

# Simple measures of ozone depletion in the polar stratosphere

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**Abstract.** We investigate the extent to which quantities that are based on total column ozone are applicable as measures of ozone loss in the polar vortices. Such quantities have been used frequently in ozone assessments by the World Meteorological Organization (WMO) and also to assess the performance of chemistry-climate models. The most commonly considered quantities are March and October mean column ozone poleward of geometric latitude 63° and the spring minimum of daily total ozone minima poleward of a given latitude. Particularly in the Arctic, the former measure is affected by vortex variability and vortex break-up in spring. The minimum of daily total ozone minima poleward of a particular latitude is debatable, insofar as it relies on one single measurement or model grid point. We find that, for Arctic conditions, this minimum value often occurs in air *outside* the polar vortex, both in the observations and in a chemistry-climate model. Neither of the two measures shows a good correlation with chemical ozone loss in the vortex deduced from observations. We recommend that the minimum of daily minima should no longer be used when comparing polar ozone loss in observations and models. As an alternative to the March and October mean column polar ozone we suggest considering the minimum of daily average total ozone poleward of 63° equivalent latitude in spring (except for winters with an early vortex break-up). Such a definition both obviates relying on one single data point and reduces the impact of year-to-year variability in the Arctic vortex break-up on ozone loss measures. Further, this measure shows a reasonable correlation ( $r = -0.75$ ) with observed chemical ozone loss. Nonetheless, simple measures of polar ozone loss must be used with caution; if possible, it is preferable to use more sophisticated measures that include additional information to disentangle the impact of transport and chemistry on ozone.

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## 1 Introduction

Since the early eighties, substantial chemical ozone loss has occurred during winter and spring each year in the Antarctic (e.g., Jones and Shanklin, 1995; Tilmes et al., 2006; Huck et al., 2007; WMO, 2007); substantial chemical loss of ozone has likewise been reported for recent cold Arctic winters (e.g., Manney et al., 2003; Tilmes et al., 2004; WMO, 2007). In the very cold Arctic winter of 2004/05, the chemical loss of ozone came closer to Antarctic values (Manney et al., 2006; Rex et al., 2006; Tilmes et al., 2006; von Hobe et al., 2006; Jin et al., 2006). Sophisticated methods have been developed to separate dynamically induced changes of polar ozone from chemical ozone depletion (e.g., Proffitt et al., 1990; Manney et al., 1994; Müller et al., 1996; Rex et al., 2002; Harris et al., 2002; Tilmes et al., 2004; Christensen et al., 2005; Goutail et al., 2005, and references therein). Furthermore, it has been shown (Rex et al., 2004; Tilmes et al., 2004, 2006) that the deduced chemical ozone loss is closely related to the potential for the formation of polar stratospheric clouds (PSCs).

However, these methods rely on rather comprehensive data sets that are not available for all winters of interest, particularly those before the 1990s. Similarly, the necessary information for applying such sophisticated methods to simulations conducted with chemistry-climate models (CCMs) is often not available for archived model output (e.g., Austin et al., 2003; Eyring et al., 2006; Lemmen et al., 2006b).

Consequently, measures of chemical polar ozone depletion are commonly employed that rely solely on total column ozone data (e.g., Newman et al., 2004; Bodeker et al., 2005; Huck et al., 2007). The two most frequently employed measures are the average total column ozone in spring (March, NH and October, SH) poleward of 63° and the minimum of daily total column ozone minima over the polar cap (e.g., Newman et al., 1997; Müller, 2003; Austin et al., 2003; WMO, 2003; Shepherd, 2003; Eyring et al., 2006, 2007;

WMO, 2007). For assessments of the World Meteorological Organization (WMO) and of the Intergovernmental Panel on Climate Change (IPCC), the average March total column ozone poleward of 63° N (Newman et al., 1997) has been regularly used (WMO, 1999, 2003; IPCC/TEAP, 2005; WMO, 2007). The value of 63° N was chosen by Newman et al. (1997), because "... the area poleward of 63° N is [...] large enough to contain the polar vortex, yet small enough to not be dominated by mid-latitude air masses in the lower stratosphere". Obviously, this condition will be better fulfilled in some Arctic winters than in others depending on the varying size and shape of the polar vortex (Waugh and Randel, 1999; Karpetchko et al., 2005). And there are winters when the Arctic vortex breaks up before March (WMO, 2007).

The minimum of daily total ozone minima poleward of a particular latitude has been used as a standard to assess the performance of CCMs (e.g., Austin et al., 2003; WMO, 2003; Eyring et al., 2006, 2007; WMO, 2007). This measure is problematic, insofar as it relies on one single measurement or on one single model grid point. Knudsen (2002) criticised the use of this parameter and pointed out that the minimum ozone in the Arctic is frequently caused by high pressure systems rather than by chemical ozone destruction. He therefore argued that, for example, the March mean total column ozone poleward of 63° N would be a more suitable measure for the development of polar ozone.

Here, we scrutinise the information that can be deduced from simple (total column ozone based) measures of chemical ozone loss and investigate circumstances where mis- and over-interpretations might occur. Our study uses a combined data set of satellite-based total column ozone measurements (Bodeker et al., 2005) from 1978 to 2007 (Sect. 2). Monthly mean column ozone poleward of 63° in spring (Newman et al., 1997; WMO, 2007) shows similar year-to-year variability as vortex mean ozone, but neither of these quantities shows a close relation to meteorological quantities that describe the potential for polar heterogeneous chlorine activation and thus ozone loss (Sect. 3). Furthermore, we argue here that the minimum of daily total ozone minima poleward of a particular latitude should no longer be used when comparing polar ozone loss between observations and CCMs because this minimum value often occurs in air *outside* of the polar vortex (Sect. 3). As a better alternative for a total ozone based measure, we suggest using the minimum of daily average total ozone poleward of a particular equivalent latitude in March or October. However, whenever possible, more sophisticated measures of chemical polar ozone loss should be used that allow the impact of transport and chemistry on ozone to be disentangled.

## 2 The total column ozone data set

The data set used here is version 2.7 of the NIWA combined ozone database which provides daily total column ozone

fields from November 1978 to March 2007. It is based on version 8 TOMS (Total Ozone Mapping Spectrometer) retrievals from four different satellites (Nimbus 7, Meteor 3, ADEOS, and Earth Probe), total column ozone retrievals from the application of the TOMS version 8 retrieval algorithm to OMI (Ozone Monitoring Instrument) measurements, version 4.0 GOME (Global Ozone Monitoring Experiment) ozone retrievals, GOME total column ozone fields from the KNMI TOGOMI algorithm, and version 8 SBUV (Solar Backscatter Ultra-Violet) retrievals from four different satellites (Nimbus 7, NOAA 9, NOAA 11, and NOAA16). It updates the data set described in Bodeker et al. (2005), extending it to the end of March 2007, and implementing a number of improvements including:

