

Differences between the 2018 and 2019 stratospheric polar vortex split events

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3 **Differences between the 2018 and 2019**
4 **stratospheric polar vortex split events**

5 **Amy H. Butler^{1*} | Zachary D. Lawrence^{2,3*} | Simon H. Lee^{4*} | Samuel P. Lillo^{2,3*} | Craig S. Long^{5*}**

¹NOAA Chemical Sciences Laboratory, 325 Broadway, Boulder, Colorado 80305

²Cooperative Institute for Research in Environmental Sciences (CIRES), University of Colorado, Boulder, Colorado 80309

³NOAA Physical Sciences Laboratory, 325 Broadway, Boulder, Colorado 80305

⁴Department of Meteorology, University of Reading, Reading, UK, RG6 6BB

⁵NOAA National Centers for Environmental Prediction (NCEP) Climate Prediction Center, 5830 University Research Drive, College Park, Maryland 20740

Correspondence

Amy H. Butler, CIRES and NOAA CSL, Boulder, CO, 80305, USA
Email: amy.butler@noaa.gov

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Two recent occurrences in February 2018 and January 2019 of a dynamic split in the Northern Hemisphere stratospheric polar vortex are compared in terms of their evolution and predictability. The 2018 split vortex was associated with primarily wavenumber-2 wave forcing that was not well predicted more than 7-10 days ahead of time, and was followed by persistent coupling to the surface with strong weather impacts. In 2019 the vortex was first displaced by slow wavenumber-1 amplification into the stratosphere, which was predictable at longer lead times, and then split; the surface impacts following the event were weaker. Here we examine the role of large-scale climate influences, such as the phase of the El Niño-Southern Oscillation, the Quasi-biennial Oscillation and Madden-Julian Oscillation, on the wave forcing, surface impacts, and predictability of these two events. Linkages between the forecast error in the stratospheric polar vortex winds with the forecast error in the Quasi-biennial Oscillation and Madden-Julian Oscillation are examined.

KEYWORDS

polar vortex, stratosphere, teleconnections, prediction, sudden stratospheric warming, stratosphere-troposphere coupling

* Equally contributing authors.

6 1 | INTRODUCTION

7 Disruptions of the wintertime Northern Hemisphere (NH) stratospheric polar vortex, marked by rapidly warming
8 temperatures and a reversal of the zonal-mean zonal winds from westerly to easterly, occur on average about every
9 other year (Butler et al., 2017). These so-called “sudden stratospheric warmings” (SSWs) involve the displacement of
10 the vortex off the pole, the split of the vortex into two vortices, or some combination of the two (Charlton and Polvani,
11 2007). These dramatic events 10-50 km above the Arctic are increasingly recognized as being potential sources of
12 sub-seasonal to seasonal (S2S) predictability of wintertime weather (e.g., Butler et al., 2019; Domeisen et al., 2020a).

13 Two recent SSWs occurred on February 12 2018 and January 2 2019 (based on the reversal of the daily-mean 10
14 hPa 60°N zonal-mean zonal wind in ERA-Interim reanalysis). In 2018 the vortex abruptly split in two, while in 2019
15 the vortex was first displaced and then split (Rao et al., 2019). While both disruptions eventually resulted in a split
16 vortex, there were important differences in how they evolved, their predictability, and their ultimate influence on the
17 troposphere in the following weeks (Karpechko et al., 2018; Rao et al., 2018, 2019, 2020). For example, while the
18 2018 SSW was predicted at lead times of 10 days or less (Lee et al. (2019); typical for mid-winter SSWs, see Domeisen
19 et al. (2020b)), the 2019 SSW could be predicted more than a week beyond that (Rao et al., 2020). This enhanced
20 predictability was linked to the ability of the S2S prediction systems to largely capture planetary-scale (wavenumber
21 1) wave driving in mid-December 2018, weeks prior to the 2019 split vortex (Rao et al., 2019).

22 Another difference in the SSWs was that while the 2018 event was followed by the canonical weather pattern
23 associated with the negative phase of the Northern Annular Mode (NAM) for up to 2 months (Greening and Hodgson,
24 2019), the 2019 event was followed by weaker and shorter-lived stratosphere-troposphere coupling, despite the split
25 vortex persisting for an extended period of time (Lee and Butler, 2019). The stronger amplitude of stratospheric wind
26 reversal in 2018 was statistically linked to increased probability of downward coupling to the surface (Rao et al., 2020).

27 The contrasting characteristics in evolution and predictability may be related to differences in large-scale circula-
28 tion patterns during these two winters (Rao et al., 2019). For example, in 2018 the tropical Pacific was in a moderate
29 La Niña state, while in 2019 there was a moderate El Niño. In general, poleward-propagating tropospheric planetary-
30 scale (i.e., wavenumbers $k=1-3$) “wave trains”, typically forced by tropical convective heating associated with the El
31 Niño-Southern Oscillation (ENSO) or the Madden-Julian Oscillation (MJO), can linearly interfere with the climatologi-
32 cal extratropical wave pattern (Domeisen et al., 2019; Garfinkel et al., 2012a). When this interference is constructive,
33 these large-scale waves can amplify into the extratropical stratosphere, leading to a weakening of the polar vortex
34 (Smith and Kushner, 2012; Fletcher and Kushner, 2013). These tropically-forced wave-trains can also modulate pat-
35 terns in the extratropical troposphere that precede SSWs (Barriopedro and Calvo, 2014). These so-called precursor
36 patterns are often different for split and displacement vortex events (Garfinkel et al., 2010; Cohen and Jones, 2011;
37 Attard et al., 2016).

38 In the tropical stratosphere, the Quasi-biennial Oscillation (QBO) can also alter where vertically-propagating
39 waves will break and deposit their momentum in the stratosphere. While the mechanism coupling the tropical strato-
40 spheric winds to the polar vortex is still uncertain (Garfinkel et al., 2012b; Anstey and Shepherd, 2014; White et al.,
41 2015), one possibility is that the QBO changes the location of the subtropical critical line, beyond which the wind is
42 easterly and where quasi-stationary planetary scale waves cannot propagate (Charney and Drazin, 1961). This mech-
43 anism, termed the Holton-Tan effect, leads to changes in wave momentum convergence, driving an associated mean
44 meridional circulation that weakens the polar vortex and increases the likelihood of mid-winter SSWs during the east-
45 erly QBO phase (Holton and Tan, 1980). Lu et al. (2020) also find evidence that the QBO is associated not only with
46 changes in the location of wave breaking but also with changes to wave structure, with greater wavenumber-1 distur-
47 bance of the polar vortex in the easterly QBO phase. In general, S2S prediction systems do not intrinsically simulate

48 the QBO, but are able to persist the initialized tropical stratospheric wind conditions for a couple of weeks before
49 degrading towards model climatology (typically a weak easterly state) (Butler et al., 2016; Garfinkel et al., 2018).

50 The two split polar vortex events in February 2018 and January 2019 were preceded by conditions with differ-
51 ent phases of the QBO and ENSO, as well as differently phased but active periods of MJO evolution. These differ-
52 ences in the large-scale climate patterns, and how well prediction systems captured their influences on stratosphere-
53 troposphere coupling, likely had implications for the predictability of these events. Here we describe the evolution of
54 the two events in Section 3.1 and their differences in predictability in Section 3.2. We then quantify how errors in the
55 QBO and MJO forecasts in particular contributed to differences in the predictability of these major SSWs in Section
56 3.3.

57 2 | DATA AND METHODS

58 We use daily-mean fields from the ERA-Interim reanalysis (Dee et al., 2011) to describe the evolution of the atmo-
59 sphere and sea surface temperatures in Section 3.1 and for forecast verification in Section 3.2-3.3. Daily anomalies
60 were computed by removing a daily climatology calculated as the mean value of each calendar day over the 1979-2018
61 period. The Northern Annular Mode (NAM) index was calculated at each pressure level by computing deseasonalized
62 daily-mean geopotential heights averaged over 65-90°N with global-mean geopotential height anomalies removed
63 (Gerber and Martineau, 2018). The resulting time series are normalized to have unit variance, and multiplied by -
64 1 such that a negative NAM index denotes an anomalously weak polar vortex (and vice versa). The QBO index is
65 described by the first and second time-extended empirical orthogonal functions (EOFs) of 30-day mean tropical (5°S-
66 5°N) zonal winds in the layer from 10 hPa to 70 hPa (using ERA-40 reanalysis back to 1960 to extend the historical
67 data). Following Fraedrich et al. (1993) and Wang et al. (1995), the time-extended matrix is built using a sliding window
68 of fifteen 30-day periods.

69 The MJO index (the Real-time Multivariate MJO, or RMM) is calculated as in Wheeler and Hendon (2004) and is
70 provided by the Australian Bureau of Meteorology [<http://www.bom.gov.au/climate/mjo/graphics/rmm.74toRealttime.txt>]. Bivariate root mean square error (RMSE) of the MJO forecasts for the S2S prediction systems in Section 3.3 is
71 calculated following Rashid et al. (2011). The amplitude of ENSO is determined from the National Centers for Environ-
72 mental Prediction (NCEP) Climate Prediction Center Oceanic Niño Index based on ERSSTv5 sea surface temperatures
73 [https://origin.cpc.ncep.noaa.gov/products/analysis_monitoring/ensostuff/ONI_v5.php].

