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Brown Carbon Production by Aqueous-Phase Interactions of Glyoxal and SO2

Abstract

Oxalic acid and sulfate salts are major components of aerosol particles. Here, we explore the potential for their respective precursor species, glyoxal and SO₂, to form atmospheric brown carbon via aqueous-phase reactions in a series of bulk aqueous and flow chamber aerosol experiments. In bulk aqueous solutions, UV- and visible-light-absorbing products are observed at pH 3-4 and 5-6, respectively, with small but detectable yields of hydroxyquinone and polyketone products formed, especially at pH 6. Hydroxymethanesulfonate (HMS), C₂, and C₃ sulfonates are major products detected by electrospray ionization mass spectrometry (ESI-MS) at pH 5. Past studies have assumed that the reaction of formaldehyde and sulfite was the only atmospheric source of HMS. In flow chamber experiments involving sulfite aerosol and gas-phase glyoxal with only 1 min residence times, significant aerosol growth is observed. Rapid brown carbon formation is seen with aqueous aerosol particles at >80% relative humidity (RH). Brown carbon formation slows at 50-60% RH and when the aerosol particles are acidified with sulfuric acid but stops entirely only under dry conditions. This chemistry may therefore contribute to brown carbon production in cloud-processed pollution plumes as oxidizing volatile organic compounds (VOCs) interact with SO₂ and water.

Keywords

Absorption, Ions, Aerosols, pH, Oxidation

Disciplines

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Notes

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2	interactions of glyoxal and SO2
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14 ABSTRACT: Oxalic acid and sulfate salts are major components of aerosol particles. Here we 15 explore the potential for their respective precursor species, glyoxal and SO₂, to form atmospheric 16 brown carbon via aqueous-phase reactions in a series of bulk aqueous and flow chamber aerosol 17 experiments. In bulk aqueous solutions, UV- and visible-light absorbing products are observed at pH 3 - 4 and 5 - 6, respectively, with small but detectable yields of hydroxyquinone and 18 19 polyketone products formed, especially at pH 6. Hydroxymethanesulfonate (HMS), C₂, and C₃ 20 sulfonates are major products detected by ESI-MS at pH 5. Past studies have assumed that the 21 reaction of formaldehyde and sulfite was the only atmospheric source of HMS. In flow chamber 22 experiments involving sulfite aerosol and gas-phase glyoxal with only 1-minute residence times, 23 significant aerosol growth is observed. Rapid brown carbon formation is seen with aqueous 24 aerosol particles at >80% RH. Brown carbon formation slows at 50 - 60% RH and when the 25 aerosol particles are acidified with sulfuric acid, but stops entirely only under dry conditions. This chemistry may therefore contribute to brown carbon production in cloud-processed pollution 26 27 plumes as oxidizing VOCs interact with SO₂ and water.

Keywords: cloud processing, sulfonate formation, secondary organic aerosol, sulfite oxidation

30 Introduction

31 Sulfur dioxide, a S(IV) compound emitted from volcanoes, coal-burning power plants, smelters, 32 and oil refineries, is oxidized to sulfate in the atmosphere with a lifetime of 4-12 h, as estimated from satellite retrievals.(1) This oxidation takes place mainly in cloudwater(2) or aqueous 33 aerosol,(3) where dissolved SO₂ reacts with dissolved oxidants, especially HOOH(4) and organic 34 35 peroxides.(5) The sulfate produced is a major component of submicron aerosol particles,(6) which 36 affects human health and climate. Atmospheric sulfate concentrations are typically correlated with 37 other products of oxidative cloud processing, such as oxalate ions,(7) produced from aqueous glyoxal oxidation. Recent observations of extremely rapid SO₂ oxidation at high RH during 38

pollution episodes over northeast China have highlighted remaining gaps in our understanding of
relevant oxidation pathways.(8-10)

41 Oxidation is not the only chemical reaction that dissolved SO₂ participates in. Glyoxal, like 42 other aldehydes, (4, 11) reacts rapidly with S(IV) compounds in aqueous media, reversibly forming 43 sulfonate adduct molecules(12) containing C-S bonds.(13) No further products are observed in 44 the dark when oxidants are excluded, leading to the conclusion that sulfonates serve only as 45 condensed-phase atmospheric reservoirs for aldehydes and S(IV), (14) thereby increasing SO₂ 46 partitioning(12) and glyoxal uptake into clouds.(15) However, when exposed to air, aqueous 47 mixtures of glyoxal and sodium sulfite (which have pH > 7 due to sulfite basicity) quickly produce 48 the redox-active, aromatic compound tetrahydroxybenzoquinone (THBQ)(16) at the air-water 49 interface. THBQ is a black precipitate, and is red when dissolved in aqueous solution, strongly 50 absorbing visible light ($\sigma = 3.7 \times 10^{-17}$ cm² at $\lambda_{max} = 485$ nm). THBQ is itself oxidized under basic 51 conditions to form the light-absorbing, redox-active species rhodizonic acid (RhA) and croconic 52 acid (CrA), of which the latter is more stable. This oxidation pathway is summarized in Scheme 53 1.