- Data from both the Dobson spectrophotometer and Brewer spectrometer global networks are now used to remove offsets and drifts between the various satellite-based total column ozone data sets used to construct the database.
- Data from the Ozone Monitoring Instrument (OMI) flown onboard NASA's Aura satellite since September 2004 are now used. Differences between OMI overpass measurements and ground-based measurements are small ( $-2.8 \pm 5$  DU) and the OMI grid data are corrected before they are combined with the other data sources, which are also corrected as in Bodeker et al. (2005).
- A better correction function for the Earth Probe TOMS ozone measurements has been derived to account for non-linearities in the drift between the Earth Probe TOMS measurements and ground-based measurements in recent years. Since September 2004 these data have only been used when OMI data are unavailable.
- For the DLR GOME retrieval, a new data version (4.0) was used and a better screening of anomalous ozone measurements at high solar zenith angles has been implemented. As the GOME 4.0 retrieval shows no bias against ground-based instruments (Balis et al., 2007), no corrections to these data were applied.
- In previous versions of the combined database, data from only one satellite-based instrument were used on any given day. In this version of the database, data from all satellite-based instruments are considered sequentially to fill each 1.25° longitude by 1.0° latitude grid cell with a priority of Nimbus 7 TOMS, Meteor 3 TOMS, OMI, Earth Probe TOMS, ADEOS TOMS, the KNMI and then DLR GOME ozone retrievals, and then the 4 SBUV data sets. For example, if most, but not all, of the globe is covered by Nimbus 7 TOMS and Meteor 3 TOMS data are available to fill the gap, then these data are used for this. Both data sets, as before, are first corrected for offsets and drifts against the ground-based

measurements. In previous versions of the database however, only the Nimbus 7 data would have been used for that day and the gap in the data would remain.

### 3 Results

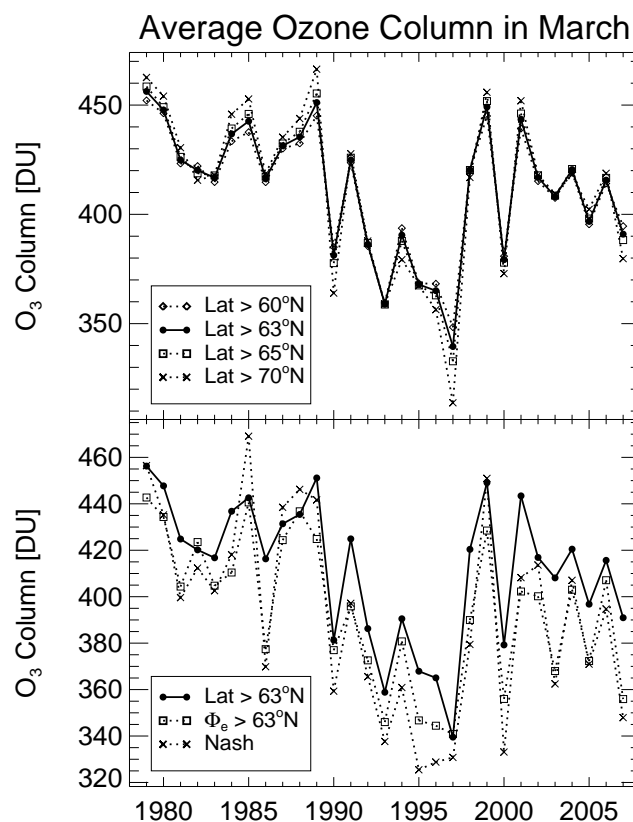
#### 3.1 Mean total column ozone over the poles in early spring

To test the sensitivity of the calculated Arctic mean ozone to the selected latitude limit, the March mean Arctic total column ozone for a latitude boundary at  $63^\circ\text{N}$ , as originally chosen by Newman et al. (1997), is compared with the mean for boundaries at  $60^\circ\text{N}$ ,  $65^\circ\text{N}$ , and  $70^\circ\text{N}$  in Fig. 1 (top panel). Note that in some winters (1987, 1999, 2001, and 2006; WMO, 2007), Santee, 2008<sup>1</sup>, the polar vortex broke up in February; for those winters, the March mean total column ozone hardly provides any information on polar ozone loss.

Clearly, the monthly spatial means are not very sensitive to the exact choice of the threshold latitude. A very similar picture is found for the average total ozone column in April (see electronic supplement <http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip>). For the eighties, and in the warmer winters during the nineties, choosing a more poleward threshold latitude yields greater average ozone columns. The opposite (a smaller ozone column for a more poleward threshold latitude) is observed for the cold Arctic winters 1994/1995, 1995/1996, and 1999/2000 when substantial chemical ozone loss was observed (e.g., Manney et al., 2003; Tilmes et al., 2004). The strongest effect arising from changing the threshold latitude is found for the winter 1996/1997, a winter where the March mean column ozone showed an unusually strong meridional gradient (Newman et al., 1997).

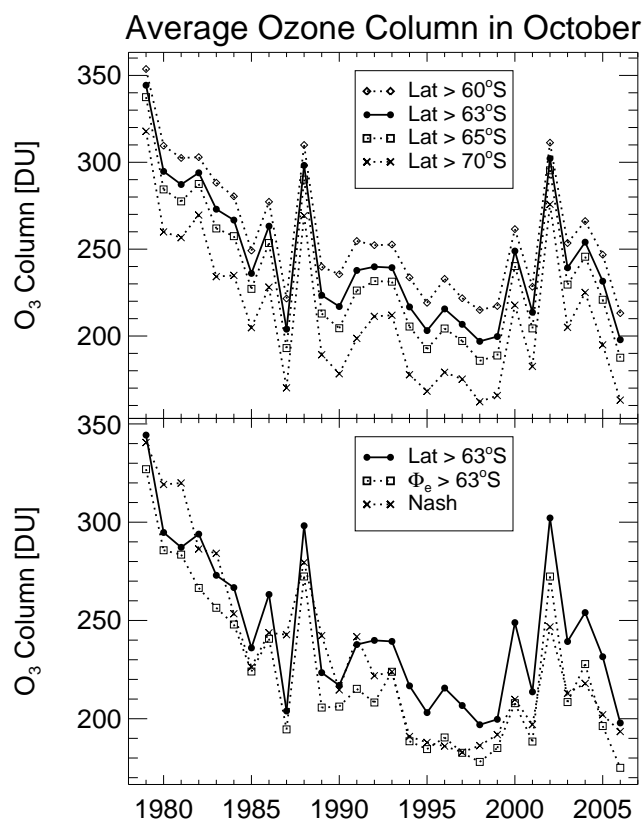
Using equivalent latitude ( $\Phi_e$ ) rather than geographic latitude as an estimate of the boundary of the vortex (e.g., Butchart and Remsberg, 1986; Lary et al., 1995) should lead to a better demarcation between polar and mid-latitude air. Therefore, in Fig. 1 (bottom panel), the total column ozone average for March in the Arctic is shown using both the latitude and the equivalent latitude poleward of  $63^\circ\text{N}$  as the threshold. Here and throughout the paper equivalent latitude and other diagnostics of the vortex edge are evaluated on the 475 K potential temperature level. Overall the two averages show a similar inter-annual pattern. When equivalent latitude is used as the threshold, Arctic winters with a polar centric vortex (e.g., 1996/1997) show little change, while winters in the nineties with substantial chemical ozone loss within a perturbed vortex show lower averages. This result is

<sup>1</sup>Santee, M. L., MacKenzie, I. A., Manney, G. L., Chipperfield, M. P., Bernath, P. F., Walker, K. A., Boone, C. D., Froidevaux, L., Livesey, N. J., and Waters, J. W.: A study of stratospheric chlorine partitioning based on new satellite measurements and modeling, *J. Geophys. Res.*, submitted, 2008.



