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Anaerobic Oxidation of Short-Chain Alkanes in Hydrothermal Sediments: Potential Influences on Sulfur Cycling and Microbial Diversity

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1 **Anaerobic oxidation of short-chain alkanes in hydrothermal sediments:**
2 **potential influences on sulfur cycling and microbial diversity**

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39

40 Short-chain alkanes play a substantial role in carbon and sulfur cycling at hydrocarbon-
41 rich environments globally, yet few studies have examined the metabolism of ethane
42 (C₂), propane (C₃), and butane (C₄) in anoxic sediments in contrast to methane (C₁). In
43 hydrothermal vent systems, short-chain alkanes are formed over relatively short
44 geological time scales via thermogenic processes and often exist at high concentrations.
45 The sediment-covered hydrothermal vent systems at Middle Valley (MV, Juan de Fuca
46 Ridge) are an ideal site for investigating the anaerobic oxidation of C₁-C₄ alkanes, given
47 the elevated temperatures and dissolved hydrocarbon species characteristic of these
48 metalliferous sediments. We examined whether MV microbial communities oxidized C₁-
49 C₄ alkanes under mesophilic to thermophilic sulfate-reducing conditions. Here we present
50 data from discrete temperature (25, 55, and 75 °C) anaerobic batch reactor incubations of
51 MV sediments supplemented with individual alkanes. Co-registered alkane consumption
52 and sulfate reduction (SR) measurements provide clear evidence for C₁-C₄ alkane
53 oxidation linked to SR over time and across temperatures. In these anaerobic batch
54 reactor sediments, 16S ribosomal RNA pyrosequencing revealed that
55 *Deltaproteobacteria*, particularly a novel sulfate-reducing lineage, were the likely
56 phylotypes mediating the oxidation of C₂-C₄ alkanes. Maximum C₁-C₄ alkane oxidation
57 rates occurred at 55 °C, which reflects the mid-core sediment temperature profile and
58 corroborates previous studies of rate maxima for the anaerobic oxidation of methane
59 (AOM). Of the alkanes investigated, C₃ was oxidized at the highest rate over time, then
60 C₄, C₂, and C₁, respectively. The implications of these results are discussed with respect
61 to the potential competition between the anaerobic oxidation of C₂-C₄ alkanes with AOM
62 for available oxidants and the influence on the fate of C₁ derived from these hydrothermal
63 systems.

64 **Keywords: hydrothermal vent, metalliferous sediments, Juan de Fuca Ridge, short-**
65 **chain alkanes, sulfate reduction**

67 **Introduction**

68 Hydrocarbon gases, including methane (C₁), ethane (C₂), propane (C₃), and *n*-butane (C₄),
69 are produced via thermogenic and biogenic processes in the deep subsurface and are
70 substantial components of the organic carbon pool across marine and terrestrial
71 ecosystems (Joye et al., 2004; Milkov, 2005; Hinrichs et al., 2006; Cruse and Seewald,
72 2006; Savage et al., 2011). Over the past decade, studies focused on the anaerobic
73 oxidation of methane (AOM) revealed the functional potential, ecological physiology,
74 and diversity of microorganisms mediating this process and the global distribution of
75 AOM as an effective benthic filter that reduces methane emissions into the oceans and
76 atmosphere (for reviews, see Conrad 2009; Knittel and Boetius, 2009; Valentine, 2011).
77 In contrast, the anaerobic oxidation of long chain alkanes (> C₆) and aromatics has also
78 been studied extensively resulting in the isolation of several bacteria, such as sulfate-
79 reducing bacteria (SRB) that oxidize crude oil anaerobically (Van Hamme et al., 2003).
80 There is a gap in our understanding of the metabolism and fate of non-methane, short-
81 chain (C₂-C₄) alkanes in deep sea sediments. Furthermore, there is growing interest in
82 determining the extent to which microorganisms mediate the anaerobic oxidation of C₂-
83 C₄ alkanes, as many studies have indicated that the degradation of these aliphatic
84 hydrocarbons may be linked to global biogeochemical cycles (Lorenson et al., 2002;
85 Formolo et al., 2004; Sassen et al., 2004; Milkov, 2005; Bowles et al., 2011; Quistad and

86 Valentine, 2011).
87
88 Recently, SRB from hydrocarbon seep sediments of the Gulf of Mexico and Guaymas
89 Basin – both of which are environments rich in short-chain alkanes – were documented to
90 oxidize short-chain alkanes to CO₂ anaerobically (Kniemeyer et al., 2007). Different
91 temperature regimens (12, 28, and 60 °C) along with multiple substrates were tested and
92 a pure culture (deemed BuS5) was isolated from mesophilic enrichments with C₃ or C₄ as
93 the sole exogenous carbon source. Through comparative sequence analysis, strain BuS5
94 was determined to cluster with the metabolically diverse *Desulfosarcina* / *Desulfococcus*
95 (DSS) cluster, which also contains the SRB found in consortia with anaerobic
96 methanotrophs (ANME) in seep sediments. Enrichments from a terrestrial, low
97 temperature sulfidic hydrocarbon seep corroborated the biodegradation mechanism of
98 complete C₃ oxidation to CO₂ with most bacterial phylotypes surveyed belonging to the
99 *Deltaproteobacteria*, particularly within the family *Desulfobacteraceae* (Savage et al.,
100 2011). Cold adapted C₃ and C₄, sulfate-reducing cultures have also been obtained from
101 Gulf of Mexico and Hydrate Ridge sediments with maximum rates of SR between 16 –
102 20 °C and dominant phylotypes allied to the DSS cluster including BuS5 (Jaekel et al.,
103 2012). In the study by Kniemeyer et al., C₄ alkane degradation linked to sulfate reduction
104 (SR) was not quantified at thermophilic temperatures, but a Guaymas Basin sediment
105 enrichment with C₃ at 60 °C was dominated by Gram positive bacteria most closely allied
106 to the *Desulfotomaculum*. Moreover, there was no evidence for C₂ degradation in
107 mesophilic (28 °C) or thermophilic (60 °C) enrichments or C₂-linked SR (albeit, there
108 was very slow C₂-dependent SR in Gulf of Mexico enrichments at 12 °C after > 200
109 days).

110
111 The Middle Valley (MV) hydrothermal vent field - located on the northern Juan de Fuca
112 Ridge - is an ideal environment for investigating mesophilic to thermophilic anaerobic
113 oxidation of C₂-C₄ alkanes, given the elevated temperatures and dissolved hydrocarbon
114 species characteristic of these sediments (Goodfellow and Blaise, 1988; Davis, 1992;
115 Cruse and Seewald, 2006). Deep sea hydrothermal vents are complex and dynamic
116 habitats characterized by steep thermal and chemical gradients, a diverse array of carbon
117 and energy sources, and high concentrations of dissolved volatiles (Butterfield et al.,
118 1990; Butterfield et al., 1994; Von Damm, 1995). In the MV system, hydrothermal vent
119 fluids interact with overlying sediments and the thermal alteration of sedimentary organic
120 matter results in the production and/or release of a number of carbon sources, including
121 short-chain alkanes (Cruse and Seewald, 2006; Cruse et al., 2008; Cruse and Seewald,
122 2010). These hydrothermally influenced sediments also contain high concentrations of
123 reduced compounds, such as H₂ and H₂S (Ames et al., 1993; Rushdl and Simonelt, 2002),
124 and metals and metal sulfides at various reduced and oxidized states (Goodfellow and
125 Blaise, 1988; Ames et al., 1993; Wankel, Adams et al., 2012). In contrast to the
126 extremely organic-rich sediments of other sedimented hydrothermal systems, e.g. the
127 Guaymas Basin hydrothermal vent fields in the Gulf of California (% OC = 2 - 4),
128 Middle Valley represents a system that is more typical of mid-ocean ridge hydrothermal
129 vents worldwide (% OC = 0.3 - 0.5) (Wankel, Adams et al., 2012). Such environments
130 could support the coupling of C₁-C₄ alkane degradation to SR in addition to alternative
131 electron acceptors, such as metal oxides, particularly when the organic carbon load and

132 associated SR rates are low (Wankel, Adams et al., 2012).