54 Redox-active quinone species have been implicated, along with transition metals ions, in the 55 widespread toxicity of atmospheric aerosol.(17-21) These compounds can generate free radical 56 oxidants in lung fluid for hours after inhalation, triggering adverse health effects.(20-25) Thus, it 57 is critical to determine to what extent reactions between glyoxal and SO₂ dissolved in aqueous 58 aerosol and cloud droplets might produce quinone species and brown carbon under 59 atmospherically relevant conditions. In this work we performed a series of aerosol and bulk 60 aqueous experiments involving glyoxal and dissolved SO₂ at pH < 7. We find that fast production 61 of brown carbon products is maximized in aqueous aerosol at high RH and slowed by acidification

62 of the aerosol phase. In bulk phase simulations at pH < 6, sulfonates with C-S bonds and odd 63 carbon numbers are major products detected by ESI-MS, along with small yields of quinones 64 generated at pH \ge 5.

65

66 Scheme 1: Summary of Reported Glyoxal + Sulfite Reaction Products Under Basic Conditions



67

ROS = reactive oxidant species. Light absorbing products (middle row): THBQ =
tetrahydroxybenzoquinone; RhA = rhodizonic acid; CrA = croconic acid. Non-light-absorbing
products (bottom row): CHH = cyclohexahexanone; CPP = cyclopenta-pentanone. All structures
from ref (*16*). The pH 5 oligomer pathway is proposed in this work.

72

74 Materials and Methods

All chemicals were used as received from Sigma-Aldrich except as noted.

Bulk studies. Sodium sulfite (Spectrum) solutions were pH-buffered with formic acid (pH 3), 76 77 acetic acid (pH 4 and 5.4) or malonic acid (pH 6) and reacted with aqueous glyoxal generated by 78 hydrolysis of solid glyoxal trimer dihydrate (GTD, Fluka) at 0.25 M concentrations. Standards of 79 tetrahydroxybenzoquinone (THBQ), rhodizonic acid (RhA, disodium salt), croconic acid (CrA, 80 disodium salt), and glyoxal bis-disulfite adduct (GBDS, disodium salt hydrate) were made in N2-81 bubbled deionized water to minimize the presence of dissolved oxidants. Sodium sulfite - glyoxal 82 reaction mixtures and standards were analyzed as a function of time by diode array UV/vis absorbance spectrometry (HP 8452A) and/or negative-mode electrospray ionization mass 83 84 spectroscopy (Thermo LTQ).

85 Aerosol flow chamber studies. A schematic of the experimental system is shown in Figure S1. 86 Aqueous aerosols were generated from 0.05 - 0.15% w/w aqueous sodium sulfite solutions (TSI 87 3076 atomizer), pH buffered in one experiment with sulfuric acid. In certain experiments the 88 aerosol flow was diffusion dried. Hydrogen peroxide gas was added in certain experiments by bubbling 0.2 L/min N2 through a 30% w/w solution. A continuous flow of glyoxal gas was 89 90 generated by flowing N2 over a mixture of solid GTD and P2O5 heated to 45 - 80 °C.(26) Glyoxal 91 production was monitored by absorbance at 405 nm using a cavity ringdown (CRD) spectrometer 92 and a cross section of 4.491×10-20 cm2 molecule-1.(27) The various inlet flows totaling 2.5 L/min 93 were mixed at the inlet to a 2.5 L Pyrex vessel, such that the average reaction time was 1 min. 94 Aerosol particles exiting the reaction vessel were diffusion dried and monitored by Q-AMS 95 (Aerodyne), CAPS-ssa (450 nm, Aerodyne), scanning mobility particle sizing (SMPS), CRD (530 96 nm), and photoacoustic spectrometers (405 and 530 nm), with total sampling flows set to match

97 the inlet flows (2.5 L/min). Two-minute averaging of 1-Hz CAPS data allows albedo to be 98 measured with precision typically between ± 0.001 and ± 0.005 , while geometric mean diameters 99 extracted from SMPS size distributions have ± 2 nm precision. RH sensors monitored humidity 100 levels at the aerosol inlet, chamber outlet, and dried chamber outlet flows.

