1 Can we project changes in fish abundance and distribution in response to climate?

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22 Keywords

- 23 Climate change, Marine fisheries, Biological feedback, Modelling, Species interactions,
- 24 Size spectrum, Model validation, Error estimation
- 25
- 26
- 27 Paper type
- 28 Technical advances
- 29

30 Abstract

Large scale and long-term changes in fish abundance and distribution in response to 31 climate change have been simulated using both statistical and process-based models. 32 However, national and regional fisheries management requires also shorter term 33 projections on smaller spatial scales, and these need to be validated against fisheries data. 34 35 A 26-year time series of fish surveys with high spatial resolution in the North East Atlantic 36 provides a unique opportunity to assess the ability of models to correctly simulate the changes in fish distribution and abundance that occurred in response to climate 37 variability and change. We use a dynamic bioclimate envelope model forced by physical-38 biogeochemical output from eight ocean models to simulate changes in fish abundance 39 40 and distribution at scales down to a spatial resolution of 0.5°. When comparing with these simulations with annual fish survey data, we found the largest differences at the 0.5° 41 scale. Differences between fishery model runs driven by different biogeochemical models 42 decrease dramatically when results are aggregated to larger scales (e.g. the whole North 43 Sea), to total catches rather than individual species or when the ensemble mean instead 44 of individual simulations are used. Recent improvements in the fidelity of biogeochemical 45 46 models translate into lower error rates in the fisheries simulations. However, predictions based on different biogeochemical models are often more similar to each other than they 47 are to the survey data, except for some pelagic species. We conclude that model results 48 49 can be used to guide fisheries management at larger spatial scales, but more caution is needed at smaller scales. 50

52 Introduction

Anthropogenic greenhouse gas emissions and associated warming strongly changes 53 ocean conditions, including temperature, salinity, ice cover, currents, oxygen, nutrients 54 and seawater acidity. These physical-biogeochemical changes affect the distribution, 55 abundance and productivity of phytoplankton, zooplankton and the fisheries that depend 56 on them (Perry et al., 2005; Pörtner, 2010; Cheung et al., 2011; Simpson et al., 2011; 57 Barange et al., 2014; Jennings and Collingridge, 2015; Fernandes et al., 2017; Maar et al., 58 59 2018; Lotze et al., 2019). Such changes are expected to continue during the 21st century under further global warming (IPCC, 2019) and have large implications for communities 60 and industries that depend on marine species for food and income (Roessig et al., 2004; 61 62 Cheung et al. 2012; Lam et al. 2012; Merino et al. 2012). A range of modelling approaches has been developed to project future changes in marine ecosystems and fisheries (e.g. 63 Stock et al. 2011; Cheung et al. 2016). These models range from considering only the 64 ocean physical dynamics and low trophic levels (Dunne et al., 2010; Butenschön et al. 65 2016; Yool et al. 2013) to high trophic levels of fisheries and conservation interest 66 (Nielsen et al., 2018; Peck et al., 2018; Tittensor et al., 2018). 67

Species process-based bioclimate envelope models are commonly applied to study 68 biological responses to global warming (Cheung et al., 2011; Jones et al. 2012). For 69 example, the Dynamic Bioclimate Envelope Model (DBEM) is a combined mechanistic-70 statistical approach that has been applied to a large number of marine species globally. 71 The DBEM projects changes in species distribution and abundance with explicit 72 consideration of known mechanisms of population dynamics, dispersal (larval and adult) 73 74 and ecophysiology, under changes in ocean temperature, salinity, oxygen, pH, upwelling, sea-ice extent and habitat type (Cheung et al. 2008; 2009, 2011). Simulations with the 75 76 DBEM model show that high-latitude regions will experience high rates of species invasion while the tropics will have high rates of local extinction by the end of the 21st 77 78 century under a high greenhouse gas emission scenario (Cheung et al., 2009; 2016). In addition, maximum catch potential is projected to decrease in the tropics while some high 79 latitude regions may experience increases in potential catch because of changes in range 80 and size of exploited marine species as well as changes in primary productivity under 81 global warming (Cheung et al. 2010). Recently, the DBEM has been combined with a size-82 83 spectrum model (Jennings et al., 2008) to evaluate the effects of inter-specific interactions

in projecting species distribution (Fernandes et al., 2013a). Size spectrum theory 84 accounts for energy transfer from primary production to individuals of different body 85 sizes to estimate abundance/biomass and their flows in marine ecosystems (Jennings et 86 al. 2008). The resulting integrated SS-DBEM (size-spectrum and DBEM model) projected 87 slower fish species shifts than in models that did not account for energy limitation. The 88 SS-DBEM has also been applied to several conservation issues (Jones et al., 2013; Queiros 89 et al., 2015) as well as socio-economic assessments in the North East Atlantic (Mullon et 90 al., 2016; Fernandes et al., 2017; Queiros et al., 2016) and developing countries 91 (Fernandes et al., 2016). However, projections of future species and fishery distributions 92 at local scales are uncertain (Payne et al., 2016, Cheung et al. 2016, Frölicher et al. 2016). 93

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The model simulations are often not well constrained by observational data as the 95 available observations are limited in time and space. However, the availability of an 96 extensive compilation of data describing the distribution and abundance of North Sea 97 fishes from 1982 to 2007 (Simpson et al., 2011) provide us the opportunity to evaluate 98 99 the accuracy of projections of simulated fish abundance and distributions from the SS-DBEM model. The data were collected by ICES co-ordinated bottom trawl surveys (ICES, 100 2012). We compare the data with simulations of fish abundance and distribution 101 conducted by linking multiple biogeochemical models and the SS-DBEM of commercial 102 fish populations. Thereby, we assess the likely reliability of future projections of fish 103 104 stocks under climate change and the impact of spread in ocean biogeochemistry simulations on the fisheries projections. In addition, the combination of observation 105 based atmospheric boundary conditions and atmospheric boundary conditions from 106 Earth system models allows for an indication as to how the projections of fish abundance 107 108 and distribution are affected by the internal variability of these systems in contrast with the more realistic variability from the reanalysis datasets. This understanding can be 109 110 applied when considering other ecological modelling approaches that occupy a similar niche to SS-DBEM, many of which are the focus of inter-comparison efforts within the 111 international FISH-MIP initiative (Nielsen et al., 2018; Peck et al., 2018; Tittensor et al., 112 2018; Heike et al., 2019). 113

