

1 **Refining palaeoenvironmental analysis using integrated quantitative**
2 **granulometry and palynology**

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12
13 **Abstract**

14 Accurate palaeoenvironmental analysis is at the heart of producing reliable
15 interpretations and depositional models. This study demonstrates a multivariate
16 statistical approach to facies analysis based on relationships between grain size and
17 quantitative palynology. Our methodology has the advantage that it can be used on
18 small amounts of sample, such as core or well cuttings, as the basis for facies
19 analysis.

20 Proof of concept studies involving collection of grain-size and palynological datasets
21 from well exposed outcrops of the Middle Jurassic, Lajas Formation of the Neuquén
22 Basin, Argentina, demonstrate that canonical correspondence analysis can be used to
23 consistently recognize facies and aid in the determination of depositional
24 environments. This study demonstrates the link between depositional facies, grain-
25 size distribution, palynomorph hydrodynamics, and assemblage taphonomy of
26 palynomorphs. This knowledge can be transferred into a semi-automated statistical
27 facies prediction technique for the subsurface in complex depositional settings,
28 particularly when calibrated against conventional sedimentary facies analysis.

29 **Supplementary material:** The full set of grain size data and statistical scores are
30 available at: <http://www.geolsoc.org.uk/SUP0000>

31 **Introduction**

32 Palaeoenvironmental interpretation is at the core of effective reservoir
33 characterization and sequence stratigraphic analysis. While significant advances have
34 been made with integrated palaeontological/sedimentological facies analysis of highly
35 heterolithic deltaic facies in outcrop and core (McIlroy 2004, 2008), the same analysis
36 from drill cutting material would be problematic. Detailed in this study is the
37 analysis of two primary datasets that can be obtained from well cuttings (grain size
38 attributes and palynology) from a well-documented outcrop framework (Brandsæter
39 *et al.* 2005; McIlroy *et al.* 2005; McIlroy *et al.* 1999).

40 Significant advances in the understanding of tide-dominated deltas have been
41 made in recent years through study of the Lajas Formation of the Neuquén Basin in
42 Argentina (McIlroy *et al.* 2005; McIlroy 2007; McIlroy 2004; Martinius *et al.* 2000;
43 Ichaso & Dalrymple 2009; Brandsæter *et al.* 2005; Rossi & Steel 2016). The
44 significance of tide-dominated and tide-influenced deltas stems from their importance
45 as petroleum reservoirs, especially offshore Mid-Norway (McIlroy 2004; Martinius *et*
46 *al.* 2000; Ichaso & Dalrymple 2009) and the Northern North Sea (Maxwell *et al.*
47 1999).

48 Facies characterization and palaeoenvironmental interpretation is notoriously
49 difficult in tide-influenced deltaic deposits, requiring the full breadth of
50 palaeontological and sedimentological tools to be applied (Martinius *et al.* 2011; van
51 Cappelle 2016). Sedimentological, ichnological and palaeobotanical data has already
52 been integrated into facies analysis in the studied sections (McIlroy *et al.* 2005;
53 Morgans-Bell & McIlroy 2005). The linkage between the palynology and
54 palaeoenvironment of the hinterland floras, and its application to the understanding of
55 the coastal depositional system has been investigated previously (Stukins *et al.* 2013).
56 This study aims to build on this earlier work by developing an understanding of the
57 depositional dynamics of palynomorphs in differing environments within a tide-
58 dominated depositional setting.

59 *Geological context*

60 Middle Jurassic strata of the Lajas Formation crop out extensively in the Sierra de
61 Chacaico of Neuquén Province, Argentina (Fig. 1a). The Lajas Formation (Fig. 2) has
62 been widely studied in this region and is interpreted as marginal marine, deltaic unit
63 with a strong tidal influence, high sediment supply, and rapid subsidence (McIlroy *et al.*
64 *1999*; McIlroy *et al.* 2005; Zavala 1996), though there is ongoing debate regarding
65 the relative contribution of fluvial, wave and tidal processes through the succession
66 (Gugliotta *et al.* 2015; Rossi & Steel 2016).

67 Four stratigraphic sections, chosen to encompass a range of depositional facies at
68 different stratigraphic positions within the Lajas Formation (Fig. 1b), were logged
69 and the fine-grained facies were sampled for palynology and granulometry (Stukins *et al.*
70 *2013*).