74 In Section 3.2-3.3, we use real-time forecasts and hindcasts from the Subseasonal to Seasonal (S2S) project
75 database (Vitart et al., 2017). We consider eight different prediction systems (CMA, ECCC, ECMWF, JMA, KMA,
76 CNRM, NCEP, and UKMO). These prediction systems are all initialized at different frequencies and provide forecasts
77 out to 4-6 weeks. They also have different reforecast periods. The details of the prediction systems used in our anal-
78 ysis, including ensemble size, forecast length and frequency are provided in Supporting Information (Table S1 and S2).
79 Note that JMA and ECCC do not provide a T+0h forecast; we have noted in figure captions where forecasts from
80 these systems are shifted by 1 day relative to other systems. We bias-correct the model forecasts by subtracting the
81 initialization date and lead-time dependent biases with respect to the ERA-Interim climatology, where the biases are
82 determined using the full hindcast dataset available for each model. We note here limitations of this method: the bias
83 correction method assumes that biases can be linearly removed, which may not be accurate especially during extreme
84 vortex events; and determining the biases is dependent on the reforecast climatology which can be noisy for periods
85 of less than 20 years.

86 We also examine moment diagnostics calculated from NCEP CFSv2 data (Saha et al., 2014) in Section 3.2 to ex-
87

amine the subseasonal prediction of the stratospheric vortex geometry. We specifically focus on the centroid latitude and aspect ratio diagnostics similar to that calculated in Seviour et al. (2013). We use bias-corrected 10 hPa geopotential height fields to calculate the parameters of the best-fit ellipse to the geometrical shape of the polar vortex. These calculations differ from those in Seviour et al. (2013) only in that the moments calculations are not weighted by the geopotential heights within the vortex edge contour; this is done so that the distortion of the vortex is better represented as the relative sizes and strengths of the sister vortices evolve during the split events. The centroid latitude and aspect ratio provide compact information about the degree of displacement of the vortex, and the degree of stretch in the vortex, respectively. They thus provide insight into how and when CFSv2 predicted the specific geometrical evolution of the 2018 and 2019 vortex splits. Note we also do not impose any persistence criteria to classify the events, as done in Seviour et al. (2013), but instead show the diagnostics for each day. We also use a region segmentation algorithm to confirm when the contour representing the vortex edge separates into more than one distinct region (e.g., Lawrence and Manney, 2018).

3 | RESULTS

3.1 | The evolution of the 2018 and 2019 polar vortex split events

The 2017-2018 and 2018-2019 NH stratospheric polar vortices evolved in a similar way through mid-December (Figure 1a). In the 2017-2018 winter, the vortex weakened in December, re-strengthened into the new year, weakened briefly in the latter half of January, and then rapidly decelerated and reversed on 12 February 2018 in a major SSW. The vortex split into two sister vortices, with the larger lobe ending up over central Canada and the smaller lobe over Europe (Figure 1b). The dominance of the Canadian vortex following this split event is not common; usually the Siberian vortex is dominant (Matthewman et al., 2009). The anomalous stratospheric warming peaked over the Atlantic sector within days of the zonal-mean zonal wind reversal. The zonal wind reversal following the February 2018 SSW was large in amplitude and persistent, lasting through the month and setting daily records for minimum zonal-mean zonal wind speeds at 60°N and 10 hPa for the period 1979-2019 in ERA-Interim. The vortex subsequently recovered, and the so-called “final warming” (the last reversal of the wintertime vortex to its easterly summer state) occurred on the climatological mean date of 15 April.

In the 2018-2019 winter, the vortex began to weaken in mid-December, ending with a major SSW on January 2 2019. This zonal wind reversal also set daily records for minimum zonal-mean zonal wind speeds for the 1979-2019 period, though the amplitude of the reversal was not as large as for the February 2018 event (Lee and Butler, 2019). In this case, the vortex first displaced over the North Atlantic region, warming the polar stratosphere; after the vortex split, the sister vortices ended up over Europe and eastern Canada (Figure 1c). These lobes of the disrupted vortex persisted in these locations with nearly equivalent size for three weeks after the SSW (Figure S1), and stayed split even after the zonal-mean zonal winds returned to westerly. After the SSW, the vortex radiatively recovered, then continued to intensify, leading to daily records for maximum zonal-mean zonal wind intensity in March 2019. Following this vortex strengthening, the vortex decelerated, ending in a final warming slightly later than average on 23 April (Lee and Butler, 2019).

While for both SSWs the polar vortex split, with one lobe of the vortex over Eurasia and the other over North America, both the drivers and the impacts of these events were actually quite different. Figure 2 shows the sea surface temperature (SST) and 500 hPa geopotential height anomalies in the month prior to the 2018 and 2019 vortex splits, respectively. In the winter of 2017-2018, a La Niña state prevailed (the DJF Oceanic Niño index was -0.9°C), with anomalously cold SSTs over the eastern and central tropical Pacific and tropical Indian Ocean. Opposite-signed tropical

128 Pacific SST anomalies emerged the following winter of 2018-2019, with El Niño conditions (the DJF Oceanic Niño
129 index was $+0.8^{\circ}\text{C}$). The pattern of 500 hPa height anomalies in the 30 days prior to the SSWs are at least partially
130 related via teleconnections to the different tropical Pacific conditions that were present in each winter. La Niña is
131 canonically associated with an anomalous ridge over the Aleutians and a negative Pacific-North American (PNA)-like
132 pattern over North America, and El Niño with an anomalous Aleutian low and a positive PNA-like pattern (Deser et al.,
133 2017).

134 The canonical La Niña teleconnection pattern was apparent in the month prior to the February 2018 SSW (Fig-
135 ure 2a), with some differences from the composite pattern: the anomalous Aleutian ridge was shifted northwest, the
136 anomalous western North American trough was largely absent, and the ridging over the southeastern US was shifted
137 northward. Persistent Aleutian blocking, mostly during La Niña winters, is a tropospheric precursor to SSWs, particu-
138 larly split vortex events (Martius et al., 2009; Castanheira and Barriopedro, 2010; Barriopedro and Calvo, 2014; Cohen
139 and Jones, 2011; Bancalá et al., 2012). Indeed, in the month prior to the February 2018 SSW, at least three strong
140 poleward eddy heat flux events (40-80°N, 300 hPa) occurred (Figure 3; positive or poleward meridional heat flux is a
141 proxy for vertically propagating waves from the troposphere to stratosphere). These predominantly $k=2$ wave pulses,
142 associated with anomalous ridging over the Aleutians and also the Ural region (Martius et al., 2009), have been found
143 to “precondition” the vortex, similar to other historical SSWs (Lawrence and Manney, 2020). The $k=2$ eddy heat fluxes
144 set 2-3 new daily record highs at 300 hPa in the month prior to the SSW (Figure 3b). The total eddy heat fluxes at
145 300 hPa also set daily record highs around 13-14 January 2018 and in the 2-3 days prior to the SSW (Figure 3b). The
146 daily record-breaking values of eddy heat flux at 10 hPa (Figure 3a) in the week leading up to the SSW are likely a
147 manifestation of the amplified $k=2$ pattern from the surface to 10 hPa as the vortex splitting occurred (Birner and
148 Albers, 2017; White et al., 2019; de la Cámara et al., 2019).

149 The canonical El Niño teleconnection was apparent prior to the 2019 SSW as an anomalous Aleutian low, another
150 “precursor pattern” for SSWs (Figure 2b; Garfinkel et al. (2010); Bao et al. (2017)). While there were two noticeable
151 $k=2$ eddy heat flux events in mid-November and early December, the vortex was primarily perturbed by a substantial
152 but relatively slow amplification of $k=1$ eddy heat fluxes associated with the anomalous Aleutian low and Ural blocking
153 high. The increase in $k=1$ eddy heat flux activity began in early to mid-December (Figure 3c,d), as the polar vortex was
154 displaced off the pole. It is interesting to note that historically El Niño SSWs are associated with more $k=1$ amplification
155 compared to La Niña SSWs (Barriopedro and Calvo, 2014). The $k=1$ eddy heat fluxes at 300 hPa and 10 hPa set daily
156 record highs in the two weeks prior to the 2019 SSW (Figure 3c,d). Finally, note that a strong $k=3$ pulse, setting a new
157 daily record high at 300 hPa (Figure 3d), occurred just days before the 2019 SSW (Rao et al., 2019).

158 The differences in wave activity in the 10 days before each SSW is highlighted in Figure 4. In 2018 (Figure 4a,b),
159 the $k=2$ pattern was nearly barotropic (vertically-stacked from the surface to the mid-stratosphere), a feature common
160 to split vortex events (Esler and Scott, 2005; Matthewman et al., 2009). The amplitude of $k=2$ exceeded $k=1$ from
161 roughly 500-30 hPa and was well above climatology from 500-1 hPa. In comparison, in 2019 (Figure 4c,d), there was a
162 strongly baroclinic (westward tilting with height) predominantly $k=1$ pattern. This anomalous wave pattern projected
163 well onto the climatological wave pattern, suggesting constructive linear interference, which would allow amplification
164 into the stratosphere (Smith and Kushner, 2012). Note also that in 2019, the amplitude of $k=2$ was below climatology
165 throughout the atmospheric column, while $k=3$ was above climatology from the surface to 30 hPa and matched $k=1$
166 amplitudes in the troposphere (see also Figure 3d). While disruptions of the polar vortex have only rarely been linked
167 to $k=3$ activity (Shi et al., 2017), we think there are two potential explanations for its presence here. One possibility is
168 that the $k=3$ pattern manifested as anomalous Pacific, Atlantic, and Ural blocking highs (Figure 4c), in a similar manner
169 to the $k=2$ pattern in 2018 (Figure 4a) but with enhanced sinuosity. As the vortex displaced over the Atlantic sector,
170 the Atlantic blocking high in particular may have helped to pull the vortex apart from below (Woollings et al., 2010,

171 Figure S1). Another possibility is that nonlinear interactions between large-scale waves and the stratospheric basic
172 state, as the vortex first displaced off the pole and then split, introduced a substantial $k=3$ component (Smith, 1983).