101

102 **Results and Discussion**

103 Bulk studies. In cloudwater (or other aqueous samples with pH between 2 and 7), acid-base 104 equilibria will cause dissolved SO₂ to exist mainly as bisulfite ions (HSO₃-) since H₂SO₃ has pK_a 105 values of 1.9 and 7.2. The initial reaction between glyoxal and bisulfite is rapid. For example, in 106 N₂-bubbled pH 4 buffer, 0.1 M HSO₃-reacted with 0.1 M glyoxal with $t_{1/2} = 9$ s (Figure S2). These 107 aqueous-phase concentrations are 40 and 80x higher, respectively, than estimated equilibrium 108 concentrations in atmospheric cloud droplets at pH 6.(12) The measured half-life implies a rate 109 constant k = 1.11 s-1M-1, which is lower than (but within a factor of 7 of) earlier measurements 110 conducted at pH = 3.26 and 0.0015 M.(12) This difference could be due to dehydration of glyoxal 111 becoming rate-determining above pH 3.26.(12) The reversible formation of bisulfite dimer ions 112 does not appear to be favorable enough to impact reaction rates even in our highest concentration 113 experiments (Figure S3).(28)

The pH-dependence of brown carbon products generated in 2 days by glyoxal + bisulfite reactions is summarized in Figure 1. At pH 3.1 or 4.2, reaction products absorb light only in the UV range, while at pH 5.4 (Figure S4) and 6.0, visible light absorbers were produced. Importantly, at no point in the glyoxal + sulfite reaction in experiments at pH \leq 5.4 do the characteristic visible absorbance bands of THBQ or RhA appear (both with $\lambda_{max} = 485$ nm, Figures S5 and S6), suggesting that if these species are formed they are either reactive intermediates or minor products. 120 At pH 6.0, CrA can be quantified by its absorbance and fluorescence bands (Figures 1 and S7) at 121 a 0.02% yield. The major products formed at mildly acidic pH are therefore different than the 122 hydroquinones that form at high yields under basic conditions.



Figure 1: UV-vis absorbance spectra of 0.25 M glyoxal + 0.25 M sodium sulfite solutions after
2 days reaction time, buffered to initial pH 6.00 (blue), pH 5.42 (green), pH 4.20 (orange), and pH
3.13 (red). Spectrum of 64 μM aqueous croconic acid standard is overlaid for comparison.

127

The observed product differences may be due to pH-dependent hydroquinone instability. Both THBQ and RhA are unstable in aqueous solution at pH 5.5, having respective lifetimes of only 42 and 70 min (Figures S5 and S6) even after initial N₂ bubbling to remove oxidants. A 1 mM THBQ solution aged at pH 3 for 24 h turned from red to yellow and showed an unmistakable oligomer pattern(*29*) when analyzed by ESI-MS, with dominant peaks in the C₁₀-C₁₁ mass range (Figure

133 S8). Thus, it is plausible that glyoxal + sulfite reactions under slightly acidic conditions produce
134 a small amount THBQ as an intermediate, which then forms other light-absorbing oligomers.

135 To better characterize glyoxal + HSO₃- chemistry, a pH 5.4 reaction sample was analyzed by 136 negative ion mode ESI-MS after 12 d reaction time in capped vials. Major peaks detected are 137 listed in Table 1, along with proposed peak assignments. Even after 12 d, glyoxal and its self-138 reaction oligomers are responsible for 5 peaks, including 3 of the 8 largest peaks. Bisulfite adducts 139 of these molecules (sulfonates) make up another 5 of the largest 15 peaks, and the HSO₄ (bisulfate, 140 m/z 97) peak is small, evidence that oxidation of the sample has been minimal. Glyoxal 141 monobisulfite (GMBS), m/z 139, and glyoxal dibisulfite, m/z 221, major products observed by 142 Olson and Hoffmann, (12) are the 2nd and 14th largest peaks detected in the mass spectrum, 143 respectively. The ESI-MS peak areas of these sulfonates were largest upon initial measurement (t 144 = 3 h) and showed a slow downward trend thereafter (Figure S9), evidence that they are quickly 145 formed and quite stable, consistent with prior work.(12)

146

Table 1: Peak Ions Detected in Solution Containing 0.25 M Glyoxal and NaHSO₃ at pH 5.42 After
13 d Reaction Time.

