# CONSTRUCTION AND RESEARCH OF FULL BALANCE ENERGY OF VARIATIONAL PROBLEM MOTION SURFACE AND GROUNDWATER FLOWS

Petro Venherskyi

Department of Applied Mathematics and Informatics Franko National University of Lviv I Universytetska str., Lviv, Ukraine, 79001 petro.vengersky@gmail.com

#### Abstract

Based on the laws of conservation of mass and momentum the basic equations of motion with unknown quantities velocity and piezometric pressure are written. These equations are supplemented with boundary and initial conditions describing the motion of compatible flows. Based on the laws of motion continuum, received conditions contact on the common border interaction of surface and groundwater flows. Variational problems formulated compatible flow. Energy norms of basic components of variational problem are analyzed. Correctness of constructing variational problem arising from construction of the energy system of equations that allow to investigate properties of the problem solution, its uniqueness, stability, dependence on initial data and more. Energy equation of motion of surface and groundwater flows are derived and investigated. It is shown that the total energy compatible flow depends on sources that are located inside the domain or on its border.

**Keywords:** surface flow, groundwater flow, watershed, incompressible fluid, velocity fluid and hydrostatic and piezometric pressure, energy equation, bilinear form, Initial and boundary conditions, interface conditions, coupling flow.

## DOI: 10.21303/2461-4262.2017.00270

© Petro Venherskyi

#### **1. Introduction**

An important role in studying the water cycle plays hydrological system. In general, research integrity of the system, taking into account all impacts, are complex and not always feasible problem for the study because only investigated some of the area involved in the water cycle [1–3] Highly likely part of the territory may be a watershed area (**Fig. 1**), which is characterized by similar climatic conditions and is influenced by such factors that affect the water movement.



Fig. 1. Two-dimensional projection watershed on the plane  $X_1OX_2$ 

At the watershed may be an interaction between flow and located above and below water-bearing layers. Models of different dimensions are used in each layer to describe the water movement and their solutions are connected by boundary conditions [4-6].

We select in solid medium (liquid) moving surface layer  $F(t) \in \mathbb{R}^3$  (Fig. 2) of such a structure

$$\Omega_{F}(t) := \left\{ \left( x_{1}, x_{2}, x_{3} \right) \in \mathbb{R}^{3}, \quad \eta(x) < x_{3} < \nu(x, t) \quad \forall x = \left( x_{1}, x_{2} \right) \in \Omega(t) \right\}.$$

$$(1)$$

Let's denote projection of its lower

$$\Omega(t) := \left\{ \left( x_1, x_2, x_3 \right) \in \mathbb{R}^3 \mid x_3 = \eta(x), \forall x = \left( x_1, x_2 \right) \in \Omega(t) \right\}$$
(2)

and upper

$$\Lambda_{F}(t) := \left\{ \left( x_{1}, x_{2}, x_{3} \right) \in \mathbb{R}^{3} \mid x_{3} = v(x, t) \quad \forall x = \left( x_{1}, x_{2} \right) \in \Omega(t) \right\}$$
(3)

bases on the plane  $0x_1x_2$ . The rest of the surface layer

$$\Gamma_{F}(t) \coloneqq \left\{ \begin{pmatrix} x_{1}, x_{2}, x_{3} \end{pmatrix} \in \mathbb{R}^{3}, \quad \eta(x) < x_{3} < \nu(x, t) \quad \forall x = (x_{1}, x_{2}) \in \Omega(t) \right\}$$
(4)

will be called the lateral surface layer F(t).

Similarly denote part of fluid that moves in the soil, so

$$\Omega_{P}(t) := \{ (x_{1}, x_{2}, x_{3}) \in \mathbb{R}^{3}, h(x) < x_{3} < \eta(x), \forall x = (x_{1}, x_{2}) \in \Omega(t) \}$$
(5)

the projection of the lower part will be written as

$$\Lambda_{P}(t) := \left\{ \left( x_{1}, x_{2}, x_{3} \right) \in \mathbb{R}^{3} \mid x_{3} = h(x), \forall x = \left( x_{1}, x_{2} \right) \in \Omega(t) \right\}.$$

$$(6)$$

Then, a layer of ground water

$$\Gamma_{P}(t) := \left\{ \begin{pmatrix} x_{1}, x_{2}, x_{3} \end{pmatrix} \in \mathbb{R}^{3}, \quad h(x) < x_{3} < \eta(x) \quad \forall x \in \Gamma_{P}(t) \right\}.$$

$$(7)$$



Fig. 2. General view of the model of flows and their cross-section

## 2. Materials and Methods

## 2. 1. Initial boundary value problem of interaction of water flows

We formulate initial boundary problem of motion of surface and groundwater flows on the surface watershed considering boundary and initial conditions [7–9].