114

- 115 Methods
- 116

We use output from three different ocean biogeochemical models (Table 1) to generate 117 the environmental and biological conditions (temperature, salinity, oxygen, pH, currents 118 and primary production) that drive the fish community model (size-spectrum dynamic 119 bioclimate envelope model; SS-DBEM; Fernandes et al., 2013a). The ocean 120 biogeochemical models are either run in fully coupled mode (i.e. coupled to a freely 121 evolving atmosphere, land, and sea-ice model) or run in hindcast mode (i.e. ocean-only 122 and forced at the ocean's surface with observed or model-derived atmospheric 123 conditions). We use multiple versions with different horizontal resolution of the three 124 ocean biogeochemical models to examine how an increase in spatial resolution and/or 125 different ocean biogeochemical model output influences the SS-DBEM simulations (Table 126 1). Furthermore, runs for the same model are used to compare different generations of 127 the same model and spatial resolutions. Fish distribution and abundance as simulated 128 between 1982 to 2007 are compared with 26 years of data from fish surveys by European 129 130 marine laboratories with a comparable spatial resolution (Simpson et al., 2011). 131

133 <u>Ocean biogeochemical models</u>

The choice of the biogeochemical model used to force the DBEM may have a significant 134 effect on forecasts of fish species abundance and distribution, especially in shelf seas 135 regions where predictions of the biogeochemical models can differ markedly. For this 136 reason, we chose to force the SS-DBEM with a diverse suite of output from three different 137 ocean biogeochemical models run under a range of different forcing modes in order to 138 cover the impact of the major sources of uncertainty in the biogeochemical forcing on the 139 fisheries mode. These forcing modes differ in the underlying model configuration used 140 and, as a result, in their relationship with real-world patterns of temporal variability. The 141 forcing modes used here are: 142

143 1. "Ocean-only hindcast mode", where an ocean-only model is driven by 144 observationally-derived surface forcing at its air-sea interface. This mode 145 necessarily includes real-world trends and patterns of variability. There are a 146 number of methods available for creating the so-called reanalysis forcing used in 147 this mode, and this study exploits model simulations using several different 148 approaches.

2. "Fully-coupled mode", where the model configuration includes an atmospheric 149 component that interacts dynamically with the ocean component. This mode 150 includes temporally-varying factors such as radiatively-active gases, so should 151 reproduce overall climate trends (e.g. global warming). But each fully coupled 152 model will show variability between runs whereas the observed data represent a 153 154 single realisation for each time point. For example, the real ocean experienced an El Niño in 1997-1998. The coupled simulations may have had a La Niña, El Niño 155 or been neutral at this time. Only the averages and variances of this internal 156 variability are consistent with the real world. 157

3. "Coupled-forced ocean-only mode", where an ocean-only model is driven by
surface forcing derived from a fully-coupled model. This mode is similar to the
first mode, but uses surface forcing output from a model running in the second
mode rather than observationally-derived forcing. As such, it may reproduce
observed long-term trends, but not specific temporal variability. This approach is
typically used where projection simulations into the future with relatively high
ocean model resolution or model downscaling are required (Yool et al., 2015).

Note that both ocean-only and coupled modes have limitations in the context of decadal-165 scale simulations: the forced ocean-only mode excludes any feedbacks between ocean 166 and atmosphere (and the associated uncertainties), while the fully-coupled mode and 167 coupled-forced ocean-only mode generate less directly-comparable data with present 168 day conditions (particularly when phasing of internal variability is relevant). The use of 169 a coupled prediction system as used in the WCRP Decadal Climate Prediction Project 170 could reduce these shortcoming, however, these systems currently focus on physical 171 climate and exclude biogeochemical components. 172

173

174 Table1 summarises the models from which output has been used to force the SS-DBEM

175 model.

Name of	Horizontal	Forcing mode (Forcing dataset)
	HUHZUIItai	Forcing mode (Forcing dataset)
model run	ocean	
	resolution	
GFDL-hindcast	1°	Hindcast (CORE2 reanalysis)
GFDL-coupled	1°	Fully coupled
GFDL-coupled-esm2m	1°	Fully coupled
MEDUSA-coupled-forced	1°	Coupled-forced (HadGEM2-ES)
ERSEM-hindcast-lowres	0.25°	Hindcast (DFS 4.1)
ERSEM-hindcast-lowres2	0.25°	Hindcast (DFS 5)
ERSEM-hindcast-highres	0.125°	Hindcast (ERA 40 & ECMWF)

176 Table 1. Characteristics of the different ocean biogeochemical simulations used in this study.

177

178 The models are:

1) The National Oceanographic and Atmospheric Administration (NOAA) and 179 Geophysical Fluid Dynamic Laboratory (GFDL) Earth System Model. GFDL is a 180 global model where: 1) GFDL CM2.1 is using MOM4 for its physics (Delworth et al. 181 2006) and TOPAZv0 for its biogeochemistry (Henson et al., 2010); and 2) GDFL 182 ESM2M is using MOM4p1 for its physics and TOPAZv2 for its biogeochemistry 183 (Dunne et al., 2010; 2012; 2013). TOPAZ simulates the cycling of carbon, nitrogen, 184 phosphorus, silicon, iron, oxygen, alkalinity and lithogenic material, and includes 185 three phytoplankton functional groups and one zooplankton. In this work the 186 GFDL CM2.1 is run in both hindcast and fully-coupled modes, whereas GFDL 187 ESM2M is run in fully coupled mode. The hindcast run here uses boundary 188 conditions of bulk air properties, incoming fluxes of radiation and freshwater, and 189

190 191 surface wind stress as prescribed by the observationally-derived CORE-2 reanalysis product (Large & Yeager, 2009).