173 We also note that the downward influence following the SSW differed in 2018 and 2019. Both SSWs had a
174 similar persistence of the zonal-mean zonal wind reversal in the stratosphere (19 days in 2018 and 21 days in 2019),
175 though the magnitude of the reversal, which was roughly -24 m/s in 2018 and only -10 m/s in 2019 (Lee and Butler,
176 2019), might have played a role in the subsequent downward coupling (Karpechko et al., 2017; Rao et al., 2020). In
177 particular, in 2018, the developing anomalous high pressure over the polar stratosphere had a downward influence
178 that reached the troposphere within a week (Figure 5a). The anomalously high geopotential height anomalies over
179 the Arctic were associated with an equatorward shift of the North Atlantic storm track and a negative NAM that
180 persisted from late February to mid-March. Significant winter storms, such as the "Beast from the East", impacted
181 Europe in late February (Greening and Hodgson, 2019); and as the North Atlantic blocking retrograded westward
182 over Greenland in March (Figure S1), the northeastern United States also saw much colder weather and a surge of
183 nor'easter extratropical cyclones (<https://www.ncdc.noaa.gov/sotc/national/201803>). The stratospheric polar vortex
184 did not reform until March.

185 In 2019, strongly positive polar cap geopotential anomalies at 10 hPa, or the negative phase of the stratospheric
186 NAM, were observed 10-12 days prior to the zonal wind reversal (Figure 5b). In the troposphere, anomalous and
187 retrograding North Atlantic blocking was apparent, as in 2018, but the blocking grew weaker in the weeks following
188 the SSW rather than stronger, and the impacts from the SSW did not project strongly onto the North Atlantic Os-
189 cillation (the regional manifestation of the NAM). Note that the NAM at 1000 hPa did change to its negative phase
190 from mid-January to early February, despite the fact that the free troposphere was in a positive NAM state for most
191 of that period (Figure 5b). It has been previously noted that the surface response to anomalous stratospheric forcing
192 is amplified compared to the troposphere (Baldwin and Thompson, 2009; Domeisen et al., 2020a). Nonetheless, the
193 stratosphere-troposphere coupling following the SSW was not particularly strong nor persistent, with few notable
194 impacts except perhaps the tropospheric "polar vortex" event that led to record cold in the central United States in
195 late January (which was vertically well-aligned with the lobe of the stratospheric polar vortex over central Canada),
196 followed by a week of anomalous cold over the eastern US in early February.

197 In essence, both the upward (troposphere to stratosphere) and downward (stratosphere to troposphere) coupling
198 were quite different for these two events. The 2018 SSW was driven by primarily $k=2$ eddy heat fluxes associated with
199 barotropic blocking patterns, and resulted in a persistently negative phase of the NAM and cold weather outbreaks.
200 The 2019 SSW was preceded by slow and steady amplification of $k=1$ that first displaced the vortex before it split,
201 and resulted in strong downward coupling to the tropopause that only briefly translated to typical surface impacts of
202 a SSW. It should be noted that in the historical record, split polar vortex events have been observed to follow both $k=1$
203 and $k=2$ tropospheric precursor patterns (Bancalá et al., 2012), with the precursor pattern strongly tied to the phase
204 of ENSO (Barriopedro and Calvo, 2014), in agreement with the 2018 and 2019 cases here. Additionally, irrespective
205 of ENSO phase, both SSWs were preceded by strong anomalous ridging over the Ural mountains in northwestern
206 Eurasia (also apparent in Figure 2 and Figure S1). Previous studies have noted the Ural blocking high as an important
207 driver of increased upward-propagating wave activity (Garfinkel et al., 2010; Karpechko et al., 2018; Peings, 2019).

208 Other large-scale circulation patterns such as the QBO and the MJO may have influenced these two events.
209 Figure 6a shows polar vortex weakening events at 60°N and 10 hPa plotted in the phase space of the QBO, as
210 described by the first and second time-extended EOFs of 30-day mean tropical (5°S - 5°N) zonal winds in the layer
211 from 10 hPa to 70 hPa. The light gray lines trace the historical QBO values; note that the QBO amplitude does not
212 vary much, with the obvious exception of the 2016 disruption (Coy et al., 2017). It is apparent that polar vortex
213 weakenings tend to occur when the tropical stratospheric winds below roughly 30 hPa are easterly (i.e., below the

214 diagonal EOF1=EOF2 line; 42 events), relative to when the stratospheric winds below 30 hPa are westerly (18 events).
215 Both the February 2018 and January 2019 SSWs occurred in the top two QBO octants for most vortex weakenings.
216 However the February 2018 event occurred during a period when the tropical winds at 50 hPa were weakly westerly
217 (with strong easterlies above 40 hPa), whereas in January 2019 the tropical winds at 50 hPa were easterly (with strong
218 westerlies above 40 hPa); i.e., the QBO was roughly 180° out of phase between the two events. The maximum
219 daily 100 hPa 45-75°N eddy heat flux anomaly in the 10 days prior to vortex weakening (shading of dots) generally
220 corresponds with the magnitude of the polar vortex weakening (size of dots). There is some preference for preceding
221 eddy heat flux anomalies to be enhanced for events occurring in phases with easterlies in the tropical stratosphere
222 below 30 hPa (Holton and Tan, 1980; Lu et al., 2020).

223 Like ENSO, the MJO can lead to poleward-propagating wave trains that shift longitudinally depending on where
224 the anomalous tropical convection occurs (Hoskins and Karoly, 1981). Figure 6b shows MJO-polar vortex states for
225 February 2018 and January 2019 in a historical context. Polar vortex weakenings at 60°N and 10 hPa (dot size) and
226 the minimum value of the 1000 hPa NAM index in the following 45 days (shading), are plotted in the phase space of
227 the RMM index. Note that only simultaneous relationships between the MJO and polar vortex strength are indicated
228 here; other studies have suggested there is a 10-20 day lag between the MJO tropical forcing and extratropical and
229 stratospheric impacts (Kang and Tziperman, 2017). Many SSWs have occurred when the MJO is inactive (i.e., within
230 the unit circle on the phase space diagram). However, for both the February 2018 and January 2019 SSW, the MJO
231 was active in phases 5-7 prior to the events, consistent with, e.g., Garfinkel et al. (2012a) and Liu et al. (2014).

232 Specifically, 15 days before the February 2018 SSW, enhanced tropical convection was evident over the western
233 Pacific as the MJO hit near record amplitudes in phase 6-7 (Figure 6b). This phase of the MJO is typically followed
234 10-15 days later by enhanced extratropical eddy heat fluxes (Garfinkel et al., 2014) with polar vortex weakening
235 (Garfinkel et al., 2012a). A week after the SSW, the MJO entered phase 8 which is associated with a negative NAO
236 pattern (Lin et al., 2009; Henderson et al., 2016). Thus it is possible that the MJO played a significant role via its extra-
237 tropical teleconnection in encouraging a North Atlantic blocking pattern by mid- to late February that was conducive
238 to downward coupling from the stratosphere (Barnes et al., 2019). By early March 2018, the MJO signal in phase 2
239 had decayed; interestingly, this is also around when the mid-troposphere showed less coupling with the stratosphere
240 (Figure 5), though the surface NAM stayed negative through mid-March. In comparison, prior to the January 2019
241 SSW, the MJO was entering phase 5-6. The MJO then moved quickly at low amplitude through phase 6-7, stalled in
242 phase 8 in early January, decayed, and then re-emerged around 18 January in phases 4-5. Overall, the MJO was in a
243 similar phase as 2018 but with weaker amplitude and faster decay to neutral state.

244 Figure 6b indicates that the MJO phase may modulate the surface response following the polar vortex deceleration
245 (Barnes et al., 2019). There is a preference for lower minimum NAM values in the 45 days following SSWs that
246 occur in active (RMM magnitude >1) MJO phases 6-1 compared to phases 2-5 (the frequency of events with minimum
247 NAM $< -2\sigma$ is 12/19 = 0.63 for phases 6-1, and 7/20 = 0.35 for phases 2-5). This suggests that the MJO enhanced
248 the likelihood of a more negative surface NAM response following both the February 2018 and January 2019 SSWs.

249 3.2 | The predictability of the 2018 and 2019 polar vortex splits

250 In general, SSWs are not predictable more than 10-20 days in advance, though this can vary by individual event
251 (Tripathi et al., 2014; Karpechko, 2018). Vortex splits are likely to be even less predictable than vortex displacements
252 (Taguchi, 2015, 2018). The predictability of the 2018 SSW has been detailed in previous studies. Karpechko et al.
253 (2018) found that S2S prediction systems forecast the 2018 SSW about 11 days in advance, and accurate prediction
254 was linked to the ability of the models to correctly capture the location of anomalous Ural blocking (e.g., Figure 2). Rao

255 et al. (2018) found similar timescales of predictive skill for the 2018 event. Lee et al. (2019) found that anticyclonic
256 Rossby wave breaking in the North Atlantic, which was tied to the amplification of the Ural blocking high and $k=2$
257 in the stratosphere, was not well forecast and thus limited how soon prediction systems forecast the event. For the
258 2019 event, Rao et al. (2019) found predictive skill at up to 18 day lead times in the S2S prediction systems.

259 We use real-time forecasts of the S2S prediction systems to compare the prediction of the 2018 and 2019 vortex
260 splits. We show the percent of ensemble members accurately detecting the first zonal-mean zonal wind reversal in the
261 forecast within ± 3 days of the observed SSW date (Figure 7a,c), and the percent of ensemble members that detect the
262 first zonal wind reversal in the forecast outside this 3-day window (Figure 7b,d) for different S2S prediction systems.
263 In 2018, most prediction systems had fewer than 75% of members accurately detecting the SSW until 10-13 days
264 prior to the SSW. Models with poor stratospheric resolution, such as CMA, had even shorter predictability of about
265 7 days. However most of the prediction systems also showed increased probability ($> 50\%$) of a "false" SSW (mostly
266 predicted to occur near the end of January, not shown) when initialized between January 7-20, during the period of
267 record-setting upper tropospheric wave activity (Figure 3). In other words the models may have tried to break down
268 the vortex too soon in response to this anomalous tropospheric eddy heat flux.

