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Stiles, William; Rowe, Edwin; Dennis, Peter

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Keywords: Nitrogen deposition; soil carbon; carbon fluxes; pollution; co-limitation; P limitation

Corresponding Author: Dr. William A. V. Stiles,

Corresponding Author's Institution: Aberystwyth University

First Author: William A. V. Stiles

Order of Authors: William A. V. Stiles; Edwin C Rowe; Peter Dennis

Abstract: Reactive nitrogen (N) deposition can affect many ecosystem processes, particularly in oligotrophic habitats, and is expected to affect soil C storage potential through increases in microbial decomposition rate as a consequence of greater N availability. Increased N availability may also result in changes in the principal limitations on ecosystem productivity. Phosphorus (P) limitation may constrain productivity in instances of high N deposition, yet ecosystem responses to P availability are poorly understood. This study investigated CO₂ and CH₄ flux responses to N and P enrichment using both short- (1 year) and long-term (16 year) nutrient addition experiments. We hypothesised that the addition of either N or P will increase CO₂ and CH₄ fluxes, since both plant production and microbial activity are likely to increase with alleviation from nutrient limitation. This study demonstrated the modification of C fluxes from N and P enrichment, with differing results subject to the duration of nutrient addition. On average, relative to control, the addition of N alone inhibited CO₂ flux in the short-term (-9%) but considerably increased CO₂ emissions in the long-term (+35%), reduced CH₄ uptake in the short term (-90%) and reduced CH₄ emission in the long term (-94%). Phosphorus addition increased CO₂ and CH₄ emission in the short term (+20% and +184% respectively), with diminishing effect into the long term, suggesting microbial communities at these sites are P limited. Whilst a full C exchange budget was not examined in the experiment, the potential for soil C storage loss with long-term nutrient enrichment is demonstrated and indicates that P addition, where P is a limiting factor, may have an adverse influence on upland soil C content.

Response to Reviewers: Reviewer #1: Comments on "Nitrogen and phosphorus enrichment effects on CO₂ and methane fluxes from an upland ecosystem"

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senior authors spend time on the representation if the leading author is reluctant for revision.

>> We thank the reviewer for taking the time to review this manuscript and are grateful for their helpful comments. We have attempted to refine the text throughout in line with the above suggestion.

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>> Entire first sentence now deleted.

3) Comparison with other studies will enhance the impact of this paper; it is same for the model implications.

>> Comparisons with the findings of other studies are drawn frequently in the discussion section of this paper - for instance, line 324/325, line 345/346, line 370, line 387, line 414/415 and others. With regard to 'model implications', as we stated in the first round of comments: We consider that the study contributes to scientific understanding and as such the results and inferences will be useful for testing and design of ecosystem models. However, reviewing the potential applications of the insights gained within specific models is beyond the scope of this paper.

4) Figure 1, a & b are missing.

>> Figure 1, parts a) and b) are in reference to the nutrient addition treatment squares (bottom right of the figure). We have also added a) and b) labels to the map to help the reader distinguish between trial plots.

5) Line 331, "sizable" not professional phrase.

>> All use of the word 'sizeable' deleted or replaced with significant, where appropriate. Similar subjective phrases were identified in other places in the text (e.g. markedly, considerable, slightly) and have been substituted with statistically validated, objective terms. We hope that this has therefore addressed the concern raised by the reviewer in the opening comment since the example given helped us to appreciate where concern lay in the presentation of the work in the earlier version of the manuscript.

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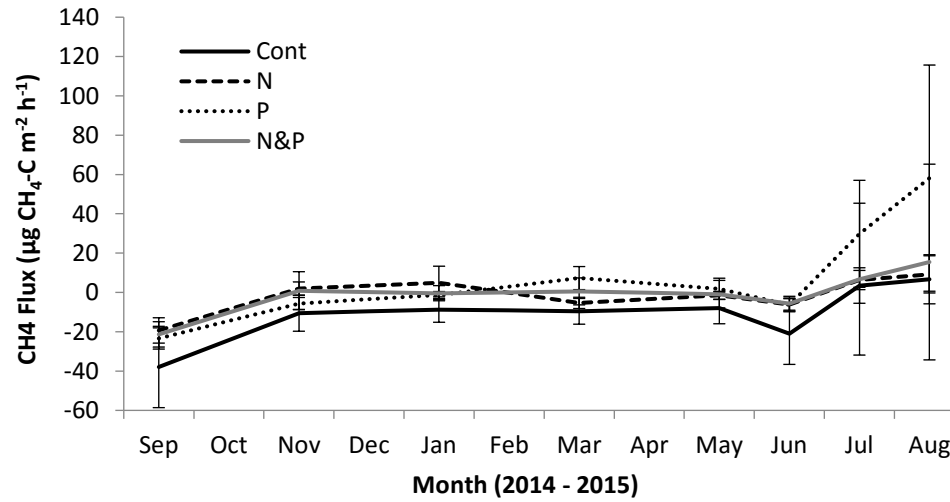
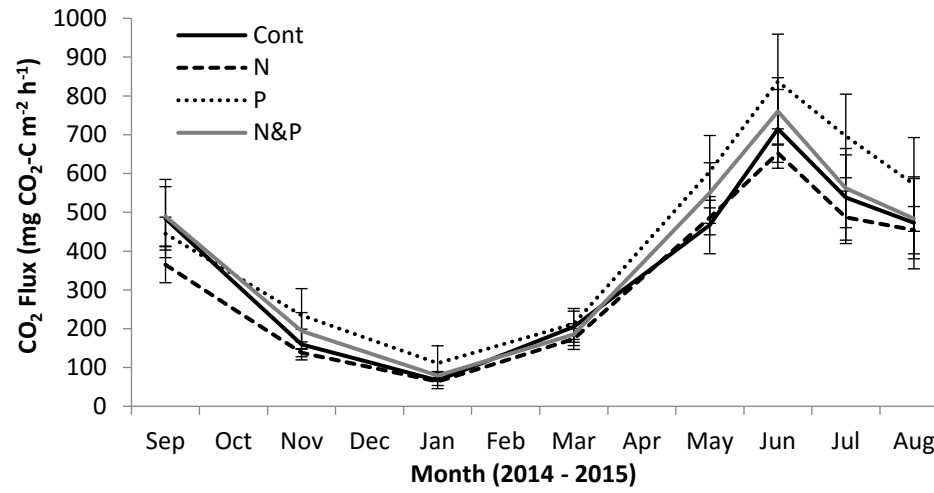
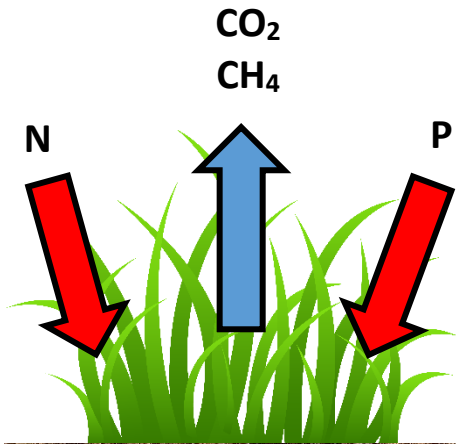
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- >> Paragraph added to the discussion as requested.

Graphical abstract:



Highlights:

- Under simulated N pollution, P addition increased CO₂ and CH₄ emissions
- This has contributed to lower soil C concentrations where P was added
- N addition initially inhibited CO₂ emission but increased emission in the long term
- Nutrient availability can significantly modify C emissions from upland soil
- P limitation may reduce potential soil C loss in situations of chronic N pollution

1 **Nitrogen and phosphorus enrichment effects on CO₂ and methane fluxes**
2 **from an upland ecosystem**

William A. V. Stiles ^{a,b}, Edwin C. Rowe ^b, Peter Dennis ^a

*^aInstitute of Biological, Environmental and Rural Sciences, Penglais Campus, Aberystwyth
University, Wales, SY23 3DD, United Kingdom*

^bCentre for Ecology & Hydrology, Bangor, Environment Centre Wales, Bangor, LL57 2UW, UK.

Corresponding author:

William A. V. Stiles

Institute of Biological, Environmental and Rural Sciences,

Penglais Campus,

Aberystwyth University,

SY23 3DD

Email: wvs@aber.ac.uk

3 ABSTRACT

4 ~~High organic matter soils are significant terrestrial reservoirs of carbon (C) which may be transformed~~
5 ~~from C sinks to sources by environmental change.~~ Reactive nitrogen (N) deposition can affect [many](#)
6 ecosystem processes, particularly in ~~these typically~~ oligotrophic habitats, and is expected to affect soil
7 C storage potential through increases in microbial decomposition rate as a consequence of greater N
8 availability. Increased N availability may also result in changes in the principal limitations on
9 ecosystem productivity. Phosphorus (P) limitation may constrain productivity in instances of high N
10 deposition, yet ecosystem responses to P availability are poorly understood. This study investigated
11 CO₂ and CH₄ flux responses to N and P enrichment using both short- (1 year) and long-term (16 year)
12 nutrient addition experiments. We hypothesised that the addition of either N or P will increase CO₂
13 and CH₄ fluxes, since both plant production and microbial activity are likely to increase with
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15 P enrichment, with differing results subject to the duration of nutrient addition. On average, relative to
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20 respectively), with diminishing effect into the ~~long-term~~[long-term](#)[long-term](#), suggesting microbial
21 communities at these sites are P limited. Whilst a full C exchange budget was not examined in the
22 experiment, the potential for soil C storage loss with ~~long-term~~[long-term](#)[long-term](#) nutrient
23 enrichment is demonstrated and indicates that P addition, where P is a limiting factor, may have an
24 adverse influence on upland soil C content.

25 Keywords:

26 Nitrogen deposition; soil carbon; carbon fluxes; pollution; co-limitation; P limitation

27 1.1 INTRODUCTION

28 Global climate change is expected to have profound impacts on natural systems, which could threaten
29 biodiversity and ecosystem processes (Walther et al. 2002; Thomas et al. 2004; Grim et al. 2013;
30 Carroll et al. 2015). In northern temperate ecosystems, carbon (C) has been accumulating in terrestrial
31 reservoirs since the end of the last ice age, sequestered mainly as organic matter in soils with low rates
32 of decomposition such as peat (Yu et al. 2010). Maintaining these C stores is important to avoid the
33 transformation from current status of C sink, to potential C source (House et al. 2010). Carbon stored
34 in soil is lost to the atmosphere in the form of CO₂ from soil organic matter mineralisation (Dawson &
35 Smith 2007) and CH₄ from the anaerobic decomposition of organic matter by methanogenic microbes
36 (Bubier & Moore 1994; Cooper et al. 2014). Emissions of CO₂ are considerably larger than CH₄

37 emissions, but the global warming potential of CH₄ is 28 times greater than CO₂, making it an
38 important GHG (IPCC 2013).

39 The effect of climate change on soil C is controversial (Worrall et al. 2004; Davidson & Janssens
40 2006; Worrall & Burt 2007; Clark et al. 2010), but as temperatures rise the rate of organic matter
41 decomposition is expected to increase, potentially resulting in a positive feedback on climate change
42 (Knorr et al. 2005). This effect is also likely to be exacerbated by current land use practices including
43 drainage, grazing and burning (Wallage et al. 2006; Ward et al. 2007 & 2013) and by the effects of
44 nutrient enrichment from nitrogen (N) deposition (Bragazza et al. 2006). Nutrient availability also
45 affects soil C fluxes via influences on plant productivity and inputs of labile C into soil.

46 Net fluxes of C from soil are affected by numerous environmental factors and ecosystem processes.
47 Soil moisture, temperature and pH have strong controlling effects on emissions of gaseous C from
48 soil, by influencing soil microbial and plant root activity, and the diffusion of gases through soil pores
49 (Smith et al. 2003; Chen et al. 2015). Vegetation composition can also affect rates of both CO₂ and
50 CH₄ emission (Raich & Tufekciogul 2000; Robroek et al. 2015). For CH₄, the effect is direct via the
51 transport of gas through the aerenchymatous tissue of some vascular plants (notably sedges), allowing
52 gaseous exchange with the atmosphere (Joabsson et al. 1999; McEwing et al. 2015). For CO₂,
53 vegetation composition has indirect effects via changes to net C input by plants, variation in
54 decomposition resistance of plant material, rates of root respiration, and influences on soil
55 microclimate and structure (Raich & Tufekciogul 2000). This has significant implications for soil C
56 flux in the context of current environmental change (Berendse et al. 2001; Ward et al. 2013; Xu et al.
57 2015) and potentially makes forecasting soil C flux highly complex, as all biotic and abiotic factors
58 not only act directly, but also indirectly through the modification of vegetation composition by soil
59 characteristics, and *vice versa* (McEwing et al. 2015).

60 Nitrogen enrichment from pollutant deposition affects ecosystem processes (Magnani et al. 2007;
61 Jones & Power 2012; Southron et al. 2013) and can affect rates of C flux through changes to
62 vegetation and microbial activity (Basiliko et al. 2006; Juutinen et al. 2010; Wu et al. 2015). The
63 effect of N on soil CO₂ flux is controversial, with studies demonstrating an inhibitory (Phillips &
64 Fahey 2007; Janssens et al. 2010; Ramirez et al. 2010) or a stimulatory (Bragazza et al. 2006;
65 Cleveland & Townsend 2006; Zhang et al. 2013) effect. This is potentially the result of differences in
66 C availability (Knorr et al. 2005). In oligotrophic environments with soils of high organic matter
67 content, nutrient enrichment from N deposition is expected to impact soil C storage potential by
68 increasing microbial activity and thus decomposition rates (Bragazza et al. 2006). It is also expected
69 that an increase in N availability will shift ecosystems towards limitation of other nutrients such as
70 phosphorus (P) (Crowley et al. 2012; Peñuelas et al. 2013). Phosphorus limitation may be an
71 important mechanism that constrains productivity in situations of high N deposition, hence release

72 from this may greatly affect many ecosystem processes and fluxes. Understanding the role of nutrient
73 availability in determining C fluxes and C storage potential is of considerable importance. The
74 availability of reactive N in upland habitats with peaty soils has been increased by anthropogenic
75 sources, and is projected to double in size globally from current levels by 2050 (Galloway et al. 2004;
76 Phoenix et al. 2006). The rate of reactive N input is reportedly in decline in Western Europe (Fowler
77 et al. 2004), but current rates remain higher than the estimated critical load for many upland habitats
78 (RoTAP 2012).

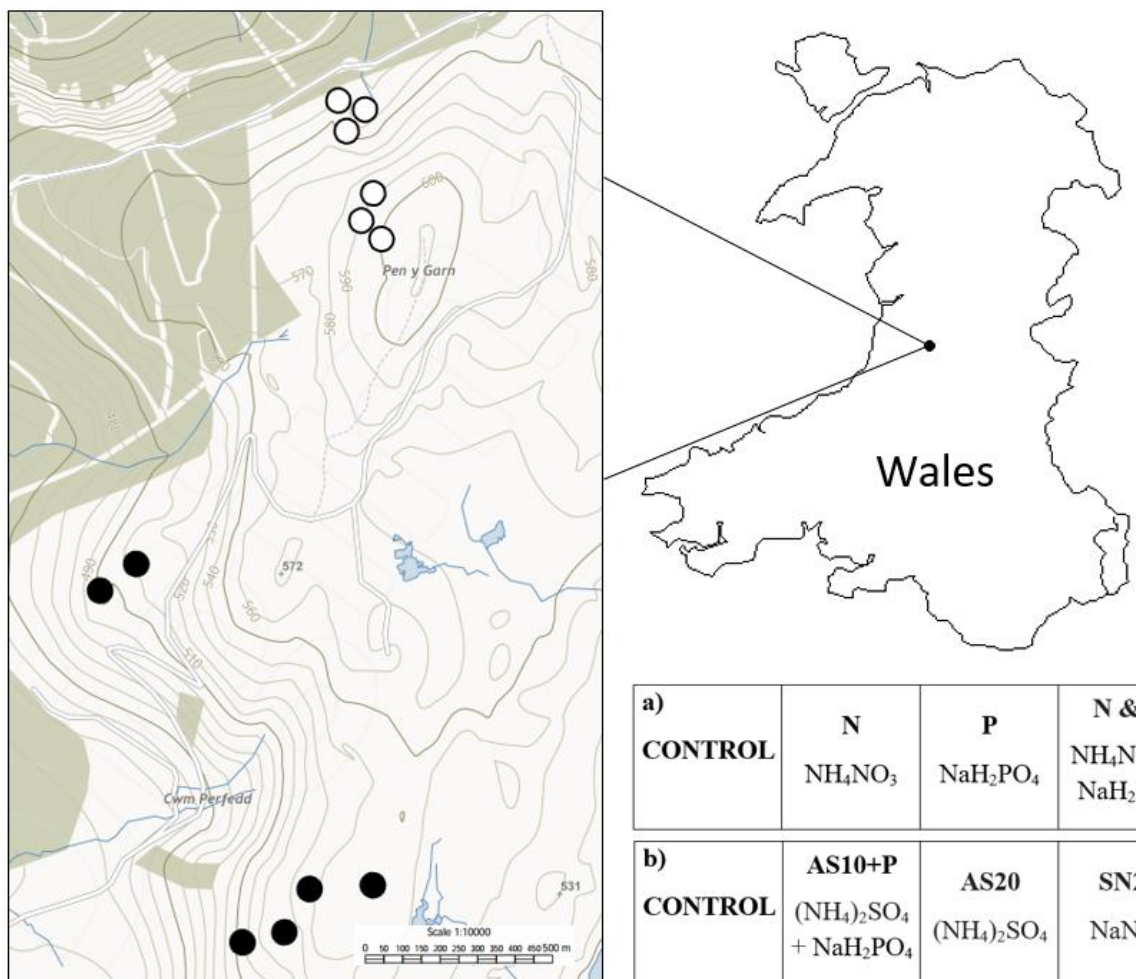
79 To date, research in this area has focussed on the environmental factors which ~~can~~ promote C flux,
80 such as vegetation composition, temperature and soil moisture. Few studies have considered the effect
81 of increased N availability on C flux within upland habitats that have been exposed to N deposition
82 rates near or above the critical load, with none considering the impact of P limitation in this context.
83 This study aimed to establish the effect of N and P enrichment on soil C flux by measuring CO₂ and
84 CH₄ emissions across two randomised block experiments, which have been run in both the short (one
85 year) and longer term (1996 - 2012, with sampling undertaken three years after ceasing N inputs and
86 15 years after a single P application). We hypothesise that the addition of nutrients (N or P) will
87 increase CO₂ flux (H1) and increase CH₄ flux (H2) at these upland sites, and N and P addition will
88 have stronger stimulation effects than N addition alone on fluxes of CO₂ (H3) and CH₄ (H4). The
89 response will be greatest in treatments where P is added, since plant growth and microbial activity are
90 expected to increase with alleviation from nutrient limitation.

