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Orthorhombic 11C pyrrhotite from Michałkowa, Góry Sowie Block, The Sudetes, Poland – preliminary report

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Abstract

This study provides the preliminary report about first occurrence of orthorhombic 11C pyrrhotite ($\text{Fe}_{(1-x)}\text{S}$) from the Sudetes, Poland. Samples of pyrrhotite-containing two-pyroxene gabbro were found in a classic pegmatite locality in Michałkowa near Walim in the Góry Sowie Block. Based on microscopic methods, pyrrhotite is associated with pentlandite, chalcopyrite, chromite, ilmenite, gersdorffite, magnetite, biotite, magnesiohornblende, clinocllore, lizardite and talc. X-Ray diffraction (XRD) indicate that pyrrhotite has orthorhombic 11C structure and it is characterized by: $a = 3.433(9) \text{ \AA}$, $b = 5.99(2) \text{ \AA}$, $c = 5.7432(5) \text{ \AA}$, $\beta = 90^\circ$ and $d_{102} = 2.06906 \text{ \AA}$. Mössbauer studies confirmed the XRD data. Pyrrhotite has three sextets with hyperfine parameter values 30.8 T for sextet A, 27.9 T and 25.8 T for sextets B and C respectively, indicating orthorhombic structure, the composition near $\text{Fe}_{10}\text{S}_{11}$ and $x = 0.0909$.

Key words: orthorhombic pyrrhotite, Polish Sudetes, Góry Sowie Block

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1. Introduction

Pyrrhotite is one of the most common sulfide minerals, widespread in many rocks and is an important component of the sulfide ore deposits, especially in those that contain Ni, Cu and Pt group elements (De Villiers and Liles 2010). The pyrrhotite group is non-stoichiometric and has the general formula of $\text{Fe}_{(1-x)}\text{S}$ where $0 \leq x < 0.125$ and is based on the hexagonal close-packed NiAs structure (Becker et al. 2011). The nonstoichiometry of pyrrhotite is caused by a deficiency of iron atoms in the Fe-layers (Kontny et al. 2000). Numerous pyrrhotite superstructures have been described in the literature, but only three of them are naturally occurring at ambient conditions (e.g. Pósfai et al. 2000; Fleet 2006; Becker et al. 2011). This includes the stoichiometric hexagonal FeS known as troilite

(Evans 1970), the ferrimagnetic 4C pyrrhotite with an ideal composition Fe_7S_8 and monoclinic crystallography (Powell et al. 2004) and non-magnetic orthorhombic or monoclinic pyrrhotite formally described as NC pyrrhotite where N represents the number of repetitions of the NiAs unit cell along the c-axis (Carpenter and Desborough 1964; Morimoto et al. 1970; Koto et al. 1975; Pósfai et al. 2000; De Villiers and Liles 2010; Becker et al. 2011; Liles and De Villiers 2012). The non-magnetic NC pyrrhotites have a range of compositions varying from Fe_9S_{10} to $\text{Fe}_{11}\text{S}_{12}$ including three stoichiometric polytypes: 5C pyrrhotite with Fe_9S_{10} formula (De Villiers et al. 2009), 11C polytype (Yamamoto and Nakazawa 1982) and 6C pyrrhotite (Koto et al. 1975) with $\text{Fe}_{10}\text{S}_{11}$ and $\text{Fe}_{11}\text{S}_{12}$ formulas respectively.

This study provides the preliminary report about the orthorhombic 11C pyrrhotite from the Polish Sudetes. Samples of two-pyroxenes gabbro were found in a classic pegmatite locality near Walim in the Góry Sowie Block (Fig. 1), on the right bank of the Młynowka River in Michałkowa (Michelsdorf), better

known as the sarcopside type locality (Websky 1868), where feldspars were exploited before the World War II in open pits and adits (Łodziński and Sitarz 2009). These rocks are commonly occurring as a veins in the Sowie Mts. gneissic block.