269 In 2019, the detection of the SSW was possible at longer lead times, with KMA, ECMWF, and UKMO prediction
270 systems showing accurate detection of the SSW for forecasts initialized up to 15-22 days prior. However, other
271 systems such as NCEP showed accurate detection at shorter lead times compared to the 2018 event (though the
272 false alarm rate peaked in NCEP just before the peak in hit rate, suggesting NCEP predicted an SSW for forecasts
273 initialized as early as Dec 11th, but not at the correct time). Across the prediction systems, there was a gradual
274 increase in false alarms for forecasts initialized in the weeks prior to the event, likely tied to the gradual increase of
275 $k=1$ amplitudes (Figure 3).

276 Note here we are showing detection rates using bias-corrected forecasts. The bias-correction affects primarily
277 extended-range forecasts; a version of Figure 7 without bias correction is available in Supplemental Information (Fig
278 S2) for comparison. We find at forecast times of greater than 10 days significant biases that can affect accurate
279 prediction of stratospheric variability for some prediction systems. While Figure 7 accounts for these biases, we
280 note that in particular CMA and NCEP CFSv2 have large (opposing) biases at forecast times of greater than 10 days
281 during parts of the winter season. For example, NCEP is biased at extended-range (10-40 days) towards a weaker-
282 than-observed vortex, particularly in early winter (Figure S3). On the other hand, CMA is biased at extended-range
283 towards a stronger-than-observed vortex for initializations throughout the winter.

284 Given that the February 2018 SSW was a predominantly $k=2$ split vortex event while the January 2019 SSW first
285 displaced by an amplifying $k=1$ before splitting, we investigate how well the NCEP CFSv2 predicted these differences
286 in morphology. NCEP CFSv2 forecasts of polar vortex geometry are derived from 10 hPa geopotential heights (Seviour
287 et al., 2013). The calculation of the aspect ratio, which measures the elongation of the vortex (where larger aspect
288 ratio means more elongation), and the centroid latitude, which designates the center latitude position of the vortex, is
289 described in Section 2. Figure 8a,c shows the predicted vortex evolution for the February 2018 SSW. Here, individual
290 ensemble-mean forecasts run diagonally, with increasing forecast day on the y-axis, so that forecast validation times
291 line up in each vertical column. During the 2018 winter season, at forecast times beyond roughly 20 days, CFSv2
292 generally predicted a tendency for the vortex to become displaced to latitudes near and below 70°N by early February
293 (Figure 8a), with little elongation (Figure 8c). Only for forecasts of less than 10-14 days was NCEP able to better
294 capture both the observed center latitude of the vortex and the elongation of the vortex. The prediction system
295 captured the observed vortex elongation that occurred between January 20-25 for forecast times of up to 10 days,
296 which was also when a false SSW was forecast (Figure 7b); and only accurately predicted the observed split vortex
297 (here defined as an aspect ratio greater than 2.4) for forecast times of less than 10 days.

298 In contrast, for the 2019 SSW (Figure 8b,d), NCEP CFSv2 consistently and accurately predicted a displaced vortex
299 to latitudes near and equatorward of 70°N by mid-December 2018 even up to 20 days in advance. However, the
300 forecasts for the end of December into January strongly indicated the event would be a displacement, even for 5-20
301 day forecasts, as NCEP struggled to capture the elongation and split of the vortex. The difficulty in capturing the
302 event evolution is likely related to failure of the models to capture the $k=3$ peak at longer forecast times (Rao et al.,
303 2019). The observed split of the vortex, which occurred December 28 2018-January 2 2019, was only accurately
304 reflected in the forecasts initialized at most a week ahead of time, even shorter lead times than in 2018.

305 Finally, we consider how well the large-scale surface response, as measured by the 1000 hPa NAM index for days
306 10-30 following each SSW, was predicted by the S2S prediction systems' ensemble-mean (Figure 9). The observed
307 NAM during this 20-day period was -1.23 standardized units following the 2018 SSW, but only -0.60 standardized
308 units following the 2019 SSW (see Figure 5). For each SSW, we show how the forecast NAM response was verified for
309 forecasts initialized two weeks prior to the event and then for the 10 days after the SSW occurred; gray dots indicate
310 that the forecast value was within ± 0.5 standard deviations of the ERA-Interim response. For most prediction systems,
311 as expected, the forecast of the surface response is better predicted for initializations closer to the 10-30 day window
312 following each SSW (i.e., at shorter lead times).

313 Prior to the 2018 SSW (Figure 9a), most prediction systems were suggesting a NAM response that was more
314 positive than observed (blue shading), particularly when initialized in late January, when the forecasts backed off
315 from a SSW occurring (Figure 7a). KMA and UKMO continued to underestimate the negative NAM response even in
316 forecasts initialized after the SSW. In 2019 (Figure 9b), again most forecasts were suggesting a more positive NAM
317 response than observed in the forecasts initialized prior to the event, though ECMWF showed little error in verification
318 even in forecasts initialized on December 20th (and also had greater than 70 % of members detecting the observed
319 SSW, Figure 7c). Other prediction systems, such as UKMO and NCEP, continued to underestimate the negative NAM
320 response until after the SSW occurred. CMA generally predicted too positive of a response prior to, and too negative
321 of a response after, both the 2018 and 2019 SSW.

322 Note that the predicted NAM response has been bias-corrected in Figure 9. Without bias correction (Figure
323 S4), CMA generally predicted too weak a NAM response while NCEP predicted a more negative NAM response than
324 observed across almost all lead times in both 2018 and 2019. This is consistent with the strong vortex bias in CMA and
325 the weak vortex bias in NCEP. The systematic bias correction thus has a substantial impact on verification following
326 these events.

327 **3.3 | Linking errors in the QBO and MJO forecasts to errors in polar vortex prediction**

328 Accurate extended-range (week 2-5) forecasts of polar vortex variability in the S2S prediction systems may be linked
329 to their ability to simulate processes and teleconnections correctly. In this section, we examine the relationship of
330 forecast errors in the QBO and MJO to errors in the polar vortex forecast leading up to the 2018 and 2019 SSWs.
331 Throughout this section we use bias-corrected forecasts, but the results are the same if we use uncorrected forecasts.

332 In 2018, the QBO at 50 hPa was weakly westerly. Averaging over forecast-times for 4 initializations (or launch
333 dates) prior to the SSW that are consistent across prediction systems (Figure 10a), we can see that although the
334 S2S prediction systems were initialized with similar tropical stratospheric wind values (with the exception of JMA and
335 ECCC), the forecast tropical winds degrade to an easterly state at 50 hPa within 10 days for almost all systems. Models
336 without an intrinsic QBO typically have weak climatological easterlies in the tropics (Butler et al., 2016). All systems
337 thus show too easterly winds in the tropics at almost all forecast times; JMA and ECCC show the greatest errors at
338 20-30 day forecast times.

339 A different situation was present in 2019, when the QBO at 50 hPa was easterly but more rapidly transitioning to
340 neutral as westerlies descended from above (Figure 10c). The prediction systems degraded the initialized easterlies
341 more rapidly than observed, so in this case, most of the systems had tropical winds that were too westerly compared
342 to ERA-Interim. The one exception was NCEP which had too easterly winds at all forecast times. UKMO had the
343 smallest errors at 10-30 day forecast times, while JMA had the largest errors at 20-30 day forecast times.

344 We examine the relationship between errors in the predicted tropical stratospheric winds for the 7-11 February
345 2018 validation dates and errors in the stratospheric polar vortex (SPV) for the 12-16 February 2018 validation dates
346 (Figure 10b; following Rao et al. (2019)), for different forecast initialization dates. We might expect via the Holton-Tan
347 relationship that a more easterly QBO error would be linked to a more easterly SPV error (i.e., a positive correlation).
348 However, at first glance it appears the opposite is true in 2018; forecasts with more easterly QBO error had too strong
349 of polar vortex winds. Upon further examination, we see that the amplitude of the relationship across the 4 launch
350 dates shown ($r=-0.31$) is largely a reflection of the decreased error, particularly for polar vortex strength, at shorter
351 lead-times. Only at the earliest launch date (18 January 2018, blue dots in Fig 10b) is the correlation of SPV error to
352 QBO error significantly positive. The systems with the greatest easterly QBO error at 20-30 day forecast times, JMA
353 and ECCO, also show the smallest (more easterly) SPV error for this launch date.

354 In 2019 (Figure 10d), because the 28 Dec-1 Jan QBO forecast error is mostly positive (i.e., the winds were too
355 westerly compared to observed), the correlation with the 2-6 Jan SPV error is positive across all launch dates ($r=0.38$).
356 Again, the amplitude of the correlation across launch dates is a reflection of the reduced errors in the SPV at shorter
357 lead-times. At the earliest launch date of December 6 (blue dots), the correlation is significantly positive as expected,
358 and the systems with the more easterly QBO errors, NCEP and UKMO, also had the smallest (more easterly) SPV
359 errors. Thus beyond 1-3 week forecast times, the biases in polar vortex winds and QBO winds may be linked as
360 expected from the Holton-Tan relationship (Garfinkel et al., 2018).

361 Note that for individual launch dates less than 1-3 weeks prior to the SSWs, the relationship between errors in
362 QBO and SPV is quite different for 2018 and 2019. Namely, the launch dates from 25 January-8 February 2018
363 all show significant negative correlations, whereas the launch dates from 13-27 December 2018 show insignificant
364 correlations and disagree in sign. This suggests that prediction systems that better predicted the QBO in the week
365 prior to the February 2018 SSW also better predicted the SPV, while in 2019 there was little gain for systems better
366 predicting the QBO. The negative correlations in 2018 are particularly puzzling because they suggest the systems
367 that predicted a more easterly QBO also predicted a more westerly SPV. We think it likely that these relationships are
368 a reflection of some systems better predicting both the QBO state and the SPV state in 2018, rather than a causal
369 linkage between the QBO and SPV errors at shorter lead-times.

