TAL-16-06-16	0.12	0.01	0.10	1.22	8.89						
TAL-16-06-17	0.12	0.01	U.11	1.34	9.19						
TAL-16-06-18	0.06	0.01	0.05	0.57	0.65						
TAL-16-06-20	0.04	0.01	0.04	0.43	0.27						
TAL-16-07-00	0.08	0.02	0.06	0.69	8.05						
TAL-10-07-02	0.11	0.03	0.08	0.93	8.34 e oo						
TAL-16-07-04	0.04	0.01	0.03	0.36	6.00						
TAL-16-07-08	0.10	0.01	0.09	1.12	9.33						

Appendix VIII: Summary table of all quantified XRF data

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mqq) oM
16-01	0	40	0.03	0	22.7	0.09	0.20	0.80	513	0.47	245	219	1.10	6.86	15.4	32.0	58.5	41.3	29.4	8.38	13.7	12.6	20.7	379	3.50	11.9
16-01	0.2	40.2	0.05	0	19.0	0.10	0.14	0.69	287	0.45	232	223	0.99	6.80	15.4	30.9	57.6	38.1	28.7	7.93	10.4	7.70	20.2	271	2.52	6.13
16-01	0.4	40.4	0.01	0	20.6	0.09	0.14	0.65	0.0	0.44	221	226	1.17	7.34	15.4	31.6	61.6	44.0	29.7	8.09	11.0	6.64	20.7	219	2.93	4.40
16-01	0.6	40.6	0.03	0	19.6	0.11	0.14	0.65	383	0.47	242	223	1.27	7.89	16.7	30.9	60.2	43.8	31.3	7.75	10.4	6.97	19.3	229	1.65	4.97
16-01	0.8	40.8	0.04	0	18.6	0.11	0.12	0.67	76.1	0.47	242	224	1.32	7.78	15.7	31.0	61.4	44.9	30.6	8.05	11.4	8.00	21.0	189	2.33	0.96
16-01	1	41	0.05	0	18.8	0.11	0.13	0.62	343	0.45	232	211	1.30	7.69	16.5	32.3	58.4	45.9	30.8	8.12	9.60	6.34	21.4	202	2.89	3.17
16-01	1.2	41.2	0	0	20.8	0.10	0.13	0.59	476	0.44	223	225	1.50	7.79	15.0	34.6	59.0	42.3	3E+12	7.91	9.74	7.71	18.8	226	2.53	3.19
16-01	1.4	41.4	0.01	0	21.7	0.09	0.15	0.57	461	0.44	216	232	1.43	7.63	14.8	31.3	52.3	41.0	28.4	8.13	9.42	8.64	18.6	250	1.66	3.76
16-01	1.6	41.6	0	0	21.1	0.10	0.14	0.56	0.0	0.43	214	229	1.13	7.37	16.3	30.6	58.1	42.3	29.9	8.32	9.62	6.82	18.5	238	1.91	6.22
16-01	1.8	41.8	0	0	20.6	0.11	0.12	0.54	562	0.43	220	211	1.28	7.45	14.1	29.7	53.1	37.7	27.9	8.21	9.29	5.53	17.5	244	1.93	5.09
16-01	2	42	0.04	0	18.7	0.11	0.11	0.56	655	0.41	206	201	0.94	6.77	14.5	28.5	51.4	39.4	26.8	7.87	8.79	2.74	16.7	225	1.92	3.71
16-01	2.2	42.2	0.03	0	19.5	0.11	0.12	0.56	449	0.40	205	203	1.05	7.03	15.8	30.4	49.0	36.9	25.3	8.46	8.25	5.73	17.3	212	1.96	4.62
16-01	2.4	42.4	0.02	0	19.5	0.09	0.12	0.57	359	0.40	203	203	0.78	6.19	13.4	27.5	47.1	36.6	24.5	8.66	9.39	7.39	16.5	219	2.13	4.27
16-01	2.6	42.6	0.03	0	19.6	0.09	0.12	0.56	456	0.40	205	206	0.93	6.96	15.1	27.3	62.2	38.3	27.5	7.75	9.44	5.19	17.9	299	1.76	7.02
16-01	2.8	42.8	0.04	0	18.3	0.10	0.11	0.58	531	0.41	209	211	0.62	6.90	18.3	39.5	624	38.7	27.7	8.19	11.1	6.08	18.9	417	2.61	10.1
16-01	3	43	0.05	0	18.0	0.10	0.11	0.61	512	0.42	216	215	0.16	6.19	18.7	46.5	694	43.7	29.6	8.33	11.4	8.72	19.7	400	2.48	10.7
16-01	3.2	43.2	0.04	0	17.8	0.10	0.11	0.64	294	0.44	222	213	1.09	7.92	18.3	42.0	645	44.9	28.3	8.56	11.4	9.01	18.3	319	2.01	7.09
16-01	3.4	43.4	0.05	0	17.5	0.11	0.10	0.65	292	0.43	217	211	0.75	6.85	15.8	39.3	182	43.3	29.8	8.08	9.04	7.65	19.2	255	3.43	6.15
16-01	3.6	43.6	0.06	0	17.8	0.11	0.11	0.67	416	0.44	225	218	0.75	6.73	16.2	31.0	58.3	46.3	31.8	8.25	10.3	8.96	20.6	302	2.52	7.69
16-01	3.8	43.8	0.05	0	17.6	0.11	0.10	0.73	321	0.44	224	220	0.96	7.09	16.3	36.4	61.4	44.7	29.8	8.22	11.7	9.51	20.4	319	2.09	8.81
16-01	4	44	0.07	0	17.0	0.11	0.10	0.69	0.0	0.45	229	213	1.47	7.84	16.2	36.5	58.8	49.5	33.7	7.97	11.9	9.23	23.0	258	3.67	7.20
16-01	4.2	44.2	0.06	0	16.9	0.11	0.10	0.66	495	0.44	229	206	1.07	7.00	17.4	36.4	54.6	47.5	34.5	8.66	11.4	8.38	19.7	194	3.65	1.17
16-01	4.4	44.4	0.04	0	17.4	0.10	0.10	0.65	394	0.44	225	205	1.48	7.64	16.7	33.8	58.6	47.9	31.3	8.32	10.8	11.4	20.9	201	2.92	2.89
16-01	4.6	44.6	0.04	0	18.3	0.10	0.10	0.64	563	0.43	222	205	1.32	7.40	16.7	30.5	58.0	42.6	29.6	7.92	11.8	9.50	20.8	225	3.13	3.19
16-01	4.8	44.8	0.04	0	18.2	0.11	0.11	0.65	0.0	0.45	230	210	1.40	7.84	17.3	34.2	53.2	43.7	28.2	8.57	9.79	7.91	21.2	282	2.74	5.25
16-01	5	45	0.04	0	18.2	0.11	0.11	0.66	148	0.44	226	210	1.23	7.19	15.7	35.2	89.4	49.4	32.2	8.49	10.5	9.59	21.0	344	2.18	1.13
16-01	5.Z	45.Z	0.05	0	17.3	0.12	0.09	0.60	218	0.46	235	210	0.98	7.09	18.4	41.2	430	50.Z	31.0	0.71 0.54	11.3	0 4 4	22.7	325	3.24	8.50
10-01	5.4	45.4	0.06	0	10.7	0.12	0.09	0.64	217	0.47	242	203	1.20	7.00	10.5	30.2	393	47.3	31.0	0.54	10.9	0.44	21.0	270	2.01	5.92
16-01	5.0 5.0	45.0	0.00	0	10.7	0.13	0.09	0.65	154	0.46	237	197	1.87	8.23 7.00	10.0	32.4	74.4	45.4	29.0	8.70 0.22	11.2	0.03	22.1	220	2.10	2.10
10-01	0.0	40.0	0.07	0	15.0	0.13	0.09	0.07	140	0.47	241	193	1.41	7.09	10.1	33.4 22.5	54.9	40.0	32.1	0.32	10.9	0.00	21.0	190	3.24	1.39
16-01	0	40	0.08	0 01	15.7	0.12	0.08	0.68	214	0.45	233	190	1.49	7.40	10.2	32.5	59.Z	47.9	30.9	0.91	11.1	9.99	20.9	220	3.40	2.35
10-01	0.2	40.2	0.09	0.01	15.2	0.12	0.10	0.71	309	0.45	232	197	1.37	7.27	10.9	34.0	02.4	50.4	32.0	0.30	12.3	7.55	20.0	220	2.07	5.10
16-01	0.4 6.6	40.4	0.10	0	14.8	0.13	0.08	0.05	305	0.44	225	195	1.00	7.19	10.0	33.0	62.7 50.5	47.8	32.1	0.00	13.4	9.19	20.7	224	2.81	5.45
10-01	0.0	40.0	0.11	0.30	14.3	0.13	0.09	0.63	0.0	0.43	222	193	1.85	7.84	14.9	35.7	59.5	47.0	31.8	8.60	11.1	8.98	22.2	201	2.41	5.87
16.01	0.0 7	40.8 47	0.11	0.22	14.5	0.13	0.08	0.03	419	0.45	232	190	1.44	1.10	10.0	35.∠ 27.4	50.1	5U.8	32.9	0.00 0 00	12.5	0.UD	21.9	200	2.98	0.3Z
10-01	70	47	0.11	0.17	14.2	0.13	0.07	0.03	525	0.45	234	184	1.47	0.93	15.8	37.1	57.4 60.7	47.0	31.0	0.00	11.0	9.40	21.5	2/1	2.19	10.0
10-01	1.2	47.2	0.10	0	15.2	0.13	0.09	0.64	44.9	0.46	230 207	181	1.65	7.41	17.5	3∠.ŏ	юU./	49.9	33.2	ö.02	12.5	0.31	22.2	401	3.06	10.1
16-01	1.4	47.4	0.09	0.08	15.4	0.13	0.10	0.67	611	0.44	227	1/1	1.88	1.66	17.7	37.5	53.3	47.5	31.4	8.63	13.3	6.35	23.8	385	2.19	9.92

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mqq) oM
16-01	7.6	47.6	0.07	0	16.2	0.12	0.11	0.66	349	0.45	231	179	1.57	7.21	18.0	34.4	58.1	48.0	30.4	8.45	12.5	8.45	24.4	391	2.02	9.45
16-01	7.8	47.8	0.07	0	16.2	0.13	0.11	0.64	317	0.45	233	187	1.84	7.92	16.9	32.0	58.9	48.0	29.9	8.57	13.0	7.66	23.1	322	3.17	8.66
16-01	8	48	0.07	0	17.1	0.12	0.12	0.61	173	0.45	229	195	1.65	8.13	19.2	34.2	60.0	48.0	30.6	8.17	12.0	11.2	23.8	201	3.71	1.68
16-01	8.2	48.2	0.05	0	18.4	0.11	0.14	0.65	0.0	0.47	241	212	1.32	7.68	17.9	36.4	60.3	53.1	33.6	8.29	13.1	10.9	23.4	219	3.10	3.59
16-01	8.4	48.4	0.05	0	17.8	0.12	0.12	0.67	514	0.46	237	220	1.19	7.79	16.8	35.8	59.5	49.8	32.9	8.45	11.2	10.1	23.4	245	2.76	5.98
16-01	8.6	48.6	0.11	0.10	14.9	0.13	0.08	0.64	296	0.43	219	224	0.26	6.45	15.9	33.5	50.1	40.9	26.2	7.67	9.16	7.98	20.6	257	1.65	3.53
16-01	8.8	48.8	0.21	1.12	10.7	0.15	0.04	0.59	366	0.40	206	219	0.61	7.23	12.3	30.8	46.5	37.1	28.2	7.73	8.33	6.84	20.2	272	2.62	6.87
16-01	9	49	0.33	2.36	5.97	0.16	0.00	0.60	680	0.40	205	217	0.00	7.14	15.2	40.8	50.6	39.0	28.7	7.77	7.47	5.18	19.3	251	2.52	9.70
16-01	9.2	49.2	0.46	3.43	2.23	0.15	0.00	0.57	0.0	0.38	195	209	0.00	5.75	14.3	67.7	59.7	37.6	28.2	10.16	9.20	4.11	21.9	254	3.53	10.37
16-01	9.4	49.4	0.46	3.48	1.87	0.16	0.00	0.68	908	0.43	222	229	0.00	6.29	19.2	82.8	63.3	37.2	25.1	8.29	6.92	1.59	13.9	220	2.95	5.56
16-01	9.6	49.6	0.19	1.16	10.7	0.14	0.01	0.64	0.0	0.45	228	239	0.00	6.98	16.6	54.1	64.5	39.8	29.2	6.83	4.06	1.89	18.4	166	2.87	0.00
16-01	9.8	49.8	0.04	0	17.3	0.11	0.09	0.62	637	0.42	215	226	0.84	6.97	12.1	39.4	56.6	43.0	27.8	7.40	6.68	4.93	20.4	201	1.11	0.68
16-01	10	50	0.02	0	20.1	0.10	0.13	0.62	503	0.42	214	216	0.92	6.96	13.7	32.8	50.3	44.0	28.1	7.86	10.1	7.27	21.6	201	2.21	0.00
16-01	10.2	50.2	0	0	19.9	0.10	0.13	0.62	392	0.42	215	220	1.24	7.45	13.4	25.1	50.2	47.6	28.9	7.87	9.11	6.56	21.7	218	1.75	2.79
16-01	10.4	50.4	0.02	0	19.8	0.11	0.13	0.60	333	0.43	223	221	1.26	7.91	16.7	31.3	54.3	47.5	30.9	7.81	9.04	8.33	21.2	218	1.93	1.02
16-01	10.6	50.6	0.02	0	19.3	0.12	0.13	0.64	177	0.45	232	221	0.75	6.92	15.8	30.4	56.5	45.7	30.3	7.85	10.5	11.0	23.9	215	2.15	2.86
16-01	10.8	50.8	0.05	0	17.9	0.12	0.12	0.62	245	0.43	224	222	1.03	7.63	16.8	28.5	52.8	46.7	31.3	8.27	9.56	8.26	23.0	194	3.34	0.00
16-01	11	51	0.05	0	17.3	0.12	0.11	0.62	661	0.43	218	214	0.67	7.01	17.3	28.3	55.1	46.0	30.2	8.06	10.7	7.03	20.6	240	2.48	4.06
16-01	11.2	51.2	0.05	0	17.1	0.12	0.10	0.62	413	0.43	218	212	0.68	6.96	16.4	30.4	50.9	45.4	30.1	7.90	8.81	5.43	22.4	300	2.15	8.39
16-01	11.4	51.4	0.06	0	17.1	0.12	0.11	0.60	588	0.43	221	216	0.84	7.56	17.8	33.9	50.7	42.9	27.3	8.10	9.86	7.28	19.9	338	1.61	8.67
16-01	11.6	51.6	0.06	0	17.7	0.12	0.12	0.63	0	0.44	226	214	0.62	6.94	16.9	28.4	55.1	43.7	28.7	8.28	9.98	6.62	19.2	310	2.04	3.47
16-01	11.8	51.8	0.05	0	17.9	0.12	0.12	0.60	180	0.44	221	210	0.18	6.37	19.0	29.3	56.1	40.3	28.6	8.01	9.46	6.76	19.4	291	1.93	6.74
16-01	12	52	0.05	0	18.4	0.11	0.12	0.58	486	0.42	209	203	0.38	6.16	14.6	28.3	49.2	37.0	27.7	7.66	6.87	2.60	19.0	317	1.54	8.46
16-01	12.2	52.2	0.0	0	0	0	0	0.56	229	0.40	198	197	0.35	5.58	11.2	24.6	46.3	30.9	22.4	8.32	5.55	3.76	17.8	312	1.61	6.96
16-01	12.4	52.4	0.11	0	16.1	0.13	0.09	0.60	0	0.41	203	201	0.20	5.73	14.9	27.4	47.3	31.1	22.3	8.03	7.01	0.83	17.0	237	2.00	2.97
16-01	12.6	52.6	0.16	0.54	13.8	0.13	0.06	0.65	0	0.42	209	200	0.30	5.91	14.0	34.2	51.1	31.3	25.6	8.55	7.45	1.27	17.4	210	2.25	4.59
16-01	12.8	52.8	0.21	1.12	11.6	0.15	0.04	0.59	813	0.40	204	203	0.53	6.19	13.0	33.3	47.2	29.2	21.4	8.03	6.48	1.52	16.2	206	2.03	4.23
16-01	13	53	0.22	1.18	11.1	0.15	0.03	0.55	236	0.42	210	207	0.03	5.67	15.3	38.1	51.9	30.9	21.6	8.22	7.82	1.04	18.4	227	3.26	6.64
16-01	13.2	53.2	0.16	0.62	13.5	0.14	0.06	0.57	126	0.39	198	206	0.28	5.61	12.1	34.8	50.7	31.1	24.4	8.47	6.17	0.25	17.0	259	2.82	6.64
16-01	13.4	53.4	0.09	0	17.8	0.12	0.10	0.58	611	0.38	193	216	0.00	5.22	13.9	33.9	46.3	33.0	25.0	7.82	5.46	0.00	18.9	337	1.94	8.89
16-01	13.6	53.6	0.05	0	20.1	0.11	0.13	0.59	507	0.40	204	194	0.00	6.83	23.7	41.3	56.4	33.9	26.0	7.73	6.35	2.41	17.2	309	2.31	7.24
16-01	13.8	53.8	0.02	0	22.0	0.11	0.16	0.56	506	0.42	215	198	0.00	6.48	23.2	38.1	56.3	32.6	23.5	7.81	6.41	3.76	18.3	244	1.83	5.30
16-01	14	54	0.05	0	22.3	0.10	0.18	0.59	645	0.42	208	228	0.00	6.94	22.3	38.1	56.1	32.3	25.4	7.43	4.87	2.88	15.5	252	1.04	3.76
16-01	14.2	54.2	0	0	23.3	0.10	0.18	0.57	363	0.38	194	210	0.00	5.85	14.5	23.1	42.9	29.8	23.6	7.31	3.75	0.08	13.6	270	1.43	3.44
16-01	14.4	54.4	0	0	22.8	0.10	0.19	0.55	261	0.37	190	196	0.89	6.24	13.1	23.4	42.0	30.4	24.6	7.68	5.14	2.70	15.1	256	1.70	2.57
16-01	14.6	54.6	0	0	21.8	0.10	0.15	0.54	96.2	0.38	193	191	0.99	6.55	14.4	24.8	45.0	33.3	25.9	7.66	7.20	2.21	15.1	225	1.35	3.04
16-01	14.8	54.8	0.04	0	18.0	0.12	0.10	0.54	601	0.39	194	191	1.34	7.21	13.9	24.9	46.1	36.6	26.5	7.92	8.07	5.47	19.7	240	2.11	2.83
16-01	15	55	0.18	0.73	12.3	0.14	0.03	0.55	275	0.37	183	185	0.95	6.10	12.0	29.8	41.5	32.2	24.6	8.35	7.37	3.57	17.4	244	2.16	7.40

