

## Lower Cretaceous (Hauterivian-Albian) ammonite biostratigraphy in the Maestrat Basin (E Spain)

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### Abstract

A review of the stratigraphic distribution of ammonoid species in the Lower Cretaceous (Hauterivian-Albian) of the Maestrat Basin (E Spain) was carried out. The specimens were mainly collected in the field by us and are stored in university or museum collections. Specimens from private collections and figured in the literature were also studied. We recognized 73 species that are distributed, in accordance with the latest version of the standard Mediterranean ammonite zonation for the Lower Cretaceous, in 14 ammonite zones: *Acanthodiscus radiatus*, *Crioceratites loryi*, *Lyticoceras nodosoplicatum* (Lower Hauterivian); *Pseudothurmannia ohmi* (Upper Hauterivian); *Imerites giraudi* (Upper Barremian); *Deshayesites oglanensis*, *Deshayesites forbesi*, *Deshayesites deshayesi*, *Dufrenoyia furcata* (Lower Aptian); *Epicheloniceras martini*, *Parahoplites melchioris*, *Acanthohoplites nolani* (Upper Aptian); *Leymeriella tardefurcata* and *Douvilleiceras mammillatum* (Lower Albian). The recognition of these biozones allows a precise age calibration of the Maestrat Basin's lithostatigraphic units that contain ammonoids as well as an associated indirect age calibration of the formations without ammonoids. Consequently, this report provides an updated, comprehensive and precise biostratigraphic framework, which aims to become a reference for the analysis of the Lower Cretaceous strata of the Maestrat Basin. The results are also relevant for the analysis of coeval ammonite-bearing sedimentary successions found in other Tethyan basins.

**Keywords:** ammonites, Lower Cretaceous, biostratigraphy, Maestrat Basin, Spain

### Resumen

En este trabajo se ha realizado una revisión detallada de la distribución estratigráfica de las especies de ammonoideos del Cretácico inferior de la Cuenca del Maestrazgo (Este de España). Los ejemplares recolectados, principalmente por los autores, han sido depositados en colecciones universitarias y museísticas. Además hemos estudiado los ejemplares de colecciones privadas y figurados en la literatura. Hemos reconocido 73 especies que se distribuyen, siguiendo la última versión de la biozonación de ammonites mediterránea estándar del Cretácico inferior, en 14 zonas de ammonoideos: *Acanthodiscus radiatus*, *Crioceratites loryi*, *Lyticoceras nodosoplicatum* (Hauteriviano inferior); *Pseudothurmannia ohmi* (Hauteriviano superior); *Imerites giraudi* (Barremiano superior); *Deshayesites oglanensis*, *Deshayesites forbesi*, *Deshayesites deshayesi*, *Dufrenoyia furcata* (Aptiano inferior); *Epicheloniceras martini*, *Parahoplites melchioris*, *Acanthohoplites nolani* (Aptiano superior); *Leymeriella tardefurcata* y *Douvilleiceras mammillatum* (Albiano inferior). El reconocimiento de estas biozonas permite precisar la edad de las unidades litoestratigráficas que contienen ammonites y también una calibración indirecta de las formaciones que no contienen ammonites. En consecuencia este trabajo proporciona un marco bioestratigráfico actualizado, exhaustivo y preciso que pretende ser una referencia para el análisis estratigráfico del Cretácico inferior de la Cuenca del Maestrazgo. Los resultados obtenidos son también relevantes para el análisis de las sucesiones sedimentarias coetáneas con ammonites existentes en otras cuencas de Tetis.

**Palabras clave:** ammonites, Cretácico Inferior, biostratigraphy, Cuenca del Maestrazgo, España

## 1. Introduction

The presence of Lower Cretaceous ammonoids in the Maestrat Basin (E Spain) has been known since the second half of the nineteenth century (e.g., Vilanova, 1859; Coquand, 1865). There are many works, meeting abstracts, and doctoral theses that are focused on and/or contain data of this ammonoid record. However, the literature remains dispersed and difficult to access. Many studies are outdated because ammonoid biostratigraphic knowledge is continually improving, as indicated in the numerous reports from other Lower Cretaceous localities of the Tethys (e.g., Cecca *et al.*, 1999; Avram, 1999; Delanoy, 1995; Dutour, 2005; Ropolo *et al.*, 2000; Aguado *et al.*, 1997; Moreno-Bedmar *et al.* 2009a, 2010a; Najarro *et al.*, 2011; Moreno-Bedmar *et al.*, 2012a), as well as in the subsequent versions of the standard Mediterranean zonation published during the last two decades (Hoedemaeker and Bulot, 1990; Hoedemaeker *et al.*, 1993; Hoedemaeker *et al.*, 1995; Rawson *et al.*, 1999; Hoedemaeker *et al.*, 2000; Hoedemaeker *et al.*, 2003; Reboulet *et al.*, 2006; Reboulet *et al.*, 2009; Reboulet *et al.*, 2011). The stratigraphic knowledge and understanding of the Lower Cretaceous of the Maestrat Basin has also evolved and improved in recent years (e.g., Salas *et al.*, 2001; Bover-Arnal *et al.*, 2010).

In the present work, we update and review the biostratigraphic ammonoid knowledge of this basin following the latest version of the standard Mediterranean ammonite zonation (Reboulet *et al.*, 2011). The ammonite zonation in Mesozoic sedimentary successions is the most useful biostratigraphic tool as it supplies the most precise zonation. Therefore, this study also provides a reference for the analysis of coeval ammonite-bearing successions found in other basins of the Tethys.

## 2. Geological setting

The Maestrat Basin is located at the eastern Iberian Peninsula (Fig. 1). It corresponds to an intraplate rift basin, which developed during the Late Oxfordian-early Late Albian as a result of the spreading Atlantic Ocean and the opening of the Bay of Biscay (Salas & Casas, 1993; Salas *et al.*, 2001). The rift stage structured the Maestrat Basin into 7 sub-basins: Galve, Aliaga, Oliete, Morella, El Perelló, Salzedella and Penyagolosa (Fig. 1) (Salas and Guimerà, 1996). During this time, a sedimentary succession of up to > 2500 m thick made up of continental to hemipelagic deposits accumulated (e.g., Canérot *et al.*, 1982). The depocenter of the basin was located at the Salzedella Sub-basin (Fig. 1). During the Paleogene the Alpine orogeny caused the inversion of the Iberian Mesozoic rifts, and the Maestrat Basin gave rise to the eastern part of the Iberian Chain (Salas and Casas, 1993; Salas *et al.*, 2001).

The ammonite biostratigraphic analysis reported in this study is focused on the Hauterivian to Albian stages (Lower Cretaceous). Ammonites occur within 8 lithostratigraphic

units named, from older to younger: Gaita Formation (Hauterivian), Morella Formation (Barremian? Aptian?), Forcall Formation (Aptian), Villarroya de los Pinares Formation (Aptian), Benassal Formation (Aptian), Dehesa Formation (Aptian), Escucha Formation (Albian) and an Albian unit of green marls lacking a formal name (Figs. 2 and 3).

The Gaita Formation is constituted by a hemipelagic succession of marls and limestones (Salas, 1987). The Morella Formation is made up of fluvial sandstones and red clays, which were occasionally influenced by tides (e.g., Canérot *et al.*, 1982). This formation is characterized by the presence of dinosaur remains. The Forcall Formation corresponds to a succession of marls, marly limestones and limestones. This is the most prolific formation in terms of ammonoid content of all the lithostratigraphic units studied (e.g., Moreno-Bedmar *et al.*, 2010a). Other macrofossils present, or locally characteristic, within this formation are orbitolinids, nautiloids, brachiopods, vermetid gastropods, solitary corals, serpulids, echinoids, *Plicatula placuna* and *Lithocodium aggregatum* (e.g., Bover-Arnal *et al.*, 2010; 2011). The Villarroya de los Pinares Formation is mainly made up of platform top carbonates, which change laterally to slope facies exhibiting marl-limestone alternations (Bover-Arnal *et al.*, 2009). Characteristic fossil content found within this formation include corals, *Chondrodonta* and rudists such as *Toucasia carinata*, *Caprina parvula* and *Polyconites hadriani* (e.g., Bover-Arnal *et al.*, 2010; Skelton *et al.*, 2010). The Benassal Formation is primarily formed by very-shallow marine to transitional alternations of clays, marls, marly limestones and limestones. Common skeletal components are orbitolinids and shells of *Toucasia* sp., nerineid gastropods and brachiopods (e.g., Canérot *et al.*, 1982; Bover-Arnal *et al.*, 2010). The Dehesa Formation is made up of a succession of marls and skeletal limestones. This formation is the lateral equivalent of the Benassal Formation in the Oliete Sub-basin (Canérot *et al.*, 1982). The Escucha Formation mainly comprises a succession of continental clays, sandstones and coal. However, its base, where the ammonites occur, presents a transgressive interval with marine calcarenitic limestones and marls (Canérot *et al.*, 1982; Moreno-Bedmar *et al.*, 2008). The Albian informal unit of green marls analysed corresponds to transitional deposits stratigraphically located between the fluvial sandstones of the Utrillas Formation and the marine limestones and marls of the Mosqueruela Formation (Canérot *et al.*, 1982).