370 We also consider links between the bivariate RMSE (Rashid et al., 2011) of the MJO forecasts and the error in
371 both the SPV and NAM forecasts (Figure 11). Since the correlation across all 4 launch dates (shown in upper right
372 corner of each panel) is largely a measure of the change in error in both SPV/NAM and MJO RMSE with shorter lead-
373 times, we instead focus on the correlations for each launch date separately. Considering first the relationship of MJO
374 RMSE to the SPV error (Figure 11a,c), it is evident that in general, for most launch dates, there is little relationship
375 between the RMSE of the predicted MJO for the week before the SSW and the error in the predicted SPV. This may be
376 a reflection of the general underestimation of the observed MJO amplitude and slower eastward propagation in the
377 extended range by the S2S systems (Vitart, 2017; Lim et al., 2018, see also Figure S5-S6 in Supporting Information),
378 which may affect the MJO teleconnections and influence on the SPV. Or, it may be that we need to consider longer-
379 lagged relationships between the MJO error and SPV error; here we have selected validation periods a week apart to
380 maximize the number of forecasts we can compare, but relationships between MJO and SPV have been found to be
381 largest at 10-20 day lags (Kang and Tziperman, 2017). For launch dates closer to the 2018 SSW, there is a stronger

382 positive relationship, showing models that had reduced MJO RMSE also had reduced error in the SPV forecasts, but
383 these may not be causally linked.

384 Finally we examine the relationship of the MJO RMSE to errors in the NAM forecasts (Figure 11b,d). Here, we only
385 consider the 4 prediction systems (CMA, ECMWF, NCEP, and UKMO) that have long enough forecasts to compare
386 the error in NAM forecasts validated for the 22 Feb-14 Mar 2018 and 12 Jan-1 Feb 2019 periods (the same NAM
387 forecasts shown in Figure 9) with the MJO RMSE validated the week prior to the observed SSW. Note that ECMWF
388 will dominate these correlations because it has many more ensemble members than the other three. The overall
389 correlation across launch dates is weak and insignificant, as the NAM error is not consistently smaller across prediction
390 systems for shorter lead-times. In 2018, for the 29 January and 5 February launch dates, the relationship between the
391 NAM error after the SSW and the RMSE of the MJO the week prior to the SSW is significantly negatively correlated.
392 In particular, we note that ECMWF and UKMO both predicted too positive of 22 Feb-14 Mar NAM response for the
393 29 January and 1 February launch dates (Figure 9), despite having low RMSE for their MJO forecast, suggesting a
394 well-predicted MJO did not mean a better NAM forecast in 2018. Only for the 8 Feb launch date is the lower MJO
395 RMSE linked to lower NAM error, which may suggest the MJO helped induce the negative NAM once the SSW had
396 occurred.

397 In 2019, no significant relationship between MJO RMSE and NAM error was found for any launch date except
398 27 December. Overall, Figure 11b,d suggests that the NAM error after the 2018 and 2019 SSW was not closely tied
399 to the ability of the prediction systems to capture the MJO the week prior to the SSW.

400 4 | CONCLUSIONS

401 In this study, we performed an in-depth comparison between two split vortex events that occurred on February 12
402 2018 and January 2 2019. There were substantial differences in the evolution of these events, their surface impacts,
403 and their predictability. Here, we have examined how those differences may have been tied to differences in large-
404 scale circulation patterns, such as ENSO, the MJO and the QBO, and their teleconnections.

405 The 2018 SSW was primarily a wavenumber-2 event preceded by anomalous Aleutian and Ural blocking highs,
406 a tropospheric pattern which commonly precedes SSWs occurring in La Niña winters such as this one. However the
407 $k=2$ forcing that ultimately split the vortex was predictable at timescales of less than 7-10 days. The 2018 SSW was
408 followed by swift and persistent coupling to the surface, with substantial impacts on late winter weather over Europe
409 and the eastern United States. The negative NAM response was generally under-predicted in amplitude by most
410 prediction systems until the SSW occurred.

411 The 2019 SSW was instead a wavenumber-1 event, which first displaced the polar vortex over the North Atlantic
412 before splitting. This SSW was preceded by an Aleutian low and a North Atlantic/Ural blocking high, which interest-
413 ingly manifested as a strong wavenumber-3 forcing days before the vortex split. The moderate El Niño in the 2019
414 winter likely contributed to the persistent precursor pattern that led to the amplification of $k=1$ into the stratosphere,
415 and ultimately to the SSW. Because this SSW was associated with a slow build-up of $k=1$, it was predictable at much
416 longer lead-times, though the prediction systems only predicted the vortex would split (rather than displace) at lead
417 times of less than a week. After this event, there was weaker coupling to the surface with only minor weather impacts.
418 The predicted NAM response was still mostly too positive for the majority of prediction systems except for forecasts
419 initialized after the SSW. To understand the differences in evolution and predictability on S2S timescales, we examined
420 in particular how the QBO and MJO may have influenced these events and their impacts. While both SSWs occurred
421 in phases of the QBO historically conducive to SSWs, the QBO was 180 degrees out of phase in 2018 and 2019. This

422 difference in QBO state played a role in the forecast biases in S2S prediction systems. Namely, in 2018 the predic-
423 tion systems were not able to maintain the weak westerly state at 50 hPa and were biased easterly at all forecast
424 times, while in 2019 they largely captured the easterlies at 50 hPa but weakened too fast, so were biased westerly.
425 Nonetheless, in both years for forecasts initialized more than 3 weeks from the SSW, a Holton-Tan relationship was
426 apparent between the error in the QBO and the error in the SPV, with more easterly errors in the QBO significantly
427 correlated to more easterly error in the SPV. A significant positive relationship for the February 2018 SSW was also
428 seen between the MJO RMSE and the error in the SPV for launch dates less than 2 weeks prior to the SSW, and for
429 the error in the NAM for the launch date closest to the SSW. On the other hand, for the 2019 SSW, little significant
430 relationships were found between the RMSE of MJO forecasts and the error in the NAM or the error in the SPV.

431 While this is only a case study of two polar vortex events, and so statistical relationships cannot be inferred or
432 confirmed, it demonstrates the large inter-event variability of SSW impacts and predictability (Gerber and Martineau,
433 2018). A substantial part of this variability may not be due to the morphology of the SSW itself as some studies
434 suggest (Mitchell et al., 2013), but to other large-scale climate patterns in which the SSW evolves. While much
435 understanding has been gained regarding why and how SSWs occur, further work remains on how we can better
436 predict the stratosphere-troposphere coupling signal given the intrinsic variability of the large-scale circulation during
437 any given SSW event.

438 Supporting Information

439 Table S1. Model versions used in the prediction analysis of the 2018 SSW.

440 Table S2. Model versions used in the prediction analysis of the 2019 SSW.

441 Figure S1. Evolution of the 10 hPa geopotential height and temperature anomalies, and 500 hPa geopotential height
442 and surface temperature anomalies, from 21 days prior to 21 days after the 2018 and 2019 SSWs.

443 Figure S2. The non-bias-corrected version of Figure 7.

444 Figure S3. Biases in the NCEP CFSv2 prediction system.

445 Figure S4. The non-bias-corrected version of Figure 9.

446 Figure S5. MJO forecasts for 4 different launch dates prior to the February 2018 SSW.

447 Figure S6. MJO forecasts for 4 different launch dates prior to the January 2019 SSW.

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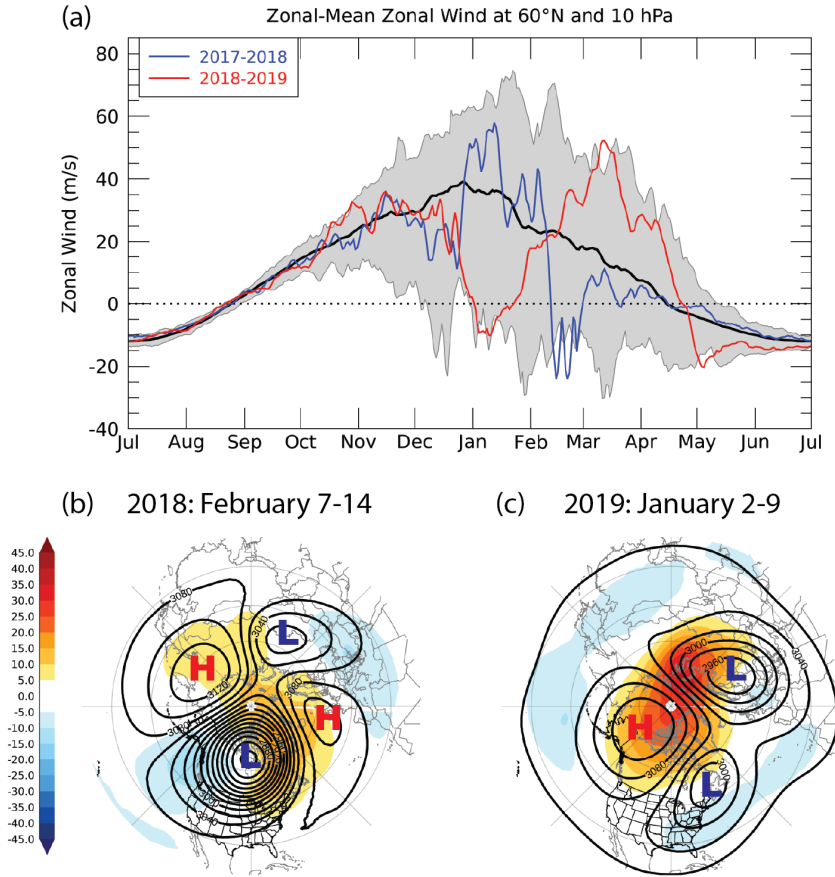


FIGURE 1 (a) Time series of the ERA-Interim 10 hPa 60°N zonal-mean zonal winds for 2017-2018 (blue line) and 2018-2019 (red line) overlying the daily mean (black line) and daily maximum and minimum (grey shading). The daily statistics are based on the 1979-2017 record. Averaged 10 hPa geopotential heights [dam] (contour) and 10 hPa temperature anomalies [K] (shading) for the (b) 7-14 February 2018, and the (c) 2-9 January 2019 period.