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mdd) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	(mqq) sA	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-01	15.2	55.2	0.31	2.11	7.20	0.16	0	0.57	616	0.36	184	192	0.84	6.02	12.3	38.1	47.2	31.6	23.3	7.60	6.66	2.18	15.5	242	1.72	4.92
16-01	15.4	55.4	0.28	1.76	9.01	0.16	0	0.53	297	0.34	174	190	0.74	5.98	12.7	38.1	41.5	29.8	22.8	7.83	5.59	2.51	16.1	231	2.30	6.54
16-01	15.6	55.6	0.19	0.66	13.2	0.15	0.02	0.47	0.0	0.30	149	181	0.60	5.77	10.4	26.6	36.9	24.3	21.5	7.75	5.45	1.48	13.3	254	1.83	5.81
16-01	15.8	55.8	0.20	0.90	12.7	0.15	0.02	0.43	271	0.30	139	167	0.99	5.85	9.31	24.6	35.9	22.0	18.6	7.49	5.57	0.62	11.9	251	1.36	5.11
16-01	16	56	0.33	2.23	7.55	0.16	0	0.43	0.0	0.28	142	171	0.30	4.32	7.56	30.3	36.9	21.1	20.1	7.75	5.75	0.00	9.63	241	2.07	6.02
16-01	16.2	56.2	0.50	3.70	1.89	0.17	0	0.45	433	0.28	149	179	0.89	5.93	11.7	49.9	47.3	25.2	21.9	8.35	4.96	0.00	12.7	233	1.28	9.00
16-01	16.4	56.4	0.56	4.29	0.16	0.18	0	0.46	0.0	0.27	146	176	0.83	5.42	9.6	88.6	48.6	22.3	17.8	10.8	4.92	0.00	14.1	190	4.07	3.55
16-01	16.6	56.6	0.60	4.63	0	0.18	0	0.44	1550	0.25	139	182	1.04	6.19	11.9	128.0	62.3	21.6	16.9	8.63	8.35	0.00	12.7	200	1.78	4.89
16-01	16.8	56.8	0.63	4.72	0	0.18	0	0.43	1286	0.21	126	177	0.61	6.18	20.1	132.4	60.9	26.6	25.3	9.75	11.67	0.00	10.7	159	5.11	8.17
16-01	17	57	0.62	4.70	0	0.18	0	0.42	0	0.26	137	172	0.74	6.52	18.7	159.4	69.4	22.0	22.4	10.2	6.24	0.00	10.3	236	2.57	10.34
16-01	17.2	57.2	0.50	3.74	2.08	0.17	0	0.48	1002	0.38	194	203	0.00	5.32	19.6	125.0	75.5	24.5	21.6	9.97	6.64	0.00	13.3	191	4.41	9.21
16-01	17.4	57.4	0.20	1.16	11.8	0.13	0	0.64	0.0	0.45	226	233	0.54	6.20	16.4	64.8	65.7	49.6	33.0	9.26	8.44	6.44	17.2	257	2.89	4.92
16-01	17.6	57.6	0.08	0	18.1	0.13	0.06	0.59	209	0.40	205	215	0.91	6.58	15.1	47.5	55.9	39.1	26.9	7.91	9.72	6.07	18.3	266	2.23	4.09
16-01	17.8	57.8	0	0	22.4	0.11	0.12	0.56	345	0.36	183	192	0.57	5.48	12.1	29.2	42.2	31.3	23.9	8.68	7.49	2.27	15.3	243	3.16	6.60
16-01	18	58	0.04	0	21.8	0.11	0.14	0.56	298	0.35	178	201	0.93	6.89	11.7	22.3	39.7	29.8	21.1	7.67	6.64	1.48	16.5	215	1.69	4.12
16-01	18.2	58.2	0.04	0	21.0	0.10	0.17	0.59	0.0	0.33	166	193	0.10	5.36	11.1	19.8	35.7	29.0	22.9	7.45	5.74	4.04	13.7	203	1.32	0.00
16-01	18.4	58.4	0.06	0	19.8	0.11	0.14	0.55	99.8	0.35	171	190	0.00	4.91	10.6	20.1	35.6	27.8	21.1	7.78	4.68	4.36	14.0	185	1.63	0.00
16-01	18.6	58.6	0.06	0	19.0	0.12	0.09	0.55	322	0.41	196	188	0.65	5.95	9.9	21.7	35.4	28.9	23.3	7.76	5.29	2.70	13.7	204	1.89	1.12
16-01	18.8	58.8	0.07	0	18.5	0.13	0.09	0.53	650	0.40	193	183	1.03	6.21	10.7	24.5	38.1	28.6	23.3	7.45	4.63	0.81	15.6	223	1.98	0.78
16-01	19	59	0.07	0	19.3	0.12	0.09	0.51	0.0	0.35	173	183	0.43	5.32	11.7	20.0	40.2	25.3	19.9	7.55	6.30	0.00	32.8	219	2.69	3.42
16-01	19.2	59.2	0.05	0	19.5	0.12	0.11	0.53	0.0	0.40	200	209	0.71	7.25	14.8	30.1	49.4	36.5	25.5	8.24	8.65	0.70	75.5	190	5.85	1.40
16-01	19.4	59.4	0.04	0	18.9	0.12	0.13	0.56	115.6	0.44	227	222	0.52	8.09	20.5	36.3	63.6	47.7	31.2	8.09	13.5	0.00	111	181	6.02	0.70
16-01	19.6	59.6	0.05	0	18.1	0.12	0.13	0.61	221	0.48	248	238	0.68	8.96	24.0	45.6	79.2	54.8	35.6	8.05	17.1	1.84	122	167	8.11	0.17
16-01	19.8	59.8	0.05	0	18.3	0.12	0.13	0.57	0.0	0.45	228	238	1.12	9.18	21.0	39.7	73.3	55.2	34.6	8.54	15.4	3.26	99.8	161	6.13	2.30
16-01	20	60	0.07	0	18.7	0.13	0.11	0.55	0.0	0.43	221	222	1.04	8.90	22.1	36.6	68.5	48.4	33.7	8.34	14.9	2.78	73.3	175	5.08	3.39
16-01	20.2	60.2	0.13	0.18	16.8	0.14	0.07	0.50	172	0.39	198	209	1.10	8.15	16.9	38.1	60.2	44.0	30.2	8.71	12.7	4.25	58.7	180	4.67	1.95
16-01	20.4	60.4	0.24	1.24	12.8	0.16	0.00	0.49	0.0	0.35	175	195	1.09	7.33	12.7	50.6	57.0	37.0	26.0	8.25	11.7	3.14	44.5	181	4.41	5.55
16-01	20.6	60.6	0.50	3.76	3.68	0.19	0	0.43	46.7	0.28	144	183	0.73	6.55	13.5	66.1	49.3	27.8	24.0	9.45	7.67	0.00	29.1	191	2.90	3.41
16-01	20.8	60.8	0.64	4.97	0	0.20	0	0.38	1452	0.17	105	155	0.04	3.71	11.1	188	57.2	14.9	14.7	10.3	4.72	0.00	2.64	108	0.92	7.64
16-01	21	61	0.66	5.08	0	0.19	0	0.36	301.0	0.11	89	159	0.29	5.83	34.3	455	141.0	27.2	22.4	10.5	4.60	0.00	2.96	110	2.10	0.01
16-02	0	49	0.24	1.40	9.8	0.12	0.04	0.56	66.0	0.42	212	206	1.44	1.13	15.9	40.1	54.4	41.3	28.0	9.19	11.6	5.81	18.3	190	2.20	5.66
16-02	0.2	49.2	0.22	1.24	10.0	0.12	0.05	0.54	213	0.43	215	201	0.55	5.94	15.6	35.6	55.4	43.6	31.8	9.01	11.2	9.27	19.1	189	2.92	3.92
16-02	0.4	49.4	0.21	1.03	10.6	0.12	0.05	0.54	139	0.42	214	203	0.97	6.87	16.0	35.7	57.3	43.3	31.2	8.49	12.2	7.97	18.0	187	3.63	3.77
16-02	0.6	49.6	0.19	0.80	11.1	0.13	0.05	0.54	195	0.43	217	205	1.29	1.31	14.3	34.7	56.8	41.1	27.2	8.34	13.5	3.64	20.2	217	2.82	5.62
16-02	8.0	49.8	0.19	0.92	11.0	0.12	0.06	0.54	403	0.45	227	204	0.74	0.31	15.3	36.3	50.0	42.9	28.8	8.09	10.1	1.02	1/./	249	2.68	8.23
16-02	1	50	0.19	0.85	11.2	0.12	0.06	0.55	0	0.44	224	202	1.19	1.31	10.5	35.8	54.0	39.9	20.1	8.62	9.89	4./1	16.0	242	3.00	3.8/
16-02	1.2	50.2	0.19	0.96	10.9	0.12	0.05	0.57	102	0.45	229	199	1.2/	6.93	12.8	37.7	56.2	41.3	21.1	8.73	8.90	5.99	16.6	231	2.79	6.68
16-02	1.4	50.4	0.21	1.02	10.3	0.13	0.03	0.61	0	0.46	232	200	0.94	6.56	14.9	35.0	54.9	44.0	31.3	7.90	9.96	7.18	17.5	233	2.49	3.43

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mdd) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	(mqq) sA	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-02	1.6	50.6	0.19	0.81	11.4	0.12	0.05	0.60	296	0.43	219	204	1.04	6.81	14.8	30.2	57.4	41.9	31.1	8.64	12.2	5.12	17.2	233	2.61	6.13
16-02	1.8	50.8	0.18	0.96	11.5	0.12	0.05	0.56	577	0.43	220	201	1.15	7.11	15.8	34.7	55.0	40.8	30.0	8.60	10.9	6.48	18.9	234	3.25	6.40
16-02	2	51	0.14	0.21	13.4	0.12	0.08	0.60	80.7	0.43	218	205	1.16	7.07	14.9	33.3	57.4	45.9	32.9	8.44	11.8	8.78	19.9	233	2.68	6.91
16-02	2.2	51.2	0.13	0.44	13.4	0.12	0.08	0.59	232	0.44	227	213	0.89	6.96	16.8	32.5	53.3	43.1	28.1	8.49	11.9	9.00	20.5	222	2.83	3.87
16-02	2.4	51.4	0.15	0.30	12.9	0.13	0.08	0.59	502	0.43	223	214	0.72	6.61	15.7	32.3	56.4	45.4	30.4	8.23	13.1	8.52	20.3	202	3.02	5.53
16-02	2.6	51.6	0.14	0.37	13.6	0.12	0.08	0.60	198	0.43	216	207	1.03	6.99	14.0	32.9	55.4	43.9	31.5	9.18	13.4	8.99	17.9	210	2.48	0.62
16-02	2.8	51.8	0.23	1.24	9.96	0.14	0.03	0.59	257	0.41	207	204	0.67	6.02	12.9	41.0	53.2	39.0	30.0	8.45	11.6	7.11	20.2	186	1.83	1.58
16-02	3	52	0.37	2.56	5.51	0.16	0	0.61	42.0	0.40	203	199	1.35	7.34	14.2	53.4	60.1	40.5	27.9	8.79	8.72	5.84	19.0	171	2.37	2.70
16-02	3.2	52.2	0.55	4.22	0.10	0.18	0	0.56	604	0.35	181	193	0.77	6.43	17.3	111	59.2	35.1	24.9	9.05	7.75	1.28	12.2	141	1.57	1.68
16-02	3.4	52.4	0.63	4.96	0	0.19	0	0.43	681	0.26	136	196	1.06	7.10	19.6	282	106.1	40.5	28.8	11.11	7.44	0	4.93	120	2.16	7.89
16-02	3.6	52.6	0.66	5.01	0	0.19	0	0.43	566	0.24	116	183	0.61	7.11	37.7	428	130.9	36.1	29.8	12.66	9.71	0	11.7	94	3.05	11.9
16-02	3.8	52.8	0.64	5.01	0	0.19	0	0.43	2121	0.21	108	177	0.69	6.58	31.6	389	152.8	33.1	22.9	10.99	6.43	0	8.49	103	2.25	8.19
16-02	4	53	0.63	4.83	0	0.18	0	0.50	949	0.25	135	179	1.35	7.97	27.9	324	144.6	36.6	30.6	12.98	16.2	0	0.00	98	4.28	8.25
16-02	4.2	53.2	0.55	4.33	0	0.18	0	0.48	0.0	0.30	158	188	1.76	7.68	14.1	147	71.1	28.8	19.6	7.52	7.41	0	12.5	162	0.07	5.84
16-02	4.4	53.4	0.44	3.26	2.95	0.16	0	0.51	115	0.33	172	188	1.13	6.81	17.0	72.2	50.4	28.3	23.6	9.09	7.21	0	13.0	187	1.39	3.99
16-02	4.6	53.6	0.26	1.64	8.27	0.14	0	0.58	436	0.38	193	207	0.86	6.72	17.9	48.1	49.0	39.3	29.2	8.45	8.87	4.72	17.0	189	2.34	7.91
16-02	4.8	53.8	0.16	0.71	12.5	0.12	0.05	0.56	127	0.37	190	194	1.15	6.39	11.8	30.6	42.4	39.7	28.5	7.91	5.81	4.79	17.2	190	2.03	3.06
16-02	5	54	0.09	0	15.8	0.12	0.09	0.57	549	0.40	204	196	1.50	7.40	13.4	30.4	45.3	42.8	29.7	8.08	9.98	4.13	20.0	227	1.63	3.65
16-02	5.2	54.2	0.09	0	15.1	0.13	0.07	0.60	10.3	0.42	214	200	1.58	7.74	14.9	27.5	50.8	43.3	29.6	8.21	10.4	5.43	21.1	239	1.87	4.05
16-02	5.4	54.4	0.07	0	16.6	0.12	0.10	0.61	0	0.44	226	211	1.28	7.71	17.2	27.0	51.8	44.3	30.7	8.34	10.6	9.16	20.9	221	2.64	3.29
16-02	5.6	54.6	0.05	0	17.5	0.11	0.13	0.63	0	0.43	219	217	1.11	7.39	15.4	31.9	54.3	44.6	32.2	8.23	10.8	7.76	20.4	204	2.09	0.52
16-02	5.8	54.8	0.05	0	17.5	0.11	0.13	0.63	157	0.43	217	214	0.94	7.27	16.9	26.8	51.0	46.4	30.1	8.10	12.7	9.39	19.7	191	3.16	1.72
16-02	6	55	0.06	0	17.1	0.11	0.11	0.63	294	0.41	210	209	1.14	7.39	15.3	28.8	50.3	45.9	32.7	8.47	11.4	6.95	21.9	228	2.58	4.19
16-02	6.2	55.2	0.08	0	16.5	0.11	0.11	0.61	0.0	0.42	210	199	1.09	7.02	16.2	26.4	48.9	43.5	30.4	8.17	10.8	8.58	21.3	259	2.56	3.94
16-02	6.4	55.4	0.07	0	16.4	0.12	0.11	0.60	171	0.41	210	201	1.21	6.93	13.8	25.5	45.3	44.4	30.5	8.10	9.75	7.59	20.6	329	2.34	8.26
16-02	6.6	55.6	0.06	0	17.2	0.12	0.11	0.58	0.0	0.43	217	201	0.93	6.42	15.1	29.0	50.7	42.9	29.3	8.55	9.28	4.37	20.5	330	1.44	8.97
16-02	6.8	55.8	0.05	0	18.1	0.12	0.11	0.63	0.0	0.43	219	203	1.09	6.66	14.9	29.3	49.4	41.9	29.8	8.19	8.75	7.85	18.4	286	3.01	5.28
16-02		56	0.06	0	17.0	0.12	0.09	0.63	326	0.44	227	207	1.66	8.01	15.3	28.1	50.2	49.9	34.2	8.34	10.3	7.20	19.7	250	2.45	3.65
16-02	7.2	56.2	0.06	0	16.7	0.12	0.10	0.64	11.1	0.44	225	215	1.56	7.95	15.5	33.4	55.6	52.0	32.8	8.16	10.9	8.22	25.4	214	2.42	1.05
16-02	7.4	56.4	0.06	0	16.7	0.12	0.09	0.68	360	0.44	229	214	1.64	8.54	19.0	33.2	55.3	51.5	34.3	1.11	11.7	11.1	25.3	192	1.92	2.45
16-02	7.6	56.6	0.05	0	17.6	0.12	0.09	0.70	241	0.43	223	207	1.29	7.55	17.8	31.3	53.2	53.6	33.4	8.64	9.95	8.69	24.8	178	2.64	0.17
16-02	7.8	56.8	0.08	0	16.2	0.12	0.10	0.70	0	0.43	219	201	1.86	8.47	17.0	29.7	50.8	46.8	32.9	7.85	10.0	9.38	23.0	184	3.11	2.16
16-02	8	57	0.04	0	17.7	0.11	0.11	0.67	0.0	0.43	218	198	1.34	6.92	13.4	24.7	47.0	50.9	32.2	8.18	9.35	6.76	22.4	206	2.66	2.12
16-02	8.2	57.2	0.07	0	16.2	0.12	0.10	0.64	/1.3	0.44	225	205	1.62	7.92	16.4	32.7	50.5	53.4	33.2	8.34	11.8	7.89	23.7	214	2.01	0.97
16-02	8.4	57.4	0.06	0	16.6	0.12	0.10	0.61	604	0.45	230	211	1.67	8.24	17.3	28.9	55.1	54.2	33.6	8.35	11.0	9.57	23.8	216	2.66	3.31
16-02	8.6	57.6	0.07	0	16.1	0.12	0.09	0.59	289	0.45	234	215	1.49	1.85	16.2	31.6	53.9	55.9	34.7	8.54	11.8	11.6	26.3	207	2.35	2.79
16-02	8.8	57.8	0.07	0	16.5	0.12	0.09	0.58	54.2	0.44	226	211	2.02	9.07	17.6	28.5	50.3	52.8	34.3	7.90	9.79	10.1	23.1	191	2.99	0.77
16-02	9	58	0.07	0	16.6	0.12	0.09	0.61	742	0.42	216	208	1.42	7.75	16.3	29.3	52.1	50.4	32.2	8.29	10.3	11.3	22.7	193	2.91	0.58