<i>m/z</i> of detected ion	assigned formula	identity	product generation	peak size ranking	proposed neutral structure
73	C2HO3-	glyoxylic acid	1 st	22	ОН
75	C2H3O3-	glyoxal monohydrate GX.H2O	reactant	7	о он он

81	HSO ₃ -	bisulfite	reactant	13	
97	HSO ₄₋	bisulfate	1st	18	
111	CH3SO4-	hydroxymethane -sulfonic acid, HMS	3rd	3	но
133	C4H5O5-	GX2.H2O	reactant oligomer	8	HO HO
139	C2H3SO5-	GX.HSO3- GMBS	1st	2	
151	C4H7O6-	GX2.2H2O	reactant oligomer	11	НО ОН ОН
155	C2H3SO6-	glyoxylic acid.HSO3-	2nd	10	OH OH OH OH
157	C2H5SO6-	GX.HSO3H2O	1 st	4	но ОН ОН ОН ОН ОН
169	C3H5SO6-	GX-HMS	3rd	1	О ОН ОН ОН ОН ОН ОН ОН
171	C2H3SO7-	HOO-GX-HSO3-	2nd	6	но он он он

209	C6H9O8-	GX3.2H2O	reactant oligomer	5	но о о он
215	C4H7SO8-	GX2.HSO3-	1st	9	HO OH HO OH
221	C2H5S2O8-	GX.2HSO3	2nd	14	
267	C8H11O10-	GX4.2H2O	reactant oligomer	16	
273	C6H9SO10-	GX3.HSO3-	1st	15	
335	C10H7O13-	CrA2.3H2O	5th	minor	

149 Notes: While m/z 171 is also the unit mass of THBQ, the dominant product at high pH, the low 150 yields of THBQ determined by spectrophotometry in the previous section means that THBQ 151 cannot contribute more than 1% to the m/z 171 ESI-MS signal. Similarly, RhA contributes less 152 than 0.05% of the signal at m/z 169. These hydroquinone products are too unstable in aqueous 153 solution to build up to high concentrations at long reaction times. The ESI-MS spectrum is shown 154 in Figure S3.

¹⁵⁶ Other detected peaks show evidence of oxidation, presumably caused by ambient oxidants 157 introduced when reaction vials were sampled (every 2 d on average). A small peak at m/z 73 is 158 likely due to glyoxylic acid, formed by oxidation of glyoxal monomer. A minor peak at m/z 335, 159 which was the largest peak in the 24-h aged THBQ solution (Figure S8), may be a light-absorbing

160 CrA dimer trihydrate formed via THBQ. If an OH radical abstracts a proton from GMBS, the 161 addition of O₂ produces a peroxyacid sulfonate molecule at m/z 171. OH radical addition converts 162 GMBS to a carboxylic sulfonic acid $(m/z \ 155)$, which could also be produced by reaction between 163 glyoxylic acid and bisulfite. OH radical oxidation of this carboxylic sulfonic acid followed by 164 decarboxylation is the likely path to two major products with odd numbers of carbon: 165 hydroxymethanesulfonate (HMS, m/z 111, 3_{rd} largest peak) and a C₃ sulfonate (m/z 169, largest 166 peak) likely formed via glyoxal addition to the HMS radical (as summarized in Scheme S1). The 167 formation of HMS was confirmed by 1H NMR through its characteristic CH₂ peak at 4.23 ppm 168 (Figure S10).

Peak areas were examined as a function of reaction times between 3 h and 15 d for three oxidized products (m/z 97, bisulfate ion; m/z 169, C3 sulfonate; and m/z 171, C2 peroxysulfonate). All 3 slowly increased with reaction time (Figure S9). In contrast, the peak areas of four 1st-generation sulfonate products generated by direct (non-oxidative) reactions slowly decreased over the same time period, presumably after being formed earlier in the reaction (t < 3 h). This is consistent with oxidation reactions converting some 1st generation sulfonates into other sulfonate species, at rates limited by the supply of oxidant species.

176 It is notable that OH radical oxidation of GMBS appears to occur much more readily on the 177 organic end of the molecule rather than at the sulfite group. The small size of the bisulfate ion 178 peak at m/z 97 suggests that sulfur is protected from oxidation via incorporation into sulfonates, 179 and that sulfonate C-S bonds are largely preserved during the limited oxidation of this sample. 180 This is consistent with earlier studies that noted that while S(IV) can be oxidized by HOOH and 181 ozone, once S(IV) is converted to HMS it does not react appreciably with either oxidant.(*11*)