2) The European Regional Seas Ecosystem Model (ERSEM, Butenschön et al. 2016) 192 coupled to the NEMO ocean model (low resolution; Madec, 2008) and the 193 POLCOMS ocean model (high resolution; Holt et al., 2001). Hindcast mode 194 simulations are used in this study both at different resolutions and under different 195 observationally-derived atmospheric boundary forcing including DFS 4.1 196 (Brodeau et al., 2010), ERA 40 & ECMWF (Uppala et al., 005) and DFS 5 (Brodeau 197 et al., 2010). ERSEM is a biogeochemical model for the lower trophic levels of the 198 pelagic and benthic ecosystem, and uses a functional-groups approach that 199 incorporates four phytoplankton, three zooplankton and bacterioplankton to 200 simulate decoupled carbon and nutrient dynamics (Blackford et al., 2004; 201 Butenschön, 2016). All ERSEM models here are regional models where the highres 202 simulation is a regional set-up for the North-West European Shelf with details on 203 the configuration, initial conditions, boundary conditions and forcings are 204 available in (Holt et al. 2012). The two ERSEM lowres simulations are from a 205 NEMO-ERSEM configuration for the Atlantic Ocean at 0.25 degree (further details 206 207 in Memery and Allen, 2011; Allen et al., 2014). The two different versions of this system reflect an update in model parametrisation and atmospheric forcing and 208 forecasting performance differences between using higher and lower results. Both 209 versions were included in the analysis in order to investigate how the resulting 210 changes propagate to the higher trophic level model. 211

3) The MEDUSA biogeochemical model (Yool et al., 2013a, 2013b) coupled to the
NEMO ocean model. Output from a coupled-forced ocean-only mode simulation is
used here, with the forcing derived from a CMIP5 simulation of the HadGEM2-ES
ESM (Collins et al., 2011). MEDUSA is lower complexity model, with two
phytoplankton, two zooplankton, three nutrients (N, Fe, Si) and slow-/fast-sinking
detritus compartments (Yool et al. 2013a).

In terms of biogeochemical complexity, MEDUSA is the simplest model considered here,

219 ERSEM is the more complex, with GFDL intermediate. Note that different versions of the

220 GDFL and ERSEM models were used here as well as different simulation configurations.

221 These differences relate to model evolution and improvement, including parameter

222 updates and the use of different spatial resolutions.

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224 <u>Fish model</u>

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The size-spectrum dynamic bioclimate envelope model (SS-DBEM) described in 226 Fernandes et al. (2013a) is used to simulate changes in abundance and distribution of fish 227 species. The SS-DBEM projects changes in species distribution and abundance with 228 explicit consideration of known mechanisms (Table 2) of population dynamics, dispersal 229 (larval and adult) and ecophysiology, under changes in ocean temperature, salinity, 230 upwelling, sea-ice extent and habitats (Cheung et al. 2011), and species interactions 231 232 based on size-spectrum theory and habitat suitability (Fernandes et al., 2013a). In SS-233 DBEM, current distributions of the studied species are first estimated based on habitat suitability (Close et al. 2006). This is done based on a global dataset of observed 234 abundance data from Cheung et al. (2008; available at fishbase.org which redirects to 235 maps hosted at aquamaps.com) overlaid with environmental data (temperature, salinity, 236 oxygen and pH at sea surface for pelagic species and at sea bottom for demersal species 237 238 as well as depth and distance to ice) from biogeochemical models described above. It is assumed that the carrying capacity of each species in each area is partly dependent on 239 the inferred preference profiles which depend on the projected biogeochemical 240 conditions (e.g. temperature, salinity, pH and currents) but limited by primary 241 production. Simultaneously, the model considers each species' physiological preferences 242 and tolerances to temperature, and sensitivity of key parameters determining the 243 species' mechanisms (mortality, growth and length-weight relationship). Natural 244 mortality rate is estimated from an empirical equation (Pauly, 1980) which considers 245 weight, growth and temperature. The model growth algorithm (Cheung et al., 2011) is 246 derived from the von Bertalanffy growth function (VBGF; von Bertalanffy, 1951). Therein, 247 growth is viewed as the difference between anabolic and catabolic processes. The 248 249 temporal and spatial patterns of pelagic larval dispersal (Cheung et al., 2008) are modelled by a two-dimensional advection-diffusion equation (Sibert et al. 1999; Gaylord 250 & Gaines 2000; Hundsdorfer & Verwer, 2003). Adult dispersal is calculated from the 251 dispersal or movement rate using an algorithm employed in an Eulerian spatial 252 ecosystem simulation model (Walters et al. 1999). 253

Table 2. Table summarizing main equations and parameters to consider the species

Mechanism	Equation	Parameters
Growth = anabolism -	$G = HW^a - kW$	H = anabolism coefficient
catabolism	$H = g[O_2] * e^{-j1/T}$	k = catabolism coefficient
(Pauly 2010; Cheung et al.,	$k = h[H^+] * e^{-j2/T}$	W = body weight
2011)		a = anabolism exponent (0.5 to 0.95)
		W∞ = asymptotic weight
		The coefficients g and h were derived from the average W1, K, and environmental temperature (T) of the species reported in the literature.
Length-Weight	W = a * L ^b	W = weight
		L = length
Size-spectrum	$P = \exp (25.22 - E/kT) *$	E = activation energy of metabolism
production	W ^{0.76}	k = Boltzmann's constant
(Jennings et al., 2008;		T = temperature in Kelvin (°C+273)
Fernandes et al., 2013)		
Intrinsic population	G=r * A * (1 – (A/KC))	r = intrinsic rate of population increase
growth rate (Hilborn &		A = the relative abundance
Walters,1992)		KC = population carrying capacity
Larval dispersal		D = diffusion parameter
(Hundsdorfer & Verwer		(u, v) = velocity parameters
2003;		LAV = larvae recruitment
Cheung et al., 2008)		
Adult movement	Cm * h-1	Cm = centimetre
		h = hour
Natural mortality	M = -0.4851 - 0.0824 *	Winf = asymptotic weight
	log(Winf) + 0.6757 * log (K)	K = von Bertalanffy growth parameter
	+ 0.4687 * log(T)	T = average water temperature in the animal

255 mechanisms in SS-DBEM. Further details are given in the associated references.

