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Rapid formation of intense haze episodes via aerosol–boundary layer feedback in Beijing

Yonghong Wang[1,2], **Miao Yu**[3], **Yuesi Wang**[1,6], **Guiqian Tang**[1], **Tao Song**[1], **Putian Zhou**[2], **Zirui Liu**[1], **Bo Hu**[1], **Dongsheng Ji**[1], **Lili Wang**[1], **Xiaowan Zhu**[1], **Chao Yan**[2], **Mikael Ehn**[2], **Wenkang Gao**[1], **Yuepeng Pan**[1], **Jinyuan Xin**[1,6], **Yang Sun**[1], **Veli-Matti Kerminen**[2], **Markku Kulmala**[2,4,5], and **Tuukka Petäjä**[2,4,5]

[1]State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry (LAPC),
Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China
[2]Institute for Atmospheric and Earth System Research/Physics, Faculty of Science, P.O. Box 64,
00014 University of Helsinki, Helsinki, Finland
[3]Institute of Urban Meteorology, China Meteorological Administration, Beijing, China
[4]Joint international research Laboratory of Atmospheric and Earth SysTem sciences (JirLATEST), Nanjing University,
Nanjing, China
[5]Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering,
Beijing University of Chemical Technology (BUCT), Beijing, China
[6]Centre for Excellence in Atmospheric Urban Environment, Institute of Urban Environment, Chinese Academy of Science,
Xiamen, Fujian 361021, China

Correspondence: Yuesi Wang (wys@mail.iap.ac.cn) and Markku Kulmala (markku.kulmala@helsinki.fi)

Abstract. Although much effort has been put into studying air pollution, our knowledge of the mechanisms of frequently occurring intense haze episodes in China is still limited. In this study, using 3 years of measurements of air pollutants at three different height levels on a 325 m Beijing meteorology tower, we found that a positive aerosol–boundary layer feedback mechanism existed at three vertical observation heights during intense haze polluted periods within the mixing layer. This feedback was characterized by a higher loading of $PM_{2.5}$ with a shallower mixing layer. Modelling results indicated that the presence of $PM_{2.5}$ within the boundary layer led to reduced surface temperature, relative humidity and mixing layer height during an intensive haze episode. Measurements showed that the aerosol–boundary layer feedback was related to the decrease in solar radiation, turbulent kinetic energy and thereby suppression of the mixing layer. The feedback mechanism can explain the rapid formation of intense haze episodes to some extent, and we suggest that the detailed feedback mechanism warrants further investigation from both model simulations and field observations.

1 Introduction

With the rapid economic growth and urbanization, an increasing frequency of haze episodes along with the air pollution has become of great concern in China during the last decade (Cao et al., 2016; Huang et al., 2014; Kulmala, 2015; Y. H. Wang et al., 2014, 2015). For example, during December 2016 a series of intense haze episodes took place in eastern China, characterized by surface $PM_{2.5}$ concentrations exceeding $500\,\mu g\,m^{-3}$ in several measurement sites in Beijing and its surrounding sites (http://www.mep.gov.cn/gkml/hbb/qt/201701/t20170102_393745.htm, last access: 11 December 2019). Severe air pollution has serious effects on human health. A recent study reported that the particulate matter has significantly decreased the life span of residents by as many as 5.5 years in northern China (Chen et al., 2013). On a global scale, the air pollution was estimated to cause over 3 million premature deaths every year (Lelieveld et al., 2015).

Increased emissions from fossil fuel combustion due to vehicle traffic, industrial activities and power generation, along with exceptionally strong secondary aerosol formation, were

thought to be responsible for these haze episodes (Cheng et al., 2016; Huang et al., 2014; Pan et al., 2016; Petäjä et al., 2016; Zhang et al., 2015; Zhao et al., 2013). Meanwhile, the formation of intense haze episodes was considered to be affected by meteorological conditions (J. Wang et al., 2014; Quan et al., 2013; Zheng et al., 2016). For example, the mixing layer height is a key parameter that constrains the dilution of surface air pollution, and the development of the mixing layer is highly related to the amount of solar radiation absorbed by the air and reaching the surface (Ding et al., 2016; Stull, 1988; Sun et al., 2013; Tang et al., 2016). By using field measurements combined with model simulation, a positive feedback between aerosol pollution, relative humidity and boundary layer was found to be important in aerosol production, accumulation and severe haze formation in Beijing (Liu et al., 2018). Wang et al. (2018) found that PBL schemes in their atmospheric chemistry models are not sufficient to describe the explosive growth of $PM_{2.5}$ concentrations in the Beijing–Tianjin–Hebei region due to the absence of an online calculation of aerosol–radiation feedback and/or a deficient description of extremely weak turbulent diffusion.

