

*Vulnerable Islands: Climate Change,
Tectonic Change, and Changing
Livelihoods in the Western Pacific*

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The people of Tuvalu and Kiribati in the north Pacific, Takuu (Mortlock Islands) and the Carteret Islands in Papua New Guinea, and Australia's Torres Strait Islanders, are now voyaging towards an uncertain future. Rapidly rising sea levels and massive king tides are encroaching on their villages and salt is affecting arable land. The mass migration of entire island communities is imminent.

—SUSAN COCHRANE,
“FLOATING LAND—RISING SEA”

There has been reluctance to leave, especially among older [Carteret] islanders, but after fighting a losing battle against the ocean for more than 20 years (building sea walls and planting mangroves), it appears the islanders have given up hope, resigned to be among the world's first “climate change refugees.”

—JOHANNES LUETZ, *Planet Prepare*

For more than two decades, small Pacific islands, and especially atolls, have been argued to be in the vanguard of climate change and sea-level rise—the “canaries in the coal mine.” The most anticipated physical impacts of sea-level rise on islands are coastal erosion, flooding, and salinity intrusion, reducing the resilience of coastal and small island ecosystems. Atolls (and atoll states) are particularly vulnerable to climate change (and to cyclones), having long sinuous coasts and small land areas (Nunn and Kumar 2006). The populations of several islands and island groups—notably the island state of Tuvalu, the Carteret Islands (Bougainville,

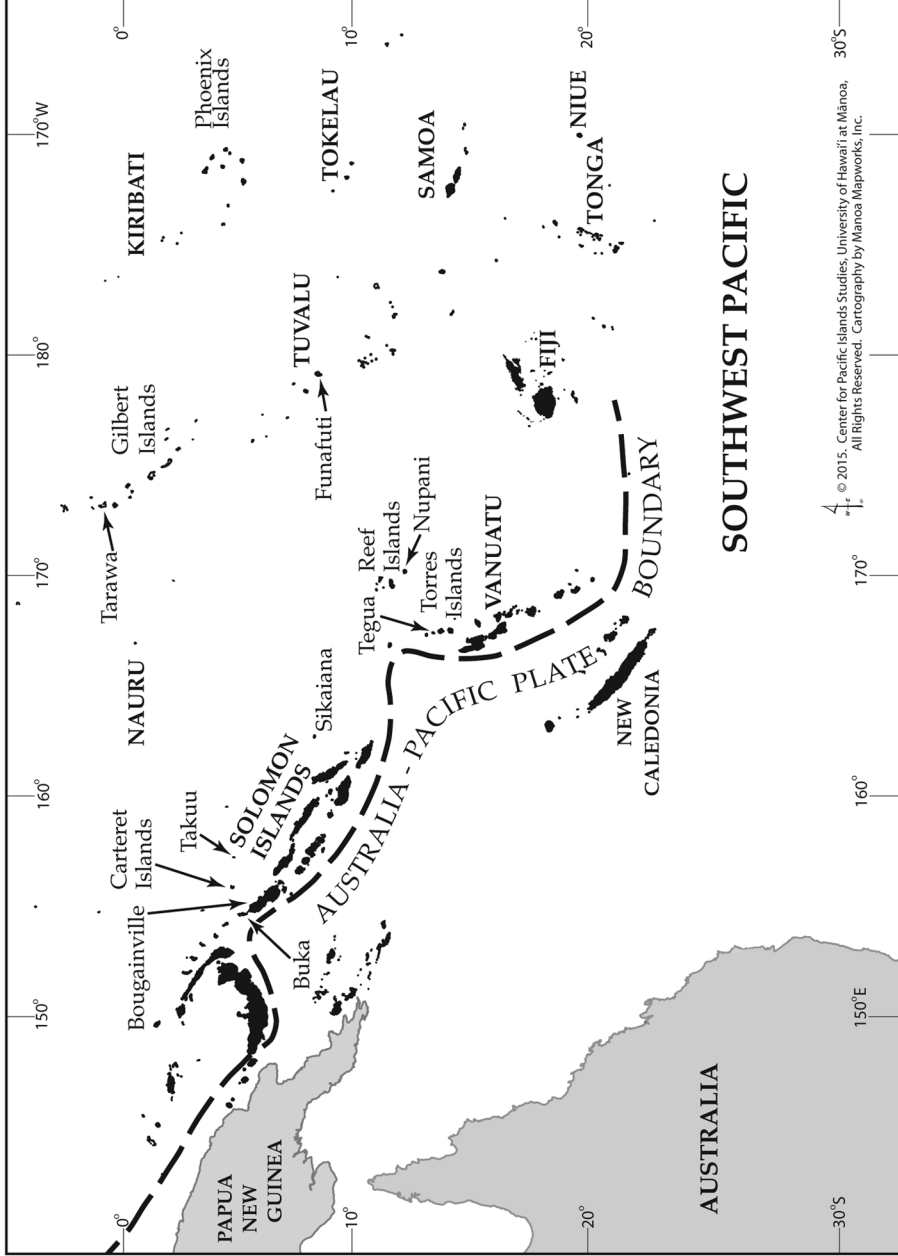
Papua New Guinea [PNG]), and Tegua (Torba, Vanuatu)—have been perceived as the first probable victims of rising seas, so that their inhabitants would become, and in some quarters already are seen to be, the first “environmental” or “climate change refugees” (see, eg, Cochrane 2010; Luetz 2008). The imminent “fate” of Tuvalu has been widely documented alongside that of other atoll states, especially Kiribati, the Maldives (in the Indian Ocean), and the Marshall Islands (Connell 1999; Farbotko 2005, 2010; Rudiak-Gould 2013a), to the extent that “Tuvalu’s iconic role as ‘poster child’ for encroaching global disaster is well established” (Chambers and Chambers 2007, 294; see also Goldsmith 2012). The smaller, less accessible Carteret Islands (Kilinailau), Takuu (Bougainville, PNG), and Tegua have largely escaped comparable detailed media coverage but have still been widely seen as early examples of the impact of sea-level rise on small islands, and thus necessitating and experiencing population resettlement (Connell 1990, forthcoming; Bohane 2005). Carteret Islanders have typically been described as “the first direct climate change refugees with islands inundated and damaged, gardens and water supplies destroyed by salt water intrusion and evacuation announced in 2005” (Roper 2009, 5). By 2002, the future of nearby Takuu was said to be “increasingly uncertain because of the rising sea levels that continue to erode the shorelines of all its islands . . . and abscond with meters of beach frontage” (Moyle 2007, 272–273). The island of Tegua “is claimed to be the first in the world to have to move its community because of rising sea levels” (Bohane 2005), and the Tegua Islanders were thus said to be “the first climatic refugees” (Siméoni and Ballu 2012). Small islands off the west coast of Buka (Bougainville, PNG), such as Petats, have experienced similar changes, with Islanders having “begun to move away from the sea to higher ground” because of encroaching seas (Luetz 2008, 18). Multiple regional claims have been made to the primacy of climate-change refugees.