## 3. Material

This study is mainly based on more than 3000 specimens of ammonites collected during the last 30 years from numerous Lower Cretaceous outcrops in the Maestrat Basin (Fig. 1). The localities sampled are the following. In the Salzedella Sub-basin: La Torre d'en Doméneç section (Moreno *et al.*, 2007), the area of Traiguera (Moreno-Bedmar *et al.*, 2008), and in the environs of the village of Sant Mateu, we stud-

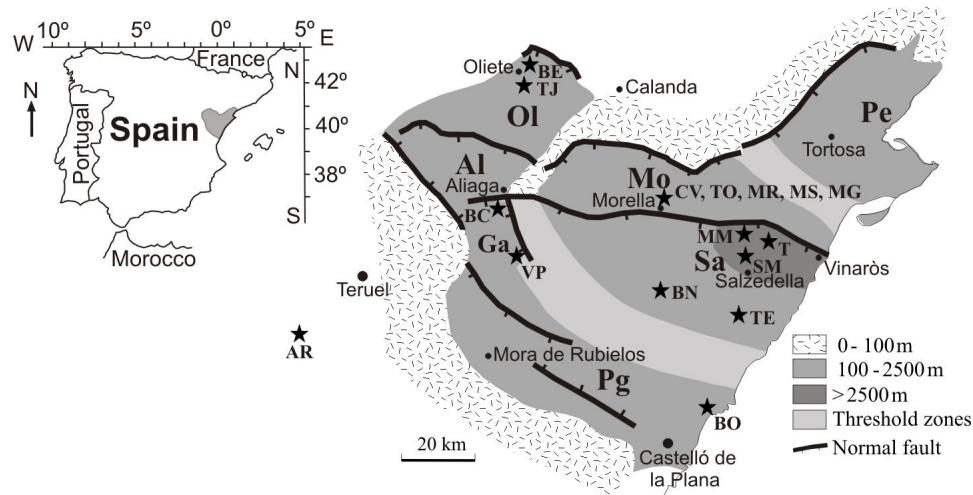


Fig. 1.- Simplified paleogeographic map of the Maestrat Basin during the Early Cretaceous (modified after Salas *et al.* 2001), showing the location of the sections and studied areas. TE: La Torre d'en Doméneç section; SM: San Mateu area; T: Traiguera area; MM: Mola Murada section; BO: Benicàssim-Orpesa area; BN: Benassal area; BC: Barranco de las Calzadas section; VP: Villarroya de los Pinares section; CV: Cap de Vinyet section; TO: Todolella section; MR: Mas del Roc section; MS: Mas Segura section; TJ: Tejería de Josa section; BE: Barranco Emilia section; AR: Arroyofrío area; Ol: Oliete Sub-basin; Al: Aliaga Sub-basin; Mo: Morella Sub-basin; Pe: Perelló Sub-basin; Ga: Galve Sub-basin; Sa: Salzedella Sub-basin; Pg: Penyagolosa Sub-basin.

ied Les Bassetes del Cocó de l'Estaca and La Gaita sections (Tomás, 2007). In the Penyagolosa Sub-basin: the Cingle del Morral section and La Ferradura area (Benicàssim-Orpesa area; Moreno-Bedmar *et al.*, 2009b), and the surroundings of Xodos and Benassal. In the Galve Sub-basin: the Barranco de las Calzadas and Villarroya de los Pinares sections (Bover-Arnal *et al.*, 2010 and Moreno-Bedmar *et al.*, 2010a). In the Morella Sub-basin: the Mola Garumba area, and the sections of Cap de Vinyet, Todolella, Mas del Roc and Mas Segura (Moreno-Bedmar *et al.*, 2010a). In Oliete Sub-basin: the Tejería de Josa and Barranco Emilia sections (Moreno-Bedmar *et al.*, 2010a). In the south of l'Espiradà we studied the area of Arroyofrío (Geyer, 1995a, b).

The ammonoids studied were deposited in museums (Mu-seu de Valltorta: MV, València; Conjunto Paleontológico de Teruel: CPT, Aragón) or in university collections (Colleccions de Paleontologia de la Universitat Autònoma de Barcelona: PUAB, Bellaterra, Catalunya). We also studied ammonoids from private and museum collections (Museu Paleontològic Juan Cano Forner in Sant Mateu; Rafael Martí: RM, Andrés Cumba: AC and Domingo Tolós: DT in La Jana, València; Museu Geològic del Seminari Conciliar de Barcelona: MGSB, Barcelona). Finally, ammonoids figured in the literature were also analyzed.

#### 4. Ammonite biostratigraphy

Ammonoids are generally rare and sparsely distributed throughout the Lower Cretaceous succession of the Maestrat Basin. Only few individual beds or marly intervals yielded abundant material. However, the large time span comprised by the ammonoid collection allows us to present a very complete record.

##### 4.1. Lower Hauterivian

Few works quote ammonoids of this age, all of which were found in the Salzedella Sub-basin (Fig. 1). Canérot (1974) listed some species from the environs of the village of Sant Mateu, La Torre d'en Besora and Massís de la Serratella localities. Neumann (1987) identified some ammonoids of the *Acanthodiscus radiatus* Zone in La Torre d'en Doméneç and others from the Lower Hauterivian of La Torre d'en Besora. In La Torre d'en Doméneç, Moreno *et al.* (2007) identified two biozones: *Acanthodiscus radiatus* and *Crioceratites loryi*. In the present work, we recognized all the ammonite zones of the Lower Hauterivian (Fig. 2).

*Acanthodiscus radiatus* Zone. This zone has been recognized by the occurrence of *Breistrofferella cf. castellanensis* (d'Orbigny) in La Torre d'en Doméneç (Fig. 4 E).

*Crioceratites loryi* Zone. Also in La Torre d'en Doméneç, we have found *Olcostephanus (Jeannoticeras) jeannotti* (d'Orbigny) and *Olcostephanus sayni* (Kilian) (Fig. 4 H) species that characterize the *Crioceratites loryi* Zone. In particular, *O. jeannotti* is currently used to define its upper subzone.

*Lyticoceras nodosoplicatum* Zone. In the same locality, this zone is represented by the index taxa *Lyticoceras nodosoplicatum* (Kilian and Reboulet) (Fig. 4 A-B, F-G, I) and *Lyticoceras kiliani* Reboulet and Atrops (Fig. 4 C-D).

##### 4.2. Upper Hauterivian

The ammonoids of this age come from the Salzedella Sub-basin as well (Fig. 1). Canérot (1974) listed some species from La Torre d'en Besora. Neumann (1987) reported some ammonites from the Tossal d'Orenga and assigned them to the *Plesiospitiidiscus ligatus* and *Balearites balearis* Zones.

