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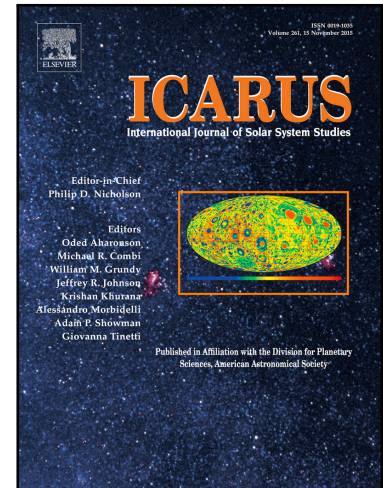
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Highlights

- Ten unique wind measurements within the Martian PBL during lander descents
- Wind measurements from displacement of Beagle 2 and Schiaparelli jettisoned hardware
- Information on the structure of the PBL from Viking landers, Phoenix and Curiosity
- Phoenix lander experienced unusually high winds during landing
- Wind direction measurements correlate well with Mars Climate Database

Measurement of Martian boundary layer winds by the displacement of jettisoned lander hardware

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Abstract

Martian boundary layer wind speed and direction measurements, from a variety of locations, seasons and times, are provided. For each lander sent to Mars over the last four decades a unique record of the winds blowing during their descent is preserved at each landing site. By comparing images acquired from orbiting spacecraft of the impact points of jettisoned hardware, such as heat shields and parachutes, to a trajectory model the winds can be measured. We start our investigations with the Viking lander 1 mission and end with Schiaparelli. In-between we extract wind measurements based on observations of the Beagle 2, Spirit, Opportunity, Phoenix and Curiosity landing sites.

With one exception the wind at each site during the lander's descent were found to be $<8 \text{ m s}^{-1}$. High speed winds were required to explain the displacement of jettisoned hardware at the Phoenix landing site. We found a tail wind ($>20 \text{ m s}^{-1}$), blowing from the north-west was required at a high altitude ($>2 \text{ km}$) together with a gust close to the surface ($<500 \text{ m}$ altitude) originating from the north. All in all our investigations yielded a total of ten unique wind measurements in the PBL. One each from the Viking landers and one each from Beagle 2, Spirit, Opportunity and Schiaparelli. Two wind measurements, one above about 1 km altitude and one below, were possible from observations of the Curiosity and Phoenix landing site.

Our findings are consistent with a turbulent PBL in the afternoon and calm PBL in the morning. When comparing our results to a GCM we found a good match in wind direction but not for wind speed. The information provided here makes available wind measurements previously unavailable to Mars atmosphere modellers and investigators.

Keywords: Mars, wind, trajectory, lander, PBL

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1 1. Introduction

2 Horizontal winds on Mars are an important property of the Martian atmo-
3 sphere as they can influence the weather (Savijärvi, 2012; Kass et al., 2008)
4 occasionally with dramatic results as observed during dust storms. The winds
5 also sculpt the surface and transport material around the planet (Day & Ko-
6 curek, 2016). The near-surface winds, i.e. less than 2 m altitude, have been
7 measured in situ by a number of surface missions (Chamberlain et al., 1976;
8 Seiff et al., 1997; Gunnlaugsson et al., 2008) however the winds within the PBL
9 have been measured only sporadically mostly by remote techniques e.g. Tam-
10 pari et al. (2010). More wind measurements are required especially those within
11 the PBL as they are valuable for verifying the physics of Martian atmospheric
12 models (Justus et al., 2004), understanding dust and volatile transport (Spiga
13 & Lewis, 2010), help with the detection of trace gases and for planning landings
14 (Michaels & Rafkin, 2008; Kauhanen et al., 2008).

15 Modelling of the Martian atmosphere on scales of 10 to 100 km (Spiga &
16 Lewis, 2010) and investigations of sand dunes (Parteli et al., 2009; Liu & Zimbel-
17 man, 2015; Jackson et al., 2015) suggest topographic control of the near-surface
18 winds. At smaller scales (1-10 km) the influence of turbulence on the wind pat-
19 terns becomes important. Dust devils, a significant form of turbulence, are seen
20 to be blown across the surface by the flow in the PBL (Stanzel et al., 2008).

21 In section 2 background on the images of the jettisoned hardware used in our
22 analysis is provided. Additionally some relevant atmospheric modelling details
23 are provided relevant to this study. In section 3 we describe our measurement
24 technique and how the measurement uncertainties are obtained. In section 4
25 we determine the wind speed, direction and some limited information on the
26 vertical structure of the Martian winds. In section 5 these results are discussed
27 in the context of the Martian PBL.

28 2. Background

29 Wind profiles in the PBL have been determined from landers on the parachute
30 descending through the Martian atmosphere using Doppler measurements such
31 as with the Viking landers (Seiff, 1993a) and Schiaparelli (Ferri et al., 2017). In-
32 strumentation has been developed specifically for making high resolution wind
33 speed and direction measurements up to an altitude of 10 km (Montmessin,
34 2014, 2017) using Doppler wind lidar and will hopefully be deployed on the
35 surface sometime soon. Moores et al. (2010, 2015) has used lander imagers to
36 investigate atmospheric dynamics over the landing site by tracking dust and
37 cloud features. Moores et al. (2016) modelled the trajectories of hardware jet-
38 tisoned by Curiosity to verify their mesoscale modelling results.