Especially but not only in atoll environments, local people and others have pointed to various changes such as erosion, flooding, and contamination of agricultural systems by seawater intrusions that have been widely attributed to global warming and sea-level rise (SLR) (see, eg, Connell 2003; Rudiak-Gould 2009, 2012; Kuruppu and Liverman 2011). Massive public discussion of climate change at international, regional, and local levels has contributed to SLR being the primary explanation for harmful, unusual, or unprecedented environmental changes. Media reports have frequently attributed major environmental problems to climate change and SLR (Connell 2003; Farbotko 2005), notably in Al Gore’s documen-

tary *An Inconvenient Truth* (Guggenheim 2006), in which Tuvaluans were said to have had “to evacuate to New Zealand.” Some academic reports are not averse to similar dramatization, such as SLR having resulted in “the evacuation of some islands in Tuvalu” (Armstrong and Read 2004, 196). Such statements are far from atypical and have resulted in perceptions, first, of environmental changes in small Pacific islands being almost synonymous with SLR and, second, of the necessity (or actuality) of migration as a response. Detaching science and social science from politics, polemics, aspirations, and needs has proved difficult. This article seeks to examine the extent to which climate change has contributed to environmental change and degradation in small island contexts, especially coral atolls, and the extent to which other factors have influenced both environmental change and perceptions of these changes. It is argued that contemporary environmental changes in most islands in the southwest Pacific (figure 1) are not necessarily the outcome of SLR but are more likely to be responses to local stresses of various kinds, from geophysical to human, including tectonic movements, cyclones, seawall construction, and sand mining. This is examined below with particular reference to relationships with contemporary Island livelihoods.

SEA-LEVEL RISE

Sea-level rise has become increasingly evident in the Pacific, with the southwestern Pacific regarded as one of the regions most vulnerable to contemporary and future changes (Nicholls and Cazenave 2010; Becker et al 2012; Meyssignac and Cazenave 2012). Recent estimates of SLR by the Intergovernmental Panel on Climate Change (IPCC) have indicated a global average of 3.1 mm/year between 1993 and 2009, representing an increase from 1.7 mm/year between 1950 and 2009, with the eventual outcome of an SLR of between 0.18 and 0.59 meters by the end of the twenty-first century. Considerable regional and local variations occur, but relatively few measures (either in time or space) exist for most Island regions. However, contemporary records for the central Pacific suggest that SLRs of around the IPCC estimates are evident (Becker et al 2012). One study has suggested a rising trend for some Pacific Island states of 3.1 mm/year, with Nauru, the Federated States of Micronesia (FSM), Fiji, Tonga, and Kiribati exhibiting even more rapid rises and thus exceeding earlier IPCC predictions (Becker et al 2012). A further study has pointed to rates as high as 7 mm/year since 1993 for both the Marshall Islands



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MAP I.

and Papua New Guinea and 5 mm/year and 1–4 mm/year for Tuvalu and Kiribati, respectively (ABM and CSIRO 2011). Trends and estimates have increased over time.

Specific data exist for some sites. Thus measures of SLR on Abaiang Atoll, Kiribati, using geomorphological techniques, suggest an average rate of SLR ranging from 0.1 to 0.2 mm/year between 1998 and 2008 (Flora, Ely, and Flora 2009) and on Tarawa, Kiribati, of 1.4 mm/year between 1968 and 1998, or 1.8 mm/year over the thirty-five years since 1974 (Biribo and Woodroffe 2013). Other measurements, however, suggest over 4.3 mm/year between 2002 and 2009 (Rankey 2011). Lautoka (Fiji) has exhibited an SLR of 4.9 mm over a twenty-year period. Recent floods in Nadi (Fiji) suggest that SLR is now having a negative impact (Nunn 2013). Funafuti (Tuvalu) has recently recorded a particularly high rate of 5.1 mm/year, about three times that of the global mean, whereas in the eastern Pacific rates of between 2.5 and 2.9 mm/year predominate (Chowdhury, Chu, et al 2010; Becker et al 2012). All of these rates are faster than earlier rates, putting increased pressure on Funafuti. Specific measures of SLR are unavailable for any of the Bougainville atolls, Solomon Islands, or most of Vanuatu.

Data exist only for short time periods in several places and are also complicated by El Niño–Southern Oscillation (ENSO) fluctuations, since during La Niña conditions SLR has increased (Hussein, Singh, and Aung 2011; Aung, Singh, and Prasad 2009; Aung, Singh, and Maryam 2011; Chowdhury, Barnston, et al 2010; Becker et al 2012). Isostatic changes to earth and ocean surfaces further complicate measurement (Woodroffe et al 2012). Nevertheless, it is reasonable to assume that the rate of SLR is increasing, that it will increase further in the future (Church and White 2006; Horton et al 2014), and that the western and southwestern Pacific are experiencing more accelerated SLR than the eastern Pacific (Merrifield 2011; Zhang and Church 2012). Regional variations, significant discrepancies, and some uncertainty remain in measures of SLR, according to sources and analytical techniques, but a distinct overall upward trend exists throughout the western Pacific.

TECTONIC INSTABILITY

In the western Pacific, SLR is complicated and compounded by tectonic instability (including both uplift and subsidence) and earthquakes. Thus

the Torres Islands (Vanuatu), at the convergence of the Australian and Pacific plates, have experienced particularly strong vertical movements while also being earthquake and cyclone prone. Between 1992 and 2010, SLR around the Torres Islands was nearly 8 mm/year, significantly above the global and regional average. That exceptional SLR was complicated by movements during and between earthquakes. Between 1997 and 2009, the outcome of vertical plate movements was that the island of Tegua subsided by about 117 mm—one of the highest recorded subsidence rates in the world, effectively almost doubling more gradual SLR (Ballu et al 2011). However, subduction rates vary considerably, even from north to south within Vanuatu, reflecting both tectonic segmentation and the transience of rates measured at particular stages of earthquake cycles (Dickinson 2013).