In this study, using unique measurements on the Beijing 325 m high meteorology tower, we show a clear relationship between mixing layer height and turbulent kinetic energy at the 140 m observation platform. We also present direct evidence of the feedback that relates the decreasing mixed layer height to increasing particulate matter concentrations, and this feedback is critical to the formation of intense haze episodes in Beijing.

2 Methods

2.1 Calculation of mixing layer height with ceilometer

The ceilometer was deployed in the yard of the IAP (Institute of Atmospheric Physics, Chinese Academy of Science), with a horizontal distance of around tens of metres from the 325 m meteorology tower. The mixing layer height was measured with the enhanced single-lens ceilometers from July 2009 to August 2012 (CL 31, Vaisala, Finland), which utilized the strobe laser lidar technique (910 nm) to measure the attenuated backscattering coefficient profiles. The detection range of the CL31 is 7.6 km with the report period of 2–120 s. Detailed information can be found in previous studies (Tang et al., 2016). Since the distribution of particle concentrations is uniform in the mixing layer and has significant differences between the mixing layer and free atmosphere, the height at which a sudden change exists in the attenuated backscattering coefficient profile indicates the top of the mixing layer height. Vaisala software product BL-VIEW was used to determine the mixing layer height by finding the position with the maximum negative gradient $(-\mathrm{d}\beta/\mathrm{d}x)$ in the attenuated backscattering coefficient profiles as the top of the mixing layer (Münkel et al., 2007).

2.2 Measurements of energy flux at the 325 m Beijing meteorology tower

The turbulent fluxes of sensible heat (Q_H), latent heat (Q_E) and turbulence kinetic energy (TKE) were measured at the 140 m level using the eddy covariance technique from July 2009 to August 2012. The raw data (10 Hz) of wind components (u, v and w) and sonic temperature (T_s) was recorded with three-dimensional sonic anemometers (Model CSAT3, Campbell Scientific Inc., Logan, Utah, USA), and of water vapour concentrations (q) with open-path infrared gas analysers (Model LI-7500, LiCor Inc., Lincoln, Nebraska, USA). The fluxes of heat (Q) were calculated as the covariance between the instantaneous deviation or fluctuations of vertical velocity (w_i') and their respective scalar (s_i') averaged over a time interval of 30 min:

$$Q = \overline{w's'} = \frac{1}{N}\sum_{i=1}^{N} w's', \tag{1}$$

where the overbar denotes a time average, N is the number of samples during the averaging time and the fluctuations are the differences between the instantaneous readings and their respective means. The TKE was calculated as follows (Stull, 1988):

$$\frac{\mathrm{TKE}}{m} = \frac{1}{2}\left(\overline{u'} + \overline{v'} + \overline{w'}\right) = \overline{e}, \tag{2}$$

where m is the mass (kg) and e is the TKE per unit mass ($\mathrm{m^2\,s^{-1}}$). A more detailed description of the calculation and post-processing of flux is provided elsewhere (Song et al., 2013).

2.3 Measurements of $PM_{2.5}$ concentration and gases at the 325 m Beijing meteorology tower

The mass concentrations of $PM_{2.5}$ at the 8, 120 and 280 m observation platforms were measured with three TEOM RP1400s simultaneously from July 2009 to August 2012 (Thermo Scientific, https://www.thermoscientific.com, last access: 11 December 2019). The resolution and precision of the instrument for 1 h measurements were 0.1 and $\pm 1.5\,\mathrm{\mu g\,m^{-3}}$, respectively. The filters were exchanged when the loading rates were approximately 40 %. The flow rate was monitored and calibrated monthly. The volume mixing ratios of ozone and NO_x were measured with 49i and 42i (Thermal Environment Instruments (TEI) Inc.), respectively (Y. H. Wang et al., 2014).