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The SS component of the model addresses resource competition between different 257 species co-occurring in any given cell by comparing the biomass that can be supported in 258 the cell, as determined from primary production and the size-spectrum model, with the 259 energy demanded by the abundance of the species predicted to inhabit this cell. This 260 allocation is based on habitat suitability considerations and a generic group (other 261 species) that can also compete for energy particularly if there is a surplus is available 262 (Fernandes et al 2013a). If the energy demanded by all species in the cell exceeds the 263 energy available, then the model allocates available energy to each species in proportion 264

to its energy demands. If the energy demanded by all the species is lower than the energy
available, the surplus energy is allocated according to the proportional energy demand of
the species present. The rate at which this energy can be assimilated is limited by
constraints on species' growth rates as described in Fernandes et al (2013a).

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The model can consider fishing pressure in relation to maximum sustainable yield
(MSY). MSY is defined as the highest average theoretical equilibrium catch that can be
continuously taken from a stock under average environmental conditions (Hilborn &
Walters, 1992). Based on a simple logistic population growth function and under
equilibrium conditions, MSY can be defined as:

275

276 MSY = $B\infty * intR / 4$

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278 where intR is the intrinsic rate of population increase and $B\infty$ is the biomass at carrying capacity (Schaefer, 1954; Sparre and Venema, 1992). In our application, the intR values 279 280 are calculated based on natural mortality (Pauly 1980; Cheung et al., 2008). This is an 281 approximation and not as reliable as estimates of biomass using survey-based methods (McAllister et al. 2001; Pauly et al., 2013). However, these estimates have proven to be 282 significantly correlated with those from aggregated stock assessments (Froese et al., 283 284 2012; Fernandes et al., 2013). This fishing mortality is applied uniformly across all the cells according to scenarios of fishing (e.g. 0.8 or 1.2 times MSY) and do not aims to 285 reproduce exact past fisheries effort distribution. Therefore, in this work no fishing 286 mortality scenario was activated in the model projections. Future work based on catch 287 and fishing effort reconstructions (Watson et al., 2017; Taconet et al., 2019) may allow 288 estimates of historical non-uniform fishing mortality to be included in models. 289

290 <u>Fisheries survey data</u>

Eleven standardised and long-term fisheries surveys from 1982 to 2007 covering all year 291 seasons (Simpson et al., 2011) were used to validate the model (Fig. 1). These surveys 292 (AFBI Iris Sea Q3, AFBI Irish Sea Q1, CEFAS Celtic Sea, CEFAS Eastern Channel, CEFAS 293 Irish Sea, CEFAS North Sea, CEFAS Western Channel, FRS NW Scotland Q1, FRS NW 294 295 Scotland Q4, ICES IBTS North Sea Q1, MBA Western Channel) were collated by Simpson et al. (2011) and now available at ICES DATRAS online database (www.ices.dk/marine-296 data/data-portals/Pages/DATRAS.aspx). To control for the differing effort between 297 surveys, the swept area for each haul (over 22 000 hauls with six different gears) was 298 calculated using estimates of wing-spread for Grande Ouverture Verticale trawls from 299 Fraser et al. (2007): 300

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302 Area swept = $(((6.85 * (log10(depth))) + 5.89) * distance) / 10^{6}$

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Where area swept is in km2, and depth and distance are in meters. A tow speed of 4 knots (7.4 km h⁻¹) was assumed for the duration of the haul, except for data from the Celtic sea collected by Agri-Food and Biosciences Institute, Belfast, UK, where the data were originally provided as number of individuals per 3 nautical miles (~5.6 km). The 10^6 scaling converts from m² to km².

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Catchability estimates (Table 3) were used to provide more robust estimates of abundance for each species by sizes (Sparholt 1990; Fraser et al., 2007). An average catchability estimate was applied to similar species, where individual species values were not available.

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315 Corrected abundance = uncorrected abundance*(1/catchability)

- 317 Table 3. Table summarizing catchability correction values by species and sizes (based on
- 318 Simpson et al., 2011).

Species	Size (cm)	Catchability correction
Chelidonichthys lucerna	1-20, 21-22, 23-25, 26-30, 31-33, 34, 25-	0.21, 0.18, 0.17, 0.16, 0.14, 0.13, 0.11
	80	0.1
Clupea harengus	1-100	0.1
Gadus morhua	1-31, 32-36, 37-39, 40-42, 43-45, 46-47,	0.16, 0.17, 0.18, 0.19, 0.2, 0.21, 0.22, 0.23,
	48-49, 50-51, 52-53, 54-55, 56, 57-58,	0.24, 0.25, 0.26, 0.27, 0.28, 0.29, 0.3, 0.31,
	59, 60-61, 62, 63-64, 65, 66, 67, 68, 69-	0.32, 0.33, 0.34, 0.35, 0.36, 0.37, 0.38,
	70, 71, 72, 73, 74, 75, 76, 77, 78, 79, 80,	0.39, 0.4, 0.41, 0.42, 0.43, 0.44, 0.45, 0.46,
	81, 82, 83, 84, 85, 86, 87, 88, 89, 90, 91,	0.47, 0.48, 0.49, 0.5, 0.52, 0.53, 0.54, 0.55,
	92, 93, 94, 95, 96, 97-200	0.56, 0.58, 0.59, 0.6, 0.61, 0.63, 0.64, 0.65,
		0.67
Glyptocephalus	1-20, 21, 22-23, 24, 25, 26-27, 28-31, 32-	0.11, 0.09, 0.08, 0.07, 0.06, 0.05, 0.04,
cynoglossus	34, 35-38, 39-60	0.03, 0.04, 0.05
Hippoglossoides	1-5, 6, 7, 8, 9, 10, 11, 12, 13, 14, 15, 16-19,	0.11, 0.16, 0.2, 0.23, 0.26, 0.29, 0.32, 0.34
platessoides	20, 21, 22, 23, 24, 25, 26, 27, 28, 29, 30-	0.36, 0.37, 0.38, 0.39, 0.38, 0.37, 0.35,
	60	0.33, 0.31, 0.29, 0.26, 0.22, 0.19, 0.15, 0.1
Lepidorhombus whiffiagonis	1-60	0.06
Melanogrammus	1-11, 12, 13, 14, 15, 16, 17, 18, 19, 20, 21,	0.01, 0.04, 0.06, 0.09, 0.11, 0.13, 0.15,
aeglefinus	22, 23, 24, 25, 26-27, 28-33, 34-35, 36-	0.17, 0.19, 0.2, 0.21, 0.23, 0.24, 0.25, 0.26
	37, 38, 39, 40, 41, 42, 43, 44, 45, 46-80	0.27, 0.28. 0.27, 0.26, 0.25, 0.24, 0.22,
		0.21, 0.2, 0.18, 0.16, 0.14, 0.12
Molva molva	1-11, 12-200	0.05, 0.1
Raja montagui	1-200	0.15
Sardina pilchardus	1-100	0.1
Scomber scombrus	1-100	0.19
Scophthalmus maximus	1-21, 22-24, 25-28, 29-31, 32-34, 35-38,	0.05, 0.06, 0.07, 0.08, 0.09, 0.1, 0.11, 0.12
	39-41, 42-200	
Scyliorhinus canicula	1-200	0.22
Solea solea	1-22, 23-60	0.04, 0.05
Sprattus sprattus	1-100	0.1
Trachurus trachurus	1-100	0.19
Trisopterus esmarkii	1-100	0.19
Trisopterus luscus	1-11, 12-60	0.05, 0.1