39 2.1. Wind-blown lander hardware

40 The distribution of lander hardware on the surface such as heat shields and
41 parachutes on the surface will be sensitive to the winds aloft in the Planetary
42 Boundary Layer (PBL), e.g. see Paton (2017). A way to determine wind speed,

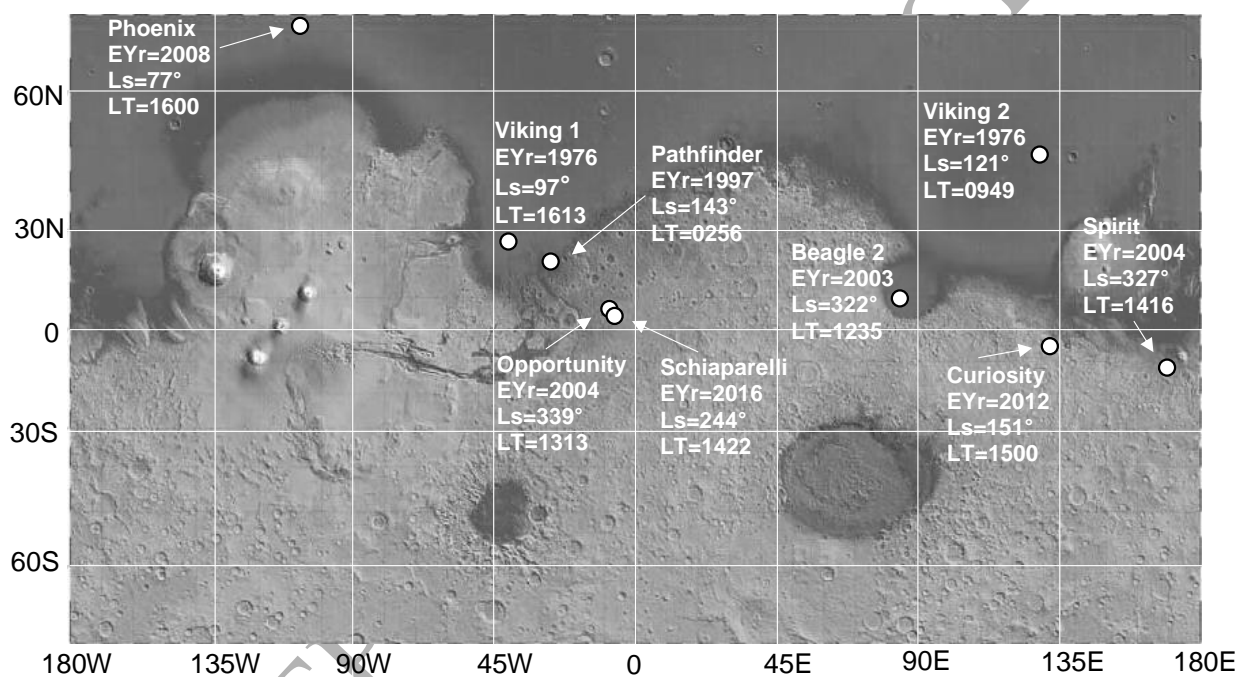


Figure 1. Lander locations on a topographical map of Mars (Smith et al., 2001). EYr is the Earth year that the lander arrived at the surface of Mars. Ls is the solar longitude. LT is the local time.

43 direction and limited information on the vertical structure on Mars is from
 44 trajectory modelling of jettisoned lander hardware Moores et al. (2016). As
 45 there have been a number of landings on Mars over a number of years and at
 46 a variety of locations, images of these sites could provide useful information
 47 on the variability of the Martian winds. Figure 1 shows the distribution of
 48 landers across Mars all of which, apart from VL-2 and Pathfinder, have landed
 49 in the afternoon when the atmosphere is most turbulent. In some cases strong
 50 turbulence (Seiff, 1993a) and gusts (Cheng et al., 2004) appear to have had a
 51 significant effect on the lander’s motion during descent.

52 Mars Global Surveyor (MGS) and Mars Reconnaissance Orbiter (MRO) have
 53 imaged landers along with their jettisoned EDLS components on the surface of
 54 Mars. The images are all available via the HiRISE or MOC websites. The landers
 55 that have been imaged are Vikings 1 and 2, Pathfinder, the MER rovers
 56 Spirit and Opportunity, Phoenix and Curiosity. Attempts to image failed mis-
 57 sions have been attempted. So far ESA’s Beagle 2 (Merrifield, 2015) and Schi-
 58 aparelli (see HiRISE website) have been spotted.