Substantial flooding in the Torres Islands, starting in the late 1990s, was probably primarily the result of a 1997 earthquake (followed by a small tsunami) that accompanied subsidence, while a 2009 earthquake reversed that subsidence, lifting the island by about 200 mm. The 1997 earthquake resulted in subsidence of at least 500 mm and was probably the main reason for subsequent sea incursions and flooding through progressive erosion of natural barriers. Despite SLR being relatively rapid, the principal cause of observed flooding in Tegua was rapid subsidence, related to plate movements and earthquakes (Ballu et al 2011; Siméoni and Ballu 2012). Between 2004 and 2006, flooding resulted in the relocation inland of the village of Lataw—sometimes described as the first migratory response to global warming in the Pacific (Bohane 2005).

Similar explanations account for parallel changes in Vanikoro (Solomon Islands) (Traufetter 2012), and probably elsewhere in Solomon Islands, both earlier in the Reef Islands—where there had been subsidence “during the last thirty years” on all four of the northern outer islands (Davenport 1975, 68; Birk 2012)—and more recently in the Bougainville islands and atolls, including the Carteret Islands (Duguman 2009), Takuu (Mortlock Islands), the Buka islands, and the Duke of York Islands, all close to the plate boundary. All these islands regularly experience strong earthquakes, which can significantly change sea levels. A single earthquake in Ranongga (Solomon Islands) in 2007 lifted parts of the island by over three meters (McAdoo et al 2008), but subsidence, as in nearby Vella Lavella and Simbo, which subsided almost one meter (Lauer et al 2013), is an alternative outcome.

EL NIÑO SOUTHERN OSCILLATION

Globally, three large-scale climate-related changes are simultaneously occurring: a rise in air- and sea-surface temperatures (projected to increase by 2°C by 2050); changes in rainfall (with expectations of less frequent but more intense rainfall events, potentially contributing to greater storminess and flooding); and consequent higher sea levels, but with significant regional variations (Lal, Harasawa, and Takahashi 2002; Mimura et al 2007). Each varies regionally in accordance with particular local changes, notably in the El Niño–Southern Oscillation (ENSO), a quasi-periodic climatic pattern that occurs across the Pacific roughly every five years, that changes sea-surface temperatures (warming with El Niño and cooling with La Niña) and air-surface pressures. At the peaks of these cycles, more extreme weather occurs. SLR is complicated further by ENSO events, being affected by large areas of water warming and expanding, or cooling and contracting. Current SLR is a fraction of the sea-level changes that occur as short-term impacts (unusually high tides and storm surges) from ENSO-related fluctuations.

ENSO fluctuations have occurred for millennia, with a considerable range of outcomes (McGregor et al 2013), and in some places they are linked to local agricultural and ceremonial cycles (Mondragón 2004). In recent decades, the number and intensity of El Niño warming events has increased and the number of La Niña events decreased—a change that may be attributable to global warming. ENSO fluctuations have been the most important influence on climate variability in recent decades, but short-term climatic shifts must be differentiated from long-term changes. Particularly high sea levels are recorded during La Niña events, for example, from July 2006 to June 2008 in most of the Pacific (Chowdhury, Barnston, et al 2010; Siméoni and Ballu 2012). Various islands, especially atolls, have experienced damage—notably erosion and flooding—caused by both perigean (astronomical) high tides and storm surges associated with particular El Niño conditions. However, fluctuations of island outlines correspond with wind changes associated with ENSO cycles (Solomon and Forbes 1999; Biribo and Woodroffe 2013). Climatic changes, at different scales, are unusually complex, while relatively few accurate measures exist to document the extent and impact of change, both directly and in the outcome of such changes in terms of water availability. How-

ever, ENSO events have contributed to rainfall patterns becoming more unpredictable, with floods and droughts being accompanying outcomes, while future wet and dry extremes—and hence floods and droughts—are expected to be more intense (Power et al 2013).

Distant-source wind waves, not associated with commonly identified sources such as cyclones or perigean tides, have also been an occasional, and increasingly less rare (Aucan, Hoeke, and Merrifield 2012), source of inundation of low-lying areas, as occurred in several Pacific Island states in December 2008. Swell generated more than four thousand kilometers from the farthest affected island created considerable damage at a time when regional sea level was already elevated, due to both La Niña conditions and SLR (Hoeke et al 2013; Rudiak-Gould 2013a). Atolls from Midway and Wake in the northern Pacific to Takuu were affected. The impact would have been greater had the swells arrived during a peak of spring tides or in stronger La Niña conditions. Flooding in the Marshall Islands in 2013 may have had similar causes.

CYCLONES

Cyclones are frequent in much of the Pacific, other than extremely close to the equator, although cyclonic activity has become more common there (Rasmussen et al 2009). Climate change is expected to result in fewer cyclones but greater individual intensity. Time series data are too limited to give more than localized indications of the extent of any increased periodicity and severity, and in the Pacific region no significant trends in either the number of cyclones or their intensity were evident from 1981 to 2007 (ABM and CSIRO 2011). There have been indications that cyclones are becoming more intense (Webster et al 2005; Elsner, Kossin, and Jagger 2008; Rasmussen et al 2009; Walsh, McInnes, and McBride 2012), perhaps because of the increased intensity of ENSO events. However, the possibility and impact of changing patterns and intensity are uncertain, and, in the Pacific, “owing to their small land masses and widely dispersed nature within a vast expanse of ocean, whether or not individual . . . islands sustain significant damage by the passage of tropical cyclones would seem to be governed largely by chance” (Terry and Etienne 2010, 193). Cyclone frequency and impact are thus highly variable. Tuvalu was hit by an average of three cyclones per decade between the 1940s and 1970s, but eight occurred in the 1980s (Nunn 1990), and several have occurred since then but with varied intensity in different parts of the country.