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-02	9.2	58.2	0.06	0	17.0	0.13	0.09	0.63	0	0.44	225	208	1.28	7.45	16.1	33.7	54.2	55.1	36.2	8.62	10.3	10.1	23.1	213	3.10	0.50
16-02	9.4	58.4	0.07	0	15.9	0.13	0.07	0.64	0	0.47	241	215	1.58	8.41	18.0	38.0	58.9	62.4	39.4	8.10	11.5	7.86	27.6	224	3.60	3.81
16-02	9.6	58.6	0.07	0	15.6	0.13	0.06	0.66	32.4	0.47	246	226	1.06	7.42	17.4	42.1	60.2	59.7	37.7	8.17	11.2	9.59	26.7	255	3.41	5.38
16-02	9.8	58.8	0.08	0	15.2	0.12	0.06	0.62	67.4	0.44	228	218	1.07	7.50	18.4	43.3	60.6	53.8	34.7	8.12	10.6	9.87	24.2	223	2.77	3.44
16-02	10	59	0.07	0	16.0	0.12	0.07	0.61	229	0.43	218	219	0.73	6.63	15.2	37.7	57.5	51.7	33.2	8.29	12.0	10.3	21.3	228	3.09	2.82
16-02	10.2	59.2	0.06	0	16.4	0.12	0.09	0.59	98.8	0.41	208	210	0.82	6.69	14.7	28.8	51.9	49.5	32.8	8.11	10.3	6.86	23.4	250	2.78	4.36
16-02	10.4	59.4	0.06	0	16.6	0.13	0.09	0.61	79.3	0.42	213	209	0.73	6.59	15.9	32.6	50.9	46.8	31.6	8.01	9.96	7.32	21.5	266	1.96	5.22
16-02	10.6	59.6	0.07	0	16.4	0.13	0.09	0.61	336	0.41	211	208	1.31	7.81	17.0	27.4	49.8	47.0	31.5	8.44	10.8	6.71	21.8	243	2.44	3.52
16-02	10.8	59.8	0.07	0	16.5	0.13	0.08	0.59	244	0.40	202	205	1.35	7.53	14.9	26.4	50.2	46.7	31.1	8.37	9.17	5.77	21.1	217	1.56	3.07
16-02	11	60	0.06	0	16.8	0.13	0.09	0.56	0.0	0.39	199	201	1.34	7.14	13.5	25.1	44.8	39.4	26.7	8.05	9.49	6.97	21.1	192	1.71	1.83
16-02	11.2	60.2	0.07	0	16.6	0.13	0.09	0.56	776	0.39	195	192	1.21	6.94	14.7	28.4	48.7	40.4	28.5	8.08	9.33	6.26	18.6	192	2.51	1.48
16-02	11.4	60.4	0.08	0	16.1	0.13	0.08	0.56	286	0.40	201	195	1.17	6.98	15.4	25.1	48.9	39.2	27.1	7.84	8.13	7.61	17.0	224	2.27	3.75
16-02	11.6	60.6	0.06	0	16.7	0.13	0.08	0.59	332	0.40	203	196	1.29	6.71	11.5	27.0	46.1	39.5	27.4	8.36	8.48	3.93	15.4	230	3.46	3.08
16-02	11.8	60.8	0.08	0	16.5	0.13	0.08	0.62	0.0	0.43	214	190	1.49	7.18	12.7	27.0	44.9	39.8	28.4	8.26	8.55	5.54	16.4	279	3.00	6.22
16-02	12	61	0.09	0	15.3	0.13	0.07	0.62	482	0.44	222	192	1.15	6.49	12.6	28.0	43.9	42.8	27.4	8.08	10.3	7.28	16.7	267	2.65	5.27
16-02	12.2	61.2	0.07	0	16.3	0.13	0.08	0.60	211	0.44	224	200	1.63	7.70	14.3	26.6	47.3	46.9	32.0	8.20	9.21	7.92	18.3	227	3.18	0.82
16-02	12.4	61.4	0.08	0	16.0	0.12	0.08	0.59	81.6	0.43	217	205	1.76	8.26	15.4	28.7	50.1	49.5	32.2	8.00	8.56	7.04	22.5	183	1.92	0.41
16-02	12.6	61.6	0.07	0.03	15.7	0.13	0.08	0.57	508	0.45	233	216	1.86	8.73	16.4	32.7	56.0	55.8	35.7	8.03	10.4	12.0	26.3	165	3.30	0.51
16-02	12.8	61.8	0.07	0	15.7	0.13	0.07	0.59	493	0.46	239	224	1.85	9.40	19.9	36.5	56.6	60.2	35.5	8.49	13.4	10.4	28.4	182	2.93	0.94
16-02	13	62	0.09	0	14.9	0.13	0.07	0.61	492	0.46	239	224	1.77	8.88	17.1	35.6	58.2	65.3	40.7	8.66	12.8	10.5	29.5	200	3.16	2.42
16-02	13.2	62.2	0.09	0	15.1	0.13	0.06	0.61	15.5	0.46	237	218	1.93	8.99	16.3	33.1	55.9	63.0	37.0	8.26	14.2	9.62	28.8	201	3.01	2.62
16-02	13.4	62.4	0.11	0.08	14.3	0.13	0.06	0.60	17.3	0.44	229	216	1.32	8.04	18.9	35.9	54.4	58.4	36.4	8.37	11.7	9.56	27.5	221	2.48	2.12
16-02	13.6	62.6	0.11	0	14.2	0.13	0.06	0.61	0.0	0.44	225	216	1.20	7.66	17.8	37.3	55.6	59.4	35.6	9.07	11.7	11.3	25.2	215	2.89	2.63
16-02	13.8	62.8	0.19	0.95	10.9	0.14	0.03	0.60	142	0.44	225	214	1.33	7.94	17.4	42.9	61.0	56.0	35.6	8.82	11.6	10.4	24.9	210	3.82	3.86
16-02	14	63	0.27	1.70	7.97	0.15	0	0.60	940	0.43	222	217	1.37	7.92	17.1	51.0	57.3	54.8	34.6	8.84	12.5	9.15	25.4	203	3.79	4.27
16-02	14.2	63.2	0.34	2.36	5.73	0.15	0	0.59	915	0.41	213	223	0.85	7.00	14.6	54.6	59.3	43.6	3E+12	9.14	9.18	6.82	21.0	178	2.99	2.16
16-02	14.4	63.4	0.24	1.47	8.73	0.15	0.01	0.57	0.0	0.42	211	209	1.23	7.62	15.0	40.8	56.2	46.9	29.6	7.98	9.52	11.8	21.4	192	3.43	1.51
16-02	14.6	63.6	0.13	0.32	13.5	0.14	0.06	0.56	590	0.42	213	203	1.20	7.30	15.4	32.7	48.3	45.2	32.0	8.53	11.6	9.36	22.5	220	2.13	7.38
16-02	14.8	63.8	0.11	0	14.4	0.13	0.07	0.57	0.0	0.42	212	205	1.02	6.95	14.8	31.5	51.3	42.7	29.8	8.45	10.1	9.20	21.9	233	3.21	5.00
16-02	15	64	0.11	0.25	14.3	0.14	0.07	0.58	174	0.42	214	211	0.91	7.12	16.9	30.6	53.2	44.9	30.9	8.26	11.0	5.53	21.5	231	2.85	5.48
16-02	15.2	64.2	0.10	0	15.3	0.13	0.07	0.57	60.1	0.40	205	206	0.79	6.56	14.1	30.4	49.8	39.6	27.2	8.20	10.1	3.90	21.0	211	2.90	5.04
16-02	15.4	64.4	0.09	0.00	16.1	0.13	0.09	0.56	0.0	0.40	203	202	0.97	6.90	14.4	28.3	49.8	44.2	30.9	8.36	8.46	5.31	21.4	193	2.08	1.20
16-02	15.6	64.6	0.10	0	15.7	0.13	0.09	0.56	875	0.40	204	196	1.17	6.93	13.8	23.6	45.4	35.3	26.0	7.88	8.68	3.87	16.4	190	1.81	0
16-02	15.8	64.8	0.09	0.23	15.6	0.12	0.10	0.60	283	0.40	203	196	0.88	6.14	11.9	26.2	48.0	41.2	30.6	7.96	9.24	4.60	20.0	197	3.62	0
16-02	16	65	0.08	0	16.3	0.12	0.09	0.62	0	0.40	203	199	1.20	6.82	12.9	26.6	47.5	40.1	28.0	8.19	8.86	7.78	19.8	225	1.80	0.87
16-02	16.2	65.2	0.09	0	15.8	0.12	0.09	0.65	132	0.41	210	200	1.03	6.49	13.0	29.7	53.4	44.3	29.8	8.33	11.2	10.2	20.7	220	2.87	1.93
16-02	16.4	65.4	0.09	0.14	15.3	0.12	0.08	0.61	74.0	0.42	215	203	1.49	7.68	14.8	28.0	50.5	46.2	31.3	7.92	9.21	10.6	21.2	202	2.51	1.81
16-02	16.6	65.6	0.11	0	14.5	0.13	0.07	0.62	173	0.42	211	201	1.79	8.16	13.5	26.6	48.7	48.4	31.5	7.77	11.6	9.08	22.8	197	2.24	0.93
			-																							

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-02	17	66	0.15	0.26	12.8	0.13	0.05	0.59	402	0.41	208	206	1.03	6.84	14.3	30.3	49.6	44.9	31.6	8.11	7.59	5.55	21.9	219	2.11	2.54
16-02	17.2	66.2	0.17	0.73	11.7	0.14	0.04	0.59	794	0.43	220	212	1.18	7.36	16.4	33.7	48.7	41.5	28.4	7.30	7.11	4.48	22.0	234	1.86	4.71
16-02	17.4	66.4	0.17	0.74	11.7	0.15	0.05	0.60	541	0.44	226	217	1.13	7.60	18.7	41.5	57.1	52.5	33.4	8.49	12.2	8.26	23.8	227	2.76	4.20
16-02	17.6	66.6	0.16	0.62	12.0	0.14	0.05	0.62	0	0.47	245	225	1.43	8.47	20.6	42.8	68.5	71.0	41.1	9.03	14.7	11.4	31.2	191	4.60	1.85
16-02	17.8	66.8	0.13	0.35	13.4	0.13	0.06	0.64	272	0.48	250	229	1.18	7.58	17.2	40.0	68.7	73.5	41.3	8.70	15.9	13.6	31.9	206	4.02	4.00
16-02	18	67	0.14	0.40	13.0	0.13	0.06	0.61	0	0.46	236	222	1.42	7.90	16.7	38.6	58.6	61.9	38.5	8.50	16.7	9.76	29.2	212	4.07	2.01
16-02	18.2	67.2	0.12	0.12	14.1	0.13	0.07	0.58	0	0.44	223	207	1.39	7.52	15.2	32.2	53.8	50.6	32.6	8.12	15.4	10.03	24.7	198	4.36	1.23
16-02	18.4	67.4	0.12	0.15	14.2	0.13	0.07	0.59	84.2	0.41	207	199	1.30	7.12	14.1	25.0	44.8	41.3	27.8	8.12	10.9	9.0	22.0	211	3.38	2.96
16-02	18.6	67.6	0.09	0.21	15.1	0.12	0.08	0.62	224	0.39	200	196	1.13	6.61	12.8	25.1	46.9	42.1	27.9	8.29	9.36	6.34	18.1	215	2.69	3.60
16-02	18.8	67.8	0.10	0.00	15.0	0.12	0.09	0.64	276	0.43	220	210	1.33	7.10	13.5	31.1	53.0	53.7	34.9	8.51	11.2	9.67	23.5	231	2.54	4.20
16-02	19	68	0.09	0.16	15.1	0.12	0.08	0.61	575	0.45	234	219	1.55	8.16	16.9	36.0	56.2	57.3	35.5	8.23	11.6	9.98	25.1	239	3.42	4.04
16-02	19.2	68.2	0.09	0	15.5	0.12	0.08	0.64	0.0	0.45	233	218	1.70	8.65	17.8	33.8	55.3	59.1	36.3	8.31	12.9	11.47	26.4	241	3.20	3.68
16-02	19.4	68.4	0.05	0	17.5	0.12	0.09	0.63	172	0.44	227	219	1.82	8.89	17.9	33.7	55.9	59.3	34.6	8.48	12.0	8.92	27.0	221	3.14	3.18
16-02	19.6	68.6	0.04	0	18.5	0.12	0.09	0.58	0	0.44	227	218	1.04	7.15	16.4	34.0	54.3	59.3	36.4	8.23	13.4	9.77	25.8	212	2.70	3.32
16-02	19.8	68.8	0.02	0	19.5	0.12	0.10	0.60	0	0.43	221	214	1.53	8.30	18.1	29.6	52.9	53.8	35.8	8.28	12.1	8.22	24.5	225	2.53	3.71
16-02	20	69	0.09	0	15.9	0.13	0.09	0.56	130	0.41	211	201	1.83	8.30	15.5	25.0	50.4	49.5	33.9	8.26	10.6	7.80	23.6	213	2.53	1.99
16-02	20.2	69.2	0.10	0.10	15.2	0.13	0.07	0.53	226	0.39	196	193	1.33	6.89	12.8	23.2	44.2	42.6	28.8	8.19	7.36	6.58	20.3	194	2.42	0.00
16-02	20.4	69.4	0.12	0.11	15.3	0.14	0.05	0.50	351	0.35	181	189	1.66	7.45	13.6	16.1	37.3	37.2	26.5	8.24	6.49	3.11	18.6	193	1.71	1.89
16-02	20.6	69.6	0.18	0.76	13.0	0.15	0.02	0.48	0	0.33	168	180	1.48	6.44	9.0	24.2	36.5	35.1	26.6	8.19	7.88	2.76	17.3	192	2.44	0.00
16-02	20.8	69.8	0.36	2.45	7.07	0.17	0.00	0.46	0	0.31	155	168	1.33	5.75	7.2	42.8	36.0	29.3	23.7	8.47	6.62	0.00	12.1	198	1.42	4.32
16-02	21	70	0.57	4.36	0.39	0.18	0.00	0.46	554	0.27	137	142	1.01	4.63	12.6	102	44.4	21.1	15.5	7.85	0.96	0.00	4.11	166	0.00	0.00
16-03	0	61	0.28	1.75	8.38	0.15	0.02	0.31	0.0	0.41	208	207	1.62	7.83	10.4	76.3	57.9	42.5	27.1	10.35	11.2	7.99	17.4	1/1	4.93	5.53
16-03	0.2	61.2	0.27	1.65	9.32	0.14	0.02	0.32	0.0	0.39	200	206	1.22	7.42	14.5	51.5	54.8	43.3	28.9	9.14	12.9	7.94	20.1	205	3.65	3.87
16-03	0.4	61.4	0.24	1.35	9.51	0.14	0.02	0.28	412	0.40	203	212	0.77	6.32	12.9	56.1	56.7	45.6	28.9	9.46	13.4	11.21	21.1	209	3.44	12.45
10-03	0.0	01.0	0.21	1.01	10.8	0.14	0.03	0.31	996	0.39	201	209	1.38	8.10	17.8	50.4	53.9	44.3	29.3	9.24	10.7	10.2	20.3	218	3.62	4.71
16-03	0.8	61.8	0.10	0.71	12.8	0.13	0.06	0.35	0	0 11	0	0	0 5 9	0	150	0 40 F	0	10 1	0	0	10.1	0	10.2	0	0	10.2
10-03	10	02 62.2	0.12	0.54	14.0	0.12	0.00	0.30	420	0.41	200	202	0.50	0.13	10.0	40.5	01.0 10.0	40.1 20.0	27.5	0.10	10.1	0.90 5 96	19.2	207	2.74	7 4 2
16 02	1.2	62.4	0.11	0.10	10.0	0.12	0.00	0.39	430	0.43	219	204	0.09	0.23	14.0	34.7 20 0	49.0	30.0	21.9	0.02	10.6	5.00	10.0	297	2.90	7.42
16.03	1.4	02.4 62.6	0.12	0.21	14.0	0.13	0.09	0.39	210	0.44	222	200	0.65	0.0Z	10.7	20.9	40.9 50.0	40.0	20.0	7.00	10.0	5.91 4.54	10.3	212	2.29	7.00 5.45
16.03	1.0 1.0	62.0	0.12	0.15	14.4	0.13	0.07	0.39	210	0.43	220	204	0.59	5.85	10.0	20.0	53.0	40.0	20.1	802	10.1	5.29	19.5	174	3.57	3.40
16.03	1.0 2	02.0 63	0.10	0.00	15.5	0.13	0.00	0.42	509	0.43	222	200	0.03	5.00	16.0	33.4	00.Z	41.0	20.7	0.9Z 9.45	0.77	J.20 7 20	19.1 21.1	174	3.10	3.00
16.03	2	63.2	0.09	0.00	15.9	0.13	0.00	0.45	201	0.42	214	205	0.55	5.99 6.55	16.7	30.0	49.5	40.9	20.2	0.40	9.77 11 0	7.20 5.77	21.1 19.5	120	3.01	1.20
16.03	2.2	63.4	0.11	0.21	14.5	0.13	0.09	0.40	1.0	0.42	210	200	0.77	6.83	10.7	30.0	40.9 51.5	41.2	20.7	0.5Z	12.1	3.77	20.6	200	3.25	3 93
16-03	2.4 2.6	63.6	0.12	0.02	14.5	0.13	0.00	0.40	177	0.41	220	202	0.54	6.40	1/ 5	37.2	55.1	40.0	20.0	0.20	12.1 8.88	7.81	20.0	200	2 30	5 27
16-03	2.0	63.8	0.11	0.18	14.0	0.13	0.09	0.42	111	0.43	220	∠11 216	0.00	6.25	15.8	34.6	53.1	40.0	20.2 20.6	8 18	10.5	6.61	20.0 21 4	108	2.00	2.21
16-03	2.0 2	64	0.12	1.21	10.1	0.13	0.00	0.40	80.1	0.44	220	210	0.41	7.06	12.0	37.3	56.6	49.5	20.0	8 63	10.5	6.83	21.4	160	2 27	2.20
16.03	30	64.2	0.25	1.24 3.21	3 17	0.14	0.02	0.29	133	0.40 0 / Q	201	220 220	0.30	6/1	16.6	58.9	70.4	-10.0 51 0	3/ 7	0.00 8 70	10.2	0.00	20.0 20.9	179	2.07 1.80	2.24 8 01
10-03	3.2	04.2	0.40	3.31	3.17	0.17	0.00	0.13	433	0.40	∠40	232	0.32	0.41	10.0	0.00	10.4	51.9	34.7	0.12	10.5	9.30	20.0	1/0	1.02	0.21