Find unknown quantities  $\{u, p, \varphi\}$  such that satisfy the following system of equations:

$$\begin{split} \frac{\partial}{\partial t} \left( \rho u_{i} \right) + \sum_{k=1}^{3} \frac{\partial}{\partial x_{k}} \left( \rho \ u_{i} u_{k} \right) - \rho \ f_{i} \ - \sum_{k=1}^{3} \frac{\partial \sigma_{ik}}{\partial x_{k}} = 0, \end{split} \tag{8}$$
$$\sigma_{ij} = -p_{F} \delta_{ij} + \tau_{ij}, \\\tau_{ij} = 2\mu \ e_{ij}, i, j = 1, 2, 3, \\e_{ij} = \frac{1}{2} \left( \frac{\partial u_{i}}{\partial x_{j}} + \frac{\partial u_{j}}{\partial x_{i}} \right), \end{split}$$

$$\frac{\partial \rho}{\partial t} + \sum_{k=1}^{3} \frac{\partial \left(\rho \ u_{k}\right)}{\partial x_{k}} = 0, \text{ in } \Omega_{F} \times (0, T], \tag{9}$$

$$m\frac{\partial\phi}{\partial t} = \sum_{j=1}^{3} \frac{\partial}{\partial x_{j}} \left( k\frac{\partial\phi}{\partial x_{j}} \right) + \varepsilon \quad \text{in} \quad \Omega_{p} \times (0;T],$$
(10)

where  $\{u_i(x,t)\}_{i=1}^3$  and  $p_{F_i} = p_F(x,t)$  - sought velocity vector of fluid and hydrostatic pressure, respectively;  $F = \left\{g \ f_i(x)\right\}_{i=1}^3$  - mass forces;  $\rho = \rho(x,t) > 0$  - density of the mass water flow;  $\mu = \mu(x) > 0$  - viscosity coefficient;  $\left\{e_{ij}\right\}_{i,j=1}^3$ ,  $\left\{\sigma_{ij}\right\}_{i,j=1}^3$  - tensors of velocities of deformation and stress of the liquid at the point x in time t;  $\delta_{ij}$  - Kronecker symbol; k = k(x,t) - filtration coefficient; m = m(x,t) - coefficient of specific water loss;  $\epsilon = \epsilon(x,t)$  - known function of sources of water influx;

$$\varphi = x_3 + \frac{p_p}{\rho g} \tag{11}$$

piezometric pressure;

$$q = -k\nabla\phi \tag{12}$$

flow (flow separation);  $\upsilon = \upsilon(x,t)$  – velocity vector of fluid in the ground;  $\upsilon = \frac{q}{\omega}$ ,  $\omega$  – volume porosity;  $\overline{n_F} = -\overline{n_P}$  – vectors normal to the boundary area  $\Omega_F$  and  $\Omega_P$  in accordance;

$$\overline{\Omega} = \overline{\Omega_{F}} \cup \overline{\Omega_{P}}, \Omega_{F} \cap \Omega_{P} = \{\emptyset\}, \overline{\Omega_{F}} \cap \overline{\Omega_{P}} = \Gamma,$$
$$\partial \Omega_{F} = \Gamma_{F} \cup \Lambda_{F} \cup \Gamma; \partial \Omega_{P} = \Gamma_{P} \cup \Lambda_{P} \cup \Gamma.$$

