



THE UNIVERSITY *of* EDINBURGH

Edinburgh Research Explorer

Skeletal carbonate mineralogy of Scottish bryozoans

Citation for published version:

Loxton, J, Spencer Jones, M, Najorka, J, Smith, AM & Porter, JS 2018, 'Skeletal carbonate mineralogy of Scottish bryozoans', *PLoS ONE*. <https://doi.org/10.1371/journal.pone.0197533>

Digital Object Identifier (DOI):

[10.1371/journal.pone.0197533](https://doi.org/10.1371/journal.pone.0197533)

Link:

[Link to publication record in Edinburgh Research Explorer](#)

Document Version:

Publisher's PDF, also known as Version of record

Published In:

PLoS ONE

General rights

Copyright for the publications made accessible via the Edinburgh Research Explorer is retained by the author(s) and / or other copyright owners and it is a condition of accessing these publications that users recognise and abide by the legal requirements associated with these rights.

Take down policy

The University of Edinburgh has made every reasonable effort to ensure that Edinburgh Research Explorer content complies with UK legislation. If you believe that the public display of this file breaches copyright please contact openaccess@ed.ac.uk providing details, and we will remove access to the work immediately and investigate your claim.



RESEARCH ARTICLE

Skeletal carbonate mineralogy of Scottish bryozoans

Jennifer Loxton^{1,2,3aa*}, Mary Spencer Jones², Jens Najorka⁴, Abigail M. Smith⁵, Joanne S. Porter^{1,2ab}

1 Centre for Marine Biodiversity and Biotechnology, School of Life Sciences, Heriot-Watt University, Riccarton, Edinburgh, United Kingdom, **2** Department of Life Sciences, Natural History Museum, London, United Kingdom, **3** University Marine Biological Station, Millport, Isle of Cumbrae, United Kingdom, **4** Core Research Laboratories, Natural History Museum, London, United Kingdom, **5** Department of Marine Science, University of Otago, Dunedin, New Zealand

^{aa} Current address: School of Geosciences, University of Edinburgh, Edinburgh, United Kingdom

^{ab} Current address: International Centre for Island Technology, Heriot-Watt University, Stromness, Orkney, United Kingdom

* Jennifer.loxton@ed.ac.uk



Abstract

This paper describes the skeletal carbonate mineralogy of 156 bryozoan species collected from Scotland (sourced both from museum collections and from waters around Scotland) and collated from literature. This collection represents 79% of the species which inhabit Scottish waters and is a greater number and proportion of extant species than any previous regional study. The study is also of significance globally where the data augment the growing database of mineralogical analyses and offers first analyses for 26 genera and four families. Specimens were collated through a combination of field sampling and existing collections and were analysed by X-ray diffraction (XRD) and micro-XRD to determine wt% MgCO₃ in calcite and wt% aragonite. Species distribution data and phylogenetic organisation were applied to understand distributional, taxonomic and phylo-mineralogical patterns. Analysis of the skeletal composition of Scottish bryozoans shows that the group is statistically different from neighbouring Arctic fauna but features a range of mineralogy comparable to other temperate regions. As has been previously reported, cyclostomes feature low Mg in calcite and very little aragonite, whereas cheilostomes show much more variability, including bimineralic species. Scotland is a highly variable region, open to biological and environmental influx from all directions, and bryozoans exhibit this in the wide range of within-species mineralogical variability they present. This plasticity in skeletal composition may be driven by a combination of environmentally-induced phenotypic variation, or physiological factors. A flexible response to environment, as manifested in a wide range of skeletal mineralogy within a species, may be one characteristic of successful invasive bryozoans.

OPEN ACCESS

Citation: Loxton J, Spencer Jones M, Najorka J, Smith AM, Porter JS (2018) Skeletal carbonate mineralogy of Scottish bryozoans. *PLoS ONE* 13 (6): e0197533. <https://doi.org/10.1371/journal.pone.0197533>

Editor: Ulrike Gertrud Munderloh, University of Minnesota, UNITED STATES

Received: January 23, 2018

Accepted: May 3, 2018

Published: June 13, 2018

Copyright: © 2018 Loxton et al. This is an open access article distributed under the terms of the [Creative Commons Attribution License](https://creativecommons.org/licenses/by/4.0/), which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

Data Availability Statement: All relevant data are within the paper and its Supporting Information file.

Funding: The authors thank the Marine Alliance for Science and Technology in Scotland (MASTS) and the Marine Environmental Research Group (MERG) at Heriot-Watt University for supporting and funding J.L.'s PhD. Natural Environment Research Council [NE/M006566/2] to Dr Jennifer Loxton. The funders had no role in study design, data collection and analysis, decision to publish, or preparation of the manuscript.

Competing interests: The authors have declared that no competing interests exist.

Introduction

Many marine organisms make carbonate shells and skeletons, which can act as repositories for information about seawater conditions. Biomineralisation can be controlled by the environment (e.g. algae [1] and corals [2,3] but many invertebrates, such as Foraminifera [4] and Mollusca [5], also exert biological control over the calcification process, managing nucleation [6–8], ultrastructure and crystal fabric [9–11], and carbonate composition [12–15]. Calcification can develop along with ontogeny/astogeny (e.g., growth rate [16–18], age [19–21]) or be determined by genetic factors and phylogenetic position [1,22–24]. Even highly controlling organisms, however, may be influenced by environmental parameters such as temperature [16–21], salinity [22–24], and pH [25,26].

Marine bryozoans are colonial filter-feeding invertebrates and are well characterised globally in terms of mineralogy [27]. Unlike some groups, where perhaps 2–3% of species have actually been measured (e.g., 29 chiton species = 3% of extant species [22]), marine bryozoans have been extensively studied at both ends of the Earth [14,15,17–24], such that at least a quarter of extant species have been characterised at least once (~1500 species [28]).

Bryozoan skeletons are mineralogically variable, and can be entirely calcitic, entirely aragonitic or bimineralic, featuring more than one CaCO_3 polymorph [27,29]. There is also a wide range of Magnesium (Mg) content in bryozoan calcite, ranging from low magnesium carbonate (LMC, < 4 wt% MgCO_3), through intermediate magnesium carbonate (IMC, >4 wt% MgCO_3 to < 8 wt% MgCO_3) to high magnesium carbonate (HMC, >8 wt% MgCO_3) [20,27] with most bryozoans featuring low to intermediate Mg-calcite [27,20,29,30]. This variation can be ascribed to environmental, physiological or phylogenetic influences [31].

Despite the strong phylogenetic signal in bryozoan mineralogy [24], it is links between mineralogy and environmental conditions that are the main driving force behind many mineralogical studies. Since the mid twentieth century [32] the skeleton composition of biological groups, such as the Foraminifera [33] and Mollusca [34], have been used in palaeoclimatology, where historic and fossil specimens have been assumed to hold a record of the seawater chemistry and temperature from the time in which it was deposited. Additionally, in recent years, it has been hoped that animals with variable calcium carbonate chemistry, such as the Bryozoa, might be able to act as a “bellweather” for climate change and ocean acidification [35]—with their skeletal chemistry proving that changes in our oceans are having a measureable effect on the biota which inhabit them. It is this spur of climate change and ocean acidification that has fuelled some of the increase in bryozoan mineralogy studies over the past decade.

Since the beginning of mineralogy studies on the Bryozoa there have been region-specific publications. The majority are of limited use due to the very low specimen numbers analysed (e.g. South Africa [36], Hawaii [37], Talbot Shoal [38], Naples [39], Malaysia [40]). There have, however, also been a handful of studies where the coverage of regional bryozoans has been great enough to draw meaningful conclusions. These studies include: the Mediterranean with 94/300 species analysed [41]; Antarctica with 21/300 species analysed [42]; New Zealand with 49/953 species analysed [43]; Chile with 23/267 species analysed [24] and the Arctic with 76/300 species analysed [20].

Scotland was chosen for this regional study for its ecological and biological diversity. Scotland lies on the convergence of the Atlantic Ocean and the North Sea, which provides both a wide range of ecological niches with regards to temperature and depth [44] and a source for both Northern and Southern bryozoan species [45]. There have been 7 published mineralogical studies [27,30,41,42,46–48] featuring bryozoan species which are present in the Scottish fauna. The earliest of these were conducted with titration [46,47], with the remainder being conducted with XRD, and, in a few cases, Raman spectroscopy [30]. These publications

provided analyses of 41 species ($n = 148$), although specimens collected from Scottish waters were only used in the analysis of 5 species ($n = 8$).

Here we report comprehensively on the variation in bryozoan skeletal carbonate mineralogy from around Scotland and how it compares to other temperate regions. We quantify the effects of phylogeny on skeletal composition, and evaluate the extent to which environment plays a role in phenotypic expression, particularly within species.

Materials and methods

Sample collection, archiving, and preparation

This study is based on bryozoan species collected from Scottish waters. Material was taken from existing collections and field sampling in order to collect samples of as many species as possible. None of the specimens in this study are endangered or protected. Full details of all specimens can be found in Table A in [S1 File](#).

The Bryozoa collections of the Natural History Museum (NHM) in London and the National Museum of Scotland (NMS) were searched for Scottish species with permission from the museum curators. 38 species were sourced from the NHM collection and 43 species were sourced from the NMS collection. The private collections of Dr Joanne Porter, Heriot-Watt University and Dr Jim Drewery, Marine Scotland (Rockall collection) were also investigated, resulting in a further 10 species. Field sampling was conducted by hand (mostly using SCUBA), in Scottish waters between 2010 and 2013. Sample collection localities ranged from 60.8°N in Shetland to 54.9°N in Stranraer, Scotland. Water depths ranged from intertidal to 35m but were mostly less than 20m. Permissions for field sampling were obtained where required (from private property or protected sites) and details can be found in Table B in [S1 File](#). For field sampling from all other sites, which were neither private property nor protected, permission was not required.

Species were identified to species level under a dissection (stereo) microscope (Zeiss) using the monographs of Hayward and Ryland [49,50]. Taxonomy was corrected to match the World Register of Marine Species [51]. Samples were extracted from the tip of erect colonies and the growing edge of encrusting colonies. A minimum of 5 zooids were extracted for each sample. As far as possible, care was taken to ensure that no substrate (e.g. coralline algae) or epibionts were included within the sample as they could potentially contaminate results. Each specimen was examined under a microscope and any substrate or epibionts were carefully scraped away with the tip of a scalpel blade; no chemical cleaning or treatment was conducted on specimens as it can have an impact on mineralogy [52]. Rare, figured, type or holotype specimens were not sampled but were analyzed whole using non-destructive micro-XRD. Vouchers for all specimens were retained, and reference specimens for each species sampled have been accessioned to the collections of the NHM, London.

Data for an additional 508 specimens of 4 Scottish species were included from scientific publications [15,52].

Skeletal mineralogy analysis

The majority of mineralogical analyses were conducted at the Imaging and Analysis Centre (NHM London) using semi-quantitative X-ray diffractometry (XRD) following methods described in Loxton et al. [14]. The XRD instrument used was a high-precision Nonius XRD with a position-sensitive detector and cobalt generated X-rays. Compositional information from XRD analysis is considered accurate to within 2% on a well-calibrated instrument [20].