2.4 Experiment design

The model used in this study is the Weather Research and Forecasting (WRF) model (ARW, version 3.8.1; Skamarock et al., 2008). The simulation domain was centered in Beijing (39.0° N, 116.0° E) and implemented with one-nested grids with a resolution of 1 km. The number of grid cells was

460×403 for the domain in the east–west and south–north directions. The model run was initialized at 00:00 UTC (or 08:00 LST) on 16 November 2010 and integrated for 131 h until 10:00 UTC on 21 November 2010, including 48 h for spin-up. The initial conditions of the model and its outermost lateral boundary conditions, as well as the soil moisture field, were taken from National Centers for Environmental Prediction/National Center for Atmospheric Research Reanalysis data (resolution: $1° \times 1°$). The model physics schemes used include Thompson microphysical parameterization (Thompson et al., 2004); BouLac boundary-layer parameterization (Bougeault and Lacarrere, 1989); the RRTMG (Iacono et al., 2008) radiation scheme; and the Building Effect Parameterization (BEP) and Building Energy Model (BEM) schemes implemented in WRF that can more accurately describe three-dimensional urban land surface features and processes, including anthropogenic heat from buildings (Martilli et al., 2002; Salamanca and Martilli, 2010). The control and test experiments were performed separately to investigate impact of aerosol direct radiative forcing on surface temperature, relative humidity and development of boundary layer height. The control run (CTL) used the RRTMG radiation scheme which ignored the direct radiation effects of aerosol input. In the sensitivity test experiment, we add the aerosol input in the RRTMG scheme using Tegen climatology and urban-type aerosols during the sensitivity test.

2.5 Other supporting measurements

Total solar radiation was measured with a direct radiometer (TBQ-2, Junzhou, China). Direct radiation was measured with a direct radiometer (TBS-2, Junzhou, China). UV radiation in the range of 220–400 nm was measured using a CUV3 radiometer (USA). The estimated experiment errors for the three instruments are 3 %, 1 % and 2 %, respectively. The original data were obtained at 1 min intervals and the hourly average values were used in this study. The chemical composition of organic, sulfate, nitrate, ammonium and chloride in non-refractory submicron aerosol were measured during several campaigns with an Aerodyne High-Resolution Time-of-Flight Aerosol Mass Spectrometer from July 2009 to August 2012 (HR-ToF-AMS, Aerodyne Research Inc., Billerica, MA, USA). Detailed information about the instrument, calibration and data process have been introduced by Zhang et al. (2016). All these measurements were conducted at the IAP station.

3 Results and discussion

A typical intense haze episode occurred during the heating season in urban Beijing during 17 to 22 November 2010. This episode was associated with synoptic stagnation in the North China Plain (Fig. S1 in the Supplement) and was characterized by low wind speeds and irregular wind direction

(Fig. 1). Several meteorological variables had distinct temporal patterns during different stages of pollution, including reduced solar radiation and increased relative humidity during the most intense presence of haze (Fig. 1). The temporal patterns of $PM_{2.5}$ concentrations were very similar at the two lower measurement heights (8 and 120 m, Fig. 1d), even though the concentration was clearly highest close to the surface. The $PM_{2.5}$ concentration measured at 280 m behaved in a different way, especially during the most intense period of the haze when the mixed layer height was very low (Fig. 1e). The decoupling of the 280 m platform from the other two lower ones at low mixed layer heights is apparent in our 3-year measurement data set, especially when comparing O_3 and NO_x concentrations between the three measurement platforms (Figs. S2 and S5). During the haze period, the maximum $PM_{2.5}$ concentrations at 8, 120 and 280 m were 505, 267 and 339 µg m^{-3}, respectively. The higher maximum concentration at 280 m compared with 120 m can be ascribed to the transport of pollutants from the surrounding regions of Hebei and Tianjin provinces typical for polluted periods (Sun et al., 2013). The mixing layer height varied from 130 to 1640 m during the haze episode, ranging between about 200 and 500 m during the most intense period of the haze period on 18 November 2010 (Fig. 1e). The TKE was quite low during this intensive haze episode from 18 to 21 November, with an average value around 0.3 m^2 s^{-2}. However, the TKE increased significantly on the morning of 21 November as surface wind increased from 1.2 to around 6 m s^{-1}, which was possible due to the movement of a cold front as shown in Fig. S1.