133

134 We studied the anaerobic oxidation of C₁-C₄ alkanes in metalliferous, organic-poor MV
135 hydrothermal sediments across environmentally relevant temperature gradients. This
136 biogeochemical investigation aimed to determine: (i) the temperature range over which
137 hydrothermal sediment communities oxidize C₁-C₄ alkanes, (ii) the degree to which the
138 anaerobic oxidation of these alkanes is coupled to SR, and (iii) the putative microbial
139 phylotypes mediating C₁-C₄ alkane oxidation. To address these aims, a series of
140 incubations were conducted using slurries of sediments collected from the MV system.
141 These anaerobic batch reactors enabled the quantification and direct comparison of C₁-C₄
142 alkane oxidation and SR rates in a closed system across a broad range of discrete
143 temperatures (25, 55, and 75 °C). Archaeal and bacterial community dynamics were
144 investigated via pyrotag sequencing in select batch reactor sediments that exhibited the
145 greatest alkane oxidation activity over the incubation time course. The overall objective
146 of this study was to advance our understanding of the nature and extent of the anaerobic
147 oxidation of short-chain alkanes in hydrothermal systems and to ascertain the potential
148 influence of these processes on other biogeochemical cycles. The data presented herein
149 shed light on the relative contribution of the anaerobic oxidation of C₂-C₄ alkanes at
150 different temperature regimes, the potential influence on AOM and the sulfur cycle, and
151 the phylotypes most likely allied to the observed metabolisms.

152

153 **Materials and methods**

154 **Study site and sample collection**

155 Sediments were collected during an expedition with the *DSV Alvin* and *R/V Atlantis* in
156 July 2010 from the Chowder Hill hydrothermal vent field in Middle Valley (48° 27.44 N,
157 128° 42.51W) at 2413 m depth. Intact sediment cores were recovered with
158 polyvinylchloride core sleeves (20 - 30 cm height, 6.35 cm ID, 0.32 cm sleeve thickness).
159 Sediment sampling sites were selected based on *in situ* temperature depth profiles
160 collected with *DSV Alvin*, the presence of chemoautotrophic microbial mats atop the
161 sediments, and shimmering water from the diffuse flow sediments. At all sites, sediment
162 temperature profiles were collected using the RTD probe, while dissolved alkanes and
163 other gases were quantified using an *in situ* mass spectrometer (or ISMS; data not shown)
164 (Wankel et al., 2011). Pushcores were collected from areas where sediments temperatures
165 ranged from 5 - 55 °C in the upper 15 cm and 57 - 75 °C at 30 cm sediment depth. Upon
166 retrieval, cores were sealed and refrigerated for transport to the laboratory. Upon return to
167 the lab, the overlying water in the sediment cores was replaced weekly with fresh, filter-
168 sterilized anoxic seawater prior to initiation of the experiments.

169

170 **Anaerobic batch reactors with C₁-C₄ alkanes**

171 In an anaerobic chamber (Coy Laboratory Products), 50 ml of homogenized whole core
172 sediment and 50 ml of sterile, anaerobic artificial “diffuse vent fluid” were aliquoted into
173 200 ml glass autoclaved serum vials for each treatment. The artificial vent fluid was
174 modified from Widdel and Bak (1992) to include 1 mM Na₂S to ensure that sediments
175 remained at reducing conditions, 50 mM Na₂SO₄²⁻ to reduce the possibility of sulfate
176 limitation, and the pH adjusted to 6 to mimic the diffuse vent fluids. For each incubation
177 temperature, the headspace was pressurized to slightly above 1 atm with the respective

178 alkane (C₁-C₄) or nitrogen (N₂) gas in duplicate batch reactors to avoid alkane limitation
179 in the aqueous phase during the incubation time series. The reactors were incubated at
180 temperatures reflecting the sea water-sediment interface (25 °C), the mid-depth average
181 temperature (55 °C), and the highest temperatures measured at the deepest depth (75 °C).
182 Flasks were shaken daily to ensure homogeneity in the slurry.

183

184 **Geochemical measurements**

185 Concentrations of the dissolved C₁, C₂, C₃, and C₄ alkanes were determined after
186 allowing the incubations to reach room temperature and by vigorously shaking samples to
187 transfer gas from the anaerobic seawater media to the batch reactor headspace. Then, a
188 0.5 ml sample of the headspace was injected into a gas chromatograph equipped with a
189 flame ionization detector (Hewlett Packard 5890 Series II) and a packed column (Restek
190 Rt-XL) to quantify all alkanes. Injections of chemically pure alkanes (Airgas East, >
191 99 % purity) were used to generate standard curves.

192

193 SR rates were determined by quantifying changes in sulfate and sulfide concentrations
194 via ion chromatography and colorimetric assays, respectively (Joye et al. 2004; Cline et
195 al. 1969). After shaking and allowing the sediment to settle, a 1 ml fluid subsample was
196 collected with a syringe from each reactor, filter-sterilized (0.2 µm) and transferred into a
197 vial, preserved with 10 µl HNO₃, and stored at 7 °C until analysis. Concentrations of
198 sulfate were determined using a Dionex ion chromatography system (Dionex Corp.
199 Sunnyvale, CA) at the University of Georgia and NaBr, a conservative tracer in the batch
200 reactors, was measured simultaneously. A 1 ml headspace sub-sample was collected and
201 mixed with an equal volume of 20 % zinc acetate to quantify gaseous hydrogen sulfide
202 (H₂S). Concentrations of H₂S were then determined colorimetrically as per Cline (1969).
203 The reported values were corrected for HS⁻ dissolved in the aqueous phase and reflect
204 both sulfide species in the serum vial headspace and sediment slurry.

205

206 **DNA extraction, massively parallel sequencing, and phylogenetic analysis**

207 At the conclusion of each incubation, sediments were sub-sampled in an anaerobic
208 chamber, and ~ 15 g of sediment slurry from each batch reactor was transferred directly
209 into a 15 ml cryovial, flash frozen in liquid nitrogen and stored at -80 °C until further
210 molecular analyses. A time zero (T₀) sub-sample was collected at the start of the
211 incubations to represent the initial community after homogenization, but prior to
212 inoculation of the batch reactors. Total genomic DNA was extracted using phenol-
213 chloroform (Barns et al., 1994; Dojka et al., 1998; Elshahed et al., 2004) modified to
214 prevent nucleic acid loss and eliminate potential inhibitors of downstream PCR (as
215 described in Webster et al. (2004)). Briefly, 0.5 g of sediment per batch reactor was
216 washed with 5 % HCl and then DNA was extracted with addition of 200 µg of poly-
217 adenylic acid (poly A) during the lysis step followed by incubation with lysozyme and
218 proteinase K, multiple freeze-thaw cycles with 5 % SDS, addition of hot phenol,
219 extraction with phenol-chloroform, and elution in 50 µl TE buffer (10 mM Tris
220 hydrochloride, 1 mM EDTA, pH 8.0). The concentration of extracts was determined
221 using the Quant-iT™ dsDNA high sensitivity Assay (Invitrogen, Carlsbad, CA).

222

223 DNA extracted from the 55 °C incubations, which represented the highest rates of
224 activity, was subjected to massively parallel sequencing of the 16S ribosomal RNA
225 (rRNA) gene using the primer pairs 27F / 519R and 340F / 806R for the bacterial V1 –
226 V3 and archaeal V3 – V4 regions, respectively (Dowd et al. 2008; Acosta-Martinez et al.
227 2010). All pyrosequence data were submitted to the NCBI Sequence Read Archive under
228 accession number SRA066151. The resulting reads were checked for sequence quality,
229 trimmed, filtered, and analyzed in the software MOTHUR (Version 1.28.0) (Schloss et al.,
230 2009). Sequences were first filtered by the presence of sequence ambiguities, long
231 homopolymers, and quality scores. The PryoNoise algorithm was then implemented in
232 MOTHUR (i.e., shhh.flows) to remove sequences likely generated by pyrosequencing
233 error (Quince et al. 2009). After selection of unique sequences, chimeras were identified
234 and removed using UCHIME (<http://www.drive5.com/uchime/>). The resulting archaeal
235 and bacterial reads were then aligned to the SILVA SEED Bacterial and Archaeal
236 databases, containing 14,956 and 2,297 sequences, respectively.