Boundary conditions [10, 11]:

$$\vec{u}_i = 0 \text{ on } \Gamma_F, i=1, 2, 3,$$
 (13)

$$\sigma_{n\tau} = \overline{\sigma}, \text{ on } \Lambda_{F},$$
 (14)

$$u_{3} + R = \frac{\partial v}{\partial t} + u_{1}^{0} \frac{\partial v}{\partial x_{1}} + u_{2}^{0} \frac{\partial v}{\partial x_{2}} \text{ in } \Omega_{F} \times (0, T],$$
(15)

where R – velocity of falling rain drops,  $u_1^0$ ,  $u_2^0$  – horizontal components of velocity on the free surface  $v(x.t) (\Lambda_F)$ ;

$$\upsilon_{\bullet} n_{\rm p} = \overline{\upsilon} \quad \text{on} \quad \Gamma_{\rm P};$$
(16)

$$\upsilon_1 = \upsilon_2 = 0 \quad \text{on} \quad \Lambda_P, \tag{17}$$

$$\upsilon_3 = -I \quad \text{on} \quad \Lambda_P, \tag{18}$$

where I – known function that describes the velocity of fluid flow through the surface  $\Lambda_{p}$ . Initial conditions:

$$\begin{split} u \big|_{t=0} &= u_0, \\ p \mid_{t=0} &= p_0, \text{ in } \Omega. \end{split} \tag{19} \\ \phi \mid_{t=0} &= \phi_0, \end{split}$$

Contact flow conditions on a common boundary  $\Gamma$  [4–6, 8]:

$$\sigma_{nn}(u, p_F) = p_p,$$

$$\sigma_{nn} = 0,$$

$$u_n = -v_n.$$
(20)

# 2. 2. Variational formulation of the problem of interaction of water flows

We introduce the following bilinear forms:

$$\begin{split} M_{v}(r;w,q) &= \int_{v} \sum_{i=1}^{3} rw_{i}q_{i}ds, \ N_{v}(w;u,q) = \int_{v} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho w_{k} \frac{\partial u_{i}}{\partial x_{k}} q_{i}ds, \\ C_{v}(w,q) &= \int_{v} 2\mu \ e(w) : e(q)ds, \\ A_{v}(w,q) &= -\int_{v} wdivqds, \ Y_{v}(w,q) = -\int_{v} wq_{n}d\gamma, \ B_{v}(p,w) = -\int_{v} \sum_{i=1}^{3} p.\nabla wds. \end{split}$$

Introduce spaces:

$$\begin{split} H_{F} &:= \left\{ \xi \in (H^{1}(\Omega_{F}))^{3} \mid \xi = 0 \quad \text{on} \quad \Gamma \right\}, \\ H_{p} &:= \left\{ \psi \in H^{1}(\Omega_{p}) \mid \psi = 0 \quad \text{on} \quad \Gamma \right\}, \quad W := H_{F} \times H_{p}, \ \mathfrak{S}_{j} : W \to R, j = \overline{1,3}, \\ &\left\langle \mathfrak{S}_{1}, \xi \right\rangle = \sum_{i=1}^{3} \int_{\Omega_{F}} \rho f_{i} \xi_{i} ds + \int_{\Lambda_{F}} (\xi_{n} p_{a} + \xi_{\tau} \cdot \widehat{\sigma}) d\gamma, \\ &\left\langle \mathfrak{S}_{2}, \theta \right\rangle = - \int_{\partial \Lambda_{F}} u_{n}^{0} \theta d\gamma, \ \left\langle \mathfrak{S}_{3}, \psi \right\rangle = \int_{\Omega_{p}} \frac{\epsilon(x, t) \rho g \psi}{\omega} dp - \int_{\partial \Lambda_{p}} \overline{\upsilon} \psi \rho g d\gamma. \end{split}$$

Let's denote

$$\tilde{\psi} = \psi \rho g, \quad \tilde{m} = \frac{m}{\omega},$$

Then, let's write the following variational problem [1-2, 10, 11]:

$$\begin{aligned} & \operatorname{Find} \left\{ u, p, \varphi \right\} \in V \times Q \times W, \\ & \operatorname{M}_{\Omega_{F}} \left( \rho; u', \xi \right) + \operatorname{N}_{\Omega_{F}} \left( u; u, \xi \right) + \operatorname{A}_{\Omega_{F}} \left( p, \xi \right) + \operatorname{C}_{\Omega_{F}} \left( u, \xi \right) + \\ & + \operatorname{Y}_{\Gamma} \left( u, \xi \right) = \left\langle \mathfrak{I}_{1}, \xi \right\rangle, \forall \xi \in V, \end{aligned} \tag{21}$$