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321 <u>Model validation</u>

All survey data were aggregated into a $0.5^{\circ} \times 0.5^{\circ}$ cell grid to match the SS-DBEM grid. To 322 compare simulated changes with observations from surveys, time-series at different 323 aggregation scales (e.g. aggregating all the cells, or all the species in a cell) were extracted 324 and projected biomass was scaled to lie between 0 and 1. Since multiple species at 325 multiple cells are considered, we ensured that results were comparable by omitting time-326 series of survey data with more than 3 years of missing data. Then, time-series output 327 from the SS-DBEM models were extracted for the remaining years, species and cells 328 where there were data from the surveys at the $0.5^{\circ} \times 0.5^{\circ}$ grid and yearly resolution. This 329 restricted the data that could be analysed to the time period 1982 to 2007 (26 years) and 330 to 18 species, including 5 pelagic species: Atlantic herring (*Clupea harengus*), Atlantic 331 332 mackerel (*Scomber scombrus*), Atlantic horse mackerel (*Trachurus trachurus*), European pilchard/sardine (Sardina pilchardus), sprat (Sprattus sprattus); and 13 demersal 333 species: haddock (Melanogrammus aeglefinus), plaice (Hippoglossoides platessoides), 334 witch (*Glyptocephalus cynoglossus*), megrim (*Lepidorhombus whiffiagonis*), cod (*Gadus* 335 morhua), common sole (Solea solea), lesser spotted dogfish (Scyliorhinus canicula), 336 Norway pout (Trisopterus esmarkii), turbot (Scophthalmus maximus), tub gurnard 337 338 (Chelidonichthys lucerna), pouting/bib (Trisopterus luscus), ling (Molva molva) and spotted/cuckoo ray (*Raja montagui*). We chose these species as they are both 339 commercially important for local fisheries and include some species with distributions 340 centred on the North Sea, Celtic Sea and Irish Sea alongside those with more northerly 341 (high latitude) and southerly (low latitude) centres of distribution. 342

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We compared model projections with data at different scales, spanning those used in 344 previously published projections of climate-driven changes in fish distribution and/or 345 abundance (e.g. Fernandes et al., 2016; Fernandes et al., 2017; Jones et al., 2013; Mullon 346 et al., 2016; Queiros et al., 2015; Queiros et al., 2016; Coccoli et al., 2018). We consider 347 the following (Fig. 1): (1) all the cells are considered for all the species (325 cells and 18 348 349 species) named "Species by cells" in a first spatial validation; (2) time series for individual species aggregated over all cells ("All species" or individual species) in a first temporal 350 validation; and, (3) time series for each of the 325 cells where all the species are 351 aggregated ("All cells" or individual cells) in a spatial and temporal validation, and, (4) a 352 time series for each species aggregating all the cells for which survey information was 353

available (species cells sum). A species could be present in between 75 (turbot) and 325

355 (Herring) cells.

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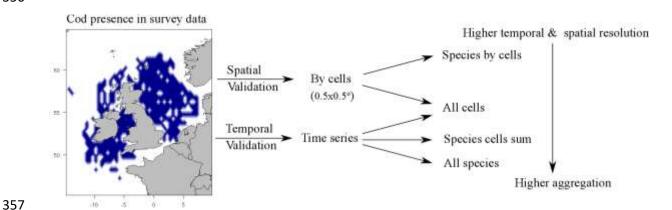


Figure 1. Example of survey data for cod (one of the species with more coverage) in Celtic Sea, Irish Sea and
North Sea (left side). Scheme of spatial and temporal validation resolutions in the right side.

The time-series 1-3 were generated for both the survey data and the model projections and comparisons between data and projections were reported as absolute error (AE):

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$$364 \quad AEj = |p_j - x_j| \tag{1}$$

365

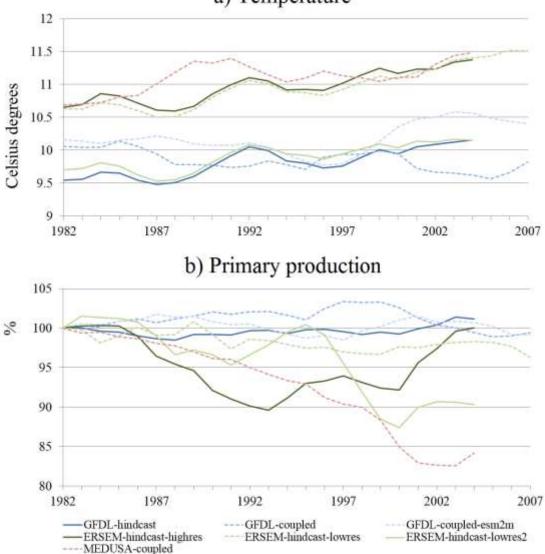
where, *p* is the scaled biomass predicted in a SS-DBEM model in a particular year for a species, and *x* is the scaled biomass from the survey. The use of scaled values enables direct comparison of data and SS-DBEM projections across species and levels of aggregation.