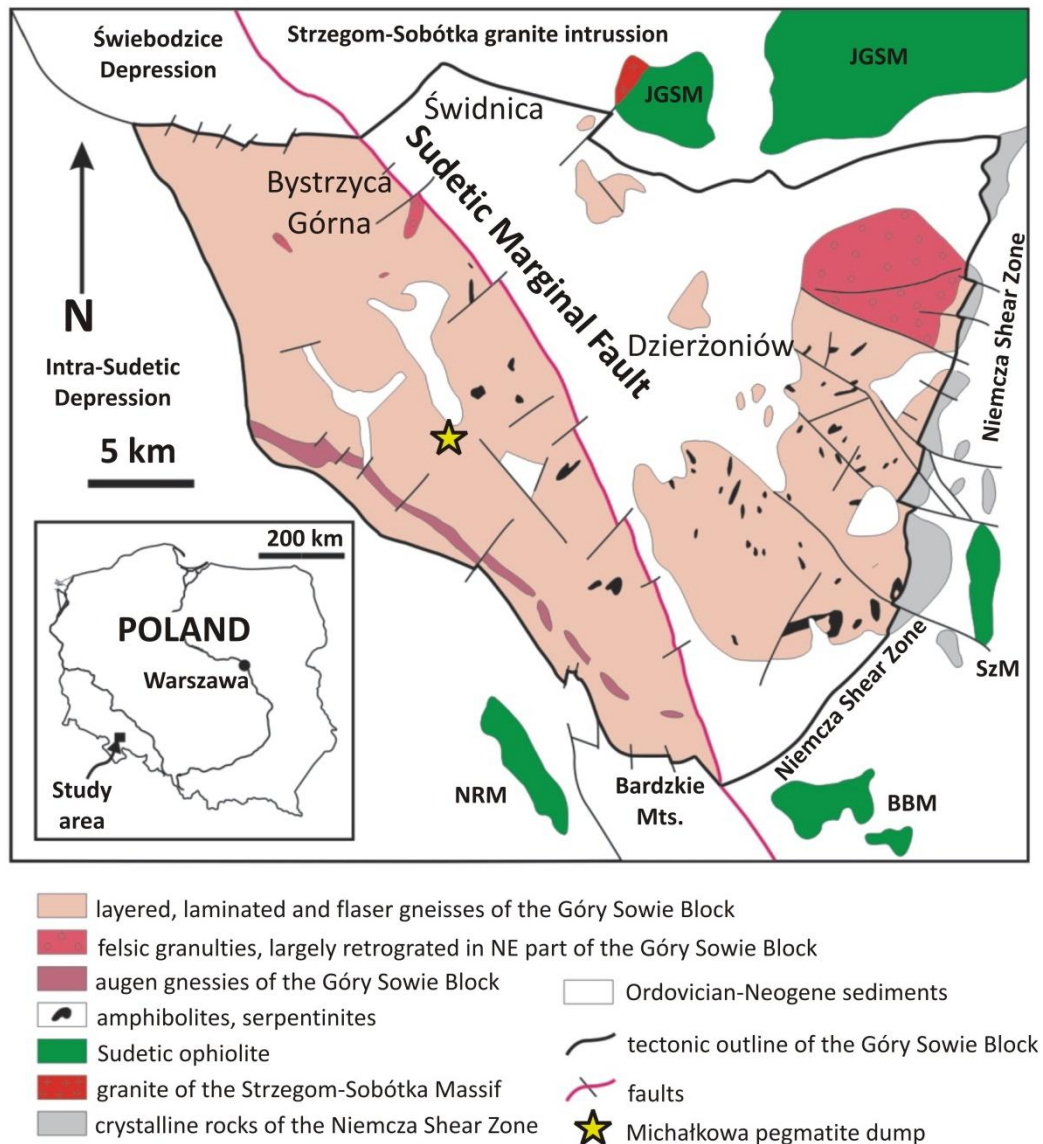


Fig.1. Geological map of the Góry Sowie Block (after Szuszkiewicz et al. 2013, modified).

BBM – Braszowice–Brzeźnica Massif, JGSM – Jordanów–Gogołów Serpentinite Massif; NRM – Nowa Ruda Massif, SzM – Szklary Massif (all constitute the Sudetic Ophiolite).

2. Samples and methods

Two samples of pyrrhotite containing two-pyroxene gabbro have been collected from the pegmatite dump in Michałkowa. Mineralogical

studies were multistage. All analyses have been performed at the University of Silesia. Mineral abbreviations utilized in this study follow the recommendations of Whitney and Evans (2010) (Tab.1).