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	(mqq) sA	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-03	3.4	64.4	0.65	4.98	0	0.18	0	0	442	0.42	216	213	1.13	7.76	18.9	89.5	65.9	46.9	27.6	9.76	13.7	6.87	18.4	149	5.62	4.20
16-03	3.6	64.6	0.67	5.04	0	0.19	0	0	1100	0.30	142	218	0.15	6.64	28.0	263	109	31.4	25.0	11.85	9.92	1.31	8.4	168	0.45	10.2
16-03	3.8	64.8	0.67	5.14	0	0.18	0	0	1591	0.16	89	178	0.14	4.80	20.0	373	120	21.1	18.7	13.72	11.4	0	7.5	80	1.80	8.12
16-03	4	65	0.68	5.01	0	0.18	0	0	627	0.10	78	157	0.10	4.89	31.4	420	150	17.0	10.1	15.60	17.5	0.14	2.6	55	1.66	0
16-03	4.2	65.2	0.64	5.01	0	0.18	0	0	0.0	0.10	73	148	0.48	4.94	28.7	442	164	21.9	19.4	13.83	9.8	0	13.7	66	0.38	0
16-03	4.4	65.4	0.48	3.46	2.54	0.17	0	0.10	1089	0.17	98	170	0.00	5.41	34.9	437	156	26.4	18.8	11.62	12.1	1.28	8.3	121	0.0	0
16-03	4.6	65.6	0.29	1.82	7.86	0.16	0	0.27	699	0.37	189	215	0.44	6.23	17.6	141	77.6	41.4	26.1	9.80	10.6	5.86	15.2	140	4.28	2.48
16-03	4.8	65.8	0.20	1.01	11.0	0.14	0.03	0.35	357	0.42	216	231	0.00	6.27	17.0	68.7	61.9	50.5	32.7	8.32	14.3	10.5	21.2	178	4.12	1.08
16-03	5	66	0.16	0.57	12.4	0.14	0.06	0.37	502	0.44	228	236	0.00	6.31	14.9	42.8	53.7	46.1	30.6	8.47	9.94	8.91	21.0	187	3.77	7.46
16-03	5.2	66.2	0.17	0.62	11.8	0.14	0.06	0.37	0.0	0.45	233	228	0.49	6.73	14.6	32.9	52.1	45.0	29.3	8.87	12.5	6.28	20.6	192	4.14	4.73
16-03	5.4	66.4	0.17	0.78	11.9	0.14	0.06	0.38	0.0	0.46	235	224	0.34	6.29	14.8	37.4	56.8	46.2	31.4	8.58	11.4	8.25	21.1	198	3.87	4.85
16-03	5.6	66.6	0.17	0.89	12.3	0.14	0.07	0.37	193	0.43	223	223	0.75	7.19	18.0	37.6	53.5	47.3	29.7	9.05	11.6	8.54	22.5	193	3.86	4.77
16-03	5.8	66.8	0.16	0.56	12.8	0.14	0.06	0.37	0.0	0.43	222	219	0.78	6.53	14.2	36.1	50.7	50.3	31.9	8.76	12.7	9.03	19.2	197	3.71	5.57
16-03	6	67	0.15	0.54	13.0	0.14	0.06	0.36	0.0	0.43	218	214	0.51	5.69	12.6	35.9	50.2	46.6	30.6	8.61	10.8	9.02	20.9	183	3.32	1.70
16-03	6.2	67.2	0.15	0.52	13.2	0.14	0.07	0.39	611	0.42	214	203	1.28	6.93	13.0	32.2	45.9	45.7	29.6	8.36	12.5	5.51	20.0	170	3.00	2.71
16-03	6.4	67.4	0.14	0.40	13.7	0.14	0.07	0.39	483	0.40	205	195	1.37	6.92	11.9	36.0	45.6	39.3	26.7	8.46	11.8	7.73	18.7	178	2.68	4.96
16-03	6.6	67.6	0.05	0	18.0	0.13	0.10	0.45	16.7	0.41	205	197	1.37	7.25	14.1	35.9	48.9	43.9	28.7	8.77	10.1	7.69	23.1	186	3.64	3.27
16-03	6.8	67.8	0.08	0	16.4	0.14	0.08	0.42	0.0	0.42	211	199	1.07	6.55	13.9	36.6	49.6	46.6	30.3	7.89	10.1	7.92	23.6	178	3.44	3.19
16-03	7	68	0.22	1.14	10.7	0.13	0.06	0.39	41.3	0.42	211	193	1.28	6.90	14.3	24.9	48.9	44.1	29.9	8.31	12.0	6.61	21.8	171	3.01	2.44
16-03	7.2	68.2	0.14	0.47	13.7	0.13	0.08	0.40	0.0	0.42	212	196	0.84	6.05	13.7	29.6	49.7	42.5	29.4	8.41	8.20	3.83	22.9	173	3.97	3.32
16-03	7.4	68.4	0.05	0	18.3	0.14	0.12	0.46	75.7	0.43	219	209	0.37	6.32	16.5	28.1	46.2	43.9	29.2	8.58	10.6	7.17	22.9	208	3.45	3.70
16-03	7.6	68.6	0.07	0	17.0	0.14	0.09	0.41	228	0.42	211	209	0.33	6.74	17.3	25.8	44.6	39.5	27.4	8.34	10.3	6.40	21.7	202	3.91	2.09
16-03	7.8	68.8	0.12	0.07	15.2	0.14	0.08	0.41	0.0	0.40	199	203	0.57	6.44	14.8	24.0	46.9	41.7	29.8	8.22	11.9	6.17	23.8	204	3.10	1.98
16-03	8	69	0.09	0.07	15.9	0.14	0.08	0.45	50.1	0.41	209	210	0.62	6.86	15.8	26.2	49.4	43.7	29.1	8.15	13.0	6.00	22.9	186	3.37	2.53
16-03	8.2	69.2	0.02	0	19.2	0.11	0.11	0.56	0.0	0.43	219	216	1.04	7.56	15.6	27.9	48.2	42.8	30.0	8.41	12.7	8.51	23.6	185	3.85	3.67
16-03	8.4	69.4	0.03	0	19.0	0.13	0.11	0.53	3.4	0.44	224	209	0.93	7.09	15.4	31.5	52.6	45.4	30.3	8.19	10.2	7.84	23.3	186	2.93	3.98
16-03	8.6	69.6	0.17	0.66	12.6	0.14	0.08	0.40	265	0.43	220	209	0.92	6.78	13.8	29.7	45.5	42.5	29.7	7.92	12.1	5.98	22.4	163	3.26	0.15
16-03	8.8	69.8	0.19	0.74	12.7	0.14	0.09	0.39	194	0.42	209	200	0.61	6.02	13.3	25.9	44.3	42.1	29.6	8.08	10.7	9.28	20.5	152	4.28	1.05
16-03	9	70	0.06	0	19.4	0.13	0.12	0.44	0.0	0.39	194	192	0.98	6.35	11.6	25.4	42.7	38.6	26.4	8.23	9.88	6.13	18.6	158	2.72	0.80
16-03	9.2	70.2	0.01	0	21.1	0.12	0.11	0.45	157	0.39	191	188	1.38	6.93	11.4	24.4	40.4	41.1	28.6	8.44	8.14	6.36	19.7	194	2.14	2.56
16-03	9.4	70.4	0.08	0	17.1	0.14	0.09	0.42	337	0.38	190	188	1.14	6.70	13.7	24.9	40.1	40.9	26.9	8.01	9.33	6.07	19.6	281	3.28	7.79
16-03	9.6	70.6	0.08	0	16.5	0.13	0.08	0.37	281	0.40	204	200	0.97	6.72	14.7	22.7	44.1	41.2	27.9	7.90	9.14	7.78	20.5	273	2.38	5.64
16-03	9.8	70.8	0.03	0	20.3	0.13	0.11	0.42	0.0	0.41	209	204	0.71	6.13	13.3	25.5	48.7	40.1	26.8	8.45	9.29	5.30	21.8	269	2.63	5.48
16-03	10	71	0.02	0	20.9	0.13	0.13	0.45	178	0.39	201	206	0.69	6.38	15.1	28.3	43.2	39.4	28.8	8.21	8.77	5.99	21.4	217	2.84	2.58
16-03	10.2	71.2	0.04	0	20.3	0.13	0.11	0.44	185	0.39	201	215	1.31	7.45	14.2	33.2	45.5	42.5	28.9	8.52	9.14	6.06	22.6	196	2.86	2.91
16-03	10.4	71.4	0.10	0	16.7	0.14	0.08	0.40	7.9	0.40	205	214	1.52	7.85	13.1	35.2	51.6	47.5	30.5	8.24	12.0	6.12	26.0	175	3.33	3.18
16-03	10.6	71.6	0.06	0	19.0	0.13	0.09	0.40	122	0.43	219	225	1.47	8.48	16.6	37.5	60.6	53.3	34.2	8.35	11.9	6.73	30.8	172	3.94	1.56
16-03	10.8	71.8	0.06	0	19.0	0.13	0.09	0.41	549	0.43	221	219	1.88	9.20	15.6	35.2	60.4	52.3	33.8	8.46	14.3	11.5	31.4	196	3.59	3.49

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-03	11	72	0.09	0	18.4	0.13	0.08	0.37	0	0.42	211	214	1.37	8.31	16.0	36.9	57.1	46.2	31.4	8.35	13.5	9.47	28.1	198	3.55	4.26
16-03	11.2	72.2	0.16	0.43	15.5	0.15	0.02	0.26	43.7	0.39	195	195	1.76	7.95	12.4	3E+12	48.5	40.3	27.0	8.93	12.2	7.87	23.9	185	3.58	2.39
16-03	11.4	72.4	0.39	2.68	7.79	0.17	0	0.15	3.1	0.30	147	164	0.88	5.19	9.86	21.7	36.1	27.4	21.5	8.26	8.17	2.61	15.5	175	2.68	4.19
16-03	11.6	72.6	0.43	3.10	5.29	0.18	0	0.09	282	0.25	114	134	1.20	4.10	4.77	27.7	22.8	17.0	15.7	7.85	5.50	0	9.01	148	1.70	5.69
16-03	11.8	72.8	0.36	2.49	7.45	0.18	0	0.11	0	0.23	110	133	1.02	4.01	8.00	49.1	25.2	15.7	12.7	8.39	3.40	0	9.25	97	3.38	2.64
16-03	12	73	0.31	1.98	8.96	0.17	0	0.16	124	0.23	121	145	0.73	4.17	10.4	55.0	28.3	20.1	17.9	8.74	5.71	0	8.10	104	3.15	5.01
16-03	12.2	73.2	0.33	2.12	8.94	0.16	0	0.19	0.0	0.23	123	146	0.48	3.24	5.76	43.6	26.5	19.3	15.0	8.06	4.25	3.22	8.29	137	4.23	3.66
16-03	12.4	73.4	0.40	2.83	6.27	0.16	0	0.17	30.0	0.23	112	139	0.77	3.48	4.23	31.3	22.7	16.8	15.7	8.56	7.09	0	5.77	133	3.10	4.12
16-03	12.6	73.6	0.38	2.58	8.38	0.17	0	0.16	0.0	0	107	135	0.83	3.38	3.81	16.5	21.6	15.6	14.8	8.38	4.66	0	7.69	114	2.91	2.77
16-03	12.8	73.8	0.36	2.41	9.91	0.18	0	0.12	127	0.18	99	130	0.78	2.92	2.50	15.9	15.5	14.2	12.8	8.52	3.83	0	6.71	89.9	1.65	4.63
16-03	13	74	0.36	2.43	9.80	0.17	0	0.10	178	0.17	92	127	0.88	2.83	0.47	16.7	17.3	13.7	11.4	8.52	3.37	0	8.40	77.7	3.00	0.00
16-03	13.2	74.2	0.55	4.10	1.61	0.19	0	0.06	46.6	0.25	129	143	1.31	5.05	10.3	54.7	40.2	24.8	17.6	9.29	6.57	0	11.0	79.5	5.26	5.72
16-03	13.4	74.4	0.47	3.59	2.47	0.19	0	0.08	0.0	0.28	149	153	1.62	6.23	14.6	55.9	45.7	25.1	21.4	8.54	6.87	0	12.6	108	2.94	1.36
16-03	13.6	74.6	0.37	2.75	5.11	0.18	0	0.13	165	0.29	146	146	1.63	5.87	12.3	42.5	43.2	29.5	19.3	7.90	6.65	0	11.7	124	1.44	0
16-03	13.8	74.8	0.27	1.73	9.56	0.16	0	0.17	0.0	0.26	134	146	1.63	5.41	6.96	26.7	31.7	27.3	18.4	7.89	4.97	0	9.29	108	1.57	0.64
16-03	14	75	0.17	0.63	14.8	0.16	0.03	0.21	0.0	0.26	123	141	1.62	5.30	6.13	16.0	25.1	20.9	17.1	7.40	3.66	0	7.05	102	1.15	0
16-03	14.2	75.2	0.13	0.28	16.8	0.16	0.05	0.21	97.2	0.24	120	140	1.39	4.69	5.02	13.5	22.1	21.3	18.9	7.81	3.81	0	7.90	111	0.76	0
16-03	14.4	75.4	0.11	0	18.4	0.15	0.05	0.24	0	0.25	124	142	1.60	5.26	6.61	15.2	25.2	23.3	18.0	7.66	3.46	0	8.17	105	2.09	0
16-03	14.6	75.6	0.18	0.69	14.6	0.15	0.02	0.21	0	0.24	126	146	1.53	5.29	6.91	25.5	30.3	23.2	19.9	7.82	6.12	0.22	12.1	119	1.69	0.25
16-03	14.8	75.8	0.31	1.95	9.34	0.16	0	0.13	162	0.26	133	144	1.99	6.07	7.21	38.1	35.3	27.1	20.2	8.27	7.10	0.12	10.9	131	1.49	2.10
16-03	15	76	0.43	3.18	4.72	0.18	0	0.08	649	0.32	160	154	2.06	7.03	12.7	57.6	45.8	37.4	25.3	8.93	6.78	0.09	11.9	143	1.93	2.05
16-03	15.2	76.2	0.47	3.47	4.58	0.17	0	0.06	0.0	0.31	154	152	1.76	6.35	12.1	55.0	43.9	35.0	25.6	9.56	9.76	3.91	13.0	125	3.01	0
16-03	15.4	76.4	0.42	2.89	9.31	0.16	0	0.06	473	0.24	118	136	1.32	4.50	6.74	46.0	35.2	27.6	22.3	8.36	3.54	0.45	9.48	103	1.97	0
16-03	15.6	76.6	0.30	1.83	13.8	0.16	0	0.09	0	0.23	113	143	1.53	5.22	5.50	34.2	29.8	22.5	17.8	8.23	2.07	0.00	9.40	106	1.64	0
16-03	15.8	76.8	0.21	0.86	18.3	0.16	0	0.14	0	0.23	117	147	1.12	4.82	8.06	25.6	24.8	22.6	19.1	7.59	2.95	0.03	7.55	133	1.24	0
16-03	16	77	0.19	0.80	19.8	0.15	0.03	0.16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16-03	16.2	77.2	0.14	0	24.3	0.15	0.04	0.19	0	0.23	112	135	1.41	4.63	5.69	16.4	20.6	20.8	17.7	7.77	2.85	0.00	7.97	130	1.59	0
16-03	16.4	77.4	0.09	0	26.1	0.14	0.07	0.24	0	0.23	111	130	1.52	4.56	5.65	13.3	20.1	22.6	18.0	7.76	4.75	2.43	8.96	124	1.82	0
16-03	16.6	77.6	0.09	0	26.9	0.13	0.09	0.29	0	0.24	120	133	1.74	5.05	5.24	7.84	20.3	23.3	19.6	7.08	2.07	1.26	9.72	141	0.90	0
16-03	16.8	77.8	0.09	0	28.0	0.12	0.13	0.33	0	0.23	116	134	1.88	5.34	4.89	7.85	20.8	21.5	17.8	6.85	3.27	0.78	6.89	151	0.95	0
16-03	17	78	0.14	0.14	23.7	0.14	0.08	0.26	0	0.23	115	135	1.77	5.18	5.15	9.77	15.6	19.7	16.6	7.40	0.00	0	5.23	155	1.11	0.49
16-03	17.2	78.2	0.25	1.27	16.4	0.14	0	0.18	0	0.21	108	139	1.61	5.03	4.06	16.5	20.8	21.0	17.9	6.99	1.84	0	2.86	161	0.00	0
16-03	17.4	78.4	0.33	2.23	10.8	0.14	0	0.13	167	0.22	120	145	1.32	5.10	9.41	26.6	22.1	22.2	19.0	8.62	4.76	0	7.29	134	2.06	2.71
16-03	17.6	78.6	0.42	3.07	6.38	0.15	0	0.10	1114	0.21	112	140	1.22	4.57	6.74	30.8	26.6	18.9	14.5	8.62	7.12	0	5.55	129	1.80	5.64
16-03	17.8	78.8	0.48	3.59	4.86	0.17	0	0.07	149	0.19	106	137	1.08	4.17	5.87	40.7	24.0	17.8	14.8	9.29	4.39	2.08	4.01	95	2.49	2.02
16-03	18	79	0.54	4.00	2.95	0.18	0	0.05	158	0.19	106	132	0.91	3.55	5.92	49.1	25.1	16.0	16.2	10.2	5.13	0	5.29	105	1.82	6.22
16-03	18.2	79.2	0.56	4.26	1.31	0.17	0	0.03	648	0.19	104	134	1.24	4.27	5.59	58.7	29.8	19.4	17.9	9.73	7.19	1.37	7.79	99.7	3.52	7.34
16-03	18.4	79.4	0.61	4.54	0.42	0.17	0	0.02	0	0.16	92	136	0.98	4.49	11.7	87.2	36.0	18.7	20.6	9.55	8.76	0	8.62	112	2.73	8.08