**Fig. 1.** Top panel: March mean Arctic total column ozone averaged poleward of  $60^\circ\text{N}$ ,  $63^\circ\text{N}$ ,  $65^\circ\text{N}$ , and  $70^\circ\text{N}$ . Bottom panel: March mean Arctic total column ozone averaged poleward of  $63^\circ\text{N}$  compared with calculations using the equivalent latitude of  $\Phi_e=63^\circ\text{N}$  and the Nash criterion (applied on the 475 K potential temperature surface) as vortex edge definitions. All averages are area-weighted averages.

not surprising because, when using equivalent latitude as the threshold, the average ozone is less likely to be influenced by air from outside the vortex where, in winters with substantial ozone loss, ozone is higher. An estimate of the location of the vortex boundary, i.e. the location of the strongest barrier to meridional transport in the polar region which is based on fluid-dynamical theory, is the maximum gradient in potential vorticity in equivalent latitude constrained by the horizontal wind velocity (Nash et al., 1996). The edge of the vortex defined in this way agrees reasonably well with observed strong tracer gradients at the vortex boundary (e.g., Greenblatt et al., 2002; Müller and Günther, 2003). Because the area of the polar vortex varies substantially from year to year in the Arctic (e.g., Karpetchko et al., 2005; Tilmes et al., 2006), a constant equivalent latitude cannot provide an accurate estimate of the vortex area. Thus, because chemical ozone loss is confined to within the vortex boundary, it is expected that the average total column ozone poleward of the vortex edge as defined by Nash et al. (1996) should show lower values than an average based on geometric or constant equivalent latitude. This



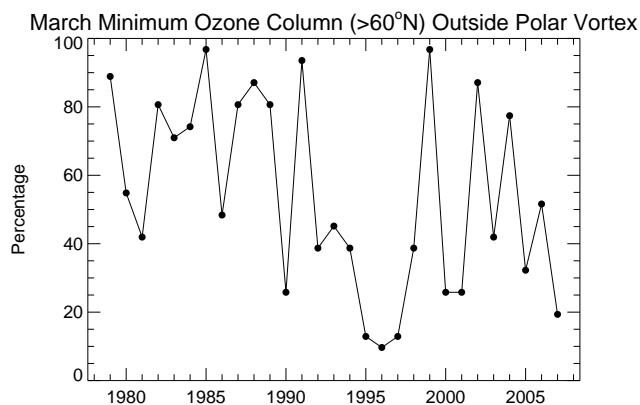
**Fig. 2.** Top panel: October mean Antarctic total column ozone averaged poleward of 60° S, 63° S, 65° S, and 70° S. Bottom panel: October mean Antarctic total column ozone averaged poleward of 63° S compared with calculations using the equivalent latitude  $\Phi_e=63^\circ$  S and the Nash criterion (applied on the 475 K potential temperature surface) as vortex edge definitions.

is indeed borne out by the analysis, with the exception of the warm winters of 1978/1979, 1984/1985, 1987/1988, and 1998/1999 (Fig. 1, bottom panel). All four of these Arctic winters show a very low PSC formation potential and thus a very small potential for chemical ozone destruction (Rex et al., 2004; Tilmes et al., 2006). For all other winters, the polar average total column ozone using a potential vorticity gradient based threshold is lower than using any other threshold and the difference is particularly pronounced for winters showing strong ozone depletion. The sole exception is the winter of 1996/1997 that showed an unusually inhomogeneous ozone distribution within the Arctic vortex with a particularly low ozone column in the vortex core (e.g., Newman et al., 1997; Manney et al., 1997; McKenna et al., 2002). The Arctic vortex, typically, is strongly eroded in April and is therefore smaller than the area encompassed by the 63° N equivalent latitude contour (e.g., Waugh and Randel, 1999). Consequently, column ozone in April averaged over the vortex area as deduced using the Nash et al. (1996) criterion is expected to be lower than the average column ozone poleward of 63° N equivalent latitude, as is the case here

(see electronic supplement, <http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip>). Because the Arctic vortex in April is unlikely to be polar-centred, computing a polar average column ozone using geometric latitude rather than equivalent latitude as the threshold should lead to even more ozone-rich mid-latitude air masses being included in the average, and thus should lead to even greater averages as is indeed the case (see electronic supplement, <http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip>).

Although originally designed for the Arctic, the concept of the average total ozone column poleward of a threshold latitude of 63° has been extended to the Antarctic, where October mean values are commonly considered (e.g., WMO, 1999; IPCC/TEAP, 2005; WMO, 2007). In the Antarctic, the choice of the threshold latitude has a much stronger impact on the October average polar column ozone compared to the Arctic; the difference between a threshold of 60° S and 70° S is noticeable in every Antarctic winter analysed here and can be as large as  $\sim 50$  DU (Fig. 2, top panel). The pattern of inter-annual change in the October mean ozone column, however, is not strongly affected by the choice of the threshold latitude. These statements are likewise valid for the average total ozone column in September (see electronic supplement, <http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip>). As for the Arctic, the sensitivity of Antarctic average total column ozone to the latitude poleward of which the average is calculated depends on the steepness of the meridional gradients in ozone across the vortex edge. For the latitude range considered in Fig. 2, the ozone gradients are steeper in the Antarctic than the Arctic (see e.g. Fig. 5 from Brunner et al. 2006) and therefore Antarctic mean total column ozone is expected to be more sensitive to the spatial limiting latitude.

In the Antarctic, the 63° S equivalent latitude contour is often located within the polar vortex in early spring (e.g., Bodeker et al., 2001, 2002). As a result, the difference between the average column ozone poleward of the Nash-defined vortex edge and the average poleward of 63° S equivalent latitude (Fig. 2, bottom panel) is smaller than in the Arctic. Using the Nash vortex edge as the limiting contour for the averaging includes air masses towards the vortex edge that are not included if a threshold of 63° S equivalent latitude is used. Because total column ozone in the Antarctic vortex increases towards the vortex boundary (Bodeker et al., 2002), using the Nash criterion will generally lead to greater polar total ozone averages (Fig. 2, bottom panel). The only obvious exception to this observation in the October time series analysed here is the winter of 2002. In this winter, at the end of September, a sudden stratospheric warming occurred in the Antarctic vortex, leading to a much smaller and weaker vortex than usual, reminiscent of Arctic conditions (e.g., Krüger et al., 2005; Newman and Nash, 2006).



**Fig. 3.** The fraction (in percent) of the daily total column ozone minima in the region  $60^{\circ}\text{N}$ – $90^{\circ}\text{N}$  occurring outside of the polar vortex in March. The polar vortex is defined here by the maximum gradient in potential vorticity (Nash et al., 1996).

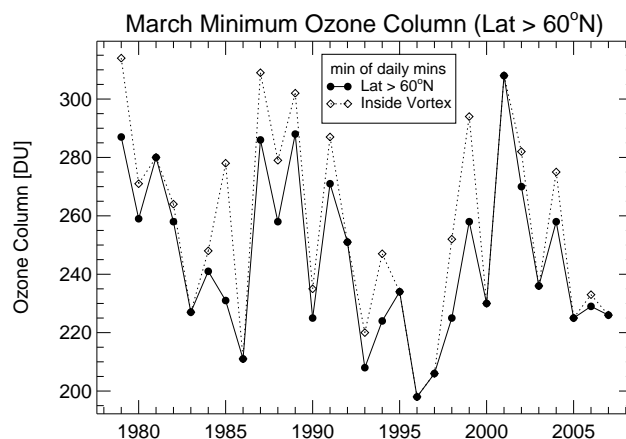
### 3.2 Minimum column ozone in the polar region

#### 3.2.1 Minimum polar total column ozone in observations

Spatially localised and transient (several days) reductions in column ozone, so-called miniholes, are frequently observed; it has been established that they are of dynamical origin (McKenna et al., 1989; Petzoldt et al., 1994; James et al., 2000; Hood et al., 2001; Weber et al., 2002). The dynamics of miniholes involves a lifting of the tropopause above a tropospheric anticyclone and poleward motion of ozone-poor subtropical air in the upper troposphere and lower stratosphere. In combination, this causes a reduction in column ozone (e.g., Reid et al., 2000; Hood et al., 2001). The dynamics of miniholes further results in equatorward excursions of polar air in the mid-stratosphere; in a situation where this air is chemically depleted in ozone by halogen chemistry in the polar vortex, particularly low ozone miniholes may occur (James et al., 2000; Weber et al., 2002; Keil et al., 2007). Furthermore, the frequency of low-ozone events changes with time; Brönnimann and Hood (2003) report that low-ozone events in winter were much more frequent in 1990–2000 than in 1952–1963 over northwestern Europe.