## Salzedella Sub-basin

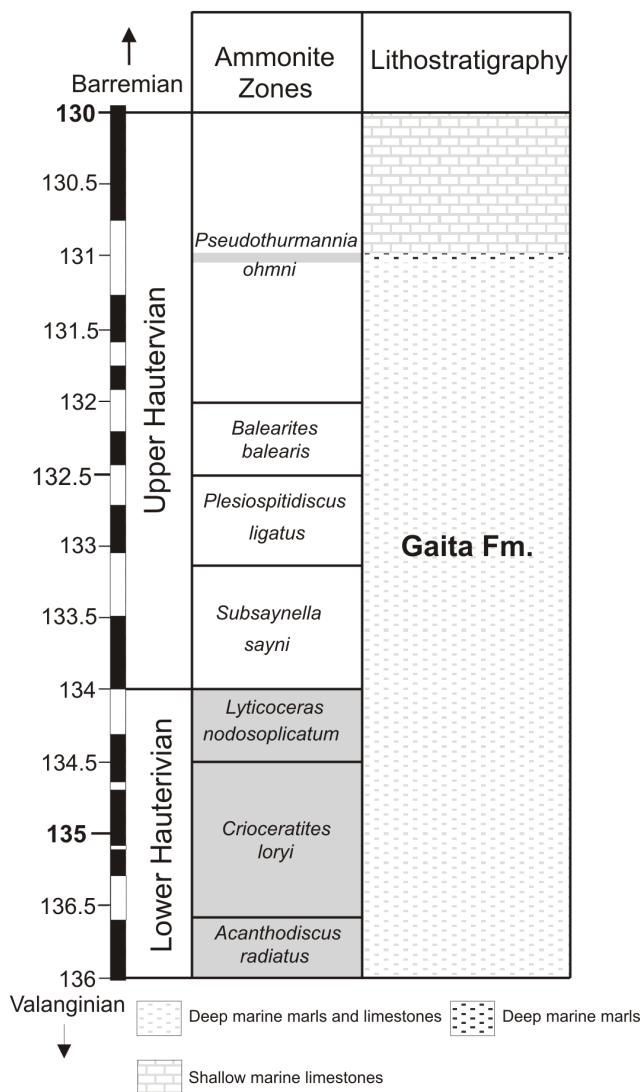


Fig. 2.- Chronostratigraphy of the Hauterivian deposits in the Salzedella Sub-basin (Maestrat Basin). The ammonite zones recognized are indicated by the shaded levels. Absolute ages after Gradstein *et al.* (2004).

Martínez *et al.* (1994) studied the ammonoid record of the Les Bassetes del Cocó de l'Estaca and also recognized the *Plesiospitidiscus ligatus* and *Balearites balearis* Zones. They also reported one specimen of the Castellfort-Coll d'Ares that they attributed to the species *Crioceratites* (*Pseudothurmannia*) *grandis*, from the *Pseudothumannia angulicostata* Zone (=*Pseudothumannia ohmi* Zone of Reboulet *et al.*, 2011).

*Pseudothumannia ohmi* Zone. In this work we have studied an interesting pyritized fauna coming from Les Bassetes del Cocó de l'Estaca, and La Gaita. We have identified: *Barremites primitivus* Cecca, Faraoni and Marini (Fig. 5 A-E, F, I-J, N-P), *Plesiospitidiscus* sp. (Fig. 5 K) and *Pseudothurmannia* sp. (Fig. 5 G-H, L-M). This fauna probably indicates the middle-upper part of the zone (*Pseudothumannia catullo* and/or *Pseudothumannia picteti* Subzones).

### 4.3. Upper Barremian

López Llorens (2007) reported the finding of a sole Barremian ammonite from the Morella Sub-basin, a juvenile heterceratid that he attributed to the *Imerites giraudi* (Kilian). The revision of this specimen for the present work indicates that the determination was probably wrong and this ammonoid is most likely an *Argvethites* sp. The genus *Argvethites* sp. is present in the uppermost Barremian (*Imerites giraudi* Zone, spanning both the *Imerites giraudi* and the *Martelites sarasini* Subzones (Delanoy, 1997).

### 4.4. Lower Aptian

The Lower Aptian ammonite record is the best-known record in the Maestrat Basin (Moreno-Bedmar *et al.*, 2010a). These authors recognized four zones and five subzones some of which are new, and were recently accepted and published in the biostratigraphic scheme of the latest standard Mediterranean zonation (Reboulet *et al.*, 2011). More recently, studies such as Moreno-Bedmar and Garcia (2011) and Delanoy *et al.* (2013) have helped to complete the knowledge about the Lower Aptian ammonoid record.

*Deshayesites oglanlensis* Zone. Moreno-Bedmar *et al.* (2010a) and Moreno-Bedmar and Garcia (2011) identified the *Deshayesites oglanlensis* Zone and proposed a new *Deshayesites lupovi* Subzone in the Morella Sub-basin (Figs. 1 and 3). In this zone, the following species were recognized: *Procheloniceras* sp., *Deshayesites antiquus* Bogdanova (Fig. 6 A), *Deshayesites lupovi* Bogdanova (Fig. 6 B), and *Deshayesites* sp. cf. *oglanlensis* Bogdanova.

*Deshayesites forbesi* Zone. Moreno-Bedmar *et al.* (2010a) proposed the inclusion of this boreal zone in the Tethyan zonation. This zone was identified in the Oliete, Morella and Galve sub-basins (Fig. 1). They also proposed a *Roloboceras hambrovi* Subzone (Fig. 3). Recently, Grauges *et al.* (2010) studied the taxonomy of the Lower Aptian desmoceratid fauna (*Pseudosaynella* and *Pseudohaploceras* genera) of this zone in the Oliete Sub-basin. Moreno-Bedmar and Garcia (2011) provided new data from the ammonoid record on the lower part of this zone in the Morella Sub-basin through the collection of new specimens, including the best specimen of the index species known to date of this biozone, *Deshayesites forbesi*. The species recognized in this zone are: *Pseudosaynella raresulcata* (d'Orbigny) (= *Pseudosaynella undulata* (Sarasin), see Grauges *et al.*, 2010, p. 9), *Pseudosaynella bicurvata* (Michelin), *Pseudosaynella* sp., *Barremites* sp., *Pseudohaploceras liptoviense* (Zeuschner), *Pseudohaploceras* sp., *Toxoceratoides royerianus* (d'Orbigny), *Ancylceras?* sp., *Proaustraliceras* sp., *Lithancylus* cf. *grandis* (Sowerby), *Cheloniceras* sp., *Roloboceras hambrovi* (Forbes), *Roloboceras hispanicum* Sornay and Marin, *Roloboceras* sp., *Megatyloceras coronatum* (Rouchadzé), *Megatyloceras transiens* Casey, *Deshayesites euglyphus* Casey, *Deshayesites* gr. *euglyphus/spathi* Casey, *Deshayesites forbesi* Casey (Fig. 6 C-D), *Deshayesites spathi* Casey and *Deshayesites* sp.