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	(mqq) nZ	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
16-03	18.6	79.6	0.62	4.76	0	0.18	0	0.01	0	0.16	103	141	0.60	4.10	15.0	129	58.3	18.7	20.2	11.0	9.25	0.00	9.26	105	3.92	6.56
16-04	0	95	0.24	1.42	8.55	0.14	0.04	0.28	343	0.57	299	0	0	8.40	24.4	91.0	103	92.7	45.3	9.70	20.6	17.3	14.7	258	4.63	10.4
16-04	0.2	95.2	0.26	1.62	8.10	0.15	0.03	0.26	465	0.55	295	158	0.80	10.1	25.5	84.9	96.7	86.4	44.3	9.76	18.7	15.7	14.8	274	3.23	10.2
16-04	0.4	95.4	0.23	1.33	9.05	0.15	0.04	0.27	590	0.56	286	240	0.75	9.78	25.8	84.7	104	87.9	43.1	8.87	20.8	16.4	14.2	266	2.87	13.2
16-04	0.6	95.6	0.22	1.14	9.19	0.15	0.04	0.27	130	0.58	301	272	1.54	11.1	26.1	74.3	98.2	92.2	46.5	9.49	21.1	16.8	14.5	242	3.70	9.65
16-04	0.8	95.8	0.23	1.26	8.97	0.14	0.03	0.26	311	0.57	300	286	0.97	9.56	25.9	83.4	103	88.1	45.9	10.0	17.7	18.2	14.0	272	3.67	11.9
16-04	1	96	0.21	1.08	9.74	0.15	0.04	0.27	28.3	0.56	279	302	1.95	10.7	23.9	80.2	102	80.9	43.6	9.81	16.1	15.6	13.1	290	4.80	11.0
16-04	1.2	96.2	0.21	1.11	9.51	0.15	0.04	0.26	0	0.56	289	274	1.15	9.31	27.6	75.4	99.4	78.1	43.8	9.90	18.6	13.6	13.2	330	3.72	11.1
16-04	1.4	96.4	0.21	0.97	9.76	0.15	0.04	0.25	0	0.55	289	249	2.06	10.6	25.9	72.6	98.3	77.5	42.5	9.59	20.3	14.3	14.3	329	3.65	13.7
16-04	1.6	96.6	0.23	1.23	9.03	0.15	0.03	0.23	391	0.55	280	226	1.36	8.87	25.7	74.9	94.1	74.4	38.5	9.25	18.0	13.8	14.0	313	5.02	12.2
16-04	1.8	96.8	0.23	1.29	9.29	0.15	0.04	0.24	358	0.54	266	215	2.34	10.3	23.2	81.7	93.2	69.1	39.3	9.06	18.0	14.8	12.8	336	4.18	11.9
16-04	2	97	0.23	1.27	9.44	0.15	0.03	0.23	46.6	0.55	271	210	2.72	11.0	22.5	81.0	97.2	70.6	39.5	8.96	18.2	11.9	14.4	358	3.90	13.9
16-04	2.2	97.2	0.23	1.20	9.17	0.15	0.03	0.23	109	0.54	261	205	1.96	8.83	18.8	81.3	98.7	66.6	37.2	9.33	19.0	11.9	12.4	341	3.44	11.2
16-04	2.4	97.4	0.24	1.49	8.54	0.15	0.02	0.24	0	0.56	259	214	1.86	9.67	25.2	88.0	94.5	62.9	37.3	9.07	18.4	14.8	10.1	309	5.01	10.6
16-04	2.6	97.6	0.26	1.62	7.82	0.16	0.02	0.22	251	0.54	246	209	2.20	9.78	21.8	87.6	102	71.8	40.8	9.46	18.5	11.8	12.8	287	3.68	12.4
16-04	2.8	97.8	0.27	1.81	7.54	0.16	0.02	0.23	0.0	0.55	259	203	1.98	9.68	24.5	93.2	104	77.9	40.5	9.29	18.9	13.5	12.6	270	5.48	8.82
16-04	3	98	0.30	1.99	6.78	0.16	0.02	0.22	0.0	0.55	254	200	2.26	9.95	23.0	90.5	99.5	65.3	35.4	9.34	21.3	14.7	15.2	280	4.39	10.2
16-04	3.2	98.2	0.29	2.01	6.45	0.16	0.02	0.21	0.0	0.56	228	213	2.12	10.3	25.6	102	106	70.9	39.0	9.30	18.9	18.2	12.6	315	2.80	14.1
16-04	3.4	98.4	0.31	2.20	5.95	0.16	0.03	0.22	592	0.59	260	218	1.79	10.5	27.7	109	118	83.2	44.2	9.09	22.6	19.7	15.4	310	4.20	12.6
16-04	3.6	98.6	0.30	2.04	6.55	0.16	0.04	0.23	55.4	0.62	284	225	1.07	9.83	29.2	132	134	80.1	41.4	10.2	23.8	23.4	17.2	294	4.35	12.0
16-04	3.8	98.8	0.32	2.15	6.81	0.16	0.04	0.23	0	0.64	317	227	0	7.86	30.4	154	146	84.9	42.1	9.06	27.7	26.5	16.7	295	3.57	9.06
16-04	4	99	0.34	2.28	6.33	0.17	0.03	0.21	339	0.64	319	239	0	8.98	36.9	195	170	80.7	41.7	11.2	31.7	29.8	21.5	267	5.98	8.08
16-04	4.2	99.2	0.38	2.59	5.74	0.17	0.03	0.18	0	0.63	279	238	0	8.79	35.9	277	215	90.5	41.6	13.8	36.2	33.7	22.3	225	5.41	10.9
16-04	4.4	99.4	0.48	3.41	3.20	0.18	0	0.12	460	0.43	153	198	2.96	11.7	33.9	324	163	59.3	35.2	12.5	20.9	20.9	2.4	185	2.45	11.5
16-04	4.6	99.6	0.46	3.40	2.52	0.18	0	0.10	530	0.50	216	195	3.70	12.5	25.2	199	123	46.2	30.5	9.80	21.5	15.0	11.7	209	4.58	9.83
16-04	4.8	99.8	0.42	3.10	3.07	0.16	0	0.11	0	0.52	229	205	2.61	10.7	24.8	141	115	48.6	33.4	8.92	23.8	16.5	16.9	203	5.94	4.63
16-04	5	100	0.30	2.05	6.40	0.16	0.04	0.18	619	0.54	245	213	2.39	11.1	28.2	132	121	49.8	33.1	9.98	25.1	22.7	15.6	236	4.27	14.4
16-04	5.2	100.2	0.27	1.68	7.69	0.16	0.07	0.23	257	0.55	265	207	2.89	11.1	21.1	125	122	54.1	31.2	9.23	24.9	22.5	18.5	246	5.04	1.44
16-04	5.4	100.4	0.26	1.67	8.11	0.15	80.0	0.24	229	0.57	276	215	2.23	11.2	28.5	111	123	59.6	36.9	9.13	24.2	25.5	19.5	254	5.24	10.5
16-04	5.6	100.6	0.28	1.91	6.98	0.15	0.07	0.22	131	0.56	266	213	2.42	11.2	26.8	119	125	59.1	33.1	9.77	25.2	21.5	16.3	261	4.90	13.1
16-04	5.8	100.8	0.31	2.14	6.23	0.16	0.04	0.20	0	0.54	241	213	2.64	11.8	28.5	116	118	60.3	34.3	9.67	25.3	20.8	17.2	232	5.58	8.30
16-04	6	101	0.33	2.37	5.46	0.16	0.02	0.19	17.6	0.56	246	216	2.25	11.0	28.3	115	124	70.0	34.9	9.35	25.9	17.5	17.0	228	5.99	14.2
16-04	6.2	101.2	0.39	2.93	3.45	0.17	0	0.15	432	0.55	249	212	2.66	11.1	25.6	112	114	62.8	36.9	9.69	21.2	17.9	15.5	239	5.25	15.2
16-04	6.4	101.4	0.40	3.04	3.12	0.17	0	0.14	806	0.51	248	199	2.85	11.3	25.5	115	113	65.2	39.1	9.17	24.8	15.6	17.3	237	5.1/	13.6
16-04	6.6	101.6	0.42	3.10	2.98	0.17	0	0.14	0	0.53	252	214	3.38	12.5	25.6	109	106	61.3	33.4	11.2	18.9	17.9	13.6	223	6.40	9.06
16-04	6.8	101.8	0.39	3.02	2.98	0.1/	U	0.14	701	0.52	249	227	3.63	13.1	25.2	11/	118	/1.9	41.2	10.5	26.6	15.6	15.0	223	6.68	16.9
16-04	/	102	0.40	3.12	2.76	0.17	0	0.14	10.3	0.54	268	239	2.24	10.6	25.4	119	121	/3.4	42.1	9.82	23.1	18.1	13.6	218	5.60	10.8
16-04	1.2	102.2	0.43	3.33	2.07	0.17	0	0.12	771	0.53	268	256	2.63	11.6	25.0	124	126	70.9	33.2	10.6	21.5	17.7	16.4	226	5.94	8.60
core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	As (ppm)	(mqq) dA	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
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16-04	7.4	102.4	0.44	3.44	1.65	0.17	0	0.11	0	0.56	268	286	2.24	11.6	28.6	142	121	83.8	40.1	10.4	21.7	19.4	15.8	272	4.81	18.9
16-04	7.6	102.6	0.41	3.12	2.57	0.17	0	0.12	9.6	0.58	303	278	2.05	11.0	26.8	131	120	90.3	50.9	10.8	20.9	19.5	17.2	300	5.39	19.4
16-04	7.8	102.8	0.38	2.83	3.66	0.16	0	0.15	445	0.56	282	289	2.06	10.6	22.6	98.4	102	71.5	43.8	10.6	24.3	19.5	17.0	328	5.65	15.2
16-04	8	103	0.36	2.62	4.73	0.16	0	0.18	688	0.56	289	235	0.83	10.5	26.7	98.6	111	84.6	45.5	10.0	23.7	17.9	19.9	316	6.39	19.9
16-04	8.2	103.2	0.34	2.39	5.45	0.16	0	0.21	1492	0.57	293	0	0.00	8.50	31.2	98.7	121	84.9	43.7	10.8	20.2	25.8	22.8	301	6.13	16.9
16-04	8.4	103.4	0.33	2.24	6.08	0.16	0	0.23	229	0.59	315	0	0.00	8.72	29.6	95.6	117	78.3	41.7	10.1	21.6	22.5	18.2	296	5.77	13.1
16-04	8.6	103.6	0.30	1.95	7.50	0.16	0	0.24	843	0.57	306	0	0.00	6.67	27.5	100	103	88.5	43.6	8.86	20.8	19.4	15.3	282	5.23	11.7
16-04	8.8	103.8	0.33	2.14	7.50	0.17	0	0.20	320	0.54	284	200	0.00	6.61	22.4	88.8	102	70.6	38.8	9.68	18.2	15.7	11.6	265	3.49	9.81
16-04	9	104	0.43	3.06	4.53	0.18	0	0.12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16-04	9.2	104.2	0.61	4.55	0	0.20	0	0.04	1845	0.48	228	152	0.00	5.61	34.1	311	146	75.7	45.6	0	0	0	0	0	0	0
14-01	0	105	0.44	3.26	3.34	0.16	0	0.82	18.8	0.37	184	215	0.79	6.67	16.7	101	60.4	39.4	26.7	9.23	12.7	0.88	0	575	5.40	20.6
14-01	0.2	105.2	0.33	2.24	7.19	0.15	0	0.85	280	0.43	213	224	2.01	8.72	15.8	66.3	48.9	44.4	29.1	10.12	11.4	1.97	20.0	843	3.52	26.2
14-01	0.4	105.4	0.20	1.04	12.1	0.14	0	0.88	329	0.46	239	208	2.26	8.84	14.0	43.1	51.3	46.5	29.4	8.96	12.8	1.24	22.1	795	4.67	23.5
14-01	0.6	105.6	0.13	0.26	16.0	0.12	0.03	0.82	135	0.44	229	210	2.02	8.60	13.6	35.5	46.7	44.5	28.6	8.50	9.03	0	18.9	686	2.55	19.5
14-01	0.8	105.8	0.10	0.09	18.1	0.13	0.03	0.76	751	0.44	226	217	2.32	9.74	16.6	34.4	47.2	46.9	29.9	8.63	8.12	0	16.9	676	1.28	21.5
14-01	1	106	0.07	0	19.8	0.12	0.06	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14-01	1.2	106.2	0.06	0	20.7	0.12	0.08	1.12	822	0.47	241	229	2.37	9.68	16.6	34.2	50.0	44.1	28.4	8.92	9.51	4.17	20.1	1273	2.76	24.7
14-01	1.4	106.4	0.07	0	21.5	0.11	0.08	1.13	719	0.51	272	223	2.26	9.58	15.8	36.3	46.9	37.5	25.4	9.23	7.41	1.08	17.8	1327	2.52	23.1
14-01	1.6	106.6	0.07	0	21.3	0.12	0.08	1.01	651	0.51	269	229	0.91	8.65	15.6	28.3	47.1	38.5	24.9	8.65	9.59	0.00	15.7	1022	1.86	26.5
14-01	1.8	106.8	0.08	0	21.5	0.12	0.07	0.88	202	0.48	248	221	0.77	8.27	14.2	28.9	45.1	38.4	26.3	9.22	8.69	1.56	17.4	994	3.22	23.4
14-01	2	107	0.08	0	20.5	0.11	0.07	1.02	1021	0.49	259	238	1.80	9.90	15.9	34.6	50.2	43.7	27.5	9.47	10.8	4.75	19.9	1417	2.69	21.2
14-01	2.2	107.2	0.09	0	18.8	0.11	0.08	1.15	821	0.54	285	237	3.17	11.4	17.7	37.6	60.4	47.9	30.1	9.47	14.5	6.55	22.8	1832	3.10	6.43
14-01	2.4	107.4	0.11	0.05	16.1	0.11	0.07	1.26	764	0.55	284	244	3.15	11.3	18.8	42.4	71.1	53.3	33.2	9.09	17.1	13.74	24.9	1739	3.18	10.2
14-01	2.6	107.6	0.11	0.20	16.7	0.11	0.10	1.23	552	0.50	258	216	3.44	11.2	16.5	39.2	64.1	53.5	33.1	8.57	12.9	14.35	21.8	1331	2.83	22.3
14-01	2.8	107.8	0.13	0.27	16.9	0.11	0.07	1.05	525	0.43	222	197	2.46	8.81	13.4	27.3	48.7	41.4	28.2	8.31	10.5	4.68	15.2	876	2.91	23.1
14-01	3	108	0.16	0.41	16.2	0.13	0.04	0.84	532	0.36	183	183	1.86	7.59	10.9	18.3	42.5	37.0	25.3	8.05	8.19	0	13.5	517	2.22	16.5
14-01	3.2	108.2	0.15	0.29	15.5	0.13	0.04	0.75	490	0.38	192	188	2.57	9.09	10.5	18.4	54.9	36.2	27.0	8.62	10.9	0.29	15.1	747	2.04	21.6
14-01	3.4	108.4	0.15	0.61	14.1	0.12	0.05	0.82	793	0.48	251	217	3.13	10.8	14.8	37.3	64.9	50.0	32.1	8.71	15.5	4.66	20.6	1615	3.00	15.0
14-01	3.6	108.6	0.13	0.38	14.9	0.12	0.07	0.91	915	0.55	294	224	3.64	11.9	16.9	41.2	67.2	53.1	32.1	9.44	17.4	8.98	23.8	2057	3.52	0
14-01	3.8	108.8	0.11	0.33	16.0	0.11	0.08	0.92	498	0.52	273	223	3.47	11.8	18.5	40.8	62.2	52.3	31.9	9.16	15.8	8.17	22.7	1951	3.27	0.17
14-01	4	109	0.11	0.31	15.8	0.12	0.08	0.93	871	0.49	254	204	3.48	11.2	15.5	32.3	60.4	53.8	33.4	8.79	16.1	7.17	23.2	1296	3.64	21.8
14-01	4.2	109.2	0.10	0.00	16.6	0.11	0.09	0.90	650	0.47	246	196	3.68	11.5	15.7	36.6	55.9	55.0	32.4	8.84	16.2	6.97	22.1	1107	2.56	24.1
14-01	4.4	109.4	0.12	0.52	15.0	0.12	0.08	0.94	1312	0.48	251	207	3.86	12.3	17.6	39.9	62.6	52.2	31.3	8.67	15.3	5.45	24.4	986	3.23	23.8
14-01	4.6	109.6	0.11	0.08	15.7	0.12	0.08	0.99	640	0.49	254	216	3.24	11.3	17.2	38.8	61.3	51.0	32.1	8.06	16.3	7.48	19.5	776	2.96	21.7
14-01	4.8	109.8	0.11	0.13	15.4	0.12	0.07	1.11	766	0.48	250	223	3.50	12.0	17.8	35.9	65.1	50.0	31.7	8.98	14.9	8.88	20.5	766	2.60	23.7
14-01	5	110	0.11	0	16.8	0.11	0.08	1.12	785	0.47	244	219	3.01	11.3	19.7	33.7	64.7	46.9	29.4	8.59	15.2	6.90	21.0	797	3.40	21.6
14-01	5.2	110.2	0.10	0.24	17.2	0.11	0.07	1.10	375	0.44	223	217	3.00	10.5	14.8	31.4	61.8	45.5	27.9	8.73	13.2	6.94	20.7	1000	3.24	24.2
14-01	5.4	110.4	0.11	0.20	16.7	0.11	0.07	1.07	218	0.45	231	217	2.86	10.3	14.4	30.8	57.8	44.4	28.3	8.99	12.4	8.25	21.3	1209	2.59	25.1