237

238 For sequence classification, bootstrap values were set to nodes that had > 80 % support in
239 a bootstrap analysis of 100 replicates, and operational taxonomic units (OTUs) were
240 defined as sequences sharing 97 % nucleotide sequence identity for further community
241 analyses. A phylogenetic tree of representative *Deltaproteobacteria* (50 unique sequences
242 selected in Mothur, i.e., sub.sample) was then generated with FastTree 2.0.0 (Price et al.
243 2010) using minimum-evolution subtree-pruning-regrafting and maximum-likelihood
244 nearest-neighbor interchanges. Local support values shown are based on the Shimodaira-
245 Hasegawa (SH) test with 1,000 resamples. Only values > 80 % are shown on the
246 branches as black circles. The tree was rooted to the 16S rRNA sequence of
247 *Archaeoglobus profundus* DSM 5631 (NR_074522).

248

249 **Results**

250 **C₁-C₄ alkane oxidation as a function of temperature in batch reactors**

251 Batch reactor incubations were conducted using MV sediment slurries with one alkane
252 gas (C₁, C₂, C₃, or C₄) as the sole exogenous hydrocarbon, and incubated in the
253 laboratory at 25, 55, and 75 °C to reflect the range of temperatures measured *in situ*.
254 Temperature affected the time required to detect alkane consumption, the percent of
255 available substrate consumed, and the absolute rates of the anaerobic oxidation of C₁-C₄.
256 In batch reactors at 55 °C, alkane consumption, defined as 10 % of pool consumption,
257 was evident after 71 days of incubation (**Figure 1, top**). In contrast, alkane consumption
258 was detectable in 25 °C batch reactors after 105 days for C₁-C₄. In 75 °C batch reactors,
259 substantial C₂-C₄ consumption was apparent after 105 days; however, C₁ consumption
260 was evident after a much shorter time period (30 days) at 75 °C.

261

262 Examining the fraction of available alkane consumed during the entire experiment (169
263 days), the greatest total consumption of C₁-C₄ occurred in the 55 °C batch reactors (~ 93,
264 75, 93, and 77 % of C₁, C₂, C₃ and C₄, respectively). In addition, C₁ was nearly depleted
265 in the 75 °C batch reactors by the end of the time series (with > 90 % substrate
266 consumed). With the exception of C₁ at 75 °C, less than half of the available short-chain
267 alkane pool was consumed during the incubation time course in 25 and 75 °C batch
268 reactors (32, 44, 37, and 46 % of C₁, C₂, C₃ and C₄ at 25 °C, respectively, and 35, 31, and

269 27 % of C₂, C₃, and C₄ at 75 °C, respectively).

270

271 Absolute rate measurements of the batch reactor sediments revealed that maximum C₁-C₄
272 oxidation occurred at 55 °C (~ 42, 36, 54, and 23 nmol cm⁻³ day⁻¹ for C₁, C₂, C₃, and C₄,
273 respectively, *n* = 2) (**Table 1**). Substantially lower rates of the anaerobic oxidation of C₂-
274 C₄ were observed in all 25 and 75 °C batch reactors (~ 21, 16, and 8 nmol cm⁻³ day⁻¹ for
275 C₂, C₃, and C₄ at 25 °C, respectively, *n* = 2, and ~ 17, 17, and 8 nmol cm⁻³ day⁻¹ for C₂, C₃,
276 and C₄ at 75 °C, respectively, *n* = 2). In contrast to the other short-chain alkanes, maximal
277 rates of AOM were also observed at 75 °C (~ 42 nmol cm⁻³ day⁻¹, *n* = 2), while rates
278 decreased to less than half of these AOM maxima at 25 °C (14 nmol cm⁻³ day⁻¹, *n* = 2).

279

280 **Table 1.** Volume-specific rate measurements of the anaerobic oxidation of methane,
281 ethane, propane and butane and sulfate reduction in batch reactors incubated at 25, 55,
282 and 75 °C. Rates were determined from the consumption of alkanes and sulfate over the
283 incubation time course. Each point represents the average of duplicate reactors with the
284 standard error. To account for background sulfate reduction (due to autochthonous carbon,
285 etc.), the rates measured in the alkane treatments have been corrected via subtraction of
286 those measured in the control (nitrogen) batch reactors.

	Anaerobic Oxidation	Sulfate Reduction
	nmol cm ⁻³ day ⁻¹	nmol cm ⁻³ day ⁻¹
Methane - 25°C	14.33 ± 2.88	15.01 ± 2.31
Methane - 55°C	41.45 ± 1.17	55.83 ± 4.91
Methane - 75°C	42.22 ± 1.91	68.03 ± 5.01
Ethane - 25°C	21.39 ± 4.77	53.61 ± 6.53
Ethane - 55°C	36.03 ± 4.46	99.30 ± 8.48
Ethane - 75°C	17.22 ± 3.59	47.94 ± 3.65
Propane - 25°C	19.93 ± 2.81	71.96 ± 7.25
Propane - 55°C	53.66 ± 2.52	238.36 ± 10.77
Propane - 75°C	17.26 ± 1.26	60.37 ± 6.18
Butane - 25°C	12.84 ± 2.81	55.27 ± 8.68
Butane - 55°C	23.07 ± 5.13	113.46 ± 15.37
Butane - 75°C	8.01 ± 1.13	34.22 ± 2.39

287

288 Sulfate reduction coupled to C₁-C₄ alkane oxidation across temperature regimes

289 In addition to a dependence on short-chain alkane length, temperature constrained SR in
290 the anaerobic batch reactors, influencing quantified changes in porewater sulfate and total
291 sulfide. Decreases in sulfate concentration were observed in all batch reactors across time
292 and temperature regimes, consistent with trends for the anaerobic oxidation of C₁-C₄
293 alkanes. Analogous to alkane consumption dynamics, sulfate consumption was
294 appreciable (defined as > 10 % substrate consumption) after 71 days of incubation in C₁-

295 C₄ batch reactors at 55 °C (**Figure 1, middle**). In contrast, there was a lag of ~ 105 days
 296 in C₂, C₃, and C₄ batch reactors prior to substantial sulfate consumption at both the lowest
 297 (25 °C) and highest (75 °C) incubation temperature. Over the span of the incubation time
 298 series (169 days), the greatest reduction in sulfate concentration was at 55 °C (~ 30, 45,
 299 92, and 49 % of total sulfate consumed in the C₁, C₂, C₃, and C₄ reactors, respectively).
 300 Sulfate consumption was also observed in the N₂-control batch reactors, albeit to a much
 301 smaller extent (~ 8, 11, and 2 % at 25, 55, and 75 °C, respectively). Sulfate reduction was
 302 also assessed by quantifying the production of gaseous and dissolved sulfide in the batch
 303 incubations (**Figure 1, bottom**). In all reactors, sulfide concentrations at the end of each
 304 incubation time period accounted for greater than 90 % of the initial total sulfate plus
 305 sulfide concentration; therefore, these mass balance estimates were within 10 % of the
 306 total sulfur species observed initially.

307
 308 Concomitant with the anaerobic oxidation of C₂-C₄ rates, maximum SR rates were
 309 observed at 55 °C for the non-methane short-chain alkanes (~ 99, 238, and 113 nmol cm⁻³
 310 day⁻¹ for C₂, C₃, and C₄, respectively, *n* = 2) (**Table 1**). However, maximum SR rates
 311 associated with AOM occurred at 75 °C (~ 68 nmol cm⁻³ day⁻¹), with lower rates at 55 °C
 312 (~ 55 nmol cm⁻³ day⁻¹, *n* = 2) and even more modest rates at 25 °C (~ 15 nmol cm⁻³ day⁻¹,
 313 *n* = 2). In comparison to maximal SR rates at 55 °C, SR rates linked to C₂-C₄ oxidation
 314 were lower at both 25 and 75 °C (~ 54, 72, and 55 nmol cm⁻³ day⁻¹ for C₂, C₃, and C₄ at
 315 25 °C, respectively, *n* = 2, and ~ 48, 60, and 34 nmol cm⁻³ day⁻¹ for C₂, C₃, and C₄ at 75
 316 °C, respectively, *n* = 2).