$$B_{\Omega_{F}}(u,\theta) + Y_{\Gamma}(\theta,u) = \langle \mathfrak{I}_{2}, \theta \rangle, \forall \theta \in Q,$$
(22)

$$M_{\Omega_{p}}(\tilde{m}; \phi', \tilde{\psi}) + A_{\Omega_{p}}(\tilde{\psi}, \upsilon) + Y_{\Gamma}(\tilde{\psi}, \upsilon) = \langle \mathfrak{I}_{3}, \tilde{\psi} \rangle, \forall \psi \in W$$
(23)

with initial conditions

$$M_{\Omega_{\rm F}}(\mathbf{u}'(0) - \mathbf{u}_0, \xi) = 0, \tag{24}$$

$$B_{\Omega_{\rm F}}(p(0) - p_0, \theta) = 0; \tag{25}$$

$$M_{\Omega_{p}}(\phi'(0) - \phi_{0}, \tilde{\psi}) = 0.$$
(26)

Let's calculate, considering initial conditions (24)–(26) and boundary conditions (13)–(18), values of variables u and p with relations (21) and (22). Then on the basis of coupling flow conditions (interface conditions) (20) and boundary condition (11) the value of the variable  $\varphi$  is calculated from (23).

# 2. 3. The properties of the components and norms of variational problem interaction water flows.

It should be noted that trilinear form

$$N_{v}(w; u, q) = \int_{v} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho w_{k} \frac{\partial u_{i}}{\partial x_{k}} q_{i} ds, \qquad (27)$$

is continuous and bilinear form

$$C_v(w,q) = \int_v 2\mu e(w) : e(q) ds$$
 (28)

continuous and symmetrical.

It is a scalar product in the space  $H_F$  and creates a norm

$$\left\|\mathbf{w}\right\|_{\mathbf{H}_{\mathrm{F}}} = \sqrt{\mathbf{C}_{\mathbf{v}}(\mathbf{w}, \mathbf{w})}, \forall \mathbf{w} \in \mathbf{H}_{\mathrm{F}}.$$

Then, let's write the scalar function  $\varphi$  bilinear forms

$$D_{v}(\phi, \psi) = \int_{v} k(x, t) \nabla \phi \cdot \nabla \psi \, dp.$$
<sup>(29)</sup>

which is continuous and integral in the space of admissible functions  $H_{p}$ . It is also symmetrical and forms a semi-norm

$$|\varphi|_{H_p} = \sqrt{D_v(\varphi, \varphi)}, \forall \varphi \in H^1(\Omega_p).$$
 (30)

Let's consider the properties of bilinear forms

$$A_{v}(w,q) = \int_{v} divwdivqds.$$
(31)

In space H<sub>F</sub>, it is continuous, integral and symmetrical, and also forms the norm

$$\left\| \left| q \right| \right\|_{\mathbf{H}_{\mathbf{F}}} = \sqrt{\mathbf{A}_{\mathbf{v}}(q,q)}, \forall q \in \mathbf{H}_{\mathbf{F}}.$$
(32)

## 3. Results of research

# 3. 1. Equation balance energy of coupling water flow

Let's write variational equations for momentum

$$M_{\Omega_{F}}(\rho; u', u) + N_{\Omega_{F}}(u; u, u) + C_{\Omega_{F}}(\mu; u, u) =$$
$$= \langle \phi_{I}, u \rangle - Y_{\Omega_{F}}(p, u) - A_{\Omega_{F}}(p, u).$$
(33)

Let's write left side of the equation:

$$\int_{\Omega_{F}} \sum_{i=1}^{3} \rho_{i} u_{i}' u_{i} ds + \int_{\Omega_{F}} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho u_{k} \frac{\partial u_{i}}{\partial x_{k}} u_{i} ds - \int_{\Omega_{F}} \sum_{k=1}^{3} \frac{\partial \sigma_{ik}}{\partial x_{k}} u_{k} ds = \int_{\Omega_{F}} \sum_{i=1}^{3} \rho_{i} u_{i}' u_{i} ds + \int_{\Omega_{F}} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho u_{k} \frac{\partial u_{i}}{\partial x_{k}} u_{i} ds + \int_{\Omega_{F}} \sum_{k=1}^{3} \sigma_{ik} \frac{\partial u_{i}}{\partial x_{k}} ds - \int_{\partial\Omega_{F}} u_{i} \sum_{k=1}^{3} \sigma_{ik} n_{F_{k}} d\gamma.$$