91 **1.2 METHODS**

92 1.2.1 EXPERIMENTAL DESIGN

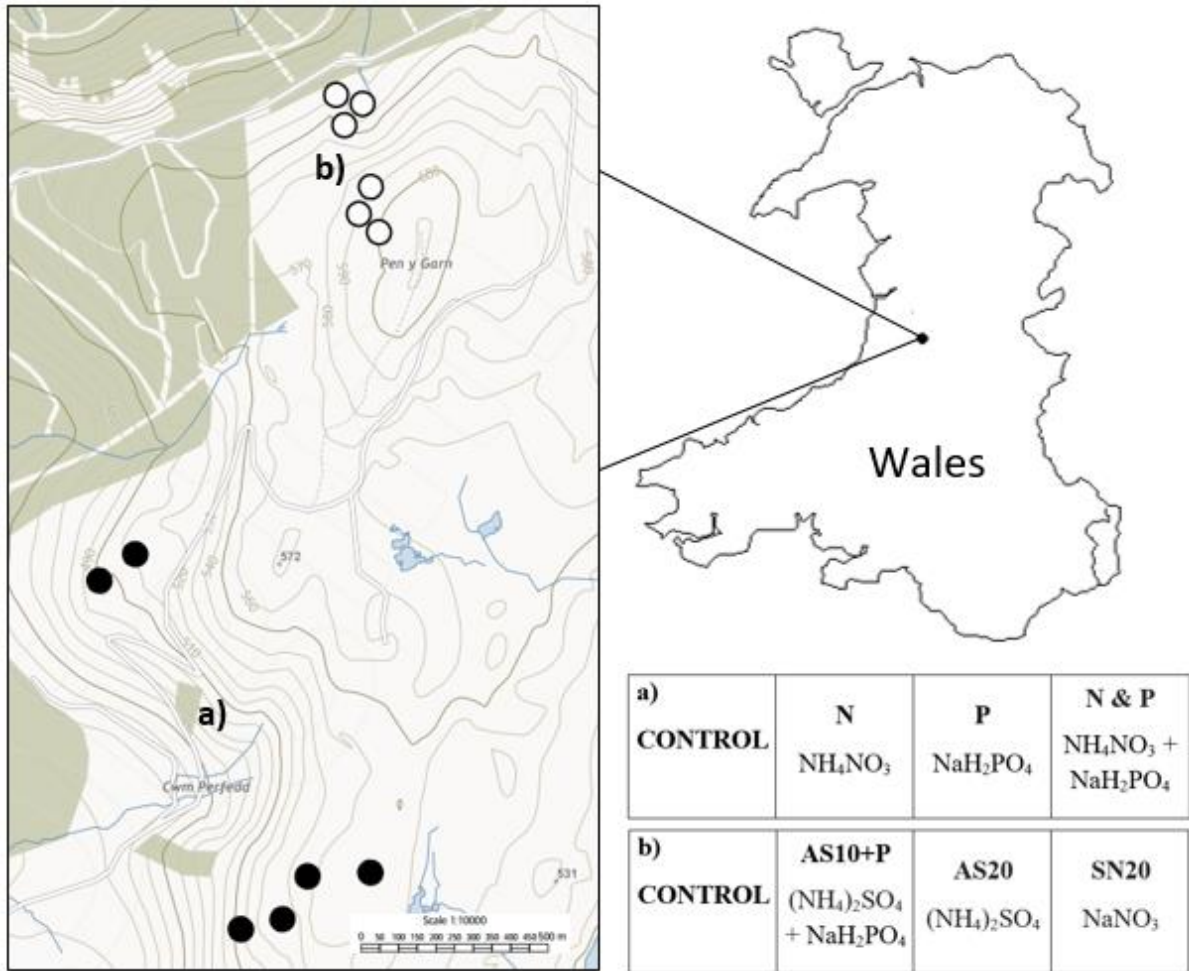
93 This study was conducted at two sites: Pen y Garn (PEN - 52° 37' N, 3° 76' W) and Pwllpeiran (PWL
94 - 52° 37' N, 3° 77' W) in the Cambrian Mountain range, mid-Wales. The sites were located within 1
95 km of each other and within an altitude range of 500 - 600 metres a.s.l., ~~and~~ ~~both~~ Both sites were on a
96 transition between NVC U4 *Festuca ovina* / *Agrostis capillaris* grassland and H18 *Vaccinium*
97 *myrtillus* / *Deschampsia flexuosa* heath, overlying mixed soils ranging from shallow ferric
98 stagnopodzol to deep peat (Emmett et al. 2007; Phoenix et al. 2012). The mean annual rainfall rate for
99 this location was 1512.2 mm (UK Meteorological Office, no date) and the background N deposition
100 rate was 22 kg N ha⁻¹ yr⁻¹ (Emmett et al. 2007). Nutrient addition experiments were established at both
101 locations, each consisting of six replicate blocks of four 3 x 3 m plots, in a randomised block design.
102 In the long-term experiment at PEN (Figure 1b), N was added fortnightly between 1996 and 2012,
103 and P was added once in 2000 to the ammonium sulphate 10 kg N ha⁻¹ yr⁻¹ (AS10+P) treatment. The
104 second experiment, at PWL, was established more recently, with different nutrient addition treatments
105 to better distinguish the effect of P addition by including a P-only treatment (Figure 1a). In both

106 experiments, the addition of N was split to avoid abrupt increases in concentration, with additions of
 107 1/14 of the total dose every two weeks between the months of April – October 2014.



a)	N	P	N & P
CONTROL	NH_4NO_3	NaH_2PO_4	$\text{NH}_4\text{NO}_3 + \text{NaH}_2\text{PO}_4$
b)	AS10+P	AS20	SN20
CONTROL	$(\text{NH}_4)_2\text{SO}_4 + \text{NaH}_2\text{PO}_4$	$(\text{NH}_4)_2\text{SO}_4$	NaNO_3

108



109

110 Figure 1. Experimental treatments at two sites, each replicated six times: black circles and a)
 111 **Pwllpeiran (PWL)**, (experiment set up in 2004): CONTROL = no addition; N = ammonium nitrate at
 112 60 kg N ha⁻¹ yr⁻¹; P = phosphorus (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹; N & P =
 113 ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ + sodium dihydrogen orthophosphate at 40 kg P ha⁻¹ yr⁻¹; white
 114 circles and b) **Pen-y-Garn (PEN)**, (experiment set up in 1996): CONTROL = no addition; AS10+P =
 115 ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus (sodium dihydrogen orthophosphate) at 20 kg P
 116 ha⁻¹ yr⁻¹; AS20 = ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹ yr⁻¹.

117 At the PWL site, P was added once in April 2014. The PWL site was grazed during the study, at a rate
 118 of 1.0 sheep/ha. The PEN site was not grazed during this study, but had been grazed by sheep
 119 between 1990 and 2007 at two levels (1.0 sheep/ha and 1.5 sheep/ha). Although there was no
 120 grazing at the PEN site between 2007 and 2015, these moorland habitats are resistant to succession
 121 and there was no change in habitat type. Three replicate blocks were established in each of two
 122 paddocks that had been grazed at the different levels, but the lack of grazing during the intervening
 123 seven years had greatly diminished effects of the grazing treatments and differences in previous
 124 stocking rate were not taken into account in the design.

125 1.2.2. VEGETATION ANALYSIS

126 Vegetation data were collected for PWL in June 2014 and June 2015 and for PEN in June 2015.
127 Cover was recorded visually for each species within each 3 x 3 m plot using the Domin scale, and
128 subsequently transformed to percentage cover for statistical analysis (Currall 1987). All plants and
129 bryophytes were identified to species level. Vegetation height was recorded as the average of five
130 measurements to the top of the canopy taken using a metric sward stick, of ca. 1 cm diameter, marked
131 at 0.5 cm intervals (Dennis *et al.* 2005). At ~~PENPEN~~, heights were taken from the centre of each plot
132 and then from the mid-point between the centre and each corner. ~~and at~~ PWL, heights were recorded
133 from within small grazing exclosures (~1 m²) that were established within each plot.

134 1.2.3 SOIL ANALYSIS

135 Soil samples for both sites were collected in June 2015 from five locations within each treatment plot,
136 using a 20 mm diameter soil corer up to a depth of 20 cm, and bulked together. The samples were air
137 dried and passed through a 2 mm sieve. Soil pH was measured with a Hydrus 400 meter (Fisherbrand,
138 Leicestershire, UK) in a slurry of 10 g fresh soil in 25 mL water. Total C was established by the
139 Dumas combustion method using an elemental analyser (Vario MAX Cube – Elementar
140 Analysensysteme, Hanau, Germany).

141 1.2.4 SOIL FLUX

142 Fluxes of CO₂ and CH₄ were measured for the different nutrient addition treatments at both PWL and
143 PEN from September 2014 to August 2015; initially every two months until March 2015, then every
144 month during the growing season (May to August 2015), under dark conditions using a non-steady
145 state, static chamber approach (Livingston & Hutchinson 1995; Parkin & Venterea 2010). Measured
146 CO₂ flux thus represents both soil and plant respiration (ecosystem respiration) but excludes the
147 effects of photosynthesis. Soil collars, 25 cm in diameter, were installed on each treatment in July
148 2014, two months prior to the first gas flux measurements to reduce disturbance. A soil knife was
149 used to cut through the initial vegetation and topsoil layer, before the collar was inserted to a depth of
150 7 cm, leaving 5 cm aboveground for chamber attachment. The location of the collar was selected
151 randomly for each treatment square and once installed, vegetation was trimmed from inside the collar
152 and maintained trimmed throughout sampling. Chamber design followed the description by Parkin &
153 Venterea (2010). Each chamber measured 19 cm in height including the soil collar and had an internal
154 volume of 9.33 L. A modified pressure vent tube (Xu *et al.* 2006) was installed to allow internal and
155 ambient air pressure equilibration. Before each measurement, chambers were ~~carefully~~ placed
156 carefully on collars, ensuring a gas tight seal. The soil CO₂ and CH₄ fluxes were calculated ~~on the~~
157 ~~basis of~~ based on changes in chamber concentrations over 30 minutes. Measurements were taken for
158 each chamber at 0, 15 and 30 minutes from chamber attachment (Parkin & Venterea 2010). The
159 duration of measurement at each sampling point was 60 seconds. CO₂ and CH₄ concentrations were
160 measured using an LGR™, Ultra- Portable Greenhouse Gas Analyser (Model 915–0011, Los Gatos

161 Research, Palo Alto, CA, USA) with a 1 Hz sampling rate. Chambers were attached to the analyser
162 via inlet and outlet tubing (2 m by 4 mm internal diameter). Chamber internal temperature, ambient
163 air temperature and soil temperature at a depth of 10 cm were also recorded. The rate of change in gas
164 concentration inside the chambers was established with linear regression. Gas flux (CO₂ and CH₄)
165 was then calculated from the rate of gas concentration change using the following equation (McEwing
166 et al. 2015):

$$167 \quad F_0 = S \frac{V M 273.16}{A V_m (273.16 + T)} 60$$

168 Where:

- 169 F_0 = Flux ($\mu\text{g CH}_4/\text{CO}_2\text{-C m}^{-2} \text{ hr}^{-1}$)
170 S = Rate of change in CH₄ and CO₂ concentration (ppm min^{-1})
171 V = Chamber volume (m^3)
172 A = Chamber area (m^2)
173 M = Molecular mass of CH₄/CO₂ (g mol^{-1})
174 V_m = Ideal gas mole volume ($0.0224 \text{ m}^3 \text{ mol}^{-1}$)

175
176 Each regression plot was ~~individually~~ assessed individually using R^2 as an indicator of accuracy, and
177 plots with $R^2 > 0.7$ were accepted for analysis. Low fluxes for CH₄ typically give a low R^2 , but should
178 be included to avoid over-estimation of mean flux (Alm et al. 2007), so CH₄ fluxes where $R^2 < 0.7$, but
179 where measurements did not exceed 0.3 ppm, were also retained as zero measurements. Two collars
180 at PWL suffered from excessive water pooling, which visibly affected the vegetation and
181 ~~soil. These soil. These~~ were removed from the dataset before analysis. Net ecosystem exchange was
182 not considered as part of this study, so whilst the results demonstrate differences in C flux, primary
183 productivity was not measured and thus the full influence of N and P availability on C budget cannot
184 be determined.

185 1.2.5 STATISTICAL ANALYSIS

186 All variables were tested visually for normality and homoscedasticity with Levene's test prior to
187 statistical analysis. Data for CO₂ and CH₄ fluxes were Log (x+1) transformed to meet the assumptions
188 of analyses. Differences in CO₂ and CH₄ flux between treatments were analysed with linear mixed
189 models (~~LMMs~~) specifying repeated measures, with treatment as a fixed factor and air temperature as
190 covariate to allow for diurnal/seasonal variance. When significant differences between treatments
191 were detected ($P < 0.05$), post-hoc tests were conducted using LSD pairwise comparisons. The
192 relationship between GHG flux and air temperature was initially tested for significance with simple
193 linear regression before inclusion in the model. Air temperatures were used in the analysis since data
194 for soil temperature were only captured from January onwards, but the two measurements were shown
195 to be correlated (PWL: $R^2 = 0.52$, $P = < 0.001$; PEN: $R^2 = 0.81$, $P = < 0.001$). Differences in mean

196 annual soil emission of CO₂ and CH₄ between experimental treatments were also investigated with
 197 one-way analysis of variance (ANOVAs). LSD pairwise comparisons were used to further investigate
 198 individual relationships. Multiple regression analysis with all-possible-subsets model procedure was
 199 used to investigate the relationships between annual average GHG flux and environmental
 200 parameters: soil pH, vegetation height, and cover values for each of five plant functional types. All
 201 statistical analyses were conducted with SPSS 21.0 (IBM SPSS Statistics for Windows, 2012).

202 1.3 RESULTS

203 1.3.1 ENVIRONMENTAL CONTROLS ON CO₂ AND CH₄ FLUX

204 Over the sampling period, ~~sizeable~~ fluctuations were observed for air and soil temperatures in line
 205 with seasonal variation. The highest air temperatures were recorded in June for PWL (22.3 °C) and in
 206 August for PEN (19.4 °C). Air temperatures were lowest at both locations in January (PWL = 5.0 °C,
 207 PEN = 6.2 °C). Soil temperatures were similar with the highest values recorded in August for PWL
 208 (13.2 °C) and in July for PEN (12.9 °C) and lowest vales for both in January (PWL = 4.6 °C, PEN =
 209 3.5 °C). Over the twelve months of sampling, the effect of changing air temperature on CO₂ flux was
 210 significant for both sites, with positive relationships observed between increasing temperature and
 211 CO₂ emission (PWL: R² = 0.68, P = <0.001; PEN: R² = 0.70, P = <0.001, Figure 2). The effect of
 212 temperature on CH₄ flux was less consistent; no relationship was observed at PWL (R² = 0.004, P =
 213 0.37), but a significant positive relationship was recorded at PEN (R² = 0.033, P = 0.012).

214 The results of the multiple regression analysis of effects of vegetation cover values, vegetation height
 215 and soil pH on mean annual CO₂ and CH₄ fluxes are shown in Table 1. This analysis revealed
 216 significant relationships between soil pH and fluxes for both CO₂ and CH₄ at PWL (Table 1, Figure
 217 3), but other factors were shown to have no effect. At ~~PENPEN~~, no environmental factor was shown
 218 to have an effect on either CO₂ or CH₄ flux despite the ~~sizeable-significant~~ differences in vegetation
 219 cover between treatments as a result of long-term nutrient addition (Figure 6).

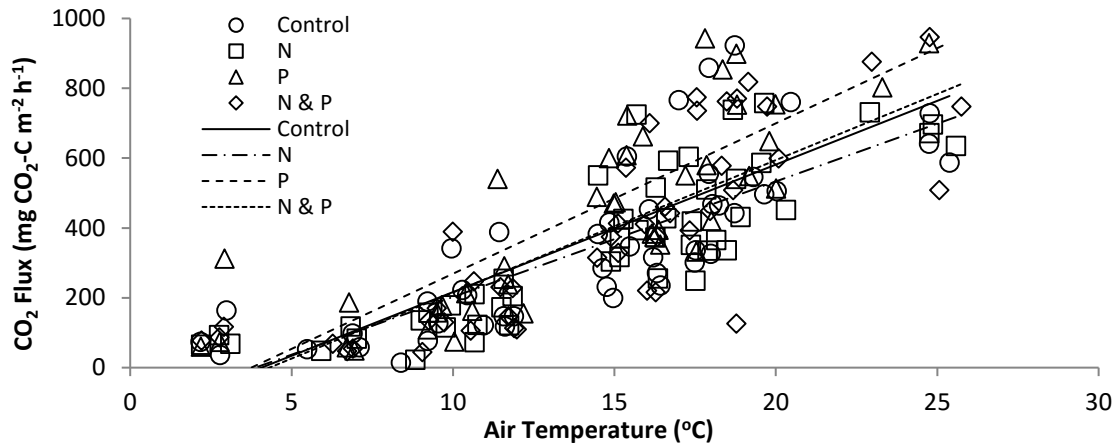
220 Table 1. Relationships between gas flux and environmental variables with potential influence on C
 221 emission.