Cyclones cause environmental changes. Storms may create waves that wash over atolls and contaminate freshwater lenses. A cyclone that swept over Pukapuka (Cook Islands) in 2005 reduced potable lens water supply to brackish water that was unsuitable for drinking for ten months; rainwater tanks and wells were so polluted that freshwater had to be imported, while fresh food supplies were wiped out by incursions of seawater into taro gardens (Terry and Falkland 2010). Similar devastation occurred in the Chuuk atolls (FSM) in 2007 (Hezel 2009), and other, less extensively documented circumstances have affected Funafuti and elsewhere. Yet the storm waves associated with cyclones have positive and negative effects. No other natural process can produce coral debris and transport it from where it is produced by the growth of corals along ocean reef fronts to where it can accumulate in more or less stable landforms. Cyclones are typically destructive in the short term, in relation to property, crops, and freshwater resources, but they can also be agents of island geomorphic enhancement.

ISLAND SHAPES AND SIZES

A number of physical influences—long-term SLR, tectonic shifts, ENSO events (resulting in storm surges), cyclones, and distant swells—all contribute to environmental changes on islands. In most contexts, the least significant of these is SLR, while the principal physical influences are recurrent events of various kinds that may not have obviously changed in periodicity in the past century. With the exception of tectonic changes (and especially subsidence), in most circumstances their short-term impacts have been substantial but their long-term influence has usually been slight.

Erosion of islands can lead to a reduction in the size (and water quality) of the freshwater aquifer of the island (Roy and Connell 1991), as on Majuro (Marshall Islands) (Terry and Thaman 2008), or its complete disappearance from smaller islands. Even where islands demonstrate physical resilience, they may be rendered uninhabitable by the diminution of available water resources (Terry and Chui 2012). However, storms have both constructional and erosional impacts on atolls, depending on the sediment composition of the islands and the orientation of the storm (see, eg, Bayliss-Smith 1988; Yates et al 2013). Climate change and SLR are widely anticipated to bring increased rates of erosion and shoreline regression, especially on coral atolls and cays. Erosion has occurred in multiple Pacific island contexts, but there is little evidence of its being consis-

tently in excess of deposition and accretion in most coastal systems where human impacts are insignificant. By contrast, where human impacts have been considerable, erosion is more likely, especially where coastal vegetation has been removed; hence, erosion occurs particularly in more densely populated urbanized coastal areas (see, eg, Solomon 1994; Pierce 1999). Erosion is more readily visible there and thus more frequently reported. Deposition may occur elsewhere.

The impact of SLR is greatest for low-lying atoll islands, which have been widely perceived to erode in response, despite accretion usually occurring simultaneously (Woodroffe 2008). Quantitative studies of physical changes in twenty-seven atoll islands in three central Pacific states (FSM, Kiribati, and Tuvalu) over a variable 19–61-year period (when SLR was measured at 2 mm/year) show that 86 percent of atoll islands either remained the same size (43 percent) or increased in area (43 percent). Only 14 percent of the islands exhibited a net reduction in area. Despite small overall changes in area, the islands exhibited larger gross changes, expressed as changes in the planform configuration and the position of islands on reef platforms. Many islands therefore increased in area and underwent diverse physical adjustments in response to changing influences, of which SLR was just one. On Funafuti, eleven islands increased in size and five decreased, while four Tarawa islands increased, primarily because of extensive reclamation (Webb and Kench 2010; Biribo and Woodroffe 2013). In French Polynesia, the atoll of Manuae has decreased in size over a fifty-year period, while Manihi has increased in size, and two other atolls remained more or less stable, with the differences attributed to variability in wave exposure rather than to SLR (Yates et al 2013; Le Cozannet et al 2013). On Wotje (Marshall Islands), over a sixty-seven-year period until 2010, accretion was more prevalent than erosion, but there was a post-2004 shift toward erosion (Ford 2013). Similarly, three out of four Reef Islands (Solomon Islands) were smaller in 1960 than in 2010 (Birk 2012). On Raine Island, a vegetated coral cay in the northern Great Barrier Reef, 34 percent of the coastline experienced net retreat and 44 percent experienced net progradation between 1967 and 2007. Despite local concerns that Raine Island was rapidly eroding, net island growth (6 percent by area and 4 percent by volume) had occurred. Perceptions of erosion probably reflected large morphological changes arising from seasonal, inter-annual, and inter-decadal patterns of sediment redistribution rather than any net sediment loss (Dawson and Smithers 2010). One

Tree Island in the Great Barrier Reef has similarly been stable, with constructional processes on the rubble flats (Shannon et al 2012). Overall, therefore, atoll islands in the Pacific have grown rather than declined in recent decades.

Local changes on atolls are complex. On Tarawa, between 1968 and 1998, when the urbanized central atoll of Kiribati experienced what appears to have been a net SLR of 1.4 mm/year, shoreline responses varied. Most ocean-facing beaches showed little detectable change, with some showing net erosion and a few showing accretion, but a significant proportion of lagoon beaches also showed accretion. The factors that most influenced change were location and human influence, while substantial fluctuations also occurred in response to El Niño conditions (Biribo and Woodroffe 2013). The same conjuncture of both accretion and erosion was evident on Maiana and Aranuka atolls (Kiribati) between 2005 and 2009. The greatest erosion was on ocean-facing coasts (where no one lived), where storm and wind impacts were considerable, or on the edges of lagoons near villages that had been most influenced by human activity, notably by the construction of groins and seawalls but also of causeways and boat channels (Gillie 1993; Rankey 2011; Woodroffe and Biribo 2011). In each of these cases—and in Majuro, where a thirty-five-year record showed that erosion has been greatest in the most urbanized and most modified parts of the atoll (Xue 2001; Ford 2012; Rudiak-Gould 2013a), and by the 1960s on Nupani (Reef Islands, Solomon Islands) (Davenport 1975, 69)—erosion was greatest where human activity was substantial (the construction of causeways, seawalls, vegetation removal, sand and gravel mining, etc). On Abaiang Atoll (Kiribati), areas of the lagoon coast had receded by up to eighty meters in four decades after a channel had been closed, blocking sand transport (Webb 2006). The largely urban concentration of erosion partly explains the widespread media assumptions of generalized erosion following SLR, which have come predominantly from urbanized atolls. Coral islands have otherwise exhibited more resilience and stability than usually assumed under ongoing SLR. Erosion of particular island shorelines takes place in the context of physical adjustments to entire island shorelines, as erosion is more or less balanced by deposition and progradation elsewhere. However, the land deposited is usually coral fragments, as in the Reef Islands (Davenport 1972), and the biological content of the eroded material (notably its soil content) is not replaced by deposition of material of comparable value.