317
 318 The observed ratio (mol/mol) of C₁-C₄ oxidation to sulfate reduction in the batch reactors
 319 was then compared to the predicted stoichiometric ratio assuming the sulfate-dependent
 320 complete oxidation of C₁-C₄ alkanes to CO₂ (from Knemeyer et al., 2007). These ratios
 321 are corrected for the consumption of sulfate in the control (N₂) batch reactors as an
 322 estimate for sulfate reduction linked to non-alkane organic carbon donors present in the
 323 sediment. The ratio of mol alkane consumed per mol sulfate reduced was 1.42, 1.11, and
 324 0.93 mmol of C₁ mmol⁻¹ sulfate; 0.59, 0.54, and 0.54 mmol of C₂ mmol⁻¹ sulfate; 0.42,
 325 0.34, and 0.43 mmol of C₃ mmol⁻¹ sulfate; and 0.35, 0.31, and 0.35 mmol of C₄ mmol⁻¹
 326 sulfate at 25, 55, and 75 °C, respectively (**Table 2**). These ratios closely mirror the
 327 predicted stoichiometric ratios of 1, 0.5, 0.4, and 0.3 for C₁-C₄, respectively.

328
 329 **Table 2.** The predicted and calculated stoichiometric ratios for the anaerobic oxidation of
 330 methane, ethane, propane, and butane coupled to the reduction of sulfate to sulfide. From
 331 closed system batch reactors, the mol alkane lost was calculated per mol sulfate reduced
 332 at 25, 55, and 75 °C, respectively. To account for background sulfate reduction (due to
 333 autochthonous carbon, etc.), ratios have been corrected via subtraction of those measured
 334 in the control (nitrogen) batch reactors.
 335

	Stoichiometric Ratio	Observed Ratio
	(mol/mol)	(mol/mol)
Methane - 25°C	1	1.42

Methane - 55°C	1	1.11
Methane - 75°C	1	0.93
Ethane - 25°C	0.5	0.59
Ethane - 55°C	0.5	0.54
Ethane - 75°C	0.5	0.54
Propane - 25°C	0.4	0.42
Propane - 55°C	0.4	0.34
Propane - 75°C	0.4	0.43
Butane - 25°C	0.31	0.35
Butane - 55°C	0.31	0.31
Butane - 75°C	0.31	0.35

336

337 **Phylogenetic diversity and distribution in sediments from batch C₁-C₄ reactors**

338 After sequence processing and denoising as previously described, a total of 5783, 6562,
 339 5307, 6985, and 8796 bacterial sequences were analyzed from sediments incubated with
 340 N₂, C₁, C₂, C₃, and C₄ alkane, respectively, and 7965 bacterial sequences from the T₍₀₎
 341 sediment. There were substantial shifts at the phyla level between the communities
 342 incubated with different alkane substrates in comparison to the control batch reactor and
 343 T₍₀₎ sediment community (**Figure 2**). From the initial sediment community, sequences
 344 allied to the *Bacteroidetes* and *Fusobacteria* decreased from ~ 9 and 40 % of T₍₀₎
 345 sequences respectively, to less than 0.5 % of sequences in all batch reactor libraries. In
 346 turn, sequences allied to the *Proteobacteria*, *Firmicutes*, Candidate Division OP8,
 347 *Chloroflexi*, and *Actinobacteria* increased in batch reactor libraries compared to T₍₀₎
 348 sequences. Notably, the *Proteobacteria*, which comprised ~ 36 % of T₍₀₎ sequences,
 349 increased in representation in the N₂, C₁, C₂, and C₃ sequence libraries (~ 49, 58, 41, and
 350 46 %, respectively). The *Firmicutes* also increased substantially from the T₍₀₎
 351 composition (~ 4 %) in N₂, C₂, C₃, and C₄ sequences (~ 11, 12, 12, and 59 %, respectively).
 352

353

354 Among the *Proteobacterial* sequences allied to known sulfate-reducing
 355 *Deltaproteobacteria*, there was a substantial increase from T₍₀₎ sequences (~15 %) in the
 356 N₂, C₁, C₂, C₃, and C₄ sequence libraries (~ 39, 87, 70, 88, and 86 %, respectively).
 357 Concurrently, there was a substantial decrease in the representation of
 358 *Epsilonproteobacteria* in the N₂, C₁, C₂, C₃, and C₄ sequence libraries (~ 39, 3, 17, 7, and
 359 9 %, respectively). Within the putative sulfate-reducing phylotypes, the C₁ library was
 360 comprised primarily (> 92 %) of sequences allied to *Desulfobulbus*, as shown in a
 361 previous study of MV sediment communities associated with AOM (Wankel, Adams et
 362 al., 2012). Analysis of 16S rRNA gene libraries revealed that a distinct lineage of SRB
 363 are the predominant *Deltaproteobacterial* phylotypes in the C₂-C₄ reactor communities,
 364 comprising ~ 93, 91, and 95 % of C₂, C₃, and C₄ sequences, respectively (**Figure 4**). The
 365 most closely related phylotypes (93 - 99 % nucleotide sequence identity) were previously
 366 recovered in two 16S rRNA-based surveys of sulfate-reducing anaerobic enrichments of

367 Guaymas Basin sediments with C₄ at 60 °C (Butane60-GuB, accession no. EF077228)
368 and with C₁ at 37 °C (Guaymas_Bac9 clone, accession no. FR682643) (Kniemeyer et al.,
369 2007; Kellermann et al., 2011).

370

371 A total of 1290, 1724, 1540, 1916, 1780, and 2846 Euryarchaeotal sequences were
372 further analyzed from the N₂, C₁, C₂, C₃, and C₄ batch reactors and T₍₀₎ sediments,
373 respectively (**Figure 3**). There were notable shifts in the sequences allied to the
374 predominant Euryarchaeotal phyla - *Archaeoglobi*, *Halobacteria*, *Methanomicrobia*,
375 *Thermococci*, and *Thermoplasmata* - from the initial sediment community and across the
376 different alkane batch incubations. Over 40 % of sequences were allied to the
377 *Halobacteria* in T₍₀₎ sediments, decreasing to comprise < 0.5 to 29 % of batch reactor
378 sequences. In contrast, *Archaeoglobi* sequences increased from ~ 2 % of T₍₀₎ sequences
379 to ~ 14, 12, 19, 36, and 29 % of N₂, C₁, C₂, C₃, and C₄ sequences, respectively. Other
380 trends in Euryarchaeotal community structure included an increase in *Methanomicrobia*
381 from 27 % of T₍₀₎ sequences to 40 % of C₁ sequences.

382

383 Within the *Methanomicrobia*, there were also substantial changes in sequences allied to
384 known methanogens and methane-oxidizing phylotypes. *Methanosarcinales* comprised >
385 97 % of T₍₀₎ sequences and ~ 19, 36, 84, 75, and 80 % of N₂, C₁, C₂, C₃, and C₄
386 sequences, respectively. In contrast, *Methanobacteriales* increased from < 0.5 % of T₍₀₎
387 sequences to ~ 35, 9, and 26 % of N₂, C₁ and C₃ sequences (there was no substantial
388 increase in C₂ or C₄ sequences). For the putative methane-oxidizing communities, over 40
389 and 12 % of C₁ and C₂ sequences were allied to ANME-1 ribotypes.