Tab.1. Mineral abbreviations used in this study (after Whitney and Evans (2010)).

Mineral name	Formula	Abbreviation
Aluminian chromite	$\text{Fe}^{2+}(\text{Cr}^{3+}, \text{Al})_2\text{O}_4$	Alc
Biotite	$\text{K}(\text{Mg}, \text{Fe}^{2+})_3\text{AlSi}_3\text{O}_{10}(\text{F}, \text{OH})_2$	Bt
Chalcopyrite	CuFeS_2	Ccp
Chromite	$\text{Fe}^{2+}\text{Cr}^{3+}_2\text{O}_4$	Chr
Clinocllore	$(\text{Mg}, \text{Fe}^{2+})_5\text{Al}(\text{AlSi}_3\text{O}_{10})(\text{OH})_8$	Clc
Gersdorffite	NiAsS	Ger
Ilmenite	$\text{Fe}^{2+}\text{TiO}_3$	Ilm
Lizardite	$\text{Mg}_3(\text{Si}_2\text{O}_5)(\text{OH})_4$	Lz
Magnesio-hornblende	$\text{Ca}_2[(\text{Mg}, \text{Fe}^{2+})_4\text{Al}](\text{Si}_7\text{Al})\text{O}_{22}(\text{OH})_2$	Mhb
Magnetite	$\text{Fe}^{2+}\text{Fe}^{3+}_2\text{O}_4$	Mag
Pentlandite	$(\text{Fe}, \text{Ni})_9\text{S}_8$	Pn
Pyrrhotite	Fe_{1-x}S	Po
Talc	$\text{Mg}_3(\text{Si}_4\text{O}_{10})(\text{OH})_2$	Tlc

2.1. Microscopic methods

Polished sections were studied by optical microscopy of reflected light and using scanning electron microscopy (SEM) in Laboratory of Scanning Microscopy at Faculty of Earth Sciences, University of Silesia with the use of environmental analytical microscope Philips ESEM XL30/TMP equipped with detector EDS (Energy Dispersive Spectrometer) / EDAX Sapphire type in low-voltage regime (low vacuum to 1 Torr).

2.2. X-Ray diffraction (XRD)

XRD patterns were obtained with the Philips PW 3710 diffractometer using $\text{CoK}\alpha$ radiation (35 and 45 kV, 30 mA). Powder diffractograms were acquired in the $3-90^\circ$ 2θ range, with 7-20 s counting per 0.01° and 0.04° 2θ step. Analyses were performed in the X-Ray Laboratory at Faculty of Earth Sciences, University of Silesia.

2.3. Mössbauer spectrometry

Mössbauer spectra were obtained using POLON mössbauer spectrometer with

a constant acceleration. The radiation source was a ^{57}Fe in Pd matrix. The velocity scales were calibrated with a $\alpha\text{-Fe}$ foil absorber at room temperature and isomer shifts are quoted with respect to the centroid of these spectra. The spectrometer was operated by MOSIEK software. Mössbauer spectra have been developed in the MOS program. Analyses were performed in the Department of Nuclear Physics and Its Applications at the Institute of Physics, University of Silesia.

3. Results and discussion

3.1. Microscopic observations

Optical and SEM studies revealed that pyrrhotite crystals up to 2 cm size are common in two-pyroxene gabbro from Michałkowa. They are often associated with pentlandite, chalcopyrite (Fig. 2, 3), ilmenite (Fig. 3c), gersdorffite (Fig. 3d) and chromite ones (Fig. 3b). These last form crystals up to 15 μm size with oscillatory concentric zoning caused by variations in the proportions of their major elements (Fe, Mg, Al, Cr) (Fig. 3b).