Site	Variable	Regression model	Degrees of freedom	R ²	P
PWL	CO ₂	pH	1,21	0.52	<0.001
	CH ₄	pH	1,21	0.25	0.019
PEN	CO ₂	-	1,23	-	ns
	CH ₄	-	1,23	-	ns

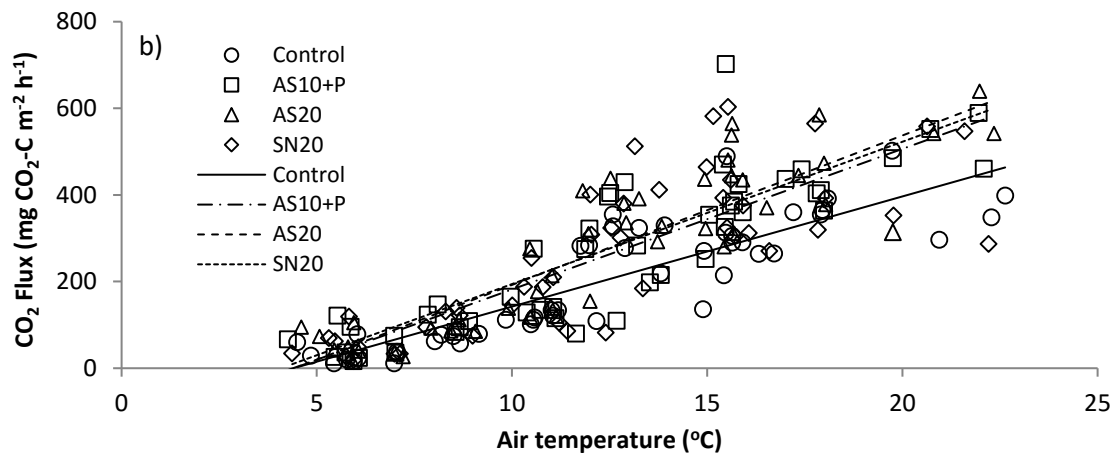
222 P-values represent the chance that the regression slope is not different from 0; bold highlights results
 223 significant at P <0.05. The best fitting regression model ~~for~~ using all-possible-subsets ~~model~~
 224 ~~procedure~~ regression analysis is shown for each parameter. Variables included in the analysis were:

225 ground cover of each plant functional type (graminoid, cryptogam, dwarf shrub, forbs), vegetation
226 height and soil pH.

227



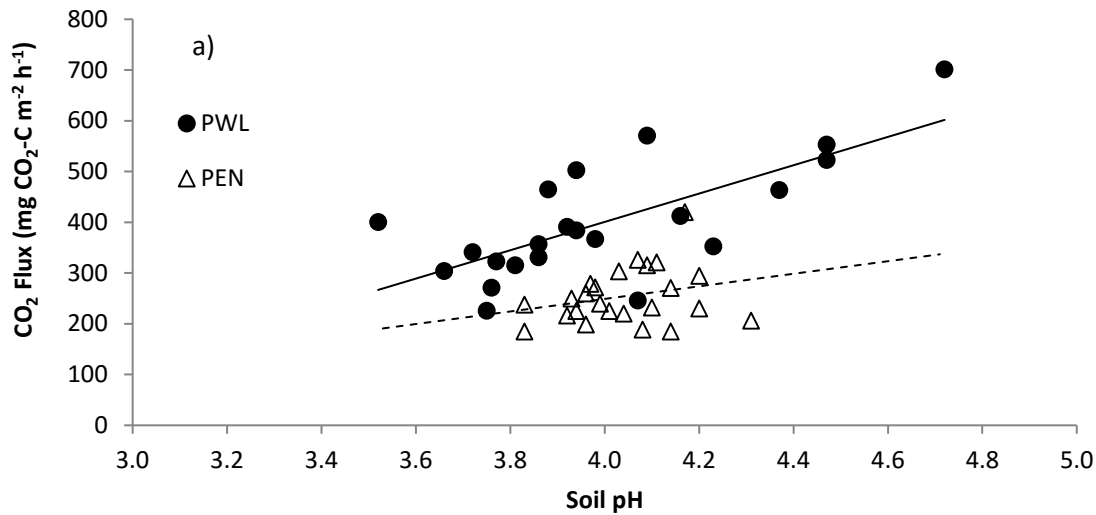
228



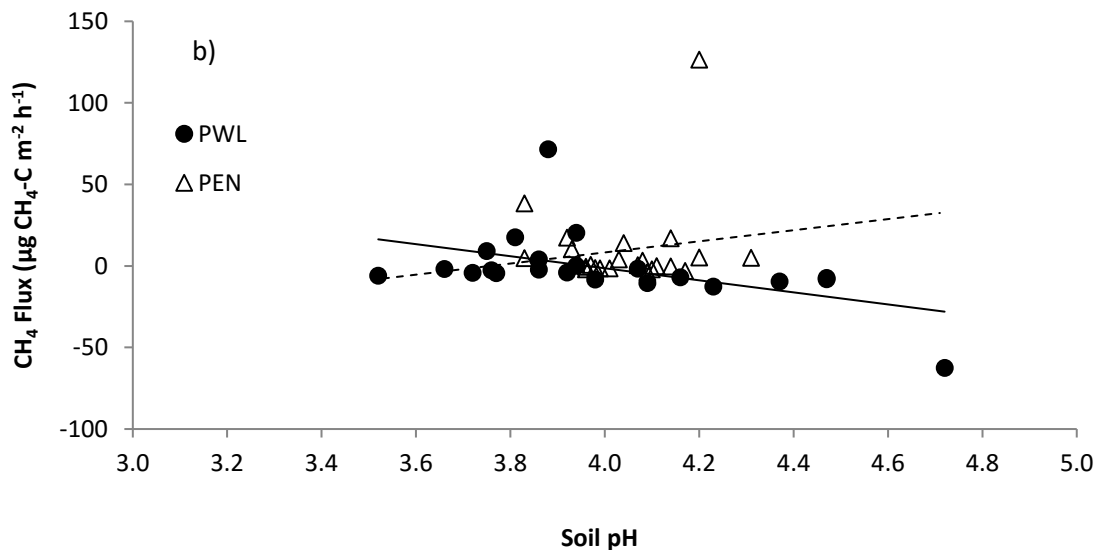
229

230 Figure 2. Relationships between CO₂ flux and air temperature at two experimental sites: a)
231 **PwllperianPWL**: Control = no nutrient addition (circles); N = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹
232 (squares); P = phosphorus (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹ (triangles); N & P
233 = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹ (diamonds); and b) **Pen-y-**
234 **GarnPEN**: Control = no nutrient addition (circles); AS10+P = ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹
235 + phosphorus at 20 kg P ha⁻¹ yr⁻¹ (squares); AS20 = ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹
236 (triangles); SN20 = sodium nitrate at 20 kg N ha⁻¹ yr⁻¹ (diamonds).

237



238



239

240 Figure 3. The relationship between soil pH and a) CO₂ flux and b) CH₄ flux at two sites: **Pwlpeiran**
 241 **PWL** (black circles) and **Pen-y-Garn****PEN** (white triangles). Solid lines indicate significant
 242 relationships, dashed lines non-significant relationships.

243 1.3.2 EFFECTS OF NUTRIENT ADDITION ON CO₂ FLUXES

244 Differences were observed in the CO₂ fluxes between nutrient addition treatments at both sites, but
 245 with differences [between sites](#) in the size of response observed [between sites](#). For PWL, the between-
 246 treatment differences (Table 2, Figure 4a) were shown by *post hoc* analysis to be driven by the largest
 247 mean annual flux recorded for the P addition treatment (464 mg CO₂-C m⁻² h⁻¹). The P treatment had
 248 consistently higher CO₂ flux in all but the first month sampled (September) ([Table 2, Figure 4a](#)). The
 249 N addition treatment had the lowest mean annual flux (353 mg CO₂-C m⁻² h⁻¹) and was consistently
 250 the lowest flux recorded across the sampling period (in all months except May). The CO₂ fluxes in the
 251 control and N + P treatments had similar mean annual fluxes (386 and 413 mg CO₂-C m⁻² h⁻¹
 252 respectively). For PEN, *post hoc* analysis revealed the differences between treatments (Table 2,

253 Figure 4b) to be due to the lowest mean annual flux recorded, which was the control (205 mg CO₂-C
 254 m⁻² h⁻¹). [The AS10+P, AS20 and SN20 treatments](#) had similar mean annual fluxes (261, 277 and 274
 255 mg CO₂-C m⁻² h⁻¹ respectively). The differences observed between treatments at PEN suggest that
 256 nutrient enrichment in the ~~long-term~~[long-term](#) significantly increases CO₂ flux relative to
 257 control, but there were no [statistically significant](#) differences among the nutrient addition treatments,
 258 suggesting limited influence of the P added to the AS10+P treatment on CO₂ flux 15 years after
 259 application.

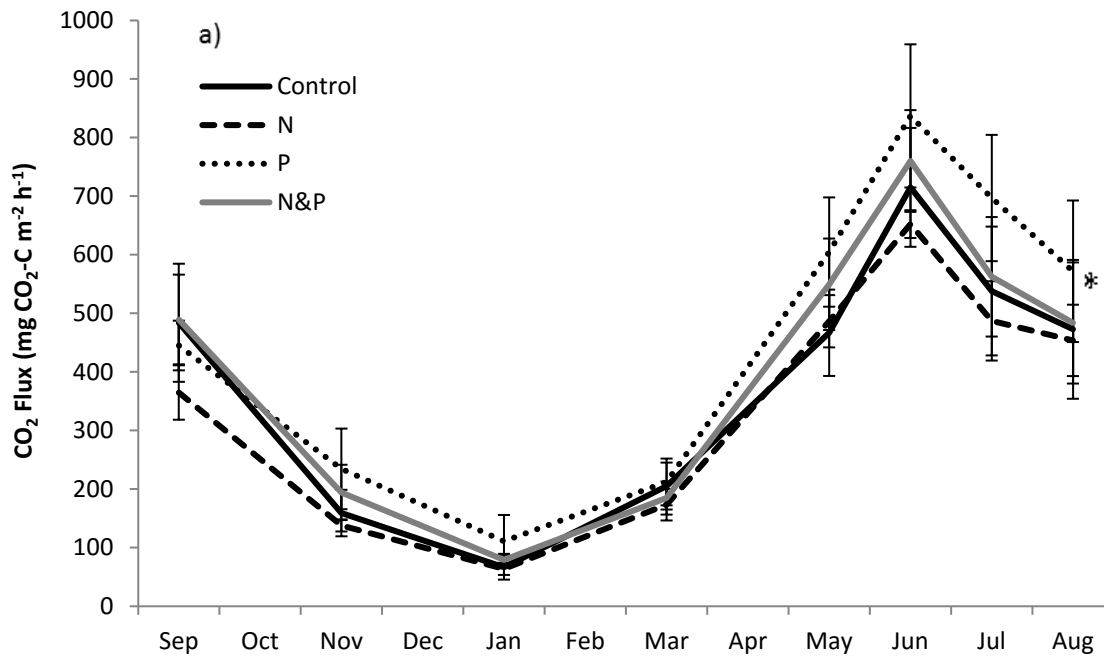
260

261 Table 2. Summary of results from [LMMs-Linear Mixed Models](#) describing soil CO₂ and CH₄ flux
 262 responses to nutrient addition treatments, with treatment as fixed factor and air temperature as
 263 covariate.

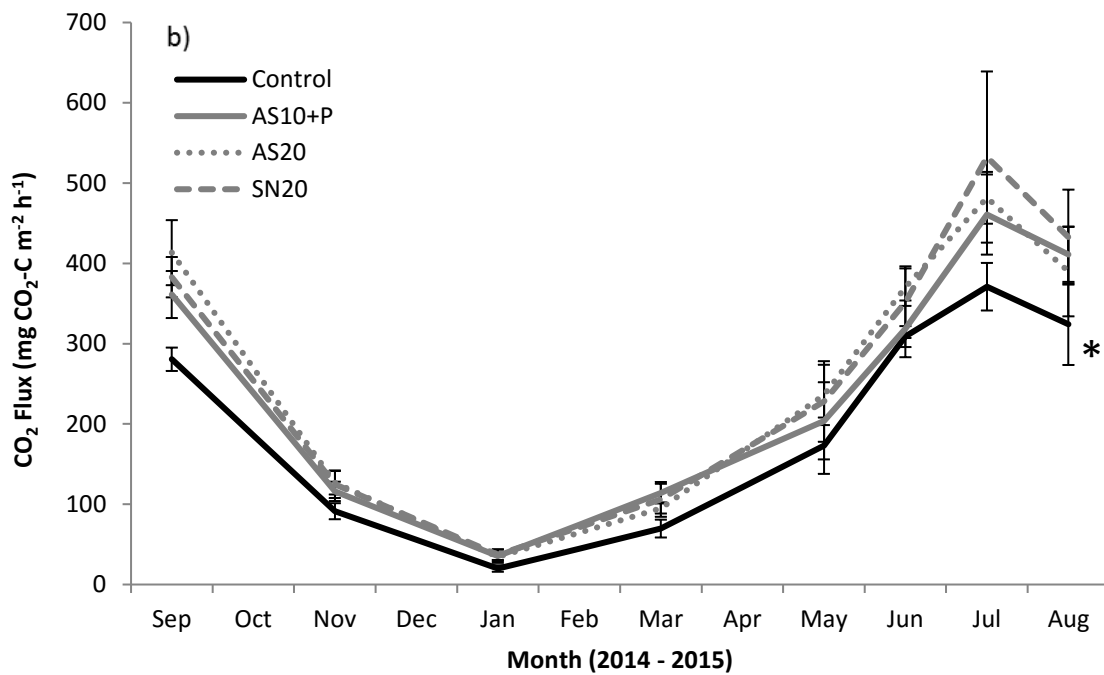
Site	Variable	Factor	Degrees of freedom	<i>F</i>	<i>P</i>
PWL	CO ₂	Treatment	3, ₁₇₁	3.69	0.013
		Air temperature	1, ₁₇₁	422.32	<0.001
	CH ₄	Treatment	3, ₁₇₁	2.86	0.041
		Air temperature	1, ₁₇₁	0.60	0.439
PEN	CO ₂	Treatment	3, ₁₈₇	4.63	0.004
		Air temperature	1, ₁₈₇	461.23	<0.001
	CH ₄	Treatment	3, ₁₈₇	4.22	0.006
		Air temperature	1, ₁₈₇	6.75	0.01

264 *P*-values refer to ~~a test of~~[a test of the likelihood that](#) the coefficient ~~being is zero~~ θ , bold highlights results
 265 significant at *P* < 0.05.

266



267

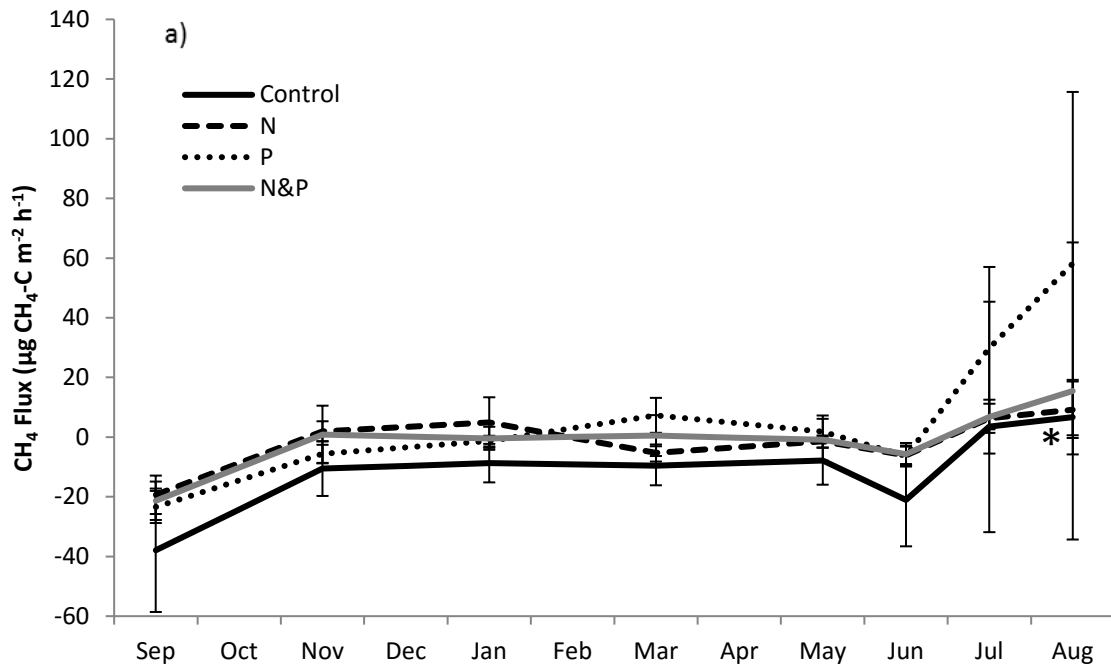


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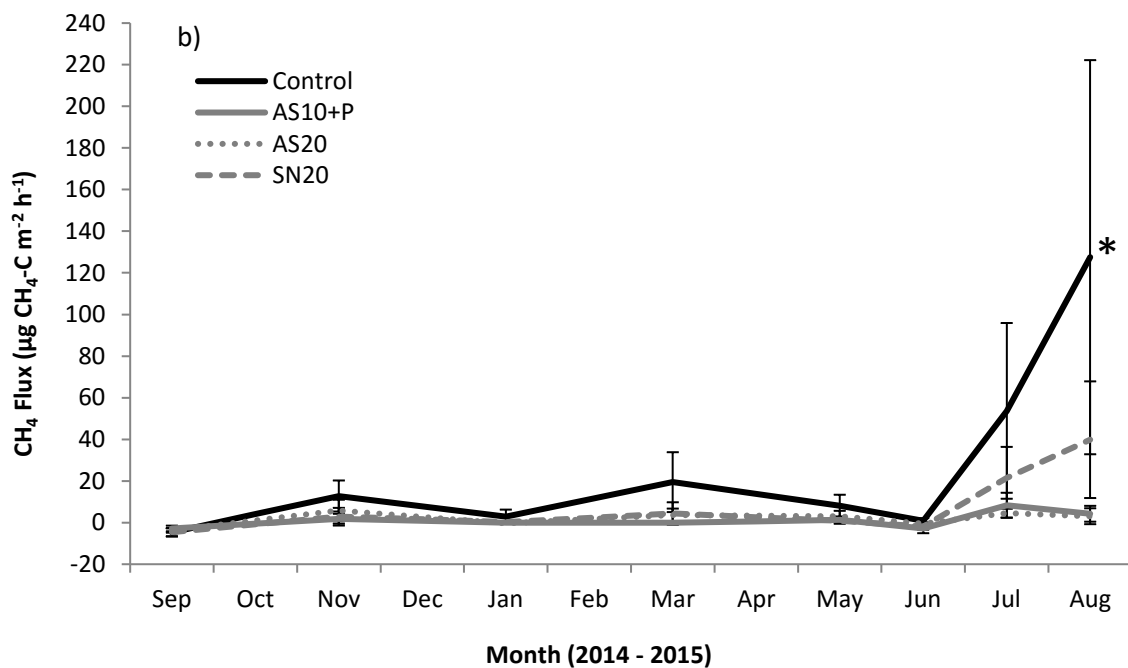
270 Figure 4. Net CO₂ emission fluxes for a) PWL and b) PEN. Error bars denote standard error. *
 271 indicates treatment with significant difference as determined by *post hoc* pairwise comparison (LSD).
 272 PWL: Control = no nutrient addition; N = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹; P = phosphorus
 273 (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹; N & P = ammonium nitrate at 60 kg N ha⁻¹
 274 yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹. PEN: Control = no nutrient addition; AS10+P = ammonium
 275 sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20 kg P ha⁻¹ yr⁻¹; AS20 = ammonium sulphate at 20 kg N
 276 ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹ yr⁻¹.