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mdd) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
14-01	5.6	110.6	0.12	0.28	16.1	0.11	0.07	1.07	412	0.44	224	216	2.89	10.4	15.1	29.4	57.6	44.6	29.1	8.84	14.1	7.58	20.79	1139	2.58	24.4
14-01	5.8	110.8	0.13	0.39	15.6	0.11	0.07	1.00	668	0.44	225	211	2.88	10.3	14.6	29.5	58.4	47.3	29.1	8.75	14.2	7.19	19.61	1051	2.41	24.1
14-01	6	111	0.12	0.16	17.8	0.11	0.07	0.91	215	0.42	216	202	2.98	9.89	10.7	23.3	48.2	41.4	27.7	8.67	11.8	5.37	17.12	1064	2.91	24.7
14-01	6.2	111.2	0.13	0.21	19.2	0.11	0.06	0.87	715	0.45	233	205	2.78	9.69	12.8	30.3	46.3	41.1	28.5	9.09	13.7	6.22	19.47	1379	2.92	22.3
14-01	6.4	111.4	0.10	0	20.4	0.11	0.10	0.97	548	0.49	256	210	3.14	10.7	15.1	31.2	51.9	46.9	29.4	8.99	13.0	8.42	22.27	1380	2.79	22.1
14-01	6.6	111.6	0.10	0	19.5	0.11	0.13	1.14	198	0.50	262	224	3.25	11.4	17.3	38.4	55.1	55.0	34.1	8.93	15.0	9.75	23.55	1218	2.92	23.9
14-01	6.8	111.8	0.10	0	18.3	0.11	0.13	1.23	586	0.52	270	234	3.55	12.4	19.2	42.9	64.1	62.3	35.5	8.57	15.2	10.87	26.11	1193	3.85	23.0
14-01	7	112	0.10	0.12	18.2	0.11	0.12	1.27	662	0.52	256	250	3.55	12.7	21.2	44.8	73.2	62.9	37.4	8.83	15.2	13.98	25.51	1444	3.50	19.3
14-01	7.2	112.2	0.09	0.22	18.3	0.11	0.10	1.36	296	0.50	258	231	3.21	11.9	21.5	43.0	68.2	54.3	31.7	8.96	15.5	13.93	24.94	1400	3.12	21.0
14-01	7.4	112.4	0.10	0.27	18.4	0.11	0.10	1.41	446	0.48	251	215	3.07	10.9	17.3	30.1	54.5	49.0	31.2	9.03	13.3	12.03	21.07	1146	2.81	25.2
14-01	7.6	112.6	0.11	0	19.2	0.11	0.11	1.29	691	0.45	233	201	3.21	10.8	15.3	30.7	49.8	46.3	29.7	8.60	12.2	12.01	15.49	910	3.03	24.8
14-01	7.8	112.8	0.11	0	19.5	0.11	0.11	1.21	578	0.43	223	206	2.82	10.1	13.7	24.7	46.8	40.5	27.7	8.59	11.7	10.60	18.48	902	2.73	24.9
14-01	8	113	0.11	0	20.8	0.11	0.13	1.15	512	0.40	205	198	2.70	9.57	12.3	23.0	45.5	39.6	27.8	8.75	10.5	11.56	16.38	738	2.62	23.5
14-01	8.2	113.2	0.11	0.08	22.5	0.11	0.12	1.11	272	0.38	194	193	2.51	8.97	10.7	23.4	40.3	39.1	26.6	8.48	9.29	6.88	15.15	591	2.13	18.1
14-01	8.4	113.4	0.10	0.15	22.5	0.11	0.10	1.00	633	0.36	185	191	2.61	9.38	12.6	18.3	39.6	37.6	26.7	8.36	10.3	8.04	11.60	540	2.86	18.6
14-01	8.6	113.6	0.12	0.03	21.6	0.12	0.07	0.87	211	0.35	176	182	2.21	8.17	9.3	18.3	37.6	35.8	25.7	8.35	9.19	3.78	13.80	523	1.79	16.2
14-01	8.8	113.8	0.11	0	22.9	0.11	0.08	0.72	383	0.35	176	180	2.55	8.88	10.3	19.3	35.1	32.3	23.8	8.13	9.15	4.12	12.42	496	2.20	16.7
14-01	9	114	0.10	0	21.9	0.12	0.08	0.67	303	0.35	173	178	2.80	9.20	9.18	17.8	37.5	34.7	24.5	8.59	11.7	3.08	11.69	471	1.67	15.1
14-01	9.2	114.2	0.11	0	21.9	0.12	0.07	0.60	0.0	0.33	161	172	2.51	8.51	9.38	20.0	33.0	32.2	24.1	8.64	9.81	2.67	11.00	382	1.47	10.3
14-01	9.4	114.4	0.12	0.08	22.3	0.12	0.05	0.49	201	0.29	146	165	1.77	6.56	6.40	10.8	27.5	24.7	19.4	8.22	8.08	2.21	8.10	353	1.17	10.3
14-01	9.6	114.6	0.11	0	22.1	0.13	0.06	0.47	179	0.28	137	161	1.74	6.43	5.78	9.9	25.0	23.4	19.1	8.23	7.28	0.00	8.81	392	0.91	10.3
14-01	9.8	114.8	0.12	0	19.9	0.13	0.07	0.52	309	0.33	156	167	2.09	7.57	8.50	13.4	33.1	29.6	21.8	8.16	8.50	1.71	10.43	378	1.99	8.80
14-01	10	115	0.13	0.18	18.7	0.13	0.08	0.57	597	0.36	178	177	2.41	8.40	9.45	10.0	40.6	34.6	25.5	8.13	10.6	3.63	9.77	332	2.87	8.90
14-01	10.2	115.2	0.14	0.10	18.1	0.13	0.07	0.63	309	0.38	186	182	2.86	9.59	10.6	18.1	43.3	39.3	23.8	8.26	12.8	6.45	11.82	317	2.35	6.33
14-01	10.4	115.4	0.15	0.62	16.0	0.13	0.07	0.70	589	0.41	209	198	2.99	10.3	11.9	26.7	51.3	46.4	31.0	8.35	13.2	6.70	14.39	280	3.03	4.72
14-01	10.6	115.6	0.16	0.73	14.0	0.12	0.06	0.78	111	0.44	224	213	2.87	10.4	12.7	29.1	60.0	54.0	32.9	8.17	15.3	6.89	14.37	267	2.56	3.81
14-01	10.8	115.8	0.17	0.72	13.1	0.11	0.07	0.81	442	0.44	226	216	2.54	10.4	16.8	28.2	59.8	52.7	33.5	8.20	15.3	6.65	12.97	301	1.71	3.25
14-01	11	116	0.16	0.74	13.0	0.10	0.08	0.76	540	0.41	206	203	2.68	10.1	13.2	26.9	55.1	44.9	29.2	8.74	11.5	8.10	15.31	370	2.83	10.8
14-01	11.2	116.2	0.17	0.67	12.9	0.10	0.06	0.71	0.0	0.37	183	185	2.08	8.31	12.6	21.8	48.4	34.9	25.4	8.61	11.0	5.03	12.33	588	1.35	15.9
14-01	11.4	116.4	0.17	0.73	12.9	0.09	0.05	0.69	350	0.36	184	1/1	2.42	8.08	11.4	23.8	43.5	30.1	22.0	8.18	8.81	3.31	12.73	/19	1.41	20.3
14-01	11.6	116.6	0.15	0.65	13.6	0.09	0.05	0.70	454	0.37	187	157	2.63	1.11	10.0	23.3	40.3	28.7	21.0	8.31	9.37	2.98	14.08	768	2.43	20.7
14-01	11.8	116.8	0.15	0.49	14.0	0.09	0.06	0.70	404	0.38	191	152	3.30	9.39	14.0	22.8	43.8	28.3	21.6	8.90	10.9	4.39	13.80	728	2.05	20.9
14-01	12	117	0.14	0.27	15.0	0.10	0.07	0.71	324	0.35	181	148	3.38	8.95	9.68	21.8	42.2	27.4	21.5	8.76	9.45	4.41	13.42	653	1.90	18.6
14-01	12.2	117.2	0.13	0.23	15.8	0.10	0.07	0.75	398	0.36	185	149	3.38	9.37	13.2	21.3	43.7	29.8	22.4	8.29	9.90	3.69	13.35	543	1.80	15.1
14-01	12.4	11/.4	0.13	0.22	16.1	0.10	0.07	0.78	0.0	0.35	181	148	3.21	8.81	11.6	21.5	41.1	30.7	22.3	8.45	8.39	4.78	13.05	432	1.79	9.88
14-01	12.6	117.6	0.14	0.35	16.0	0.10	0.07	0.//	110	0.38	192	152	3.60	10.0	14.4	24.2	40.2	35.0	24.4	8.28	9.36	4.32	13.65	326	1.50	8.00
14-01	12.8	117.8	0.16	0.66	14.7	0.10	0.08	0.75	252	0.37	188	150	3.54	9.43	10.4	24.2	43.6	35.7	24.3	8.05	10.3	3.58	14.37	326	2.08	6.51
14-01	13	118	0.16	0.56	13.5	0.10	0.07	0.68	510	0.40	202	154	3.45	9.63	12.3	24.9	47.5	39.3	25.4	8.20	10.8	8.29	14.11	370	2.45	9.84