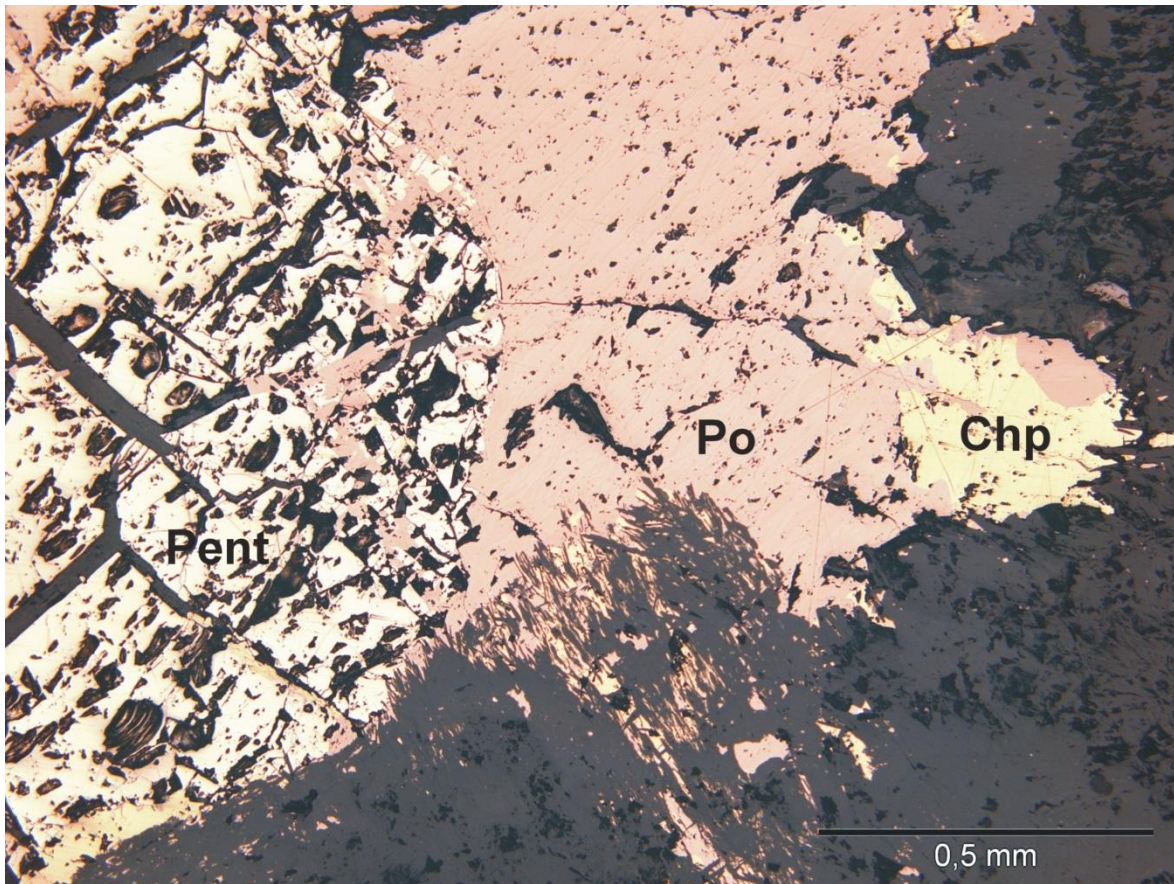


Fig.2. Microphotograph in reflected light of pyrrhotite associated with pentlandite and chalcopyrite (IN).