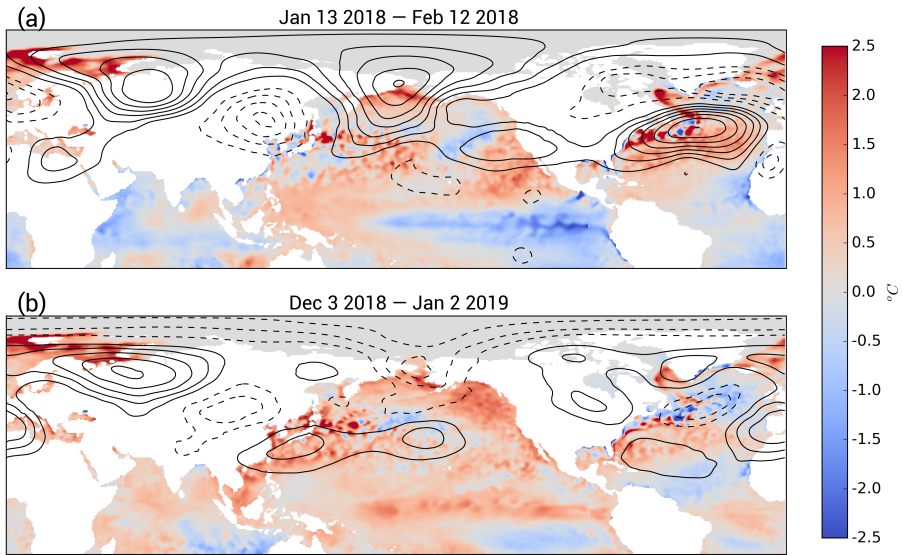


FIGURE 2 Sea surface temperature anomalies (shading, [degrees C]) and 500 hPa geopotential height anomalies (contoured every 30 m, omitting the 0 contour) from ERA-interim for 30 days preceding the vortex split on (a) 12 February 2018 and (b) 2 January 2019.

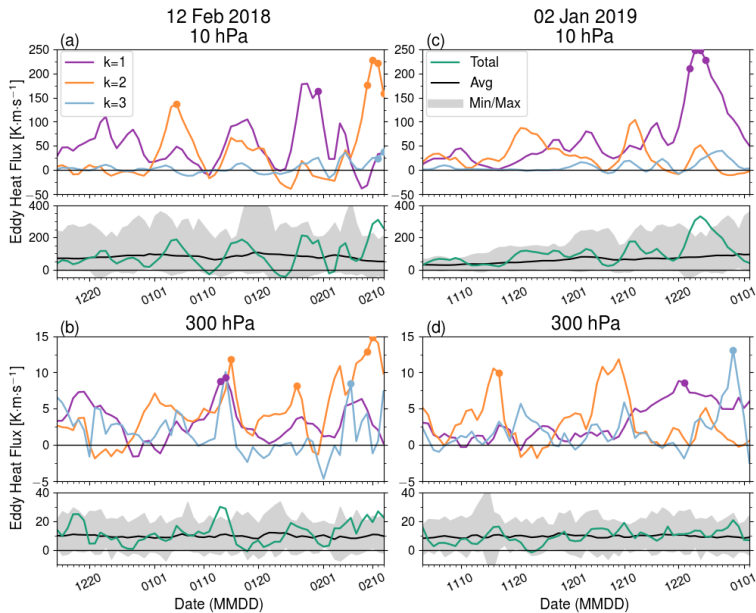


FIGURE 3 Time series of 40-80°N averaged eddy heat flux due to wavenumbers $k=1$, 2, and 3 on top of total eddy heat fluxes (green line) and their daily climatological extremes for the 1979-2017 period (gray shading) for 60 days before (a,b) the 2018 event and (c,d) the 2019 event, for (a,c) 10 hPa and (b,d) 300 hPa. Times when the wave-component heat fluxes are maxima in the ERA-Interim record are denoted with dots.

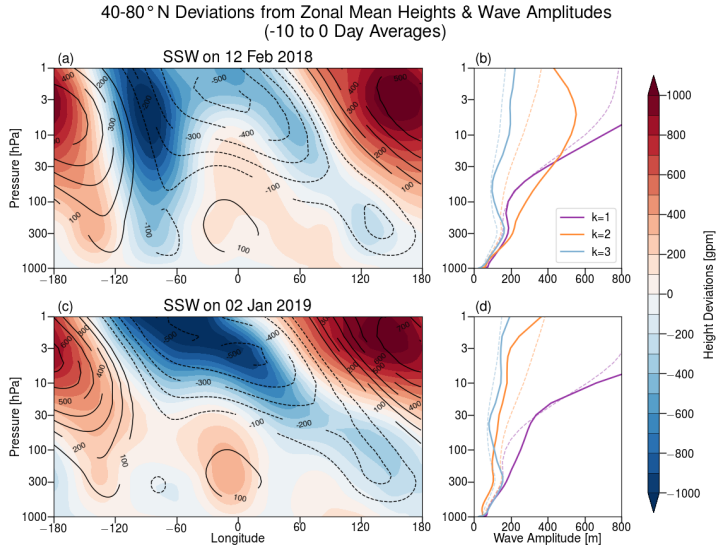


FIGURE 4 (a,c) Longitude by height cross-sections of 40-80°N deviations from zonal mean geopotential heights, and (b,d) the corresponding wave amplitudes for wavenumbers $k=1-3$, composited over days -10 to 0 prior to the (a,b) 2018 SSW and (c,d) 2019 SSW. In (a,c), the climatological height deviations from zonal mean for the same days are shown by black contours. In (b,d) the climatological wave amplitudes are shown by light dashed lines.

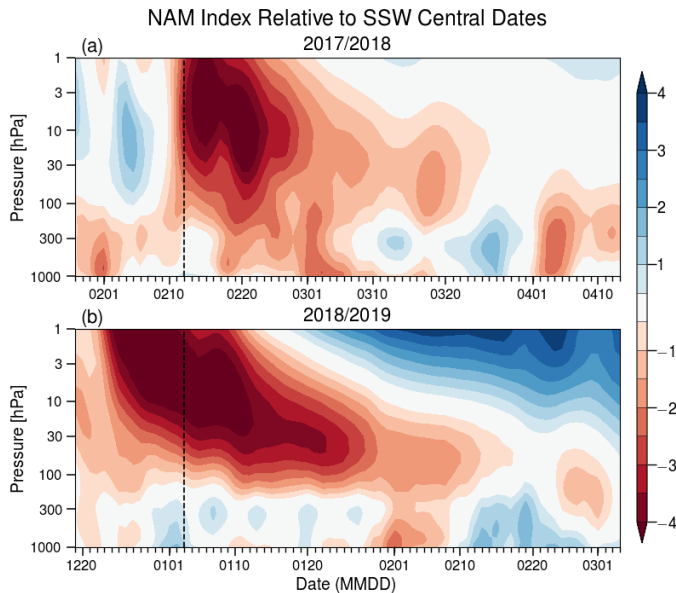


FIGURE 5 Time series of the Northern Annular Mode index as a function of pressure, for 15 days before and 60 days following the (a) 12 February 2018 and (b) 2 January 2019 sudden warming events.

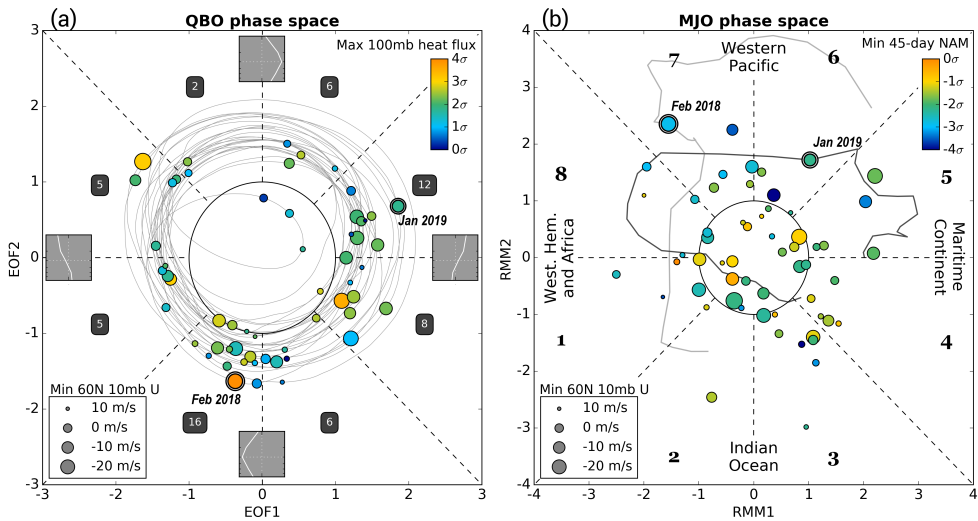


FIGURE 6 a) QBO phase space, with grey lines tracing the historical QBO and colored dots representing vortex weakening events (using ERA-40 reanalysis from 1960-1979, ERA-Interim from 1979-2019). The marker size corresponds to the magnitude of zonal-mean zonal wind reversal at 10 hPa and 60°N, and the color corresponds to the maximum 100 hPa 45-75°N eddy heat flux anomaly in the 10 days prior to each vortex event. For visualization, the grey box every 90 degrees shows the vertical profile of tropical (5°S-5°N) zonal-mean zonal winds, as a function of pressure, for that location in phase space. The horizontal white dashed line in the box represents 30 hPa. The number of vortex weakening events in each octant of the QBO phase space is provided in a dark grey box. b) State of the RMM index for the two events as compared with vortex weakening events since 1979 (based on ERA-Interim only), with the marker size corresponding to intensity and the color corresponding to the minimum 1000 hPa NAM value in the 45 days following the vortex events. Gray lines trace the MJO for 15 days prior to 15 days after the 12 February 2018 SSW (light gray) and 2 January 2019 SSW (dark gray).