277



278

279



280

281 Figure 5. Effects of nutrient addition treatments on net methane emissions for a) PWL and b) PEN
 282 sites. Error bars denote standard error. * indicates treatment with significant difference as determined
 283 by *post hoc* pairwise comparison (LSD). PWL: Control = no nutrient addition; N = ammonium nitrate
 284 at 60 kg N ha⁻¹ yr⁻¹; P = phosphorus (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹; N & P =
 285 ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹. PEN: Control = no nutrient
 286 addition; AS10+P = ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20 kg P ha⁻¹ yr⁻¹; AS20 =
 287 ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹ yr⁻¹.

288

289 1.3.3 EFFECTS OF NUTRIENT ADDITION ON CH₄ FLUXES

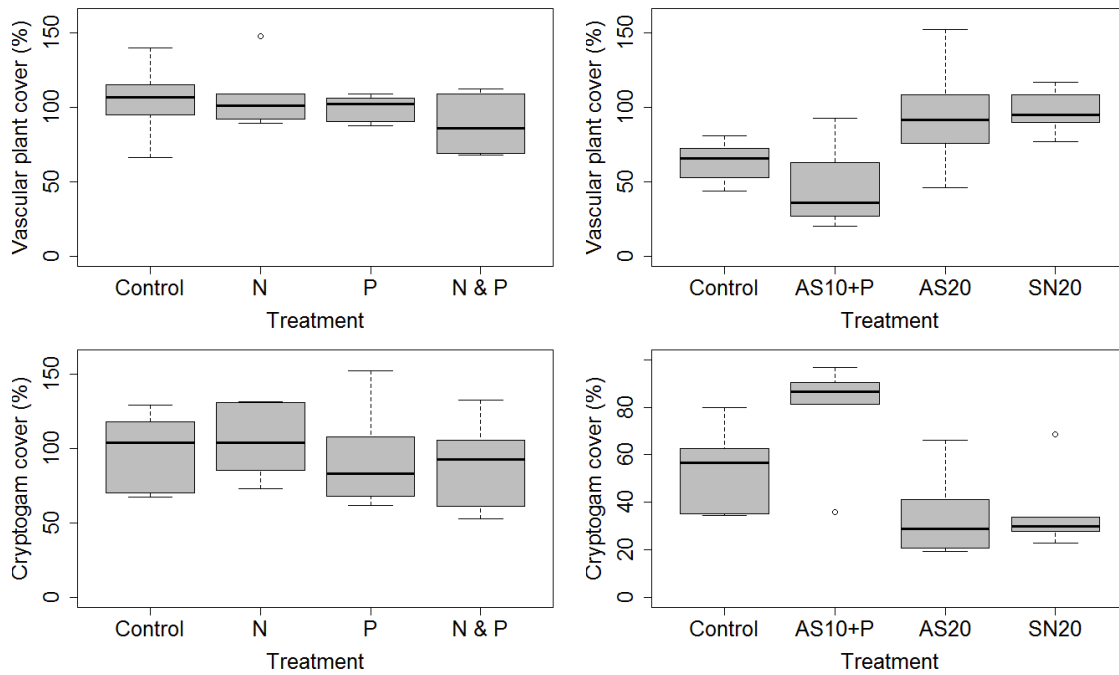
290 ~~The extent of~~ to which differences in CH₄ fluxes could be attributed to treatment effects varied
291 according to experimental site ~~Considerable variation was observed between sites for CH₄ flux.~~ At
292 PWL there were significant differences between treatments (Table 2, Figure 5a), independent of
293 seasonal variations in temperature (Table 2), which *post hoc* tests revealed to be driven by the
294 difference in CH₄ emissions between the control (no addition) and nutrient addition treatments (N, P
295 and N + P). At this site, CH₄ uptake was ~~greater~~ greatest in control plots, with this treatment behaving
296 as a small sink (on average -10.68 µg CH₄-C m⁻² h⁻¹), only emitting CH₄ during the last two months
297 (June and August), at rates just above zero µg C CH₄ m⁻² h⁻¹. The N, P and N + P treatments were
298 highly variable in emission rate, with fluctuations recorded across the sampling range. The P addition
299 treatment had the largest mean annual flux (9.00 µg CH₄-C m⁻² h⁻¹), which can chiefly be attributed to
300 an increase in emission rate during the last two months (Figure 5a). The N and N + P treatments ~~both~~
301 ~~had~~ were intermediate in response, with slightly negative emission ~~emission rates just below zero~~ (N =
302 -1.05 µg CH₄-C m⁻² h⁻¹, N + P = -0.77 µg CH₄-C m⁻² h⁻¹) on a mean annual basis (P < 0.05). At PEN
303 there were also significant differences observed between treatments (Table 2, Figure 5b), which *post*
304 *hoc* analysis revealed to be likewise driven by the difference between control and all nutrient addition
305 treatments. However, at this site the CH₄ flux was significantly higher in control plots than nutrient
306 added treatments, with mean annual emissions of 27.64 µg CH₄-C m⁻² h⁻¹ (P < 0.01). All of the
307 nutrient addition treatments at PEN otherwise had similar flux rates for the full sampling range, until
308 the last two months, when the CH₄ flux from the sodium nitrate treatment (SN20) increased ~~markedly~~
309 comparatively (Figure 5b).

310 1.3.4 EFFECTS OF NUTRIENT ADDITION ON VEGETATION COMPOSITION

311 The addition of nutrients resulted in ~~sizeable~~ significant differences in the observed vegetation cover,
312 but only after sufficient time had elapsed allowing vegetation community modification. At PWL,
313 where nutrients were added in the ~~short term~~ short-term ~~short term~~ only, no differences in vegetation
314 cover were observed. At PEN, long-term nutrient additions resulted in higher ~~greater~~ graminoid
315 species ~~cover~~ and lower ~~lesser~~ cryptogam species ~~cover~~ where N alone was added and higher ~~greater~~
316 cryptogam and lower ~~lesser~~ graminoid species ~~cover~~ in treatments where P was added (AS10+P).
317 Control plots were intermediate in composition (Figure 6).

318 PWL

PEN



319

320

321 Figure 6. Ground cover of vascular plants and cryptogams recorded for nutrient addition treatments at
 322 two sites, PWL (L-hand plots) and PEN (R-hand plots). PWL: Control = no nutrient addition; N =
 323 ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹; P = phosphorus (sodium dihydrogen orthophosphate) at 40 kg
 324 P ha⁻¹ yr⁻¹; N & P = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹. PEN:
 325 Control = no nutrient addition; AS10+P = ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20
 326 kg P ha⁻¹ yr⁻¹; AS20 = ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹
 327 yr⁻¹. Box midline indicates median, box edges indicate interquartile range. Whiskers indicate range of
 328 data within 1.5 x the interquartile range; points indicate data outside 1.5 x the interquartile range.

329

1.4 DISCUSSION

330 The addition of N and P had significant influence on CO₂ and CH₄ fluxes, with differences in the
 331 direction and magnitude of the effects between sites as a consequence of duration of nutrient addition
 332 treatment. In the short-term trial, N addition inhibited CO₂ and CH₄ flux, whereas in the long-term
 333 trial, N addition significantly increased CO₂ emissions but inhibited CH₄ flux. The addition of P
 334 significantly increased CO₂ and CH₄ flux in the short-term trial, but this effect was reduced in the
 335 long-term trial, after 14-15 years since P addition.

336

1.4.1 NUTRIENT ADDITION EFFECTS ON CO₂ FLUXES

337 The addition of N and P influenced fluxes of CO₂, however there were substantial differences in effect
 338 between treatments, and also differences in response between short-term (PWL = 1 year for
 339 N & P) and long-term nutrient additions (PEN = additions between 1996 – 2012, sampling
 340 undertaken three years after the cessation of N addition and 15 years after a single P application). At
 341 PWL, the addition of P increased CO₂ flux, which supported hypothesis H3 and is in line with
 342 findings from similar studies in other environments (Cleveland & Townsend 2006; Liu et al. 2013).
 343 Phosphorus limitation is a significant mechanism constraining ecosystem processes, particularly in

344 systems suffering from the effects of chronic N deposition (Cleveland et al. 2011; Crowley et al.
345 2012). The addition of P stimulates decomposition, with observed rises in CO₂ flux associated with
346 increases in heterotrophic (bacteria and fungi) biomass and activity, and thus respiration (Liu et al.
347 2012), rather than through increased fine-root biomass (Cleveland & Townsend 2006). Phosphorus
348 addition results in the modification of microbial community structure and reduces the ratio between
349 Gram-positive and Gram-negative bacteria, resulting in more copiotrophic communities (Fanin et al.
350 2015). At PEN, the role of P was less clear. The CO₂ flux in the N + P treatment (AS10+P) was larger
351 than in the control plots, but was mostly lower across the sampling period than CO₂ fluxes from the
352 two N-only treatments (AS20 and SN20). Phosphorus is highly persistent in soil (Nye & Tinker
353 1977), and was presumably responsible for driving shifts observed in vegetation species composition
354 (Figure 6), thus the comparatively weak effect of P on CO₂ flux at this site was unexpected. The N-
355 only treatments received N at a higher rate than the AS10+P treatment, and associated larger CO₂
356 fluxes may reflect the effect of greater N availability. In addition, while the effects of P on plant
357 species composition were still visible at the site, P may no longer have been stimulating plant
358 production of labile C. Liu et al. (2013) observed a diminishing effect of P on microbial biomass after
359 four years, which was attributed to C limitation, where the exhaustion of available soil C had a
360 constraining effect on respiration (Fanin et al. 2015).

361 The addition of N had an inhibitory effect on soil CO₂ flux at PWL and refutes hypothesis H1 that
362 nutrient addition would increase CO₂ efflux. Similar inhibition of decomposition by N has been found
363 in other studies (Ramirez et al. 2010; Chen et al. 2015), and may result from high N availability. In
364 soils where N is not a limiting factor for microbial growth and activity, the addition of N can constrain
365 organic matter decomposition (Janssens et al. 2010). In such instances, reductions in CO₂ flux may be
366 driven by shifts in C allocation from belowground to aboveground biomass (Litton et al. 2007), which
367 reduces rhizosphere and microbial respiration (Phillips & Fahey 2007; Bae et al. 2015). The
368 background N deposition for these sites is 22 kg N ha⁻¹ yr⁻¹ (Emmett et al. 2007), which is greater than
369 the critical load limit for this habitat (10 – 15 kg N ha⁻¹ yr⁻¹; APIS 2014), suggesting that these soils
370 are unlikely to be N limited. In addition, reductions in soil pH associated with N addition potentially
371 have a limiting effect (Chen et al. 2015).

372 This contrasts with the effects observed at PEN, where N addition treatments had higher CO₂ flux in
373 comparison to control (no addition) treatments, which does support hypothesis H1. PEN was sampled
374 three years after ceasing N addition, which could indicate that inhibitory effects were reduced due to
375 N losses from leaching in the intervening three years. This explanation was deemed unlikely however,
376 as the size and activity of microbial populations exposed to N enrichment can remain elevated 6–8
377 years after cessation of additions, suggesting a prolonged effect on the rate of nutrient cycling (Power
378 et al. 2006). Instead, the higher CO₂ flux where N was added is most likely the result of greater soil
379 organic matter decomposition from increased microbial abundance and activity (Mack et al. 2004;

380 Bragazza et al. 2006), and by reduced production of more decomposition resistant species (Bragazza
381 et al. 2012). This influence was not translated to a reduction in soil C levels in N treatment plots
382 however; these plots in fact had the highest soil C content, albeit not significantly greater than control
383 (Stiles et al. 2017). ~~Which~~ This suggests a potential balance between increased rates of both plant
384 production and decomposition with greater N availability ~~and thus productivity~~ (Mack et al. 2004).

385 The variability observed in CO₂ flux between sites with N enrichment could be explained by
386 differences in the N addition rate, which at PWL was three times the maximum annual application
387 rate at PEN. Janssens et al. (2010) demonstrated that the inhibitory effect of N on soil CO₂ flux was
388 larger in sites receiving N at dose rates greater than 50 kg ha⁻¹ yr⁻¹, which may account for some of the
389 between-site disparity. The inhibitory effect of N at PWL apparently offset the stimulatory effect of P,
390 in that the CO₂ flux response to N + P addition was intermediate between responses to P and N
391 addition. Thus, although the results were not consistent with a general stimulation of CO₂ efflux by
392 nutrient addition (H1), there was support for hypothesis H3 that P would have a greater stimulating
393 effect.

394 1.4.2 NUTRIENT ADDITION EFFECTS ON CH₄ FLUXES

395 The addition of N and P altered CH₄ flux, although different effects were observed in the ~~short~~
396 ~~term~~ ~~short term~~ ~~short term~~ short term after enrichment (PWL) and after a longer period (PEN). The CH₄ flux
397 results represent net emissions, and effects on CH₄ production and oxidation cannot be distinguished.
398 At PWL, all nutrient addition treatments had less negative net CH₄ emissions than the control,
399 supporting hypothesis H2. The less negative net emissions with N addition are presumably due to
400 decreases in the rate of CH₄ oxidation (Aerts & Toet 1997; Aerts & Caluwe 1999). This is primarily
401 caused by competition for the CH₄ mono-oxygenase enzyme, which is affected when N addition
402 increases rates of nitrification, inhibiting CH₄ oxidation rate and reducing the amount of methane
403 consumed by methanotrophs (Bodelier 2011). Phosphorus addition also increased the rate of emission
404 for CH₄, which supports hypothesis H4, but is contrary to similar research conducted in other
405 environments (Zhang et al. 2011; Song et al. 2012). These studies attributed the effects observed to
406 the inhibition of methanogenesis, the stimulation of methanotrophic potential, and ~~the elevation of~~
407 ~~plant root growth, which~~ increased plant growth and water uptake reducing soil water content and
408 thus increasing ~~oxidation~~ ~~methanotrophy~~ through greater aeration. This last effect would seem
409 unlikely to operate in the PWL and PEN experiments because ~~these are rather wet sites and the~~
410 ~~treatments~~ ~~vegetation community~~ would not be expected to increase soil aeration after soil water
411 availability significantly between treatments greatly, even with the potentially augmented growth
412 associated with P enrichment. The stimulation of methanogenesis is likely to be the principal
413 mechanism and has been previously observed for some peat types (Keller et al. 2006), which suggests

414 the methanogen community at PWL is P limited. Further tests would be necessary to rule out any
415 inhibitory effect of P on methane oxidation, but this is beyond the scope of the current study.

416 At PEN, the highest CH₄ flux recorded was in the control ~~and was larger than CH₄ fluxes in the~~
417 ~~nutrient addition treatments~~, which was contrary to hypothesis H2 that nutrient addition would
418 stimulate CH₄ fluxes. Fluxes of CH₄ from the nutrient addition treatments remained near zero for the
419 majority of the sampling period until the last two months (July and August), when the fluxes from the
420 sodium nitrate treatment (SN20) increased substantially. Nitrate inhibits CH₄ emission, but only in
421 very high concentrations (Bodelier & Laanbroek 2004). ~~whereas~~ ammonium has been shown to
422 inhibit CH₄ emission more strongly ~~The effects of N are more severe with ammonium than with~~
423 ~~nitrate addition~~ (Crill et al. 1994). This ~~and~~ may explain why emissions were consistently inhibited in
424 the ammonium treatments (AS10+P and AS20) in contrast to the spike in emission observed in the
425 nitrate treatment towards the end of the study period. Overall, N addition inhibited CH₄ emission,
426 which ~~is~~ can be attributed to N-induced increases in population size and activity of methane-oxidising
427 bacteria (Bodelier & Laanbroek 2004). The addition of P (AS10+P) had no obvious effect at PEN, 15
428 years after application, contrary to the H4 hypothesis. As with the CO₂ flux results, this was
429 unexpected. Whilst the exact mechanisms remain unclear, the composition of the microbial
430 community may have changed over the intervening years (as with the diminishing effect of P over
431 time on CO₂ emission, discussed above) which may have reduced the size of any initial effect. This
432 may also account for the substantial differences observed in the results of the two experiments.

