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1	Measurements of the Evaporation and Hygroscopic Response
2	of Single Fine-Mode Aerosol Particles Using a Bessel Beam
3	Optical Trap
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8	
9	A single horizontally-propagating zeroth order Bessel laser beam with a counter-propagating gas flow
10	was used to confine single fine-mode aerosol particles over extended periods of time, during which
11	process measurements were performed. Particle sizes were measured by the analysis of the angular
12	variation of light scattered at 532 nm by a particle in the Bessel beam, using either a probe beam at
13	405 nm or 633 nm. The vapour pressures of glycerol and 1,2,6-hexanetriol particles were determined
14	to be 7.5 $\pm$ 2.6 mPa and 0.20 $\pm$ 0.02 mPa respectively. The lower volatility of hexanetriol allowed better
15	definition of the trapping environment relative humidity profile over the measurement time period,
16	thus higher precision measurements were obtained compared to those for glycerol. The size evolution
17	of a hexanetriol particle, as well as its refractive index at wavelengths 532 nm and 405 nm, were
18	determined by modelling its position along the Bessel beam propagation length while collecting phase
19	functions with the 405 nm probe beam. Measurements of the hygroscopic growth of sodium chloride
20	and ammonium sulfate have been performed on particles as small as 350 nm in radius, with growth
21	curves well described by widely used equilibrium state models. These are the smallest particles for
22	which single-particle hygroscopicity has been measured and represent the first measurements of
23	hygroscopicity on fine mode and near-accumulation mode aerosols, the size regimes bearing the most
24	atmospheric relevance in terms of loading, light extinction and scattering. Finally, the technique is
25	contrasted with other single particle and ensemble methods, and limitations are assessed.
20	

### 27 I. Introduction

Atmospheric aerosols play important roles in regulating the Earth's climate and influencing 28 29 atmospheric composition, as well as impacting on human health, yet many of the fundamental physicochemical properties of aerosol remain poorly understood and quantified.<sup>1</sup> As well as their 30 ability to scatter solar radiation directly, atmospheric aerosols indirectly affect climate by providing 31 32 cloud condensation nuclei (CCN) for cloud droplets, impacting on the size, composition and lifetime of clouds.<sup>1,2</sup> Understanding the condensational growth of an aerosol and the partitioning of water 33 34 between the gas and particle phases is critical to describe accurately the size distribution and composition of atmospheric aerosol.<sup>3</sup> Measurements of the response of the equilibrium particle size 35 and composition to changes in relative humidity (RH) can be used to quantify the hygroscopicity of 36 the aerosol, the capacity of a particle to absorb water from the surrounding vapour.<sup>4-7</sup> Further, 37 changes in the partitioning of volatile and semi-volatile solutes (organic and inorganic) between the 38 particle and gas phases can accompany changes in the RH.<sup>8</sup> Measurements of the decreasing radius of 39 a particle evaporating into a gas phase devoid of the volatile or semi-volatile component can allow 40 the vapour pressure of the component to be determined.<sup>9-11</sup> 41

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43 Ensemble measurements have been used extensively to probe aerosol systems in the fine (< 2.5  $\mu$ m 44 diameter) and accumulation (0.1 - 1 µm) mode regimes, providing measurements of quantities 45 averaged over a large number of particles with insufficient resolution to quantify aerosol processes 46 robustly at the fundamental level of individual particles. Single particle techniques can be employed 47 successfully in the measurement of aerosol optical and thermodynamic properties, as well as their kinetic response, and have the potential to resolve the fundamental detail of physicochemical 48 processes.<sup>11</sup> The use of electrodynamic balances (EDBs) to isolate and study single particles is well 49 50 established, with recent improvements in experimental techniques allowing highly accurate measurements of particle hygroscopicity, but has been largely limited to studying coarse mode 51 particles  $> 5 \mu m$  in radius.<sup>9,4</sup> Optical tweezers are routinely used to isolate and study aerosol particles 52 1 - 10 µm in radius, still much larger than the accumulation mode particles that are of direct 53 atmospheric importance.<sup>12,13,2,14</sup> Thus, the size regimes examined by ensemble and single-particle 54

techniques do not overlap; there is a need for new techniques to control and characterise individual fine mode particles in order to reconcile single particle and ensemble measurements.

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The need for single aerosol particle studies in the fine mode is augmented by the fact that some 58 59 significant and atmospherically-relevant aerosol processes occur for diameters well below 1 micron. For example, the Kelvin equation shows that the vapour pressure for semi-volatile organic aerosol 60 (SVOA) increases exponentially as the particle radius decreases (and the surface curvature 61 increases).<sup>15</sup> However, the influence of the Kelvin effect is negligible for the coarse mode particles 62 that are probed in existing single particle techniques. Although measurements of the kinetics of 63 aerosol transformation (for example, heterogeneous oxidative aging or water condensation/ 64 evaporation) can be made routinely on single coarse mode particles, the accuracy of scaling the 65 66 observed kinetics to the much smaller fine mode particles found in the atmosphere is frequently debated.<sup>11</sup> Further, and of particular relevance for the direct influence of aerosol on climate, particle 67 68 diameters in the range 500 nm to 1 µm are of sufficiently high number concentration and optical 69 cross-section to play a significant role in atmospheric optics. Direct measurements of aerosol optical 70 properties in this size range would be particularly advantageous. Single particle techniques that can access the fine mode and push the lower size limit down well into the accumulation mode would 71 72 expand the range of particle sizes accessible to single particle techniques and allow the robust 73 assessment and development of thermodynamic, kinetic and optical models.