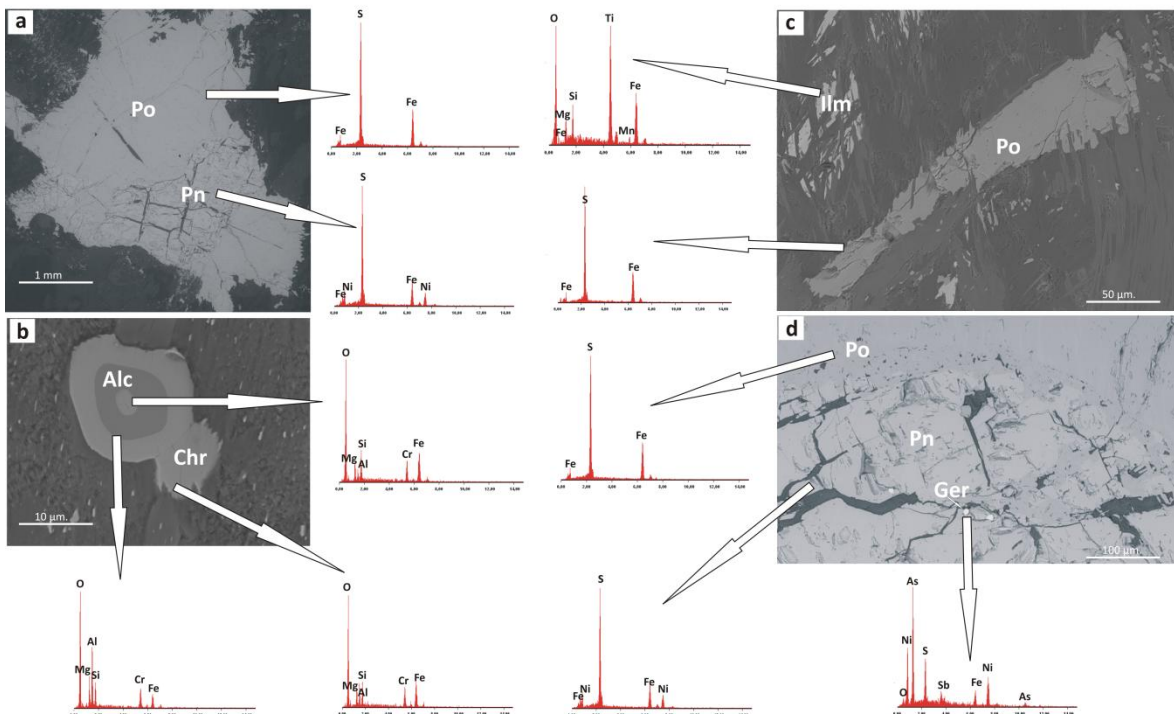


Fig.3. SEM images (BSE) and EDS spectra of studied samples. A) Pyrrhotite with pentlandite; B) Zonal structured mineral (oscillatory concentric zoning) of the spinel group with ferrous-chromium zones inside and outside (chromite) and aluminian-ferrous-chromium between them (aluminian chromite); C) Pyrrhotite with ilmenite; D) Pyrrhotite with pentlandite and gersdorffite.

3.2. X-Ray powder diffraction

Pyrrhotite unit cell dimensions vary with changes in the chemical composition (Ramsdell 1927; Grönvold and Haraldsen 1952). Arnold and Reichen (1962) have shown that the relation of interplanar distance d_{102} and iron content in pyrrhotites can be useful in estimating the chemical composition of pyrrhotites and calculating the x parameter. The most Fe-deficit phases have the smallest d_{102} values, while the increase of the iron content results in greater interplanar distance d_{102} which changes in the following ranges: from $d_{102} = 2.0591 \text{ \AA}$ in monoclinic 4C pyrrhotite from Chichibu Mine, Akaiwa, Japan with a composition of Fe_7S_8 (Morimoto et al. 1975) to $d_{102} = 2.0952 \text{ \AA}$ in troilite from Long Island meteorite, Kansas, USA (Carpenter and Desborough 1964).

XRD analysis indicate that studied pyrrhotite has orthorhombic 11C structure with $a = 3.433(9) \text{ \AA}$, $b = 5.99(2) \text{ \AA}$, $c = 5.7432(5) \text{ \AA}$, $\beta = 90^\circ$, very similar to that described by Morimoto et al. (1970). The d_{102} value was obtained to be 2.06906 \AA indicating the composition near $\text{Fe}_{10}\text{S}_{11}$ and $x = 0.0909$.

In addition to the phases identified by microscopic methods, pyrrhotite occurs with magnetite, biotite and magnesio-hornblende. These last two phases indicating the hydration of igneous interstitial alloys or subsolidus processes – amphibolization or biotitization (Fig. 4). Moreover, the presence of clinocllore, lizardite and talc indicates that chloritization and serpentinization processes have had some impact on the studied samples (Fig. 4).