HUMAN INTERVENTIONS

Physical changes to atoll topography have both been more evident where human impact has been greatest and most likely to have been observed where population concentrations are greatest (and media interest concentrated). The harmful impacts of SLR have been most frequently reported from Funafuti, the highly urbanized main atoll of Tuvalu, where flooding of the interior of the main island (Fongafale) and the disappearance of the tiny islet of Te Puka Savilivili are widely cited as evidence of the contemporary impact of SLR. The islet's disappearance has been "held up as evidence of the effects of climate change: encroaching tides have erased sandbanks and coconut trees, which die and collapse into the sea when soil becomes too salty" (Williams 2001, 28; quoted in Connell 2003, 100). However, at the end of the nineteenth century, when the geologist Edgeworth David was examining reef changes, the islet had just three coconut trees, compared with thirteen in 1976 (Roger McLean, pers comm, 2002), suggesting that it had gone through cyclical changes and was then growing. In Tarawa, the disappearance of the similar tiny island of Bikeman in Tarawa Lagoon has been likewise incorrectly attributed to SLR (Howorth 2000). Again, the most dramatic changes have been reported for the most urbanized atolls. Much the same appears to be true for Kiribati and the Marshall Islands (Rudiak-Gould 2012). However, while attribution of environmental damage to global warming and SLR remains primarily an urban phenomenon, coastal erosion and flooding have become more widely apparent.

Each of the key urbanized islands in the three atoll states of Tuvalu, Marshall Islands, and Kiribati—Funafuti, Majuro, and Tarawa—has experienced very substantial post-independence rural-urban migration, resulting in extremely high urban population densities and intense pressure on scarce resources including land, firewood, sand, and gravel. Their human landscapes have changed much more substantially than those of outer islands. Contemporary flooding in Funafuti has been exacerbated by the construction of a permanent airstrip (and its subsequent sealing), which occupies as much as a quarter of the main island; borrow pits (dug during the war for gravel); sand and gravel mining; vegetation removal (mainly for firewood); depletion of the water lens (and resultant compaction); groin and causeway construction; "one-off" reclamation; lagoon eutrophication and the loss of sediment for reef construction; and other

human-induced changes (Connell 2003; Xue 2005; Woodroffe 2008; Lazrus 2009). Floods in 2006 occurred in parts of the main island lying below high-water level, which had previously been wetlands and were only populated in the 1970s during a phase of rapid migration to Funafuti; flooding may also have been affected by excavation of a storm ridge (Yamano et al 2007). In Tarawa and Majuro, erosion and flooding have similarly followed modification of shorelines by reclamation, housing, causeways, and sand and gravel mining, with flooding in Majuro occurring only in recently populated areas (Donner 2012; Ford 2012). Sand and gravel mining in lagoons has been particularly damaging, scouring lagoon floors and contributing to coastal erosion, again notably in Funafuti, Tarawa, and Majuro, where demand is greatest (Ambroz 2009; Smith and Collen 2004; McKenzie, Woodruff, and McClennen 2006; Babinard et al 2014). Even on non-urban atolls such as Vaitupu (Tuvalu), gravel mining has been the primary cause of erosion (Xue 1997). Fewer such changes have occurred on outer islands, where populations have usually declined rather than increased in the post-independence era, hence pressures on the environment have also not increased.

In many atolls, including Woleai and Moch (FSM), Rongelap (Marshall Islands), and elsewhere, Islanders are conscious that erosion is part of a “normal process”—more of the same—and had long taken steps to modify and defend eroded areas and secure land elsewhere, at least as early as the 1970s (Bridges and McClatchey 2009; Lipset 2011; Pam and Henry 2012; Schneider 2012). On several Kiribati atolls, many erosion sites had experienced “one or two generations of seawalls” by the 1990s (Gillie 1993). On Petats, people were aware of “rising water” in the 1980s when they began to build seawalls (Luetz 2008, 18). In Tegua, older Islanders recounted how the area that had flooded in the 2000s had once been a lagoon before reclamation (Siméoni and Ballu 2012). Similarly, parts of Pororan (Buka) have been reclaimed while, as one Islander pointed out, “on this side, the island is growing, but at other spots, the sea is coming inside” (Schneider 2012, 11; see also Schneider 2012, 189). The uninhabited interior of Funafuti was subject to flooding in the 1890s (Yamano et al 2007). In the Reef Islands, one group of older men suggested that not only was the contemporary erosion rate not unusual but that erosion was often balanced by deposition, with one man claiming that there had been no SLR but that environmental problems were self-inflicted: “We cut trees; we take sand from the beach” (quoted in Birk and Rasmussen 2014, 9). Cook Islands atoll residents ponder beach erosion and query, “Is this

climate change—or is it our own work—or both?” (Rubow 2013, 65). In very many islands, there is inherited knowledge of dangers from the ocean. In Tegua, villages and gardens were formerly located on a one-hundred-meter platform to avoid tsunamis; villagers were only encouraged to migrate to the coast in colonial times, but the main coastal villages were not established until at least the 1930s (Ballu et al 2011). Thus in Funafuti, Tegua, Tarawa, and also Majuro (Spennemann 2006), Islanders have only very recently moved to sites that were avoided in the past and have subsequently proved to be cyclone and flood prone.

It is not surprising that some slow-onset changes such as climate change and SLR have become political issues, because human influences on them occur far from the island Pacific and are the subject of debate and argument and because perceived physical changes that may have alternative explanations can be attributed to climate change and SLR: a “garbage can anarchy” where once separate and complex phenomena have become systematically interrelated (see, eg, Connell 2003; Corlett 2008; Hulme 2010). Small islands offer sites where the great global narratives of climate change can be comprehended and made tangible and visible, attributed to distant sources, and confirmed by local “green governmentality discourses” based on indigenous knowledge, memory, and eyewitness accounts (see, eg, Farbotko 2010). News media and occasional nongovernmental organization accounts are replete with romanticized statements such as “[Matsungan] Island chief John Kela doesn’t understand the science of climate change. But he sees that the ocean surrounding his island is rising” and “Even as the last veneers of organic matter are pushed out to the ocean, Tobasi [in the largest Carteret island] prays daily for his island. He knows that life on the atoll is coming to an end” (Luetz 2008, 17, 20). Visibilism and populism have triumphed over science (Rudiak-Gould 2013b; Swyngedouw 2010). Single-factor interpretations are limited modes of explanation in the face of evolving scientific knowledge, deny local agency, and invoke powerlessness.