183 Flow chamber studies. The ability of glyoxal gas to cause rapid browning of Na₂SO₃ aerosol 184 particles was tested in a series of 7 flow chamber experiments with 1-min residence times (Table 185 2). Significant browning was observed at 405 and 450 nm by photoacoustic and cavity-attenuated 186 phase-shift spectroscopy, respectively, in experiment 1 (Figure 2), where a constant flow of 187 aqueous Na2SO3 aerosol was mixed in this chamber with a smaller flow of dry N2 / glyoxal to 188 achieve 80% RH and 100 ppb glyoxal after mixing. In experiment 6, the only experiment with a 189 gas-phase source of S(IV) and the only experiment with NaCl instead of Na₂SO₃ seed particles, no 190 gas-phase reaction between SO₂ and glyoxal was observed. However, NaCl seeds have been 191 shown to catalyze SO₂ uptake, (30) and once aqueous NaCl seed particles were added, the second-192 lowest 450 nm albedo of 0.96 was achieved (Figure S11). AMS organic aerosol signals were also 193 clearly present as soon as NaCl seeds were added. These observations suggest that both SO2 and 194 glyoxal were taken up by aqueous NaCl aerosol particles, where they reacted to form measurable 195 brown carbon within on a 1-min timescale. In all other experiments (2-5 and 7), less browning 196 was observed despite generally higher glyoxal concentrations and high SOA growth via glyoxal 197 uptake. Browning in these experiments was likely slowed by much lower levels of aerosol-phase 198 water (Expts. 3-5 and 7), or by the presence of 60% mole fraction aerosol-phase sulfuric acid 199 (Expt. 2, Figure 3), which converted SO₃₂- ions to a mixture of less reactive(31) HSO₃- and H₂SO₃. 200 Experiments 1 and 2 are discussed in detail below.

expt.	figure	PAS _{abs} , 405 nm (Mm-1)	albedo, 450 nm	[GX] (ppb)a	[Na2SO3] (µg/m3)	% RH	AMS Org/SO4 ratio	growth, (µg/m3)	notes
1	2	160	0.940	110	80 wet	80	0.9	55	
2	3	2	0.986	90	40 wet	80	0.4	24	Acidified pH 4b (2:3 Na2SO3 / H2SO4 mole ratio)
3	S12	15	0.988	450	200 wet	50	0.9	150	
4	S12	10	0.980	500	180 wet	60	1.4	310	in air
5	S13	0.2	0.983	700	40 wet	50	1.0	52	HOOH (g) present
6	S11	С	0.960	~100d	Wet NaCl	80	Org/Cl = 3	С	SO2 (g) from Na2SO3 (aq) bubbler
7	S14	<0.1	0.983	2500	420 dry	60-80	2.0	210	In 7% O ₂ , 93% N ₂

203 Table 2: Summary of Flow Chamber Experiments: Sodium Sulfite Aerosol + Glyoxal (g)

*Not measured. PAS = photoacoustic spectroscopy. GX = peak glyoxal conc. RH = relative
humidity. AMS = aerosol mass spectrometer. Aerosol particles had 1 min lifetime in flow
chamber. a: peak glyoxal concentrations are shown. b: before atomization into aerosol. c: no
measurement. d: estimated from glyoxal source temperature.

208

In experiment 1 (Figure 2), deliquesced Na₂SO₃ aerosol were sent through the flow chamber in humidified N₂ from 2:45 until 3:36 pm. At 2:50 pm the dry N₂ flow was routed through a heated glyoxal source, rapidly reaching a calculated steady-state concentration of ~110 ppb glyoxal in the mixing chamber. (Actual glyoxal concentrations may have been significantly lower due to uptake by chamber walls at 80% RH.) Glyoxal addition caused SMPS particle diameters to increase by 20% and particle masses to increase by 70%, indicating significant and rapid glyoxal uptake. At the same time, albedo at 450 nm declined from 0.97 to 0.94, indicating brown carbon formation

216 by glyoxal + sulfite reactions on a 1-min timescale. Interestingly, AMS signals for both organics 217 and sulfate rose upon glyoxal addition, such that their ratio remains near 1. While sulfonates are efficiently broken down in the AMS inlet into organic and sulfate fragments,(3) the formation and 218 219 destruction of sulfonates would not be expected to increase sulfate signals. Instead, this increase 220 in sulfate signal indicates that the uptake of glyoxal (and formation of sulfonates) caused a 221 significant shift in the effective Henry's Law equilibrium of S(IV) (SO₂ and sulfite) towards the 222 aqueous phase, as predicted by Olson & Hoffmann.(12) The glyoxal source was turned off at 3:15 223 pm (Figure 2), but albedo at 450 nm, aerosol absorbance at 405 nm, particle size, and AMS particle 224 chemistry all remained constant, likely due to the preservation of glyoxal steady-state concentrations caused by equilibration from the glass walls in the humid chamber. (32, 33)225