The vertical distributions of attenuated backscatter density obtained from ceilometer measurements indicate vertical mixing conditions accompanied by an inversion layer and high relative humidity in the surface as shown in Fig. 2. The strong inversion and high relative humidity occurred on the morning of 18 November 2010, with a lapse rate of 2 K / 100 m, relative humidity of 78 % and a northerly wind speed of around 2 m s^{-1} detected by the vertical sounding. The turbulent kinetic energy at 140 m was reduced to around 0.1–0.7 m^2 s^{-2} due to decreased solar radiation, as presented in Fig. 1a. In this manner, the development of a mixing layer was significantly suppressed during the intense haze episode.

In order to demonstrate how the $PM_{2.5}$ modifies the surface temperature, relative humidity and development of the mixing layer height, we performed two numerical simulation experiments, using the WRF model as a tool. We took the measurements during the intensive haze episode shown in Fig. 1 as an example. As shown in Fig. 3a, the variation of temperature and relative humidity showed pronounced daily variations, with higher and lower values, respectively, during daytime in both test and control experiments. However, the presence of aerosol in the test experiment clearly showed decreased surface temperature and increased relative humidity. The presence of aerosol reduces downward radiation reaching the surface, as a result of which the surface
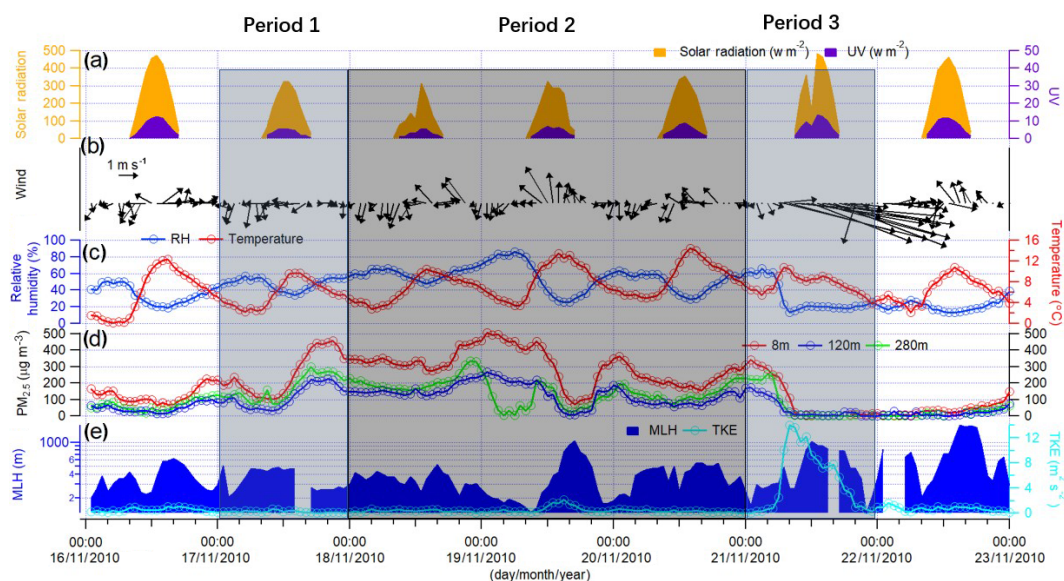