390

391 **Discussion**

392 The microbial degradation of short-chain alkanes under oxic conditions and the anaerobic
393 oxidation of methane and heavier hydrocarbons have been extensively studied in diverse
394 terrestrial and marine environments. Despite studies indicating short-chain alkane
395 degradation in anoxic deep sea sediments (Sassen et al., 2004; Mastalerz et al., 2009;
396 Quistad and Valentine, 2011) and the abundance of short-chain alkanes in hydrocarbon-
397 rich ecosystems (Milkov, 2005; Cruse and Seewald, 2006), relatively little is known
398 about the biogeochemical importance of these processes or the diversity of anaerobic
399 short-chain alkane degrading microorganisms in marine hydrothermal sediments. The
400 data here provide a deeper glimpse into the anaerobic oxidation of C₂-C₄ in metalliferous
401 hydrothermal sediments and reveal that rates of the anaerobic oxidation of C₂-C₄ alkanes
402 in hydrothermal vent sediment are heavily influenced by temperature and coupled to
403 sulfate reduction, though the rates presented herein are derived from conditions not likely
404 to be present *in situ*, and as such care should be taken when extrapolating these rates to
405 natural processes. In batch reactor sediments that exhibited the most substantial activity,
406 changes in the representation of phylotypes in libraries generated via high throughput
407 sequencing implicate *Deltaproteobacteria* in C₂-C₄ alkane degradation, and shifts in
408 microbial community composition indicate that other members of the community respond
409 to the presence of short chain alkanes (though the mechanisms underlying this response
410 remain unknown).

411

412 These data revealed a preferential consumption of C₂-C₄ at 55 °C, suggesting that the

413 active alkane degraders in these hydrothermal vent sediments are thermophilic.
414 Furthermore, these *ex situ* calculated rates for the anaerobic oxidation of C₂-C₄ were in
415 the same range (nmol cm⁻³ day⁻¹) as the recently-reported anaerobic oxidation of C₃ in
416 marine hydrocarbon seep sediments and as AOM rates measured in organic-rich coastal
417 sediments at the sulfate-methane transition zone (Quistad and Valentine, 2011; Alperin et
418 al., 1988; Hoehler et al., 1994; Girguis et al., 2003; Wegener et al., 2008). Based on lag
419 time and total alkane degraded over time, C₃ appeared to be the preferred substrate in the
420 55 °C incubations, followed by C₁, C₄, and C₂, respectively. Similar trends in the
421 biodegradation of short-chain alkanes have been found in stable isotopic studies of
422 hydrocarbon reservoirs at temperatures below 60 °C with a preference for C₃ followed by
423 C₄ and then C₂ (Boreham et al., 2001; Wenger et al., 2002; Larter et al., 2005).

424
425 Various physicochemical and biotic parameters may impact the degree of C₂-C₄
426 consumption in *ex situ* studies and in the natural environment. Notably, the gaseous
427 alkanes were maintained at above saturation conditions for the liquid phase of the batch
428 incubations until the end of the time series to ensure substrate availability (dissolved
429 concentrations of ~1.42, 1.89, 0.91, and 1.05 mM for C₁, C₂, C₃, and C₄, respectively).
430 Under elevated hydrostatic pressure in the deep sea, hydrothermal vent fluids at Middle
431 Valley reach C₁ concentrations of ~ 20 mM, while the other short-chain alkanes are an
432 order of magnitude lower (~ 220, 55, and 6 μM for C₂, C₃, and C₄, respectively) (Cruse
433 and Seewald, 2006). Although C₁ is most likely more abundant than C₃ in MV
434 hydrothermal sediments, the *in situ* rates of C₃ degradation may be appreciable due to the
435 inherent reactivity of secondary C-H bonds (Rabus et al., 2001; Schink and Freidrich,
436 1994; Van Hamme et al., 2003).

437
438 Our results also suggest that, at the highest incubation temperatures, AOM in MV
439 sediments occurs at higher rates than the anaerobic oxidation of C₂-C₄ alkanes. In the
440 higher temperature (75 °C) incubations, C₁ consumption was evident after 30 days and
441 reached near deplete concentrations (90 % substrate consumed), while there was a much
442 longer lag period until C₂-C₄ degradation (105 days) and much less of the substrates were
443 consumed by the completion of the time series (27 – 35 %) (**Figure 1, top**). In contrast, a
444 greater proportion of C₂ and C₄ (44 and 46 %, respectively) were consumed than C₁ and
445 C₃ (32 and 37 %, respectively) in the lower temperature incubations (25 °C). The
446 increased AOM activity at the higher end of the temperature range in MV sediments is
447 consistent with our previous observations of AOM in these metalliferous sediments
448 (Wankel, Adams, et al. 2012), and is also consistent with the growth temperatures of
449 archaeal communities (such as ANME phylotypes) from hydrothermal vents, which
450 indicate that many archaea live at their maximum growth temperature *in situ* (Kimura et
451 al., 2013). Another line of evidence for thermophilic AOM was also provided in a recent
452 16S rRNA based-study identifying a putatively high temperature-adapted ANME
453 subgroup in both hydrothermal sediments from Guaymas Basin and diffuse vent fluids
454 from Axial Volcano and the Endeavor Segment of Juan de Fuca Ridge (Merkel et al.,
455 2013).

456
457 Notably, the anaerobic oxidation of C₁-C₄ was coupled to SR across temperature
458 gradients in MV sediment batch reactors. Sulfate loss (~ 2 - 6 mM) was also observed

459 over the time series in alkane-free control batch reactors (**Figure 1, middle**). In
460 comparison to SR linked to the oxidation of short-chain alkanes, this modest sulfate
461 consumption relates to the oxidation of endogenous substrates, particularly organic
462 carbon, by the sediment communities (Gieg et al., 1999). The sediment organic carbon
463 pool of MV sediments (% OC = ~ 0.5 in this study) is low in comparison to the high
464 amounts of organic matter that characterize other deep sea environments with known
465 short-chain alkane degraders, such as the organic-rich Guaymas Basin hydrothermal
466 sediments (Jorgensen et al., 1992; Kniermeyer et al., 2007). The observed SR rates in C₁-
467 C₄ batch reactors of MV sediments demonstrate the potential for organic carbon-poor,
468 high temperature mid-ocean ridge systems to support the anaerobic oxidation of short-
469 chain alkanes coupled to SR.

470
471 Our results further indicate that short-chain alkane degradation linked to SR might
472 considerably influence sulfate cycling at these sedimented hydrothermal vents. In
473 accordance with the observed stoichiometries, SR coupled to the anaerobic oxidation of
474 C₂, C₃, and C₄ proceeded at a faster rate than AOM at mesophilic and thermophilic
475 temperatures (25 and 55 °C, respectively). However, the SR rates in anaerobic batch
476 reactors were observed under sulfate-replete conditions, while the sulfate pool *in situ*
477 depends on the downward advection of seawater and the activity of sulfide-oxidizing
478 microbial communities (Bowles et al., 2011). Sulfate availability will become limiting at
479 greater sediment depths from the seawater surface. Therefore, the C₂, C₃, and C₄-
480 degrading, sulfate-reducing microbial communities likely compete for available sulfate
481 and might indirectly limit AOM in the temperature range from ~ 25 – 55 °C. As
482 previously discussed, the anaerobic oxidation of these aliphatic hydrocarbons coupled to
483 the reduction of sulfate to sulfide yields greater energy per unit substrate than AOM.
484 Such processes could constrain methane release from the deep-sea with a critical impact
485 on the global carbon cycle and climate. Furthermore, if AOM activity peaks at greater
486 sediment depths and higher temperatures *in situ* as predicted by rate measurements, then
487 sulfate will most likely have been depleted in these sediment horizons. Sulfate limitation
488 may thus result in the coupling of AOM to alternative electron acceptors (i.e., iron
489 oxides), as indicated in previous studies of MV high temperature sediment incubations
490 (Wankel, Adams et al., 2012).