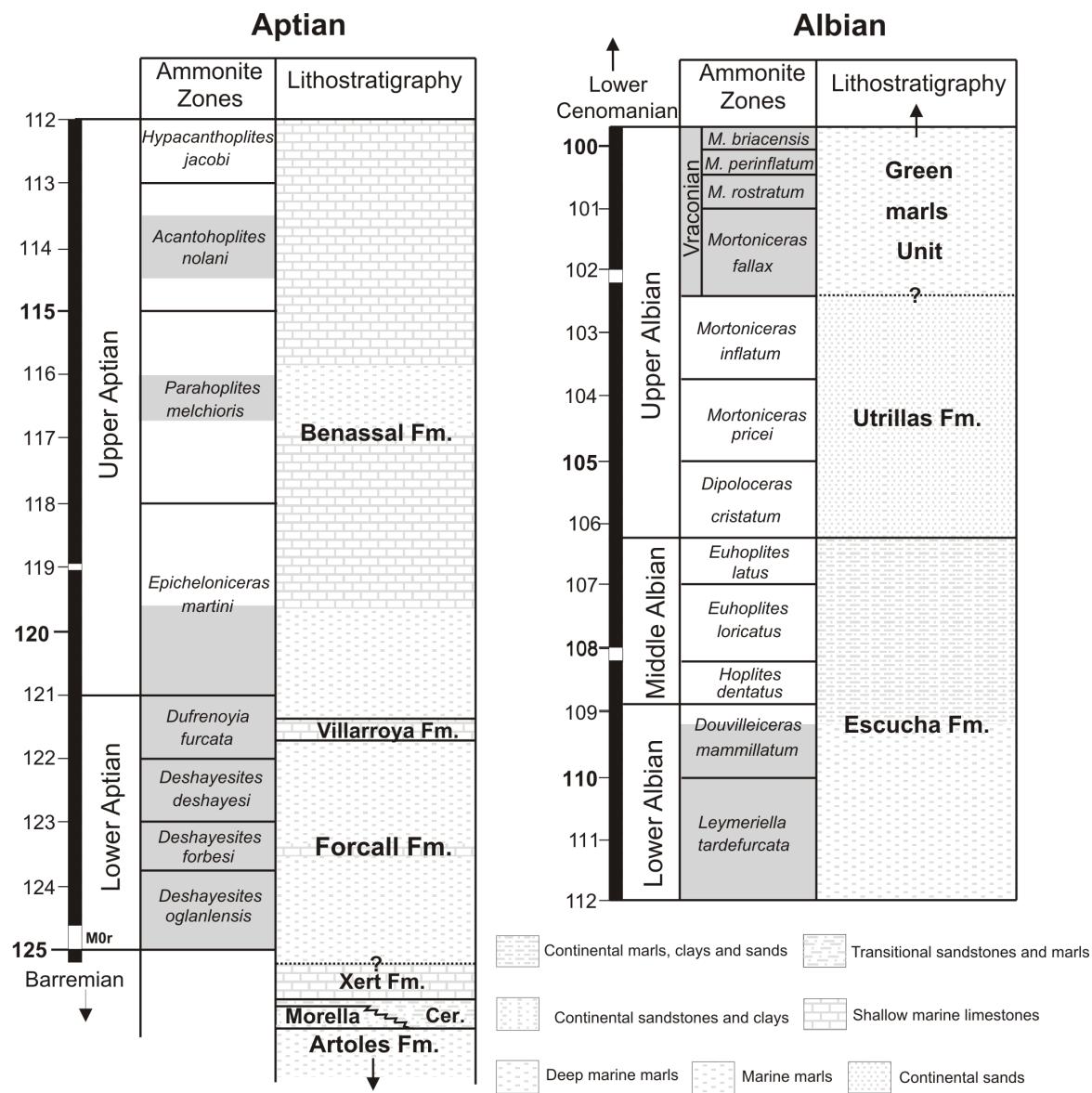


Fig. 3.- Chronostratigraphy of the Aptian-Albian deposits in the Maestrazgo Basin. The ammonite zones recognized are indicated by the shaded levels. Absolute ages after Gradstein *et al.* (2004).

*Deshayesites deshayesi* Zone. Moreno-Bedmar *et al.* (2010a) identified the *Deshayesites deshayesi* Zone and its Subzone, *Deshayesites grandis*, in the Morella, Galve, Salzedella and Penyagolosa sub-basins (Figs. 1 and 3). This zone contains the species: *Phylloceras* sp., *Lytoceras bellisepatum*, *Anthula*, *Pseudosaynella raresulcata* (d'Orbigny), *Pseudosaynella bicurvata* (Michelin), *Pseudohaploceras* sp., *Ammonitoceras* (*Epancyloceras*) sp., *Proaustraliceras* cf. *gigas* (Sowerby), *Toxoceratoides royerianus* (d'Orbigny), *Cheloniceras cornuelianum* (d'Orbigny), *Deshayesites deshayesi* (d'Orbigny) (Fig. 6 E-G), *Deshayesites vectensis* (Spath) (Fig. 6 H), *Deshayesites latilobatus/involutus* group, *Deshayesites grandis* (Spath) and *Deshayesites* sp. Recently a new genus and species was described in Salzedella Sub-basin, *Xerticeras salasi* Delanoy, Moreno-Bedmar, Ruiz, Tolós Lladser, 2013. This new taxon is present in *Deshayesites deshayesi* and *Dufrenoya furcata* zones.

*Dufrenoya furcata* Zone. Moreno-Bedmar *et al.* (2010a, 2010b) identified this zone in the Perelló, Galve, Salzedella and Penyagolosa sub-basins (Fig. 1). They also recognized the two subzones proposed by Dutour (2005) in the Vocontian Basin (France): *Dufrenoya furcata* and *Dufrenoya dufrenoyi*. Moreno-Bedmar *et al.* (2010b) and Garcia and Moreno-Bedmar (2010) contributed to the knowledge of the *Dufrenoya furcata* Zone in the Perelló Sub-basin where they studied a rich record of pyritized ammonites. Moreno-Bedmar *et al.* (2011, 2012b) studied some species of the *Dufrenoya* genus from the uppermost part of this zone in the Morella and Galve sub-basins. The following species were recognized in this zone: *Phyllopachyceras baborens* (Coquand), *Salfeldiella* sp., *Pseudosaynella raresulcata* (d'Orbigny), *Pseudohaploceras* sp., *Macroscaphites striasulcatus* (d'Orbigny), *Aconeoceras nisum* (d'Orbigny), *Cheloniceras cornuelianum* (d'Orbigny), *Dufrenoya furcata*

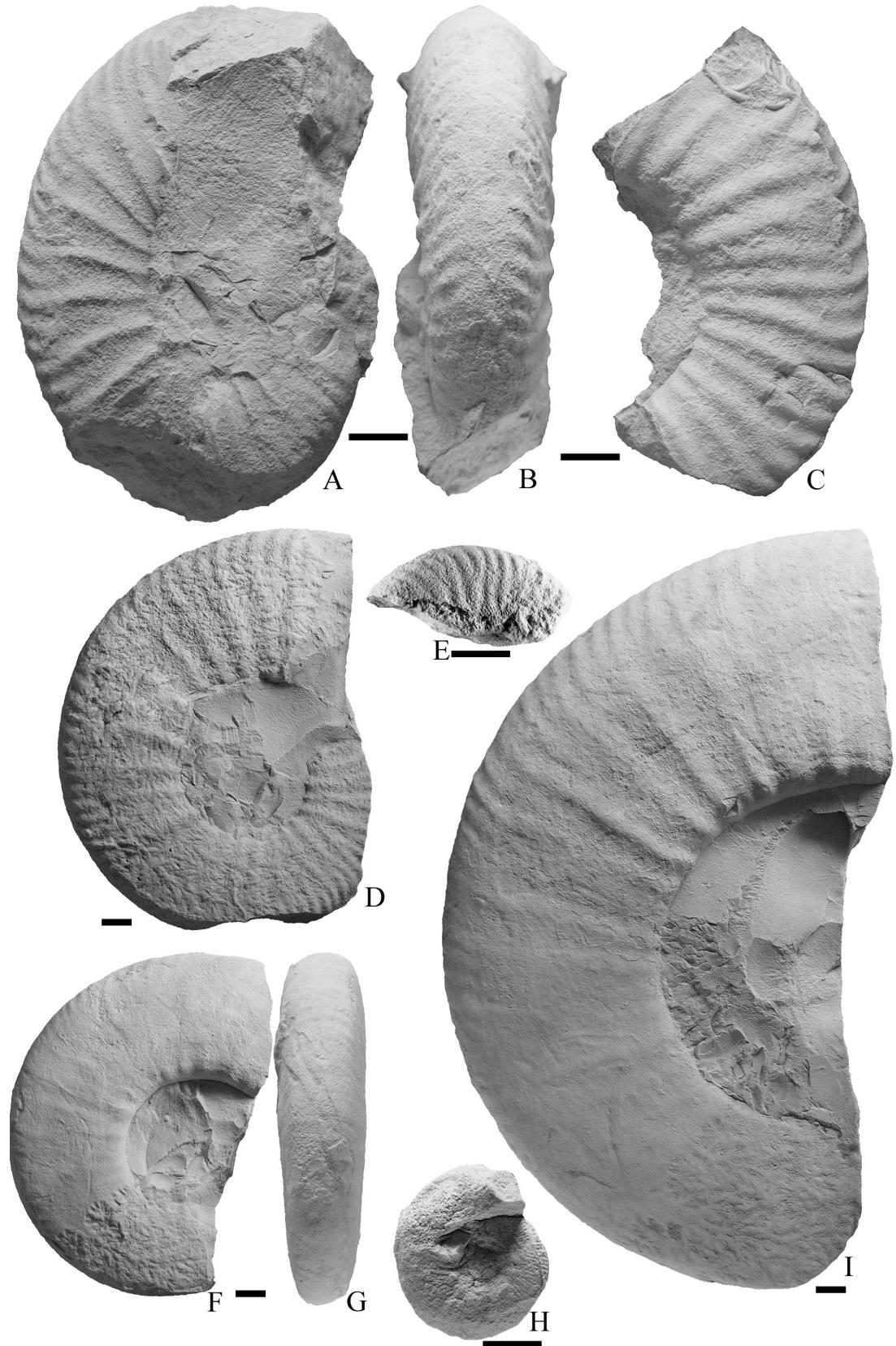


Fig. 4.- A-B: *Lyticoceras nodosoplicatum* lateral and ventral views of the specimen PUAB 88777, La Torre d'en Doméneç. C: *Lyticoceras kiliani* lateral view of the specimen PUAB 88776, La Torre d'en Doméneç. D: *Lyticoceras kiliani* lateral view of the specimen PUAB 88780, La Torre d'en Doméneç. E: *Breistrofferella* cf. *castellanensis* lateral view of the specimen PUAB 48651, La Torre d'en Doméneç. F-G: *Lyticoceras nodosoplicatum* lateral and ventral views of the specimen PUAB 88778, La Torre d'en Doméneç. H: *Olcostephanus sayni* lateral view of the specimen PUAB 48650, La Torre d'en Doméneç. I: *Lyticoceras nodosoplicatum* lateral view of the specimen PUAB 88781, La Torre d'en Doméneç. Scale bar=10 mm.