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A carefully engineered Bessel laser beam<sup>16,17</sup> may be used as an optical trap for single, as well as 75 multiple, aerosol particles across a range of particle sizes, from 10s of microns to a few hundred 76 nanometres in diameter.<sup>18–20</sup> The intensity profile along the transverse direction (*i.e.* in the direction 77 78 perpendicular to the optical propagation axis) of a zeroth order Bessel beam (BB) is circularly 79 symmetric with a central core of intensity surrounded by multiple rings. It can be shown by Fourier 80 optics that such a beam is realisable by propagating the wave vectors of an incident beam along a cone, which may be achieved using a conical-shaped lens, or axicon.<sup>21-24</sup> The axicon method 81 generates a BB over an extended distance due to the large region over which waves may interfere. If 82

83 an imaging plane is placed at any point along this interference region, a Bessel-squared intensity 84 distribution is observed. Crucially, the core of a BB appears non-diffracting and retains its size and 85 shape over a macroscopic length far larger than the characteristic Rayleigh range of a Gaussian beam with the same focal beam waist as the BB core diameter.<sup>16</sup> If a particle is irradiated by the core of a 86 87 BB, it will experience radiation pressure force in the direction of the propagation axis of the beam, while optical gradient forces act in the transverse directions to retain the particle within the core. To 88 confine single particles along the BB propagation length, the radiation pressure force needs to be 89 balanced. If two counter-propagating BBs are used, a radiation pressure trap is created in which the 90 trapping of both spherical and non-spherical sub-micron particles over macroscopic distances may be 91 achieved.18 92

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94 Previous work has shown that BBs can be used to trap and characterise multiple aerosol particles 500 to 2000 nm in radius, with their sizes resolved from the angular variation of elastically scattered 95 light.<sup>18,19</sup> We have previously shown that the accuracy of such sizing methods can be enhanced by 96 sizing with multiple-wavelength probe beams, which is particularly valuable for determinations of 97 98 size for particles below 1 micron in radius. Furthermore, the size range of particles trapped by the beam can be coarsely tuned by varying the core size and laser power of the BB, and may be 99 100 optimised for trapping fine mode particles. More recently, we have reported the use of a vertically 101 propagating BB optical trap coupled with cavity ring-down spectroscopy to measure directly the optical extinction by single aerosol particles<sup>20</sup> 102

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The purpose of this paper is to demonstrate that BBs can be used to study the physiochemical properties of fine mode aerosols over extended periods of time. A single horizontally propagating BB is coupled with a counter-propagating gas flow to balance the radiation pressure force exerted on confined aerosol particles. This arrangement allows the environment around a single particle to be varied easily without perturbing the trapping stability. Measurements of particle hygroscopicity and gas-particle partitioning of semi-volatile components are reported.

### 111 **II. Experimental Description**

112 The experimental apparatus used is summarised in Figure 1. A 532 nm Gaussian laser beam (Laser Quantum 3W Opus) was expanded using a pair of lenses (L1 and L2) to illuminate the back-aperture 113 114 of a 2° axicon (Altechna), generating a region of interference of approximately one meter in length 115 over which a BB was formed. Immediately after the axicon, the BB core diameter was measured to be 55 µm, which is too large to trap sub-micron aerosols. Hence, a further pair of lenses (L3 and L4) 116 was used to reduce the core diameter by a factor of 12.5 times to  $4.7 \pm 0.3 \mu m$ , propagating over a 117 distance of approximately 5 mm. This beam was directed into the trapping cell. The radiation 118 pressure exerted by this horizontally propagating BB on a trapped particle was balanced by a 119 humidified nitrogen gas flow of 50 sccm, the RH of which could be controlled. This gas flow was 120 121 also used to purge the cell of excess aerosol. The speed of the gas flow combined with the geometry 122 of the trapping cell yields a Reynolds number Re close to 1, much lower than the critical value for Re of approximately 3000 for turbulent flow. The RH inside the trapping cell was monitored using a 123 124 capacitance probe (Honeywell). A camera (C1) coupled to a  $20 \times \log$  working distance objective 125 with a numerical aperture (NA) of 0.42 was used to capture the phase function from a single trapped aerosol,<sup>25,19</sup> while a second camera (C2) coupled to a 10× objective tracked the position of the particle 126 127 along the full propagation length of the BB. C2 also indicated whether multiple particles were 128 trapped. Aerosol was introduced into the cell via a medical nebuliser (Omron). If multiple particles 129 were trapped, the trapping cell was evacuated and fresh aerosol was nebulised into the cell. This 130 process was repeated until only a single particle was trapped.