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mdd) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
14-01	13.2	118.2	0.16	0.52	12.7	0.10	0.07	0.65	196	0.41	211	157	3.87	10.8	14.0	31.1	49.6	42.2	29.7	7.90	11.3	7.91	17.8	410	2.31	10.1
14-01	13.4	118.4	0.17	0.74	12.1	0.10	0.06	0.62	0	0.40	207	158	3.76	11.1	17.5	24.1	50.1	43.4	26.8	8.33	10.2	6.53	16.2	402	2.20	10.4
14-01	13.6	118.6	0.17	0.74	12.3	0.10	0.07	0.63	632	0.42	218	161	3.84	11.2	16.9	30.0	56.7	46.7	29.2	8.17	11.8	9.58	13.9	404	2.03	9.71
14-01	13.8	118.8	0.16	0.59	12.6	0.10	0.07	0.65	434	0.43	224	165	4.04	11.4	15.0	29.5	52.3	46.2	28.7	8.84	13.1	8.73	13.5	461	2.28	12.9
14-01	14	119	0.15	0.30	13.3	0.09	0.07	0.61	324	0.44	228	164	3.54	10.5	15.4	34.5	52.4	47.5	28.7	8.00	10.9	9.20	17.0	505	1.55	16.4
14-01	14.2	119.2	0.13	0.49	14.0	0.10	0.09	0.59	156	0.42	213	157	3.70	10.9	17.2	32.1	52.9	45.8	31.4	8.78	14.3	10.3	15.8	503	2.37	14.0
14-01	14.4	119.4	0.11	0.37	14.9	0.10	0.10	0.60	359	0.42	214	155	3.67	10.9	18.5	31.7	54.3	44.6	29.0	8.52	12.7	9.37	14.7	477	2.70	11.6
14-01	14.6	119.6	0.12	0.15	14.8	0.10	0.11	0.59	42.1	0.43	219	156	4.20	11.6	15.2	30.4	54.7	43.7	28.7	8.23	13.9	11.2	13.6	443	2.54	12.8
14-01	14.8	119.8	0.12	0.14	14.4	0.09	0.10	0.60	524	0.45	231	159	3.85	11.2	17.0	32.9	56.1	45.3	28.3	8.41	12.9	13.6	14.7	468	2.55	13.2
14-01	15	120	0.13	0.25	13.8	0.09	0.09	0.61	507	0.45	229	159	3.99	11.4	16.1	33.7	57.3	49.4	29.5	8.74	14.4	11.4	16.9	477	2.78	13.5
14-01	15.2	120.2	0.14	0.51	13.4	0.10	0.08	0.63	13.7	0.44	223	159	3.54	10.6	16.7	36.9	55.3	46.1	29.3	8.51	12.7	11.0	14.1	468	2.55	13.2
14-01	15.4	120.4	0.15	0.57	12.9	0.11	0.07	0.60	302	0.43	220	158	3.85	10.8	13.6	34.5	53.7	47.0	30.7	8.61	11.7	9.55	16.0	497	2.09	12.9
14-01	15.6	120.6	0.16	0.72	12.5	0.11	0.07	0.60	977	0.42	215	157	4.02	11.3	15.5	33.2	51.2	47.0	30.0	8.44	14.4	11.3	13.8	492	2.58	14.7
14-01	15.8	120.8	0.15	0.58	13.1	0.11	0.07	0.62	340	0.42	215	162	3.29	10.2	17.3	32.0	53.2	43.1	28.7	8.65	14.0	11.4	14.9	515	2.13	15.3
14-01	16	121	0.14	0.53	13.7	0.11	0.07	0.65	0.0	0.42	216	168	3.07	9.73	16.6	31.6	51.5	44.7	27.5	8.29	12.5	11.6	14.9	483	1.99	13.7
14-01	16.2	121.2	0.12	0.35	14.6	0.09	0.08	0.65	325	0.42	214	171	3.20	9.92	15.5	33.1	49.3	42.3	28.0	8.46	12.0	14.6	13.7	496	2.60	14.5
14-01	16.4	121.4	0.11	0.07	15.3	0.08	0.10	0.62	890	0.42	217	174	3.28	10.1	14.8	31.8	51.0	43.3	28.7	8.59	12.9	13.9	13.2	527	2.04	14.3
14-01	16.6	121.6	0.09	0.12	16.0	0.07	0.12	0.61	123	0.42	217	175	3.38	10.3	14.6	25.7	50.9	44.0	29.2	8.57	9.88	10.1	13.7	569	2.21	14.5
14-01	16.8	121.8	0.09	0.32	16.0	0.07	0.11	0.58	625	0.42	215	178	3.38	10.3	13.8	32.6	52.3	42.1	28.1	8.25	10.7	16.1	13.0	558	2.00	16.5
14-01	17	122	0.10	0.05	16.2	0.07	0.11	0.59	116	0.42	214	186	3.26	10.4	15.2	28.0	52.9	47.4	30.0	8.24	11.7	18.3	12.9	552	1.58	16.2
14-01	17.2	122.2	0.10	0	16.2	0.07	0.11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14-01	17.4	122.4	0.10	0.29	15.6	0.08	0.11	0.59	685	0.43	220	192	3.10	10.5	16.8	31.7	55.5	46.0	28.7	8.06	12.6	23.7	12.1	430	1.23	11.2
14-01	17.6	122.6	0.10	0.33	15.6	0.08	0.11	0.62	430	0.44	222	191	2.88	10.0	15.7	35.7	55.6	45.9	30.4	8.30	12.0	27.5	14.3	455	2.25	12.5
14-01	17.8	122.8	0.10	0	15.7	0.08	0.11	0.61	0.0	0.42	214	190	3.27	10.9	17.2	31.7	59.6	48.8	31.2	7.92	12.2	21.4	11.4	422	2.11	10.2
14-01	18	123	0.10	0.07	15.5	0.08	0.10	0.58	0.0	0.42	215	192	3.47	11.1	15.0	32.0	60.3	53.3	33.2	8.35	13.1	18.8	10.7	371	0.81	8.46
14-01	18.2	123.2	0.10	0.28	15.4	0.09	0.09	0.57	164	0.41	212	193	3.39	11.3	17.7	33.1	60.1	55.9	34.8	8.22	13.4	14.4	9.29	339	1.22	9.64
14-01	18.4	123.4	0.10	0	15.4	0.09	0.08	0.59	1/6	0.44	228	198	3.02	10.4	15.9	36.3	59.6	59.0	34.8	8.49	13.9	11.2	9.10	258	2.57	3.65
14-01	18.6	123.6	0.10	0	15.7	0.09	0.09	0.60	258	0.42	217	197	3.00	10.7	18.4	35.7	58.4	59.0	35.9	8.42	13.7	7.96	11.0	300	2.37	3.46
14-01	18.8	123.8	80.0	0.04	16.0	0.09	0.09	0.61	625	0.42	216	196	3.14	10.8	17.2	35.2	57.5	55.2	33.0	8.33	15.5	7.75	9.89	323	1.71	1.79
14-01	19	124	0.09	0.11	15.6	0.09	0.10	0.60	605	0.44	226	204	2.98	10.5	16.4	34.5	61.1	50.9	31.9	8.58	14.2	7.92	9.96	353	2.22	6.92
14-01	19.2	124.2	0.11	0.11	14.9	0.10	0.09	0.60	358	0.45	229	205	2.69	10.1	17.0	33.3	60.3	52.9	32.6	8.53	13.0	8.21	9.99	326	2.52	7.11
14-01	19.4	124.4	0.11	0.09	15.2	0.10	0.09	0.62	345	0.44	228	207	2.94	10.5	16.6	37.8	64.0	50.1	31.6	8.11	13.1	8.18	11.1	320	2.53	8.27
14-01	19.6	124.6	0.13	0.13	14.3	0.10	0.08	0.61	88.9	0.45	232	194	2.85	9.85	15.8	38.3	59.7	46.3	30.3	8.29	13.2	7.90	11.8	342	1.77	9.88
14-01	19.8	124.8	0.14	0.45	13.2	0.10	0.06	0.60	120	0.44	220	199	2.99	10.4	17.0	34.1	01.3	49.8 49.0	31.2	0.3U	14.5	5.65	10.2	406	2.15	11.7
14-01	20	125	0.15	0.49	12.8	0.10	0.07	0.57	95.0	0.45	232	194	∠.ŏU	9.90	10.9	44.0 20.2	02.9 62.5	48.9	32.9 21.2	0./0	12.0	0.11 7.25	0.20	517	∠.5ŏ	14.3
14-01	20.2	125.2	0.10	0.59	12.4	0.10	0.07	0.50 0.57	197	0.47	241	192	১.∠⊺ ১.০০	10.0	10.2	39.∠ 42.0	03.5 74.0	5U.3	31.3 200	0.00	13.9	1.35	9.30	547 570	2.0Z	10.0
14-01	20.4	125.4	0.10	0.59	12.7	0.10	0.00	0.57	192	0.47	240	208	2.98	10.9	19.2	43.0	/ 1.Z	47.1	20.2	9.34	14.3	1.13	12.1	51Z	2.15	19.1
14-01	20.6	125.6	0.18	0.87	12.3	0.12	0.05	0.55	U	0.46	231	220	2.96	10.4	14.4	39.9	68.8	46.4	30.6	ö.57	13.5	8.35	11.8	477	2.33	17.3

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	As (ppm)	(mqq) dq	Se (ppm)	(mqq) dЯ	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mdd) oM
14-01	20.8	125.8	0.21	1.11	12.4	0.12	0.04	0.51	331	0.43	222	213	2.26	9.12	15.0	36.9	59.5	41.8	28.4	8.60	10.9	8.25	10.9	465	2.52	15.2
14-01	21	126	0.22	1.20	12.1	0.12	0.04	0.49	0	0.39	199	190	2.16	8.37	13.8	34.9	53.7	35.3	25.2	8.56	9.60	4.50	7.88	412	2.20	12.5
14-01	21.2	126.2	0.19	0.98	12.6	0.12	0.04	0.50	0	0.37	191	186	2.22	8.16	12.4	34.0	53.6	33.1	23.6	8.29	9.69	4.63	6.87	444	2.03	12.8
14-01	21.4	126.4	0.15	0.41	14.2	0.11	0.09	0.46	0	0.38	192	172	2.16	7.36	9.56	27.5	46.9	26.6	20.7	7.87	8.92	0	4.11	442	1.19	14.4
14-01	21.6	126.6	0.12	0.34	16.4	0.11	0.15	0.43	187	0.34	166	153	1.59	5.63	7.58	17.0	35.7	16.4	15.7	7.67	3.88	0	0.60	340	0	10.6
14-01	21.8	126.8	0.13	0.21	17.2	0.11	0.13	0.37	253	0.28	140	140	1.20	4.05	3.81	12.8	24.2	11.2	11.5	7.69	0	0	1.40	270	0	6.69
14-01	22	127	0.17	0.67	15.6	0.11	0.08	0.39	86	0.27	138	152	0.93	4.12	3.96	16.7	32.1	12.8	12.4	7.55	2.65	0	8.94	450	0.32	12.8
14-01	22.2	127.2	0.22	1.15	12.3	0.12	0	0.46	156	0.33	172	201	1.39	7.07	9.28	30.8	51.5	26.4	20.5	8.24	4.44	0.72	23.6	703	1.79	21.0
14-01	22.4	127.4	0.24	1.36	9.76	0.11	0	0.57	344	0.47	245	228	2.10	10.7	18.09	51.3	86.6	39.9	26.4	8.91	9.11	10.4	36.2	1008	3.26	24.1
14-01	22.6	127.6	0.23	1.27	10.0	0.09	0.03	0.70	237	0.49	252	245	1.47	11.1	21.2	60.6	124.4	46.3	30.3	9.25	13.8	41.5	29.6	929	3.58	23.2
14-01	22.8	127.8	0.22	1.27	10.1	0.07	0.07	0.77	685	0.49	254	214	0.72	11.3	24.6	57.2	141.2	44.1	30.4	9.54	14.8	51.2	28.1	1020	3.38	23.6
14-01	23	128	0.21	1.05	10.6	0.06	0.06	0.75	702	0.47	243	195	0.00	10.4	23.2	61.5	136.3	42.8	27.9	8.78	14.1	58.0	29.9	973	4.41	22.9
14-01	23.2	128.2	0.21	1.15	10.5	0.05	0.03	0.73	137	0.49	239	274	1.31	10.7	20.4	62.0	130.3	46.6	31.0	9.38	15.3	56.3	31.5	973	3.35	23.6
14-01	23.4	128.4	0.18	0.69	11.5	0.03	0.03	0.71	430	0.46	226	254	2.42	10.6	18.4	52.1	102.2	48.3	30.8	9.45	13.2	57.5	26.8	780	3.53	20.4
14-01	23.6	128.6	0.20	1.00	10.22	0.04	0.02	0.71	140	0.43	213	210	2.47	9.63	19.0	39.2	76.4	42.2	30.5	9.53	13.3	57.3	17.6	520	3.49	17.2
14-01	23.8	128.8	0.23	1.39	9.02	0.05	0.01	0.72	0	0.42	212	192	2.50	8.78	13.7	45.0	61.9	41.2	27.0	9.38	11.1	49.4	17.4	425	3.79	13.1
14-01	24	129	0.29	1.97	6.67	0.07	0	0.72	0	0.38	194	180	1.98	7.14	9.6	41.0	59.6	36.6	25.4	8.60	10.9	46.9	14.0	397	2.40	11.8
14-02	0	137	0.54	4.82	0.00	0.18	0	0.05	0	0.29	147	193	1.35	8.41	30.8	239	101.2	29.8	24.6	11.9	11.5	8.48	15.0	500	2.61	19.7
14-02	0.2	137.2	0.62	4.40	0.11	0.18	0	0.13	85.3	0.43	215	225	0.60	6.75	20.7	157	82.2	48.1	32.0	11.2	12.0	10.0	15.0	892	3.57	25.1
14-02	0.4	137.4	0.57	3.39	2.76	0.16	0	0.30	291	0.47	235	240	1.27	8.06	19.5	101	79.1	50.2	30.6	10.2	15.5	8.55	18.4	888	3.96	22.9
14-02	0.6	137.6	0.46	2.10	6.18	0.15	0	0.52	49	0.47	240	241	1.53	8.41	18.0	63.2	72.2	62.5	37.5	9.71	15.4	8.17	23.8	726	3.81	22.2
14-02	0.8	137.8	0.31	1.48	8.23	0.15	0.01	0.65	316	0.49	253	238	2.05	9.49	18.4	55.9	74.4	58.1	35.2	8.80	18.3	7.77	23.6	784	4.66	22.4
14-02	1	138	0.25	1.04	9.87	0.14	0.03	0.72	0	0.52	272	241	1.86	9.70	22.4	47.7	79.4	68.3	38.5	9.81	18.0	10.2	26.8	945	3.36	23.8
14-02	1.2	138.2	0.20	0.79	11.1	0.14	0.04	0.81	0	0.52	273	242	2.23	10.8	24.9	50.4	75.8	63.8	35.9	9.31	16.4	9.95	25.5	1354	3.37	21.9
14-02	1.4	138.4	0.17	0	0	0	0	0	220	0.51	256	256	2.26	10.2	19.5	47.6	75.3	59.2	35.9	9.10	20.1	12.6	27.9	1573	4.44	14.2
14-02	1.6	138.6	0	0.73	12.5	0.13	0.05	1.05	555	0.50	259	242	2.50	10.4	17.7	43.5	62.6	57.1	35.0	9.26	15.2	10.4	24.2	1557	3.39	15.7
14-02	1.8	138.8	0.17	0.57	13.6	0.12	0.05	1.26	341	0.45	231	226	2.25	9.79	17.2	36.0	63.2	48.0	31.1	8.77	14.0	13.9	22.1	1177	3.51	24.1
14-02	2	139	0.17	0.63	15.0	0.13	0.06	1.41	702	0.41	210	211	2.21	9.36	15.4	29.5	53.8	46.6	29.7	8.56	14.4	11.6	20.4	1088	3.44	23.6
14-02	2.2	139.2	0.17	0.48	16.2	0.12	0.08	1.48	963	0.46	235	221	2.21	9.19	14.3	34.3	60.8	48.3	30.7	9.61	16.2	12.8	23.1	1411	3.52	0
14-02	2.4	139.4	0.15	0.45	15.9	0.12	0.10	1.39	181	0.51	263	238	2.75	11.0	19.8	46.0	/1.0	56.3	34.0	9.35	16.8	14.1	27.3	1807	3.73	6.47
14-02	2.6	139.6	0.13	0.29	15.8	0.12	0.09	1.24	174	0.54	261	263	2.75	11.0	20.1	46.2	72.9	62.7	37.9	9.68	20.0	12.0	28.2	2005	3.84	0
14-02	2.8	139.8	0.13	0.27	14.8	0.12	0.08	1.22	423	0.53	270	247	2.64	10.4	18.5	44.6	69.6	61.3	34.8	9.56	16.4	12.5	27.0	1794	3.60	7.26
14-02	3	140	0.13	0.34	13.8	0.13	0.07	1.21	465	0.51	247	239	3.00	10.7	16.9	36.9	69.0	55.4	33.3	9.10	14.4	12.8	27.8	1686	3.83	11.3
14-02	3.2	140.2	0.14	0.46	13.1	0.13	0.05	1.11	422	0.52	266	234	3.17	11.6	20.6	42.6	/1.1	57.6	34.5	9.23	15.3	13.4	28.1	15/8	4.05	14.3
14-02	3.4	140.4	0.15	0.39	13.0	0.13	0.05	1.02	517	0.54	277	250	3.27	11.9	20.3	46.3	/6.4	61.8	36.0	8.97	16.5	10.3	27.8	1547	3.96	15.6
14-02	3.6	140.6	0.15	0.63	12.7	0.12	0.05	0.98	452	0.56	289	250	3.42	11.9	17.9	48.4	/4.8	61.5	36.5	9.69	17.0	10.8	26.0	1628	3.99	13.2
14-02	3.8	140.8	0.16	0.76	12.6	0.13	0.04	0.97	442	0.53	279	230	2.76	10.4	17.9	48.6	69.4	58.4	33.8	9.38	15.9	9.76	25.4	1628	2.94	14.5
14-02	4	141	0.17	0.63	12.5	0.13	0.04	1.04	427	0.51	267	217	2.84	10.4	18.3	45.8	64.0	52.3	33.1	9.50	15.2	8.63	24.7	1619	3.01	14.4