Additionally, some rare, figured, type or holotype samples, as identified by underlining in [Table 1](#), were analysed whole using non-destructive micro-XRD. Qualitative phase identification and mineralogical analyses of whole specimens were conducted at the NHM using Micro-

Table 1. Skeletal carbonate mineralogy of 154 species of bryozoans. See Table A in [S1 File](#) or source literature [15, 52] for full details of each specimen.

Species Bold = first analysis of species Underline = first analysis of genus	No. of specimens (* = Micro XRD)	Wt.% MgCO ₃ in calcite. Mean if >1 specimen, Range (mean)	Wt% calcite range (mean)
<i>Aetea anguina</i> (Linnaeus, 1758)	1*	7.9	100
<i>Aetea sica</i> (Couch, 1844)	1*	-	100
<i>Aetea truncata</i> (Landsborough, 1852)	1*	-	100
<i>Alderina imbellis</i> (Hincks, 1860)	1	6.1	100
<i>Amphiblestrum auritum</i> (Hincks, 1877)	1	5.6	100
<i>Amphiblestrum flemingii</i> (Busk, 1854)	1	3.6	100
<i>Amphiblestrum solidum</i> (Packard, 1863)	1	7.2	100
<i>Anarthropora monodon</i> (Busk, 1860)	3	-	0–0 (0)
<i>Annectocyma major</i> (Johnston, 1847)	1*	6.4	100
<i>Beania mirabilis</i> Johnston, 1840	1	7.2	100
<i>Bicrisia abyssicola</i> Kluge, 1962	1	2.4	100
<i>Bicellariella ciliata</i> (Linnaeus, 1758)	1	5.0	82
<i>Bicellarina alderi</i> (Busk, 1859)	1	2.8	57
<i>Bugula neritina</i> (Linnaeus, 1758)	1	5.1	100
<i>Bugulina avicularia</i> (Linnaeus, 1758)	1	4.8	100
<i>Bugulina flabellata</i> (Thompson, in Gray, 1848)	1	3.9	58
<i>Bugulina fulva</i> (Ryland, 1960)	1	6.6	100
<i>Bugulina simplex</i> (Hincks, 1886)	1	3.9	49
<i>Bugulina turbinata</i> (Alder, 1857)	1	2.2	100
<i>Buskea dichotoma</i> (Hincks, 1862)	1	4.5	100
<i>Buskea nitida</i> Heller, 1867	1	4.6	100
<i>Caberea ellisii</i> (Fleming, 1814)	1	4.9	100
<i>Callopora craticula</i> (Alder, 1856)	1	4.3	100
<i>Callopora dumerilii</i> (Audouin, 1826)	1	7.5	100
<i>Callopora lineata</i> (Linnaeus, 1767)	1	5.2	83
<i>Callopora rylandi</i> Bobin and Prenant, 1965	1	6.9	100
<i>Cauloramphus spiniferum</i> (Johnston, 1832)	1	3.8	85
<i>Cellaria fistulosa</i> (Linnaeus, 1758)	16	3.5–7.1 (5.1)	100–100 (100)
<i>Cellaria salicornioides</i> Lamouroux, 1816	1	3.7	100
<i>Cellaria sinuosa</i> (Hassall, 1840)	1	3.9	92
<i>Cellepora pumicosa</i> (Pallas, 1766)	1	4.7	89
<i>Celleporella hyalina</i> (Linnaeus, 1767)	1	1.7	100
<i>Celleporina caliciformis</i> (Lamouroux, 1816)	1	4.9	100
<i>Celleporina pygmaea</i> (Norman, 1868)	1	7.2	100
<i>Carbasea carbasea</i> (Ellis and Solander, 1786)	1	6.3	100
<i>Chartella barleei</i> (Busk, 1860)	1	9.6	28
<i>Chartella papyracea</i> (Ellis and Solander, 1786)	1	9.0	100
<i>Chorizopora brongniartii</i> (Audouin, 1826)	1	5.5	100
<i>Conopeum reticulum</i> (Linnaeus, 1767)	1	5.0	100
<i>Conopeum seurati</i> (Canu, 1928)	1	9.1	95
<i>Coronopora truncata</i> (Fleming, 1828)	1	7.8	100
<i>Cradoscrupocellaria reptans</i> (Linnaeus, 1758)	3	3.8–6.8 (5.1)	65–94 (82)
<i>Cribrilina annulata</i> (O. Fabricius, 1780)	1	5.0	89
<i>Cribrilina cryptoecium</i> Norman, 1903	1	2.9	100
<i>Cribrilina punctata</i> (Hassall, 1841)	1	5.0	100

(Continued)

Table 1. (Continued)

Species Bold = first analysis of species Underline = first analysis of genus	No. of specimens (* = Micro XRD)	Wt.% MgCO ₃ in calcite. Mean if >1 specimen, Range (mean)	Wt% calcite range (mean)
<i>Crisia aculeata</i> Hassall, 1841	1	4.8	90
<i>Crisia denticulata</i> (Lamarck, 1816)	1	3.3	100
<i>Crisia eburnea</i> (Linnaeus, 1758)	2	3.3–3.4 (3.3)	100
<i>Crisia ramosa</i> Harmer, 1891	1	2.3	100
<u><i>Crisidia cornuta</i></u> (Linnaeus, 1758)	1	5.0	100
<i>Crisularia plumosa</i> (Pallas, 1766)	1	7.4	100
<i>Crisularia purpuroincta</i> (Norman, 1868)	1	3.9	89
<i>Cryptosula pallasiana</i> (Moll, 1803)	1	4.6	59
<i>Cylindroporella tubulosa</i> (Norman, 1868)	1*	7.0	100
<i>Dendrobeatia murrayana</i> (Bean, in Johnston, 1847)	1	4.7	100
<i>Diplosolen obelia</i> (Johnston, 1838)	1*	5.8	100
<i>Disporella hispida</i> (Fleming, 1828)	1	5.2	100
<u><i>Doryporellina reticulata</i></u> (Ryland, 1963)	1*	7.2	
<i>Einhornia crustulenta</i> (Pallas, 1766)	1	4.1	100
<i>Electra monostachys</i> (Busk, 1854)	1	7.3	91
<i>Electra pilosa</i> (Linnaeus, 1767)	14	6.8–9.9 (8.6)	53–100 (80.7)
<u><i>Entalophoroecia deflexa</i></u> (Couch, 1842)	1	4.5	100
<i>Escharella abyssicola</i> (Norman, 1869)	1	7.5	94
<i>Escharella immersa</i> (Fleming, 1828)	147 ^a	4.3–6.9(5.7)	25–99(69.1)
<i>Escharella labiosa</i> (Busk, 1856)	1*	5.8	100
<i>Escharella laqueata</i> (Norman, 1864)	1*	6.0	100
<i>Escharella octodentata</i> (Hincks, 1880)	1	7.3	100
<i>Escharella variolosa</i> (Johnston, 1838)	1	6.6	88
<i>Escharella ventricosa</i> (Hassall, 1842)	1	5.6	100
<i>Escharina alderi</i> (Busk, 1856)	1	7.0	100
<i>Escharina dutertrei haywardi</i> Zabala, Maluquer and Harmelin, 1993	1	7.4	100
<i>Escharina johnstoni</i> (Quelch, 1884)	1*	7.2	100
<i>Escharoides coccinea</i> (Abildgaard, 1806)	1	3.3	64
<i>Escharoides mamillata</i> (Wood, 1844)	1	7.0	100
<i>Eucratea loricata</i> (Linnaeus, 1758)	1	7.7	85
<u><i>Eurystrotos compacta</i></u> (Norman, 1866)	1	6.3	100
<i>Fenestrulina malusii</i> (Audouin, 1826)	1	2.8	88
<i>Filicrisia geniculata</i> (Milne Edwards, 1838)	1	3.9	86
<i>Flustra foliacea</i> (Linnaeus, 1758)	97 ^b	6.2–13.5(9.4)	100
<i>Haplopoma graniferum</i> (Johnston, 1847)	1	3.1	75
<i>Haplopoma impressum</i> (Audouin, 1826)	1	1.8	100
<i>Haplopoma planum</i> Ryland, 1963	1	6.1	100
<i>Haplopoma sciaphilum</i> Silén and Harmelin, 1976	1*	7.5	
<u><i>Hemicyclopora polita</i></u> (Norman, 1864)	1	7.3	100
<i>Herentia hyndmanni</i> (Johnston, 1847)	1*	6.2	100
<i>Hippoporina pertusa</i> (Esper, 1796)	1*	6.3	100
<i>Hippothoa divaricata</i> Lamouroux, 1821	1*	-	0
<i>Hornera lichenoides</i> (Linnaeus, 1758)	1	5.6	94
<i>Idmidronea atlantica</i> (Forbes, in Johnston, 1847)	1	3.8	100

(Continued)

Table 1. (Continued)

Species Bold = first analysis of species Underline = first analysis of genus	No. of specimens (* = Micro XRD)	Wt.% MgCO ₃ in calcite. Mean if >1 specimen, Range (mean)	Wt% calcite range (mean)
<i>Lagenipora lepralioides</i> (Norman, 1868)	1*	5.2	100
<i>Larnacicus corniger</i> (Busk, 1859)	1	7.7	94
<i>Lepraliella hippopus</i> (Smitt, 1867)	1	7.7	96
<i>Marguetta lorea</i> (Alder, 1864)	1	8.5	100
<i>Megapora ringens</i> (Busk, 1856)	1	6.5	44
<i>Membranipora membranacea</i> (Linnaeus, 1767)	39	0-0(0)	0-0 (0)
<i>Membraniporella nitida</i> (Johnston, 1838)	140 ^a	2.2-7.9(6.2)	100
<i>Microporella ciliata</i> (Pallas, 1766)	146 ^a	4.5-8.7(6.9)	34-99(77.2)
<i>Neolagenopora eximia</i> (Hincks, 1860)	1*	5.0	100
<i>Notoplites jeffreysii</i> (Norman, 1868)	1	5.9	13
<i>Omalosecosa ramulosa</i> (Linnaeus, 1767)	15	4.8-13.6 (9.8)	91-100 (98.2)
<i>Oncousoecia diastoporides</i> (Norman, 1869)	1	6.9	87
<i>Oncousoecia dilatans</i> (Johnston, 1847)	1	7.8	100
<i>Oshurkovia littoralis</i> (Hastings, 1944)	1	6.0	48
<i>Palmicellaria elegans</i> Alder, 1864	1*	7.4	100
<i>Palmiskenea skenei</i> (Ellis and Solander, 1786)	1	5.9	100
<i>Parasmittina trispinosa</i> (Johnston, 1838)	1	7.7	80
<i>Pentapora fascialis</i> (Pallas, 1766)	1	8.4	54
<i>Phaeostachys spinifera</i> (Johnston, 1847)	1	5.7	63
<i>Plagioecia patina</i> (Lamarck, 1816)	1	6.2	100
<i>Porella alba</i> Nordgaard, 1906	1	4.0	94
<i>Porella compressa</i> (J. Sowerby, 1805)	1	8.2	100
<i>Porella concinna</i> (Busk, 1854)	2	8.1-8.2 (8.1)	99-100 (99.5)
<i>Porella laevis</i> (Fleming, 1828)	1	10.2	100
<i>Porella struma</i> (Norman, 1868)	1*	5.9	100
<i>Pseudoflustra virgula</i> Hayward, 1994	1	5.6	100
<i>Puellina innominata</i> (Couch, 1844)	1	8.8	100
<i>Pyripora catenularia</i> (Fleming, 1828)	1	8.7	90
<i>Ragionula rosacea</i> (Busk, 1856)	1*	7.2	100
<i>Ramphonotus minax</i> (Busk, 1860)	1	6.9	100
<i>Reteporella beaniana</i> (King, 1846)	1	8.5	100
<i>Reteporella incognita</i> Hayward and Ryland, 1996	1	8.4	100
<i>Reteporella watersi</i> (Nordgaard, 1907)	1	5.8	87
<i>Rosseliana rosselii</i> (Audouin, 1826)	1*	6.5	100
<i>Schizomavella (Schizomavella) auriculata</i> (Hassall, 1842)	1	5.2	27
<i>Schizomavella (Schizomavella) linearis</i> (Hassall, 1841)	1	6.4	1
<i>Schizoporella japonica</i> Ortmann, 1890	19	3.9-7.4 (5.3)	19-86 (56.9)
<i>Schizoporella patula</i> Hayward and Ryland, 1995	2	7.6-8.1 (7.9)	1-13 (7)
<i>Schizoporella unicornis</i> (Johnston, in Wood, 1844)	1	6.0	59
<i>Scruparia ambigua</i> (d'Orbigny, 1841)	1	8.1	100
<i>Scruparia chelata</i> (Linnaeus, 1858)	1	7.9	100
<i>Scrupocellaria scrupea</i> Busk, 1852	1	1.4	100
<i>Scrupocellaria scruposa</i> (Linnaeus, 1758)	5	3.3-6.3 (4.6)	57-100 (91.4)
<i>Securiflustra securifrons</i> (Pallas, 1766)	1	9.8	91