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Results 372

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374 The main environmental drivers of SS-DBEM and other fisheries models are temperature and net primary production, which are obtained from the ocean biogeochemical models. 375 Most forcing models show an increasing trend in temperature over the 1982-2007 period 376 for which we have fish survey data (Fig 2). There is spread in the absolute temperature 377 with some models simulating temperatures of 10.5 to 11.5 °C in cells included in the fish 378 surveys, while others projecting temperatures of 9.5 to 10.5°C (Fig. 1). Trends in primary 379 380 production are more uncertain than those in temperature, with some models showing no trends, while others project decreases of 10-15% over the 1982-2007 period. 381



a) Temperature

Figure 2. Simulated annual sea surface temperature and net primary production changes from 1982 to 384 2007 of different biogeochemical models. The reported values are the average across cells where fish 385 survey data is considered. The time-series has been smoothed with a five-year moving average. Solid lines 386 indicate reanalysis projections and dashed lines fully coupled projections.

There are multiple examples where there is a good fit between SS-DBEM projections and survey data with high correlation (Pearson) among projections and survey data (Fig. 3): herring between 0.73 (GFDL-hindcast) and 0.96 (GFDL-coupled), Cod 0.76 (MEDUSAcoupled-forced), haddock between 0.92 (ERSEM-hindcast-lowres2) and 0.95 (GFDLcoupled), sprat between 0.88 (GFDL-hindcast) and 0.96 (ERSEM-hindcast-lowres), sole between 0.81 (ERSEM-hindcast-highres) and 0.91 (GFDL-coupled),

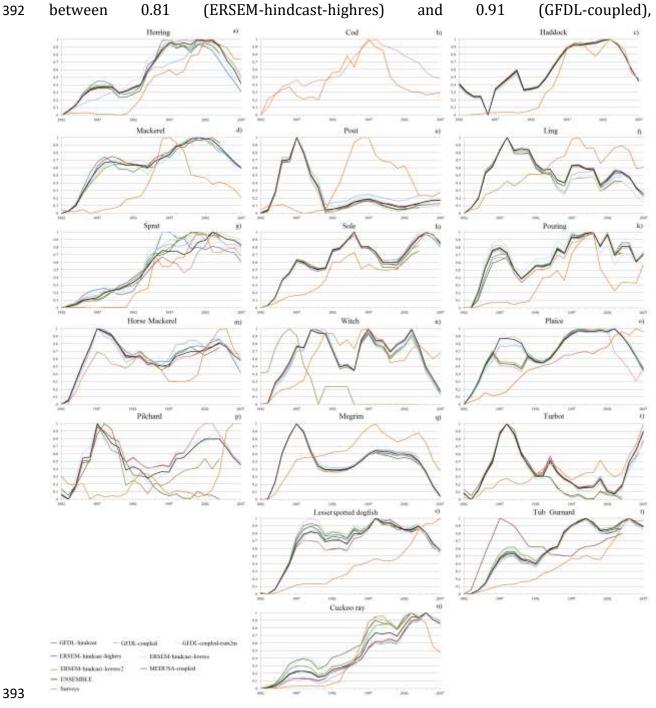


Figure 3. Scaled biomass projections (5 years moving average) for 5 pelagic and 13 demersal species with
 different biogeochemical model forcings in cells where fish survey abundance data were also available.
 Note that Cod projections failed with several inputs from biogeochemical models.

pouting between 0.65 (GFDL-coupled-esm2m) and 0.78 (ERSEM-hindcast-highres), 398 horse mackerel between 0.69 (ERSEM-hindcast-lowres) and 0.76 (MEDUSA-coupled-399 forced), plaice between 0.49 (MEDUSA-coupled-forced) and 0.87 (ERSEM-hindcast-400 highres), or cuckoo ray between 0.79 (GFDL-coupled-esm2m) and 0.99 (GFDL-coupled). 401 Witch is another species where most models seem to perform well with correlations 402 between 0.59 (ERSEM-hindcast-lowres2) and 0.75 (GFDL-coupled), but where one model 403 performs very bad showing high negative correlation of -0.74 (ERSEM-hindcast-lowres). 404 Model runs for mackerel and megrim do not perform so well, but still show competitive 405 correlations between 0.58 (GFDL-hindcast) and 0.66 (ERSEM-hindcast-lowres) for 406 mackerel, and between 0.45 (ERSEM-hindcast-highres) and 0.54 (GFDL-coupled) for 407 Megrim. Model runs for dogfish and tub can achieve high correlations of 0.74 (GFDL-408 409 coupled) and 0.87 (GFDL-coupled) respectively, however some of the models show also 410 very low performances of 0.36 (ERSEM-hindcast-lowres2) and 0.3 (MEDUSA-coupledforced) respectively. Finally, a few species show very low correlations for all model runs 411 (Ling between 0.01 and 0.28) or even negative correlations: turbot between -0.17 and 412 0.44, pout -0.41 and -0.47, and pilchard between -0.14 and -0.54. 413

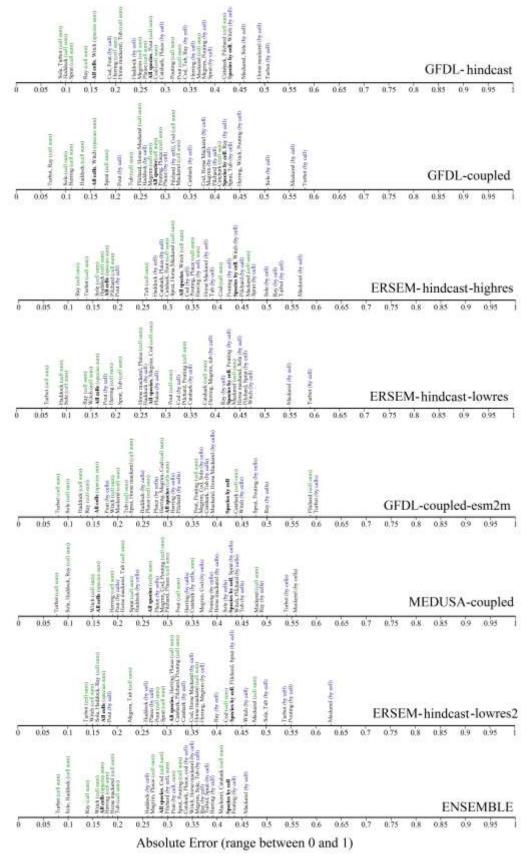
414 These results suggest that an increase in spatial resolution does not necessarily mean a higher performance in species biomass trends. The highest resolution model (ERSEM-415 hindcast-highres) sometimes shows higher correlation with survey data for some 416 demersal species (pouting and plaice) than the lowest resolution models. However, for 417 other pelagic and demersal species often the lowest resolution model runs (GFDL) are 418 419 the ones showing the highest correlations (e.g. herring, haddock, sprat, sole). In other cases, although the higher resolution model runs may show superior correlations these 420 may be very close to lower resolution model correlations and not be statistically 421 significant different from them (paired t-test). 422