Nonetheless, the minimum of daily total column ozone between  $60^{\circ}$  and the pole for the period March/April in the Arctic and September–November in the Antarctic has been frequently employed as measure of polar chemical ozone loss for the validation of CCMs (Austin et al., 2003; WMO, 2003; Eyring et al., 2006, 2007; WMO, 2007). The minimum of daily total column ozone used in this study is defined as the minimum value of the daily minima over the respective periods, i.e. one single measurement or one single model grid point within a two-month period.

In Fig. 3 it is shown that for winters with little chemical ozone destruction a substantial fraction of the daily minimum ozone columns occurs *outside* the polar vortex. For



**Fig. 4.** The minimum total column ozone in the Arctic poleward of  $60^{\circ}\text{N}$  in March (solid line) computed as the minimum of daily minima. The dashed line shows the same calculation but excluding any minima occurring outside of the polar vortex. The polar vortex is defined here by the maximum gradient in potential vorticity (Nash et al., 1996).

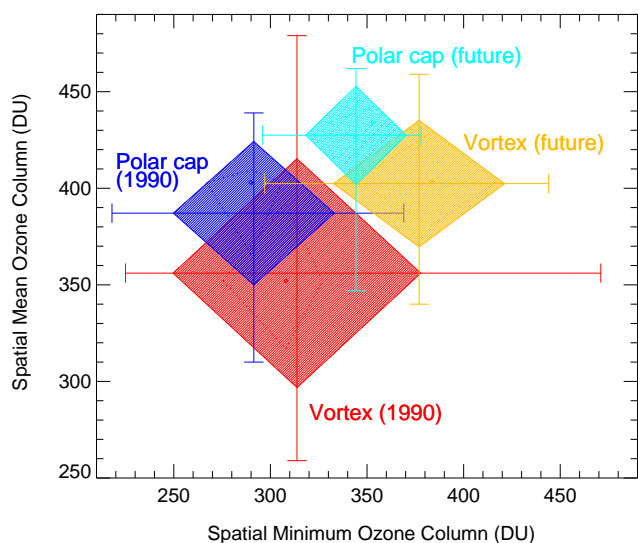
half of the winters considered here about 50% of the daily minima occur outside of the vortex, and for only four winters is the fraction of minima occurring outside below 25%. Minima occurring outside of the vortex must be due to dynamically caused low ozone events (see above) and cannot be linked with photochemical ozone loss in the polar vortex arising from heterogeneous activation of chlorine-containing molecules on polar stratospheric clouds.

If the minimum value of daily total ozone minima poleward of  $60^{\circ}\text{N}$  (as used, e.g., in WMO, 2003, 2007) is computed, it provides information from within the polar vortex in only 12 out of the 29 winters considered here (Fig. 4). The deviation between the vortex minimum total ozone and the minimum total ozone poleward of  $60^{\circ}\text{N}$  is significant, with a maximum deviation of  $\sim 50$  DU in the winter of 1984/1985.

In the Antarctic, because of the stronger reduction in column ozone due to chemical ozone destruction, the situation is different. For the period considered here, the minimum of daily column ozone between  $60^{\circ}\text{S}$  and the pole for the period September–November (not shown) is always located within the vortex.

#### 3.2.2 Minimum column polar ozone in model results

Differences between different simple measures of ozone loss and between simple and more sophisticated analyses also occur when analysing results from model simulations, as will be shown below. Since the prediction of the recovery of polar ozone is based on simulations with CCMs (e.g. Austin et al., 2003; WMO, 2003, 2007; Eyring et al., 2007), the identification of a recovery period may depend on the analysis methodology chosen for ozone loss. In the last two ozone assessments (WMO, 2003, 2007), the spatial minimum of the daily



**Fig. 5.** Simulated 1 April total ozone column from two CCM 20-year ensemble (time slice) experiments with 1990 and near-future boundary conditions for greenhouse gases and sea surface temperatures. For each time slice and for each analysis region (polar cap north of  $60^\circ$  N, or vortex according to the maximum gradient in potential vorticity, Nash et al. 1996) the ensemble statistics of spatial mean ozone column are contrasted with those of the spatial minimum ozone column. Each diagram indicates mean, range (solid lines), and one standard deviation (filled areas); quartiles are indicated as dotted lines centred around the medians which are shown as open diamonds.

spring ozone minimum poleward of  $60^\circ$  served as a quantifier for the state of the polar ozone layer for comparison with simulations from different CCMs.

For simulated ozone columns, here we use as an example two time slice experiments from the chemistry-climate model ECHAM4.DLR(L39)/CHEM (hereafter abbreviated as E39/C Hein et al., 2001; Schnadt et al., 2002). Each ensemble experiment consists of 20 recurrent simulations with constant boundary conditions (sea surface temperature, greenhouse gases), one for the 1990s and one for the near-future. The near future conditions originally aimed at the year 2015, but in the interim it has been found that the sea surface temperatures used are too high. Therefore, results from the future simulation should not be considered as a reliable projection of a specific period but rather as being indicative of possible future conditions. The 1990 time slice results, in contrast, are in agreement with the results derived from the most recent transient ensemble run for 1960 to 1999 (Dameris et al., 2005; Eyring et al., 2006). A detailed description of both time slice experiments and the simulation of ozone therewith is given elsewhere (Schnadt, 2001; Lemmen, 2005).

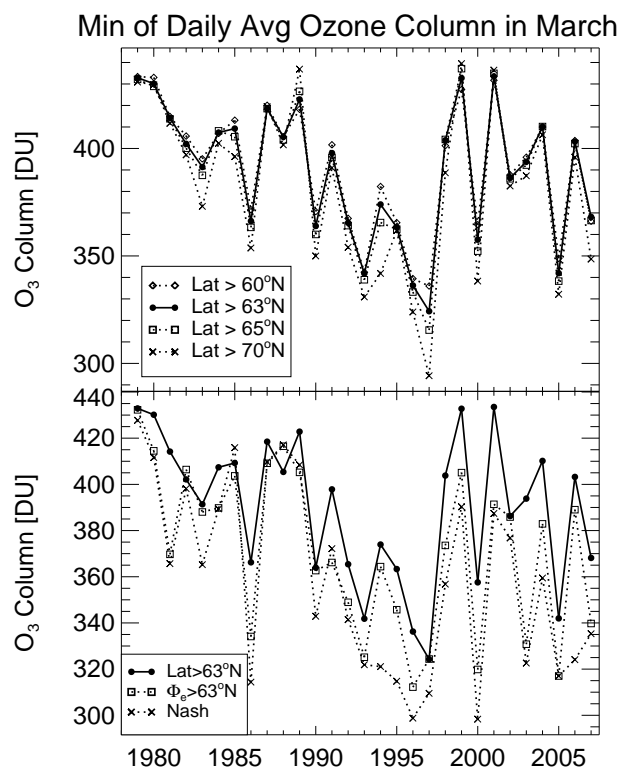
Figure 5 compares ensemble statistics of both time slices analysed separately for the polar cap and the dynamically defined vortex, all data reported for 1 April. For example, in the