ISLANDS UNDER PRESSURE

Small islands and especially atolls have often experienced difficult times. In precolonial times, islands achieved sustainable development partly through extended geographical ties, typified by dispersed clans, and linkages across atolls and between clusters of islands in order to secure social relations and claim and use land elsewhere, as on such densely populated

islands as Paama (Vanuatu) and Kotu (Tonga) (Perminow 1993) or more generally in Micronesian atolls (Petersen 2009). Linkages were often marked by elaborate exchange systems and complex regional and local reciprocal socio-economic-political linkages to ensure sustainability, such as the kula ring in the Trobriand Islands (PNG) and the hierarchical and asymmetrical sawei obligations that tied atoll outer Islanders to the high island of Yap in what is now the Federated States of Micronesia (Alkire 1978; Pam and Henry 2012). Islands were rarely isolated for long, and even the tiny island of Tikopia (Solomon Islands) was part of a Pacific “world system” for most of its three-thousand-year human history (Kirch 1986). Survival necessitated external ties. Small islands could not afford to be insular.

Population mobility was inherent to small islands. Without migration, life could be particularly difficult. In extreme form, in small, remote high islands such as Tikopia, achieving sustainability had a high human cost, involving celibacy, infanticide, and other extreme practices to stabilize population size and growth rates; even in contemporary times, island chiefs regulated the population to ensure that it did not pass the number assumed to be sustainable (Kirch 1997). On Marshallese atolls, populations were once restricted to two-child families; land was so scarce that only chiefs were buried on land, with commoners buried at sea (Spennemann 2009; Rudiak-Gould 2009). Local and regional contacts sometimes involved feuds, warfare, and violent conflicts over land, resources, and fishing grounds (D’Arcy 2006, 2009). In the more modern era, such practices gave way to dependence on distant governments or migration or both and to a reduction in local sustainability.

By the end of the nineteenth century, in remote atolls like those of Tokelau, “the idea of permanent migration, involving a severance of many ties with the home island and of seeking one’s fortune elsewhere, is well established in Tokelau life and thought. For the past 70 years or so it appears to have been accepted . . . that some of nearly every group of siblings must *take* ‘emigrate’ simply because the local resources are seen as insufficient” (Hooper and Huntsman 1973, 403–404). Drought in 1916 resulted in the resettlement of some Kapingamarangi Atoll residents on Pohnpei (Lieber 1977). More recently, migrants left small eastern Fijian islands where their homes were “beautiful, but not places to live” (Bedford 1980, 57), while boredom drove migrants from Bikini (Marshall Islands) and Sikaiana (Solomon Islands) (Kiste 1974; Donner 2002). The only three Pacific islands depopulated in the twentieth century have been

small atolls (Connell 2013a). In Solomon Islands outlier atolls, migration gradually became a household strategy, by which migrants and their families (including those staying behind) diversify sources of incomes in order to minimize risks, such as loss of income and crop failures (Birk and Rasmussen 2014; Christensen and Gough 2012). That has only intensified as migration has been prolonged, expectations have risen, and aspirations have increased. In many Pacific islands a culture of migration had become established in which migration was anticipated and normative (Donner 2002; Connell 2008). Small islands became increasingly dependent on remittances.

In postcolonial times, the populations of the main atoll states—Kiribati, Tuvalu, and the Marshall Islands—became greater than at any time in the past, and a new phase of internal migration emphasized urbanization. Internal migration accompanied international migration, as Islanders sought wage employment, preferably with the government, and bureaucracies expanded after independence. The population of Funafuti grew five-fold in three decades after independence, and in Tarawa and Majuro growth was just as rapid. Urban management in atolls has proved exceptionally difficult. In Tarawa, “the impacts of unmanaged urbanisation, a continued crisis of inadequate sanitation, a lack of solid waste disposal controls and ineffective freshwater management offer[ed] . . . threats to sustainability” equal to the much-touted issues of climate change and SLR (Storey and Hunter 2010, 167). Urban atolls were placed under exceptional and unprecedented pressure.

Remittances became even more crucial after the decline and virtual collapse of the copra trade in the 1980s and the stagnation of urban employment opportunities, largely ensuring some economic stability on small islands (Chambers and Chambers 2001; Connell 2013a). By then, Kiribati and Tuvalu had been conceptualized as MIRAB states where international migration generated remittances that were the core of the economy (Bertram and Watters 1985). In Nanumea (Tuvalu), as world copra prices slumped, remittances grew from being about half the island income in the 1970s and 1980s to constituting some 75 percent in the 1990s. In such islands, having large families in the hope that one or more children might eventually migrate and provide remittances became an even more conscious economic strategy (Chambers and Chambers 2001). On Eauripik (FSM), while everyone on the island was “useful,” people could become “worthless” if they migrated and sent nothing back (Scourse and Wilkins

2009). Demand for wages and salaries put intense pressure on urban centers even where international migration was also significant.

BEYOND CORAL ATOLLS

As early as the mid-nineteenth century, population pressures on scarce resources on some of the Gilbert Islands were severe enough for the local people to be described in 1865 by one missionary as “genuine Malthusians” (quoted in Munro and Bedford 1980, 3). Although the phrase had no empirical validity, achieving an adequate livelihood was unusually difficult, and warfare was not uncommon. Nukumanu Atoll (PNG) experienced a considerable population reduction in the 1870s following overpopulation (Bayliss-Smith 1975, 312–313). In several atoll contexts, twentieth-century resettlement has occurred in response to diverse human induced changes. The instigator of the first Gilbertese resettlement in the Phoenix Islands (Kiribati) in the 1930s argued that the optimum population density of the Gilbert Islands (the central island chain in Kiribati) had actually been reached by the year 1840 (Maude 1968). Nearly a century later, population densities were much higher, food shortages had occurred, poverty was common, and the colonial administration proceeded with resettlement. Two decades later, in the 1950s, after recurrent droughts, the Phoenix Islands settlers and others from the Gilbert Islands were relocated in the Solomon Islands. Many Vaitupu (Tuvalu) Islanders resettled on Kioa Island in Fiji around the same time. In Kiribati, in the year of independence, the settlers in the Solomon Islands, once pitied, were now envied: “Settling overseas, beyond the ocean of our islands, is something to be sought after. Why? Because our population is still growing. . . . So now many consider them, the resettled ones, as the fortunate ones and they consider us to be the unfortunate ones” (Schutz and Tenten 1979, 127). In the 1980s, similar circumstances resulted in the resettlement of I-Kiribati from the main island chain (especially from drought-prone islands such as Tamana) to the thinly populated Line Islands in the east. More recently, various Tuvaluans have settled on Niue. A further similar movement, explicitly justified as a response to “future problems of overcrowding” and following a severe cyclone, was instigated by New Zealand for the Tokelau atolls and was initially envisaged to include the whole population. Government-assisted migration to New Zealand began in 1963, resulting in a third of the population migrating, with others migrating