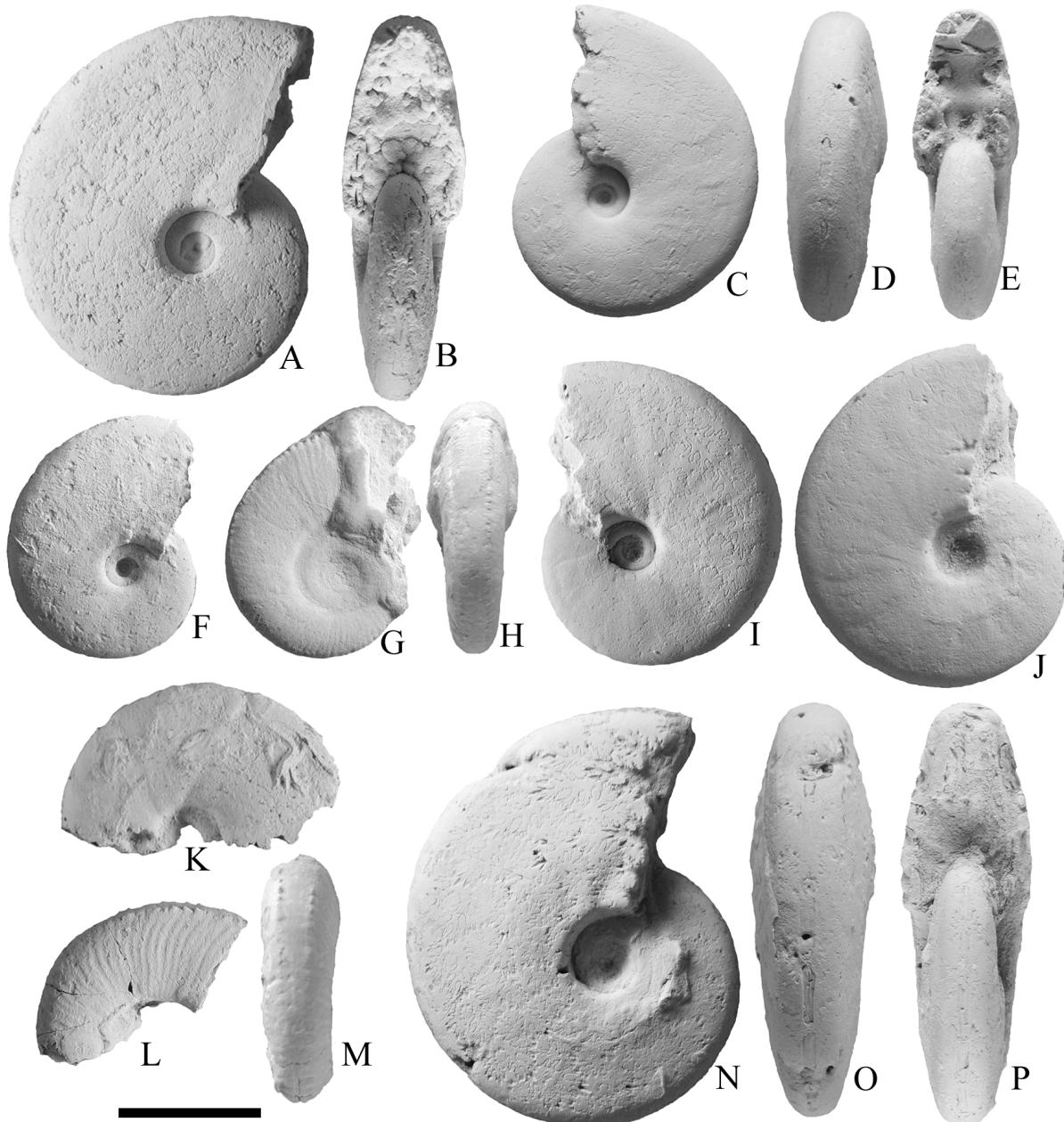


Fig. 5.- A-B: *Barremites primitivus* lateral and frontal views of the specimen PUAB 12533, La Gaita. C-E: *Barremites primitivus* lateral, ventral and frontal views of the specimen PUAB 12529, Les Bassetes del Cocó de l'Estaca. F: *Barremites primitivus* lateral view of the specimen PUAB 88798, La Gaita. G-H: *Pseudothurmannia* sp. juv. lateral and ventral views of the specimen PUAB 12530, Les Bassetes del Cocó de l'Estaca. I: *Barremites primitivus* lateral view of the specimen PUAB 88795, Les Bassetes del Cocó de l'Estaca. J: *Barremites primitivus* lateral view of the specimen PUAB 88794, Les Bassetes del Cocó de l'Estaca. K: *Plesiospitidiscus* sp. lateral and ventral views of the specimen PUAB 12530, Les Bassetes del Cocó de l'Estaca. L-M: *Pseudothurmannia* sp. juv. lateral and ventral views of the specimen PUAB 88790, Les Bassetes del Cocó de l'Estaca. N-P: *Barremites primitivus* lateral, ventral and frontal views of the specimen PUAB 88793, Les Bassetes del Cocó de l'Estaca. Scale bar=10 mm.

(Sowerby), *Dufrenoyia* cf. *furcata* (Sowerby) (Fig. 6 I-N), *Dufrenoyia dufrenoyi* (d'Orbigny), *Dufrenoyia* cf. *dufrenoyi* (d'Orbigny), *Dufrenoyia* cf. *transitoria* Casey, *Dufrenoyia* cf. *justinae* (Hill) (Fig. 6 O-P) and *Colombiceras* sp.

#### 4.5. Upper Aptian

*Epicheloniceras martini* Zone. This zone has been identified only in the Benicàssim-Orpesa area, Penyagolosa Sub-

basin (Fig. 1) (Moreno-Bedmar et al., 2009). The ammonites reported are: *Caseyella* sp. (Fig. 7 C), *Colombiceras* sp., *Epicheloniceras* sp. (Fig. 7 D-E) and *Epicheloniceras* cf. *debile*.

*Parahoplites melchioris* Zone. Ammonites characteristic of this zone have been cited in several previous papers. Sornay and Marin (1971) reported the occurrence of *Acanthohoplites* sp. and *Parahoplites* sp. from Cañizar del Olivar (Oliete Sub-basin) and of *Parahoplites* cf. *cunningtoni* from Cantavieja (Galve Sub-basin). Martínez et al. (1994) identified the am-