1990 simulation poleward of  $60^\circ$  N (dark blue colour), the spatial mean ozone column (ordinate) was 387 DU and the range was 310–439 DU. In half of the years, the spatial mean column was larger than 403 DU. For the same analysis, the mean and range of the spatial minimum column (abscissa) are 291 DU and 218–369 DU, respectively. Column values for “future” (cyan, dark yellow) are consistently higher than for “1990” (blue, red) due to a lower chlorine loading and more dynamical heating leading to a stronger downward transport (Schnadt et al., 2002; Eyring et al., 2006).

For both time slices, the spatial mean total column ozone is higher over the polar cap than over the vortex, and the spatial minimum column is lower over the polar cap than over the vortex; this difference in spatial analysis is more pronounced in the future experiment (with less chemical ozone destruction). For the 1990 experiment, 50% of the years have a minimum column of less than 310 DU analysed within the vortex, but 75% of these winters have a minimum column of less than 310 DU analysed within the polar cap. Evidently, for many years the minimum column ozone is located outside the vortex in spite of the, on average, lower ozone column within the vortex.

The erroneous association of the minimum column ozone with high chemical ozone loss becomes obvious when chemical ozone loss deduced from methane-ozone tracer correlations (Lemmen, 2005; Lemmen et al., 2006b) is compared with minimum column ozone (see electronic supplement <http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip> for further discussion). For all winters of both time slice experiments, the maximum chemical column ozone loss is located within (or, in two cases, on) the polar vortex edge. But only for three winters in each time slice experiment is the polar ozone minimum located within the vortex. On the ensemble average, the vortex edge is located at  $\Phi_e \approx 74^\circ \pm 8^\circ$  in 1990 ( $78^\circ \pm 6^\circ$  for “future”), and the location of the polar cap minimum is around  $62^\circ \pm 15^\circ$  ( $57^\circ \pm 13^\circ$ ), i.e., clearly outside of the vortex.

Arguably, the simulated vortex area,  $\Phi_e > 74^\circ$  on 1 April for many of the analysed winters is smaller than observed climatological Arctic vortex areas (e.g., Karpetchko et al., 2005, report  $\Phi_e \approx 69^\circ$  for the climatological Arctic vortex edge for mid-March). The fact that the size of the Arctic vortex in E39/C is smaller than in reality is highlighted in the electronic supplement <http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip>, where the strength of the barrier to meridional transport is compared for analysed meteorological data and results from E39/C (Dameris et al., 2005) on the 550 K surface (averaged over the years 1990–1999 and for 1–10 April). The remaining vortex possibly no longer encompasses all chemically depleted air masses at this time. Still, even when a generous vortex boundary definition such as  $\Phi_e = 63^\circ$  is considered, for more than half of all winters the polar cap minimum ozone is located outside this rather generously defined vortex.

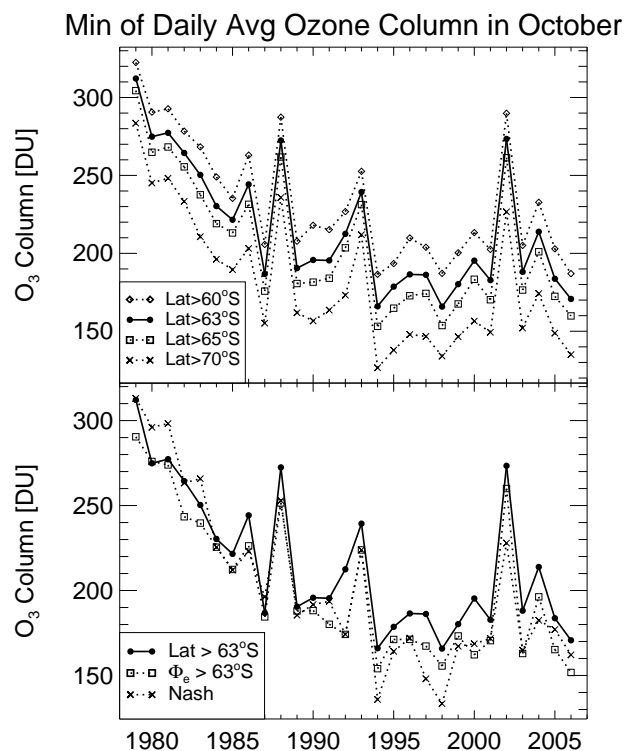


**Fig. 6.** The minimum of the daily average ozone for March poleward of  $63^{\circ}$  N (open squares), poleward of  $63^{\circ}$  N equivalent latitude (solid circles), and poleward of the vortex edge according to the Nash criterion applied on the 475 K potential temperature surface (crosses).

To avoid the misinterpretations arising from the use of minimum column ozone as a measure of chemical polar ozone loss, Lemmen et al. (2006b) recommended that a more sophisticated measure should be applied to CCM simulations to isolate chemical (halogen-induced) ozone loss from total ozone change; they suggested using ozone-tracer correlations (e.g., Proffitt et al., 1990; Tilmes et al., 2004; Müller et al., 2005). They demonstrated the applicability of this technique to output from a model simulation (Lemmen et al., 2006b) and applied it to a recent 40-year transient CCM simulation (Lemmen et al., 2006a). Similarly, Tilmes et al. (2007) applied ozone-tracer correlations to results from the WACCM3 model; they report a good reproduction of chemical ozone loss in the Antarctic vortex core ( $140 \pm 30$  DU) whereas the WACCM3 Arctic chemical ozone loss only reaches 20 DU for cold winters, which is much lower than observed values.

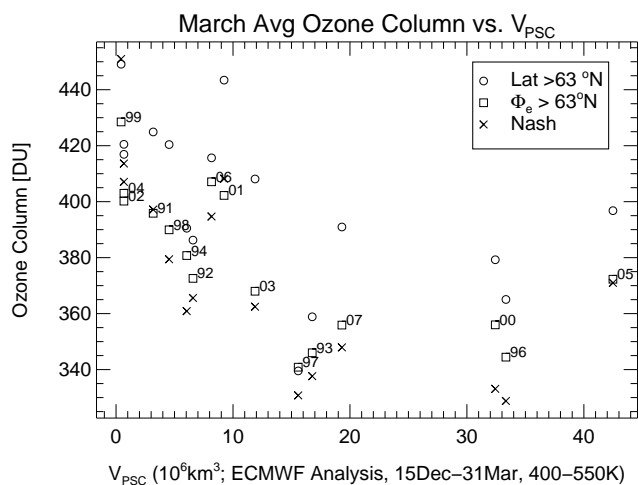
### 3.2.3 Minimum of the daily average column ozone

As an alternative measure for the maximum chemical impact on column ozone over the polar region, we suggest considering daily mean area weighted ozone over the polar region and then selecting the minimum value reached in March in