freely (Wessen et al 1992; Huntsman and Kalolo 2007). The atoll population has stabilized and slowly declined as the Tokelauan population has grown in New Zealand.

Similar factors, though less well documented, have stimulated migration from the Carteret Islands and Takuu and have been exacerbated by erosion and flooding. At least since the 1960s, the Carteret Islands periodically experienced food shortages of varying severity, associated with steadily growing human and pig populations and the increased conversion of coconuts into copra to generate income. Resettlement to Bougainville was considered as early as the 1960s (Mueller 1972; Connell 1990). Attempts to find appropriate land on nearby Bougainville were thwarted by local landowners, and fewer than ten households have moved (Connell 2012, forthcoming). In the 1950s, when Australian administration officers were similarly concerned about the growing population, overcrowding, and limited garden resources of Takuu, they sent thirty unmarried men to work on plantations elsewhere (Moyle 2007, 25). Such pressures on resources were compounded in 2006 by cyclone damage to the islands, and, in 2008, a series of tidal surges (from wind waves) at the end of a La Niña period washed away kitchens, flooded homes, and destroyed churches. Provincial shipping was unable to resume for several weeks, as hazards emphasized atoll vulnerability and locational disadvantage. However, in 1999, when the provincial government raised the idea of resettlement on Bougainville, it was widely seen as a “fearsome possibility” (Moyle 2007, 16). Both on Takuu and in the Carterets, there was a typical ambivalence between the desire for migration and superior living standards and a preference for holding on to home.

Similar circumstances exist on Nupani. Islanders there “have never been and are not now self-sufficient,” so around the early 1960s they began to use land on much larger Ndende Island, some forty kilometers away, because of intense pressure on land, to the extent that “every square foot of ground suitable for cultivation has long been planted” (Davenport 1969, 173, 174). Reef Islanders were also gradually settling on the neighboring high volcanic island of Tinakula because of land being lost in the 1940s and 1950s (Davenport 1975; Birk 2012). In the 1970s, “due to geological changes over the past two decades the food situation has become more critical. . . . the eastern edge of Nupani’s reef has subsided slightly, probably no more than a few inches [resulting in the erosion of] between one-third and one-half of the islet’s best agricultural land” (Davenport 1972, 15; see also Davenport 1975, 69). Pororan Islanders have likewise

colonized the nearby mainland for agriculture (Schneider 2012). Similar changes have been observed on other islands off the coast of Buka, where four islands (Matsungan, Petats, Torotsian, and Pororan) are said to be shrinking in size, experiencing erosion, and being increasingly affected by more frequent and stronger storm surges and high tides, resulting in the loss of fertile soil and the increased salinity of well water (Luetz 2008). At the turn of the century, the resettlement of Duke of York Islanders was announced since, for their island and “for Takuu, the threat represented by the rising seas is compounded by a violent clash of underwater tectonic plates that triggers periodic earthquakes, tidal waves and volcanic eruptions” (*The Independent* 2000). But no formal scientific studies appear to have been undertaken (or at least published) on any of the PNG islands and atolls or in the Reef Islands.

The category of “climate-change refugee” (or more generally “environmental refugee”), introduced by well-meaning nongovernmental organizations, has become widely used not just for Pacific Islanders but seemingly globally. It has been resisted in many places, including atoll states, both for its evident denial of sovereignty and independence and because Islanders have no wish to be seen as powerless, passive, or tragic victims, even if that elicits empathy and support elsewhere (McNamara and Gibson 2009; Farbotko 2005; Farbotko and Lazrus 2012; Pam and Henry 2012). Becoming climate-change refugees has been strongly contested in Tuvalu (Farbotko and Lazrus 20012) and much disliked elsewhere, as in Kiribati, where the government has sought “migration with dignity.” Even where global warming and environmental degradation have been highly politicized, as in Tuvalu, the key objective of migration nonetheless remains one of superior access to employment opportunities and the social and economic benefits derived from them, such as access to housing, health, and education services (Mortreux and Barnett 2009; Paton and Fairbairn-Dunlop 2010; Chevalier 2010; Shen and Binns 2012; Shen and Gemenne 2011; Smith 2013). Yet, in Tuvalu at least, fears of change are such that over a decade ago emigration had already occurred in anticipation of environmental threats (Connell 2003), and some present residents have deliberately prepared for an “out-of-country” option by having particular family members migrate and create a base abroad (Paton and Fairbairn-Dunlop 2010).

For many years, small Pacific islands, especially atolls, have exhibited a considerable degree of poverty (Connell 1985, 2013a), and Islanders have long regarded migration as a means of achieving diversification of

livelihoods and acquiring remittances. Extreme population pressure on resources was evident on many islands in the nineteenth century (Bayliss-Smith 1975; Green and Green 2007), long before any consideration of environmental change influenced island livelihoods. Islanders as well as colonial and postcolonial authorities often recognized migration as a means of alleviating poverty and food insecurity, and, in Kiribati and Tokelau, formal resettlement schemes were put in place. In many other islands, such as the Bougainville and Solomon Islands atolls and the Reef Islands, Islanders developed their own strategies for migration, colonization, or resettlement elsewhere.

VULNERABLE ISLANDS?