71 1) The ~40m thick progradational “Perigrina” parasequence set is dominated by
72 mudstones and siltstones and inter-bedded with tabular sandstones in upward
73 coarsening parasequences capped by distributary channel sandstones. These fine-
74 grained facies are interpreted as pro-delta mudstones inter-bedded with prodelta
75 turbidites associated with the first pulse of progradation of the Lajas Formation into
76 the offshore facies of the Los Molles Formation (McIlroy *et al.* 2005).

77 2) The Quilmez Parasequence is dominated by a mud-rich facies succession that
78 varies in stratigraphic thickness laterally from 2.8 m to 11.1 m, and is associated with
79 heterolithic sandstones and mudstones. The thick, mudstone-dominated bayfill facies
80 passes stratigraphically and laterally into tidal flat and abandoned channel fill
81 deposits related with the bay-margin, and intertidal depositional environments
82 (McIlroy *et al.* 2005).

83 3) The Norwegian Parasequence is a highly heterolithic package of sandstones and
84 mudstones that is capped by oyster banks, which are in some places reworked by
85 processes concurrent with the Owl Sequence boundary (McIlroy *et al.* 1999). The
86 tidal flat dominated parasequence (*sensu* McIlroy *et al.* 1999) includes facies from
87 sand/mud/mixed tidal flats, tidal channel fills, abandoned tidal channel fills, bayfills,
88 mouthbar, as well as both autochthonous and allochthonous shell beds (McIlroy *et al.*
89 2005; McIlroy *et al.* 1999; Brandsæter *et al.* 2005).

90 4) The Dagna Parasequence marks the transition between the marginal marine Lajas
91 Formation and the dominantly fluvial Challacó Formation at the northern end of the
92 Sierra de Chacaico. The studied section is composed of upward coarsening
93 heterolithic sandstones and mudstones inferred to have been deposited in river-
94 dominated mouthbar settings. These mouthbar deposits pass distally into organic rich
95 micaceous mudstones of lagoonal/bayfill character (McIlroy *et al.* 2005). The fine-
96 grained mouthbar facies are cut by very coarse grained sandstones interpreted as delta
97 top weakly tide-influenced river channels cutting older progradational
98 lagoonal/bayfill facies. High in the section, but below the main Challacó Formation
99 fluvial sandstones, there is a palaeosol horizon with a rooted horizon overlain by a
100 thin coal.

101 Previous studies, based on macroflora and palynofloral data (Stukins *et al.* 2013;
102 Quattrocchio *et al.* 2007), have proposed palaeoecological models for the floras of the
103 Jurassic of the Neuquén Basin. The palaeofloral ecological dynamics have previously
104 been published for the studied sections (Stukins *et al.* 2013) and are used as a point of
105 reference for this study.

106 **Methods**

107 For palynological analysis the samples were processed using 40% hydrofluoric acid
108 and the resulting residue sieved through a 7 μ m mesh then, where necessary, boiled in
109 40% hydrochloric acid to remove any precipitate. The residues were applied to cover
110 slips using 2% polyvinyl alcohol and then onto slides using two part epoxy medium.
111 The sample slides were counted using transmitted light microscopy for a minimum of
112 250 specimens per slide, abundance permitting, to create a raw data-set. The raw
113 count data was converted into abundance percentages to normalize the data.

114 All of the fine-grained samples analysed for their palynological content had their
115 grain size distributions determined using a Beckman Coulter LS 13 320 LS Particle
116 Size Analyzer. Grain size and grain sorting were identified as the most useful factors
117 for comparison with the sedimentological facies recognised in the field, and for
118 linkage with the palynological data collected herein. The parameters used as proxies
119 for grain size and grain sorting are the <50 percentile (%<50) and the mean/median
120 ratio (m/m ratio) respectively (adapted from (Friedman 1962; Inman 1952)).

121 *Canonical correspondence analysis (CCA)*

122 The possibility of linking granulometric analysis results with palynological data
123 through canonical correspondence analysis (CCA) to help determine
124 palaeoenvironment of deposition is explored in this study using the integrated
125 palynological and granulometric datasets. Correspondence analysis (CA) is an
126 ordination method that can display ecological information inferred from axes of
127 eigenvalues representing directions of variation in the dataset from given species
128 assemblage data. The eigenvalues derived from the statistical analysis are a
129 coefficient reflecting the degree of species dispersion along an axis, hence, a measure
130 of relative importance of the given axis in explaining the species variation (Ter Braak
131 1986; Kovach 1993; Dale *et al.* 2005). Similar eigenvalues are interpreted to show a
132 relationship based on the variant of the given axis.