Figure 1. Measurements of (a) solar radiation and ultraviolet radiation at 8 m, (b) wind speed and direction at 8 m, (c) relative humidity and air temperature at 8 m, (d) mass concentration of $PM_{2.5}$ at 8, 120 and 280 m, and (e) mixing layer height at 8 m and turbulence kinetic energy at 140 m in the Beijing 325 m meteorology tower during an intensive air pollution episode in November 2010. The evolution of the air pollution episode can be divided into period 1 (clean period to air pollution accumulation period), period 2 (pollution period) and period 3 (pollution to clean period).
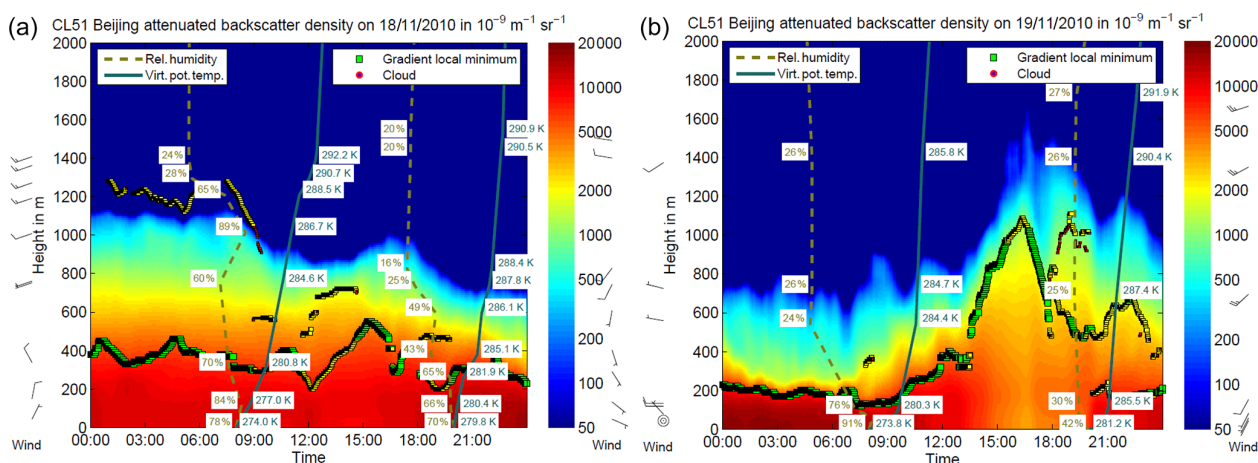
Figure 2. Observed attenuated backscatter density, calculated mixing layer height using ceilometer and vertical wind speed, wind direction, relative humidity, and virtual potential temperature using sounding data during 18 November (a) and 19 November (b). The black flags on the left- and right-hand sides of the figures stand for vertical wind speed and wind direction obtained from sounding measurements at 08:00 and 20:00 Beijing time, respectively. The circle on the left-hand side of the figure represents calm wind. The dotted yellow lines and solid green lines represent the vertical distribution of virtual potential temperature and relative humidity from sounding at 08:00 and 20:00, respectively. The yellow square and green square represent the first layer and second layer, respectively, and usually the first layer was used as the mixing layer height. The mixing layer height was determined from the local minimum of the backscatter density gradient, and the colour in the figure stands for backscatter density from the ceilometer. From both figures, we can clearly see that the mixing layer has an important role in regulating the distribution of air pollutants.