(Continued)

Table 1. (Continued)

Species Bold = first analysis of species Underline = first analysis of genus	No. of specimens (* = Micro XRD)	Wt.% MgCO ₃ in calcite. Mean if >1 specimen, Range (mean)	Wt% calcite range (mean)
<i>Setosella vulnerata</i> (Busk, 1860)	1	-	0
<i>Smittina bella</i> (Busk, 1860)	1*	-	0
<i>Smittina crystallina</i> (Norman, 1867)	1	8.1	67
<i>Smittoidea marmorea</i> (Hincks, 1877)	1*	-	0
<i>Smittoidea reticulata</i> (MacGillivray, 1842)	1	6.3	16
<i>Stigmatoechos violacea</i> (M. Sars, 1863)	1	7.4	100
<i>Stomacrustula cruenta</i> (Busk, 1854)	1*	9.0	-
<i>Stomacrustula sinuosa</i> (Busk, 1860)	1*	6.5	-
<i>Stomatopora gingrina</i> Jullien, 1882	1	7.2	93
<i>Tegella unicornis</i> (Fleming, 1828)	1	3.6	100
<i>Temachia microstoma</i> (Norman, 1864)	1*	6.3	100
<i>Tervia irregularis</i> (Meneghini, 1844)	1	8.0	100
<i>Tessarado ma boreale</i> (Busk, 1860)	1	6.4	100
<i>Tricellaria inopinata</i> d'Hondt and Occhipinti Ambrogi, 1985	1	2.4	100
<i>Tricellaria ternata</i> (Ellis and Solander, 1786)	1	5.0	100
<i>Tubulipora liliacea</i> (Pallas, 1766)	2	2.6–2.7 (2.6)	97–100 (98.5)
<i>Tubulipora penicillata</i> (O. Fabricius, 1780)	1	4.7	100
<i>Tubulipora phalangea</i> Couch, 1844	1	1.9	100
<i>Tubulipora plumosa</i> Thompson, in Harmer, 1898	1	1.0	100
<i>Turbicellepora avicularis</i> (Hincks, 1860)	1	9.0	100
<i>Turbicellepora boreale</i> Hayward and Hansen, 1999	1	7.0	100
N	790	147	150
Min	1	0	0
Mean	5	6	87
Max	147	10	100
Range	146	9	100

^a includes analyses from Loxton et al., 2014 [15]—*Membraniporella nitida* (n = 139), *Microporella ciliata* (n = 145) and *Escharella immersa* (n = 146)

^b includes analyses from Loxton et al., 2017 [52]—*Flustra foliacea* (n = 78)

<https://doi.org/10.1371/journal.pone.0197533.t001>

XRD with a GeniX High Flux Beam Delivery System and an INEL 120⁰ position-sensitive-detector. FOX2D mirror optics (XENOCs) focussed the X-ray beam of copper radiation to a spot size of 230µm. Protruding and flat skeletal surfaces of the specimens were targeted using an AxioCam MRc5 microscope camera. The error associated with this method is higher than with traditional XRD due to potential non-random orientation of crystallites in calcified skeletons and minor sample displacement. In this study uncertainties were estimated to be approximately 10% based on the duplicate analysis of the same sample using both micro-XRD and conventional XRD. Both XRD and micro-XRD instruments were calibrated daily using pure silica (Si) and silver behenate (AgC₂₂H₄₃O₂) on a quartz substrate [53,54].

Statistics and data analysis

Wt% MgCO₃ in calcite data for all species was tested for normality using Anderson-Darling normality tests. Criteria for parametric *posthoc* testing were not met due to unequal sample sizes and heterogeneous variance between datasets.

Phylogenetic data and branch lengths were taken from the publication of Waeschenbach et al. [55]. Branch length of the phylogeny indicates the number of substitutions per site based on Bayesian multi-gene analysis of a concatenated 7-gene dataset (*ssrDNA*, *lsrDNA*, *rrnL*, *rrnS*, *cox1*, *cox3*, *cytb*) constructed using BEAST under the random logical clock and GTR+I+G model [55]. The phylogeny was input into the R statistical language [56] using Newman coding. Kembel's [57] methodology for Comparative Phylogenetic Methods was followed. Comparative phylogenetic diversity and phylogenetic distance between cyclostomes and cheilostomes were calculated using the "pd" and "cophenetic" functions in the Picante package for R [58,59]. Blomberg's K [60] values for cyclostomes, cheilostomes and all Scottish bryozoans were calculated as a measure of underlying phylogenetic signal in the traits of wt% MgCO₃ in calcite and wt% calcite using the "multiPhyloSignal" function in the Picante package for R [58,59]. Blomberg's K was also calculated as a measure of underlying phylogenetic signal for the trait of morphology.

To elucidate differences between taxonomic, spatial, evolutionary and ecological groupings, non-parametric Mann-Whitney U-tests were carried out. Data from three temperate regional studies (Scotland, this study; Chile, [24]; New Zealand, [27]) were each analysed using a generalised linear model (GLM) ANOVA. The factor used was region (fixed) and the response wt% MgCO₃ in calcite. Criteria for parametric *posthoc* testing were not met due to unequal sample sizes between datasets, therefore potential differences between regions were elucidated by Mann-Whitney *Post-hoc* testing.

Results

The skeletons of 790 bryozoan specimens from a total 154 species from temperate Scottish waters were analysed or collated from literature (Table 1) and had a mean wt% calcite of 84. The most common mineralogy type was 100% calcite; 93 species (60%) were formed entirely from calcite. Bimineralic species, containing a mixture of calcite and aragonite polymorphs, were the second most common mineralogy type. Fifty-five species, or 36% of those tested had this type of mineralogy and the proportion of the two types was spread fairly evenly across the possible proportional range, from 1% to 99% calcite. Entirely aragonitic species were the least common with just 6 species (4%) featuring this type of mineralogy.

All specimens were analysed for wt% MgCO₃ in calcite, which ranged from 0 to 14, with a mean of 6. The majority (69%) of species featuring calcite were classed as intermediate-Mg calcite (4–8 wt% MgCO₃). A further 18% formed low-Mg calcite (0–4 wt% MgCO₃) and the remaining 12% of species featured high-Mg calcite (8–12 wt% MgCO₃).

See Table A in S1 File for full specimen details.

Range of carbonate mineralogy in a group of bryozoan specimens is described as biomineral "space", a term introduced by Smith et al [27]. Biomineral space is the area described by the ranges of wt% MgCO₃ in calcite and wt% calcite for a particular species of group of species. It is usually expressed as a percentage of the possible space available for biomineralization (0–22 wt% MgCO₃ in calcite and 0–100 wt% calcite, a possible space of 2178wt%²). The biomineral space for Scottish bryozoans is shown in Fig 1 with Scottish species covering 42% of the total available mineral space.

Comparison to other regions

Table 2 compares the mineralogical profile of Scottish bryozoan species to those reported from other temperate regions, New Zealand and Chile, and to the neighbouring Arctic. This shows consistency among the mineralogical means reported from the temperate regions of Scotland, New Zealand and Chile although variation can be found among the ratios of species found in different mineralogical categories. The Arctic was the only region to show nearly 100% calcite

