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## Recent Sedimentation Rate and Sediment Ages Determination of Kemaman-Chukai Mangrove Forest, Terengganu, Malaysia

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**Abstract:** Two 20 cm sediment cores were collected with a D-section corer in the Kemaman-Chukai mangrove forest and analyzed for  $^{210}\text{Pb}_{\text{ex}}$  (excess  $^{210}\text{Pb}$ ) in order to understand the sedimentation pattern. The activities of  $^{210}\text{Pb}$  along with the  $^{208}\text{Po}$  tracer were measured by Alpha Spectrometry on a silicon surface barrier detector connected to a multi channel analyzer. Sedimentation rates were determined applying  $^{210}\text{Pb}$  dating method and found to vary between  $0.94\text{--}1.11\text{ cm year}^{-1}$ . The highest accumulation rates were recorded near the river mouth and the lowest rates at the downstream. Assuming that the sedimentation rate values are accurate, this implies that the sediments in the upper 100 cm at Kemaman-Chukai mangrove forest were deposited during the past 90 years ago.

**Key words:** mangrove forest, sedimentation rate, sediment age,  $^{210}\text{Pb}$

### INTRODUCTION

Mangrove forests are a buffer zone between the coast and the ocean. One of their presumed important functions is to provide a mechanism for trapping sediment. In terms of their biological and chemical aspects, mangrove forests are highly productive source of organic matter, from which there is a net outwelling of energy that supports the complex estuarine and nearshore food web. Geologists, on the other hand, view mangroves as sediment sinks, characterized by long-term import of sediments as indicated by the substantial accretion of recent sediments, which underlie mangrove forests and adjacent coastal plains<sup>[1]</sup>. Physically, by virtue of being in the intertidal areas, they can act as a recorder of environmental changes via sedimentological characteristics and in the preservation of spores and pollens.

Despite the acceptance that mangrove ecosystems are important sinks for sediments, few studies have addressed sediment accretion in this environment. Several authors<sup>[2,3]</sup> have studied some aspects of the sedimentology of mangroves and quote different sedimentation rates, which is probably a reflection of the non-representative sampling techniques employed. Spencely<sup>[4]</sup> and Shahbuddin *et al.*<sup>[5]</sup> have introduced a simple method for measuring accretion by simulating pneumatophores using rods and stakes and an artificial horizon marker method, respectively. Meanwhile, long-

term accretion rates using radionuclides have been well documented<sup>[6-8]</sup> but use of this approach has been not many published<sup>[5,9]</sup> and limited in the mangrove ecosystems. In the present study, we use  $^{210}\text{Pb}$  to estimate sediment accretion rates.

The concentration of  $^{210}\text{Pb}$  within sediment is governed by processes such as source input from overlying waters through sedimentation, in situ production and radioactive decay. Many scientists have successfully demonstrated  $^{210}\text{Pb}$  as a sensitive tracer for understanding the geochemical changes which have occurred during the century in different ecosystems, including coastal marine sediment, due to industrialization.  $^{210}\text{Pb}$  has proved to be a valuable tracer of sediment mixing and accumulation in a variety of environments<sup>[3,6,7,10,11]</sup>.  $^{210}\text{Pb}$  are naturally occurring radionuclides of the  $^{238}\text{U}$  decay series with a 22.3 year.  $^{210}\text{Pb}$  is used to examine sediment mprocesses on a 100 year time scale.  $^{210}\text{Pb}$  is supplied by its effective parent  $^{226}\text{Ra}$  in seawater and from  $^{222}\text{Rn}$  in the atmosphere. The atmospheric source is produced as  $^{222}\text{Rn}$ , a short-lived (3.8 days) intermediate daughter of  $^{226}\text{Ra}$ , escapes from the earth's crust, decay to  $^{210}\text{Pb}$  and is deposited back to the ground. In most shallow water environment such as mangrove ecosystem, atmospheric input is the major source. In Malaysia, studies relating to mangroves are not well documented and only little information is known concerning the sedimentation of the mangrove forest<sup>[9,12]</sup>. Furthermore, geochemical

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studies of sediments from the Malaysian mangrove forests have received little attention and only limited studies have been done regarding their role in the process of sedimentation<sup>[9]</sup>. In view of the importance of the mangrove to various aspects of the environment, research on the sediment accretion as well as the sediment age were carried out.

## MATERIALS AND METHOD

**Sampling sites:** The Kemaman-Chukai mangrove is located in the Kemaman district, inside one of the most extensive mangrove area (298.2 ha) on the east coast of Peninsular Malaysia (Fig. 1). The dominant species of mangrove tree found are *Rhizophora* sp., *Avicennia* sp. and *Xylocarpus* sp. The mangroves which are relatively undisturbed have been gazetted as a mangrove reserve forest by the Terengganu Forestry Department. The mangrove area lies on the southern bank of the Kemaman estuary, where both the Kemaman and Chukai rivers flow into. The tide floods the area twice daily as it is semi-diurnal with a mean range of 1.8 m. In this study, a 20 cm sediment core from both the Kemaman and Chukai mangrove were collected with a D-section core sampler from this area (Fig. 1) and was cut into segments of approximately 5 cm interval, labelled and stored in acid cleaned bottle for analysis.