temperature and sensitive heat flux decrease and the development of the mixing layer height is suppressed (M. Li et al., 2017; Z. Li et al., 2017; Miao et al., 2016). Statistical results showed that the average relative humidity, surface temperature and mixing layer height were $8.2 \pm 3.4\,°C$, $40.5 \pm 11.6\,\%$

and $377.7 \pm 499\,m$, respectively, without the consideration of aerosol direct radiative forcing, whereas the consideration of aerosol directive radiative forcing changed these values to $7.1 \pm 3.1\,°C$, $40.6 \pm 11.7\,\%$ and $326.7 \pm 470.1\,m$, respectively. Our model results clearly demonstrate the pronounced role of
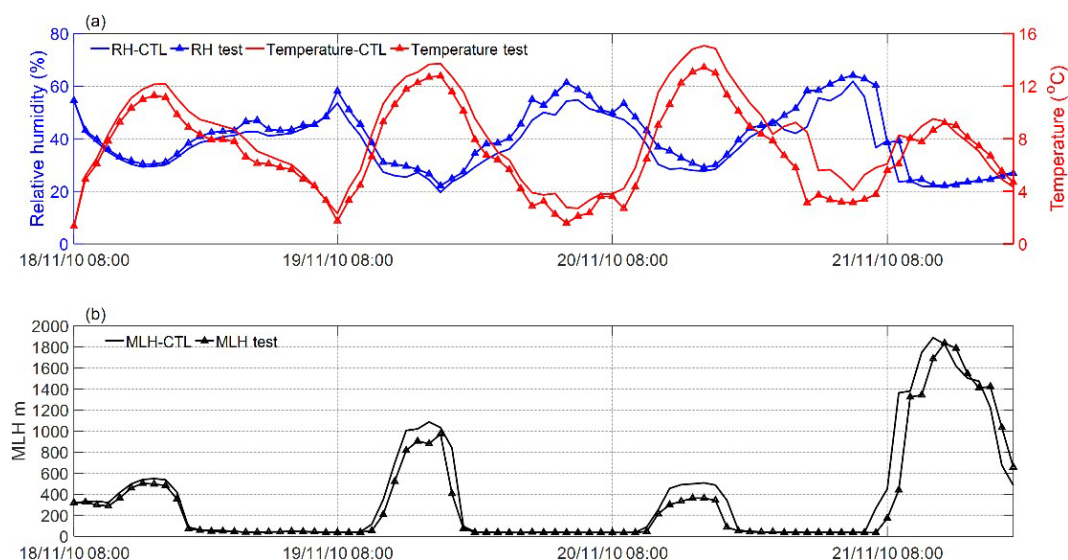
Figure 3. (a) Modelled variation of surface relative humidity and temperature and **(b)** mixing layer height during the intensive haze episode from 18 to 21 November 2010. The lines with triangles on them represent results from the test experiment, while the lines without triangles represent results from the control experiment. The control experiment was performed with the absence of aerosol direct radiative forcing in the RRTMG radiation scheme, while the test experiment was conducted with the presence of aerosol direct radiative forcing considered.

aerosol particles in reducing the mixing layer height during this haze pollution episode.

In order to further illustrate how the mixing layer height modifies $PM_{2.5}$ concentrations, we used 3 years of simultaneous winter-time air pollutant measurements in Beijing. We divided the observed $PM_{2.5}$ concentrations into highly polluted and less polluted conditions using a threshold value of $75 \, \mu g \, m^{-3}$ for $PM_{2.5}$ to distinguish between these conditions. This is consistent with the Chinese Environment Protection Bureau definition of a haze pollution event. With this threshold value, we found that 31 % and 69 % of the total measurement time corresponded to highly polluted and less polluted conditions, respectively. We plotted the $PM_{2.5}$ data as a function of the mixing layer height at the three observation heights (8, 120 and 280 m) during both highly polluted and less polluted conditions and fitted an exponential curve to these data based on best fitting (Fig. 4). The $PM_{2.5}$ concentration has a clear anti-correlation with the mixing layer height during the intense haze episodes. At all the measurement heights, the $PM_{2.5}$ concentration increased as the mixing layer height decreased, and this pattern was very strong under polluted conditions (Fig. 4). We also tested the reciprocal fitting function for the data (Fig. S8). It overestimated the $PM_{2.5}$ concentration when the mixing layer height was very low, as compared to the exponential fitting function (Fig. 4). This also indicates that a much higher $PM_{2.5}$ concentration is needed in order to obtain a very low mixing layer height without the positive feedback. This can also be supported by the root-mean-square error (RMSE) of these two fitting methods. The RMSE of the exponential fitting is

much smaller than the reciprocal fitting in any case (Table S1 in the Supplement).

It is worth noting that the increase was mainly from the $PM_{1-2.5}$ fraction that increased from 42 % to 65 % as the mixing layer height decreased from more than 1400 m to lower than 300 m (Fig. S4). A major portion of particulate mass between 1 and 2.5 µm originates from secondary aerosol formation processes in urban air (Y. H. Wang et al., 2015; Zhang et al., 2015). As shown in Fig. S7, the concentration of NR-PM_1 increased significantly from 12.1 to 56.4 $\mu g \, m^{-3}$ when the variation of MLH decreased from more than 1400 m to less than 200 m. The reduction in solar radiation reaching the surface due to fine particle matter reduces the turbulent kinetic energy and the development of the mixing layer, as shown in Fig. 5. An exponential function between the turbulent kinetic energy at 140 m and mixing layer height was fitted. Based on this fit, the MLH roughly doubles from about 400 to 800 m when TKE increases from 0.1 to 1 $m^2 \, s^{-2}$. These are typical values of MLH during polluted conditions in Beijing.