226 At 3:36 pm, the aerosol source was routed through a diffusion dryer, generating effloresced 227 instead of aqueous Na₂SO₃ aerosol even as RH remains between 60 and 84% for 20 more min due 228 to water equilibrium from the walls. (Solid Na₂SO₃ aerosol deliquesces above 85% RH.(31)) The 229 change in particle phase from liquid to solid caused SMPS particle diameter to increase by an 230 additional 11%. This suggests that glyoxal uptake is enhanced by oligomer formation in the high 231 concentration environment of the surface water layer, such that particle growth is greater than it 232 was on deliquesced aerosol. This effect has been seen in prior studies involving glyoxal.(34) At 233 the same time, the switch to solid aerosol particles coincided with a 450 nm albedo rise from 0.94 234 to 0.96, and a 60% decline in absorbance at 405 nm, both indicative of less brown carbon formation 235 during the 1 min chamber residence time. This effect is likely due to aerosol-phase diffusion 236 limitations that limit the supply of sulfite ions to the surface water layer where glyoxal is reacting. 237 These observations show that glyoxal and sulfite ions can react in aqueous aerosol (and to a lesser 238 extent in surface water layers on solid aerosol particles) to rapidly form brown carbon products.



2: Figure Gas-phase glyoxal uptake experiment 1: aqueous Na₂SO₃ aerosol (before 3:35 pm) and effloresced Na₂SO₃ aerosol in flowing chamber (1 min residence time). Top: CAPS single-scattering albedo of dried aerosol at 450 nm (red triangles), recorded immediately after instrument baseline taken through filter; photoacoustic absorbance (Mm-1, blue line) of dried aerosol at 405 nm. Middle: **SMPS** geometric mean diameters aerosol of dried (black circles); CAPS-ssa inletdilution-corrected 2-minaveraged extinction (green line), scattering (red line). **Bottom**: Glyoxal (g) concentrations in ppb, calculated based on dilution and cavity ringdown spectroscopic measurements made at 405 nm at the chamber inlet and no wall equilibria (purple line, left humidity axis): relative measured at chamber outlet (black line): dilution

274 corrected aerosol mass spectrometer loadings by category (left axis; water, light blue; sulfite, red; 275 organics, green); and organic / sulfite ratio (gray dotted line, right axis). Vertical dotted lines mark 276 beginning and ending of glyoxal addition, the switch from aqueous to deliquesced Na₂SO₃ seed 277 aerosol, and the start of 2nd glyoxal addition, as labeled. AMS water signals are qualitative due 278 to drying in the low-pressure inlet; water signals are included to show that particles pass through 279 the chamber and the AMS sampling line fully dry only after ~4:10 pm.

4:00 PM

Time

239

100

50

0

280

3:00 PM

PAS abs,

405 nm (Mm-1)

450 nm Ext.

Scat.

Organic / SO4 Ratic

273

3

2

5:00 PM

281 After 27 min. of dry gas and aerosol input (*i.e.*, \sim 4:00 pm in Figure 2), the flowing reaction cell 282 surfaces were depleted of water, and RH in the chamber outflow declined to < 10%. Under these 283 fully dry conditions, Na₂SO₃ aerosol passed through the glyoxal-containing chamber without 284 browning: aerosol albedo (450 nm) promptly rose to 1.00 and aerosol absorbance (405 nm) 285 dropped to zero. At the same time, AMS sulfite signals declined by approximately 50%, back to 286 pre-glyoxal-addition levels. These changes indicate that glyoxal and sulfite cannot react in the dry 287 aerosol phase on a 1 min timescale, such that sulfite gas-particle partitioning is no longer perturbed 288 and brown carbon is no longer formed.

289 Glyoxal uptake to Na2SO3 particles did not end at this point, however. Although SMPS 290 geometric mean diameters of the dried aerosol declined from 81 to 66 nm upon chamber drying, 291 66 nm is still 6% larger than initial pre-glyoxal diameters. Furthermore, the organic AMS signal 292 actually rises after the RH drops to below 10%, peaking at 16:09 just as the AMS water signal 293 drops to zero. AMS data collected during this period (Figure S15) shows prominent increases at 294 m/z 18, 44, 45, and 46 (likely oxalic acid), and relative increases in minor peaks at m/z 80 and 98 295 (sulfuric acid). These increases suggest that glyoxal and sulfite ions are more susceptible to 296 oxidation under dry conditions when they do not react with each other. When glyoxal addition 297 was restarted in the fully dried chamber (100 ppb, 4:33 pm), a statistically significant 5% increase 298 in aerosol diameters was observed, but albedo did not change. This is further evidence that a 299 small amount of glyoxal can be taken up by residual surface water on dry Na₂SO₃ aerosol, but no 300 brown carbon is formed without an aqueous phase.