133 Canonical correspondence analysis is an extension of CA, in which an additional
134 matrix of environmental data is incorporated into the analysis. Ordination axes
135 (represented by vectors) are added for these known environmental variables (Ter
136 Braak 1986; Kovach 1993). In this study, these ordination axes are the
137 sedimentological properties of the fine-grained sediment that the palynomorphs were
138 recovered from.

139 In this study, a matrix of samples against palynomorph assemblage data and also the
140 sedimentological properties provide the raw data. CCA of this data was analysed
141 using the software package MVSP (Kovach 1993). The output of which are
142 eigenvalues (degrees of variance) for the samples and palynomorph species identified
143 within them. Canonical coefficient scores were produced for the sedimentological
144 properties to form linear combinations of the environmental variables. These form the
145 environmental vectors that can be projected onto cross-plots of the eigenvalues.

146 We approach the development of an understanding of the dataset by first compiling
147 the whole Sierra de Chacaico dataset in order to analyse broad-scale trends in
148 granulometric and palynomorph assemblage data. Once the coarse-scale trends have
149 been determined, we then focus in on a subset of data from the Norwegian
150 Parasequence in Rhea Gorge (cf. Brandsæter *et al.* 2005) to identify more detailed
151 relationships between potential sedimentological processes in a progradational tidal
152 flat rich succession and the palynological assemblages.

153 **Grouped Sierra de Chacaico dataset analysis**

154 Canonical correspondence analysis, applied to the palynological and granulometric
155 datasets, calculated a five axes solution for the data, where the third and fourth axes
156 have the greatest percentage variance. In this analysis however, constrained by the
157 software, the known palaeoenvironmental variables are constrained only to the first
158 two axes. Nevertheless, there is enough confidence in the statistics that there is
159 adequate variance in the first two axes (eigenvalues of 0.027 and 0.008 but
160 cumulative constructive percentages of 76.041% and 100%) to represent grain size
161 related relationships for analysis.

162 When the data from the first two axes are presented on a cross plot (Fig. 3), the
163 relationships between the samples/inferred facies, palynological data and the grain
164 size data can be determined. CCA reveals that most of the samples from the same
165 reported depositional environment cluster together as is expected (Fig. 3).

166 There is a facies-related characteristic represented in the multivariate dataset in the
167 form of the linear trend within samples identified from the field as being bayfill
168 mudstone deposits (Fig. 3). The linear trend is the approximate bisector of the two
169 environmental vectors. It is not possible to determine whether this trend implies that
170 grain size and grain sorting have an almost equal control on the palynological
171 assemblages of bayfill mudstones, or whether that there is another unknown
172 palaeoenvironmental factor that cannot be described directly from the data collected.

173 A trend is also observable in the data from the prodelta mudstone samples. The
174 samples from all prodelta settings trend with an inverse relationship relative to the
175 environmental vectors. This implies that neither grain size nor grain sorting controls
176 on the sediment-palynomorph relationship. Fine-grained sediment was delivered to
177 the prodelta environments of the Lajas Formation either by suspension settling from
178 buoyant plumes, or from hyperpycnal slurries flowing down the delta slope.

179 The other facies described from the Sierra de Chacaico generally follow trends
180 relating to the grain size proxy vector. The mudstones from mouth bar deposits show
181 the strongest positive correlation with the grain size vector, and have seemingly little
182 relationship to the grain-sorting vector, which could be considered an inverse
183 relationship. Mouthbars as a whole are typically composed of well-sorted sand-grade
184 sediment, though the processes responsible for depositing mud-rich drapes are less

185 well constrained. Since there is little relationship between palynology and grain
186 sorting, it may be that different processes operate at different portions of mouthbars.
187 The mudstones sampled from the base of these mouthbar deposits, for example may
188 be more prone to suspension settling than those on the accreting foresets of the
189 mouthbar, and as such the sampling may need to be refined to determine real trends.