Fits tend to improve in later years in the time series for several species, perhaps reflecting the reduction in fishing mortality (Simpson et al., 2011) and its effect on survey data distributions in later years (Fig. 3). However, this could also be the result of a more representative area being covered by the surveys since their geographical extend has increased over time (Baudron et al., 2020).

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The different biogeochemical models generate rather similar error, with more variation 429 between models for the pelagic fish species and for sharks and rays (Fig. 4). In "Species 430 by cells" comparisons (1) median error is 0.42-0.44, with similar results between runs. 431 At the "All species" or individual species level (2), the median of the error is 0.27-0.31. 432 This means that when looking at the species aggregated biomass, the SS-DBEM is more 433 reliable and that there is rather little impact of the biogeochemical chosen for 434 environmental conditions projection. The lowest error values correspond to the 435 reanalysis run (GFDL) and the latest runs (ascending order in Table 1) showing that 436 improvements in the forcing models (e.g. new data or reparameterization) are translating 437 into more reliable projections from the SS-DBEM. However, the performance differences 438 are small in most species and levels of aggregations which highlight the higher role of the 439 440 model uncertainty in the fisheries model. With "All cells" (3) the error range is 0.16-0.18 441 which shows the ability of the forcing and SS-DBEM model to forecast variations in total biomass at 0.5° resolution. Ensemble model results are within those error ranges and do 442 not improve performance when working with "All cells" (Fig. 2i). 443

In general, errors are smaller for "All species" (2) than for the species in each cell 444 445 ("Species by cells", 1). The mean and standard deviation for the former are 0.25±0.11 and 0.40±0.14 for the latter (Fig. 4). This confirms higher model performance for total 446 biomass than for species biomass in each cell. The results by species can help to identify 447 where the SS-DBEM needs improvement. In general, widely distributed pelagic species 448 (e.g. Herring, Horse mackerel, mackerel, sprat) have among the highest errors at the 449 450 "Species by cells" (1) level, but a better performance when the total species biomass in each cell is considered ("All species", 2). This indicates that there is something 451 systematically wrong with the way that the SS-DBEM handles those species, even though 452 the general allocation of biomass to pelagics is reasonable. Note, however, that the 453 bottom trawl surveys are not designed to sample pelagic species (ICES, 2016) and 454 relatively short tows with limited time in the water column can provide a misleading 455 456 picture of abundance and distribution for predominantly shoaling species (Battaglia et al. 2006). At the "Species by cells" (1) level the species for which the models perform best 457 are pout, haddock, plaice and lesser-spotted dogfish, all bottom-dwelling species that are 458 effectively sampled by the survey gear (Fraser et al 2007; Walker et al. 2017). 459



461 Figure 4. Median of the absolute error at different scales of analysis for species and overall (bold) where (1) all the 462 cells are considered for all species (325 cells and 18 species) named "Species by cells"; (2) time series for each species 463 where all the cells are considered aggregated ("All species"); (3) time series for each of the 325 cells where all the 464 species are aggregated ("All cells" (species sums)); and, (4) a time series is produced for each species aggregating all 465 the cells with survey information (species cells sum). See text for additional explanation.

466 **Discussion**

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468 While there was generally good agreement among biogeochemical models on the sign and extent of the temperature increase, the trends in primary production were less 469 consistent. These differences can be attributed to a variety of factors such as model 470 uncertainty and internal model variability (Hawkins and Sutton, 2009; Naujokaitis-Lewis 471 et al., 2013; Payne et al., 2015, Cheung et al. 2016, Frölicher et al. 2016). Especially on 472 local-to-regional scale (i.e. North-East Atlantic), model uncertainty and internal 473 474 variability may play a dominant role in differences across model runs (Hawkins and Sutton, 2009; Chust et al. 2014; Frölicher et al. 2016). Here it seems that model 475 476 uncertainty/resolution may play a bigger role than internal variability given that uncertainty in NPP trends is as large or larger for models forced in the same way 477 (reanalysis) than for models that are forced differently (i.e. coupled or forced), but share 478 479 the same biogeochemical model. For example, Laufkötter et al. (2015) have identified 480 how uncertainty in the representation of underlying physiological processes influences the trends in net primary production. Here, we can observe in the multiple runs of 481 historical and forced GFDL the uncertainty due to internal variability seems to be smaller 482 than model uncertainty in marine biogeochemical models. Detailed comparisons of six 483 different biogeochemical models (including MEDUSA and ERSEM used here) in 484 physically-identical model simulations have shown the importance of the underlying 485 modelled biogeochemistry on biogeochemical indicators including primary production 486 487 (Kwiatkowski et al., 2014). However, the main characteristics of the trends of the two sets of environmental variables from the biogeochemical models used in this study to 488 drive the fisheries simulations are in line with the results from global model data sets 489 490 (Steinacher et al. 2010, Laufkötter et al. 2015, Frölicher et al. 2016).