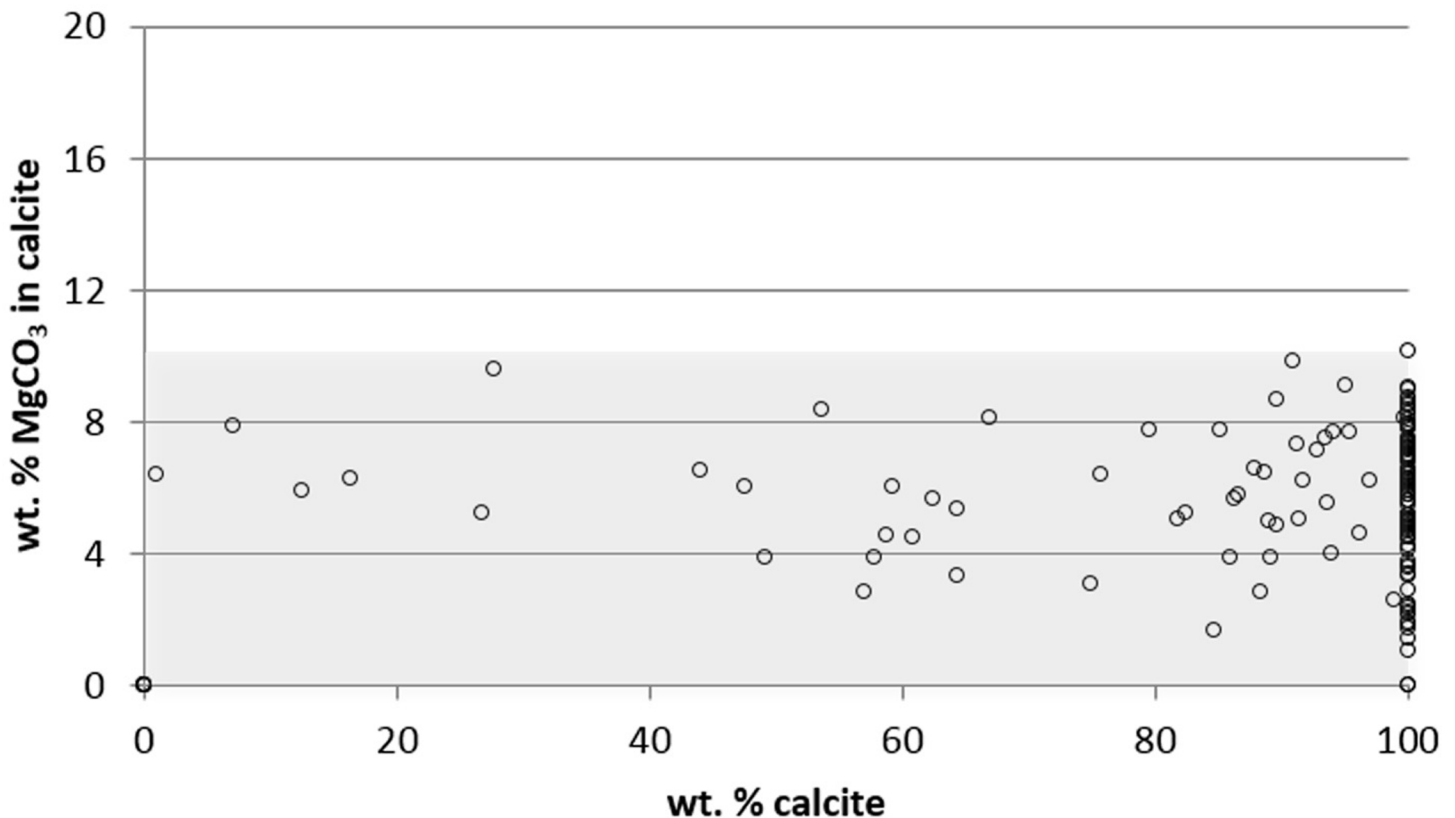


Fig 1. The biomineral space occupied by Scottish bryozoan species. Skeletal carbonate mineralogy of 154 species of Scottish marine bryozoans. Shaded box indicates maximum biomineral space occupied by the phylum in Scotland (42% of the total available biomineral space).

<https://doi.org/10.1371/journal.pone.0197533.g001>

across all species and lower incorporation of MgCO₃ in calcite than was found in any of the temperate regions.

There is a statistical difference in species wt% Mg in calcite between Scotland, New Zealand and Chile (ANOVA, $F = 3.12$, $P = 0.046^*$). *Post-hoc* Mann-Whitney analysis shows Scotland to be statistically different to New Zealand ($P = 0.030^*$) but not Chile and no statistical difference between Chile and New Zealand. Wt % MgCO₃ in calcite in Scottish species ($n = 154$) is

Table 2. Comparison of regional studies of bryozoan mineralogy. The ABC (aragonitic, bimineralic, calcitic) index quantifies the ratio of bryozoan species with particular mineralogies as first introduced by Borszcz et al [61].

	Scotland (this study, incl. data from [15,52])	New Zealand [43]	Chile [24]	Arctic [20]
Number of specimens	790	412	93	149
Number of species	154	49	23	76
Number of families	50	29	15	31
Distribution (latitude)	55–60°N	25–52°S	42–52°S	69–79°N
ABC index (aragonitic, bimineralic, calcitic ratio)	6:55:93	3:13:33	2:2:19	0:3:73
Aragonite: Bimineral: Calcite (% of species)	4:36:60	6:27:67	9:9:82	0:4:96
HMC: IMC: LMC (% of species)	6: 65: 29	0:86:14	18:57:25	0:57:43
Mean wt% calcite	84	85	84	100
Mean wt% MgCO ₃ in calcite	5.1	4.3	5.5	4.0

<https://doi.org/10.1371/journal.pone.0197533.t002>

statistically different (Mann-Whitney U-test, $P < 0.001^*$) to those from its geographical neighbour, the Arctic ($n = 76$) [20].

Taxonomic patterns

1. Class/Order. 125 species of Scottish cheilostomatous Bryozoa were found to contain a wide range of mineralogical compositions. 57% ($n = 73$) of species were found to be entirely calcitic with a further 39% ($n = 50$) featuring bimineralic mixing in some degree and only 5% ($n = 6$) found to be pure aragonite. The measured range for this group spanned the entire range from 0–100 wt% calcite, with a mean of 84 ($n = 125$). The majority (70%, $n = 85$) of species featured intermediate-Mg calcite (IMC) with the remaining species equally split between low-Mg calcite (LMC) and high-Mg calcite (HMC) (15%, $n = 18$ for both IMC and HMC). The mean of 6.0 wt% MgCO_3 in calcite ($n = 125$) reflects the dominance of IMC in Scottish cheilostomatous species.

27 specimens of 25 cyclostomatous species were analysed with 20 (80%) found to be entirely calcitic. The remaining five species showed some aragonite with the measured range showing between 82 and 99 wt% calcite. The mean wt% calcite for this class in Scotland is thus 98.5. This mean value for the Order Cyclostomatida is significantly higher than the mean for the Order Cheilostomatida (mean = 84.3, $n = 125$) (Mann-Whitney U-test, $P = 0.016^*$).

2. Families. The specimens studied come from 47 families in the phylum Bryozoa, almost all (94%) of the Recent families reported to occur in Scotland. Many of these families are only represented by one or two species although some include as many as 14 species. Although some families contain only one analysed specimen and conclusions should, therefore, be approached with caution, the mean number of specimens analysed per family is 16, and the maximum is 180, allowing some generalizations to be made. Coverage of Scottish genera within analysed families range from 33–100% with a mean of 81% of Scottish genera included here.

3. Species. The phylogenetic position for 39 (11 cyclostomes and 28 cheilostomes) Scottish bryozoan species were extracted from the bryozoan phylogeny published by Waeschenbach et al. [55] (Fig 2). There is no phylogenetic signal for the mineralogical trait of Mg content in calcite (Blomberg's $K = 0.17$, $P = 0.264$) although a strong and significant phylogenetic signal relating to calcite percentage was found (Blomberg's $K = 0.37$, $P = 0.022^*$). A similar analysis on the 28 cheilostomatous species for which phylogeny data is available revealed no phylogenetic signal associated with wt% MgCO_3 in calcite (Blomberg's $K = 0.38$, $P = 0.25$) but a statistically significant, strong phylogenetic signal associated with wt% calcite (Blomberg's $K = 0.82$, $P = 0.02^*$). Phylogenetic data was available for 11 cyclostome bryozoans and analysis of this revealed that there is no phylogenetic signal within mineralogy for either wt% MgCO_3 in calcite (Blomberg's $K = 0.75$, $P = 0.2$) or calcite (Blomberg's $K = 0.2$, $P = 0.99$).

Most bryozoan species were only tested once, but nine common species had more than 10 replicates: *Cellaria fistulosa* ($n = 16$), *Electra pilosa* ($N = 14$), *Flustra foliacea* ($n = 97$), *Membranipora membranacea* ($n = 39$), *Omalosecosa ramuosa* ($n = 15$), *Escharella immersa* ($n = 147$), *Membraniporella nitida* ($n = 140$), *Microporella ciliata* ($n = 146$) and *Schizoporella japonica* ($n = 19$), allowing some consideration of within-species environmental plasticity.

Cellaria fistulosa ($n = 16$) consistently exhibited two phases of calcite within its skeleton; both phases were reported in all specimens in an approximate ratio of 1:1. The first phase of LMC has a mean of 1.9 wt% MgCO_3 in calcite (range 0.5–3.7), while the second IMC phase showed a similar range of variation around a mean of 7.6 wt% MgCO_3 in calcite (range 5.9–9.9). *Electra pilosa* ($n = 14$) is also bimineralic, with a mean aragonite content of 19.3% (range 47%–0%) and the dominant HMC varying around a mean of 8.6 wt% MgCO_3 in calcite (range 6.8–9.9).

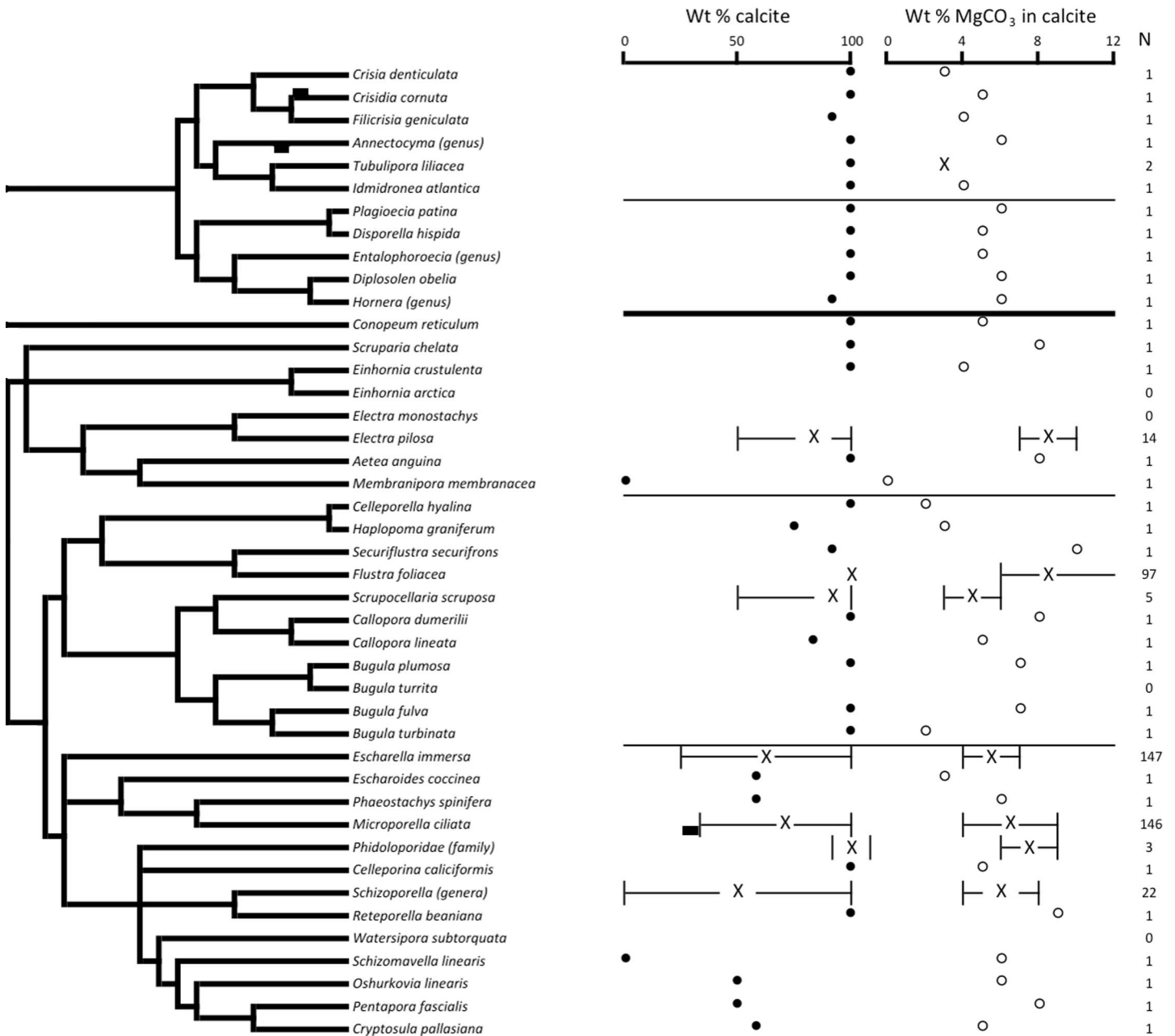


Fig 2. The phylogenetic distribution of the skeletal carbonate mineralogy of Scottish bryozoans. Phylogenetic tree adapted from Waeschenbach et al [55]. Crosses are means, with the range delineated by tails. Single measurements are represented by black (wt% MgCO₃ in calcite) or hollow (wt% calcite) circles.