**Fig. 7.** As in Fig. 6, but for the southern hemisphere polar vortex in October.

the Arctic and in October in the Antarctic. This value should most clearly reflect the maximum impact of chemical loss on the ozone column before the vortex breaks down or before substantial mixing into the vortex occurs. The time series of this quantity for March in the Arctic (Fig. 6) shows substantial year-to-year variation, with values below 350 DU reached in several years. Generally, the lowest values are reached if ozone is averaged over the polar vortex (with the edge determined from the gradient in PV (Nash et al., 1996) on the 475 K potential temperature level) and the greatest values if the average is taken poleward of geometric latitude  $63^{\circ}$  N. Averages taken poleward of an equivalent latitude of  $63^{\circ}$  N range between the two extremes. This indicates that throughout the period considered column ozone is generally lower within the boundary of the Arctic vortex than outside. The minimum daily average polar ozone in October in the Antarctic (Fig. 7) shows less year-to-year variability and clearly lower ozone values than in the Arctic. All values after 1980 are lower than 300 DU and a decline in the values between  $\sim 1980$  and 1995 is noticeable. Compared to the Arctic, there is less variation in this quantity depending on whether latitude/equivalent latitude of  $63^{\circ}$  S or the vortex boundary is chosen as the limit of the region over which averages are calculated. This observation is consistent with the Antarctic vortex being approximately polar centric and with a vortex boundary close to  $63^{\circ}$  S (e.g., Bodeker et al., 2002; Karpetchko et al., 2005).



**Fig. 8.** The relation between  $V_{PSC}$  and the total column ozone in March averaged over the Arctic polar region. Circles show averages poleward of  $63^\circ$  N, squares averages poleward of  $63^\circ$  N equivalent latitude, and crosses averages poleward of the vortex edge defined by the maximum gradient in potential vorticity.

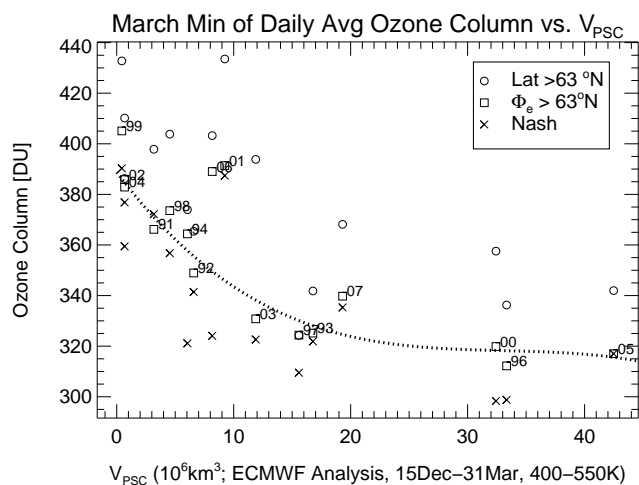
### 3.3 Relation between the mean polar ozone column and meteorological conditions

In the Arctic, chemical ozone loss is closely related to the particular meteorological conditions in each year. Rex et al. (2004) reported that Arctic chemical loss is linearly related to a measure of the potential of polar stratospheric cloud occurrence referred to as  $V_{PSC}$ .  $V_{PSC}$  is defined as the volume of stratospheric air below the threshold temperature for the existence of nitric acid trihydrate, averaged over a certain time period and altitude range (Rex et al., 2004).

$V_{PSC}$  is an absolute measure and is therefore strictly comparable only with absolute measures of ozone loss like the ozone mass deficit (e.g., Huck et al., 2007). Vortex mean ozone loss should be related to the fraction of the vortex that is activated. This effect becomes particularly important if Arctic and Antarctic conditions are compared, where the vortices have rather different sizes. Therefore, Tilmes et al. (2006) extended the concept of  $V_{PSC}$  introducing the PSC formation potential of the polar vortex (PFP),<sup>2</sup> a measure that takes into account the size of the vortex. Here we investigate whether the simple measures of polar ozone loss discussed above show any relation to the measures of chlorine activation in the vortex such as  $V_{PSC}$  and PFP.

Figure 8 shows a scatterplot of  $V_{PSC}$  against column ozone averaged over the area poleward of geometric or equivalent latitude  $63^\circ$  N and within the vortex. Obviously, in

<sup>2</sup>PFP is calculated in the following way. For all days when a vortex exists and for a defined altitude range,  $V_{PSC}$  is divided by the volume of the polar vortex, and these values are then integrated over a defined time period. Finally, this value is divided by the total number of days in the period (Tilmes et al., 2006).



**Fig. 9.** The relation between  $V_{PSC}$  and the minimum of March daily average polar column ozone (symbols as in Fig. 8). The dotted line shows an empirical fit through the minimum ozone values for averages poleward of an equivalent latitude of  $63^\circ$  N with the exception of the years 1999, 2001, and 2006 for which the final warming occurred before March (WMO, 2007); Santee, 2008<sup>1</sup>. Not shown (and excluded from the fit) is the year 1995 because of poor data quality; only 3% of the data points poleward of  $63^\circ$  equivalent latitude are valid. The polynomial describing the fit is given by  $M = 388.6 - 6.10 \cdot V + 0.178 \cdot V^2 - 0.00176 \cdot V^3$ , where  $V$  is  $V_{PSC}$  in  $10^6 \text{ km}^3$  and  $M$  is the minimum of March daily average polar column ozone in Dobson units (this relation being valid for  $V_{PSC} < 42.5$ ).

contrast to the compact relation between  $V_{PSC}$  and chemical ozone loss (Rex et al., 2004; Tilmes et al., 2004), there is no close relation between  $V_{PSC}$  and March average ozone. This holds irrespective of the horizontal boundary specified as the limit for the averaging. Apparently, the March average ozone is a measure that does not adequately differentiate between chemical loss and dynamical resupply of ozone that both change substantially from year to year. Furthermore, the March average ozone poleward of geometric latitude of  $63^\circ$  N is particularly sensitive to the strong year-to-year variability in the lifetime and shape of the Arctic vortex (Waugh and Randel, 1999; Karpetchko et al., 2005). We argued above that the minimum of March daily average polar column ozone should be a quantity more closely related to chemical ozone loss than those shown in Fig. 8. However, no clearly compact and especially no linear relation of the minimum of March daily column ozone with  $V_{PSC}$  emerges.

Some outliers, e.g. the years 1999, 2001, and 2006, can be understood. For these years, the final warming was very early so that no vortex existed during March (WMO, 2007); Santee 2008<sup>1</sup>, with the consequence of a larger ozone column caused by the influx of mid-latitude ozone-rich air. Concentrating on data for the years when the polar vortex in March was intact, we find a tighter relation between  $V_{PSC}$  and the minimum of daily average ozone column poleward



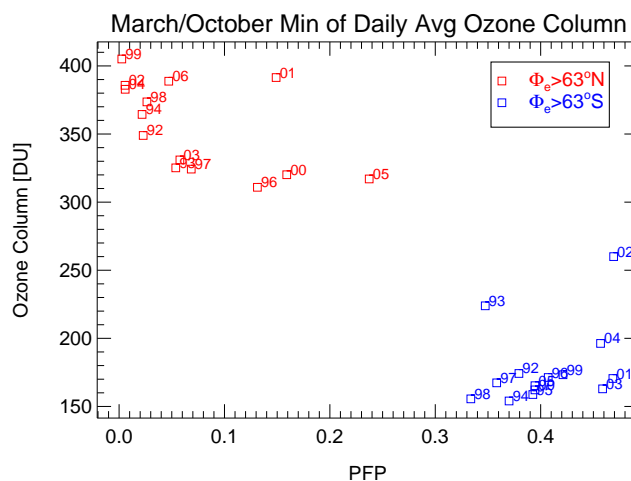
of  $\Phi_e=63^\circ$  that is indicated by the dotted line in Fig. 9. From the three different quantities shown, the polynomial fit with this quantity has the lowest deviation ( $1\sigma=5.5$  DU) from the observations and shows a (nearly) monotonic decrease with increasing  $V_{PSC}$ .

