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¹³C fractionation in transformations at the interface between roots, microorganisms, soil organic matter and soil respiration

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Introduction

Natural variations of the ¹³C/¹²C ratio have been frequently used over the last three decades to trace C sources and fluxes between plants, microorganisms, and soil. Most of these studies have used the natural ¹³C labelling approach, i.e. natural δ^{13} C variation after $C_3 - C_4$ vegetation changes. In this mini-review, we focus on ^{13}C fractionation in processes at the interface between roots, microorganisms, and soil: respiration, microbial respiration, root formation of dissolved organic carbon by root exudation, as well as microbial uptake and utilization of soil organic matter (SOM).

Root respiration

Based on literature data (Wedin et al., 1995; Klumpp et al., 2005; Schnyder and Lattanzi, 2005; Werth et al., 2006; Werth and Kuzyakov, 2006; Gessler et al., 2007;

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Larsen et al., 2007; Werth and Kuzyakov, 2009), we estimated that, on average, the roots of C₃ and C₄ plants are ¹³C enriched compared to shoots by $-1.2\pm0.6\%$ and - $0.3\pm0.4\%$, respectively. A very few studies showed that CO₂ released by root respiration was ¹³C depleted by about $+2.7\pm1.8\%$ for C₃ plants and $+1.3\pm2.4\%$ for C₄ plants compared to root tissue. This urgently calls for further studies to reliably estimate ¹³C fractionation by root respiration.

Microbial utilization

In soils developed under C₃ vegetation (Qian and Doran, 1996; Qian et al., 1997; Liang et Santruckova et al., al., 1999; 2000; Gregorich et al., 2000: Formánek and Ambus. 2004; Hamer et al.. 2004: Kristiansen et al., 2004; Pelz et al., 2005; Stevenson et al., 2005; Piao et al., 2006; Dijkstra et al., 2006; Werth et al., 2006; Crow et al., 2006; Boström et al., 2007; Werth and Kuzyakov, 2009), the microbial biomass was ¹³C enriched by -1.2±2.6‰ and microbial CO₂ was ¹³C enriched by -0.7±2.8‰ compared to SOM (Fig. 1). This discrimination pattern suggests preferential utilization of a ¹³C-enriched soil fraction by microorganisms, but a respiration of lighter compounds from this fraction. Preferential consumption of easily decomposable substrates and less negative δ^{13} C values were common for substances with low C/N ratios. For C₄ soils (Santruckova et al., 2000; Dijkstra et al., 2006), preferential substrate utilization was less important than for C₃ soils. This is because, here, microbial respiration strictly followed kinetics, i.e. microorganisms incorporated heavier carbon $(\Delta = -1.1\%)$ and respired lighter carbon ($\Delta =$ +1.1‰) than SOM (Fig. 1).



Fig. 1: ¹³C discrimination processes between soil organic matter (SOM) and the soil carbon pools: dissolved organic carbon (DOC), microbial biomass (MB), and SOM-derived CO₂ for C₃ and C₄ soils.



Temperature and precipitation had no significant effect on the ¹³C fractionation in these processes in C₃ soils. Increasing temperature and decreasing precipitation led, however, to increasing δ^{13} C values of soil C pools (Fig. 2).

CO₂ partitioning

In the framework of standard isotope mixing models, we calculated CO₂ partitioning using the natural ¹³C labelling approach at a vegetation change from C₃ to C₄ plants. A root-derived fraction of 25%, 50%, and 75% to total soil respiration was assumed. ^{13}C anv Disregarding fractionation processes (Eq. (13)), the calculated results deviated by between 2% and 7% from the presumed fractions (Fig. 3). Accounting for fractionation processes in the standard errors of the C₄ source pool and the mixing pool (Eq. (13) Fract. in SE) did not improve the exactness of the partitioning results; rather, it doubled the standard errors of the CO_2 pools. Including ¹³C fractionations directly into the mass balance equations (Eq. (20) Fract. in SE) reproduced the CO₂ partitioning exactly.



Fig. 2: Soil organic matter δ^{13} C vs. mean annual temperature (MAT) (a) and soil-derived CO₂ δ^{13} C vs. mean annual precipitation (MAP) (b) from sites with C₃ vegetation. The solid line shows the regression line, the dashed lines show the 95% confidence interval of the regression. For references see the beginning of this section.

Fig. 3: Partitioning of soil CO_2 efflux into root-derived (grey) and SOM-derived (white) carbon. Natural ¹³C labelling was used by planting a C_4 plant on soil developed under C_3 vegetation. Root-derived respiration from the C_4 plant (root respiration and rhizomicrobial respiration) was assumed to contribute 25%, 50%, and 75% to the CO_2 efflux. ¹³C fractionation between root-derived sources and CO_2

was either disregarded (Eq. (13)), accounted for in the standard error of the C_4 'end member' and the mixing pool (Eq. (13) Fract. in SE), or accounted for in the partitioning equation and in the standard error of the C_4 'end member' and the mixing pool (Eq. (20) Fract. in SE). Standard errors of the contributions are shown in the centre of the columns with their values below.

Conclusions

Most C transformations such as root respiration, formation of DOC, microbial utilization of soil organic matter as well as microbial respiration are correlated with significant changes of C isotopic signatures compared to that of the C source. The ¹³C fractionation within individual steps of C transformation is highly variable and is comparable with the δ^{13} C difference between the δ^{13} C values of 'end members' commonly used in natural ¹³C labelling studies. This makes it inappropriate to about possible accept literature data changes of δ^{13} C within the processes. the discrimination Rather. should be measured for the specific conditions of the experiment.

Simple calculations of partitioning errors in studies based on the natural ¹³C labelling approach showed high uncertainties of the results. This reflects small differences of δ^{13} C values between the 'end members', natural variation of δ^{13} C values within the the 'end members' and mixed pool. uncertainties of ¹³C fractionation, and preferential substrate utilization. This calls for caution in interpreting the results obtained using the natural ¹³C labelling approach.

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