In every island in the southwestern Pacific, a combination of physical factors (tectonic movements, SLR, ENSO events, cyclones) and human factors (particularly intense where populations are growing) has contributed to environmental change. Disaggregating them is rarely possible. Along the Australia-Pacific plate boundary especially, flooding and coastal erosion have been the outcome of tectonic changes, seismic events, ENSO-related tidal and storm surges, and human modifications rather than of SLR. Variations in SLR have long been influenced by a range of geophysical, gravitational, and oceanographic processes, but diverse measuring techniques (and relatively few of them) and short-term records complicate analysis and explanation. Beyond “visibilism,” climate change is unusually difficult to assess because it is nonlinear and cross-scalar and has diverse physical impacts. Yet atolls have exhibited considerable stability over the past half century.

However, rates of SLR are increasing, and greater concentrations of carbon dioxide—ocean acidification—may reduce the calcification of coral reefs and their ability to offer protection, hence threats to atolls are growing. Temperature fluctuations following ENSO events have contributed to reef bleaching, with the consequent reduction in the resilience of coral atolls to other environmental changes and the decline of fisheries potential (Williams et al 2010). Weak evidence suggests that, at least in parts of Solomon Islands, the fruiting patterns of breadfruit have changed (Lal, Kinch, and Wickham 2009). Marine life is moving southward and eastward (Bell et al 2013; Poloczanska et al 2013). This is both a greater threat to island livelihoods than contemporary terrestrial changes—eco-

conomic or environmental—and a metaphor and forerunner for human migration. The rate of SLR is likely to increase in any future scenario, but the scale and impacts are necessarily uncertain, in part because it is impossible to know how society will evolve in this century. While most atolls are not being eroded overall (though ENSO events, cyclones, and storm surges pose critical problems), physical viability is under increased pressure. It is difficult to attribute specific ecological and environmental changes to climatic change, however much of an “aggravating factor” it appears, even in the western Pacific. Contemporary climate change from global warming is yet to have a more significant impact on environmental change in Pacific Islands than other physical and human factors. That tipping point is yet to be reached. Still, if local and regional factors are presently more significant and “local practices are identified as contributing to . . . problems, they cannot absolve external forces from the contribution of global climate change” (Kelman 2010, 611).

Environmental changes are evident in most island contexts, but particularly in the two iconic contexts: urbanized atolls (like Funafuti) and islands close to the Australia–Pacific plate (notably the Bougainville atolls). In both contexts, they are frequently attributed to climate change by local people who have been exposed to dominant external discourses of climate-change impacts (Rudiak-Gould 2010, 2011, 2013a, 2013b; Kuruppu and Liverman 2011; Farbotko and Lazrus 2012) and by the media. The absence of local scientific studies contributes to this interpretation. Consequently, the more extreme and unusual examples of environmental change that have occurred in urbanized atolls and near the plate boundary have come to represent the norm. Significant physical changes were occurring on coral atolls long before the 1990s (and Islanders were conscious of them and taking steps to manage them) when the first inklings of a “greenhouse effect,” climate change, and SLR reached the Pacific region. The impact of natural hazards—whether SLR, cyclones, or tidal surges—has been accentuated by intensified population pressure and human modification of coastal land. In most small island contexts, people are more concerned about immediate daily needs than about the long-term implications of climate change and SLR (Storey and Hunter 2010; Gaillard 2012; Connell 2013b; Smith 2013). Indeed, in such contexts, as in Kiribati, a focus on climate change can be a “harmful distraction” from meeting present-day livelihood challenges (Gaillard 2012; Kelman 2014).

A combination of physical and human influences has placed consid-

erable pressure on small islands and particularly atolls. Achieving sustainable and acceptable livelihoods has become unusually and often increasingly difficult, as aspirations and expectations have increased. For centuries, Islanders have developed extended geographical ties and livelihoods with mobility and migration enabling security. A culture of migration has emerged; transnational corporations of kin have evolved, with “anchor populations” at home, as a form of “homeland security” for possible return migration (Marcus 1981; Rauchholz 2012), creating multiple versions of the manner in which livelihoods and identity are “determined by both rootedness around the central place and controlled mobility” (Bonnemaison 1984, 133). The majority of small islands in the Pacific—where populations were never more than a thousand—probably now have over half their population in urban centers or overseas (Connell 2013a). Many atolls and all three independent Pacific atoll states have very high population densities and significant concentrations of vulnerable people, and they share common characteristics of small size, limited land and skilled human resources, threatened biodiversity, and scarce financial resources. Economic development is severely constrained.

Small islands and atolls especially are particularly vulnerable to environmental change, partly because they have limited economies, with little flexibility or diversity, that have long resulted in interisland ties and an increased orientation to migration. As contemporary environmental changes have been widely attributed to climate change and SLR, a discourse of climate change–induced migration has become widely accepted in and for atoll environments with the emergence of “climate-change refugees” supposedly imminent (Connell 2013b). While official state discourses may reject the latter and focus on mitigation and adaptation (McNamara and Gibson 2009), atoll states have simultaneously negotiated migration agreements with metropolitan nations, which have sustained and emphasized Islander aspirations for migration. The conjuncture of enhanced and interlinked socioeconomic and environmental influences on migration, though seemingly intensifying island vulnerability, has created a new discourse of migration and human rights, where local agency is significant and a diasporic future more probable, even if migration is never entirely in circumstances of Islanders’ own choosing.

* * *

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Abstract

Small Pacific islands, especially atolls, have been widely argued to be in the forefront of climate change. Recent degradation of island environments has primarily been attributed to the impact of sea-level rise. However, physical changes to several small islands can be linked to a range of physical influences and to human modification. La Niña events, cyclones, and wind waves have caused localized flooding and storm damage. Most atoll islands have not significantly changed in size, as deposition balances erosion. Many islands have experienced broadly similar environmental problems in earlier times, at different scales, and over different time periods, now accentuated by human pressures on scarce land areas and resources. Local human factors (including construction and mining), tectonic subsidence, and La Niña events have created some iconic sites that have become sym-

bols of sea-level rise, sometimes erroneously attributed solely to global warming. Limited economic prospects in most small islands, rising expectations, and growing populations have contributed to a culture of migration, marked by international migration and urbanization, that has diversified impoverished livelihoods, extended island geographies, and resulted in accentuated population concentrations. Contemporary climate change exacerbates present environmental changes, stimulates further migration, and points to diasporic futures.

KEYWORDS: atolls, climate change, sea level, tectonics, urbanization, migration