131

Aerosol radii typically have dimensions on the order of the wavelength of visible light. Semiempirical models cannot describe the angular variation in the light scattering intensity, the phase function (PF), and Mie theory must instead be used. The elastic scattering intensity distribution is affected not only by particle radius but also by the polarisation of the incident light and the refractive index (RI) of the particle, *m*. Here, we denote the real and imaginary components of the RI by *n* and *k* respectively. Measuring the angular variation in the scattered intensity and simulating the measured distribution using Mie theory provides a method of obtaining the particle size even for sub-micron 139 aerosol particles.<sup>26</sup> This simulation is obtained by computing a library of Mie PFs, comparing each 140 one with the measured PF and performing a least-squares fit. The library of PFs can be computed by 141 floating the variables of particle radius and RI. Typically, for a single measured PF there may be 142 multiple fit solutions with different values of radius for different RI values. Thus, in our 143 measurements, the RI is modelled using literature data and only the radius is fitted.

144

In addition to the recorded PFs at 532 nm from elastically scattered light from the BB, two further 145 probe beams of different wavelength were employed. The first was from a 35 mW 405 nm diode laser 146 (IO Series, Power Technology Incorporated) and the second was from a 2 mW 633 nm HeNe laser 147 (LHP073, Melles Griot). These additional probe beams were combined using a polarisation beam 148 splitter (PBS) and aligned anti-parallel to the BB direction. The probe beams were gently focussed 149 150 into the trapping plane using a 400 mm lens (L5) in order to obtain sufficient scattered light for detection by C1, although their irradiances in the trapping plane were low enough not to affect the 151 particles position and considerably less than that from the trapping 532 nm laser. When collecting 152 153 PFs from the additional probe beams, various band pass filters were placed between the objective and 154 CCD in C1 to allow collection of light from the respective probe beams and block the more intense 155 532 nm scattered light from the BB.

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### 157 III. Measurements of Component Vapour Pressures

Volatilities (vapour pressures) of organic compounds are key thermodynamic properties required for 158 quantifying semi-volatile organic aerosol (SVOA).<sup>27</sup> Vapour pressures can be estimated from single 159 160 particle measurements by recording the change in radius over time. Glycerol is a good benchmark system to estimate the validity of the BB approach for studying aerosol volatilities; the relatively low 161 hygroscopicity means glycerol absorbs little water, so may be assumed to have a well-defined 162 composition. Numerous glycerol droplets were studied using the BB apparatus, collecting the time-163 dependent PFs as a droplet was allowed to evaporate. Note that for these experiments, PFs were only 164 165 collected at 532 nm.

167 Dry nitrogen was flowed into the trapping cell at a flow rate of 50 sccm until the in-cell RH probe 168 reported a stable, low RH, typically between 5 - 10 %. When an aqueous glycerol solution was nebulised into the cell, the in-cell RH increased rapidly to 25 - 30 % before returning to a stable 5 -169 170 10 % after 6 - 8 minutes. Because of the small size of the droplets (less than 2 µm) and their fast 171 evaporation (under 15 minutes), leaving the measured RH to equilibrate and stabilise at low RH before data acquisition was not possible. However, the low hygroscopicity of glycerol (particularly 172 when compared with inorganic salt systems) suggests the non-zero and varying RH will have a 173 negligible effect on the particle size and evaporation profile. When fitting PFs to Mie theory 174 simulations, the RI was assumed to be constant and equal to the pure component value of m = 1.475 +175  $i 10^{-8}$  over the entire experiment.<sup>28</sup> Maxwell-Garnett and volume-weighting mixing rules show that 176 the difference in the real component of the RI at 0 % and 20 % RH is 0.012, a 0.8 % deviation. This 177 178 corresponds to a 0.8 % difference in the Mie theory fit radius, which is of a similar magnitude to the uncertainties in PF measurements of radius. Other sources of uncertainty include noise in the PF 179 180 measurements, not using a precise angular range, and from not using plane wave illumination of the 181 particle, which is a fundamental assumption of Mie scattering theory.

182

Figure 2(a) shows the radius changes observed for four glycerol particles as they evaporate over time. One measurement is for a coarse mode particle of ~2.9  $\mu$ m in radius, evaporating to a radius of ~1.7 µm. Further data sets show that fine mode glycerol particles may be trapped and characterised, with one particle evaporating from a radius of ~800 nm to 500 nm. Droplet vapour pressures, *p*, may be obtained by fitting the Maxwell equation (1) to the time-dependent data.<sup>11</sup>

188 
$$\frac{da^2}{dt} = \frac{2D_{ij}M_i}{R\rho} \left(\frac{p_{\infty}}{T_{\infty}} - \frac{p}{T}\right)$$
(1)

189  $D_{ij}$  is the gas diffusion coefficient of species *i* in the surrounding phase *j*,  $M_i$  is the molecular weight, 190 *a* is the particle radius, *R* is the molar gas coefficient,  $\rho$  is the density, *p* and  $p_{\infty}$  are the vapour 191 pressures at the droplet surface and at infinite distance from the droplet respectively, and *T* and  $T_{\infty}$  are 192 the corresponding temperatures, assumed to be the same in this case for slow evaporation of low volatility components. By taking the vapour pressure of component *i* at infinite distance to be negligible and including the effect of non-zero RH, equation (2) can be derived where  $m_i$  is the mass fraction of solute *i*.