433 1.4.3 ENVIRONMENTAL CONTROLS ON CARBON FLUXES

434 Temperature influenced the rate of CO₂ flux at both PWL and PEN, which is consistent with other
435 studies (Dorrepaal et al. 2009; Briones et al. 2010; Imer et al. 2013). Higher temperatures allow
436 increased rates of soil decomposition and root respiration, which are temperature-dependent chemical
437 and biochemical reactions (Knorr et al. 2005; Davidson & Janssens 2006). The role of temperature in
438 controlling CH₄ emission is less definite and was variable between sites. At PEN, there was a positive
439 effect of temperature on CH₄ flux, whereas at PWL, no effect was observed. The effect of temperature
440 on CH₄ production is reported to be variable and driven by the influence of site-specific factors such
441 as differences in soil type (Chin et al. 1999; Van Winden et al. 2012). The effect of temperature is
442 likely to be closely linked with that of soil moisture and aeration (Smith et al. 2003). If high
443 temperatures relate to drought and therefore dry soil, it would be expected that CH₄ emission would
444 be reduced, but CO₂ emission would increase (Sundh et al. 2000). Plant functional type composition
445 and above ground biomass quantity did not influence CO₂ or CH₄ fluxes, which is contrary to
446 previous studies (McNamara et al. 2008; Green & Baird et al. 2012; Cooper et al. 2014; McEwing et
447 al. 2015). This was also contrary to expectation, as differences in vegetation composition exist
448 between treatments, particularly at PEN, where long-term nutrient additions have resulted in divergent

449 vegetation composition (Figure 6). Nitrogen addition (AS20 and SN20 treatments) has driven an
450 increased ground cover of graminoid species, whereas N + P addition (AS10+P) has driven an
451 increased ground cover of cryptogam species. Vegetation impacts on CH₄ flux are particularly
452 associated with *Eriophorum* spp. presence and cover (Green & Baird et al. 2012), as these sedges
453 have aerenchymatous tissue, which can facilitate the transport of CH₄ from the anaerobic zone to the
454 atmosphere, and actively produces ~~substrates which~~ substrates that encourage methanogenesis (Cooper
455 et al. 2014). The relative scarcity of these species at either site could potentially account for the
456 limited effect of vegetation ~~coverage composition changes~~ on CH₄ flux ~~between treatments~~.

457 The positive relationship for CO₂, and negative relationship for CH₄, observed between fluxes and soil
458 pH at PWL, can be explained by changes in microbial activity. Soil acidity may be increased by N
459 enrichment (Bobbink et al. 2010; Phoenix et al. 2012), an effect which is reported to have a stronger
460 controlling effect on C flux than stimulation ~~from~~ by increased N availability (Chen et al. 2015). Soil
461 pH can affect the activity rate and composition of microbial communities, with decreases of 1.5 pH
462 units shown to reduce activity by 50% (Fernández-Calviño & Bååth 2010). Recruitment of bacterial
463 species more tolerant of lower pH ultimately modifies the community, but with an intervening lag in
464 activity (Rousk et al. 2010), a hysteresis which may explain the observed reduction in the CO₂
465 emission. Low soil pH may reduce CH₄ emissions through inhibition of methanotrophic bacterial
466 communities (Dedysh 2002). These effects potentially explain some of the differences observed
467 between the two sites, in that ~~where~~ changes to pH driven by recent nutrient additions at PWL may
468 have resulted in short-term fluctuations in microbial activity. ~~This is an effect which~~ would be
469 expected to reduce ~~in over~~ time as the system shifts to a new equilibrium, ~~such as~~ in the longer-term
470 experiment at PEN. The absence of any association with other environmental variables at PEN
471 indicates that, for this study at least, the availability of nutrients over an extended period (even three
472 years ~~post-after~~ application for N and 15 years ~~post-after~~ application for P) has a greater influence on
473 CH₄ flux rate than other background environmental factors.

474 1.5 CONCLUSION

475 The addition of N and P had significant effects on CO₂ and CH₄ fluxes but the direction and
476 magnitude of effects were different in the two experiments, most probably due to the difference in
477 ~~time since~~ duration of additions ~~ceased at~~ between the two sites. The CO₂ flux was inhibited by recent
478 additions of N, but plots that had received prolonged N additions up until three years previously
479 showed considerably increased CO₂ emissions. The long-term effects of P addition are less clear. In
480 the ~~short term~~ short term ~~short term~~, P stimulated CO₂ emissions via release from nutrient limitation,
481 but this effect appeared to diminish with time, perhaps due to a decrease in readily available C
482 substrates. This implies a negative effect of P addition on soil C storage, although this may be offset
483 at some sites by plant productivity increases. The emission rate of CH₄ for treatments where nutrients

484 were added was of similar size between sites, despite ~~sizeable~~ differences in the CH₄ flux observed in
485 control treatments at either site. This demonstrated a similar pattern of modification by nutrient
486 addition, but with different factors controlling the response relative to either control. The differences
487 between emission rates in control (no addition) treatments at the two sites indicates the influence of
488 other site characteristics on CH₄ flux rate.

489 The contrast in results from the two experiments demonstrates the differences between initial
490 ecosystem responses to nutrient addition and responses after ~~sufficient time for~~ ecosystem processes
491 and microbial assemblages have had sufficient time to adjust to more nutrient-rich conditions.
492 Perturbation in ecosystems often drives gradual change, and many experiments do not last long
493 enough for the system to reach a new equilibrium ~~requiring temporal scales larger than most~~
494 ~~experimental durations~~ (Bubier et al. 2007; Vitousek et al. 2008). In this study the differences in
495 response between experimental additions in the short and ~~long term~~ long term ~~are were~~
496 considerable, which has important implications for the use of short-term evidence to infer responses
497 of future environmental modification and potential for C storage and other ecosystem processes to,
498 with continuing atmospheric N pollution, or ~~with~~ P fertiliser additions that persist in the soil. ~~The~~
499 ~~disparity between effects at different times after exposure also has implications for current~~
500 ~~understanding of ecosystem processes.~~ Short-term studies investigating impacts of nutrient
501 enrichment may misrepresent longer-term ecosystem responses.

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1 **Nitrogen and phosphorus enrichment effects on CO₂ and methane fluxes**
2 **from an upland ecosystem**

William A. V. Stiles ^{a,b}, Edwin C. Rowe ^b, Peter Dennis ^a

*^aInstitute of Biological, Environmental and Rural Sciences, Penglais Campus, Aberystwyth
University, Wales, SY23 3DD, United Kingdom*

^bCentre for Ecology & Hydrology, Bangor, Environment Centre Wales, Bangor, LL57 2UW, UK.

Corresponding author:

William A. V. Stiles

Institute of Biological, Environmental and Rural Sciences,

Penglais Campus,

Aberystwyth University,

SY23 3DD

Email: wvs@aber.ac.uk

3 **ABSTRACT**

4 Reactive nitrogen (N) deposition can affect many ecosystem processes, particularly in oligotrophic
5 habitats, and is expected to affect soil C storage potential through increases in microbial
6 decomposition rate as a consequence of greater N availability. Increased N availability may also result
7 in changes in the principal limitations on ecosystem productivity. Phosphorus (P) limitation may
8 constrain productivity in instances of high N deposition, yet ecosystem responses to P availability are
9 poorly understood. This study investigated CO₂ and CH₄ flux responses to N and P enrichment using
10 both short- (1 year) and long-term (16 year) nutrient addition experiments. We hypothesised that the
11 addition of either N or P will increase CO₂ and CH₄ fluxes, since both plant production and microbial
12 activity are likely to increase with alleviation from nutrient limitation. This study demonstrated the
13 modification of C fluxes from N and P enrichment, with differing results subject to the duration of
14 nutrient addition. On average, relative to control, the addition of N alone inhibited CO₂ flux in the
15 short-term (-9%) but considerably increased CO₂ emissions in the long-term (+35%), reduced CH₄
16 uptake in the short term (-90%) and reduced CH₄ emission in the long term (-94%). Phosphorus
17 addition increased CO₂ and CH₄ emission in the short term (+20% and +184% respectively), with
18 diminishing effect into the long term, suggesting microbial communities at these sites are P limited.
19 Whilst a full C exchange budget was not examined in the experiment, the potential for soil C storage
20 loss with long-term nutrient enrichment is demonstrated and indicates that P addition, where P is a
21 limiting factor, may have an adverse influence on upland soil C content.

22 **Keywords:**

23 Nitrogen deposition; soil carbon; carbon fluxes; pollution; co-limitation; P limitation

24 **1.1 INTRODUCTION**

25 Global climate change is expected to have profound impacts on natural systems, which could threaten
26 biodiversity and ecosystem processes (Walther et al. 2002; Thomas et al. 2004; Grim et al. 2013;
27 Carroll et al. 2015). In northern temperate ecosystems, carbon (C) has been accumulating in terrestrial
28 reservoirs since the end of the last ice age, sequestered mainly as organic matter in soils with low rates
29 of decomposition such as peat (Yu et al. 2010). Maintaining these C stores is important to avoid the
30 transformation from current status of C sink, to potential C source (House et al. 2010). Carbon stored
31 in soil is lost to the atmosphere in the form of CO₂ from soil organic matter mineralisation (Dawson &
32 Smith 2007) and CH₄ from the anaerobic decomposition of organic matter by methanogenic microbes
33 (Bubier & Moore 1994; Cooper et al. 2014). Emissions of CO₂ are considerably larger than CH₄
34 emissions, but the global warming potential of CH₄ is 28 times greater than CO₂, making it an
35 important GHG (IPCC 2013).

36 The effect of climate change on soil C is controversial (Worrall et al. 2004; Davidson & Janssens
37 2006; Worrall & Burt 2007; Clark et al. 2010), but as temperatures rise the rate of organic matter
38 decomposition is expected to increase, potentially resulting in a positive feedback on climate change
39 (Knorr et al. 2005). This effect is also likely to be exacerbated by current land use practices including
40 drainage, grazing and burning (Wallage et al. 2006; Ward et al. 2007 & 2013) and by the effects of
41 nutrient enrichment from nitrogen (N) deposition (Bragazza et al. 2006). Nutrient availability also
42 affects soil C fluxes via influences on plant productivity and inputs of labile C into soil.

43 Net fluxes of C from soil are affected by numerous environmental factors and ecosystem processes.
44 Soil moisture, temperature and pH have strong controlling effects on emissions of gaseous C from
45 soil, by influencing soil microbial and plant root activity, and the diffusion of gases through soil pores
46 (Smith et al. 2003; Chen et al. 2015). Vegetation composition can also affect rates of both CO₂ and
47 CH₄ emission (Raich & Tufekciogul 2000; Robroek et al. 2015). For CH₄, the effect is direct via the
48 transport of gas through the aerenchymatous tissue of some vascular plants (notably sedges), allowing
49 gaseous exchange with the atmosphere (Joabsson et al. 1999; McEwing et al. 2015). For CO₂,
50 vegetation composition has indirect effects via changes to net C input by plants, variation in
51 decomposition resistance of plant material, rates of root respiration, and influences on soil
52 microclimate and structure (Raich & Tufekciogul 2000). This has significant implications for soil C
53 flux in the context of current environmental change (Berendse et al. 2001; Ward et al. 2013; Xu et al.
54 2015) and potentially makes forecasting soil C flux highly complex, as all biotic and abiotic factors
55 not only act directly, but also indirectly through the modification of vegetation composition by soil
56 characteristics, and *vice versa* (McEwing et al. 2015).

57 Nitrogen enrichment from pollutant deposition affects ecosystem processes (Magnani et al. 2007;
58 Jones & Power 2012; Southron et al. 2013) and can affect rates of C flux through changes to
59 vegetation and microbial activity (Basiliko et al. 2006; Juutinen et al. 2010; Wu et al. 2015). The
60 effect of N on soil CO₂ flux is controversial, with studies demonstrating an inhibitory (Phillips &
61 Fahey 2007; Janssens et al. 2010; Ramirez et al. 2010) or a stimulatory (Bragazza et al. 2006;
62 Cleveland & Townsend 2006; Zhang et al. 2013) effect. This is potentially the result of differences in
63 C availability (Knorr et al. 2005). In oligotrophic environments with soils of high organic matter
64 content, nutrient enrichment from N deposition is expected to affect soil C storage potential by
65 increasing microbial activity and thus decomposition rates (Bragazza et al. 2006). It is also expected
66 that an increase in N availability will shift ecosystems towards limitation of other nutrients such as
67 phosphorus (P) (Crowley et al. 2012; Peñuelas et al. 2013). Phosphorus limitation may be an
68 important mechanism that constrains productivity in situations of high N deposition; hence, release
69 from this may greatly affect many ecosystem processes and fluxes. Understanding the role of nutrient
70 availability in determining C fluxes and C storage potential is of considerable importance. The
71 availability of reactive N in upland habitats with peaty soils has been increased by anthropogenic

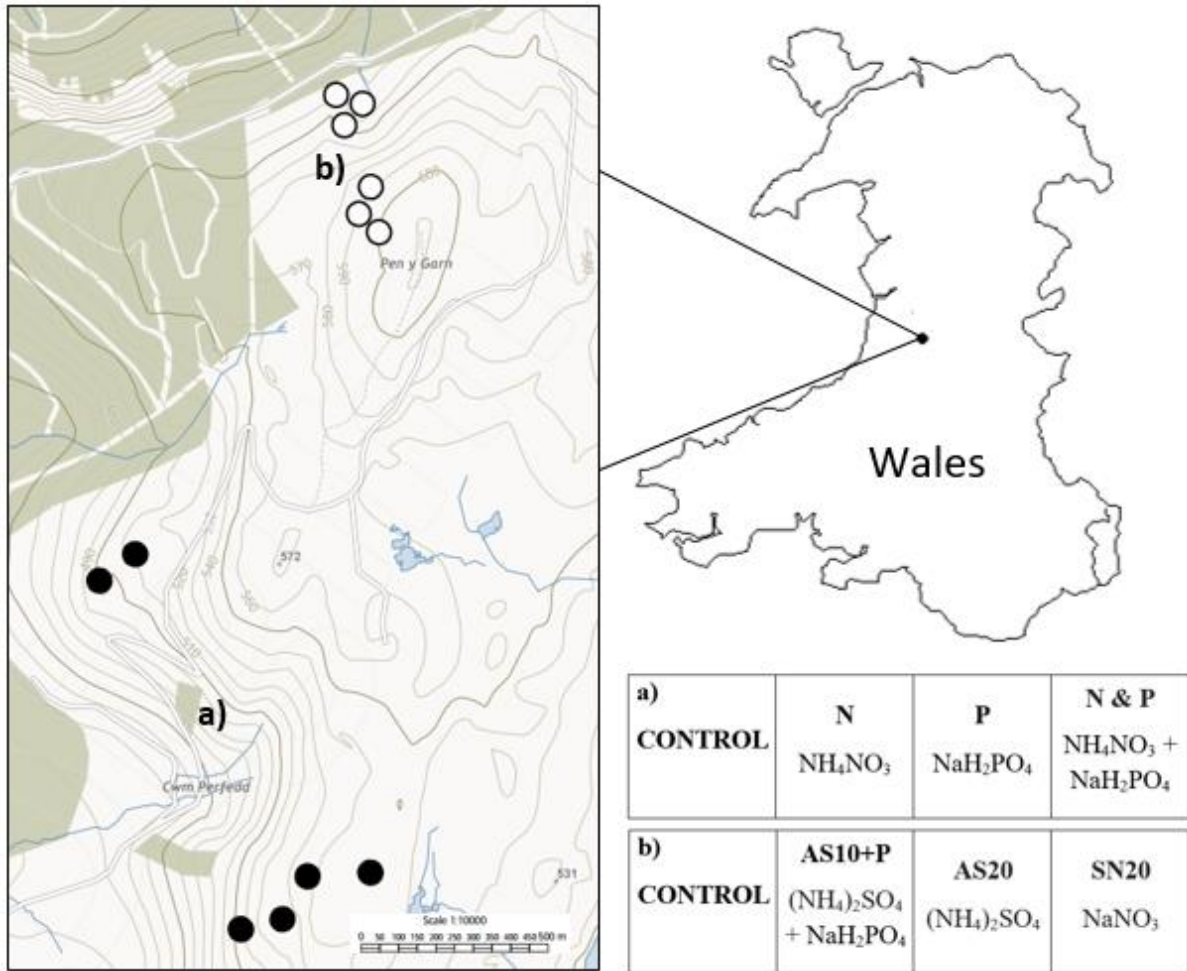
72 sources, and is projected to double in size globally from current levels by 2050 (Galloway et al. 2004;
73 Phoenix et al. 2006). The rate of reactive N input is reportedly in decline in Western Europe (Fowler
74 et al. 2004), but current rates remain higher than the estimated critical load for many upland habitats
75 (RoTAP 2012).

76 To date, research in this area has focussed on the environmental factors which promote C flux, such as
77 vegetation composition, temperature and soil moisture. Few studies have considered the effect of
78 increased N availability on C flux within upland habitats that have been exposed to N deposition rates
79 near or above the critical load, with none considering the impact of P limitation in this context. This
80 study aimed to establish the effect of N and P enrichment on soil C flux by measuring CO₂ and CH₄
81 emissions across two randomised block experiments, which have been run in both the short (one year)
82 and longer term (1996 - 2012, with sampling undertaken three years after ceasing N inputs and 15
83 years after a single P application). We hypothesise that the addition of nutrients (N or P) will increase
84 CO₂ flux (H1) and increase CH₄ flux (H2) at these upland sites, and N and P addition will have
85 stronger stimulation effects than N addition alone on fluxes of CO₂ (H3) and CH₄ (H4). The response
86 will be greatest in treatments where P is added, since plant growth and microbial activity are expected
87 to increase with alleviation from nutrient limitation.