$$(34)$$

Given that

 $\sigma_{ik} = -p\delta_{ik} + 2\mu e_{ik}$ 

rewrite (34) in the following form

$$\int_{\Omega_{F}} \sum_{i=1}^{3} \rho_{i} u_{i}' u_{i} ds + \int_{\Omega_{F}} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho u_{k} \frac{\partial u_{i}}{\partial x_{k}} u_{i} ds - \int_{\partial\Omega_{F}} p u_{n} d\gamma + \int_{\Omega_{F}} u \nabla p ds + \int_{\Omega_{F}} 2\mu e(u) : e(u) ds - \int_{\partial\Omega_{F}} u_{i} \sum_{k=1}^{3} \sigma_{ik} n_{F_{k}} d\gamma.$$
(35)

Let's write left side of variational equations for the law of conservation of mass flow

$$\int_{\Gamma} u.n_{\rm F} p d\gamma - \int_{\Omega_{\rm F}} u.\nabla p ds = 0.$$
(36)

Substituting (23) in place of  $\psi$  the function  $\phi$  will be

$$\int_{\Omega_{p}} m \frac{\partial \phi}{\partial t} \phi \frac{\rho g}{\omega} dp - \int_{\partial \Omega_{p}} k \frac{\partial \phi}{\partial n_{p}} \phi d\gamma +$$
  
+ 
$$\int_{\Omega_{p}} \sum_{j=1}^{3} m \frac{\partial \phi}{\partial x_{j}} \frac{\partial \phi}{\partial x_{j}} dp - \int_{\Omega_{p}} \epsilon \phi dp = 0.$$
(37)

Let's multiply (37) on the expression  $\frac{\rho g}{\omega}$ , then

$$\frac{1}{2} \int_{\Omega_{p}} m \frac{\partial \phi^{2}}{\partial t} \frac{\rho g}{\omega} dp - \int_{\partial\Omega_{p}} \sum_{j=1}^{3} \frac{k(x,t)}{\omega} \frac{\partial \phi}{\partial \overline{n}} \phi \rho g d\gamma +$$
$$+ \int_{\Omega_{p}} \sum_{j=1}^{3} \frac{k(x,t)}{\omega} \frac{\partial \phi}{\partial x_{j}} \frac{\partial \phi}{\partial x_{j}} \rho g dp - \int_{\Omega_{p}} \frac{\rho g}{\omega} \varepsilon \phi dp = 0.$$
(38)

Let's estimate the term in (38) on the boundary  $\Omega_{\rm p}$ 

$$-\int_{\partial\Omega_{p}} \frac{k}{\omega} \frac{\partial \varphi}{\partial n_{p}} \varphi \rho g d\gamma = -\int_{\Gamma_{p}} \frac{k}{\omega} \frac{\partial \varphi}{\partial n_{p}} \varphi \rho g d\gamma - \int_{\Gamma} \frac{k}{\omega} \frac{\partial \varphi}{\partial n_{p}} \varphi \rho g d\gamma - \int_{\Lambda_{p}} \frac{k}{\omega} \frac{\partial \varphi}{\partial n_{p}} \varphi \rho g d\gamma =$$
$$= -\int_{\Gamma} k \varphi \rho g \frac{\nabla \varphi n_{p}}{\omega} d\gamma - \int_{\Gamma_{p}} k \frac{\nabla \varphi n_{p}}{\omega} \varphi \rho g d\gamma =$$
$$= \int_{\Gamma} p_{p} \upsilon_{n_{p}} d\gamma - \int_{\Gamma_{p}} \overline{\upsilon} \varphi \rho g d\gamma.$$
(39)

Simplifying a term on the border  $\,\Omega_{\!_{F}}\,$  in the form (35), we obtain