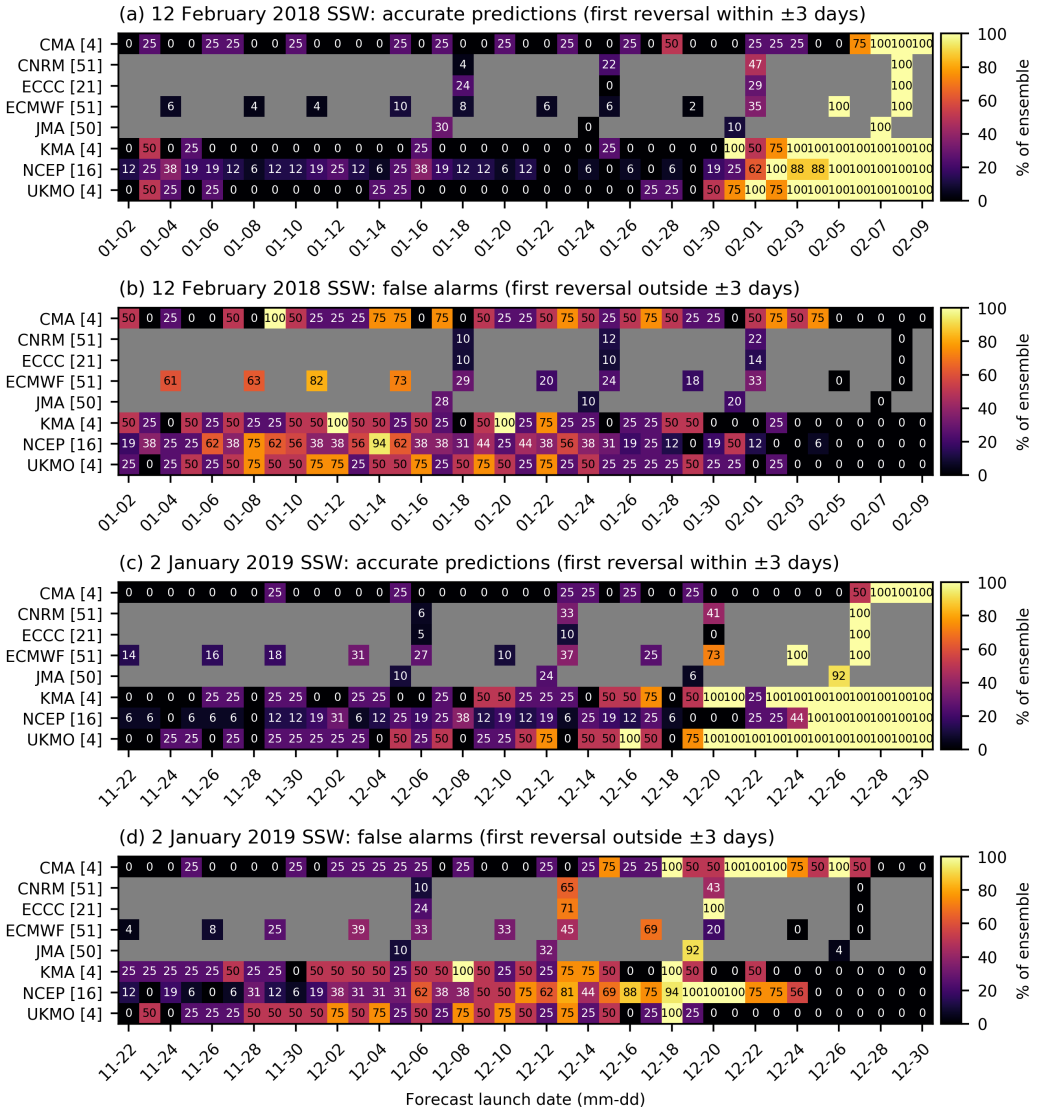


FIGURE 7 Percent of bias-corrected ensemble members from prediction systems participating in the S2S database that (a,c) correctly detected the first zonal wind reversal at 10 hPa and 60°N within ± 3 days of the observed event, and (b,d) detected the first zonal wind reversal at 10 hPa and 60°N outside the ± 3 day window, for the (a,b) 2018 SSW and the (c,d) 2019 SSW.

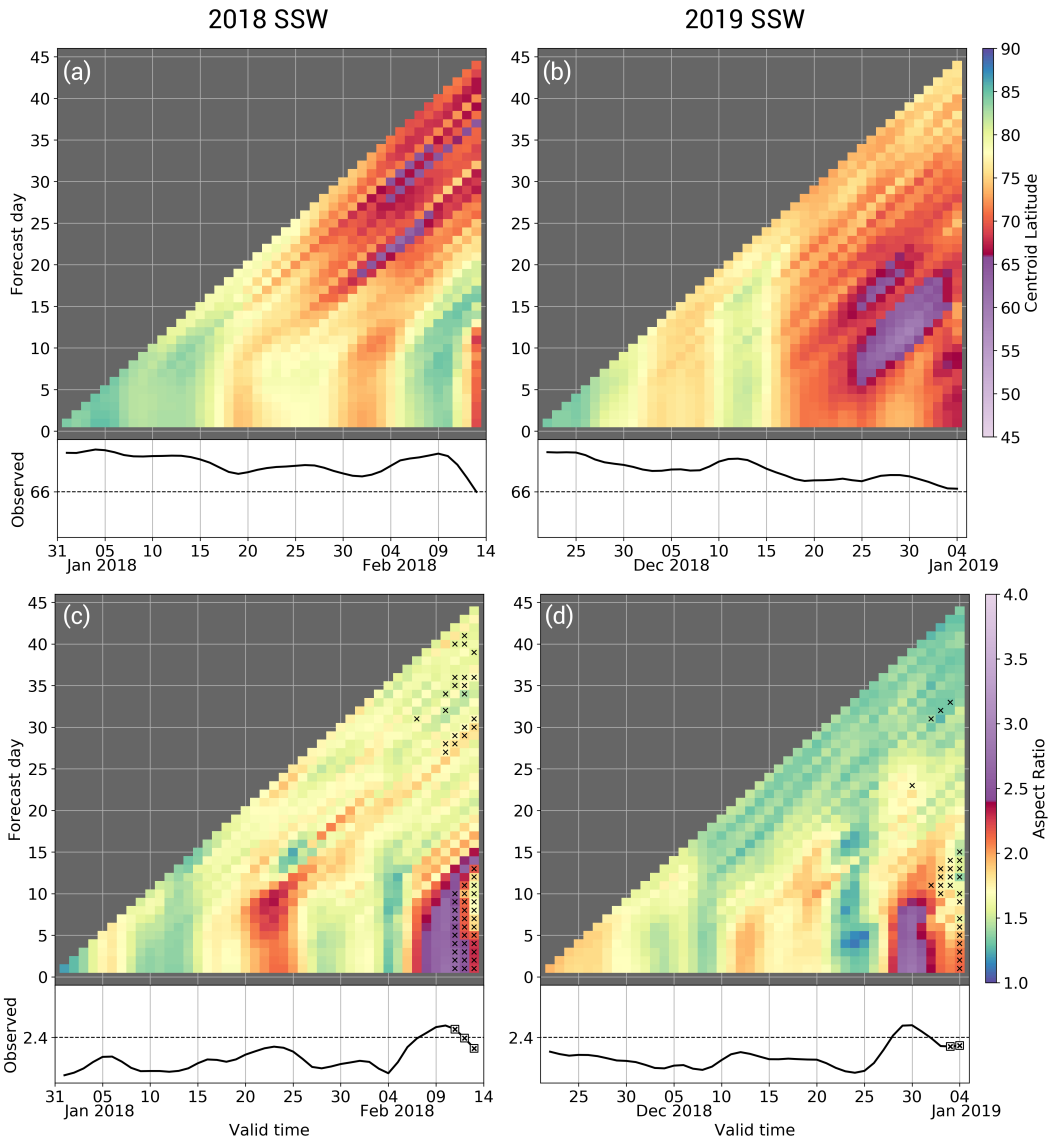


FIGURE 8 NCEP CFSv2 forecasts, as a function of initialization date (x-axis) and forecast day (y-axis), of (a,b) the vortex centroid latitude and (c,d) the aspect ratio, for the (a,c) 2018 SSW and the (b,d) 2019 SSW. The ERA-Interim moment diagnostics for each day are shown beneath each chiclet plot. The vortex can be said to be “displaced” when the centroid latitude is less than 66 degrees latitude, and to be “split” when the aspect ratio is greater than 2.4. The x’s indicate where a region segmentation algorithm detected more than 1 distinct vortex region in ERA-Interim, or where the algorithm detected splits in more than half of the CFSv2 ensemble members.