196 
$$\frac{da^2}{dt} = -\frac{2D_{ij}M_ip}{R\rho Tm_i}$$
(2)

197 Because of non-ideality in the interactions between different species, the pure component vapour pressure,  $p_0$ , is related to the measured vapour pressure by  $p = p_0 \gamma_i x_i$  where  $\gamma_i$  is the activity 198 coefficient and  $x_i$  is the mole fraction of species *i* in the droplet. Hence, by plotting the radius-199 squared as a function of time, a straight-line fit should be achieved where the gradient is related to  $p_0$ . 200 Figure 2(b) summarises the result of measuring 11 pure component vapour pressures in this way, 201 202 having assumed that both the activity coefficient and the mole fraction of glycerol are equal to unity. 203 The average glycerol vapour pressure is  $7.5 \pm 2.6$  mPa, agreeing satisfactorily with the literature value of 11.6 mPa.<sup>29,28</sup> This level of agreement is quite reasonable when the more significant 204 205 variations in vapour pressure determinations for components with similar vapour pressures such as dicarboxylic acids are considered.<sup>11</sup> It will now be shown that the wide variation in the measured 206 207 vapour pressures for glycerol droplets is primarily due to the large variation in the in-cell RH at the 208 beginning of droplet evaporation, and the relatively high volatility of glycerol.

209

210 One difficulty in calculating a precise vapour pressure for glycerol is in determining an accurate 211 measure of the RH surrounding the glycerol particles. Nebulising aerosol into the trapping cell causes a large increase in the in-cell RH. Ideally, following nebulisation and particle trapping, the cell would 212 be left for ~20 minutes, passing dry  $N_2$  gas into the cell until the RH stabilises at low (5 – 10 %) RH. 213 1,2,6-Hexanetriol droplets evaporate more slowly than glycerol, owing to the much lower vapour 214 pressure of the triol and so more accurate measurements of vapour pressure might be expected. 215 216 Hence, multiple hexanetriol droplets were studied using the BB apparatus. The measurements were performed at an average RH in the range 5 - 10 %, with an RH profile that was more stable than for 217 the previous glycerol studies. Unlike the glycerol experiments, PFs were instead collected at 405 nm 218

219 while still employing the 532 nm BB for optical trapping, allowing us to investigate simultaneously 220 the accuracy with which an unknown RI can be determined. Here, we show that the quasi-equilibrium variation in position of hexanetriol droplets along the length of the BB ( $\lambda = 532$  nm), observed 221 222 experimentally using camera C2 (Figure 1), can be modelled by Mie theory to obtain both the size 223 evolution of the particle and the RI at 532 nm. Once the size evolution of the particle is determined, 224 the PFs ( $\lambda = 405$  nm) collected with camera C1 are then fitted, while varying the RI at 405 nm so as to give the same size evolution as given by modelling of the positional data. Thus, these 225 226 measurements yield simultaneously particle size and refractive index at two wavelengths.

227

In aerosol optical tweezers, where particles are typically  $2 - 10 \,\mu\text{m}$  in radius, Raman spectroscopy is 228 used to determine simultaneously radius and refractive index with high accuracy.<sup>25</sup> Whispering 229 230 gallery modes (WGMs) superimposed on spontaneous Raman bands arise because certain frequencies 231 couple efficiently with the resonant modes of droplet microcavities. However, this method of size 232 determination cannot be employed when probing single particles in the fine mode as there are too few 233 WGMs within the Raman bandwidth for reliable size analysis. A Raman spectrum is a form of 234 resonance spectrum, which traditionally describes scattering intensity as a function of the size 235 parameter, the ratio of the circumference of a spherical particle to the incident wavelength. Measuring 236 scattering intensity variation with size at a fixed angle is an alternative method of obtaining a 237 resonance spectrum. For example, Ray and co-workers measured resonances in the scattering intensity for an evaporating glycerol droplet.<sup>30</sup> The position of a confined particle along the length of 238 239 the core of a BB when subjected to a counter-propagating gas flow is highly dependent on the 240 radiation pressure efficiency of a particle and, as we will show, follows the form of a typical resonance spectrum. Thus, measuring the relative position of an evaporating particle along the length 241 of a BB provides an alternative method of deriving the size evolution of a particle and can be 242 243 exploited even for sub-micron aerosol particles.

244

The position of a particle along a single Bessel beam with a counter-propagating gas flow depends on two forces: (i) the force exerted by the laser beam through the action of radiation pressure  $(F_{pr})$ ; (ii) 247 the Stokes drag force exerted by the gas flow. While the Stokes drag force is simply dependent on the gas flow velocity and particle radius,  $F_{pr}$  is more complicated and is proportional to the product of the 248 on-axis core intensity and the radiation pressure efficiency,  $Q_{pr}$ , of the particle. The latter term is a 249 measure of how efficiently and asymmetrically a particle scatters light; a particle back-scattering a 250 higher fraction of light has a higher  $Q_{pr}$  value.<sup>31</sup> Figure 3(a) shows the variation in position of a 251 hexanetriol particle along the length of the single BB optical trap as a function of evaporation time. 252 The sharp resonances in positional data directly correspond to the resonance structures in  $Q_{pr}$  and 253 occur at unique sizes. Hence, only Mie theory simulations of  $Q_{pr}$  are required to derive the size 254 255 evolution of a particle and the effects of Stokes drag and variations in the BB core intensity can be 256 neglected.