**Analytical methods for <sup>210</sup>Pb:** The samples were cut into segment approximately 2 cm interval using plastic knife and was transferred to petri disc. The samples were dried in an oven at 60°C to constants weights. To facilitate geochemical analysis the sample were ground into a fine powder using a pestle and mortar. The process of <sup>210</sup>Pb excess dating method is divided into two process, digestion process and Po plating process<sup>[10,13]</sup>. The excess <sup>210</sup>Pb in a different core was measured by the same procedure describes by

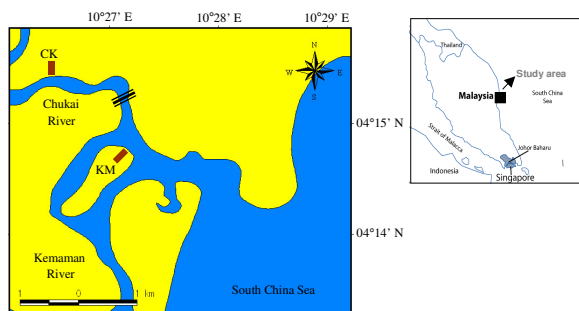


Fig. 1: Sampling location where 2 cores were collected at Kemaman (KM) and Chukai (CK) mangrove forest, Terengganu, Malaysia

Carpenter *et al.*<sup>[10]</sup> and Carpenter *et al.*<sup>[13]</sup>. Activities of <sup>210</sup>Po along with the <sup>208</sup>Po tracer were measured by Alpha Spectrometry on a silicon surface barrier detector connected to a multi channel analyzer.

**Data analysis:** Total <sup>210</sup>Pb activity was determined indirectly by the measurement of its alpha-emitting grand daughter nuclide, <sup>210</sup>Po<sup>[14]</sup>. Measurement of ratio <sup>210</sup>Po and <sup>209</sup>Po activity will provide an adequate figure of supported <sup>210</sup>Pb, as these two elements are assumed to be in equilibrium<sup>[15]</sup>. Subtracting of supported <sup>210</sup>Pb from total <sup>210</sup>Pb will determine unsupported <sup>210</sup>Pb. The activity of <sup>210</sup>Pb is obtained by the formula below:

$$\text{Activity } ^{210}\text{Po}(^{210}\text{Pb}) = A(\text{dpm/g}) \\ = \frac{\text{Actual } ^{210}\text{Po}(^{210}\text{Pb})}{\text{Actual } ^{209}\text{Po}} \times ^{209}\text{Po}(24.74 \text{ dpm/g}) \times \frac{\text{tracer weight } ^{209}\text{Pb}(\text{g})}{\text{sample weight}(\text{g})}$$

Where:

$A = A_0 e^{-\lambda t}$  (equal to the accumulative residual unsupported <sup>210</sup>Pb below the sediment age of t)

$A_0 = A/e^{-\lambda t}$  (equal to the total unsupported Pb in the sediment column)

$\lambda = \text{Ln}2/t_{1/2} = 0.639$  (decay constant of <sup>210</sup>Pb)

$T_{1/2} = \text{Half-life (22.3 years)}$

$t = \text{Depth (cm)/sedimentation rate in years}$

Inventories (I) of <sup>210</sup>Pb (unsupported) are expressed in dpm-2 and were calculated according to Turekian<sup>[16]</sup>:

$$I = \sum A_i \rho_i h_i$$

Where:

$A_i = \text{The } ^{210}\text{Pb}_{xs} (\text{dpm g}^{-1})$

$\rho_i = \text{The bulk density interval } i (\text{g cm}^{-3})$

$h = \text{The thickness of the interval (cm)}$

And finally the sedimentation rate can be calculated using the formula:

$$A = A_0 e^{-\lambda t}$$

$$A = A_0 e^{-\lambda(x/S)}$$

$$\ln A = -(\lambda/S)x + \ln A_0$$

Where:

$A = \text{Activity of excess } ^{210}\text{Pb} \text{ in the sediment at any depth}$

$A_0 = \text{Activity of excess } ^{210}\text{Pb} \text{ in the freshly deposited sediment at depth } = 0 \text{ (the sediment-water interface)}$