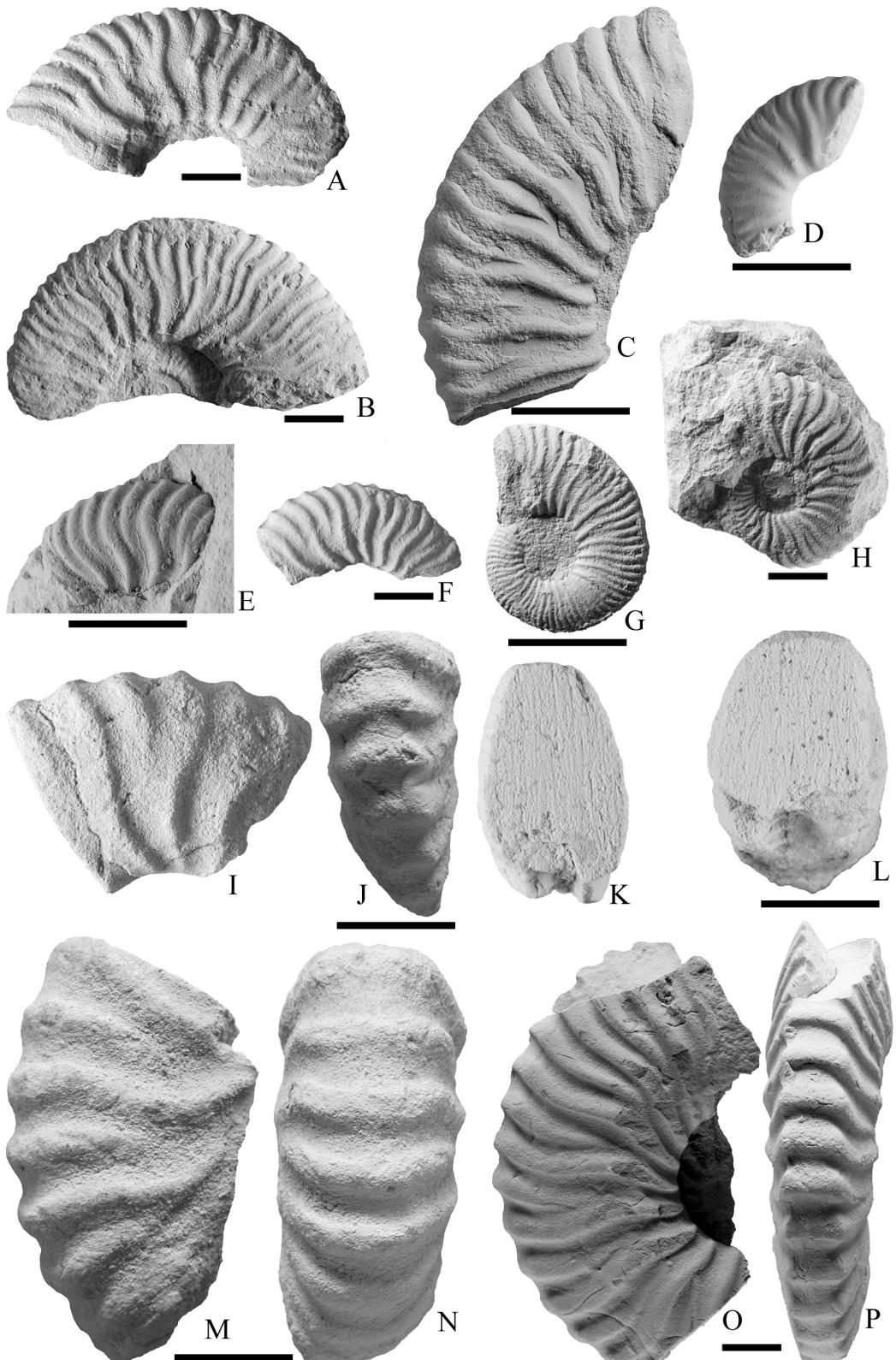


Fig. 6.- A: *Deshayesites antiquus* lateral view of the specimen PUAB 12535, Cap de Vinyet. B: *Deshayesites lupppovi* lateral view of the specimen MV 900.1, Mas Segura section. C: *Deshayesites forbesi* lateral view of the specimen PUAB 88765, Cap de Vinyet section. D: *Deshayesites forbesi* lateral view of the specimen PUAB 48071, Tejería de Josa section. E: *Deshayesites deshayesi* lateral view of the specimen MV 425.2, Mola Murada section. F: *Deshayesites deshayesi* lateral view of the specimen MV 471.1, Todolella section. G: *Deshayesites deshayesi* lateral view of the specimen MV 425.1, Mola Murada section. H: *Deshayesites vectensis* lateral view of the specimen CPT-3599, Barranco de las Calzadas section. I-K: *Dufrenoyia* cf. *furcata* lateral, ventral and whorl section views of the specimen PUAB 88788, Mola de la Garumba. L-N: *Dufrenoyia* cf. *furcata* whorl section, lateral and ventral views of the specimen PUAB 88789, Mola de la Garumba. O-P: *Dufrenoyia* cf. *justinae* lateral and ventral views of the specimen CPT-3728, Barranco de las Calzadas section. Scale bar=10 mm.

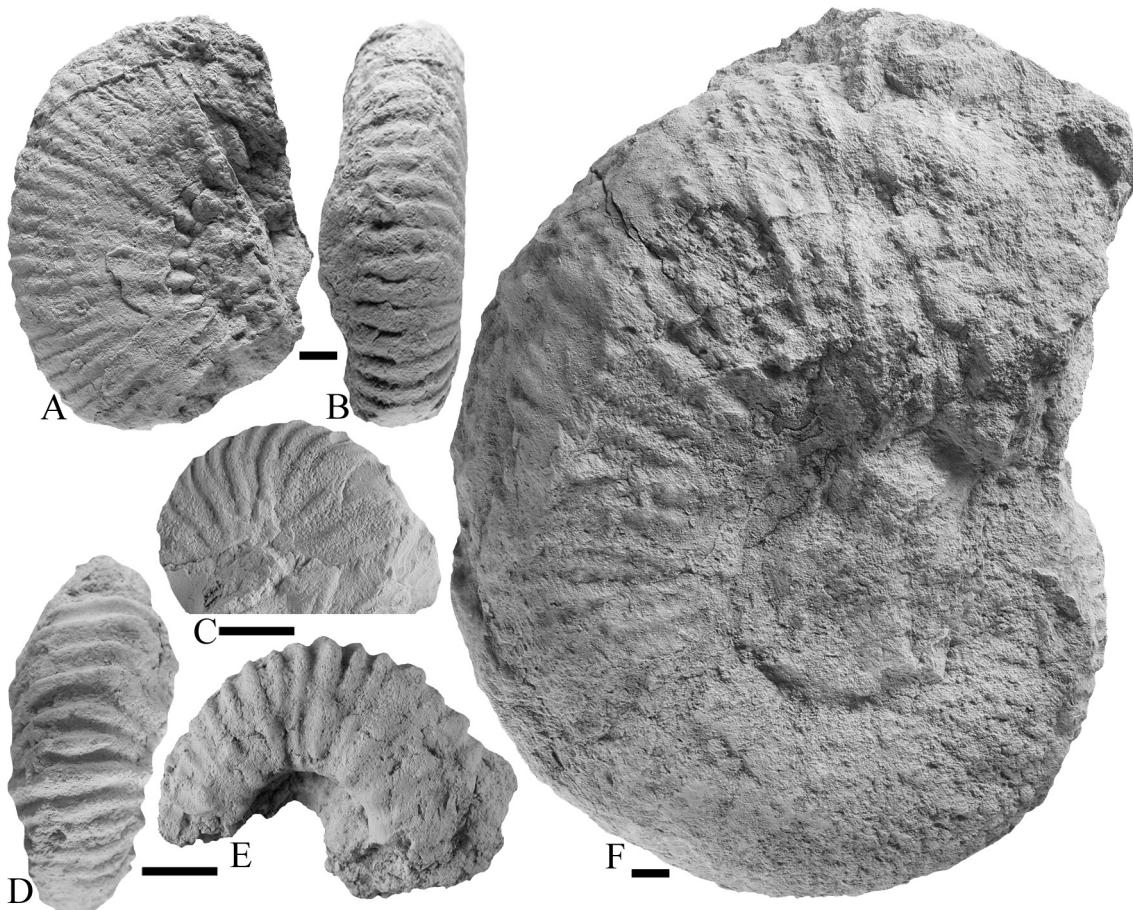


Fig. 7.- A-B: *Parahoplites* cf. *vectensis* lateral and ventral views of the specimen PUAB 68597, La Ferradura. C: *Caseyella* sp. lateral view of the specimen PUAB 68593, Cingle del Morral. D-E: *Epicheloniceras* sp. ventral and lateral views of the specimen PUAB 68596, Cingle del Morral. F: *Parahoplites nutfieldiensis* lateral view of the specimen PUAB 12307, Xodos. Scale bar=10 mm.

monoid *Parahoplites maximus* in the Oliete Sub-basin and *Parahoplites nutfieldiensis* in Xodos (Penyagolosa Sub-basin). Moreno-Bedmar *et al.* (2009b) recorded *Parahoplites* cf. *vectensis* in the Benicàssim-Orpesa area (Penyagolosa Sub-basin). The revision of ammonoids of this zone for the present work resulted in the following identifications: *Parahoplites nutfieldiensis* (Sowerby) (Fig. 7 F) from Xodos, *Parahoplites* cf. *vectensis* Casey (Fig. 7 A-B) from the Benicàssim-Orpesa area and *Parahoplites* sp. in the environs of the village of Benassal (Salzedella Sub-basin).