190 The tidal flat mudstone samples are not tightly grouped as a whole, but generally
191 share a relationship with the grain size trend. The strongly progradational tidal flat
192 environments of the Lajas Formation were extensive in aerial extent widespread and
193 concomitantly diverse in character (McIlroy *et al.* 2005). Mudstones are typically
194 present either as slack-water drapes (McIlroy *et al.* 2005) or as widespread fluid mud
195 events (Ichaso & Dalrymple 2009). Tidal flats encompass a diverse range of
196 depositional environments with subtle differences in depositional mechanisms (e.g.
197 variable degrees of river, wave and tide influence). The absence of a distinctive
198 relationship between the palynology and grain size variations is therefore perhaps to
199 be expected.

200 Lagoonal mudstone samples from the Lajas Formation are weakly correlated with the
201 grain size vector, but with some possible relationship to the grain-sorting vector since
202 the entire samples plot between the two vectors (Fig. 3). Fine-grained sediment is
203 inferred to have been transported to, partially restricted, coastal depositional
204 environments associated with the Lajas Formation deltas by a combination of
205 suspension settling and lateral transportation from hyperpycnal mud flows.

206 Abandoned tidal channel deposits have a complex relationship that is not simply
207 described by the two granulometric parameters. There appears to be a greater
208 concentration of data along the grain size vector, but there are definite outliers
209 plotting with a relationship towards the grain-sorting vector. Mud-rich abandoned
210 tidal channel environments typically fall along a continuum that grades either:
211 laterally into tidal flats (once the fluvial influence of the tidal channel has been
212 removed by channel abandonment); or distally into the basinal waters. After
213 abandonment, tidal channels gradually fill with a combination of sediment from the
214 adjacent tidal flats, and pelagic material from the marine system (McIlroy 2007).

215 The palynomorph data, when presented on a cross plot (Fig. 4), can show trends that
216 can be used to compare and relate to the facies. The most obvious relationship

217 between palynomorphs and sedimentological proxies is that *Classopollis* spp. plots
218 with an inverse relationship to both palaeoenvironmental vectors. This trend
219 corresponds to the same position as the prodelta mudstone samples (Fig. 3). The
220 prodelta samples also have the highest dominance of *Classopollis* of all facies. We
221 consider that this is due to the buoyancy of *Classopollis* pollen grains (Carvalho *et al.*
222 2006; Traverse 2008) that allows them to bypass the higher energy shallow marine
223 settings before eventually settling, often as intact tetrads, in the relatively quiescent
224 prodelta (Fig. 5). Therefore, transportation and deposition can be demonstrated to
225 systematically control the distribution of majority of palynomorphs, with the
226 exception of the *Classopollis*, which dominates prodelta mudstones.

227 *Practical applications*

228 | When minimal sample material is available, such as from drill cuttings, thise
229 | integration of palynology and grain-size analysis provides a semi-automated
230 methodology for determining depositional facies from a suite of samples within
231 complex marginal marine settings. Given that facies by their very nature grade into
232 one another, there is bound to be overlap in the statistically derived facies groupings
233 (Fig. 6). Where there is overlap, it may be that depositional facies/palaeoenvironment
234 may be better determined by detailed palynological/palaeoecological characterization
235 (cf. (Stukins *et al.* 2013).

236

237 **Taphonomy and depositional processes on tidal flats**

238 The CCA technique was performed on the subset of data from the Norwegian
239 parasequence at Rhea Gorge (Brandsæter *et al.* 2005), to characterize potential trends
240 in a tidal flat dominated succession. The resulting cross plot (Fig. 7) shows a distinct
241 correlation between the tidal flat deposit samples and the grain size proxy vector with
242 the samples forming a tight cluster. To investigate whether a model could be created
243 that links depositional environment with palynological assemblage, the palynomorph
244 data was plotted against the environmental variables (Fig. 8).

245 Two groupings are identified in the tidal flat dataset that relate closely to the
246 environmental vectors. *Inaperturopollenites hiatus* and *Deltoidospora* spp. are found
247 to closely relate to grain sorting (m/m ratio) and Araucariaceae pollen, *Vitreisporites*

248 *pallidus* and *Classopollis* spp. are in close relation to grain size (%<50). No other
249 taxa have a significant relationship to either of the two vectors.