88 **1.2 METHODS**

89 1.2.1 EXPERIMENTAL DESIGN

90 This study was conducted at two sites: Pen y Garn (PEN - 52° 37' N, 3° 76' W) and Pwllpeiran (PWL
91 - 52° 37' N, 3° 77' W) in the Cambrian Mountain range, mid-Wales. The sites were located within 1
92 km of each other and within an altitude range of 500 - 600 metres a.s.l.. Both sites were on a transition
93 between NVC U4 *Festuca ovina* / *Agrostis capillaris* grassland and H18 *Vaccinium myrtillus* /
94 *Deschampsia flexuosa* heath, overlying mixed soils ranging from shallow ferric stagnopodzol to deep
95 peat (Emmett et al. 2007; Phoenix et al. 2012). The mean annual rainfall rate for this location was
96 1512.2 mm (UK Meteorological Office, no date) and the background N deposition rate was 22 kg N
97 ha⁻¹ yr⁻¹ (Emmett et al. 2007). Nutrient addition experiments were established at both locations, each
98 consisting of six replicate blocks of four 3 x 3 m plots, in a randomised block design. In the long-term
99 experiment at PEN (Figure 1b), N was added fortnightly between 1996 and 2012, and P was added
100 once in 2000 to the ammonium sulphate 10 kg N ha⁻¹ yr⁻¹ (AS10+P) treatment. The second
101 experiment, at PWL, was established more recently, with different nutrient addition treatments to
102 better distinguish the effect of P addition by including a P-only treatment (Figure 1a). In both
103 experiments, the addition of N was split to avoid abrupt increases in concentration, with additions of
104 1/14 of the total dose every two weeks between the months of April – October 2014.



105

106 Figure 1. Experimental treatments at two sites, each replicated six times: black circles and a) PWL,
 107 (experiment set up in 2004): CONTROL = no addition; N = ammonium nitrate at $60 \text{ kg N ha}^{-1} \text{ yr}^{-1}$; P
 108 = phosphorus (sodium dihydrogen orthophosphate) at $40 \text{ kg P ha}^{-1} \text{ yr}^{-1}$; N & P = ammonium nitrate at
 109 $60 \text{ kg N ha}^{-1} \text{ yr}^{-1} + \text{ sodium dihydrogen orthophosphate at } 40 \text{ kg P ha}^{-1} \text{ yr}^{-1}$; white circles and b) PEN
 110 (experiment set up in 1996): CONTROL = no addition; AS10+P = ammonium sulphate at 10 kg N ha^{-1}
 111 $\text{ yr}^{-1} + \text{ phosphorus (sodium dihydrogen orthophosphate) at } 20 \text{ kg P ha}^{-1} \text{ yr}^{-1}$; AS20 = ammonium
 112 sulphate at $20 \text{ kg N ha}^{-1} \text{ yr}^{-1}$; SN20 = sodium nitrate at $20 \text{ kg N ha}^{-1} \text{ yr}^{-1}$.

113 At the PWL site, P was added once in April 2014. The PWL site was grazed during the study, at a rate
 114 of $1.0 \text{ sheep ha}^{-1}$. The PEN site was not grazed during this study, but had been grazed by sheep
 115 between 1990 and 2007 at two levels ($1.0 \text{ sheep ha}^{-1}$ and $1.5 \text{ sheep ha}^{-1}$). Although there was no
 116 grazing at the PEN site between 2007 and 2015, these moorland habitats are resistant to succession
 117 and there was no change in habitat type. Three replicate blocks were established in each of two
 118 paddocks that had been grazed at the different levels, but the lack of grazing during the intervening
 119 seven years had greatly diminished effects of the grazing treatments and differences in previous
 120 stocking rate were not taken into account in the design.

121 1.2.2. VEGETATION ANALYSIS

122 Vegetation data were collected for PWL in June 2014 and June 2015 and for PEN in June 2015.
123 Cover was recorded visually for each species within each 3 x 3 m plot using the Domin scale and
124 subsequently transformed to percentage cover for statistical analysis (Currall 1987). All plants and
125 bryophytes were identified to species level. Vegetation height was recorded as the average of five
126 measurements to the top of the canopy taken using a metric sward stick, of ca. 1 cm diameter, marked
127 at 0.5 cm intervals (Dennis *et al.* 2005). At PEN, heights were taken from the centre of each plot and
128 then from the mid-point between the centre and each corner. At PWL, heights were recorded from
129 within small grazing exclosures ($\sim 1 \text{ m}^2$) that were established within each plot.

130 1.2.3 SOIL ANALYSIS

131 Soil samples for both sites were collected in June 2015 from five locations within each treatment plot,
132 using a 20 mm diameter soil corer up to a depth of 20 cm, and bulked together. The samples were air
133 dried and passed through a 2 mm sieve. Soil pH was measured with a Hydrus 400 meter (Fisherbrand,
134 Leicestershire, UK) in a slurry of 10 g fresh soil in 25 mL water. Total C was established by the
135 Dumas combustion method using an elemental analyser (Vario MAX Cube – Elementar
136 Analysensysteme, Hanau, Germany).

137 1.2.4 SOIL FLUX

138 Fluxes of CO_2 and CH_4 were measured for the different nutrient addition treatments at both PWL and
139 PEN from September 2014 to August 2015; initially every two months until March 2015, then every
140 month during the growing season (May to August 2015), under dark conditions using a non-steady
141 state, static chamber approach (Livingston & Hutchinson 1995; Parkin & Venterea 2010). Measured
142 CO_2 flux thus represents both soil and plant respiration (ecosystem respiration) but excludes the
143 effects of photosynthesis. Soil collars, 25 cm in diameter, were installed on each treatment in July
144 2014, two months prior to the first gas flux measurements to reduce disturbance. A soil knife was
145 used to cut through the initial vegetation and topsoil layer, before the collar was inserted to a depth of
146 7 cm, leaving 5 cm aboveground for chamber attachment. The location of the collar was selected
147 randomly for each treatment square and once installed, vegetation was trimmed from inside the collar
148 and maintained trimmed throughout sampling. Chamber design followed the description by Parkin &
149 Venterea (2010). Each chamber measured 19 cm in height including the soil collar and had an internal
150 volume of 9.33 L. A modified pressure vent tube (Xu *et al.* 2006) was installed to allow internal and
151 ambient air pressure equilibration. Before each measurement, chambers were placed carefully on
152 collars, ensuring a gas tight seal. The soil CO_2 and CH_4 fluxes were calculated based on changes in
153 chamber concentrations over 30 minutes. Measurements were taken for each chamber at 0, 15 and 30
154 minutes from chamber attachment (Parkin & Venterea 2010). The duration of measurement at each
155 sampling point was 60 seconds. CO_2 and CH_4 concentrations were measured using an LGR™, Ultra-
156 Portable Greenhouse Gas Analyser (Model 915–0011, Los Gatos Research, Palo Alto, CA, USA)

157 with a 1 Hz sampling rate. Chambers were attached to the analyser via inlet and outlet tubing (2 m by
158 4 mm internal diameter). Chamber internal temperature, ambient air temperature and soil temperature
159 at a depth of 10 cm were also recorded. The rate of change in gas concentration inside the chambers
160 was established with linear regression. Gas flux (CO₂ and CH₄) was then calculated from the rate of
161 gas concentration change using the following equation (McEwing et al. 2015):

$$162 \quad F_0 = S \frac{V M 273.16}{A V_m (273.16 + T)} 60$$

163 Where:

- 164 F_0 = Flux ($\mu\text{g CH}_4/\text{CO}_2\text{-C m}^{-2} \text{ hr}^{-1}$)
165 S = Rate of change in CH₄ and CO₂ concentration (ppm min^{-1})
166 V = Chamber volume (m^3)
167 A = Chamber area (m^2)
168 M = Molecular mass of CH₄/CO₂ (g mol^{-1})
169 V_m = Ideal gas mole volume ($0.0224 \text{ m}^3 \text{ mol}^{-1}$)

170

171 Each regression plot was assessed individually using R^2 as an indicator of accuracy, and plots with R^2
172 >0.7 were accepted for analysis. Low fluxes for CH₄ typically give a low R^2 , but should be included
173 to avoid over-estimation of mean flux (Alm et al. 2007), so CH₄ fluxes where $R^2 <0.7$, but where
174 measurements did not exceed 0.3 ppm, were also retained as zero measurements. Two collars at PWL
175 suffered from excessive water pooling, which visibly affected the vegetation and soil. These were
176 removed from the dataset before analysis. Net ecosystem exchange was not considered as part of this
177 study, so whilst the results demonstrate differences in C flux, primary productivity was not measured
178 and thus the full influence of N and P availability on C budget cannot be determined.

179 1.2.5 STATISTICAL ANALYSIS

180 All variables were tested visually for normality and homoscedasticity with Levene's test prior to
181 statistical analysis. Data for CO₂ and CH₄ fluxes were Log (x+1) transformed to meet the assumptions
182 of analyses. Differences in CO₂ and CH₄ flux between treatments were analysed with linear mixed
183 models specifying repeated measures, with treatment as a fixed factor and air temperature as covariate
184 to allow for diurnal/seasonal variance. When significant differences between treatments were detected
185 ($P < 0.05$), post-hoc tests were conducted using LSD pairwise comparisons. The relationship between
186 GHG flux and air temperature was initially tested for significance with simple linear regression before
187 inclusion in the model. Air temperatures were used in the analysis since data for soil temperature were
188 only captured from January onwards, but the two measurements were shown to be correlated (PWL:
189 $R^2 = 0.52$, $P = <0.001$; PEN: $R^2 = 0.81$, $P = <0.001$). Differences in mean annual soil emission of CO₂
190 and CH₄ between experimental treatments were also investigated with one-way analysis of variance
191 (ANOVAs). LSD pairwise comparisons were used to further investigate individual relationships.

192 Multiple regression analysis with all-possible-subsets model procedure was used to investigate the
 193 relationships between annual average GHG flux and environmental parameters: soil pH, vegetation
 194 height, and cover values for each of five plant functional types. All statistical analyses were conducted
 195 with SPSS 21.0 (IBM SPSS Statistics for Windows, 2012).

196 1.3 RESULTS

197 1.3.1 ENVIRONMENTAL CONTROLS ON CO₂ AND CH₄ FLUX

198 Over the sampling period, fluctuations were observed for air and soil temperatures in line with
 199 seasonal variation. The highest air temperatures were recorded in June for PWL (22.3 °C) and in
 200 August for PEN (19.4 °C). Air temperatures were lowest at both locations in January (PWL = 5.0 °C,
 201 PEN = 6.2 °C). Soil temperatures were similar with the highest values recorded in August for PWL
 202 (13.2 °C) and in July for PEN (12.9 °C) and lowest vales for both in January (PWL = 4.6 °C, PEN =
 203 3.5 °C). Over the twelve months of sampling, the effect of changing air temperature on CO₂ flux was
 204 significant for both sites, with positive relationships observed between increasing temperature and
 205 CO₂ emission (PWL: R² = 0.68, P = <0.001; PEN: R² = 0.70, P = <0.001, Figure 2). The effect of
 206 temperature on CH₄ flux was less consistent; no relationship was observed at PWL (R² = 0.004, P =
 207 0.37), but a significant positive relationship was recorded at PEN (R² = 0.033, P = 0.012).

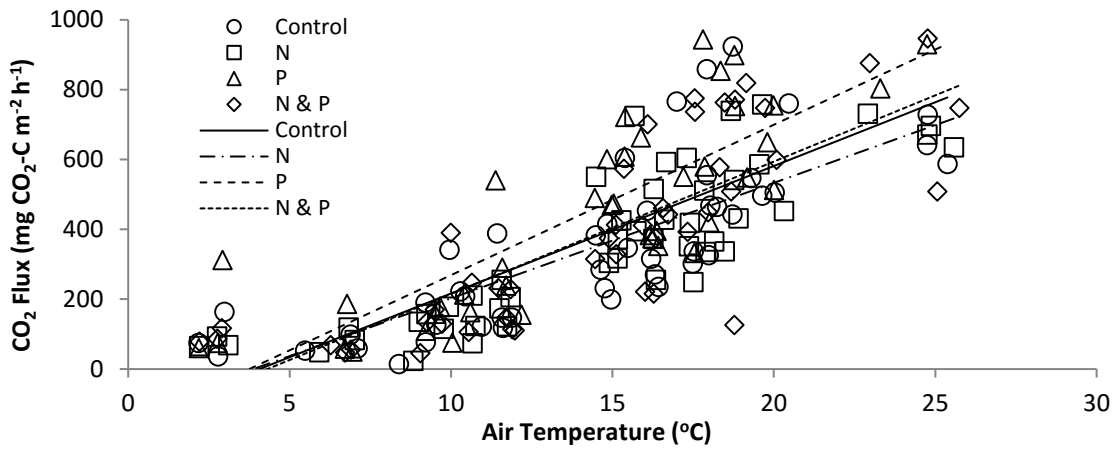
208 The results of the multiple regression analysis of effects of vegetation cover values, vegetation height
 209 and soil pH on mean annual CO₂ and CH₄ fluxes are shown in Table 1. This analysis revealed
 210 significant relationships between soil pH and fluxes for both CO₂ and CH₄ at PWL (Table 1, Figure
 211 3), but other factors were shown to have no effect. At PEN, no environmental factor was shown to
 212 have an effect on either CO₂ or CH₄ flux despite the significant differences in vegetation cover
 213 between treatments as a result of long-term nutrient addition (Figure 6).

214 Table 1. Relationships between gas flux and environmental variables with potential influence on C
 215 emission.

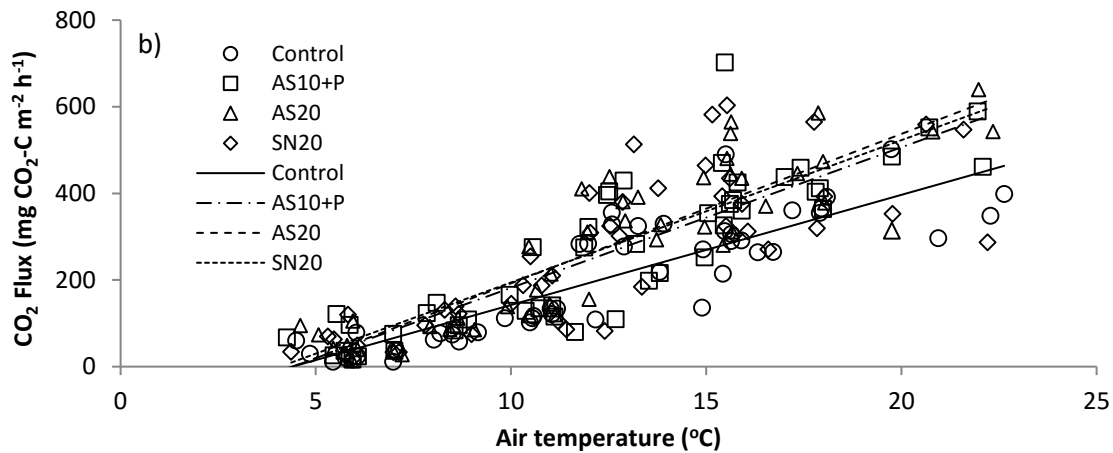
Site	Variable	Regression model	Degrees of freedom	R ²	P
PWL	CO2	pH	1,21	0.52	<0.001
	CH4	pH	1,21	0.25	0.019
PEN	CO2	-	1,23	-	ns
	CH4	-	1,23	-	ns

216 *P*-values represent the chance that the regression slope is not different from 0; bold highlights results
 217 significant at *P* < 0.05. The best fitting regression model using all-possible-subsets regression analysis
 218 is shown for each parameter. Variables included in the analysis were: ground cover of each plant
 219 functional type (graminoid, cryptogam, dwarf shrub, forbs), vegetation height and soil pH.

220



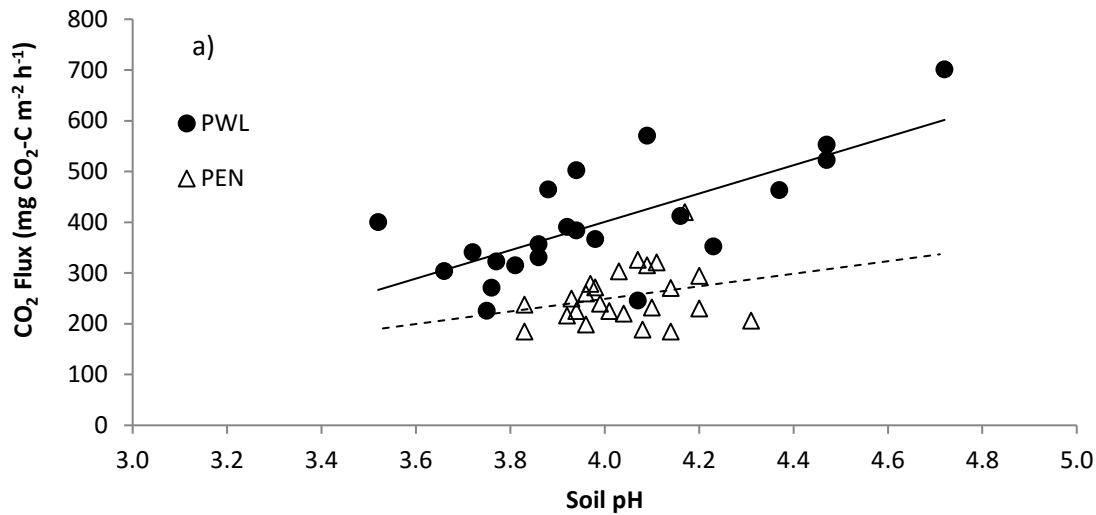
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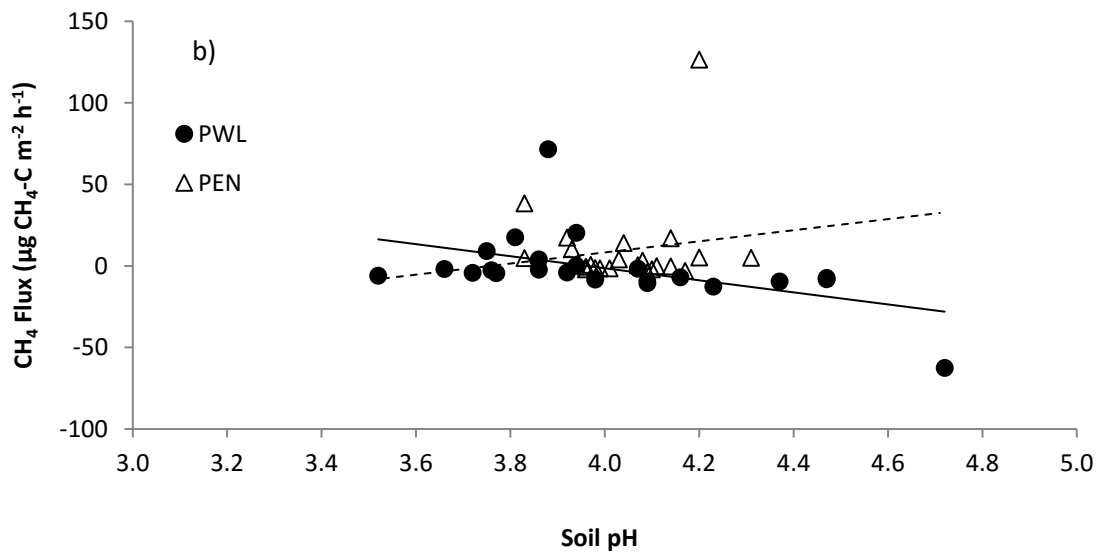
222

223 Figure 2. Relationships between CO₂ flux and air temperature at two experimental sites: a) PWL:
 224 Control = no nutrient addition (circles); N = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ (squares); P =
 225 phosphorus (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹ (triangles); N & P = ammonium
 226 nitrate at 60 kg N ha⁻¹ yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹ (diamonds); and b) PEN: Control = no
 227 nutrient addition (circles); AS10+P = ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20 kg P
 228 ha⁻¹ yr⁻¹ (squares); AS20 = ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹ (triangles); SN20 = sodium nitrate
 229 at 20 kg N ha⁻¹ yr⁻¹ (diamonds).