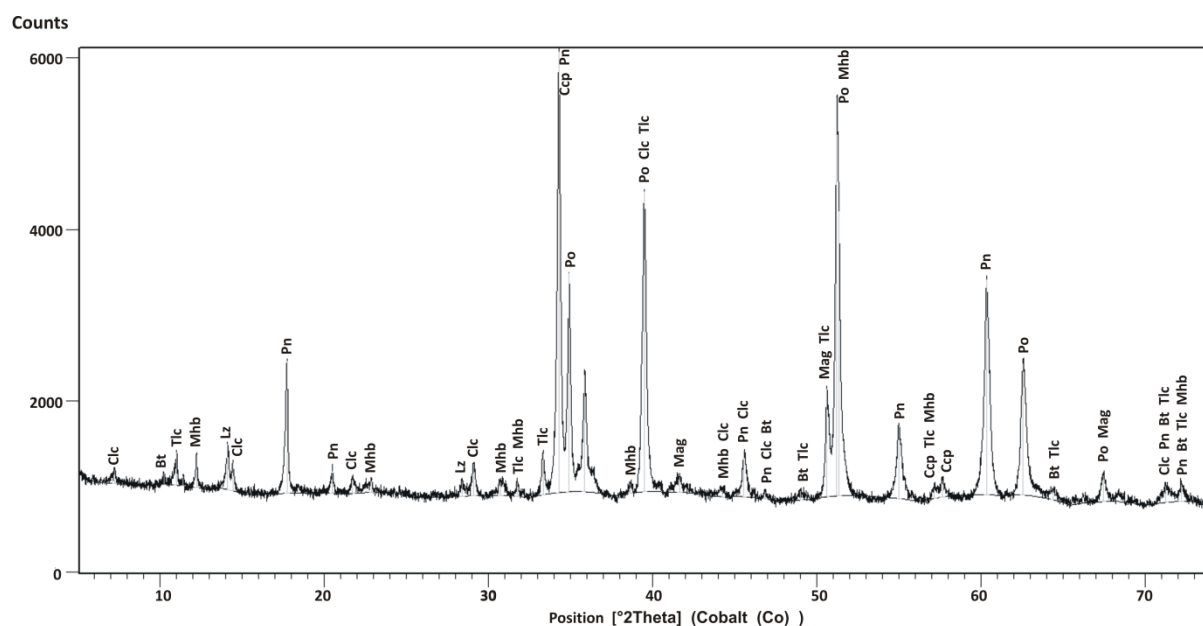


Fig.4. Diffractogram of the studied sample. Mineral abbreviations such as in Table 1.

3.3. Mössbauer data

Three pyrrhotite sextets (A, B, C) are typical for low iron content pyrrhotites (with high values of x). For small values of x , only one sextet (indicating one type of Fe position) is reported up to $x = 0.042$, the interval $0.054 < x$

< 0.079 is characterized by broad peaks, while in phases with values $0.106 < x < 0.143$ three well-resolved sextets are observed (Kruse 1990).

Analysis of the hyperfine splitting parameter (H), which is proportional to the total local magnetic field in the core

surrounded by the magnetic field, gave the information about the magnetic arrangement and structure of the studied pyrrhotite. The higher H values are observed for the more disordered structures, which means a higher symmetry (Kruse 1990; Rybicki 2011).

In Michałkowa pyrrhotite three sextets with hyperfine parameter values 30.8 T for sextet A, 27.9 T and 25.8 T for sextets B and C respectively, were observed (Fig. 5, Tab. 2). These values are higher in comparison to the 4C monoclinic pyrrhotites from the Polish Sudetes and lower than in the troilite from

Morasko meteorite (see Rybicki 2011), indicating that pyrrhotite from Michałkowa two-pyroxene gabbro is most likely orthorhombic.

Besides the pyrrhotite's sextets, two doublets have been identified based on data from McCammon (1995). The doublet D with isomer shift (IS) = 0.38 mm/s and quadrupole splitting (QS) = 0.14 mm/s fits the spectrum of pentlandite (Fig. 5, Tab. 2), while doublet E with IS = 1.12 mm/s and QS = 1.29 mm/s is most likely derived from magnesio-hornblende (Fig. 5a, b, Tab. 2).