250 When the abundance of the two groups of palynomorphs are plotted against each
251 other (Fig. 9), and the samples are sorted by their inferred depositional
252 environments—based on field characterization of facies—there is a strong trend
253 within the tidal flat samples. In samples from tidal flat facies there is a higher
254 abundance of the grain size related palynomorph taxa, and lower abundance of grain-
255 sorting related taxa. The pattern shown in this assemblage is considered to reflect the
256 sedimentation processes that deposit fine-grained sediment in tidal flat environments.
257 The grain size related taxa both have good buoyancy since Araucariaceae pollen have
258 a large air sac, and *Classopollis* have buoyant pollen morphologies (Carvalho *et al.*
259 2006; Traverse 2008). In contrast, the grain sorting group of taxa are dominated by
260 *Deltoidospora* spores, that are simple trilete spores with little to no airborne or
261 hydrodynamic transportation potential (Salter *et al.* 2002). Such spores are
262 considered to rely almost entirely on fluvial processes to act as transport mechanisms
263 into the shallow marine environment. Due to the minimal influence of distributary
264 channel systems and the dominance of tidal reworking as the primary sedimentary
265 process in tidal flat deposits, the grain sorting related taxa are represented in a far
266 smaller proportion than the grain size related taxa.

267 **Conclusions**

268 Our exploration of the relationship between palynomorph assemblage and grain size
269 of mudstones in the Lajas Formation shows that different facies, when plotted with
270 grain size data proxies for grain sorting (m/m ratio) and grain size (%<50), have
271 consistent spatial relationships (Fig. 6). This insight could be used to allow the
272 palynology and grain-size data from samples, such as drill cuttings of unknown
273 facies, to aid in palaeoenvironmental interpretation using the canonical
274 correspondence analysis method.

275 A link between the palynomorphs present in fine-grained samples from tidal flat
276 dominated successions has been established through canonical correspondence
277 analysis of the Norwegian parasequence data from Rhea Gorge. There is a greater
278 proportion of palynomorphs related to the grain size proxy (%<50) than grain sorting
279 (m/m ratio). This indicates that the process of deposition within tidal flat

280 environments is dominated by fallout from suspension, inferred from the
281 palynomorph physiology, rather than current controlled processes that could be
282 expected from other environments such as channels and mouthbars. This
283 methodology could be directly translated for use in any tidal flat dominated
284 subsurface reservoir or outcrop. The results of this study also demonstrate a link
285 between depositional facies, grain-size distribution, hydrodynamics, and assemblage
286 taphonomy of palynomorphs. This knowledge can be transferred into a semi-
287 automated statistical facies prediction technique for the subsurface in complex
288 marginal marine settings. With further investigation, this technique also has the
289 potential to be applied to a broader range of depositional settings.

290 **Acknowledgements**

291 SS would like to acknowledge a Statoil (UK) funded studentship at the University of
292 Aberdeen

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398

399 Figure 1a: Location diagram of study area. A shows the location of map B in the
400 context of South America. B shows the Neuquén Province and map C. C is a map
401 south of Zapala showing the Sierra de Chacaicó study area.

402 Figure 1b: Location of sampled sections within the Sierra de Chacaico outcrops.

403

404 Figure 2: Simplified stratigraphic column of the Cuyo Group (Based on Howell *et al.*
405 (2005), McIlroy *et al.* (2005) and dates from Gradstein *et al.* (2004)). Lajas
406 Formation sequences and parasequences after McIlroy *et al.* (2005).

407

408 Figure 3: CCA cross plot of eigenvalue scores for the samples calculated from
409 palynological and granulometric data from all Sierra de Chacaico sections. The
410 samples are coded by facies interpretation.

411

412 Figure 4: CCA cross plot of eigenvalue scores calculated from data from all Sierra de
413 Chacaico sections showing the palynomorph variables.

414

415 Figure 5: *Classopollis* spp. tetrad from the “Peregrina” parasequence set (Sample 4o9
416 Leica DMRX coordinates 42.8 103.0)

417

418 Figure 6: Model of the relationship between lithofacies/palaeoenvironment and
419 palynological and granulometric data using CCA.

420

421 Figure 7: CCA plot of eigenvalue scores from the Norwegian parasequence samples
422 and relationship with palynological and granulometric data. Samples are facies coded.

423

424 Figure 8: CCA cross plot of eigenvalue scores of the palynomorphs variables from
425 the Norwegian parasequence and their relationship to the granulometric parameters.
426 Circled are the two taxa groupings associated with the environmental vectors (m/m
427 ratio and %<50).

428

429 Figure 9: Percentage abundance of the grain size (%<50 taxa) and grain sorting (m/m
430 ratio) taxa per sample, grouped in lithofacies/palaeoenvironment.