230



231



232

233 Figure 3. The relationship between soil pH and a) CO₂ flux and b) CH₄ flux at two sites: PWL (black
 234 circles) and PEN (white triangles). Solid lines indicate significant relationships, dashed lines non-
 235 significant relationships.

236 1.3.2 EFFECTS OF NUTRIENT ADDITION ON CO₂ FLUXES

237 Differences were observed in the CO₂ fluxes between nutrient addition treatments at both sites, but
 238 with differences between sites in the size of response observed. For PWL, the between-treatment
 239 differences (Table 2, Figure 4a) were shown by *post hoc* analysis to be driven by the largest mean
 240 annual flux recorded for the P addition treatment (464 mg CO₂-C m⁻² h⁻¹). The P treatment had
 241 consistently higher CO₂ flux in all but the first month sampled (September) (Table 2, Figure 4a). The
 242 N addition treatment had the lowest mean annual flux (353 mg CO₂-C m⁻² h⁻¹) and was consistently
 243 the lowest flux recorded across the sampling period (in all months except May). The CO₂ fluxes in the
 244 control and N + P treatments had similar mean annual fluxes (386 and 413 mg CO₂-C m⁻² h⁻¹
 245 respectively). For PEN, *post hoc* analysis revealed the differences between treatments (Table 2,

246 Figure 4b) to be due to the lowest mean annual flux recorded, which was the control (205 mg CO₂-C
 247 m⁻² h⁻¹). The AS10+P, AS20 and SN20 treatments had similar mean annual fluxes (261, 277 and 274
 248 mg CO₂-C m⁻² h⁻¹ respectively). The differences observed between treatments at PEN suggest that
 249 nutrient enrichment in the long term significantly increases CO₂ flux relative to control, but there were
 250 no statistically significant differences among the nutrient addition treatments, suggesting limited
 251 influence of the P added to the AS10+P treatment on CO₂ flux 15 years after application.

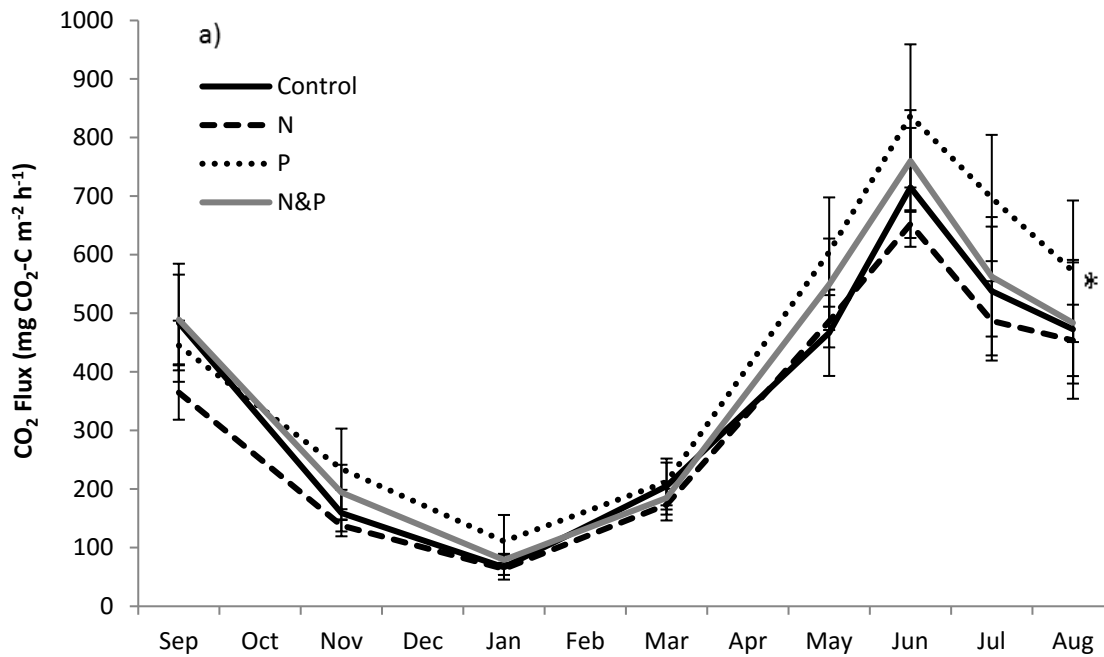
252

253 Table 2. Summary of results from Linear Mixed Models describing soil CO₂ and CH₄ flux responses
 254 to nutrient addition treatments, with treatment as fixed factor and air temperature as covariate.

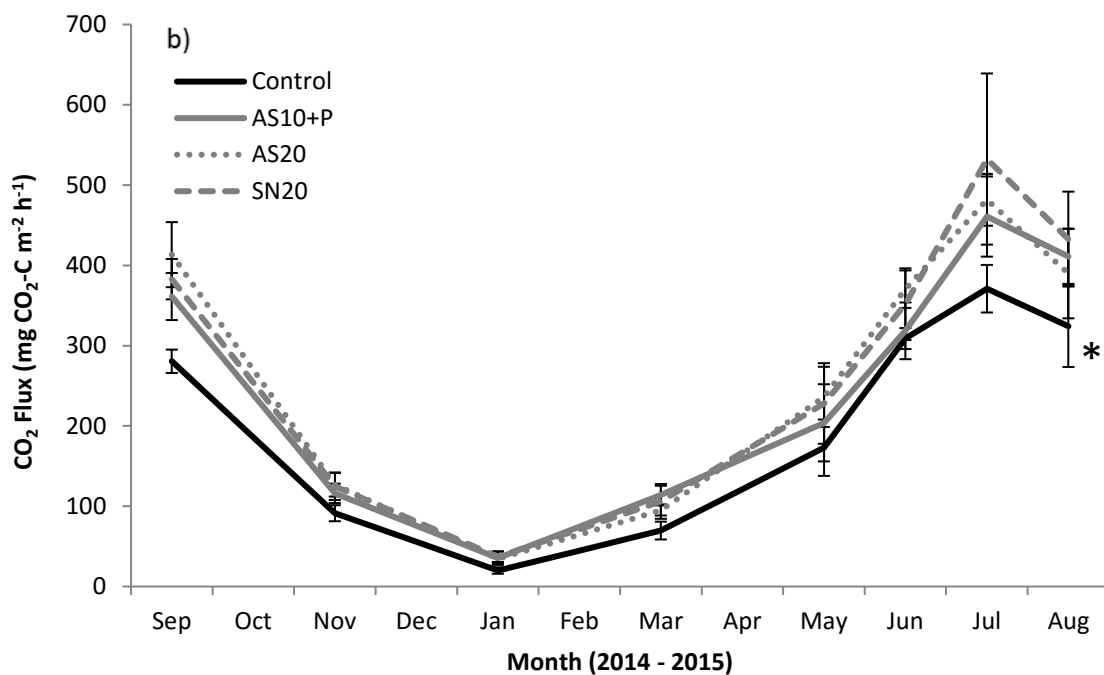
Site	Variable	Factor	Degrees of freedom	<i>F</i>	<i>P</i>
PWL	CO ₂	Treatment	3, 171	3.69	0.013
		Air temperature	1, 171	422.32	<0.001
	CH ₄	Treatment	3, 171	2.86	0.041
		Air temperature	1, 171	0.60	0.439
PEN	CO ₂	Treatment	3, 187	4.63	0.004
		Air temperature	1, 187	461.23	<0.001
	CH ₄	Treatment	3, 187	4.22	0.006
		Air temperature	1, 187	6.75	0.01

255 *P*-values refer to of the likelihood that the coefficient is zero, bold highlights results significant at *P*
 256 <0.05.

257



258

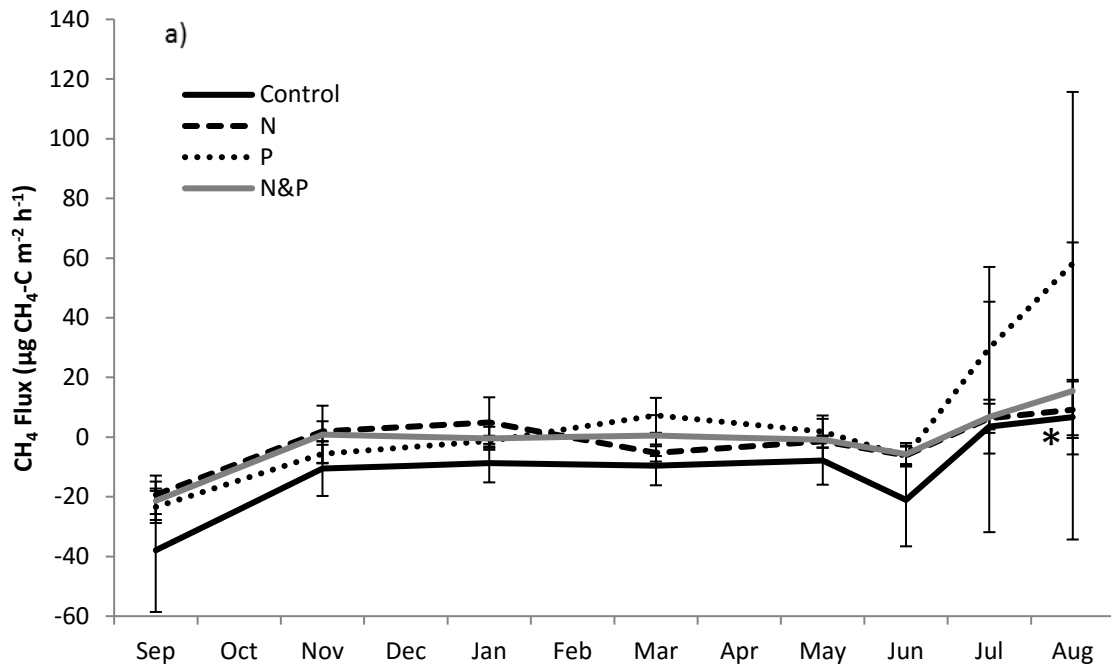


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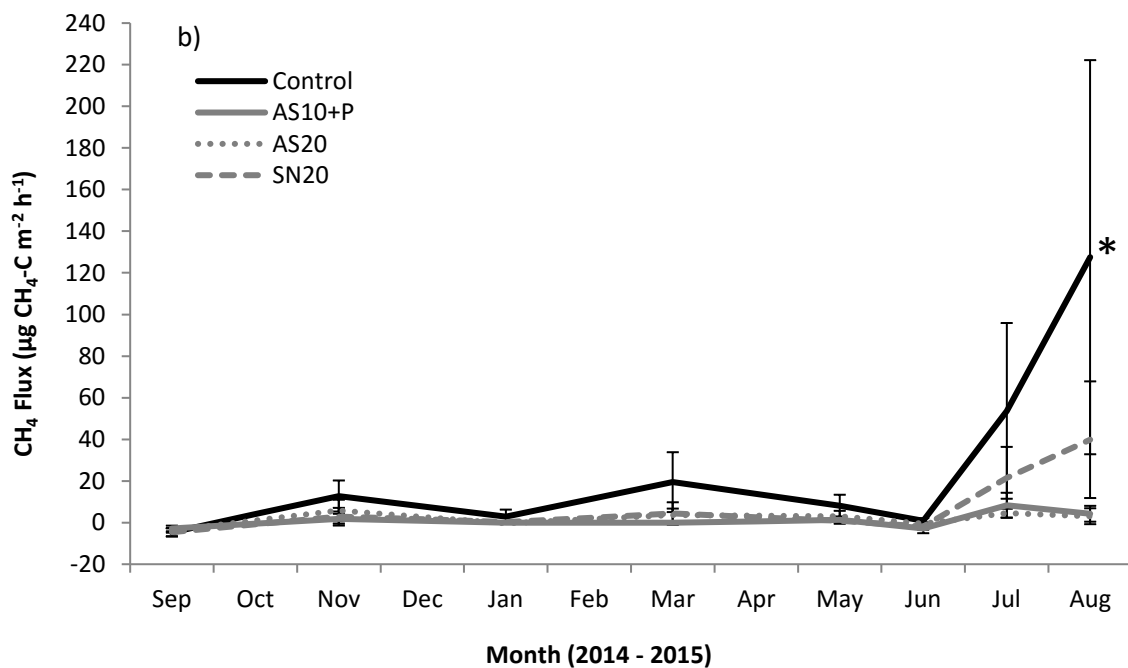
261 Figure 4. Net CO₂ emission fluxes for a) PWL and b) PEN. Error bars denote standard error. *
 262 indicates treatment with significant difference as determined by *post hoc* pairwise comparison (LSD).
 263 PWL: Control = no nutrient addition; N = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹; P = phosphorus
 264 (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹; N & P = ammonium nitrate at 60 kg N ha⁻¹
 265 yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹. PEN: Control = no nutrient addition; AS10+P = ammonium
 266 sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20 kg P ha⁻¹ yr⁻¹; AS20 = ammonium sulphate at 20 kg N
 267 ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹ yr⁻¹.

268



269

270



271

272 Figure 5. Effects of nutrient addition treatments on net methane emissions for a) PWL and b) PEN
 273 sites. Error bars denote standard error. * indicates treatment with significant difference as determined
 274 by *post hoc* pairwise comparison (LSD). PWL: Control = no nutrient addition; N = ammonium nitrate
 275 at 60 kg N ha⁻¹ yr⁻¹; P = phosphorus (sodium dihydrogen orthophosphate) at 40 kg P ha⁻¹ yr⁻¹; N & P =
 276 ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹. PEN: Control = no nutrient
 277 addition; AS10+P = ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20 kg P ha⁻¹ yr⁻¹; AS20 =
 278 ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹ yr⁻¹.

279

280 1.3.3 EFFECTS OF NUTRIENT ADDITION ON CH₄ FLUXES

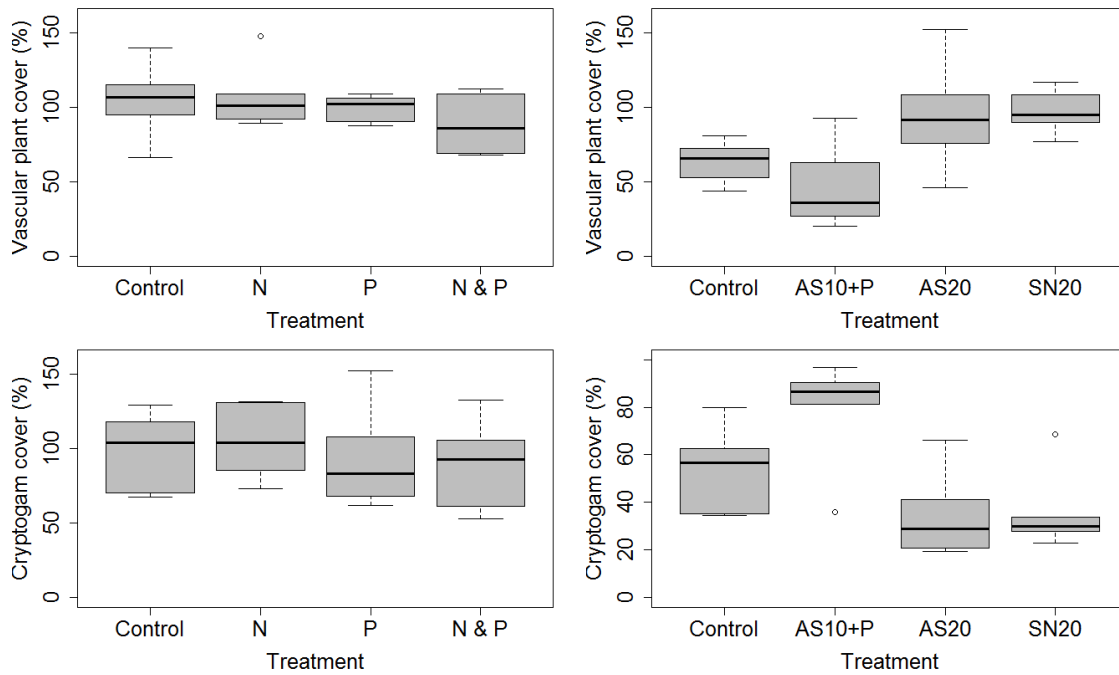
281 The extent to which differences in CH₄ fluxes could be attributed to treatment effects varied
282 according to experimental site. At PWL there were significant differences between treatments (Table
283 2, Figure 5a), independent of seasonal variations in temperature (Table 2), which *post hoc* tests
284 revealed to be driven by the difference in CH₄ emissions between the control (no addition) and
285 nutrient addition treatments (N, P and N + P). At this site, CH₄ uptake was greatest in control plots,
286 with this treatment behaving as a small sink (on average -10.68 μg CH₄-C m⁻² h⁻¹), only emitting CH₄
287 during the last two months (June and August), at rates just above zero μg C CH₄ m⁻² h⁻¹. The N, P and
288 N + P treatments were highly variable in emission rate, with fluctuations recorded across the sampling
289 range. The P addition treatment had the largest mean annual flux (9.00 μg CH₄-C m⁻² h⁻¹), which can
290 chiefly be attributed to an increase in emission rate during the last two months (Figure 5a). The N and
291 N + P treatments were intermediate in response, with emission rates just below zero (N = -1.05 μg
292 CH₄-C m⁻² h⁻¹, N + P = -0.77 μg CH₄-C m⁻² h⁻¹) on a mean annual basis (*P* < 0.05). At PEN there were
293 also significant differences observed between treatments (Table 2, Figure 5b), which *post hoc* analysis
294 revealed to be likewise driven by the difference between control and all nutrient addition treatments.
295 However, at this site the CH₄ flux was significantly higher in control plots than nutrient added
296 treatments, with mean annual emissions of 27.64 μg CH₄-C m⁻² h⁻¹ (*P* < 0.01). All of the nutrient
297 addition treatments at PEN otherwise had similar flux rates for the full sampling range, until the last
298 two months, when the CH₄ flux from the sodium nitrate treatment (SN20) increased comparatively
299 (Figure 5b).