The reduced sensible heat and TKE due to aerosol particles reduce the entrainment of relatively dry air into the mixing layer from above, which makes the air more humid within the mixing layer. This, together with the decreased surface temperature, increases the relative humidity (Z. Li et al., 2017). The increased relative humidity enhances the aerosol water uptake and promotes the formation of secondary organic and inorganic aerosol via aqueous-phase reactions (Liu et al., 2018; Wang et al., 2019), enhancing light scattering and causing further reduction in the intensity of radiation reaching the surface. All these factors suppress the
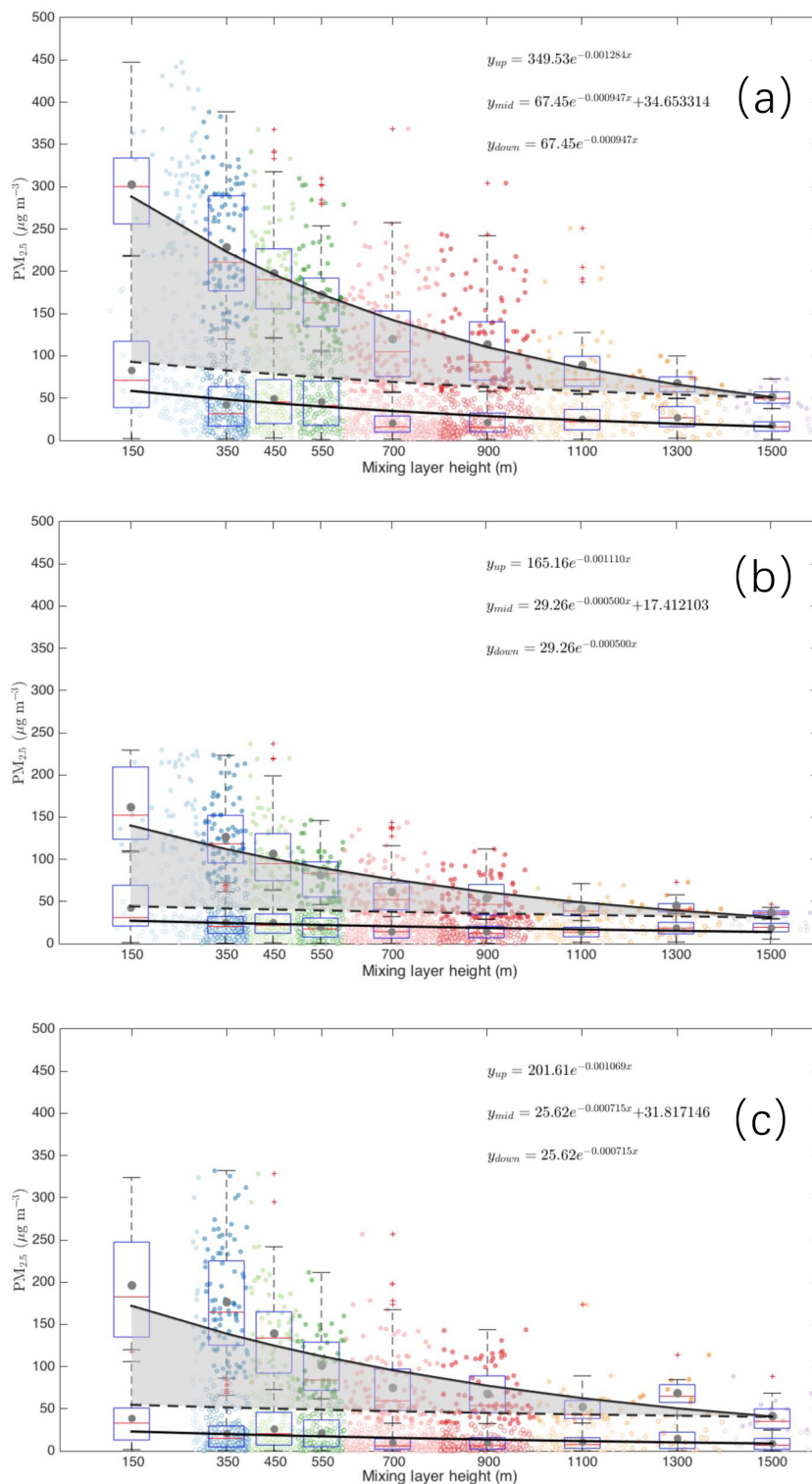
Figure 4. The variability of the $PM_{2.5}$ mass concentration as a function of the mixing layer height at 8 m **(a)**, 120 m **(b)** and 280 m **(c)**. The data related to the upper fitting line represent $PM_{2.5}$ concentrations larger than 75 $\mu g \, m^{-3}$, while the data related to the lower fitting line represent $PM_{2.5}$ concentrations less than 75 $\mu g \, m^{-3}$. The dark grey points represent mean values; the red line represents median values. The shadowed area corresponds to an increased amount of $PM_{2.5}$ with decreased mixing layer height assuming that $PM_{2.5}$ has the same variation pattern under highly polluted conditions as in less polluted times.
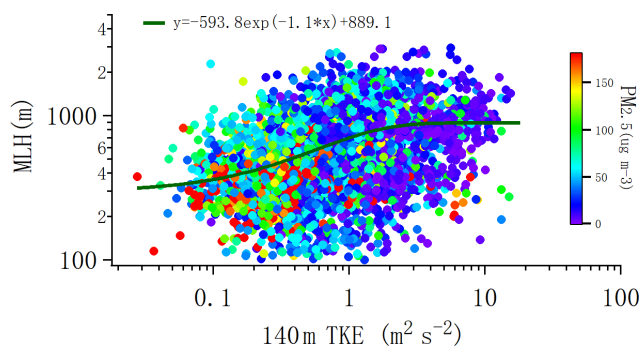
Figure 5. Turbulent kinetic energy at 140 m as a function of mixing layer height and $PM_{2.5}$ concentrations at 120 m from July 2009 to August 2011. An exponential function was fitted based on best fitting.