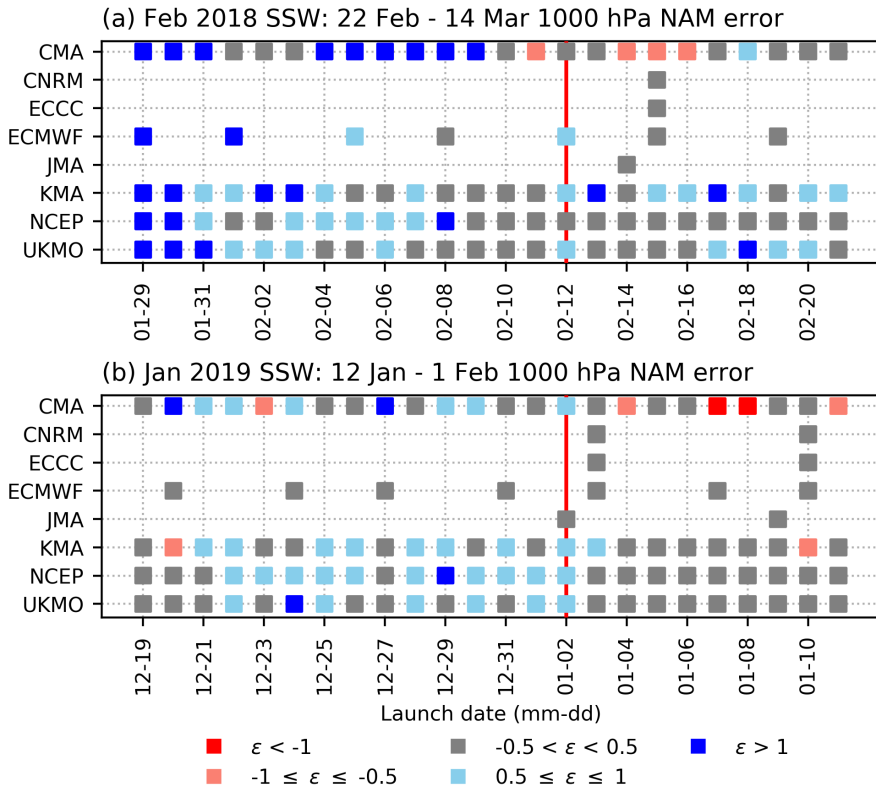


FIGURE 9 Difference between S2S ensemble mean forecasts for the average 1000 hPa NAM for 10-30 days following the 2018 and 2019 SSWs and verification from ERA-Interim. In (a) this is for the period 22 February–14 March, in (b) this is 12 January–1 February. Gray indicates the magnitude of the difference to the observed value was <0.5 , light red (light blue) indicates the forecast NAM was 0.5–1.0 standard deviations more negative (positive) than verification, and red (blue) indicates this difference exceeded 1 standard deviation. The vertical red line in each panel indicates the date of the SSW.

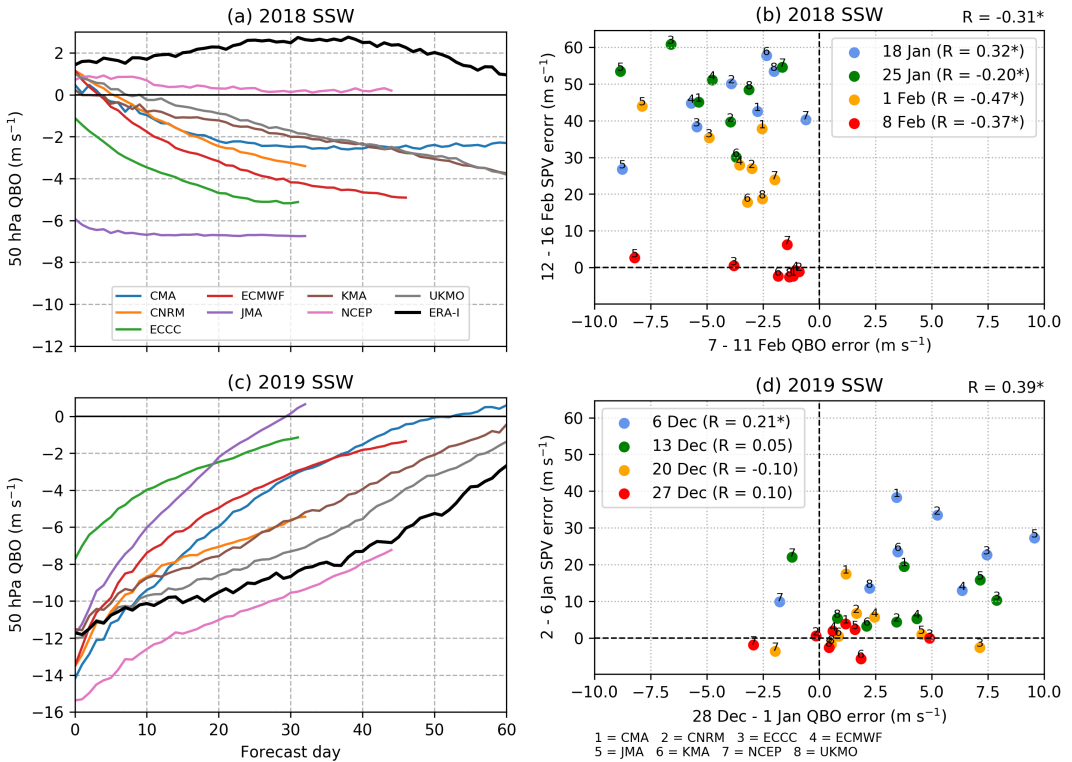


FIGURE 10 Tropical zonal-mean zonal winds averaged from 5°N-5°S at 50 hPa prior to (a) 2018 SSW and (b) 2019 SSW. These values are created by averaging forecasts over 4 launch dates consistent across prediction systems: 18 Jan, 25 Jan, 1 Feb, and 8 Feb 2018; and 6 Dec, 13 Dec, 20 Dec, and 27 Dec 2018 (note: JMA launches a day earlier). The black-line is ERA-Interim sub-sampled and averaged in an identical manner. Scatterplots of the averaged error in the QBO forecasts versus error in the zonal-mean zonal winds at 60°N and 10 hPa (SPV), for (b) 2018: QBO 7-11 Feb (8-11 Feb for 8 Feb launch date, except ECCC 9-11 Feb), SPV 12-16 Feb, and (d) 2019: QBO 28 Dec-1 Jan, SPV 2-6 Jan [i.e. 5 days before/after SSW]. Correlation values are calculated using the multi-model ensemble of 201 members. Only ensemble-mean values are shown. In all panels, these are hindcast bias-corrected. An asterisk indicates the correlation is significant at the 95% confidence level according to a bootstrap re-sampling test (10,000 repeats).

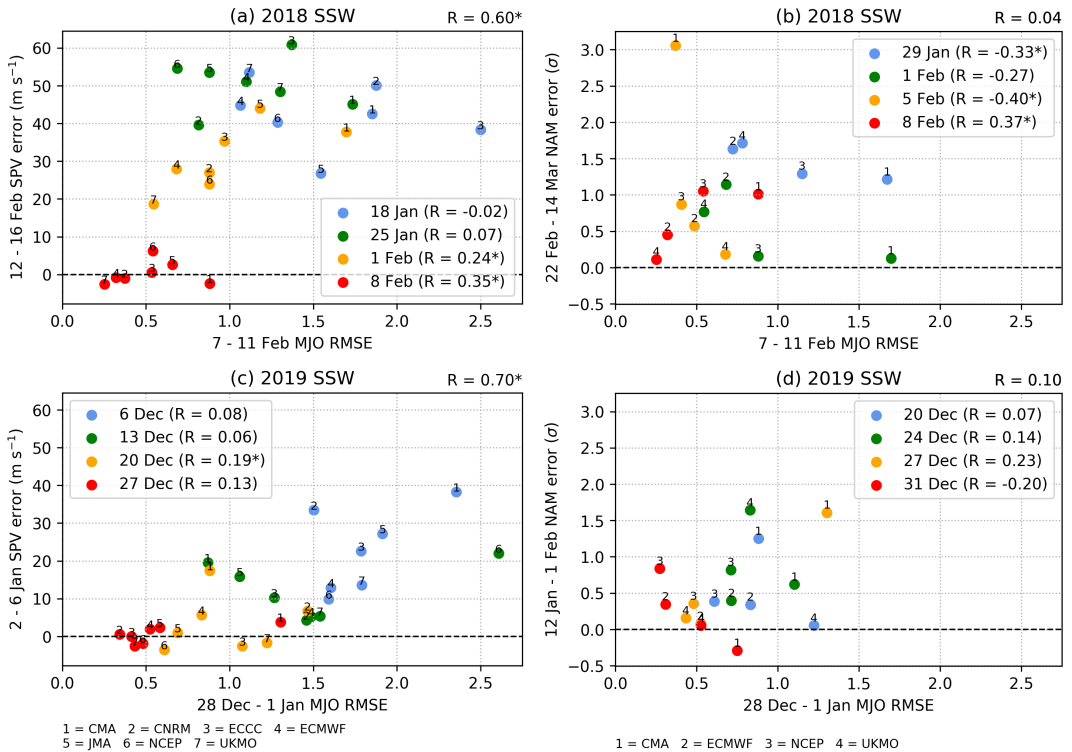


FIGURE 11 Scatterplots of the averaged bivariate RMSE in the MJO forecasts for the 5 days before each SSW [(a, b) 7 - 11 Feb (8 - 11 Feb for 8 Feb launch date except ECCS 9 - 11 Feb), (c, d) 28 Dec - 1 Jan] versus (a, c) error in the zonal-mean zonal winds at 60°N and 10 hPa (SPV) 5 days after the SSW and (b, d) error in the 1000 hPa NAM 10-30 days after the SSW (as in Figure 9). Correlation values are calculated using the multi-model ensemble of (a,c) 201 and (b,d) 75 members. Only ensemble-mean values are shown. In all panels, these are hindcast bias corrected. An asterisk indicates the correlation is significant at the 95% confidence level according to a bootstrap re-sampling test (10,000 repeats).