300 1.3.4 EFFECTS OF NUTRIENT ADDITION ON VEGETATION COMPOSITION

301 The addition of nutrients resulted in significant differences in the observed vegetation cover, but only
302 after sufficient time had elapsed allowing vegetation community modification. At PWL, where
303 nutrients were added in the short term only, no differences in vegetation cover were observed. At
304 PEN, long-term nutrient additions resulted in greater graminoid cover and lesser cryptogam cover
305 where N alone was added and greater cryptogam and lesser graminoid cover in treatments where P
306 was added (AS10+P). Control plots were intermediate in composition (Figure 6).

307 PWL

PEN



308

309

310 Figure 6. Ground cover of vascular plants and cryptogams recorded for nutrient addition treatments at
 311 two sites, PWL (L-hand plots) and PEN (R-hand plots). PWL: Control = no nutrient addition; N =
 312 ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹; P = phosphorus (sodium dihydrogen orthophosphate) at 40 kg
 313 P ha⁻¹ yr⁻¹; N & P = ammonium nitrate at 60 kg N ha⁻¹ yr⁻¹ + phosphorus at 40 kg P ha⁻¹ yr⁻¹. PEN:
 314 Control = no nutrient addition; AS10+P = ammonium sulphate at 10 kg N ha⁻¹ yr⁻¹ + phosphorus at 20
 315 kg P ha⁻¹ yr⁻¹; AS20 = ammonium sulphate at 20 kg N ha⁻¹ yr⁻¹; SN20 = sodium nitrate at 20 kg N ha⁻¹
 316 yr⁻¹. Box midline indicates median, box edges indicate interquartile range. Whiskers indicate range of
 317 data within 1.5 x the interquartile range; points indicate data outside 1.5 x the interquartile range.

318 1.4 DISCUSSION

319 The addition of N and P had significant influence on CO₂ and CH₄ fluxes, with differences in the
 320 direction and magnitude of the effects between sites as a consequence of duration of nutrient addition
 321 treatment. In the short-term trial, N addition inhibited CO₂ and CH₄ flux, whereas in the long-term
 322 trial, N addition significantly increased CO₂ emissions but inhibited CH₄ flux. The addition of P
 323 significantly increased CO₂ and CH₄ flux in the short-term trial, but this effect was reduced in the
 324 long-term trial, after 14-15 years since P addition.

325 1.4.1 NUTRIENT ADDITION EFFECTS ON CO₂ FLUXES

326 The addition of N and P influenced fluxes of CO₂, however there were substantial differences in effect
 327 between treatments, and also differences in response between short-term (PWL = 1 year for N & P)
 328 and long-term nutrient additions (PEN = additions between 1996 – 2012, sampling undertaken three
 329 years after the cessation of N addition and 15 years after a single P application). At PWL, the addition
 330 of P increased CO₂ flux, which supported hypothesis H3 and is in line with findings from similar
 331 studies in other environments (Cleveland & Townsend 2006; Liu et al. 2013). Phosphorus limitation
 332 is a significant mechanism constraining ecosystem processes, particularly in systems suffering from

333 the effects of chronic N deposition (Cleveland et al. 2011; Crowley et al. 2012). The addition of P
334 stimulates decomposition, with observed rises in CO₂ flux associated with increases in heterotrophic
335 (bacteria and fungi) biomass and activity, and thus respiration (Liu et al. 2012), rather than through
336 increased fine-root biomass (Cleveland & Townsend 2006). Phosphorus addition results in the
337 modification of microbial community structure and reduces the ratio between Gram-positive and
338 Gram-negative bacteria, resulting in more copiotrophic communities (Fanin et al. 2015). At PEN, the
339 role of P was less clear. The CO₂ flux in the N + P treatment (AS10+P) was larger than in the control
340 plots, but was mostly lower across the sampling period than CO₂ fluxes from the two N-only
341 treatments (AS20 and SN20). Phosphorus is highly persistent in soil (Nye & Tinker 1977), and was
342 presumably responsible for driving shifts observed in vegetation species composition (Figure 6), thus
343 the comparatively weak effect of P on CO₂ flux at this site was unexpected. The N-only treatments
344 received N at a higher rate than the AS10+P treatment, and associated larger CO₂ fluxes may reflect
345 the effect of greater N availability. In addition, while the effects of P on plant species composition
346 were still visible at the site, P may no longer have been stimulating plant production of labile C. Liu et
347 al. (2013) observed a diminishing effect of P on microbial biomass after four years, which was
348 attributed to C limitation, where the exhaustion of available soil C had a constraining effect on
349 respiration (Fanin et al. 2015).

350 The addition of N had an inhibitory effect on soil CO₂ flux at PWL and refutes hypothesis H1 that
351 nutrient addition would increase CO₂ efflux. Similar inhibition of decomposition by N has been found
352 in other studies (Ramirez et al. 2010; Chen et al. 2015), and may result from high N availability. In
353 soils where N is not a limiting factor for microbial growth and activity, the addition of N can constrain
354 organic matter decomposition (Janssens et al. 2010). In such instances, reductions in CO₂ flux may be
355 driven by shifts in C allocation from belowground to aboveground biomass (Litton et al. 2007), which
356 reduces rhizosphere and microbial respiration (Phillips & Fahey 2007; Bae et al. 2015). The
357 background N deposition for these sites is 22 kg N ha⁻¹ yr⁻¹ (Emmett et al. 2007), which is greater than
358 the critical load limit for this habitat (10 – 15 kg N ha⁻¹ yr⁻¹; APIS 2014), suggesting that these soils
359 are unlikely to be N limited. In addition, reductions in soil pH associated with N addition potentially
360 have a limiting effect (Chen et al. 2015).

361 This contrasts with the effects observed at PEN, where N addition treatments had higher CO₂ flux in
362 comparison to control (no addition) treatments, which does support hypothesis H1. PEN was sampled
363 three years after ceasing N addition, which could indicate that inhibitory effects were reduced due to
364 N losses from leaching in the intervening three years. This explanation was deemed unlikely however,
365 as the size and activity of microbial populations exposed to N enrichment can remain elevated 6–8
366 years after cessation of additions, suggesting a prolonged effect on the rate of nutrient cycling (Power
367 et al. 2006). Instead, the higher CO₂ flux where N was added is most likely the result of greater soil
368 organic matter decomposition from increased microbial abundance and activity (Mack et al. 2004;

369 Bragazza et al. 2006), and by reduced production of more decomposition resistant species (Bragazza
370 et al. 2012). This influence was not translated to a reduction in soil C levels in N treatment plots
371 however; these plots in fact had the highest soil C content, albeit not significantly greater than control
372 (Stiles et al. 2017). This suggests a potential balance between increased rates of both plant production
373 and decomposition with greater N availability (Mack et al. 2004).

374 The variability observed in CO₂ flux between sites with N enrichment could be explained by
375 differences in the N addition rate, which at PWL was three times the maximum annual application
376 rate at PEN. Janssens et al. (2010) demonstrated that the inhibitory effect of N on soil CO₂ flux was
377 larger in sites receiving N at dose rates greater than 50 kg ha⁻¹ yr⁻¹, which may account for some of the
378 between-site disparity. The inhibitory effect of N at PWL apparently offset the stimulatory effect of P,
379 in that the CO₂ flux response to N + P addition was intermediate between responses to P and N
380 addition. Thus, although the results were not consistent with a general stimulation of CO₂ efflux by
381 nutrient addition (H1), there was support for hypothesis H3 that P would have a greater stimulating
382 effect.

383 1.4.2 NUTRIENT ADDITION EFFECTS ON CH₄ FLUXES

384 The addition of N and P altered CH₄ flux, although different effects were observed in the short term
385 after enrichment (PWL) and after a longer period (PEN). The CH₄ flux results represent net emissions,
386 and effects on CH₄ production and oxidation cannot be distinguished. At PWL, all nutrient addition
387 treatments had less negative net CH₄ emissions than the control, supporting hypothesis H2. The less
388 negative net emissions with N addition are presumably due to decreases in the rate of CH₄ oxidation
389 (Aerts & Toet 1997; Aerts & Caluwe 1999). This is primarily caused by competition for the CH₄
390 mono-oxygenase enzyme, which is affected when N addition increases rates of nitrification, inhibiting
391 CH₄ oxidisation rate and reducing the amount of methane consumed by methanotrophs (Bodelier
392 2011). Phosphorus addition also increased the rate of emission for CH₄, which supports hypothesis
393 H4, but is contrary to similar research conducted in other environments (Zhang et al. 2011; Song et al.
394 2012). These studies attributed the effects observed to the inhibition of methanogenesis, the
395 stimulation of methanotrophic potential, and increased plant growth and water uptake reducing soil
396 water content and thus increasing methanotrophy through greater aeration. This last effect would
397 seem unlikely to operate in the PWL and PEN experiments because these are rather wet sites and the
398 treatments would not be expected to increase soil aeration greatly, even with the potentially
399 augmented growth associated with P enrichment. The stimulation of methanogenesis is likely to be
400 the principal mechanism and has been previously observed for some peat types (Keller et al. 2006),
401 which suggests the methanogen community at PWL is P limited. Further tests would be necessary to
402 rule out any inhibitory effect of P on methane oxidation, but this is beyond the scope of the current
403 study.

404 At PEN, the highest CH₄ flux recorded was in the control, which was contrary to hypothesis H2 that
405 nutrient addition would stimulate CH₄ fluxes. Fluxes of CH₄ from the nutrient addition treatments
406 remained near zero for the majority of the sampling period until the last two months (July and
407 August), when the fluxes from the sodium nitrate treatment (SN20) increased substantially. Nitrate
408 inhibits CH₄ emission, but only in very high concentrations (Bodelier & Laanbroek 2004), whereas
409 ammonium has been shown to inhibit CH₄ emission more strongly (Crill et al. 1994). This may
410 explain why emissions were consistently inhibited in the ammonium treatments (AS10+P and AS20)
411 in contrast to the spike in emission observed in the nitrate treatment towards the end of the study
412 period. Overall, N addition inhibited CH₄ emission, which can be attributed to N-induced increases in
413 population size and activity of methane-oxidising bacteria (Bodelier & Laanbroek 2004). The addition
414 of P (AS10+P) had no obvious effect at PEN, 15 years after application, contrary to the H4
415 hypothesis. As with the CO₂ flux results, this was unexpected. Whilst the exact mechanisms remain
416 unclear, the composition of the microbial community may have changed over the intervening years
417 (as with the diminishing effect of P over time on CO₂ emission, discussed above) which may have
418 reduced the size of any initial effect. This may also account for the substantial differences observed in
419 the results of the two experiments.

420 1.4.3 ENVIRONMENTAL CONTROLS ON CARBON FLUXES

421 Temperature influenced the rate of CO₂ flux at both PWL and PEN, which is consistent with other
422 studies (Dorrepaal et al. 2009; Briones et al. 2010; Imer et al. 2013). Higher temperatures allow
423 increased rates of soil decomposition and root respiration, which are temperature-dependent chemical
424 and biochemical reactions (Knorr et al. 2005; Davidson & Janssens 2006). The role of temperature in
425 controlling CH₄ emission is less definite and was variable between sites. At PEN, there was a positive
426 effect of temperature on CH₄ flux, whereas at PWL, no effect was observed. The effect of temperature
427 on CH₄ production is reported to be variable and driven by the influence of site-specific factors such
428 as differences in soil type (Chin et al. 1999; Van Winden et al. 2012). The effect of temperature is
429 likely to be closely linked with that of soil moisture and aeration (Smith et al. 2003). If high
430 temperatures relate to drought and therefore dry soil, it would be expected that CH₄ emission would
431 be reduced, but CO₂ emission would increase (Sundh et al. 2000). Plant functional type composition
432 and above ground biomass quantity did not influence CO₂ or CH₄ fluxes, which is contrary to
433 previous studies (McNamara et al. 2008; Green & Baird et al. 2012; Cooper et al. 2014; McEwing et
434 al. 2015). This was also contrary to expectation, as differences in vegetation composition exist
435 between treatments, particularly at PEN, where long-term nutrient additions have resulted in divergent
436 vegetation composition (Figure 6). Nitrogen addition (AS20 and SN20 treatments) has driven an
437 increased ground cover of graminoid species, whereas N + P addition (AS10+P) has driven an
438 increased ground cover of cryptogam species. Vegetation impacts on CH₄ flux are particularly
439 associated with *Eriophorum* spp. presence and cover (Green & Baird et al. 2012), as these sedges

440 have aerenchymatous tissue, which can facilitate the transport of CH₄ from the anaerobic zone to the
441 atmosphere, and actively produces substrates that encourage methanogenesis (Cooper et al. 2014).
442 The relative scarcity of these species at either site could potentially account for the limited effect of
443 vegetation composition changes on CH₄ flux.

444 The positive relationship for CO₂, and negative relationship for CH₄, observed between fluxes and soil
445 pH at PWL, can be explained by changes in microbial activity. Soil acidity may be increased by N
446 enrichment (Bobbink et al. 2010; Phoenix et al. 2012), an effect which is reported to have a stronger
447 controlling effect on C flux than stimulation by increased N availability (Chen et al. 2015). Soil pH
448 can affect the activity rate and composition of microbial communities, with decreases of 1.5 pH units
449 shown to reduce activity by 50% (Fernández-Calviño & Bååth 2010). Recruitment of bacterial species
450 more tolerant of lower pH ultimately modifies the community, but with an intervening lag in activity
451 (Rousk et al. 2010), a hysteresis which may explain the observed reduction in the CO₂ emission. Low
452 soil pH may reduce CH₄ emissions through inhibition of methanotrophic bacterial communities
453 (Dedysh 2002). These effects potentially explain some of the differences observed between the two
454 sites, in that changes to pH driven by recent nutrient additions at PWL may have resulted in short-
455 term fluctuations in microbial activity. This effect would be expected to reduce over time as the
456 system shifts to a new equilibrium, as in the longer-term experiment at PEN. The absence of any
457 association with other environmental variables at PEN indicates that, for this study at least, the
458 availability of nutrients over an extended period (even three years after application for N and 15 years
459 after application for P) has a greater influence on CH₄ flux rate than other background environmental
460 factors.

461 **1.5 CONCLUSION**

462 The addition of N and P had significant effects on CO₂ and CH₄ fluxes but the direction and
463 magnitude of effects were different in the two experiments, most probably due to the difference in
464 duration of additions between the two sites. The CO₂ flux was inhibited by recent additions of N, but
465 plots that had received prolonged N additions up until three years previously showed considerably
466 increased CO₂ emissions. The long-term effects of P addition are less clear. In the short term, P
467 stimulated CO₂ emissions via release from nutrient limitation, but this effect appeared to diminish
468 with time, perhaps due to a decrease in readily available C substrates. This implies a negative effect of
469 P addition on soil C storage, although this may be offset at some sites by plant productivity increases.
470 The emission rate of CH₄ for treatments where nutrients were added was of similar size between sites,
471 despite differences in the CH₄ flux observed in control treatments at either site. This demonstrated a
472 similar pattern of modification by nutrient addition, but with different factors controlling the response
473 relative to either control. The differences between emission rates in control (no addition) treatments at
474 the two sites indicates the influence of other site characteristics on CH₄ flux rate.

475 The contrast in results from the two experiments demonstrates the differences between initial
476 ecosystem responses to nutrient addition and responses after ecosystem processes and microbial
477 assemblages have had sufficient time to adjust to more nutrient-rich conditions. Perturbation in
478 ecosystems often drives gradual change, and many experiments do not last long enough for the system
479 to reach a new equilibrium (Bubier et al. 2007; Vitousek et al. 2008). In this study the differences in
480 response between experimental additions in the short and long term were considerable, which has
481 important implications for the use of short-term evidence to infer responses of C storage and other
482 ecosystem processes to continuing atmospheric N pollution, or P additions that persist in the soil.
483 Short-term studies investigating impacts of nutrient enrichment may misrepresent longer-term
484 ecosystem responses.

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