257

258 The procedure for comparing positional data with Mie simulations to retrieve droplet size is similar to the fitting of Mie resonance spectra in previous single particle measurements by Chylek et al. 32-34 and 259 Ray et al.<sup>30,35</sup>, and is now briefly described.  $Q_{pr}$  is simulated as a function of particle radius and 260 repeated for fixed single values of RI over a range of RIs. Sharp resonance peaks in the  $Q_{pr}$ 261 simulations (in the radius domain) can be assigned by matching to peaks in the positional data (which 262 263 are obtained in the time domain), while also ensuring that the underlying contour reflecting higher order Mie terms is consistent with the data. Plotting these assigned radii against their respective times 264 of observation gives the size evolution of the particle, connecting time, the particle radius, and its 265 position along the length of the BB. 266

267

Figure 3(b) shows the correspondence between the resonant sizes in  $Q_{pr}$  and the resonant excursions in the particle position along the length of the BB. Differences in the signal amplitude between the positional data and the  $Q_{pr}$  simulations are likely because the effects of Stokes drag and variations in the core intensity with propagation have been neglected. It can be seen in Figure 3(b) that the positions of resonance peaks in the position data do not agree exactly with the positions of resonance peaks in the  $Q_{pr}$  simulation. The residual differences between the positions of resonance peaks in the 274 Mie simulation  $(a_{i,sim})$  and the experimental measurement  $(a_{i,exp})$  provide a measure of the 275 agreement between the simulation (for a given RI) and the position measurements. The total residual 276 of a fit,  $D_{fit}$ , is given by equation (3).

$$D_{fit} = \sum_{i} \left| a_{i,\exp} - a_{i,sim} \right| \tag{3}$$

278 Repeating this procedure and obtaining the fit residuals for simulations using different RIs allows a 279 determination of the RI at 532 nm. This outcome is demonstrated in Figure 3(c) for three studies, 280 which indicate an RI of  $1.482 \pm 0.001$  at 532 nm. Refractometer measurements give a bulk refractive 281 index of 1.4758 at a wavelength of 589.6 nm. A higher refractive index at a lower wavelength is 282 expected, given typical trends in dispersion for other materials. No literature data at the wavelength 283 of 532 nm can be found.

284

Figure 4(a) shows how the radii evolve with time for five droplets, with droplet sizes determined from 285 their positional data as well as from their 405 nm PF data. When performing the PF analysis, the RI at 286 405 nm was adjusted until only the initial radius (i.e. the radius at t = 0 s) matched that given from 287 288 modelling the positional data. Subsequent trends during evaporation illustrate the consistency that is 289 then obtained between the two size-determination methods at the two wavelengths. Assuming the imaginary component of the RI, k, to be equal to  $10^{-8}$ , the real component, n, at 405 nm was 290 consistently determined to be  $1.485 \pm 0.006$  from measurements on multiple droplets. Again, no 291 292 literature value of RI can be found specifically at the wavelength of 405 nm which to compare our measured value. The three real RIs now obtained for hexanetriol (1.485 at 405 nm, 1.482 at 532 nm, 293 1.476 at 589.6 nm) can be fitted to refractive index models such as Cauchy's equation. RI's at further 294 wavelengths are required to quantitatively assess the accuracy of such models in describing the 295 dispersion in the RI of hexanetriol. The possibility of retrieving a more detailed wavelength 296 297 dependence of RI will feature in future work.

298

The five hexanetriol droplets remained optically trapped for much longer than glycerol droplets, with one droplet isolated for 10 hours. As experiments were performed at a low RH of ~5 %, the mass fraction of water in the droplets is assumed to be negligible and the activity coefficient of hexanetriol is taken to be unity. Figure 4(b) shows calculated vapour pressures for five droplets with an average value of  $0.20 \pm 0.02$  mPa. The 10% variation in the measured hexanetriol  $p_0$  compared to the 35% variation for glycerol demonstrates that greater reproducibility can be achieved in vapour pressure measurements when the in-cell RH can be well defined. No comment can be made on the accuracy of our measurement of  $p_0$  for hexanetriol as there are no current literature values for this compound with which to compare.

#### 309 IV. Hygroscopic Growth Measurements

310 Hygroscopicity studies require measurements of the changes in the equilibrium partitioning of water 311 between the gas and particle phases with variation in RH, most often through measurements of 312 relative size.<sup>11</sup> Due to non-ideal interaction of water with solutes, thermodynamic models are required to describe the activity coefficients of solutions, particularly for the supersaturated droplets probed in 313 314 aerosol systems. Here, we compare hygroscopic growth measurements of involatile inorganic salt 315 systems to simulations of growth from the Extended Aerosol Inorganic Model (E-AIM), an 316 equilibrium state model which can describe the relationship between particle growth and changes in the RH.<sup>36</sup> Vapour pressure and hygroscopicity measurements aid in the parameterisation and 317 318 validation of such thermodynamic models.<sup>11</sup>

319

320 As an aerosol particle takes up water with an increase in RH, both the RI and size of the particle change. Fitting the PFs obtained from binary component droplets (water and solute) in hygroscopicity 321 experiments is more complicated than fitting those from vapour pressure experiments and it is not 322 possible to retrieve both the RI and size from a single PF, with several solution combinations of RI 323 and particle radius possible. Instead, the RI has to be described as a function of particle radius, a. For 324 a binary system with volume fractions  $\varphi_1$  and  $\varphi_2$  for pure water and pure solute, respectively, the 325 linear volume-fraction average mixing rule is often encountered in studies of aerosols. However, this 326 treatment of refractive index is not applicable in many cases, as shown by Liu and Daum.<sup>37</sup> Thus, we 327

<sup>308</sup> 

have opted to use the parameterisations of Tang and Munkelwitz<sup>38</sup> for various inorganic aerosol systems, which is now briefly described. The real component of the RI, *n*, for a binary aerosol can be modelled in terms of particle radius, *a*, using equation (4).