491
492 Comparison of bacterial communities in batch reactor sediments with maximum rates of
493 C₁-C₄ degradation, via massively parallel pyrosequencing, suggests that members of the
494 sulfate-reducing *Deltaproteobacteria* mediate the anaerobic oxidation of short-chain
495 alkanes in MV hydrothermal vent sediments (**Figure 2**). As these sequence data are based
496 on PCR amplification of 16S rRNA genes and are semi-quantitative, an order of
497 magnitude difference in phyla should represent shifts in community composition. Within
498 the *Proteobacteria*, there was a substantial increase of *Deltaproteobacteria* in C₁-C₄
499 sequences compared to the initial T₍₀₎ bacterial composition dominated by
500 *Epsilonproteobacteria*. Phylogenetic analyses revealed a novel subgroup of SRB that
501 comprised > 90 % of these *Deltaproteobacteria* in C₂-C₄ batch reactor sequences (**Figure**
502 **4**). This lineage of *Deltaproteobacteria* is most closely related to C₄-degrading SRB from
503 Guaymas Basin, and therefore, may be a thermophilic short-chain alkane degrader group
504 (Kniermeyer et al., 2007). Intriguingly, the predominant phylum in C₄ batch reactor

505 sequences is the *Firmicutes*, which contains sulfate-reducing members of previous
506 enrichments with C₃ and C₄ (Kniemeyer et al. 2007; Savage et al., 2010). However, the
507 majority of *Firmicutes* sequences were most closely related (98 – 99 %) to uncultured
508 *Bacillus* clones from hydrocarbon-contaminated soils (Wang et al., 2010). The greater
509 proportion of this uncharacterized *Bacillus* group in comparison to SRB may have also
510 affected the lower rates of the anaerobic oxidation of C₄ in comparison to C₂ or C₃ in
511 batch incubations. Future studies should determine if this putative thermophilic short-
512 chain alkane degrader group of SRB is widespread in other hydrothermally influenced
513 environments.

514

515 Amongst the batch reactor sediment communities, shifts in archaeal phylogenetic
516 composition were also revealed via 16S rRNA pyrosequencing. There was a substantial
517 increase of *Methanomicrobia* sequences in C₁-incubated sediments compared to the
518 initial community and C₂ - C₄ batch reactor sediments, with an order of magnitude
519 enrichment of ANME-1 phylotypes within the *Methanomicrobia* (**Figure 3**). The
520 microbes known to catalyze AOM form three phylogenetically distinct Euryarchaeota
521 clusters (ANME-1, ANME-2, and ANME-3) that often appear to live in consortia with
522 SRB (Boetius et al., 2000; Hoehler et al., 1994; Orphan et al., 2001). However, ANME-1
523 phylotypes are also found as single cells in sediments, and recent studies have shown that
524 AOM can occur in the absence of SR and that some ANME are not directly dependent on
525 SRB activity (Beal et al., 2009; Milucka et al., 2012; Wankel, Adams et al., 2012). There
526 was also a notable increase in *Archaeoglobus* sequences in C₁ – C₄ batch reactors from
527 the initial community composition (**Figure 3**), which contain hyperthermophilic species
528 known to mediate sulfate reduction (Shen et al., 2004). Based on microbial isolates and
529 enrichments from both deep sea and terrestrial ecosystems, no evidence to date indicates
530 that non-methane short-chain alkanes are anaerobically oxidized by microbial consortia
531 or sulfate-reducing archaea (Kniemeyer et al., 2007; Savage et al., 2011; Jaekel et al.,
532 2012). The data presented herein lack the resolution to conclusively address whether
533 archaeal phylotypes directly mediate or are members of consortia that perform the
534 anaerobic oxidation of short-chain alkanes other than AOM.

535

536 The collective results presented here shed light on the potential anaerobic metabolism of
537 short-chain alkanes linked to SR in the hydrothermal vent sediments of Middle Valley,
538 Juan de Fuca Ridge. Substantial oxidation of C₁-C₄ occurs up to 75 °C. The coupling of
539 C₂-C₄ with sulfate reduction over the *in situ* temperature range may impact AOM and the
540 oxidation of other hydrocarbons, as highlighted by the preferential degradation of C₃ at
541 55 °C. Such microbial communities may play a substantial role in carbon and sulfur
542 cycling at hydrothermal systems on a global-scale. Future studies should expand upon
543 other environmental conditions that may regulate the anaerobic oxidation of C₂-C₄
544 alkanes in hydrothermal sediments and should further characterize the *in situ* abundance
545 and activity of the putative thermophilic alkane-degrader SRB lineage.

546

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560

561 **Author Contributions**

562 M. M. A., P. R. G., and A. B. designed the research. M. M. A., P. R. G., and A. L. H.
563 directed the *in situ* collections and measurements. M. M. A., A. L. H., and A. B.
564 conducted the batch reactor incubations and geochemical analyses. S. B. J. performed the
565 sulfate consumption measurements. M. M. A. directed and analyzed the alkane
566 consumption, sulfide production, and all rate calculations. M. M. A. performed the
567 molecular analyses. M.M.A. and P.R.G. wrote the manuscript with input from S. B. J., A.
568 B., and A. L. H. The authors declare no conflict of interest.

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570

571 **References**

572 Alperin, M. J., Reeburgh, W. S., and Whiticar, M. J. (1988). Carbon and hydrogen
573 isotope fractionation resulting from anaerobic methane oxidation. *Global*
574 *Biogeochem. Cyc.* 2, 279–88.

575 Ames, D. E., Franklin, J. M., and Hannington, M. H. (1993). Mineralogy and
576 geochemistry of active and inactive chimneys and massive sulfide, Middle Valley,
577 northern Juan de Fuca Ridge: An evolving hydrothermal system. *Can. Mineral.*
578 31, 997–1024.

579 Barns, S. M., Fundyga R. E., Jeffries M. W., and Pace, N. R. (1994). Remarkable
580 archaeal diversity detected in a Yellowstone National Park hot spring
581 environment. *Proc. Nat. Acad. Sci.* 91, 1609–1613.

582

583 Beal, E.J., House, C.H., and Orphan, V.J. (2009) Manganese- and Iron-Dependent
584 Marine Methane Oxidation. *Science* 325, 184-187.

585

586 Boetius, A., Ravensschlag, K., Schubert, C. J., Rickert, D., Widdel, F., Gieseke, A.,
587 Amann, R., Jørgensen, B. B., Witte, U., and Pfannkuche, O. (2000). A marine
588 microbial consortium apparently mediating anaerobic oxidation of methane.
589 *Nature* 407, 623-626.

590

591 Boreham, C. J., Hope, J. M., and Hartung-Kagi, B. (2001). Understanding source,
592 distribution and preservation of Australian natural gas: a geochemical perspective.
593 *Aust. Prod. Petr. Explor. Assoc. J.* 41, 523–547.

594

- 595 Bowles M. W., Samarkin. A., Bowles K. M., and Joye, S. B. (2011). Weak coupling
596 between sulfate reduction and the anaerobic oxidation of methane in methane-rich
597 seafloor sediments during *ex situ* incubation. *Geochim. Cosmochim. Acta* 75,
598 500–519.
599
- 600 Butterfield D. A., Massoth G. J., McDuff R. E., Lupton J. E., and Lilley, M. D. (1990).
601 Geochemistry of hydrothermal fluids from Axial Seamount hydrothermal
602 emissions study vent field, Juan-de-Fuca Ridge - Subseafloor boiling and
603 subsequent fluid-rock interaction. *J. Geophys. Res. Solid.* 95, 12895–12921.
604
- 605 Butterfield D. A., McDuff R. E., Mottl M. J., Lilley M. D., Lupton J. E., and Massoth, G.
606 J. (1994). Gradients in the composition of hydrothermal fluids from the
607 Endeavour segment vent field: Phase separation and brine loss. *J. Geophys. Res.*
608 99, 9561–9583.
609
- 610 Cline, J. D. (1969). Spectrophotometric determination of hydrogen sulfide in natural
611 waters. *Limnol. Oceanogr.* 14, 454–458.
612
- 613 Cruse A. M., and Seewald, J. S. 2006. Geochemistry of low-molecular weight
614 hydrocarbons in hydrothermal fluids from Middle Valley, northern Juan de Fuca
615 Ridge. *Geochim. Cosmochim. Acta* 70, 2073–2092.
616
- 617 Cruse, A.M., Seewald, J. S., Saccocia, P. J., and Zierenberg, R. A. (2008) Geochemistry
618 of hydrothermal fluids from Middle Valley, northern Juan de Fuca
619 Ridge: Temporal variability, subsurface conditions and equilibration during
620 upflow. In *Magma to Microbe: Modeling Hydrothermal Processes at Oceanic
621 Spreading Ridge*, (ed. R. Lowell and J. S. Seewald) *AGU Monograph Series* 178,
622 145–166.
- 623 Cruse, A. M. and Seewald, J. S. (2010). Low-molecular weight hydrocarbons in vent
624 fluids from the Main Endeavour Field, northern Juan de Fuca Ridge. *Geochim.*
625 *Cosmochim. Acta* 74, 6126–6140.
- 626 Conrad, R. (2009). The global methane cycle: recent advances in understanding the
627 microbial processes involved. *Environ. Microbiol. Rep.* 1, 285–292.
- 628 Davis, E. E. and Fisher, A. T. (1994). On the nature and consequences of hydrothermal
629 circulation in the Middle Valley sedimented rift: Inferences from geophysical and
630 geochemical observations, Leg 139. In *Proc. ODP Sci. Results*, eds Mottl, M. J.,
631 Davis, E. E., Fisher, A. T., and Slack, J. F. 139, 695–717.
632
- 633 Dhima, A., de Hemptinne, J-C., and Moracchini, G. (1998). Solubility of light
634 hydrocarbons and their mixtures in pure water under high pressure. *Fluid Phase
635 Equilibr.* 145, 129–150.
636