<https://doi.org/10.1371/journal.pone.0197533.g002>

Flustra foliacea (n = 97) was found to be a variable species with mean HMC of 9.4 wt% MgCO₃ in calcite and a range of 7.3 (range 6.2–13.5). *Membranipora membranacea* (n = 39), in contrast, had absolutely no mineralogical variability with a consistent 100% aragonite skeleton. *Omalosecosa ramulosa* (n = 15) had the most variable calcite of all species analysed with calcite ranging from IMC to HMC and a mean of 9.8 wt% MgCO₃ in calcite and a range of 8.8 (range 4.8–13.6). *Escharella immersa* (n = 147) is bimineralic with a mean wt% MgCO₃ in calcite of 5.7 (range 4.3–6.9) and a mean wt% calcite of 69.1 (mean 25–99). *Membraniporella nitida* (n = 140)

contained no aragonite and has a mean wt % MgCO_3 in calcite of 6.2 (range 2.2–7.9). *Microporella ciliata* featured a mean of 6.9 wt% MgCO_3 in calcite (range 4.5–8.7) and a mean wt% calcite of 77.2 (range 34–99). *Schizoporella japonica* was the only non-native species with multiple analyses in this study and exhibited a bimineralic skeleton with a high degree of both calcitic and aragonitic variation. The mean of 43% aragonite in the skeleton varied widely (range 14–81%), and the IMC had a mean of 5.3 wt% MgCO_3 in calcite (range 3.9–7.4).

Discussion

Mineralogy of the phylum Bryozoa

This regional study of Scottish bryozoans contributes 282 measurements, and collates a further 508 measurements, of 154 species to existing knowledge of bryozoan mineralogy, including 110 species never before measured. These data increase the number of genera studied by 26, and the number of families by 4. In terms of both scope and environment, it is a Northern Hemisphere equivalent to published studies of New Zealand bryozoans [43] and Chilean bryozoans [24]. In this study we compared the taxonomic patterns discerned here with these two Southern Hemisphere temperate communities, and for contrast, with the neighbouring Arctic community. Table 2 shows the remarkable consistency among the three temperate regions with Scotland, New Zealand and Chile featuring means of 5.1, 4.3 and 5.5 wt% MgCO_3 in calcite and 84, 85, 84% calcite respectively. This dominance of IMC in bryozoan skeletons from temperate waters has been reported in previous publications investigating global [27] and regional [24,43] mineralogical patterns in the Bryozoa. Despite the similarity of the means, Scotland shows more variety in wt% MgCO_3 in calcite than New Zealand or Chile. This may be accounted for by the greater number of species presented in this study compared to the other regional studies (Table 2). The wide range of mineralogical variety and the solid presence of aragonite, LMC and HMC in Scottish bryozoan fauna may also be a combined reflection of the inclusion of Polar/Boreal and Lusitanian species in the Scottish fauna, due to Scotland's open geographical situation, and the diverse range of habitats and seasonal conditions which it offers.

Arctic species feature lower wt% MgCO_3 in calcite and less aragonite than Scottish species. Polar species feature slow growth rates [16, 62, 63] caused by low nutrient levels, low temperature and increased seasonality [16,64–66]. Links between mineralogy and growth rate have been shown with slower animals generally depositing less wt% MgCO_3 in calcite [67]. In addition the low temperatures and intrinsically low saturation of carbonate ions [68] in Polar waters means that deposition of calcite is favoured over aragonite [18,20,68]. The Arctic fauna features a few endemic species but predominantly consists of species with Pacific or Atlantic origin which entered Arctic waters following the last glacial maximum approximately 25 thousand years ago [69]. Arctic species exhibit a lower mean wt% MgCO_3 in calcite than Scotland, showing some adaptation to the Polar environment [18], although this mean is closer to that exhibited in Scotland than the mean for Antarctica [42], and possibly reflects the relatively young bryozoan fauna in the Arctic.

Mineralogy of classes/orders in the Bryozoa

Previous studies, including Smith et al., 2006 [27] and Boardman and Cheetham, 1987 [70], have found cyclostomes to be almost entirely calcitic. Boardman and Cheetham went as far as to state “All stenolaemates have calcareous skeletons. All skeletons are calcitic except for one reportedly aragonitic species from the Triassic.” [70] Cyclostomatida is the only extant order of the five comprising the Class Stenolaemata; it is only Recent specimens from this order which have been analysed in this Scottish study. Like Smith et al. [27] we find Boardman and Cheetham's statement to be mostly accurate although, like Smith et al. [27], we also found

trace amounts of aragonite in a few cases. Cheilostomes were found to be much more variable than cyclostomes and further investigation is recommended into the evolutionary radiation of cheilostomes and how that may have contributed to their greater biomineral space.

Mineralogy of families in the bryozoa

Smith et al. [27] surmised that bryozoan families fall into three general groups; those containing mostly aragonite, those containing mixed mineralogy and those containing mostly calcite. In general the data presented in Tables 1 and 3 concurs with this, although the distribution of Scottish species within these categories. (4% mostly aragonite, 32% mixed mineralogy, 64% mostly calcite) varies slightly from the Global data presented by Smith et al. [27] (5% mostly aragonite, 20% mixed mineralogy, 75% mostly calcite).

Most families, where more than one specimen has been analysed, show some level of variation within their mineralogy. Some of the IMC families that exhibit little or no variation are Aetidae ($n = 3$), Licheniporidae ($n = 2$) and Oncousoeciidae ($n = 3$). There are no families that consistently produce LMC or HMC. Families which are exclusively aragonitic (Setosellidae ($n = 1$), Exochenellidae ($n = 3$) and those which are aragonite-dominated (Stomachetosellidae, $n = 2$; Schizoporellidae, $n = 22$; Chaperiidae, $n = 1$) are found to be a mixture of flustrinid and ascophoran cheilostomes. The most variable family studied from Scotland is the Membraniporidae (18.64% of potential mineralogical space, $n = 41$).

There is no statistically significant difference between wt% $MgCO_3$ in calcite in families that appeared during aragonitic and calcitic seas (Table 3). Families that first appeared during periods of aragonitic seas (approx 50 Ma-Recent and during the Triassic), do however, almost exclusively feature intermediate to high Mg-calcite (Mean = 6.0, $n = 21$). This fits with the sea water chemistry of the time where Mg/Ca ratios were at their highest. Families forming LMC all evolved during times of calcitic seas where the Mg/Ca ratio in seawater was much lower, however more research is needed to confirm this potential pattern.

As has been highlighted by phylogenetic studies [55,72], there is a high degree of taxonomic uncertainty surrounding many bryozoan families which means that great care should be taken when applying generalisations at the family level. It should also be considered that many families are represented by low numbers of specimens in this dataset. Due to limited metadata in the case of some museum specimens, it has not been possible to consider variability, which may be caused by collection location, depth, season and developmental stage of individuals; all factors which are known to influence mineralogy.

Phylomineralogy in Scottish bryozoan species

Evolutionary origin of different species can be quantified using phylogenetic trees that often use branch length to approximate genetic distance of a species from their next nearest neighbour and offer an alternative to the use of First Appearance Date (FAD). A relatively recent advance in mineralogical methodology is the analysis of skeletal mineralogy alongside phylogenetic position of taxa, also termed “phylomineralogy” [1]. In many mineralogical studies samples are taken from multiple bryozoan lineages, and therefore do not represent statistically independent samples [60,73,74]. Some of the mineralogical differences between different species may be due to the divergent evolutionary history of the species and taxa [75], as shown in Fig 2. With the increasing availability of phylogenetic data mineralogical patterns can be assessed for a “phylogenetic signal” which can be quantified using measures such as Blomberg’s K [60] and taken into account during subsequent analysis and discussion. Phylogenetic data can also be included in Bayesian computational tools [76] where a variable such as phylogeny can be assessed as a potential explanation for observed species traits. Examples of

Table 3. Mineralogical characteristics of 50 bryozoa families from Scotland. The FAD stage details the first appearance datum, which is the first (oldest) appearance of the family in the geological record; it is reported here on the geologic time scale. The age at top (Ma) details how many million years ago these geological time periods were where they commenced.

Family	Genera	Species	n	Mean wt.% calcite	Mean wt.% MgCO ₃ in calcite	FAD stage	Age at top (Ma)
Aeteidae	1	3	3	100.0	7.9	Pliensbachian*	190
Annectocymidae	1	1	1	100.0	6.4	Albian***	110
Antroporidae	1	1	1	100.0	6.5	Maastrichtian**	65
Beaniidae	1	1	1	100.0	7.5	Bartonian**	37
Bitectiporidae	3	4	4	45.3	6.6	Ypresian**	49
Bryocryptellidae	3	8	9	99.2	7.2	Ypresian**	49
Bugulidae	4	11	11	85.0	4.6	Recent**	0
Calloporidae	7	12	12	92.6	5.4	Albian**	99
Candidae	4	7	13	87.5	4.2	Maastrichtian**	65
Cellariidae	1	3	18	97.3	5.3	Santonian**	83.5
Celleporidae	7	10	24	98.6	6.2	Priabonian**	33.7
Chaperiidae	1	1	1	5.9	7.7	Maastrichtian**	65
Chorizoporidae	1	1	1	100.0	5.5	Servallian**	11.2
Cribrilinidae	3	5	5	97.8	4.8	Cenomanian**	93.5
Crisiidae	4	7	8	96.5	3.6	Maastrichtian**	65
Cryptosulidae	1	1	1	58.8	4.6	Tortonian**	7.1
Diaperoeciidae	1	1	1	100.0	4.5	Hauterivian**	127
Diastoporidae	1	1	1	100.0	5.8	Pliensbachian**	190
Doryporellidae	1	1	1		7.2	Albian***	110
Electridae	3	4	16	89.2	6.6	Tithonian*	163.5
Escharinidae	3	5	152	92.5	6.7	Ypresian*	49
Eucrateidae	1	1	1	85.2	7.7	Recent**	0
Exochenellidae	1	1	3	0.0	0.0	Ypresian*	49
Exochellidae	1	2	2	82.2	5.1	Santonian*	83.5
Flustridae	4	5	101	83.7	8.7	Recent**	0
Haplopomidae	1	4	4	91.6	4.6	Lutetian**	90
Hippoporidridae	0	0	0			Chattian*	23.03
Hippothoidae	2	2	2	50.0	1.7	Coniacian**	85.8
Horneridae	1	1	1	93.7	5.6	Barremian**	121
Lacernidae	0	0	0			Priabonian**	33.7
Lepraliellidae	1	1	1	95.5	7.7	Santonian**	83.5
Lichenoporidae	2	2	2	100.0	6.5	Cenomanian**	93.5
Membraniporidae	2	3	180	65.1	7.1	Priabonian**	33.7
Microporellidae	2	2	147	74.6	3.7	Aquitania**	20.5
Microporidae	0	0	0				
Oncousoeciidae	2	3	3	100.0	7.0	Sinemurian*	199.3
Phidoloporidae	1	3	3	95.6	7.6	Dunian**	61
Plagioeciidae	1	1	1	100.0	6.2	Pliensbachian**	190
Romancheinidae	4	10	10	96.8	6.5	Campanian**	71.3
Schizoporellidae	1	3	22	43.5	6.4	Ypresian**	49
Scrupariidae	1	2	2	100.0	8.0	Maastrichtian*	70
Setosellidae	1	1	1	0.0	0.0	Thanetian**	54.8
Smittinidae	5	7	7	51.8	6.5	Ypresian**	49
Stigmatoechidae	1	1	1	100.0	7.4	Camponian*	83.6
Stomachetosellidae	1	2	2		7.7	Maastrichtian*	70
Stomatoporidae	1	1	1	92.8	7.2	Carnian*	235