*Acanthohoplites nolani* Zone. Weisser (1959) described a specimen of *Acanthohoplites bergeroni* (Seunes) from Vilaloya de los Pinares (Galve Sub-basin) and Martínez *et al.* (1994) reported the same taxon from the Oliete Sub-basin. We have not been able to revise these specimens or collect other ammonites from these sub-basins.

#### 4.6. Lower Albian

The Albian ammonoid occurrences are essentially restricted to the environs of Traiguera (Salzedella Sub-basin) (Fig. 1). Canérot and Collignon (1981) first reported these

ammonoids and ascribed them to the Upper Aptian. Later, Martínez *et al.* (1994) assigned the same record to the Lower Albian, and this age-attribution has been recently confirmed by Moreno-Bedmar *et al.* (2008).

*Leymeriella tardefurcata* Zone. Martínez *et al.* (1994) attributed the ammonoid record of the Traiguera area to the *Douvilleiceras mammillatum* Zone. However, they noted that some of the species could also be assigned to the *Leymeriella tardefurcata* Zone. This record was studied by Moreno-Bedmar *et al.* (2008), who confirmed that the lower part of the ammonoid record of Traiguera belongs to the *Leymeriella tardefurcata* Zone. This zone is represented by the species: *Douvilleiceras* gr. *leightonense* Casey (Fig. 8 D-E), *Parenconoceras basei* (Batallier), *Parenconoceras* cf. *ebrayi* (Loriol) (Fig. 8 C), *Parenconoceras caneroti* Collignon (Fig. 8 G, J), “*Hypacanthoplites*” sp. 1 (Fig. 8 A-B) and “*Hypacanthoplites*” sp. 2 (Fig. 8 F).

*Douvilleiceras mammillatum* Zone. Martínez *et al.* (1994) identified *Douvilleiceras monile* (Sowerby) in the Traiguera area and *Douvilleiceras mammillatum* (Schlotheim) (Fig. 8 H-I) in Benassal (Salzedella Sub-basin). We had access to these specimens and confirm these identifications.

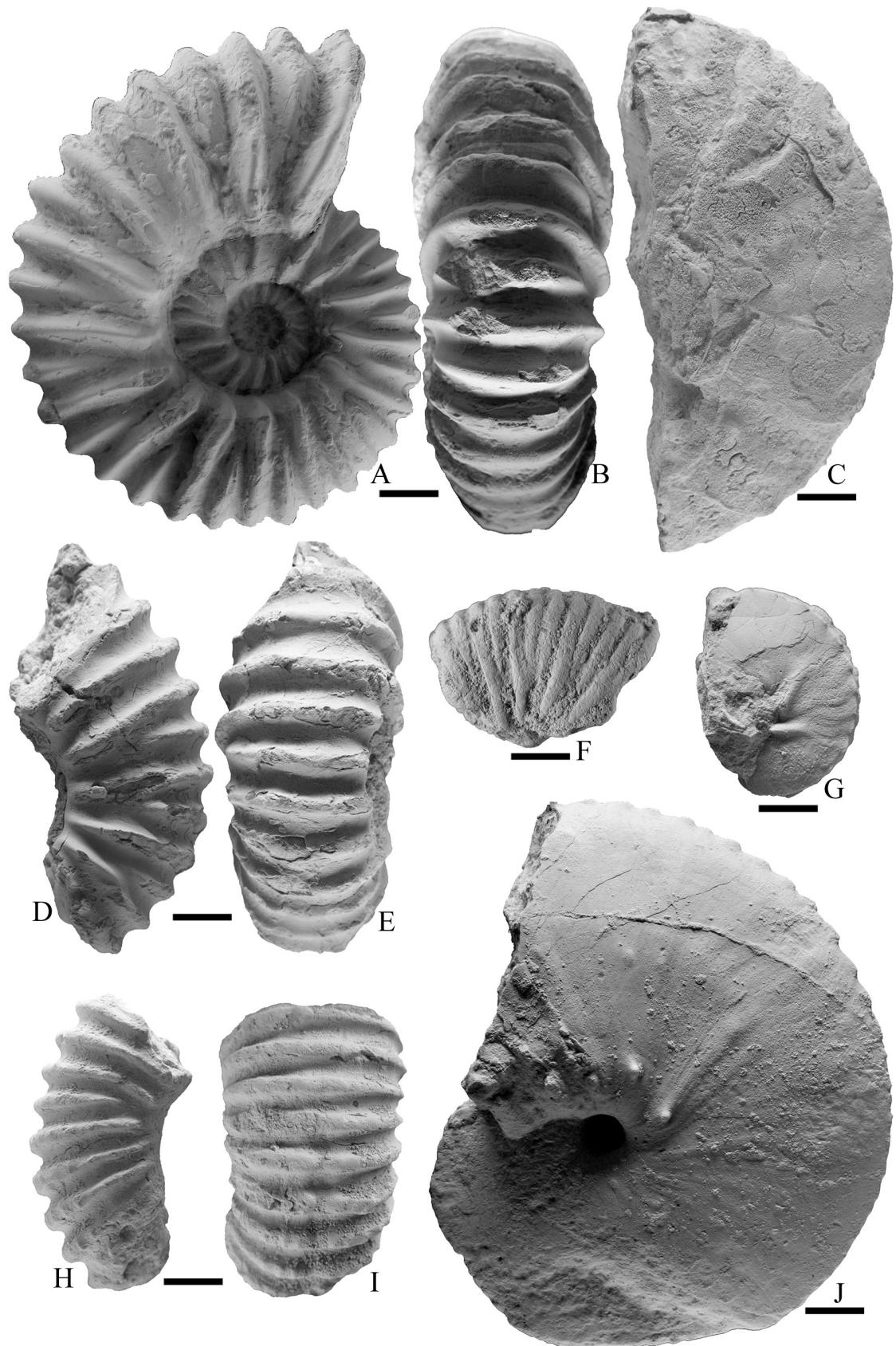
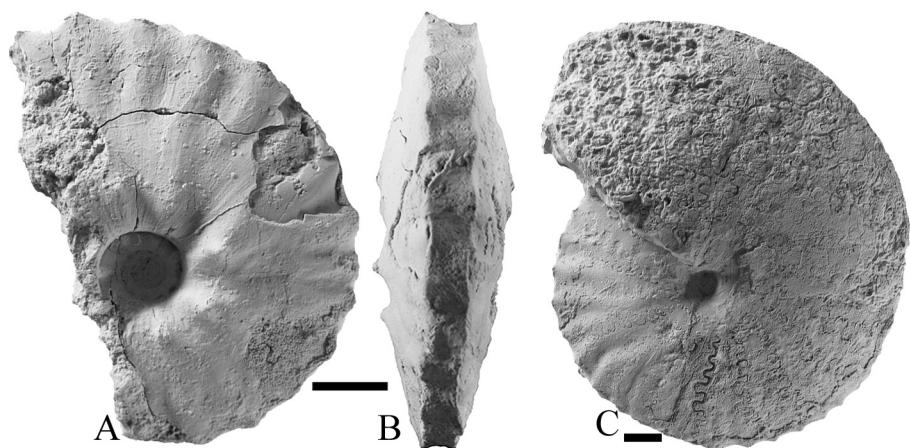


Fig. 8.- A-B: "*Hypacanthoplites*" sp.1 lateral and ventral views of the specimen RM 1, Traiguera. C: *Parengonoceras cf. ebrayi* lateral view of the specimen AC 10, Traiguera. D-E: *Douvilleiceras gr. leightonense* lateral and ventral views of the specimen AC 5, Traiguera. F: "*Hypacanthoplites*" sp.2 lateral view of the specimen AC 2, Traiguera. G: *Parengonoceras caneroti* lateral view of the specimen RM 2, Traiguera. H-I: *Douvilleiceras mammillatum* lateral view of the specimen MGSB 41829, Benassal. J: *Parengonoceras caneroti* lateral view of the specimen AC 3, Traiguera. Scale bar=10 mm.

Fig. 9.- A-B: *Engonhoplitoides* cf. *vicorpense* lateral and ventral views of the specimen MGSB 55884, Arroyofrio. C: *Engonhoplitoides* cf. *vicorpense* lateral view of the specimen MGSB 55886, Arroyofrio. Scale bar=10 mm.