331 
$$n = n_w + gc_2 + hc_2^2 = n_w + g\left[\frac{a_0^3 c_{2,0}}{a^3}\right] + h\left[\frac{a_0^3 c_{2,0}}{a^3}\right]^2$$
(4)

 $n_w$  is the real component of RI for pure water,  $c_2$  is the solute concentration, and  $a_0$  is the particle 332 radius at some reference concentration of solute  $c_{2,0}$ . Tang and Munkelwitz<sup>38</sup> reported the RIs at 633 333 nm of several aqueous inorganic aerosols and their dependence on the solute concentration,  $c_2$ . Using 334 their parameterisations, the constants g and h in equation (4) may be calculated for a particular 335 336 inorganic species. Figure 5 shows the variation of n with  $c_2$  for both a sodium chloride and an ammonium sulphate aerosol, as determined by Tang and Munkelwitz. A polynomial fit to their 337 338 parameterisation gives the constants g and h. A small wavelength correction is also applied to the RI to account for dispersion; we use the correction given by Millard and Seaver.<sup>39</sup> The parameters 339  $c_{2,0}$  and  $a_0$  remain to be determined. The E-AIM model can relate RH to solute concentration, so 340 341 using the measurement of the in-cell RH (which has a  $\pm 2\%$  accuracy) E-AIM is used to determine the solute concentration  $c_{2,0}$  at a reference RH (usually taken to be the RH at t = 0 s). Subsequently, 342 equation (4) is used to determine the RI of the initial droplet at  $c_{2,0}$ . The PF recorded at this reference 343 344 RH is then fitted using this value of RI to determine  $a_0$ . A library of phase functions can now be 345 generated in the same way as in previous PF analyses, except that the RI is varied with radius in the form given by equation (5), as opposed to having a constant value. 346

347

Aqueous sodium chloride droplets were nebulised into the optical trapping cell and the initial RH was kept high (between 85 - 95 %). The laser power was set at powers in the range 1 - 2 W, corresponding to core intensities in the range 0.3 - 0.6 mW  $\mu$ m<sup>-2</sup>. After an equilibration time of ~20 minutes, the RH was lowered to 30-40% over 1 hour, during which time the particle axial positions along the propagation distance and the PFs at 532 nm were recorded. The changes in RH are 353 sufficiently slow that the droplet is always at equilibrium during measurement and, thus, water transport kinetics can be neglected. An example data set is shown in Figure 6(a). As expected, the 354 particle radius decreases as the in-cell RH is reduced and effloresces in the region 37 - 41 % RH (for 355 comparison, literature values range from 39 to 54% RH).<sup>11</sup> Unlike in optical tweezers experiments, 356 357 the effloresced particle remains stably trapped. Predictions of the radial growth factor (RGF) calculated from the E-AIM model can be compared to the pre-efflorescence data. The RGF is the ratio 358 of the particle radius to its radius at 0% RH (dry radius). E-AIM calculations of RGF were fitted to 359 the experimental data using a least-squares method and floating the dry radius parameter. The 360 distribution of normalised least-squares residuals with dry radius can be fitted to a Gaussian 361 distribution, from which an uncertainty in the dry radius can be obtained. This uncertainty is shown in 362 363 Figure 6(a) by the shaded grey area. The experimental pre-efflorescence data show good agreement 364 with E-AIM, and the dry radius obtained by the fit is indicated by the dashed red line.

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366 Interestingly, the PFs from the supposedly crystalline NaCl particle closely resemble those expected 367 from a sphere, with distinct dark and bright fringes, although they were observed to 'shimmer' at 368 times with bright fringes appearing broken in a regular pattern. The fact that a PF resembling a 369 spherical particle is obtained may suggest that the effloresced NaCl particle has not adopted the 370 expected equilibrium cubic structure immediately. Modelling post-efflorescence PFs with Mie theory gives dry particle radii very close to the values expected from the E-AIM fit to the pre-efflorescence 371 372 data. This is shown by the experimental data points in Figure 6(a) distributed just below the dashed red line, implying the PFs obtained are indeed for an effloresced NaCl particle but for a period during 373 which the effloresced particle remains approximately spherical. Although effloresced NaCl particles 374 are generally assumed to have a cubic shape, nearly spherical NaCl particles have been observed by 375 376 others in the literature.<sup>40</sup>

377

Further information is given by the particle position data along the propagation length of the BB (Figure 6(b)). The pattern of periodic resonances expected for a particle that is changing size ends after efflorescence, as marked in the figure by the label 'crystallisation'. However, position 381 fluctuations continue following efflorescence, but are seemingly random in nature. The reason for 382 these fluctuations is unclear. One possibility is that the effloresced particle is not completely crystalline and initially retains at least some of the solution phase. This structure may be evolving 383 over time, tending towards the equilibrium cubic structure of NaCl as further residual water 384 385 evaporates from the particle. Such structural changes will impact on the optical properties and radiation pressure efficiency of a particle, affecting its position along the length of the BB. 386 Instantaneous fluctuations in the orientation of the particle in the optical trap as it tumbles, if non-387 spherical, can also lead to large and rapid deviations in particle position as the magnitude of the 388 389 radiation pressure changes.