(Continued)

Table 3. (Continued)

Family	Genera	Species	n	Mean wt.% calcite	Mean wt.% MgCO ₃ in calcite	FAD stage	Age at top (Ma)
Terviidae	1	1	1	100.0	8.0	Ypresian**	49
Tessaradomidae	1	1	1	100.0	6.4	Danian**	61
Tubuliporidae	2	5	6	99.8	2.8	Campanian**	71.3
Umbonulidae	1	1	1	47.6	6.0	Lutetian**	41.3

First appearance datum (FAD) taken from the following sources

* Taylor, 1993 [71]

**Smith et al., 2006 [27]

*** Taylor et al., 2009 [30]

<https://doi.org/10.1371/journal.pone.0197533.t003>

where this methodology has been applied to mineralogical traits of other taxa include Cairn et al.'s work on cnidarians [23] and recent publications by Smith et al on serpulids [77] and coralline algae [1]. To date no bryozoan mineralogical analyses have taken into account the phylogenetic signal, however, the recent publication of bryozoan phylogenies by Tsyganov-Bodounov et al. [78], Fuchs et al. in 2009 [79] and Waeschenbach et al in 2012 [55,72] allow this concept to be tested for Bryozoa for the first time.

Mineralogical variation within species

Evolutionary/phylogenetic theories help to explain the differences in base mineralogy between taxonomic groups and species, however in all species skeletal variation continues between different specimens within the same species.

The two main controls which are exerted on bryozoan mineralogy within species are summarised by many authors as biological and environmental control [20,27,29,32,80–84], also known as “active” and “passive” control in some publications [85,86]. Biological or “active” control usually refers to factors such as astogeny, the thickening of secondary calcification in older zooids [20,21,43], although it could also be considered to include growth rates [16,18,17], breeding cycles, food availability [19], physiological “wellness” [87] and directed mineralization to confer a competitive advantage [21,88,89]. Environmental control, or “passive” control, suggests that skeletal mineralogy is driven by the seawater from which the bryozoan is forming its skeleton with little or no physiological involvement from the animal itself. The main environmental control usually discussed is temperature [14,15,20,29,80,90], with higher temperatures reported to drive higher Mg-calcite and aragonite deposition. Other environmental factors which have also been shown to influence mineralogy include salinity [47], depth [40,61], aragonite compensation depth (ACD) [91], Mg/Ca ratio in seawater [80,81,88,90,92] and general seawater chemistry [29,87]. Many of these environmental factors, such as temperature, salinity, Mg/Ca ratio and ACD vary with latitude and have resulted in reported correlations between latitude and skeleton mineralogy [20,43,93]. Depth may be an explanatory factor for some of the variation seen within this study as samples were often collected from varying depths. Patterns of variation have been previously shown in bryozoan studies of both Scottish and Arctic species [15,61].

A sub-group of cheilostomes that feature both a wide geographical range and a corresponding wide range in mineralogy are those that have become successful as non-native (alien) species. An example from this study is *Schizoporella japonica* a species which was first found in Scotland in 2011 [94] and has since been found to inhabit marinas and harbours across Scotland and further afield [95]. This species has been reported from Norway to Malaysia and has highly plastic skeletal mineralogy (Table 1). It could be that the wide latitudinal range, and the wide variety in seawater encompassed, explains the correspondingly wide range of wt%

MgCO₃ in calcite in skeletons of this species. Alternatively, it could be that mineralogical plasticity itself enables the species to survive in a greater range of habitats enabling it to increase its distribution and making it well suited to colonising new marine habitats.

A more likely explanation for the variability seen in non-native species, however, is misidentification and cryptic species. An example is *Schizoporella unicornis* which is a non-native species in a number of countries and has been widely reported in bryozoan mineralogical publications for its mineralogical plasticity [27,32,38,47,96,97] with authors using and recommending it as an ideal species for environmental correlations with mineralogy [32] and potential palaeoenvironmental interpretation [27]. Recent publications have highlighted that many records and museum specimens previously identified as *S. unicornis* are actually misidentified *S. japonica* [94,95], *S. errata*, *S. dunkeri* or other *Schizoporella* species [98–100]. Similar studies have been recently conducted on invasive species of *Watersipora* [101,102] and *Bugula* [103] which also identified taxonomic misidentifications and cryptic speciation. With this taxonomic doubt cast on past distribution and mineralogical records of many non-native species, caution should be exercised when interpreting patterns relating mineralogy to distributional range.

Conclusion

This study describes the mineralogy of 156 species of the phylum Bryozoa within Scotland, representing 79% of the species that inhabits Scotland, a greater number and proportion of extant species than any other regional study. Analysis of the skeletal composition of Scottish bryozoans shows that the group is statistically different from the neighbouring Arctic fauna but features a range of mineralogy comparable to other temperate regions around the world.

Analysis of the mineralogical composition of Scottish bryozoans shows that the group reflects both reported patterns in evolutionary/ genetically “pre-programmed” mineralogy superimposed by variation driven by a combination of environment and biological/physiological factors. In such a variable region as Scotland, open to biological and environmental influx from all directions, it is perhaps no surprise that bryozoans reflect this diversity in the wide range of mineralogy they present.

These data add to a growing database of bryozoan mineralogical analyses from around the world, but we also highlight that more study is needed for a better understanding of the influence of genetic/ evolutionary, environmental and biological factors at play in bryozoan mineralogy.

Supporting information

S1 File. Full sample data and results. This file contains two tables: Table A contains sample data; Table B specifies sampling sites and permissions.
(XLSX)

Acknowledgments

The authors would like to thank the NHM for their support with both access to collections and the IAC unit. The authors thank the Marine Alliance for Science and Technology in Scotland (MASTS) and the Marine Environmental Research Group (MERG) at Heriot-Watt University for supporting and funding J.L.’s PhD. J.L. would like to acknowledge support from the UK research council knowledge exchange fellowship [NE/M006566/2] which allowed her to complete this manuscript. The authors would like to thank all those who contributed specimens for analysis, especially Dr Jim Drewery, the Natural History Museum and the National Museum of Scotland.

Author Contributions

Conceptualization: Jennifer Loxton, Joanne S. Porter.

Data curation: Jennifer Loxton.

Formal analysis: Jennifer Loxton, Jens Najorka.

Funding acquisition: Joanne S. Porter.

Investigation: Jennifer Loxton.

Methodology: Jennifer Loxton, Jens Najorka.

Project administration: Jennifer Loxton.

Resources: Mary Spencer Jones, Jens Najorka.

Software: Jens Najorka.

Supervision: Mary Spencer Jones, Joanne S. Porter.

Validation: Mary Spencer Jones, Jens Najorka.

Visualization: Jennifer Loxton, Abigail M. Smith.

Writing – original draft: Jennifer Loxton, Abigail M. Smith.

Writing – review & editing: Mary Spencer Jones, Jens Najorka, Abigail M. Smith, Joanne S. Porter.

References

1. Smith AM, Sutherland JE, Kregting L, Farr TJ, Winter DJ. Phylomineralogy of the coralline red algae: correlation of skeletal mineralogy with molecular phylogeny. *Phytochemistry*. 2012; 81: 97–108. <https://doi.org/10.1016/j.phytochem.2012.06.003> PMID: 22795764
2. Cuif JP, Dauphin Y. The Environment Recording Unit in coral skeletons—a synthesis of structural and chemical evidences for a biochemically driven, stepping-growth process in fibres. *Biogeosciences*. 2005; 2: 61–73. <https://doi.org/10.5194/bg-2-61-2005>
3. Trotter J, Montagna P, McCulloch M, Silenzi S, Reynaud S, Mortimer G, et al. Quantifying the pH “vital effect” in the temperate zooxanthellate coral *Cladocora caespitosa*: Validation of the boron seawater pH proxy. *Earth Planet Sci Lett*. 2011; 303: 163–173. <https://doi.org/10.1016/j.epsl.2011.01.030>
4. Chang VT-C, Williams RJP, Makishima A, Belshaw NS, O’Nions RK. Mg and Ca isotope fractionation during CaCO₃ biomineralisation. *Biochem Biophys Res Commun*. 2004; 323: 79–85. <https://doi.org/10.1016/j.bbrc.2004.08.053> PMID: 15351704
5. Piwoni-Piórewicz A, Kukliński P, Strekopytov S, Humphreys-Williams E, Najorka J, Iglkowska A. Size effect on the mineralogy and chemistry of *Mytilus trossulus* shells from the southern Baltic Sea: implications for environmental monitoring. *Environ Monit Assess*. 2017; 189: 197. <https://doi.org/10.1007/s10661-017-5901-y> PMID: 28361486
6. Keene EC, Evans JS, Estroff L a. Matrix Interactions in Biomineralization: Aragonite Nucleation by an Intrinsically Disordered Nacre Polypeptide, n16N, Associated with a β -Chitin Substrate. *Cryst Growth Des*. 2010; 10: 1383–1389. <https://doi.org/10.1021/cg901389v>
7. Thompson JB, Paloczi GT, Kindt JH, Michenfelder M, Smith BL, Stucky G, et al. Direct observation of the transition from calcite to aragonite growth as induced by abalone shell proteins. *Biophys J*. 2000; 79: 3307–12. [https://doi.org/10.1016/S0006-3495\(00\)76562-3](https://doi.org/10.1016/S0006-3495(00)76562-3) PMID: 11106633
8. Nudelman F, Gotliv B-A, Addadi L, Weiner S. Mollusk shell formation: mapping the distribution of organic matrix components underlying a single aragonitic tablet in nacre. *J Struct Biol*. 2006; 153: 176–87. <https://doi.org/10.1016/j.jsb.2005.09.009> PMID: 16413789
9. Watabe N. Crystal growth of calcium carbonate in biological systems. *J Cryst Growth*. 1974; 24/ 25: 116–122.
10. Hall SR, Taylor PD, Davis S a, Mann S. Electron diffraction studies of the calcareous skeletons of bryozoans. *J Inorg Biochem*. 2002; 88: 410–9. Available: <http://www.ncbi.nlm.nih.gov/pubmed/11897358> PMID: 11897358