Obviously, the minimum ozone column in the three coldest winters (1996, 2000, and 2005) is not much lower than the minimum in moderately cold winters such as 1993 or 2003. This effect is mostly responsible for the nonlinear relation between  $V_{PSC}$  and the minimum of March column ozone.

When PSC occurrence or the potential for chlorine activation is compared for the Arctic and Antarctic,  $V_{PSC}$  is no longer a suitable measure, because the polar vortex is much larger in the Antarctic than in the Arctic; for such a comparison, the PFP should be used instead of  $V_{PSC}$  (Tilmes et al., 2006). In Fig. 10, the PFP is plotted against the minimum of daily average polar column ozone in spring, where only data points for averages poleward of  $63^\circ$  equivalent latitude are shown. The well-known fact that winter/spring temperatures are lower in the Antarctic than in the Arctic (leading to a greater PFP, Tilmes et al., 2006) and that polar column ozone is greater in the Arctic than in the Antarctic (Dobson, 1968; WMO, 2007) is reflected in this plot. However, for this combination of meteorological and ozone measures, again no clearly compact relation emerges.

### 3.4 Relation between polar column ozone and chemical ozone loss

To diagnose chemical ozone destruction in the polar vortices, several methods have been developed that bring in information beyond ozone measurements which allows the impact of transport and chemistry on ozone to be disentangled (Harris et al., 2002; WMO, 2007). Here (Fig. 11), we compare chemical ozone loss between 380–550 K potential temperature in March deduced for the Arctic winters 1992–2005 based on the vortex average method (Chipperfield et al., 2005; Rex et al., 2006) with the two most commonly used simple measures based on total ozone, namely the average polar ozone poleward of geometric latitude  $63^\circ$  N in March and the minimum of all total ozone measurements in March poleward of geometric latitude  $60^\circ$  N (e.g., Newman et al., 1997; Austin et al., 2003; WMO, 2007). We also performed a similar comparison (not shown) using chemical ozone loss deduced for the Arctic winters 1992–2003 by applying the ozone-tracer correlation method (Tilmes et al., 2004) and obtained very similar results. Clearly, there is some relation between these measures and chemical ozone loss. The stronger the chemical ozone loss, the lower both the average March and the minimum March total ozone are (Fig. 11, top and middle panel). However, in these two cases, the correlation is not very strong and in both cases outliers are noticeable. Further, the stronger the link between chemical ozone loss and a simple measure, the closer should the slope  $m$  of the linear fit approach minus one (this would be the value for a year-

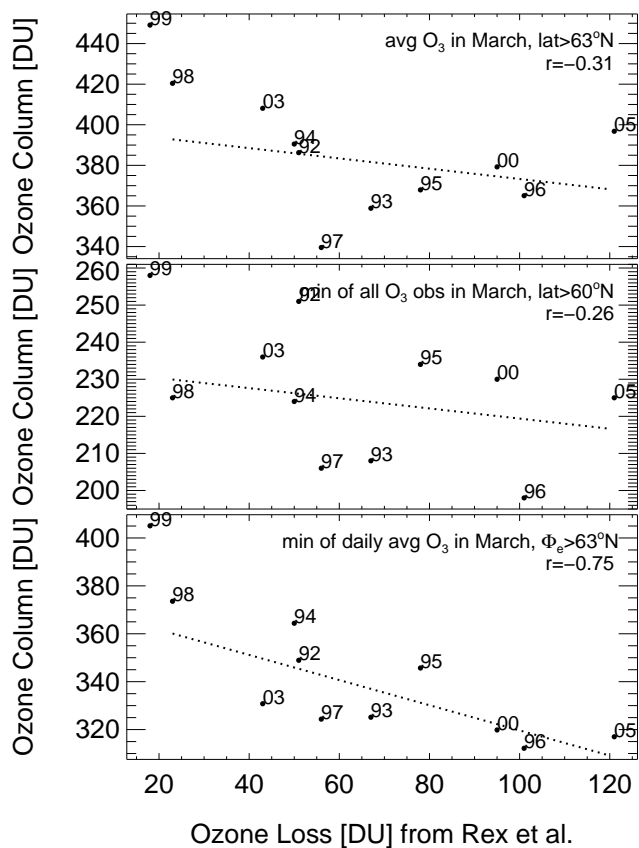


**Fig. 10.** The relation between PFP and the minimum of daily average polar column ozone in spring for both the Arctic (March, red symbols) and the Antarctic (October, blue symbols). Only data points for averages poleward of  $63^\circ$  equivalent latitude are shown. PFP values are calculated on the basis of MetO analysis (Swinbank and O'Neill, 1994), averaged over the altitude region 400–550 K for the period 15 December to 31 March and 15 June to 30 September for the Arctic and Antarctic, respectively. Note that the PFP values shown here are slightly improved compared to those reported by Tilmes et al. (2006), see also electronical supplement (<http://www.atmos-chem-phys.net/8/251/2008/acp-8-251-2008-supplement.zip>).

to-year variability in total ozone that is solely due to chemical loss). The slope calculated for the average polar ozone in March (top panel in Fig. 11) is  $m = -0.25$  and for the minimum column ozone (middle panel)  $m = -0.14$ . For the minimum of daily average ozone poleward of equivalent latitude  $63^\circ$  N (bottom panel) the strongest slope is obtained,  $m = -0.53$ . Further, for this measure, the most compact relation and the best correlation emerges. That means, of all the simple measures considered, the new measure suggested here (Sect. 3.2.3) shows the closest correlation with chemical ozone loss.

## 4 Discussion

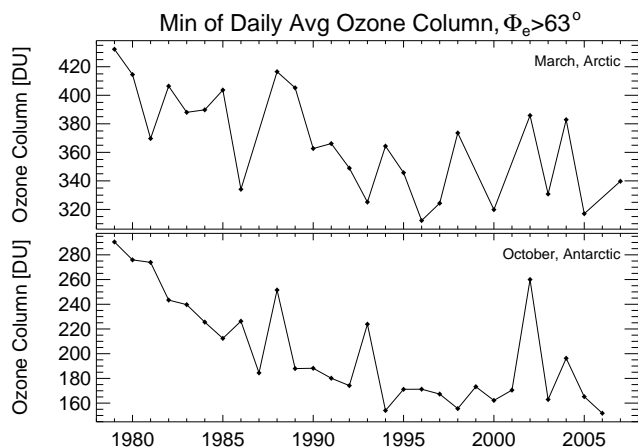
Among the quantities describing stratospheric ozone, total column ozone is the one most easily measured. Indeed, the longest atmospheric time series exist for total ozone (e.g., Dobson, 1968; Brönnimann et al., 2003). Long time series are a prerequisite for deriving correlation properties of total ozone and for a reliable trend analysis (e.g., Staehelin et al., 1998; Kiss et al., 2007; Vyushin et al., 2007; WMO, 2007). Moreover, the Antarctic ozone hole was discovered, and the discovery corroborated, through measurements of total column ozone (Chubachi, 1984; Farman et al., 1985; Stolarski et al., 1986). However, variations in total ozone are caused



**Fig. 11.** The relation between the observed chemical ozone loss in the Arctic (Chipperfield et al., 2005; Rex et al., 2006) and simple measures based on total ozone; compared are the two most commonly used measures (top and middle panel) and the measure (see section 3.2.3) recommended here (bottom panel). Top panel shows the average polar ozone poleward of geometric latitude  $63^\circ$  N in March, middle panel the minimum of all total ozone measurements in March poleward of geometric latitude  $60^\circ$  N and bottom panel the minimum of daily average ozone poleward of equivalent latitude  $63^\circ$  N. Correlation coefficients from the linear fits ( $r$ ) are shown in all panels. No ozone loss values are reported for 2001 and 2006. In all fits, 1999 was excluded as the vortex broke up before March in this year.

by both chemical change and by transport, and the different impacts of these processes are often difficult to disentangle.