390

391 So far, the particles studied are far larger than the fine mode particles we wish to probe using Bessel 392 beams. Carruthers and co-workers showed that the size range of particles isolated using BB optical traps can be preferentially selected by changing the BB core diameter.<sup>19</sup> Thus, reducing the core 393 diameter should favour the trapping of smaller particles. The post-axicon BB core diameter was 55.5 394 395  $\pm$  2.9  $\mu$ m and may be tuned in the optical trapping region by tweaking the focal length ratio of lenses L3 to L4 in Figure 1. The lens ratio L3/L4 was increased to 20 from the value of 12.5 used so far, 396 397 resulting in the measured core diameter decreasing from  $4.7 \pm 0.3 \ \mu m$  to  $2.6 \pm 0.1 \ \mu m$ . Also, the solute concentration in the nebuliser can be reduced so as to produce a size distribution of nebulised aerosol 398 peaking at smaller size. Subsequent hygroscopicity studies performed on aqueous NaCl droplets 399 400 clearly showed the trapping and reliable sizing of NaCl droplets down to radii of 600 nm, with the 401 hygroscopic growth curves again broadly consistent with predictions from E-AIM. Smaller particles 402 were optically trapped, but their PFs recorded at 532 nm were featureless so the particles could not be 403 sized.

404

In an attempt to size aerosol truly within the fine mode reliably, use of a lower (visible) probe
wavelength is logical. To this end, the 35 mW 405 nm probe beam (alignment as shown in Figure 1)
was used to collect PFs. Figure 7 shows an example of a hygroscopicity measurement for an aqueous

408 NaCl droplet, using the 405 nm probe beam to determine the radius for a droplet which could not be 409 reliably sized using the 532 nm beam. The figure also shows the corresponding PFs at 405 nm and 532 nm for the droplet at t = 0 s (corresponding to 80 % RH). The 405 nm PF clearly has features 410 which can be used to fit to Mie theory simulations, whereas the 532 nm PFs are completely 411 412 featureless (neglecting the side peaks from optical aberrations). The radius of the droplet was determined to be 510 nm from the PF at the probe wavelength of 405 nm. Droplets have been reliably 413 sized down to radii of ~350 nm, well into the fine mode regime relevant to atmospheric chemistry and 414 physics. These results represent the smallest single-particle measurements of hygroscopicity made. 415 Indeed, particles around this size present the largest total optical cross-section to visible light in the 416 417 atmosphere, with large geometric cross-section, extinction and scattering efficiency and sufficient 418 number density.

419

420 Ammonium sulfate is another atmospherically-relevant aerosol system and hygroscopicity 421 measurements have been performed using the BB apparatus, measuring PFs using the 405 nm probe beam. A summary plot, averaging the RGF from hygroscopic growth measurements of seven 422 accumulation mode droplets (<500 nm in radius), is presented in Figure 8(a). Errors are calculated by 423 424 taking the standard deviation in the RGF values, and assuming  $a \pm 2$  % standard error in the measured RH. Figure 8(b) also shows a summary of seven studies on sodium chloride accumulation mode 425 particles. Evidently, for both ammonium sulfate and sodium chloride, the hygroscopic growth of 426 427 accumulation mode particles is well described by E-AIM.

428

### 429 V. Conclusions

The capability of BB optical traps to measure the vapour pressure and hygroscopicities of fine and accumulation mode aerosol has been successfully demonstrated. Vapour pressures of organic aerosol have been measured, with a demonstrated reproducibility within 16 % precision if the optical trapping cell RH can be well defined and maintained over the entire course of the experiment. A first attempt has been presented to model the position of particles subjected to radiation pressure from the BB and 435 a counter-propagating gas flow. Analysis of the sharp resonances in particle position data determines both the size evolution of hexanetriol particles and the RI at 532 nm. The accuracy of the RI value 436 obtained is limited by the resolution of the position of the particle along the length of the beam. This 437 uncertainty can be reduced by using a longer BB core length and by reducing the exposure on the 438 439 camera recording the particle position giving higher time resolution of particle position. The size evolution of the particle from position data agrees with that determined by analysis of PFs recorded 440 441 using a 405 nm probe beam. This agreement is only achieved when using a fixed real component of 442 the RI at 405 nm of  $1.485 \pm 0.006$ . Future work will involve repeating measurements for further probe beam wavelengths (e.g. 633 nm) to see how well dispersion models describe the wavelength 443 444 dependence of the RI of hexanetriol.