637 Dojka, M. A., Hugenholtz, P., Haack, S. K., and Pace, N. R. (1998). Microbial diversity
638 in a hydrocarbon- and chlorinated-solvent-contaminated aquifer undergoing
639 intrinsic bioremediation. *Appl. Environ. Microbiol.* 64, 3869–3877.
640

641 Elshahed, M. S., Najar, F. Z., Roe, B. A., Oren, A., Dewers, T. A. and Krumholz, L. R.
642 (2004). Survey of archaeal diversity reveals an abundance of halophilic archaea in
643 a low-salt, sulfide- and sulfur-rich spring. *Appl. Environ. Microbiol.* 70, 2230–
644 2239.
645

646 Formolo, M. J., Lyons, T. W., Zhang, C., Kelley, C., Sassen, R., Horita, J., and Cole, D.
647 R. (2004). Quantifying carbon sources in the formation of authigenic carbonates
648 at gas hydrate sites in the Gulf of Mexico. *Chem. Geol.* 205, 253-264.
649

650 Gieg, L. M., Kolhatkar, R. V., McInerney, M. J., Tanner, R. S., Harris, S. H., Sublette K.
651 L., and Suflita, J. M. (1999). Intrinsic bioremediation of petroleum hydrocarbons
652 in a gas condensate-contaminated aquifer. *Environ. Sci. Technol.* 33, 2550–2560.
653

654 Girguis, P. R., Orphan, V. J., Hallam, S. J., and DeLong, E. F. (2003). Growth and
655 methane oxidation rates of anaerobic methanotrophic archaea in a continuous
656 flow reactor bioreactor. *Appl. Environ. Microbiol.* 69, 5492–5502.

657 Goodfellow, W. D., and Blaise, B. (1988) Sulfide formation and hydrothermal alteration
658 of hemipelagic sediment in Middle Valley, northern Juan de Fuca Ridge. *Can*
659 *Mineral.* 26, 675–696.

660 Hinrichs, K., Hayes, J. M., Sylva, S. P., Brewer, P. G., and DeLong, E. F. (1999).
661 Methane-consuming archaeobacteria in marine sediments. *Nature* 398, 802–805.
662

663 Hinrichs, K., Hayes, J., Bach, W., Spivack, A., Hmelo, L., Holm, N., Johnson, C., and
664 Sylva, S. (2006). Biological formation of ethane and propane in the deep marine
665 subsurface. *Proc. Nat. Acad. Sci.* 103, 14684–14689.
666

667 Hoehler, T., Alperin, M., and Albert, D. (1994). Field and laboratory studies of methane
668 oxidation in anoxic marine sediment: evidence for a methanogen-sulfate reducer
669 consortium. *Global Biogeochem. Cyc.* 8, 451–463.
670

671 Merkel, A.Y., Huber, J.A., Chernyh, N.A., Bonch-Osmolovskaya, E.A., and Lebedinsky,
672 A.V. (2013). Detection of putatively thermophilic anaerobic methanotrophs in
673 diffuse hydrothermal vent fluids. *Appl. Environ. Microbiol.* 79, 915-23.
674

675 Milucka, J., Ferdelman, T. G., Polerecky, L., Franzke, D., Wegener, G., Schmid, M.,
676 Lieberwirth, I., Wagner, M., Widdel, F., and Kuypers, M. M. (2012) Zero-valent
677 sulphur is a key intermediate in marine methane oxidation. *Nature* 491, 541-546.
678

679 Jaekel, U., Musat, N., Adam, B., Kuypers, M., Grundmann, O., & Musat, F. (2012).
680 Anaerobic degradation of propane and butane by sulfate-reducing bacteria
681 enriched from marine hydrocarbon cold seeps. *ISME J.* 1 – 11.

682
683 Joye, S. B., Boetius, A., Orcutt, B., Montoya, J., Schulz, H., Erickson, M., and Lugo, S..
684 (2004). The anaerobic oxidation of methane and sulfate reduction in sediments
685 from Gulf of Mexico cold seeps. *Chem. Geol.* 205, 219–238.
686
687 Jorgensen, B. B., Isaksen, M. F., and Jannasch, H. W. (1992). Bacterial sulfate reduction
688 above 100 °C in deepsea hydrothermal vent sediments. *Science* 258, 1756–1757.
689
690 Kellermann, M. Y., Wegener, G., Elvert, M., Yoshinaga, M. Y., Lin, Y. S., Holler, T.,
691 Mollar, X. P., Knittel, K., and Hinrichs, K. U. (2012). Autotrophy as a
692 predominant mode of carbon fixation in anaerobic methane-oxidizing microbial
693 communities. *Proc. Nat. Acad. Sci.* 109, 19321-19326.
694
695 Kinnaman, F. S., Valentine, D. L., and Tyler, S. C. (2007). Carbon and hydrogen isotope
696 fractionation associated with the aerobic microbial oxidation of methane, ethane,
697 propane and butane. *Geochim. Cosmochim. Acta* 71, 271– 283.
698
699 Kimura, H., Mori, K., Yamanaka, T., and Ishibashi, J. I. (2013). Growth temperatures of
700 archaeal communities can be estimated from the guanine-plus-cytosine contents
701 of 16S rRNA gene fragments. *Environ. Microbiol. Rep.* In press.
702
703 Kniemeyer, O., Musat, F., Sievert, S. M., Knittel, K., Wilkes, H., Blumenberg, M.,
704 Michaelis, W., Classen, A., Bolm, C., Joye, S. B., and Widdel, F. (2007).
705 Anaerobic oxidation of short-chain hydrocarbons by marine sulphate-reducing
706 bacteria. *Nature* 449, 898–901.
707
708 Knittel, K., and Boetius, A. (2009). The anaerobic oxidation of methane - progress with
709 an unknown process. *Annu. Rev. Micro.* 63, 311–334.

710 Larter, S. R., Head, I. M., Huang, H., Bennett, B., Jones, M., Aplin, A. C., Murray, A.,
711 Erdmann, M., Wilhelms, A., and Di Primio, R. (2005). Biodegradation, gas
712 destruction and methane generation in deep subsurface petroleum reservoirs: an
713 overview. *Q. J. Geol. Soc. Lond.* 6, 633–639.
714
715 Lorenson, T. D., Kvenvolden, K. A., Hostettler, F. D., Rosenbauer, R. J., Orange, D. L.,
716 Martin, J. B. (2002). Hydrocarbon geochemistry of cold seeps in the Monterey
717 Bay National Marine Sanctuary. *Mar. Geol.* 181, 285-304
718
719 Mastalerz, V., de Lange, G. J., and Dahlmann, A. (2009). Differential aerobic and
720 anaerobic oxidation of hydrocarbon gases discharged at mud volcanoes in the
721 Nile deep- sea fan. *Geochim. Cosmochim. Acta* 73, 3849–3863.
722
723 Milkov, A. V. (2005). Molecular and stable isotope compositions of natural gas hydrates:
724 A revised global dataset and basic interpretations in the context of geological
725 settings. *Org. Geochem.* 36, 681–702.
726