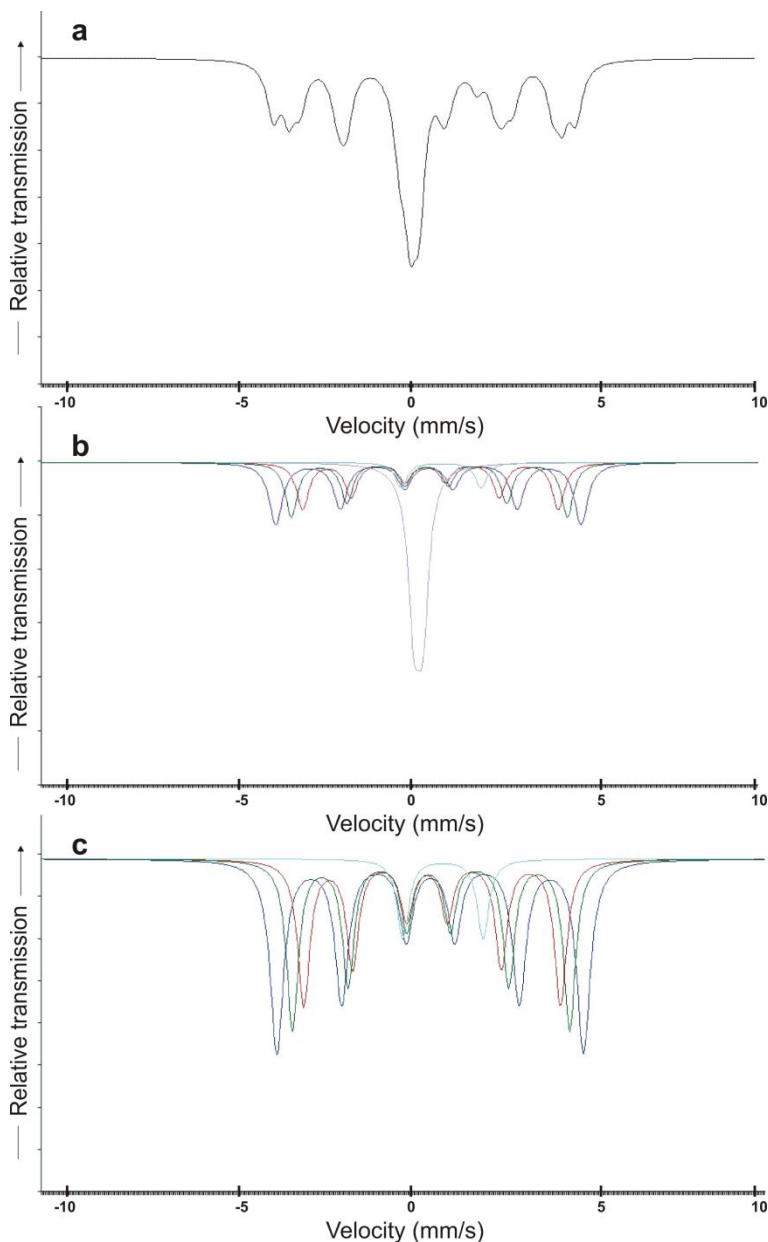


Fig.5. Mössbauer spectra of the studied sample. The subspectra colors are the same as in Table 2. A) Mössbauer total spectrum; B) Mössbauer spectrum divided into subspectra; C) Mössbauer spectrum divided into subspectra without magnesio-hornblende subspectrum

Tab.2. Mössbauer data for studied sample.

Mineral	Subspectrum	IS (mm/s)	QS (mm/s)	H (T)
Pyrrhotite	A	0.71	0.01	30.8
	B	0.70	-0.04	27.9
	C	0.69	-0.08	25.8
Pentlandite	D	0.38	0.14	-
Magnesio-hornblende	E	1.12	1.29	-

4. Conclusions

This study provides the first description of orthorhombic pyrrhotite from Polish Sudetes. It occurs in two-pyroxene gabbro with pentlandite, chalcopyrite, chromite, aluminium chromite, ilmenite, gersdorffite, magnetite, biotite, magnesio-hornblende, clinocllore, lizardite and talc.

Pyrrhotite has orthorhombic 11C structure with $a = 3.433(9) \text{ \AA}$, $b = 5.99(2) \text{ \AA}$, $c = 5.7432(5) \text{ \AA}$, $\beta = 90^\circ$ and $d_{102} = 2.06906 \text{ \AA}$.

Mössbauer studies confirm XRD data. Pyrrhotite has three sextets with hyperfine parameter values 30.8 T for sextet A, 27.9 T and 25.8 T for sextets B and C respectively, indicating orthorhombic structure, the composition near $\text{Fe}_{10}\text{S}_{11}$ and the value of x parameter = 0.0909.

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