445

Hygroscopicity measurements on sodium chloride and ammonium sulfate aerosols have been 446 447 performed for particles as small as 350 nm in radius. These results represent the first single-particle 448 measurements of hygroscopicity on fine mode and accumulation mode aerosols, the size regime 449 which has the most atmospheric significance in terms of loading and light extinction/scattering 450 efficiencies. The variation in radial growth factor with RH agrees well with the results from the E-451 AIM thermodynamic models. There are two principal factors that raise questions over the reliability of our measurements. Firstly, the RH within the cell is poorly defined with a precision of  $\pm 2$  % at 452 best. It is further assumed that the in-cell RH measured by a probe near to the wall of the cell is the 453 454 same as that experienced by an optically trapped particle at the centre of the trapping cell. We have 455 attempted to ensure that this is true by positioning the RH probe as close to the trapping region as 456 possible given the physical constraints of the trapping cell. If the probe is saturated there could be a 457 delay in equilibration of the probe with the environment of the rest of the cell. Secondly, the fitting of PFs relies on having an accurate description of the RI and assumes a dependence on radius. Although 458 the parameterisations of Tang and Munkelwitz<sup>38</sup> are commonly used to describe RI as a function of 459 concentration, we have to use our measurements of RH to obtain the starting droplet concentration, 460 which suffers from an uncertainty of  $\pm 2$  %. Further, Tang and Munkelwitz's parameterisations are 461

valid at 633 nm only. So, when describing the RI of sodium chloride or ammonium sulfate for PFs
recorded from a 405 nm probe beam, a dispersion model formulated for sea salt aerosol is applied.<sup>39</sup>
However, such dispersion models may not be valid over such large wavelength ranges.

465

466 EDB techniques are now at a point at which the RH of the trapping cell can be defined to better than 0.2 %<sup>4</sup> However, EDB methods also employ PF analysis to extract size information so they still rely 467 on modelling the RI. As well as being able to extract aerosol radius to a precision better than 1 nm, 468 optical tweezers methods employing Raman spectroscopy can simultaneously measure the RI with an 469 uncertainty better than 0.1 %, and can quantify the dispersion over a narrow wavelength range.<sup>12</sup> 470 However, RH measurements are performed by capacitance probe methods so suffer the same level of 471 uncertainty as our measurements. Thus, to exploit the high level of accuracy in quantifying droplet 472 473 size and RI in such measurements, comparative measurements on multiple droplets must be performed.<sup>41</sup> Indeed, we have shown here that the uncertainty with which hygroscopic growth 474 475 measurements can be made on fine mode inorganic aerosol using the BB optical trapping approach ( $\pm$ 476 3% error in RGF for a given RH) is comparable to the typical uncertainty associated with ensemble accumulation mode measurements made using the hygroscopic tandem differential mobility analysis 477 technique, typically  $\pm$  3% for inorganic aerosol as seen in Figure 1 of reference <sup>42</sup> with errors in RH 478 479 that are also similar.

480

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Figure 1 – Schematic of the apparatus used to generate a Bessel beam and optically trap single aerosol particles. C1
 and C2 represent 'camera 1' and 'camera 2' respectively, PBS is a 'polarisation beam splitter', and L1-L5 represent



- 563 Figure 2 (a) Four data sets measuring the radius as a function of time for optically-trapped glycerol droplets as they
- 564 evaporate at an RH of ~15-20 %. (b) Vapour pressures from eleven studies give an average p<sub>0</sub> of 7.5±2.6 mPa. The
- 565 crosses indicate individual measurements and the circle and error bar indicate the mean and standard deviation.



- 568 Figure 3 (a) Measured position of the hexanetriol droplet in Study 1 along the length of a Bessel beam as it
- evaporates over time. (b) Measured position data from Study 1 (black) in the radius domain, accompanied by the best
- 570 fit simulation of Q<sub>pr</sub> (red) which used a real component of the RI of 1.483. (c) Normalised residual in Q<sub>pr</sub> fits as a
- 571 function of RI at 532 nm for three different studies. Lines between data points are to guide the eye only.



- 574 Figure 4 (a) Five hexanetriol studies measuring the radius as a function of time using position data (red) and
- 575 subsequent measurement from phase function data (black). (b) Vapour pressures from five hexanetriol studies give
- 576 an average  $p_0$  of  $0.20 \pm 0.02$  mPa.



579 Figure 5 – The calculated variation in the real component of refractive index, n, with solute concentration,  $c_2$ , using

the parameterisations of Tang and Munkelwitz.<sup>38</sup> Black dots show data points, while the red line is a second-order
polynomial fit to data points, the equation for which is given.



Figure 6 – (a) An example NaCl hygroscopicity data set using a 4.7  $\mu$ m core diameter Bessel beam. Experimental data points shown by black dots, accompanied by a comparison to the E-AIM model (red line) which has an associated uncertainty (shaded grey area). Dashed red line indicates the dry radius estimated from a fit of E-AIM model to experimental data points. Typical uncertainty in relative humidity is ± 2%. (b) The particle position along the Bessel beam length for the data set in (a).



Figure 7 – (a) Example hygroscopicity study on an aqueous NaCl aerosol droplet in the accumulation mode regime,
with the radius determined from 405 nm phase functions. (b) Phase functions at probe wavelengths 405 nm (blue line)
and 532 nm (green line) for the initial droplet in (a) at high relative humidity. Inset are the raw phase function images
at 405 nm and 532 nm.



- 598 Figure 8 Summary of hygroscopic growth measurements on: (a) ammonium sulfate; (b) sodium chloride. In each
- 599 case, the radial growth factor is an average from seven hygroscopic growth measurements on accumulation mode
- 600 (<500 nm radius) aerosol particles.