development of mixing layer height and enhance the accumulation of air pollutants within the mixing layer. We ascribe part of the observed increase in $PM_{2.5}$ and simultaneous decrease in the mixing layer height to the positive feedback associated with the particulate matter–mixing layer interaction (Petäjä et al., 2016; Ding et al., 2016), occurring at the same time as primary emissions and secondary formation are confined into a smaller volume of air. The feedback occurred at all three observation platforms and appeared to be most intensive at 8 m. In an urban environment, NO_x originates mainly from local anthropogenic emissions, whereas the sources of particulate matter include both primary emissions and secondary formation (Ehn et al., 2014; Jimenez et al., 2009; Zhang et al., 2015; Zhao et al., 2013). As shown in Fig. S6, the median NO_x concentration at 8 m was 250 % higher under highly polluted conditions compared with less polluted conditions as the mixing layer height decreased to 100–200 m, while the corresponding number for the $PM_{2.5}$ concentration was 360 %.

The increase in the $PM_{2.5}$ concentration from less polluted to highly polluted conditions is mainly due to concentrated particulate matter caused by a decreased mixing layer height, which is accompanied by primary particle emissions, secondary aerosol formation and feedback from particulate matter–mixing layer height interactions. Compared with the increased amounts of NO_x, we can roughly estimate that a maximum of 110 % of the increased $PM_{2.5}$ originates from secondary aerosol formation processes in this study. Of the remaining 250 % of the $PM_{2.5}$ increase, potentially a large fraction originates from particulate matter–mixing layer height interactions, but we cannot quantify this fraction at the moment.

4 Conclusions

The development of a mixing layer height in an urban city is affected by the intensity of incoming solar radiation. Our measurement at the 325 m meteorology tower showed that the solar and ultraviolet radiation reaching the surface decreases considerably at increased pollution levels, which leads to a decreased TKE and, consequently, the suppression of mixing layer development. In turn, the shallowed mixing layer height further favours the enhancement of $PM_{2.5}$ concentration and its precursor gases from both direct emissions and secondary formation. This feedback mechanism may be an important reason for the rapid increase in particulate matter from moderately polluted conditions to periods of intense pollution in an urban atmosphere as the strength increased when the $PM_{2.5}$ concentration increased, although we cannot quantify the feedback amount exactly by observations currently. The particulate matter–mixing layer height feedback is probably a critical factor in the formation of intense haze periods from moderately polluted periods in Beijing and other polluted cities.

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