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Existing studies have made projections of fish distribution and abundance at scales ranging from higher resolution (Cheung et al., 2011; Burrows et al., 2014; Molinos et al., 2015) to higher level of aggregation (Mullom et al., 2016; Fernandes et al., 2016). Many studies, including those for fisheries management, only need information at relatively low level of aggregation such as LMEs, FAO areas, seas, ICES areas, EEZs or subregions (Blanchard et al., 2012; Fernandes et al., 2016; Mullon et al., 2016; Queiros et al., 2015; Queiros et al., 2018). Applications not dependent on high resolution data are often used

for management and economic research, linked with long term scenarios (Mullon et al. 499 2016; Queiros et al., 2018) or ecological studies looking at overall impacts on specific 500 species or habitats (Queiros et al., 2015). However, other studies and management 501 research aims for or requires a higher level of detail such as cells of 0.5° x 0.5° or smaller 502 as in the case of ICES rectangles or even 1x1 km² for local marine spatial planning or 503 studies of shifts of species abundance centroids (Jones et al., 2013; Queiros et al., 2016; 504 Fernandes et al., 2017; Coccoli et al., 2018). Some applications requiring higher resolution 505 are marine spatial planning (Queiros et al., 2016, Coccoli et al., 2018) or studies about 506 507 shifts of species abundance centroid (Jones et al., 2013).

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The widely reported regime shift in the late 1980's in the North Sea (Reid and Edwards 509 510 2001, Beaugrand 2004, Weijerman, Lindeboom et al. 2005) suggested to have affected several species in this analysis (Fig. 4) providing evidence of biogeochemical drivers for 511 512 the shifts. The reported shifts are sparse such as the increase in horse mackerel, sprat and reduction of cod in abundance (Reid, Borges et al. 2001, Alheit, Mollmann et al. 2005) 513 despite other species shifts might have occur. However, a recent publication shows 514 species distribution shifts for 17 main widely distributed species in the North East 515 Atlantic (Baudron et al., 2020). The comparison of the performance of the models at 516 different resolutions shows, in general, that outputs are more reliable when aggregated 517 at larger time- and space- scales. As such, these would be favoured for reporting. 518 However, users of projections will often seek projections for individual species of 519 conservation or fisheries significance and small areas that reflect those accessible to, or 520 521 used by, a defined fishing fleet. For example projections at small space and time scales may be requested to assess the abundance of "choke" species under climate change, given 522 their potential effects on the capacity of a fishery to access available quota (Baudron and 523 Fernandes, 2015). 524

525

Predicted latitudinal shifts of species are difficult to compare with empirical data since data collection is often focused in areas where species have been distributed in the past (ICES, 2016). This is the case of our survey data (Simpson et al., 2011) where sampling centres on the Celtic Sea, Irish Sea and North Sea (up to ICES area IVa) or sprat surveys that are focused in the Baltic Sea. Sparse data are available for ICES area IVa (ICES, 2016)

and North of the Faroe Islands. However, species latitudinal shifts have been observed in 531 the global catch data (Cheung et al., 2013). There is limited knowledge North of ICES area 532 IVa due to limited international cooperation to share survey data and even South of ICES 533 area IVa the surveys mostly target higher value demersal species (ICES, 2016). However, 534 the aim of models such as SS-DBEM is not to predict accurately (Planque et al., 2016; 535 Dickey-Collas et al., 2014) where the species are present or will be present, but to 536 highlight the species and areas where changes are more likely to happen, and generate 537 uncertainty estimates (Planque et al., 2016; Payne et al., 2016). There is a trade-off 538 between goodness-of-fit and generalization power (Fernandes et al., 2015). Models that 539 very precisely represent the present have a performance that deteriorates faster as 540 projections are made further into the future (Rutterford et al., 2015). 541

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543 At the highest output resolution we considered $(0.5 \times 0.5 \text{ degrees by year and species})$ error was relatively high. Rutterford et al. (2015) also compared observed data with 544 predicted values using a GAM to project changes in distribution and abundance of 545 demersal species, obtaining correlations >0.5 for a 10 year forecast, but decreasing to 546 \leq 0.5 at 10 year or longer-term forecasts. Statistical methods based on observations may 547 548 have limited value in long-term forecasting of systems that are expected to depart markedly from their past state due to the long-term impacts of climate change (Barnsley 549 et al., 2007; Queiros et al., 2015; Payne et al., 2016). In addition, statistical models that 550 consider interactions between species are rare (Fernandes et al., 2013b), whereas this is 551 more common in mechanistic models (Blanchard et al. 2012; Fernandes et al., 2013a; 552 553 Thorpe et al., 2015). The survey data used in statistical models are relatively costly to obtain and in many regions such observations are sparse or non-existent (e.g. 554 Bangladesh; Fernandes et al., 2016). 555

556

Variation between biogeochemical models is limiting predictions for several pelagic species, whereas for most of the remaining species the choice of the biogeochemical model makes little difference. Therefore, for most of the species the validation performed here can guide improvement in fisheries models. Another alternative is the use of model ensembles as a means of increase the reliability of projections (Araújo & New, 2007; Jones et al., 2012; McKenna et al., 2013; Scales et al., 2015). In this study, the ensemble of forcing-model runs did not affect performance of the SS-DBEM on average, but reduced

the probability of extreme errors, which would be an important consideration for some applications. Outcomes from any given ensemble are highly dependent on the diversity of constituent models, so it is unlikely our result can be generalised. Recent research is moving towards ensembles of biological models which could benefit from similar validation exercises (Lotze et al., 2019; Hermann et al., 2019) and collaborative protocols for integration and comparison of multiple fisheries models (Tittensor, et al. 2018).

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576 Acknowledgements

577 Simon Jennings provided valuable inputs to the early stages of this work. The research was funded by the European Union's 7th Framework and the Horizon 2020 programs 578 under the EURO-BASIN (Grant Agreement No. 264933) and CERES (Grant Agreement No. 579 678193) projects. The article does not necessarily reflect the views of the Commission. 580 581 T.L.F. acknowledges support from the Swiss National Science Foundation under grant PP00P2_170687 and the European Union's Horizon 2020 research and innovation programme 582 under grant agreement No 820989 (project COMFORT, Our common future ocean in the Earth 583 system -- quantifying coupled cycles of carbon, oxygen, and nutrients for determining and 584 achieving safe operating spaces with respect to tipping points). A.Y. acknowledges support 585 from the UK Natural Environment Research Council (NERC) under the Regional Ocean 586 Acidification Modelling project (ROAM; grant NE/H017372/1) and National Capability in 587 Ocean Modelling funding, and European Union Horizon 2020 support provided by the 588 CRESCENDO project (grant number 641816). Jose A. Fernandes received further funding 589 through the Gipuzkoa Talent Fellowships programme, by the Gipuzkoa Provincial 590 Council, Spain. 591

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