Experiment 2 (Figure 3) was conducted at the same RH as experiment 1, but with aqueous-phase sulfite ions converted to an equimolar mix of HSO₃- and H₂SO₃ by the addition of sulfuric acid to pH 4.0. (Na₂SO₃ / H₂SO₄ were mixed at a 2:3 mole ratio). When 90 ppb glyoxal was added to the

304 flowing chamber at 2:28 pm, once again we observed a decline in albedo, an increase in SMPS 305 particle mean diameter, and proportional (4×) increases in AMS signals for sulfate and organics, all occurring on a 1-min timescale (the residence time of aerosol particles in the flowing chamber). 306 307 While the increases in particle size and AMS signals are comparable in magnitude to experiment 308 1 with non-acidified sulfite aerosol, the albedo decline is at least a factor of 2 smaller. This 309 suggests that glyoxal uptake and reaction with HSO3- to form sulfonates is still fast, but brown 310 carbon formation is slowed, consistent with the pH-sensitivity observed in bulk-phase 311 experiments.

312 The change in detected ions upon glyoxal addition is summarized in Figure S16. Because of the 313 proportional 4× increase in AMS sulfate and organics, most signals fall on a new 4:1 line, with 314 m/z 29, 30, 31 and 58 (glyoxal and its EI major fragments) being prominent organic ions detected. 315 A few low-abundance ion signals increased by an order of magnitude, however, including m/z 80 316 and 97 (bisulfate ion) and 111 and 112 (HMS). These are some of the same products detected in 317 bulk, acidified glyoxal + S(IV) solutions after very long reaction times, especially given that the 318 glyoxal hydrates and oligomers (and perhaps adducts with S(IV)) detected by ESI-MS would 319 undergo thermal breakdown to glyoxal monomer(35) during vaporization in the high-temperature 320 AMS inlet.

The addition of HOOH gas from a bubbler at 2:45 pm (Figure 3) caused further growth in particle size, but brown carbon formation declined slightly (450 nm albedo increased). Starting at 3:20 pm, three pulses of very high (ppm) levels of glyoxal were introduced to the flowing chamber, but these additions caused only minor changes in physical or optical properties of the aerosol (max change in albedo = -0.007, max change in particle mass = +10%). It appears that the effects of



Figure 3: Gas-phase glyoxal uptake experiment 2: aqueous pH 4 Na2SO3 / H2SO4 aerosol in flowing chamber (1 min residence time). Top: CAPS single-scattering albedo of dried aerosol 450 at nm (red triangles), recorded immediately after instrument baseline taken through filter, bars show error standard deviation of the 2-min. averages; Middle: **SMPS** geometric mean diameters of dried aerosol (black circles); CAPS-ssa inletdilution-corrected 2min-averaged extinction (green scattering squares), (red line). **Bottom**: Glyoxal (g) concentrations in ppb after dividing by 10,



364

In summary, fast particle growth due to glyoxal uptake is most pronounced in the surface water layer of solid Na₂SO₃ aerosol at high RH, but brown carbon formation on a 1-min timescale is

326 glyoxal exposure on S(IV)-containing aerosol particles became saturated due to chamber wall 327 uptake / wall equilibrium at 80% RH.

367 maximized when Na₂SO₃ aerosol particles are deliquesced. Some glyoxal uptake onto solid 368 Na₂SO₃ aerosol is observed even under completely dry conditions (< 5% RH), but brown carbon 369 is not formed. In aqueous Na2SO3 aerosol particles acidified to pH 4, aerosol growth and 370 partitioning of S(IV) to the aqueous aerosol phase are observed that are similar in magnitude to 371 unacidified experiments, but brown carbon formation is lessened. On the other hand, while in bulk 372 liquid studies no visible light absorbers were produced at pH 4.2 or less over 2 d reaction times, in 373 aerosol particles generated from pH 4 solution measurable absorbance at 450 nm was generated in 374 only 1 min. Thus, it appears that glyoxal uptake, reaction with S(IV) to form sulfonic acids, and 375 to a lesser extent brown carbon formation from these precursors can all occur rapidly at acidic pH 376 in aqueous aerosol. Reaction acceleration in aerosol particles (relative to bulk liquid experiments) 377 has been observed in many other systems, (30, 36-40) and is likely caused by surface-reactive 378 species. Such reactions are likely to show a particle size dependence. We note that dried particle 379 sizes in this study (50-80 nm) were slightly smaller than atmospheric accumulation mode particles 380 (>100 nm).