11. Fu G, Valiyaveetil S, Wopenka B, Morse DE. CaCO₃ biomineralization: acidic 8-kDa proteins isolated from aragonitic abalone shell nacre can specifically modify calcite crystal morphology. *Biomacromolecules*. 2005; 6: 1289–98. <https://doi.org/10.1021/bm049314v> PMID: 15877344
12. Dauphin Y, Denis A. Structure and composition of the aragonitic crossed lamellar layers in six species of Bivalvia and Gastropoda. *Comp Biochem Physiol A Mol Integr Physiol*. 2000; 126: 367–377. Available: <http://www.ncbi.nlm.nih.gov/pubmed/10964031> PMID: 10964031
13. Tanur AE, Gunari N, Sullan RM a, Kavanagh CJ, Walker GC. Insights into the composition, morphology, and formation of the calcareous shell of the serpulid *Hydroides dianthus*. *J Struct Biol.*; 2010; 169: 145–60. <https://doi.org/10.1016/j.jsb.2009.09.008> PMID: 19766190
14. Loxton J, Kuklinski P, Barnes DKA, Najorka J, Jones MS, Porter JS. Variability of Mg-calcite in Antarctic bryozoan skeletons across spatial scales. *Mar Ecol Prog Ser*. 2014;507. <https://doi.org/10.3354/meps10826>
15. Loxton J, Kuklinski P, Najorka J, Jones MS, Porter JS. Variability in the skeletal mineralogy of temperate bryozoans: The relative influence of environmental and biological factors. *Mar Ecol Prog Ser*. 2014;510. <https://doi.org/10.3354/meps10889>
16. Barnes DKA, Webb KE, Linse K. Growth rate and its variability in erect Antarctic bryozoans. *Polar Biol*. 2007; 30: 1069–1081. <https://doi.org/10.1007/s00300-007-0266-2>
17. Smith AM. Age, growth and carbonate production by erect rigid bryozoans in Antarctica. *Palaeogeogr Palaeoclimatol Palaeoecol*. 2007; 256: 86–98. <https://doi.org/10.1016/j.palaeo.2007.09.007>
18. Kuklinski P, Taylor PD. Are bryozoans adapted for living in the Arctic? Proceedings of the 14th International Bryozoology Association conference, Boone, California. Virginia Museum of Natural History Publications; 2008. pp. 101–110.
19. Bone Y, James N. Bryozoans as carbonate sediment producers on the cool-water Lacedpede shelf, Southern Australia. *Sediment Geol*. 1993; 86: 247–271.
20. Kuklinski P, Taylor PD. Mineralogy of Arctic bryozoan skeletons in a global context. *Facies*. 2009; 55: 489–500. <https://doi.org/10.1007/s10347-009-0179-3>
21. Smith AM, Girvan E. Understanding a bimineralic bryozoan: Skeletal structure and carbonate mineralogy of *Odontionella cyclops* (Foveolariidae: Cheilostomata: Bryozoa) in New Zealand. *Palaeogeogr Palaeoclimatol Palaeoecol*. 2010; 289: 113–122. <https://doi.org/10.1016/j.palaeo.2010.02.022>
22. Peebles BA, Smith AM, Spencer HG. Valve microstructure and phylomineralogy of New Zealand chitons. *J Struct Biol*. 2016; 197: 250–259. <https://doi.org/10.1016/j.jsb.2016.12.002> PMID: 27940093
23. Cairns SD, Macintyre IANG. Phylogenetic Implications of Calcium Carbonate Mineralogy in the Stylasteridae Cnidaria: Hydrozoa). *Sediment Geol*. 1992; 7: 96–107.
24. Smith AM, Clark DE. Skeletal Carbonate Mineralogy of Bryozoans From Chile: an Independent Check of Phylogenetic Patterns. *Palaios*. 2010; 25: 229–233. <https://doi.org/10.2110/palo.2009.p09-106r>
25. Chilanger G V. Dependence on temperature of Ca/Mg ratio of skeletal structures in organisms and direct chemical precipitates of sea water. *Bull South Calif Acad Sci*. 1962; 61: 45–61.
26. Katz A. The interaction of magnesium with calcite during crystal growth at 25–90C and one atmosphere. *Geochim Cosmochim Acta*. 1973; 37: 1563–1586.
27. Smith AM, Key MM, Gordon DP. Skeletal mineralogy of bryozoans: Taxonomic and temporal patterns. *Earth-Science Rev*. 2006; 78: 287–306. <https://doi.org/10.1016/j.earscirev.2006.06.001>
28. Krzeminska M, Kuklinski P, Najorka J, Iglukowska A. Skeletal mineralogy patterns of Antarctic Bryozoa. *J Geol*. 2016;124.
29. Lombardi C, Cocito S, Hiscock K, Occhipinti-Ambrogi A, Setti M, Taylor PD, et al. Influence of seawater temperature on growth bands, mineralogy and carbonate production in a bioconstructional bryozoan. *Facies*. 2008; 54: 333–342. <https://doi.org/10.1007/s10347-008-0143-7>
30. Taylor PD, James NP, Bone Y, Kuklinski P, Kyser TK. Evolving mineralogy of cheilostome bryozoans. *Palaios*. 2009; 24: 440–452. <https://doi.org/10.2110/palo.2008.p08>
31. Swezey DS, Bean JR, Ninokawa AT, Hill TM, Gaylord B, Sanford E. Interactive effects of temperature, food and skeletal mineralogy mediate biological responses to ocean acidification in a widely distributed bryozoan. *Proc R Soc B Biol Sci*. 2017;284. <https://doi.org/10.1098/rspb.2016.2349> PMID: 28424343
32. Lowenstam HA. Environmental relations of modification compositions of certain carbonate secreting marine invertebrates. *Proc Natl Acad Sci U S A*. 1954; 40: 39–48. Available: <http://www.pubmedcentral.nih.gov/articlerender.fcgi?artid=527935&tool=pmcentrez&rendertype=abstract> PMID: 16589423
33. Dowsett HJ, Haywood AM, Valdes PJ, Robinson MM, Lunt DJ, Hill DJ, et al. Sea surface temperatures of the mid-Piacenzian Warm Period: A comparison of PRISM3 and HadCM3. *Palaeogeogr Palaeoclimatol Palaeoecol*. Elsevier B.V.; 2011; 309: 83–91. <https://doi.org/10.1016/j.palaeo.2011.03.016>

34. Cohen AL, Branch GM. Environmentally controlled variation in the structure and mineralogy of *Patella granularis* shells from the coast of southern Africa: implications for palaeotemperature assessments. *Palaeogeogr Palaeoclimatol Palaeoecol*. 1992; 91: 49–57.
35. Fabry VJ, McClintock JB, Mathis JT, Grebmeier JM. Ocean acidification at high latitudes: The bellwether. *Oceanography*. 2009; 22: 160–171.
36. Siesser WG. Carbonate mineralogy of Bryozoans and other selected South African organisms. *S Afr J Sci*. 1972; 71–74.
37. Agegian CR, Mackenzie FT. Calcareous organisms and sediment mineralogy on a mid-depth bank in the Hawaiian Archipelago. *Pacific Sci*. 1989; 43: 56–66.
38. Poluzzi A, Sartori R. Carbonate Mineralogy of some bryozoa from Talbot Shoal. *G di Geol*. 1973; 39: 11–15.
39. Walther J. Die gesteinsbildenden Kalkagen des Golfes von Neapel und die Entstehung structuloser Kalke. *Dtsch Geol Gesellschafen Zeitschrift*. 1885; 37: 329–357.
40. Taylor PD, Tan Shau-Hwai A, Kudryavstev AB, Schopf JW. Carbonate mineralogy of a tropical bryozoan biota and its vulnerability to ocean acidification. *Mar Biol Res*. 2016; 12: 776–780. <https://doi.org/10.1080/17451000.2016.1203951>
41. Poluzzi A, Sartori R. Report on the carbonate mineralogy of Bryozoa. *Docum Lab Geol Fac Sci Lyon*. 1974; 3: 193–210.
42. Borisenko Y., Gontar V. Biogeochemistry of skeletons of coldwater Bryozoa. *Biol Morya* 1. 1991; 80–90 (in Russian).
43. Smith AM, Nelson C, Spencer H. Skeletal carbonate mineralogy of New Zealand bryozoans. *Mar Geol*. 1998; 151: 27–46. [https://doi.org/10.1016/S0025-3227\(98\)00055-3](https://doi.org/10.1016/S0025-3227(98)00055-3)
44. Scottish Government. Chpter 2: Physical characteristics. *Scotland's Marine Atlas: Information for the National Marine Plan*. 2011.
45. Rouse S. The Distribution and Biodiversity of Bryozoans in Scotland. M.Sc. Thesis; Heriot-Watt University; 2010.
46. Clarke FW, Wheeler WC. The inorganic constituents of marine invertebrates. *United States Geol Surv Prof Pap*. 1917; 102: 1–56.
47. Clarke FW, Wheeler WC. The inorganic constituents of marine invertebrates. *United States Geol Surv Prof Pap*. 1922; 124: 34–36.
48. Schopf TJM, Allan JR. Phylum Ectoprocta, Order Cheilostomata: microprobe analysis of calcium, magnesium, strontium and phosphorus in skeletons. *Science* (80-). 1970; 169: 280–282.
49. Hayward PJ, Ryland JS. Cheilostomatous Bryozoa, Pt 1: Aeteoidea-Cribrillinoidea. 2nd ed. Shrewsbury: Field Studies Council; 1998.
50. Hayward PJ, Ryland JS. Cheilostomatous Bryozoa, Pt 2: Hippothooidea-Celleporoidea. 2nd ed. Shrewsbury: Field Studies Council; 1999.
51. Appeltans W, Bouchet P, Boxshall GA, De Broyer C, de Voogd NJ, Gordon DP, et al. World Register of Marine Species (WoRMS) [Internet]. 2012 [cited 10 Jan 2013]. Available: www.marinespecies.org
52. Loxton J, Najorka J, Humphreys-Williams E, Kuklinski P, Smith AM, Porter JS, et al. The forgotten variable: Impact of cleaning on the skeletal composition of a marine invertebrate. *Chem Geol*. 2017; <https://doi.org/10.1016/j.chemgeo.2017.10.022>
53. Blanton TN, Huang TC, Toraya H, Hubbard CR, Robie SB, Louer D, et al. JCPDS—International Centre for Diffraction Data round robin study of silver behenate. A possible low-angle X-ray diffraction calibration standard. *Powder Diffr*. 1995; 10: 91–95.
54. Blanton TN, Barnes CL, Lelental M. Preparation of silver behenate coatings to provide low to mid-angle diffraction calibration. *J Appl Crystallogr*. 2000; 33: 172–173.
55. Waeschenbach A, Taylor PD, Littlewood DTJ. A molecular phylogeny of bryozoans. *Mol Phylogenet Evol*. 2012; 62: 718–35. <https://doi.org/10.1016/j.ympev.2011.11.011> PMID: 22126903
56. R Core Team. R: A Language and Environment for Statistical Computing [Internet]. Vienna, Austria: R Foundation for Statistical Computing; 2013. Available: <http://www.r-project.org>
57. Kembel S. An introduction to the picante package Installing picante Data formats in picante. 2010; 1–16.
58. Kembel SW, Cowan P, Helmus MR, Cornwell WK, Morlon H, Ackerley DD, et al. Picante: R tools for integrating phylogenies and ecology. *Bioinformatics*; 2010. pp. 1463–1464. <https://doi.org/10.1093/bioinformatics/btq166> PMID: 20395285
59. Kembel ASW, Ackerly DD, Blomberg SP, Cornwell WK, Cowan PD, Hel- MR, et al. Package “picante.” 2013;