•

$$-\int_{\partial\Omega_{F}} u_{i} \sum_{k=1}^{3} \sigma_{ik} n_{F_{k}} d\gamma = -\int_{\Lambda_{F}} (u_{n} \sigma_{nn} + u_{\tau} \sigma_{n\tau}) d\gamma - \int_{\Gamma_{F}} (u_{n} \sigma_{nn} + u_{\tau} \sigma_{n\tau}) d\gamma - \int_{\Gamma} (u_{n} \sigma_{nn} + u_{\tau} \sigma_{n\tau}) d\gamma.$$
(40)

Adding expressions (35), (39), (40), after simple transformations, we have

Computer Sciences and Mathematics

$$\int_{\Omega_{F}} \sum_{i=1}^{3} \rho_{i} u_{i}' u_{i} dx + \int_{\Omega_{F}} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho u_{k} \frac{\partial u_{i}}{\partial x_{k}} u_{i} dx -$$

$$-\int_{\Lambda_{F}} p u_{n} d\gamma - \int_{\Gamma_{F}} p u_{n} d\gamma - \int_{\Gamma} p u_{n} d\gamma + \int_{\Omega_{F}} u \nabla p ds + \int_{\Omega_{F}} 2\mu e(u) : e(u) dx -$$

$$-\int_{\Lambda_{F}} (u_{n} p_{a} + u_{\tau} \sigma_{n\tau}) d\gamma - \int_{\Gamma} (u_{n} \sigma_{nn} + u_{\tau} \sigma_{n\tau}) d\gamma + \int_{\Gamma} u.n_{F} p d\gamma - \int_{\Omega_{F}} u.\nabla p ds +$$

$$+ \frac{1}{2} \int_{\Omega_{P}} \frac{m\rho g}{\omega} \frac{\partial \phi^{2}}{\partial t} dx + \int_{\Omega_{P}} \sum_{j=1}^{3} \frac{k(x, t)}{\omega} \left(\frac{\partial \phi}{\partial x_{j}}\right)^{2} \rho g dp - \int_{\Omega_{P}} \frac{\rho g}{\omega} \varepsilon \phi dx -$$

$$-\int_{\Gamma} p_{p} \upsilon_{n_{p}} d\gamma + \int_{\Gamma_{P}} \overline{\upsilon} \phi \rho g d\gamma = 0.$$
(41)

Rewriting the previous expression (41) in a more convenient form, including property incompressible environment and boundary conditions (13)–(18), we obtain

$$\int_{\Omega_{F}} \sum_{i=1}^{3} \rho_{i} u_{i}' u_{i} dx + \int_{\Omega_{F}} \sum_{k=1}^{3} \sum_{i=1}^{3} \rho u_{k} \frac{\partial u_{i}}{\partial x_{k}} u_{i} dx + \int_{\Omega_{F}} 2\mu e(u) : e(u) dx - -\int_{\Lambda_{F}} p u_{n}^{0} d\gamma - \int_{\Lambda_{F}} (u_{n} p_{a} + u_{\tau} \overline{\sigma}) d\gamma - -\int_{\Gamma} (u_{n} \sigma_{nn} + u_{\tau} \sigma_{n\tau}) d\gamma + \frac{1}{2} \int_{\Omega_{P}} \frac{m\rho g}{\omega} \frac{\partial \phi^{2}}{\partial t} dx + \int_{\Omega_{P}} \sum_{j=1}^{3} \frac{k(x, t)}{\omega} \left( \frac{\partial \phi}{\partial x_{j}} \right)^{2} \rho g dp - -\int_{\Omega_{P}} \frac{\rho g}{\omega} \varepsilon \phi dx - \int_{\Gamma} p_{p} v_{n_{p}} d\gamma + \int_{\Gamma_{P}} \overline{v} \phi \rho g d\gamma = 0.$$

$$(42)$$

Let's analyze the terms on joint border  $\boldsymbol{\Gamma}$ 

$$\int_{\Gamma} (u_n p_F + u_\tau \sigma_{n\tau}(u) - p_p \upsilon_{n_p}) d\gamma.$$

Given the terms the coupling (20), integral to the common border  $\Gamma$  is zero. From the expression (36) given kinematic condition (15) for equation of continuity will be

$$\int_{\partial\Omega_{\rm F}} u.n_{\rm F} p d\gamma = \int_{\Lambda_{\rm F}} u_{\rm n}^0 p d\gamma.$$
(43)