381

382 Atmospheric Significance

In the flowing aerosol experiments with deliquesced Na₂SO₃ aerosol at 80% RH, glyoxal concentrations were \sim 110 ppb, and reaction times were 1 minute. If the initial reaction between glyoxal (g) and sulfite ions is the rate-limiting step in brown carbon formation, and since that step is first order with respect to glyoxal,(*12*) then we can scale down to atmospheric conditions using an "integrated glyoxal" concept that is analogous to "integrated OH" in oxidation studies using OH reactors. Using this concept, the 1-minute lifetime in our aerosol experiments corresponds to about 18 h of glyoxal exposure at atmospheric concentrations of 100 ppt. Thus, given that glyoxal is efficiently scavenged by aqueous aerosol, it appears that glyoxal exposure could reasonably depress the albedo of atmospheric aerosol particles in humid, polluted regions where aerosol particles are liquid and contain significant quantities of dissolved SO₂.

393 The pH dependence seen in our aerosol and bulk studies suggests that brown carbon formation 394 by glyoxal-SO₂ reactions would be most pronounced in aerosol that are neutralized by ammonia 395 and amines, a situation that is becoming more common in the atmosphere as ammonia and amine 396 emissions are uncontrolled and rising. It should also be noted that aqueous reactions between 397 glyoxal and ammonia and amines under neutral conditions have also been identified as a source of 398 brown carbon.(41-45) Aqueous-phase glyoxal + S(IV) reactions produce hydroquinones at very 399 low yields at pH \leq 6, but oligomer products in this pathway may nevertheless be contributing 400 significantly to light absorption.

401 Recent work has suggested that HMS is the most abundant organosulfate compound in the 402 aerosol phase during Chinese winter haze events.(3) Because it can be easily converted to sulfate 403 by typical AMS and ion chromatography methods, HMS may be responsible for 1/3 of the 404 unexplained "sulfate" formation reported there. In this work we have shown that HMS, long 405 assumed to be formed only from formaldehyde and S(IV), (3, 46) is also an important product of 406 aqueous phase glyoxal + S(IV) reactions under mildly acidic conditions. Cloudwater 407 concentrations of formaldehyde and glyoxal in the atmosphere are comparable, with measured 408 formaldehyde:glyoxal ratios between 0.7 and 4.(47-50) Given that bisulfite reaction rates with 409 formaldehyde in bulk aqueous solution are 6000x faster than with glyoxal, (14) it seems at first 410 glance unlikely that glyoxal + S(IV) could be a significant source of atmospheric HMS. However, 411 glyoxal + S(IV) reactions begin with a nucleophilic attack by the S(IV) lone pair on a non-hydrated 412 carbonyl functional group. Glyoxal preferentially forms a monohydrate (with a non-hydrated

413 carbonyl) at air-water interfaces, causing nucleophilic attack reaction rates on glyoxal to be 414 enhanced by orders of magnitude in aerosol experiments, compared with bulk solution 415 measurements.(*36*) Even in Los Angeles, where cloud formaldehyde levels are typically in excess 416 of S(IV) levels,(*51*) Richards *et al.* determined that only 1/3 of the S(IV) in clouds and fog had 417 reacted with formaldehyde to form HMS, based on measurements of free formaldehyde.(*52*) It 418 thus appears that there is enough S(IV) in clouds and aqueous aerosol to react with glyoxal, and it 419 is at least possible that glyoxal + S(IV) reactions contribute to observed HMS concentrations.

Field measurements of the glyoxal sulfonate C₂ adduct molecule could establish an upper limit on the size of this HMS source. Based on comparative ESI-MS and NMR peak heights in this work, the ratio of glyoxal sulfonate C₂ adduct to HMS produced in glyoxal + sulfite reactions at pH 5 is 9 ± -4 even after long reaction times. Production of HMS at moderate yield by glyoxal + S(IV) reactions may help explain the correlation between HMS, oxalic acid, and sulfate noted in field measurements, where elevated levels are associated with aged, cloud-processed pollution plumes.(*13*)

428	Supporting Information Available: Experimental schematic, additional UV/vis, fluorescence,
429	NMR, and mass spectra for reaction mixtures and standard compounds, AMS spectra comparisons,
430	summaries of flow chamber experiments $3 - 7$, and reaction scheme.

- 431
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