$S = \text{Sedimentation rate in cm year}^{-1}$

$\lambda = \text{Radioactive decay constant (0.0311/year)}$

$t = \text{Time in year}$

## RESULTS AND DISCUSSION

Figure 2 shows the activities of  $^{210}\text{Pb}_{\text{xs}}$  were generally decreasing exponentially with depth. Some higher value of  $^{210}\text{Pb}$  activity on the surface might be caused by the settling particles which were basically derived from the atmosphere and land or from the older sediment issued from the resuspension process<sup>[17]</sup>. It is clear that  $^{210}\text{Pb}$  increases monotonically with depth, suggesting of the particulate scavenging<sup>[18]</sup>. The determination of average sedimentation rate is based on the assumption that the  $^{210}\text{Pb}_{\text{xs}}$  is incorporated into the sediments with a constant rate<sup>[18]</sup>. As shown in Fig. 2, applying the formula as above, the sedimentation rate of Kemaman and Chukai mangrove forest were estimated to 1.11 and 0.94 cm year<sup>-1</sup>, respectively. The sedimentation rate was estimated by selecting the 'best-curve' from the establishment  $^{210}\text{Pb}$  (dpm g<sup>-1</sup>) distribution with depth. The  $p > 0.01$  calculated from the statistical test, ANOVA two factor prove that the sedimentation rate among the study area are not significantly different. In this study, Chukai mangrove which is situated about 2 km from river mouth had the lowest accumulation rate while the highest sediment accumulation rate occurs at Kemaman mangrove which is situated near the river mouth. Rivers are believed to

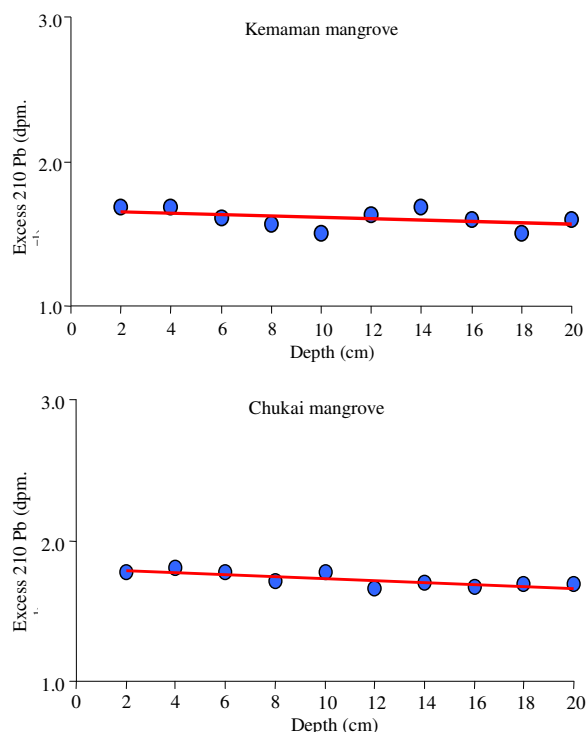


Fig. 2:  $^{210}\text{Pb}_{\text{excess}}$  distribution with depth in Terengganu mangrove forest

be the dominant source of sediment to the Kemaman-Chukai mangrove forest.

In Kemaman mangrove forest, the sedimentation rate obtained is consistent with the result obtained by Shabuddin *et al.*<sup>[5]</sup>, ranging from 0.64-1.46 cm year<sup>-1</sup> with artificial horizontal marker method but is slightly higher compared with Kamaruzzaman *et al.*<sup>[19]</sup> and Kamaruzzaman *et al.*<sup>[20]</sup>, 0.68 cm year<sup>-1</sup> with  $^{230}\text{Th}_{\text{ex}}$  and  $^{230}\text{Th}_{\text{ex}}/^{232}\text{Th}$  method. In comparison with other studies in Malaysia, sedimentation rate at Kemaman-Chukai mangrove is slightly higher than Pahang mangrove, 0.65 cm year<sup>-1</sup><sup>[21]</sup> and Johor mangrove, 0.84 cm year<sup>-1</sup><sup>[22]</sup> but lower compared to Matang mangrove, Perak<sup>[23]</sup>, 1.08 cm year<sup>-1</sup> using  $^{230}\text{Th}_{\text{excess}}$  method.

From the sedimentation rates,  $^{210}\text{Pb}$  dating technique was used to determine the sediments age for the major estuarine mangrove forest in Terengganu region. Sediment age can be determined by dividing the sediment depth with sedimentation rate. The lower the sedimentation rate indicates the older age of sediment. Assuming that the sedimentation rate values are accurate, this implies that the sediments in the upper 100 cm at Kemaman-Chukai mangrove forest were deposited during the past 90 years. From the ANOVA test, the sediment age does not differ significantly among the study areas with  $p < 0.05$ . Briefly, Chukai mangrove which had the lowest sedimentation rate had the oldest mangrove with 106.6 years while the youngest mangrove, 90.2 years, occurred at Kemaman mangrove which had the highest sedimentation rate.

## CONCLUSION

The sedimentation rates of Kemaman-Chukai mangrove forest, based on the  $^{210}\text{Pb}$  dating techniques were shown to be similar to the known local sea-level rise rate. The immobility of Pb in sediments was confirmed with stable  $^{210}\text{Pb}$  at the mangrove. It is suggested that the stable  $^{210}\text{Pb}$  is a powerful tracer in mangrove sediment studies on the determination of changing sedimentation rates, postdepositional mobilization and historical Pb sources.

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