Transport contributions to polar ozone variability are closely controlled by the strength of the middle atmospheric (Brewer-Dobson) circulation that is driven by tropospheric wave forcing. For both the Arctic and Antarctic, a stronger-than-average planetary wave forcing in winter leads to more transport of ozone to high latitudes because of a stronger circulation and a higher-than-average polar lower stratospheric temperature and a weaker vortex in early spring, whereas a weaker wave forcing leads to less transport, lower polar temperatures in spring and to a stronger vortex (Fusco and Salby, 1999; Newman et al., 2001; Weber et al., 2003). The variability in polar ozone due to the variability in wave forcing is



**Fig. 12.** Time series of minimum of daily average column ozone poleward of  $63^\circ$  equivalent latitude for March in the Arctic (top panel) and October in the Antarctic (bottom panel). Winters in which the vortex broke up before March (1987, 1999, 2001, and 2006) are not shown for the Arctic time series.

much larger in the Arctic than in the Antarctic, while chemical loss is more persistent in Antarctica.

Here we argue that measures of chemical ozone loss based on monthly averages of total column ozone over the polar cap, although they contain information about chemical ozone loss, must be interpreted with caution. The particular value of such measures will depend on the selected definition of the equatorward boundary of the region over which averages are calculated. In the Arctic, the year-to-year variability of the size of the polar vortex has a particularly strong impact. Further, there are winters when the Arctic vortex breaks up early (e.g., WMO, 2007), a fact that is often neglected in the interpretation of monthly mean total column ozone time series. Moreover, averages over the polar cap neither show compact relationships with meteorological measures of the extent of chlorine activation (and thus the potential for ozone destruction) in the polar vortices such as  $V_{PSC}$  and PFP nor a good correlation with observed chemical ozone loss. Although frequently employed, the minimum value of daily total column ozone minima over the polar region is a particularly problematic measure, insofar as it relies on a single measurement or on a single model grid point. We have shown here that for the Arctic, both in a CCM and in observations, a significant fraction of the minimum values of daily total ozone minima occur *outside* the polar vortex. Clearly, if the minimum ozone value on a particular day occurs outside the vortex, it does not provide information on halogen-driven chemical ozone loss initiated by heterogeneous chlorine activation. Because of the strong chemical ozone loss in the Antarctic, this problem is much less pronounced there. It should, however, become increasingly relevant for simulations of future ozone loss, when much less chemical ozone loss is expected. Based on this analysis, we must question the applicability of

a simple measure such as a minimum polar cap ozone value for both the E39/C model and observations, and strongly caution against application of this simple measure to results from other CCM simulations. It is remarkable that there is substantial variation in the magnitude of minimum total column ozone in the Arctic in current model simulations ranging roughly from 220 to 320 DU whereas the observations lie in the range of 290 DU to below 200 DU (WMO, 2007, Fig. 6–13).

Clearly, employing sophisticated measures of polar chemical ozone loss that show a compact correlation with meteorological measures of chlorine activation (Rex et al., 2004; Tilmes et al., 2006) is the best method to assess the temporal evolution of polar ozone loss in both models and observations. However, when only total column ozone data are available, we propose that the minimum of March (or October) daily average total ozone in the vortex (where the vortex boundary can be determined by the maximum gradient in PV or by an equivalent latitude criterion) should be considered. This quantity is not strongly influenced by low column ozone outside the vortex and does not rely on a single measurement or model grid point. Further, in contrast to monthly averages, its year-to-year variability is not substantially affected by varying dates of vortex breakdown in the Arctic. If equivalent latitude  $63^\circ$  is applied as the criterion for polar air, then the minimum of daily average total ozone in spring also shows a reasonable correlation with observed chemical ozone loss (Fig. 11).

In Fig. 12, a time series is shown of the spring minimum of daily average total ozone poleward of  $63^\circ$  equivalent latitude for March in the Arctic and October in the Antarctic. Compared to the frequently shown figure of the temporal development of the monthly mean total ozone poleward of  $63^\circ$  geographic latitude (WMO, 2007, Fig. 4–7), Fig. 12 shows a different picture, with a stronger contrast for the Arctic. However, in both cases, the time series of minimum daily average column ozone avoids the impression of an apparent “recovery” of polar ozone since the late nineties conveyed by the customary monthly mean total ozone figure.

## 5 Conclusions

Quantities deduced from measurements of total column ozone are frequently used as measures of polar ozone loss. One of the most common measures is monthly mean column ozone poleward of  $63^\circ$  for March and October in the Arctic and Antarctic, respectively (WMO, 2007). For the Arctic, a latitude of  $63^\circ$  is a reasonable boundary for polar air, except for winters when the vortex breaks up early. In contrast, for the Antarctic, the values of the October means (but not their year-to-year variability) are sensitive to the exact choice of  $63^\circ$ . A better definition of the polar vortex boundary can be obtained using the gradient in PV (Nash et al., 1996) or equivalent latitude.

The spring minimum of daily total ozone minima poleward of a given latitude is a further measure that is often employed, in particular, when comparing polar ozone loss in observations and models. This is a problematic measure, insofar as it relies on one single measurement or on one single model grid point; for Arctic conditions, it is not unlikely that this minimum value occurs in air outside the polar vortex. We suggest that this concept should no longer be used as a measure of polar ozone loss.

Neither for the monthly mean column ozone poleward of  $63^\circ$  for March and October nor for the spring minimum of daily total ozone minima can a close relation be obtained with meteorological quantities that describe the potential for polar heterogeneous chlorine activation. Likewise, neither of the two measures shows a good correlation with chemical ozone loss in the vortex deduced from observations (Rex et al., 2006; Tilmes et al., 2004).

Considering the minimum of daily average total ozone poleward of  $63^\circ$  equivalent latitude (or in the vortex) in spring, avoids the problem of relying on one single data point and reduces the impact of year-to-year variability in the Arctic vortex break-up on ozone loss measures (see Sect. 3.2.3). This measure also shows a reasonable correlation ( $r = -0.75$ ) with observed chemical ozone loss.

We propose the minimum of daily average total ozone poleward of  $63^\circ$  equivalent latitude as a candidate for a useful simple measure of polar ozone loss, when winters with an early vortex break-up (before March in the Arctic) are excluded. In any event, it is always preferable to employ more sophisticated measures of chemical polar ozone loss (e.g., Harris et al., 2002; Rex et al., 2004; Tilmes et al., 2006; Lemmen et al., 2006b; WMO, 2007) that bring in additional information to disentangle the impact of transport and chemical change on ozone.

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