60. Blomberg SP, Garland T, Ives AR. Testing for phylogenetic signal in comparative data: behavioral traits are more labile. *Evolution* (N Y). 2003; 57: 717–745.
61. Borszcz T, Kukliński P, Taylor PD. Patterns of magnesium content in Arctic bryozoan skeletons along a depth gradient. *Polar Biol.* 2013; 36: 193–200. <https://doi.org/10.1007/s00300-012-1250-z>
62. Barnes DK a., Dick MH. Overgrowth competition in encrusting bryozoan assemblages of the intertidal and infralittoral zones of Alaska. *Mar Biol.* 2000; 136: 813–822. <https://doi.org/10.1007/s002270000253>
63. Barnes D, Arnold R. A growth cline in encrusting benthos along a latitudinal gradient within Antarctic waters. *Mar Ecol Prog Ser.* 2001; 210: 85–91. <https://doi.org/10.3354/meps210085>
64. Barnes D, Webb K, Linse K. Slow growth of Antarctic bryozoans increases over 20 years and is anomalously high in 2003. *Mar Ecol Prog Ser.* 2006; 314: 187–195. <https://doi.org/10.3354/meps314187>
65. Peck LS, Barnes DK a. Metabolic flexibility: the key to long-term evolutionary success in Bryozoa? *Proc Biol Sci.* 2004; 271 Suppl: S18–21. <https://doi.org/10.1098/rsbl.2003.0053>
66. Brockington S, Clarke a. The relative influence of temperature and food on the metabolism of a marine invertebrate. *J Exp Mar Bio Ecol.* 2001; 258: 87–99. Available: <http://www.ncbi.nlm.nih.gov/pubmed/11239627> PMID: 11239627
67. Burton EA, Walter LM. Relative growth rates and compositions of aragonite and Mg calcite in seawater: effects of temperature and sulfate. *Geol Soc Am.* 1985; 17: 536.
68. Chierici M, Fransson A. Calcium carbonate saturation in the surface water of the Arctic Ocean: undersaturation in freshwater influenced shelves. *Biogeosciences.* 2009; 6: 2421–2431. <https://doi.org/10.5194/bg-6-2421-2009>
69. Dunton K. Arctic biogeography: the paradox of the marine benthic fauna and flora. *Trends Ecol Evol.* 1992; 7: 183–189. [https://doi.org/10.1016/0169-5347\(92\)90070-R](https://doi.org/10.1016/0169-5347(92)90070-R) PMID: 21236004
70. Boardman R., Cheetham A. Phylum Bryozoa. In: Boardman R., Cheetham A., Rowell A., editors. *Fossil Invertebrates*. Palo Alto: Blackwell; 1987. pp. 497–549.
71. Taylor PD. Bryozoa. In: Benton MJ, editor. *The Fossil Record 2*. Chapman & Hall, London; 1993. pp. 465–489.
72. Waeschenbach A, Cox CJ, Littlewood DTJ, Porter JS, Taylor PD. Molecular Phylogenetics and Evolution First molecular estimate of cyclostome bryozoan phylogeny confirms extensive homoplasy among skeletal characters used in traditional taxonomy. *Mol Phylogenet Evol.* 2009; 52: 241–251. <https://doi.org/10.1016/j.ympev.2009.02.002> PMID: 19236933
73. Grafen A. The phylogenetic regression. *Philos Trans R Soc B Biol Sci.* 1989; 326: 119–157.
74. Felsenstein J. Phylogenies and the comparative method. *Am Nat.* 1985; 125: 1–15.
75. Taylor, 1990, Paleontology—Bioimmuration review.pdf.
76. Hoff PD. *A First Course in Bayesian Statistical Methods*. New York: Springer Science and Business Media; 2009.
77. Smith AM, Riedi MA, Winter DJ. Temperate reefs in a changing ocean: skeletal carbonate mineralogy of serpulids. *Mar Biol.* 2013; <https://doi.org/10.1007/s00227-013-2210-z>
78. Tsyganov-Bodounov A, Hayward PJ, Porter JS, Skibinski DOF. Bayesian phylogenetics of Bryozoa. *Mol Phylogenet Evol.* 2009; 52: 904–10. <https://doi.org/10.1016/j.ympev.2009.05.010> PMID: 19460450
79. Fuchs J, Obst M, Sundberg P. The first comprehensive molecular phylogeny of Bryozoa (Ectoprocta) based on combined analyses of nuclear and mitochondrial genes. *Mol Phylogenet Evol.* 2009; 52: 225–233. <https://doi.org/10.1016/j.ympev.2009.01.021> PMID: 19475710
80. Davis KJ. The Role of Mg²⁺ as an Impurity in Calcite Growth. *Science.* 2000; 290: 1134–1137. <https://doi.org/10.1126/science.290.5494.1134> PMID: 11073446
81. Ries JB. Aragonite production in calcite seas: effect of seawater Mg/Ca ratio on the calcification and growth of the calcareous alga *Penicillus capitatus*. *Paleobiology.* 2005; 285: 323–458. [https://doi.org/10.1666/0094-8373\(2005\)031\[0445:APICSE\]2.0.CO;2](https://doi.org/10.1666/0094-8373(2005)031[0445:APICSE]2.0.CO;2)
82. Lowenstam HA. Coexisting calcites and aragonites from skeletal carbonates of marine organisms and their strontium and magnesium contents. In: Miyake Y, Koyama T, editors. *Recent researches in the field of hydrosphere, atmosphere and nuclear chemistry*. Maruzen, Tokyo; 1964. pp. 373–404.
83. Smith AM, Key MM. Controls, variation, and a record of climate change in detailed stable isotope record in a single bryozoan skeleton. *Quat Res.* 2004; 61: 123–133. <https://doi.org/10.1016/j.yqres.2003.11.001>
84. Steger KK, Smith AM. Carbonate mineralogy of free-living bryozoans Bryozoa: Otionellidae), Otago shelf, southern New Zealand. *Palaeogeogr Palaeoclimatol Palaeoecol.* 2005; 218: 195–203. <https://doi.org/10.1016/j.palaeo.2004.12.013>

85. Schafer P, Bader B. Geochemical composition and variability in the skeleton of the bryozoan *Cellaria sinuosa* (Hassall): Biological versus environmental control. In: Hageman SJ, Key MM, Winston JE, editors. Proceedings of the 14th International Bryozoology Association conference. Martinsville, Virginia: Virginia Museum of Natural History Publications; 2008. pp. 269–279.
86. Bader B, Schafer P. Impact of environmental seasonality on stable isotope composition of skeletons of the temperate bryozoan *Cellaria sinuosa*. *Palaeogeogr Palaeoclimatol Palaeoecol*. 2005; 226: 58–71. <https://doi.org/10.1016/j.palaeo.2005.05.007>
87. Stanley S, Hardie L. Secular oscillations in the carbonate mineralogy of reef-building and sediment-producing organisms driven by tectonically forced shifts in seawater chemistry. *Palaeogeogr Palaeoclimatol Palaeoecol*. 1998; 144: 3–19. [https://doi.org/10.1016/S0031-0182\(98\)00109-6](https://doi.org/10.1016/S0031-0182(98)00109-6)
88. Loste E. The role of magnesium in stabilising amorphous calcium carbonate and controlling calcite morphologies. *J Cryst Growth*. 2003; 254: 206–218. [https://doi.org/10.1016/S0022-0248\(03\)01153-9](https://doi.org/10.1016/S0022-0248(03)01153-9)
89. Borzęcka-Prokop B, Weselucha-Birczyńska a., Koszowska E. MicroRaman, PXRD, EDS and microscopic investigation of magnesium calcite biomineral phases. The case of sea urchin biominerals. *J Mol Struct*. 2007; 828: 80–90. <https://doi.org/10.1016/j.molstruc.2006.05.040>
90. Morse JW. Influences of temperature and Mg: Ca ratio on CaCO₃ precipitates from seawater. *Geology*. 1997; 2: 85–87.
91. Berger WH. Deep-sea carbonate: pteropod distribution and the aragonite compensation depth. *Deep Res*. 1978; 25: 447–452.
92. Astilleros JM, Fernández-díaz L, Putnis A. The role of magnesium in the growth of calcite: An AFM study. *Chem Geol*. 2010; 271: 52–58. <https://doi.org/10.1016/j.chemgeo.2009.12.011>
93. Taylor PD. Seawater Chemistry, Biomineralization and the Fossil Record of Calcareous Organisms. In: Okada H, Mawatari SF, Suzuki N, Gautam P, editors. Origin and Evolution of Natural Diversity, Proceedings of International Symposium “The Origin and Evolution of Natural Diversity.” 2007. pp. 21–29.
94. Ryland JS, Holt R, Loxton J, Spencer ME, Porter JS. First occurrence of the non-native bryozoan *Schizoporella japonica* Ortmann 1890) in Western Europe. *Zootaxa*. 2014;
95. Loxton J, Wood CA, Bishop JDD, Porter JS, Spencer Jones M, Nall CR. Distribution of the invasive bryozoan *Schizoporella japonica* in Great Britain and Ireland and a review of its European distribution. *Biol Invasions*. 2017; 19: 2225–2235. <https://doi.org/10.1007/s10530-017-1440-2> PMID: 28798542
96. Schopf TJM, Manheim FT. Chemical composition of Ectoprocta (Bryozoa). *J Paleontol*. 1967; 41: 1197–1225.
97. Carver RE, Rucker JB. A survey of the carbonate mineralogy of cheilostome bryozoa. *J Paleontol*. 1969; 43: 791–799.
98. Hayward PJ, Ryland JS. The British species of *Schizoporella* Bryozoa: Cheilostomatida). *J Zool*. 1995; 1: 37–47.
99. Tompsett S, Porter JS, Taylor PD. Taxonomy of the fouling cheilostome bryozoans *Schizoporella unicornis* (Johnston) and *Schizoporella errata* (Waters). *J Nat Hist*. 2009; 43: 2227–2243. <https://doi.org/10.1080/00222930903090140>
100. Clarke Murray C, Pakhomov E a., Therriault TW. Recreational boating: a large unregulated vector transporting marine invasive species. *Divers Distrib*. 2011; 17: 1161–1172. <https://doi.org/10.1111/j.1472-4642.2011.00798.x>
101. Ryland JS. Recent discoveries of alien *Watersipora* (Bryozoa) in Western Europe. *Zootaxa*. 2009; 59: 43–59.
102. Mackie JA, Darling JA, Geller JB. Ecology of cryptic invasions: latitudinal segregation among *Watersipora* (Bryozoa) species. *Sci Rep*. 2012; 2: 871. <https://doi.org/10.1038/srep00871> PMID: 23213354
103. Ryland J, Bishop J, De Blauwe H, El Nagar A, Minchin D, Wood C, et al. Alien species of *Bugula* (Bryozoa) along the Atlantic coasts of Europe. *Aquat Invasions*. 2011; 6: 17–31. <https://doi.org/10.3391/ai.2011.6.1.03>