## 5. Novel and relevant biostratigraphic considerations

The La Gaita Formation (Fig. 2), which crops out in the Salzedella Sub-basin (Fig. 1), comprises the whole Hauterivian. The relative sea-level rose during the Late Valanginian to the Early Hauterivian, thus allowing the record of all of the Lower Hauterivian ammonite zones. During the Late Hauterivian, the relative sea-level remained low except in the middle part of the *Pseudothumannia ohmi* Zone, and it is thus the only ammonoid record that exists in the Upper Hauterivian of the Maestrat Basin. This age corresponds with the maximum flooding interval of the Ha6 global sequence of Handerbol *et al.* (1998) that it is associated with the Faraoni level (Company *et al.*, 2005; Gradstein *et al.*, 2004). This important transgression is recorded in the La Gaita Formation as a marly interval 10 to 20 meters thick that contains a well-preserved pyritized ammonite record.

*Argyethites* sp. is the only ammonoid found in the Morella Formation. According to this, at least the lower and middle parts of the formation should be assigned to the Upper Barremian. This would support the hypothesis exposed by Moreno-Bedmar and García (2011), who proposed that the Barremian-Aptian boundary is located in the basal part of the Forcall Formation. This proposal is based on the fact that the species *Deshayesites antiquus* Bogdanova and *Deshayesites* sp. cf. *oglanlensis* Bogdanova, which are characteristic of the lower part of the first Aptian ammonite zone (*Deshayesites oglanlensis* Zone), are found in the lower, non-basal part of the Forcall Formation. Moreno-Bedmar and García (2011) also compared the position of the Barremian-Aptian boundary with the Provençal Platform (France) and the Organyà Basin (Spain), both of which have a similar marine transgressive event starting in the uppermost Barremian. The basal marine transgressive deposits in both areas have an ammonoid record of the *Pseudocrioceras* genus that belongs to the uppermost Barremian. Unfortunately, to date no ammonoids have been found in the basal materials of the Forcall Formation and the possible position of the Barremian-Aptian boundary in the lower part of this formation remains only a

plausible hypothesis at the present time. Moreover, the possibility that the ammonite collected within the Morella Formation by López Llorens (2007) corresponds to a reworked specimen should not be ruled out.

All of the Lower Aptian ammonite biozones are present in the Forcall Formation (Fig. 3). At least in the different sub-basins, the Forcall Formation recorded the ammonoids during the maximum flooding interval of the Ap3 global sequence of the Handerbol *et al.* (1998) that is associated with the Oceanic Anoxic Event 1a (Bover-Arnal *et al.*, 2010). This Lower Aptian record varies between the different sub-basins. For example, in the Oliete Sub-basin only one ammonoid zone is recorded while the Morella Sub-basin recorded the four zones. These differences depended on the depositional space available, which was controlled by the eustatic fluctuations and the local subsidence of each sub-basin.

In the Villarroya de los Pinares Formation, ammonoid occurrence was only reported from Las Mingachas locality (Galve Sub-basin). Specimens of *Dufrenoyia* sp. of the *Dufrenoyia furcata* Zone were found within a bed with a rudstone texture interpreted by Bover-Arnal *et al.* (2009) as a forced regressive wedge, and in the overlying slopes of a lowstand normal regressive carbonate platform (Bover-Arnal *et al.*, 2009; 2012). Throughout the Maestrat Basin, the age of the Villarroya de los Pinares Formation ranges between the *Deshayesites deshayesi* and the *Dufrenoyia furcata* zones, but it is usually restricted to the middle part of the *Dufrenoyia furcata* Zone (Fig. 3).

The Benassal Formation recorded two large-scale transgressive-regressive sequences (Bover-Arnal *et al.*, 2010). The ammonoids are present in the transgressive units, which are dominated by marls and marly limestones. The first transgressive unit corresponds to the Ap4 global sequence of Handerbol *et al.* (1998). This transgressive unit recorded two ammonite zones: the uppermost part of the *Dufrenoyia furcata* Zone and the *Epicheloniceras martini* Zone (Fig. 3). The basal beds of the Benassal Formation have been historically assigned to the Upper Aptian (Canérot *et al.*, 1982; Salas, 1987; Bover-Arnal *et al.*, 2010). However, in the Galve and

the Morella sub-basins we recently assigned these beds to the uppermost part of the Lower Aptian, specifically to the uppermost part of the *Dufrenoyia furcata* Zone (Moreno-Bedmar, 2010; Moreno-Bedmar et al., 2011, 2012b; Bover-Arnal et al., in press). The second transgressive unit recorded the *Parahoplites melchioris* Zone (Martín-Martín et al., 2013). In the upper part of the Benassal Formation in the Galve and the Oliete sub-basins a significantly scarce ammonoid record of the *Acanthohoplites nolani* Zone is preserved. This biozone comprises the maximum flooding interval of the Ap5 global sequence of Handerbol et al. (1998). In the Oliete Sub-basin, the Dehesa Formation, which is a lateral equivalent of the Benassal Formation, contains the *Parahoplites melchioris* and *Acanthohoplites nolani* zones.

The occurrence of Lower Albian ammonoids in the lower part of the Escucha Formation is probably linked to the strong transgressions corresponding to the global sequences Ap6, Al1, Al2 and/or Al3 of Handerbol et al. (1998). The continental deposits of the middle and upper parts of the formation have been assigned to the upper part of the Lower Albian and the entire Middle Albian based on the palynological studies of Villanueva-Amadoz et al. (2010, 2011). The fluvial sandstones of the Utrillas Formation overlie the Escucha Formation. Above the Utrillas Formation an informal unit of green marls crops out, which mark the transition between the Utrillas Formation and the marine limestones and marls of the Mosqueruela Formation (Canérot et al., 1982). These Green Marls were studied by Villanueva-Amadoz et al. (2011) who assigned them to the uppermost Albian. In the Maestrat Basin no ammonoids were found in these marls. However, some ammonoids were collected by Geyer (1995a, b) in the Arroyofrío locality (Montes Universales de Albaracín, Teruel, South Iberian Basin). We reviewed these ammonoids for the present study and assigned them to the species *Engonhoplitoides* cf. *vicorpense* (Basse) (Fig. 9 A-C). Although the taxonomic and biostratigraphic knowledge of this taxon is poor, these ammonoids can be assigned without doubt to the Upper Albian. This is in agreement with the results by Villanueva-Amadoz et al. (2011).

## 6. Conclusions

The analysis of the stratigraphic distribution of ammonites in the Maestrat Basin allowed us to establish the faunal succession for the Lower Cretaceous (Hauterivian-Albian) in this region. We recognized the following zones: *Acanthodiscus radiatus*, *Crioceratites loryi*, *Lyticoceras nodosoplicatum* (Lower Hauterivian); *Pseudothurmannia ohmi* (Upper Hauterivian); *Imerites giraudi* (Upper Barremian); *Deshayesites oglanensis*, *Deshayesites forbesi*, *Deshayesites deshayesi*, *Dufrenoyia furcata* (Lower Aptian); *Epicheloniceras martini*, *Parahoplites melchioris*, *Acanthohoplites nolani* (Upper Aptian); *Leymeriella tardifurcata* and *Douvilleiceras mammillatum* (Lower Albian). We also identified the following subzones: the *Olcostephanus (Jeannoticeras)*

*jeannoti* in the upper half of the *Crioceratites loryi* Zone, the *Deshayesites luppovi* in the upper half of the *Deshayesites oglanensis* Zone, the *Roloboceras hambrovi* in the upper half of the *Deshayesites forbesi* Zone, the *Deshayesites grandis* in the upper half of the *Deshayesites deshayesi* Zone and the twofold subdivision of the *Dufrenoyia furcata* Zone in two subzones, *Dufrenoyia furcata* and *Dufrenoyia dufrenoyi*.

Consequently, this study accurately determines the age of the lithostratigraphic units defined in this basin. In the lower part of the La Gaita Formation, we found recorded the three zones of the Lower Hauterivian. The uppermost part of this latter formation corresponds to the *Pseudothurmannia ohmi* Zone. The *Imerites giraudi* Zone occurs in the Morella Formation. In the Forcall Formation, we found the four zones of the Lower Aptian. The ammonite record suggests that the base of this formation could be of Late Barremian age. The age of Villarroya de los Pinares Formation ranges, in different areas of the basin, from the *Deshayesites deshayesi* to the *Dufrenoyia furcata* Zones. The base of the Benassal Formation corresponds to the upper part of the *Dufrenoyia furcata* Zone in most parts of the Maestrat Basin. The rest of the Benassal Formation is Late Aptian in age. The lower part of the Escucha Formation recorded the two Lower Albian biozones.

## 7. Acknowledgements

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