- 727 Muyzer, G., and van der Kraan, G. (2008). Bacteria from hydrocarbon seep areas
728 growing on short chain alkanes. *Trends Microbiol.* 16, 138–141.
729
- 730 Orphan, V., House, C., and Hinrichs, K. (2001). Methane-consuming archaea revealed by
731 directly coupled isotopic and phylogenetic analysis. *Science* 293, 484-487.
732
- 733 Price, M. N., Dehal, P. S., and Arkin, A. P. (2010). FastTree 2--approximately maximum-
734 likelihood trees for large alignments. *PLOS ONE* 5, e9490.
- 735 Quince, C., Lanzén, A., Curtis, T. P., Davenport, R. J., Hall, N., Head, I. M., Read, L.F,
736 and Sloan, W. T. (2009). Accurate determination of microbial diversity from 454
737 pyrosequencing data. *Nat. Methods* 6, 639–641.
738
- 739 Quistad S. D., and Valentine D.L. (2011). Anaerobic propane oxidation in marine
740 hydrocarbon seep sediments. *Geochim. Cosmochim. Acta* 75, 2159-2169.
741
- 742 Rabus, R., Wilkes, H., Behrends, A., and Armstroff, A. (2001). Anaerobic initial reaction
743 of n-alkanes in a denitrifying bacterium: Evidence for (1-methylpentyl)succinate
744 as initial product and for involvement of an organic radical in *n*-hexane
745 metabolism. *J. Bact.* 183,1707–1715.
746
- 747 Rushdl, A. I., and Simonelt, B. R. T. (2002). Hydrothermal alteration of organic matter in
748 sediments of the Northeastern Pacific Ocean: Part 1. Middle Valley, Juan de Fuca
749 Ridge. *Appl. Geochem.* 17, 1401–1428.
- 750 Sassen, R., Roberts, H. H., Carney, R., Milkov, A. V., DeFreitas, D. A., Lanoil, B., and
751 Zhang, C. (2004). Free hydrocarbon gas, gas hydrate, and authigenic minerals in
752 chemosynthetic communities of the northern Gulf of Mexico continental slope:
753 relation to microbial processes. *Chem. Geol.* 205, 195—217.
754
- 755 Savage, K. N., Krumholz, L. R., Gieg, L. M., Parisi, V. A., Suflita, J. M., Allen, J., Philp,
756 R. P., and Elshahed, M .S. (2010). Biodegradation of low-molecular-weight
757 alkanes under mesophilic, sulfate-reducing conditions: metabolic intermediates
758 and community patterns. *FEMS Microbiol. Ecol.* 72, 485–95.
- 759 Schink, B., and Friedrich, M. (1994). Energetics of syntrophic fatty acid oxidation. *FEMS*
760 *Microbiol. Rev.* 15, 85–94.
761
- 762 Schloss, P. D., Westcott, S. L., Ryabin, T., Hall, J. R., Hartmann, M., Hollister, E. B.,
763 Lesniewski, R. A., Oakley, B. B., Parks, D. H., Robinson, C. J., Sahl, J. W., Stres,
764 B., Thallinger, G. G., Van Horn, D. J., and Weber, C. F. (2009). Introducing
765 mothur: open-source, platform-independent, community-supported software for
766 describing and comparing microbial communities. *Appl. Environ. Microbiol.* 75,
767 7537–7541.
768
- 769 Shen, Y., and Buick, R. (2004). The antiquity of microbial sulfate reduction. *Earth-Sci.*
770 *Rev.* 64, 243–272.

- 771
772 Stein, J. S., and Fisher, A. T. (2001). Multiple scales of hydrothermal circulation in
773 Middle Valley, northern Juan de Fuca Ridge: Physical constraints and geological
774 models. *J. Geophys. Res.* 106, 8563–8580.
775
776 Valentine, D. L. (2011). Fates of Methane in the Ocean. *Annual Rev. Mar. Sci.* 3, 147–
777 171.
- 778 Van Hamme, J., Singh, A., and Ward, O. (2003). Recent advances in petroleum
779 microbiology. *Microbiol. Mol. Biol. Rev.* 67, 503–509.
780
- 781 Von Damm, K. L., Oosting, S. E., Kozlowski, R., Buttermore, L. G., Colodner, D. C.,
782 Edmonds, H. N., Edmond, J. M., and Grebmeier, J. M. (1995). Evolution of East
783 Pacific Rise hydrothermal vent fluids following a volcanic eruption. *Nature* 375,
784 47–50.
785
- 786 Wang Z., Xu, Y., Zhao, J., Li, F., Gao, D., and Xing, B. (2011). Remediation of
787 petroleum contaminated soils through composting and rhizosphere degradation. *J.*
788 *Haz. Mat.* 190, 677–685.
789
- 790 Wankel, S. D., Adams, M. M. (co-first author), Johnston, D. T., Hansel, C. M., Joye, S.
791 B., and Girguis, P. R. (2012). Anaerobic methane oxidation in metalliferous
792 hydrothermal sediments: influence on carbon flux and decoupling from sulfate
793 reduction. *Environ. Microbiol.* 14, 2762–2740.
794
- 795 Wankel, S. D., Germanovich, L. N., Lilley, M. D., Genc, G., DiPerna, C. J., Bradley, A.
796 S., Olsen, E. J., and P. R. Girguis. (2011). Influence of subsurface biosphere on
797 geochemical fluxes from diffuse hydrothermal fluids. *Nature Geosci.* 4, 461–468.
798
- 799 Wegener, G., Niemann, H., Elvert, M., Hinrichs, K., and Boetius, A. (2008).
800 Assimilation of methane and inorganic carbon by microbial communities
801 mediating the anaerobic oxidation of methane. *Environ. Microbiol.* 10, 2287–
802 2298.
- 803 Wenger, L. M., Davis, C. L. and Isaksen, G. H. (2002). Multiple controls on petroleum
804 biodegradation and impact on oil quality. *Soc. Petrol. Eng. Reserv. Eval. Eng.* 5,
805 375–383.
806
807

808 Legends

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810 **Figure 1.** Time course of the anaerobic oxidation of short-chain alkanes (top),
811 consumption of sulfate (middle), and production of sulfide (bottom) in anaerobic batch
812 reactors of Middle Valley hydrothermal sediments incubated with methane (**A**), ethane
813 (**B**), propane (**C**), and butane (**D**) at three discrete temperatures (25, 55, and 75 °C). Each
814 time point represents the average of duplicate reactors with bars indicating the data range.
815 For sulfate and sulfide measurements, the control in each panel represents the average

816 sulfate consumption and sulfide production in the batch reactors with nitrogen for each
817 incubation temperature as described in the text.

818

819 **Figure 2.** Relative abundance (percentage) of bacteria determined from massively
820 parallel sequencing of DNA recovered from anaerobic batch reactor sediments incubated
821 with methane, ethane, propane, butane and nitrogen at 55 °C and pre-incubation ($T_{(0)}$)
822 sediments. Left and right side panels show the taxonomic breakdown of sequences at the
823 phylum and class level, respectively. Legend indicates operational taxonomic units
824 (OTUs), defined as sequences sharing 97 % nucleotide sequence identity.

825

826 **Figure 3.** Relative abundance (percentage) of archaea determined from massively
827 parallel sequencing of DNA recovered from anaerobic batch reactor sediments incubated
828 with methane, ethane, propane, butane and nitrogen at 55 °C and pre-incubation ($T_{(0)}$)
829 sediments. Left and right side panels show the taxonomic breakdown of sequences at the
830 class and order level, respectively. Legend indicates operational taxonomic units (OTUs),
831 defined as sequences sharing 97 % nucleotide sequence identity.

832

833 **Figure 4.** Maximum-likelihood phylogenetic tree illustrating the relationships of selected
834 16S rRNA *Deltaproteobacterial* sequences recovered from Middle Valley sediments to
835 *Deltaproteobacterial* sequences and uncultured environmental phylotypes from NCBI
836 non-redundant database. Representative sequences from Middle Valley sediments
837 incubated in batch reactors at 55 °C with ethane (Ethane55-MV), propane (Propane55-
838 MV), and butane (Butane55-MV) in bold. The most closely related environmental
839 phylotypes from Guaymas Basin sediment sequences have been marked (★). The tree
840 was rooted to *Archaeoglobus profundus* DSM 5631 (NR_074522). Scale = 0.01
841 substitutions per site.

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