Thus, the energy balance equation of compatible motion of surface and groundwater flow is written:

Rewriting (44) through a total derivative, we have

$$\frac{1}{2}\rho\int_{\Omega_{F}}\frac{d}{dt}(u^{2}) ds + \int_{\Omega_{F}}2\mu e(u):e(u)dx -$$

$$+\frac{1}{2}\int_{\Omega_{P}}\frac{m\rho g}{\omega}\frac{\partial\phi^{2}}{\partial t}dx + \int_{\Omega_{P}}\sum_{j=1}^{3}\frac{k(x,t)}{\omega}\left(\frac{\partial\phi}{\partial x_{j}}\right)^{2}\rho g dp =$$

$$\sum_{i=1}^{3}\int_{\Omega_{F}}\rho f_{i}u_{i}ds + \int_{\Lambda_{F}}p u^{0}{}_{n}d\gamma -$$

$$-\int_{\Gamma_{e}}\overline{\upsilon} \phi\rho g d\gamma + \int_{\Lambda_{F}}(u_{n}p_{a} + u_{\tau}\overline{\sigma})d\gamma + \int_{\Omega_{P}}\frac{\rho g}{\omega}\varepsilon\phi dx.$$
(45)

As we see from (45), the total energy flow depends on the energy sources that are located within the region or within its boundaries.

#### 4. Conclusions

On the basis of conservation laws basic equations and boundary and initial conditions are derived describing the compatible motion flow of surface and ground water with unknown values of velocity and piezometric pressure. Variational problems of compatible flow are formulated and the contact conditions on the common border are obtained based on the laws of motion continuum. Energy standards of basic components of variational problem are analyzed. Full energy equation of energy balance for coupling motion of surface and groundwater flows are constructed and studied that makes it possible to investigate the properties of solutions of the problem, such as stability, regularity, existence, convergence and so on.

#### References

[1] Shlychkov, V. A. (2007). Numerical simulation of currents and admixture transport in a multi – arm river channel. Bull. Nov. Comp. Center, Num. Model in Atmosph., etc., 11, 79–85.

[2] Kuchment, L. S., Gelfan, A. N. (2002). Estimation of Extreme Flood Characteristics Using Physically Based Models of Runoff Generation and Stochastic Meteorological Inputs. Water International, 27 (1), 77–86. doi: 10.1080/02508060208686980

[3] Panday, S., Huyakorn, P. S. (2004). A fully coupled physically-based spatially-distributed model for evaluating surface/subsurface flow. Advances in Water Resources, 27 (4), 361–382. doi:10.1016/j.advwatres.2004.02.016

[4] Lions, J. L., Temam, R., Wang, S. (1993). Models for the coupled atmosphere and ocean. (CAO I, II). Computational Mechanics, 1 (1), 120.

[5] Discacciati, M., Quarteroni, A., Valli, A. (2007). Robin-Robin Domain Decomposition Methods for the Stokes–Darcy Coupling. SIAM Journal on Numerical Analysis, 45 (3), 1246–1268. doi: 10.1137/06065091x

[6] Cesmelioglu, A., Chidyagwai, P., Riviere, B. (2013). Continuous and discontinuous finite element methods for coupled surface-subsurface flow and transport problems. Rice University, 23.

[7] Venherskyi, P. S. (2014). Numerical investigation mathematical models coupled flow of surface and ground water from the catchment area. Mathematical and computer modeling, 10, 33–42.

[8] Venherskyi, P. S. (2014). About the problem of coupled motion of surface and ground water from the catchment area. Bulletin of Lviv University. Series Applied Mathematics and Computer Science, 22, 41–53.

[9] Venherskyi, P. S., Demkovych, O. R. (2002). Mathematical modeling of ground water in the saturated zone. 9-th National Conference Modern Problems of Applied Mathematics and Informatics, 1–36.

[10] Temam, R. (1995). Navier-Stokes equations and nonlinear functional analysis. SIAM, 148. doi: 10.1137/1.9781611970050

[11] Trushevskyi, V. M., Shynkarenko, H. A., Shcherbyna, N. M. (2014). Finite element method and artificial neural networks: theoretical aspects and application. Lviv: LNU Ivan Franko, 396.