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	As (ppm)	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mqq) oM
14-02	4.2	141.2	0.18	0.57	12.9	0.12	0.05	1.04	1061	0.46	233	212	2.75	10.1	17.7	36.9	62.9	50.9	31.9	9.00	13.3	8.61	23.0	1349	3.82	20.2
14-02	4.4	141.4	0.17	0.58	12.8	0.12	0.04	0.98	503	0.43	0	206	2.40	9.24	16.0	31.1	62.8	43.2	29.8	8.68	11.7	9.56	20.9	1019	2.52	23.2
14-02	4.6	141.6	0.17	0.80	12.1	0.13	0.04	0.76	513	0.41	211	203	2.39	9.23	15.7	30.8	66.7	43.0	27.5	8.46	10.5	6.66	22.7	706	3.00	19.2
14-02	4.8	141.8	0.19	0.79	11.7	0.12	0.02	0.62	0	0.40	203	196	2.00	8.46	14.6	28.5	83.7	38.5	26.2	8.77	9.76	4.10	23.0	566	2.71	16.7
14-03	0	169	0.49	3.78	1.25	1.25	0	0.26	877	0.39	193	210	1.06	6.80	16.7	85.7	63.3	41.1	31.6	10.28	11.3	4.29	18.4	912	3.17	24.2
14-03	0.2	169.2	0.46	3.49	2.24	2.24	0	0.35	853	0.40	202	207	1.13	7.03	17.0	70.5	58.1	37.4	28.3	9.67	14.3	4.63	19.5	701	4.30	24.5
14-03	0.4	169.4	0.34	2.45	6.46	6.46	0	0.75	625	0.50	262	221	1.01	7.02	15.9	53.4	62.9	49.2	33.1	9.51	14.5	8.85	22.0	767	4.52	24.9
14-03	0.6	169.6	0.23	1.22	10.9	10.9	0.04	1.08	42.7	0.55	282	258	0.84	6.94	16.7	54.4	64.1	54.3	35.5	9.40	16.5	16.3	22.0	971	4.20	24.5
14-03	0.8	169.8	0.16	0.48	13.3	13.3	0.07	1.15	32.1	0.59	302	277	0.81	7.31	19.6	46.2	77.7	62.3	38.3	9.76	16.2	16.7	25.3	1364	3.62	19.9
14-03	1	170	0.13	0.31	14.0	14.0	0.08	1.10	259	0.63	320	311	1.06	8.39	22.7	57.3	84.3	69.6	40.6	10.54	18.4	18.3	28.5	1837	4.45	3.84
14-03	1.2	170.2	0.14	0.42	13.5	13.5	0.06	1.02	109	0.62	283	324	1.13	8.81	24.4	58.6	87.5	78.5	43.7	10.24	20.8	16.0	31.0	2247	5.06	0.00
14-03	1.4	170.4	0.15	0.41	14.0	14.0	0.07	1.19	296	0.58	287	290	0.93	8.07	22.5	57.2	82.0	70.0	39.9	10.20	22.0	14.1	29.4	2156	5.22	0.00
14-03	1.6	170.6	0.16	0.55	14.8	14.8	0.08	1.32	0	0.51	262	246	1.18	8.08	20.5	48.2	72.6	61.5	37.0	9.64	19.0	17.8	28.2	1811	3.91	5.89
14-03	1.8	170.8	0.17	0.67	14.9	14.9	0.08	1.22	234	0.46	239	219	1.22	7.42	16.5	38.2	63.4	55.7	35.2	9.68	15.4	15.5	25.0	1285	4.35	20.8
14-03	2	171	0.16	0.55	14.8	14.8	0.05	0.90	0.0	0.43	219	203	1.08	6.60	13.8	37.6	56.3	49.7	33.6	9.10	14.7	13.4	19.3	900	4.44	21.7
14-03	2.2	171.2	0.20	0.82	12.3	12.3	0.02	0.58	2.3	0.43	221	224	0.92	7.40	15.7	37.4	58.5	44.9	30.2	8.52	15.6	11.9	17.8	761	3.02	20.5
14-03	2.4	171.4	0.17	0.52	12.8	12.8	0.03	0.52	222	0.44	228	237	0.35	6.74	15.3	38.2	62.5	46.6	29.8	9.04	14.9	9.59	20.5	848	3.91	22.9
14-03	2.6	171.6	0.17	0.69	12.4	12.4	0.04	0.54	426	0.45	231	239	0.68	7.40	16.7	48.1	79.3	46.1	30.5	9.18	13.7	11.8	22.3	881	4.22	22.0
14-03	2.8	171.8	0.16	0.53	12.5	12.5	0.03	0.62	0.0	0.45	233	238	0.31	6.71	19.2	43.0	74.5	47.7	30.5	9.19	14.0	30.1	23.2	903	3.12	22.9
14-03	3	172	0.15	0.49	13.1	13.1	0.03	0.64	0.0	0.46	230	238	0.73	6.81	19.1	43.1	74.4	41.7	30.3	8.83	12.0	48.1	24.4	800	3.46	22.5
14-03	3.2	172.2	0.14	0.25	13.6	13.6	0.03	0.60	0.0	0.47	233	232	0.64	6.14	16.5	37.3	64.5	44.5	31.0	8.88	11.5	43.9	24.4	584	3.28	17.6
14-03	3.4	172.4	0.14	0.33	13.9	13.9	0.03	0.47	395.6	0.45	225	221	1.50	7.23	12.6	34.0	51.7	36.9	25.9	8.60	12.7	25.1	19.1	503	3.39	14.3
14-03	3.6	172.6	0.13	0.28	14.4	14.4	0.03	0.44	254.8	0.45	228	219	1.45	7.13	13.0	34.4	50.4	35.7	25.4	8.83	9.05	1.95	21.0	397	3.72	10.1
14-03	3.8	172.8	0.09	0.00	17.2	17.2	0.04	0.51	22.0	0.40	203	189	1.44	6.94	12.3	28.4	45.2	31.7	24.5	8.26	6.84	0	18.2	370	2.44	11.5
14-03	4	173	0.07	0.00	18.7	18.7	0.04	0.55	287	0.36	176	175	1.22	6.08	10.1	17.4	33.8	26.2	21.1	7.33	5.38	0	14.0	268	3.07	5.49
14-03	4.2	173.2	0.07	0.00	19.4	19.4	0.02	0.54	686	0.32	154	174	0.95	5.73	8.21	23.1	30.9	26.5	21.0	8.02	4.77	0	12.9	213	2.00	3.97
14-03	4.4	173.4	0.14	0.38	17.4	17.4	0	0.46	0	0.26	126	167	0.05	3.92	5.01	19.3	23.3	21.5	18.9	7.27	1.60	0	8.50	186	0.87	2.45
14-03	4.6	173.6	0.30	1.86	11.4	11.4	0	0.30	55.7	0.22	120	175	0.04	4.22	4.25	22.4	20.0	20.0	16.5	8.05	1.06	0	7.74	143	1.32	0.41
14-03	4.8	173.8	0.33	2.15	11.2	11.2	0	0.28	0.0	0.20	110	160	0.47	4.25	5.98	22.9	18.6	15.4	12.4	7.75	2.37	0	4.09	92	1.74	1.31
14-03	5	174	0.25	1.45	13.1	13.1	0	0.37	48.9	0.22	120	161	0.64	4.04	2.03	20.4	18.3	18.3	17.6	8.23	4.88	0	7.86	113	2.88	1.28
14-03	5.2	174.2	0.20	1.05	14.1	14.1	0	0.46	411	0.31	153	173	0.73	4.78	5.19	24.5	30.8	25.5	19.6	7.73	5.34	0	13.7	163	2.93	3.86
14-03	5.4	174.4	0.18	0.70	13.4	13.4	0	0.53	56.3	0.37	186	193	1.08	6.53	11.8	31.6	39.9	33.8	27.0	7.84	4.43	0	15.3	255	1.81	8.61
14-03	5.6	174.6	0.21	1.17	11.6	11.6	0	0.55	68.7	0.40	202	206	0.80	6.62	14.8	31.9	47.0	40.6	26.9	8.61	6.11	0	14.9	279	3.02	7.37
14-03	5.8	174.8	0.24	1.23	12.2	12.2	0	0.53	230	0.36	177	186	0.99	6.22	11.2	32.2	39.6	31.3	24.7	8.10	5.12	0	13.5	247	1.51	4.06
14-03	6	175	0.26	1.45	12.7	12.7	0	0.43	43.8	0.33	166	174	1.29	6.01	8.82	23.1	31.9	26.9	21.7	7.99	4.11	0	12.1	229	1.94	4.99
14-03	6.2	175.2	0.25	1.34	13.0	13.0	0	0.40	0	0.31	157	166	0.68	4.43	7.91	18.5	24.4	24.5	20.6	8.25	4.10	0	11.2	222	3.00	4.44
14-03	6.4	175.4	0.26	1.55	11.3	11.3	0	0.35	0	0.34	175	175	0.99	4.96	6.53	18.5	28.9	27.0	21.5	8.60	4.43	0	11.1	235	3.03	5.84
14-03	6.6	175.6	0.26	1.64	10.3	10.3	0	0.36	145	0.34	175	184	1.00	5.57	8.69	30.5	34.7	30.7	22.5	7.78	4.00	0	14.0	245	2.09	7.86

core	core length (cm)	depth relatie to datum (cm)	Na (wt%)	AI (wt%)	Si (wt%)	P (wt%)	K (wt%)	Ca (wt%)	Ba (ppm)	Ti (wt%)	Cr (ppm)	(mqq) nM	Fe (%)	Co (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	(mqq) sA	(mqq) dq	Se (ppm)	Rb (ppm)	Sr (ppm)	Y (ppm)	Zr (ppm)	(mqq) dN	(mqq) oM
14-03	6.8	175.8	0.27	1.66	8.55	8.55	0	0.35	0	0.35	181	199	1.01	6.63	14.5	33.0	41.0	30.6	24.2	7.87	3.57	0	15.4	223	2.12	5.42
14-03	7	176	0.27	1.69	8.35	8.35	0	0.36	0	0.36	184	197	1.11	7.08	14.4	34.3	45.3	41.5	28.6	8.15	6.34	0	17.0	178	2.41	3.27
14-03	7.2	176.2	0	0	0	0	0	0	232	0.38	195	207	1.49	8.00	15.8	33.5	55.2	40.4	26.1	8.41	6.21	0	18.8	181	3.46	4.21
14-03	7.4	176.4	0.25	1.52	9.13	9.13	0	0.36	0	0.37	192	203	1.37	7.13	12.2	36.0	49.2	40.0	28.6	8.69	8.24	0.55	19.1	207	3.20	3.86
14-03	7.6	176.6	0.26	1.55	8.83	8.83	0	0.36	391	0.37	192	207	1.18	7.36	15.1	33.7	45.8	40.9	26.1	8.58	5.75	0	16.5	196	2.01	4.43
14-03	7.8	176.8	0.24	1.41	9.44	9.44	0.01	0.40	140	0.38	192	200	1.03	6.95	13.4	38.5	47.0	39.5	29.2	8.39	8.42	0	19.7	174	3.43	1.46
14-03	8	177	0.24	1.38	9.52	9.52	0	0.43	0	0.40	204	223	1.60	8.52	16.4	35.3	52.4	41.3	29.1	8.48	8.51	0	20.5	181	3.21	4.94
14-03	8.2	177.2	0.23	1.16	9.65	9.65	0.01	0.43	164	0.41	207	224	1.75	8.86	16.9	37.6	50.1	40.4	29.1	8.61	10.2	1.78	16.6	190	2.91	4.96
14-03	8.4	177.4	0.23	1.23	9.61	9.61	0.01	0.42	267	0.42	215	217	2.10	9.23	16.8	39.3	51.6	44.5	28.5	8.59	9.61	4.73	19.1	191	3.64	4.79
14-03	8.6	177.6	0	0	0	0	0	0	171	0.40	203	201	1.61	7.71	12.5	38.0	52.0	41.5	28.2	8.46	7.40	0	18.0	179	2.87	4.08
14-03	8.8	177.8	0.23	1.20	9.72	9.72	0.01	0.43	0	0.39	194	191	1.16	6.80	13.8	31.9	43.4	40.1	26.6	8.23	6.53	0	18.5	168	2.64	0.71
14-03	9	178	0.23	1.36	9.96	9.96	0	0.42	90	0.36	175	181	1.48	7.10	12.4	27.8	40.4	35.0	22.6	8.27	7.67	0	14.8	175	2.14	1.56
14-03	9.2	178.2	0.24	1.38	9.88	9.88	0	0.40	0	0.35	173	183	1.58	7.25	11.4	27.8	40.7	36.6	26.1	8.01	5.98	0	13.7	211	2.40	2.44
14-03	9.4	178.4	0.25	1.45	10.0	10.0	0	0.42	0	0.33	168	182	1.09	6.42	12.3	24.0	36.8	31.8	24.2	8.23	3.99	0	16.3	261	1.83	4.87
14-03	9.6	178.6	0.25	1.43	10.1	10.1	0	0.39	436	0.33	168	189	0.84	6.01	9.7	30.6	38.5	34.6	27.2	7.67	5.72	0	13.3	245	2.22	5.54
14-03	9.8	178.8	0.27	1.61	11.0	11.0	0	0.38	113	0.33	161	182	0.75	5.91	10.7	25.3	37.7	31.9	24.2	8.01	5.62	0	12.4	256	1.63	4.78
14-03	10	179	0.27	1.65	12.0	12.0	0	0.39	0	0.33	166	184	0.71	5.76	10.2	25.3	31.1	28.6	21.8	8.12	6.19	0	13.4	299	2.17	10.8
14-03	10.2	179.2	0.26	1.53	14.0	14.0	0	0.48	480	0.34	169	180	0.88	5.70	9.29	19.3	29.8	30.5	22.4	8.09	6.11	2.34	12.5	350	2.74	9.96
14-03	10.4	179.4	0.24	1.25	15.2	15.2	0.01	0.53	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14-03	10.6	179.6	0.24	1.33	13.5	13.5	0.02	0.54	0	0.37	190	191	1.19	6.29	9.28	27.2	41.1	36.3	24.4	8.72	11.2	4.10	12.1	545	2.87	19.7
14-03	10.8	179.8	0.23	1.35	12.6	12.6	0.03	0.54	383	0.39	199	195	0.89	5.72	8.75	25.5	41.1	37.7	26.8	8.39	11.9	4.09	12.0	551	3.29	17.0
14-03	11	180	0.23	1.19	12.7	12.7	0.02	0.52	351	0.36	186	195	0.77	5.90	11.6	28.2	38.5	35.7	24.6	9.08	11.7	5.27	11.4	549	2.86	16.7
14-03	11.2	180.2	0.24	1.38	12.8	12.8	0.01	0.53	487	0.33	166	179	0.86	5.36	7.66	29.9	39.3	29.8	22.2	8.69	8.53	2.79	11.9	460	2.45	17.9
14-03	11.4	180.4	0.24	1.26	12.4	12.4	0.02	0.53	53.6	0.33	167	173	1.20	5.83	9.52	27.7	37.6	29.9	23.1	8.81	8.56	3.65	1E+12	540	2.83	16.8
14-03	11.6	180.6	0.21	1.19	11.7	11.7	0.03	0.51	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14-03	11.8	180.8	0.19	0.79	11.3	11.3	0.03	0.47	0.0	0.45	232	192	1.33	6.63	14.1	31.4	55.4	43.9	28.3	8.45	12.3	29.5	17.8	672	4.10	20.5
14-03	12	181	0.18	0.66	11.5	11.5	0.03	0.46	5.6	0.46	235	201	1.14	6.44	14.1	40.8	53.7	45.1	30.2	9.44	11.7	51.3	16.8	800	3.20	22.3
14-03	12.2	181.2	0.19	0.94	11.3	11.3	0.03	0.45	437	0.45	233	206	0.85	6.07	13.8	40.7	54.1	50.5	30.6	8.96	11.1	47.2	18.9	721	2.37	20.9
14-03	12.4	181.4	0	0	0	0	0	0	163	0.45	234	210	1.24	7.08	15.2	34.9	53.2	45.4	30.3	8.30	12.6	33.8	16.9	690	2.50	19.6
14-03	12.6	181.6	0.15	0.35	12.9	12.9	0.04	0.47	212	0.46	236	202	0.89	6.24	14.4	38.1	51.8	40.8	28.3	8.65	14.0	10.1	17.4	576	3.41	18.1
14-03	12.8	181.8	0.15	0.56	13.2	13.2	0.05	0.46	248	0.44	226	193	1.15	6.38	13.1	33.1	45.9	37.7	25.2	8.45	13.3	7.29	13.6	515	2.99	16.2
14-03	13	182	0.14	0.33	13.9	13.9	0.04	0.49	0	0.44	222	190	1.09	6.17	12.5	36.0	46.4	37.6	27.0	8.78	14.3	5.16	14.7	516	3.47	14.7
14-03	13.2	182.2	0.15	0.34	13.6	13.6	0.04	0.50	392	0.43	219	192	1.25	6.63	13.7	31.9	48.8	39.0	26.9	8.66	11.1	6.52	15.3	527	2.79	16.5
14-03	13.4	182.4	0.20	1.05	11.3	11.3	0.00	0.47	0.0	0.44	222	201	1.16	6.70	13.6	38.0	55.8	44.4	30.6	9.17	11.2	8.64	15.7	585	3.38	19.4













Appendix XII: All quantified XRF data displayed relative to depth of the collected cores from TAL-16-04

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Appendix XIII: README for the GitHub repository dedicated to all coding scripts used for this project

Repository location: https://github.com/ardowney/Masters thesis

Physical and chemical properties of clastic cave sediment from the northern karst region of Puerto Rico

R scripts include:

```
Organizing raw particle size data from a Beckman Coulter single wavelength LS13-320 particle size analyzer from the instrument into a csv format
```

post-processing fluorescence spectroscopy data from a Horiba Fluorolog-3

PARAFAC modeling of EEM data

PCA analysis

Pearson Correlation Matrix

Prerequisites:

Communication with Unix Shell

Necessary R packages:

```
- Particle size
install.packages("tidyverse")
install.packages("dplyr")
install.packages("readxl")
install.packages("ggplot2")
install.packages("colorspace")
install.packages("scales")
. . .
- Fluorescence data
install.packages("eemR")
install.packages("ggplot2")
install.packages("staRdom")
install.packages("dbplyr")
install.packages("dplyr")
- PARAFAC model
install.packages(dplyr)
install.packages(multiway)
```

```
install.packages(staRdom)
install.packages(eemR)
install.packages(parallel)
install.packages(devtools)
install.packages(tidyverse)
install.packages(stats)
install.packages(pracma)
install.packages(tidyr)
install.packages(gtools)
install.packages(tibble)
install.packages(stringr)
install.packages(grDevices)
install.packages(plotly)
install.packages(GGally)
- PCA
install.packages(readxl)
install.packages("devtools")
install.packages("factoextra")
· · · ·
- Pearson Correltion Matrix
install.packages(readxl)
install.packages("Hmisc")
install.packages("devtools")
install.packages("factoextra")
install.packages(corrplot)
install.packages(rstatix)
install.packages(ggplot2)
install.packages(readxl)
install.packages(gridExtra)
install.packages(ggpubr)
install.packages(ggpmisc)
```

script name and function:

```
particle size:
- cut_loop.sh - Unix shell script organizing raw data from the
instrument. transforms proprietary file extension to flat text
file (.txt) and sends to a new organized directory.
```

- Particle size.R - Applies particle size bins from the instrument to the organized data created by cut loop.sh.

- Averaging script.R - Takes the average of the three replicates ran for each subsample. Fluorescence data: - Scatter code.R - Plots EEM data stored as a csv. Removes 15 nm 1st and 2nd order Rayleigh scattering. Smooths the EEM by interpreting intensity through the Rayleigh regions. Sends corrected files to a modified csv file. - Percent function.R - Normalizes corrected EEM to the max intensity of each scan. PARAFAC model: - Model functions.R - Functions needed to run the PARAFAC model - PARAFAC code.R - Script for creating 2-7 component models using user defined EEM data. PCA: - PCA.R - Creates PCA output for user defined data. Outputs include biplots, variance plots, eigenvalues, and standard PCA scatter plots. Pearson Correlation Matrix: - Cor matrix function.R - Creates correlation matrix with R and p values exported to a csv file. R values and p values plots are included in the code.