59 As can be seen in Table 1 the identification of jettisoned lander hardware
 60 are more certain when imaged up to a year or two after landing, i.e. from Spirit
 61 onwards in table 1. Older lander parts are presumably covered in Martian
 62 dust which makes spotting them difficult. For example the Viking lander 1
 63 and 2 backshells have been observed but the parachutes have proved difficult
 64 to identify presumably because they are covered in dust. In some more recent
 65 cases EDL hardware has been imaged in situ by the lander. The HiRISE images
 66 of the Phoenix hardware, its heat shield and parachute are particularly clear
 67 and the certainty of their origin has been strengthened with images returned
 68 from the surface by the lander itself (image PIA11172).

Table 1: Lander hardware imaged on the surface of Mars. In the second column are listed the dates of Mars landings. The third column lists the image file names used for our analysis in this paper. Column four lists the dates when the lander site images were acquired. The fifth column, headed with ‘Ca’, lists the camera used. Hi stands for HiRISE aboard MRO and M stands for MOC aboard MGS. In the last three columns the headings of HS, P and B stand for Heat Shield, Parachute and Backshell respectively. In these column Y indicates a positive identification, N indicates no identification, m indicates possible identification of hardware. All the landers themselves were identified in all the images used.

Mission	Landed	Image ID	Acquired	Cam	HS	P	B
VL-1	20/07/76	PSP_001521.2025	22/11/06	Hi	m	m	Y
VL-2	03/09/76	PSP_001501.2280	21/11/06	Hi	m	N	Y
Pathfinder	04/07/97	PSP_001890.1995	21/12/06	Hi	m	Y	Y
Beagle 2	25/12/03	ESP_039308.1915	15/12/14	Hi	Y	m	Y
Spirit	04/01/04	PIA05248	19/01/04	M	Y	Y	Y
Opportunity	25/01/04	PIA05230	09/02/04	M	Y	Y	Y
Phoenix	25/05/08	PSP_08591.2485	26/05/08	Hi	Y	Y	Y
Curiosity	06/08/12	ESP_028401.1755	17/08/12	Hi	Y	Y	Y
Schiaparelli	19/10/16	ESP_048041.1780	25/10/16	Hi	Y	Y	Y

69 Figure 2 shows a generic EDL architecture representing a lander-type using
70 powered descent. Similar lander architectures were used for the Viking landers,
71 Phoenix and Curiosity. For Viking the parachute was deployed at around 5
72 km above the landing site (Cooley & Lewis, 1977), 13 km altitude for Phoenix
73 (Desai et al., 2011) and 12 km altitude for Curiosity (Cruz et al., 2014). The heat
74 shield is jettisoned soon after the parachute is deployed, triggered by a timer,
75 and then falls relatively quickly to the surface impacting the surface before the
76 touchdown of the lander. The next relevant event for our analysis is when the
77 lander separates from the backshell-parachute. This normally occurs around an
78 altitude of 1 km. After being released the lander performs a powered descent
79 while the parachute and the connected backshell drift down to the surface. Other
80 types of landing system have been used on Mars, i.e. Pathfinder, MER rovers.
81 These differ in that they used a combination of solid propellant retrorockets and
82 airbags for the final stage of the landing to bring the lander to rest. Beagle 2
83 was light enough that retrorockets were not required and airbags could be used
84 to absorb the energy from the impact with the surface. See table A5 and A6
85 for lander EDLS properties and EDL trajectory parameters respectively.

86 Figure 3 compares the distribution of components from the landers imaged
87 by HiRISE. The distances between the components were obtained from the
88 HiRISE and MOC images. The direction of approach of the landers to the
89 landing site has been normalised to an approach from the west. This then
90 shows, that even with winds blowing, the heat shield always overshoots the
91 landing site. The parachutes, being more sensitive to the winds, are dispersed
92 in a more random fashion.

93 2.2. Mars atmosphere models

94 Climate, mesoscale and microscale models are used independently and to-
95 gether to study the Martian atmosphere at different scales. As well as 3D
96 models, 2D and 1D column models (Savijärvi, 1995, 1999; Savijärvi et al., 2004,
97 2005; Savijärvi & Kauhanen, 2008; Savijärvi & Määttänen, 2010) are used for
98 studying the atmosphere of Mars.

99 General Circulation Models (GCMs) are tools used to understand and pre-
100 dict the state of the Martian climate. As they are global in scope GCMs gener-
101 ally are restricted, due to computing resources, to a cell size of the order of 100
102 km in size. In these cells the state of the atmosphere is calculated, i.e. pressure,
103 temperature and winds. Processes that operate on smaller scale than the cell
104 size, such as convection, are parameterised so their contribution can be included
105 in the calculations making the results more accurate.

106 The Mars Climate Database (MCD) is a database (Millour et al., 2009; Mil-
107 lour et al., 2015) containing results from simulations run with the Laboratoire de
108 Météorologie Dynamique (LMD) GCM. Included in the database are a number
109 of scenarios that take into account two most variable forcings of the atmosphere
110 which are due to UV input from the Sun and the amount and distribution of
111 dust. The default "standard atmosphere" setting for the online web interface,
112 which we used, contains results from a GCM run using dust distributions re-
113 constructed from five Martian years of observations. The solar input for this

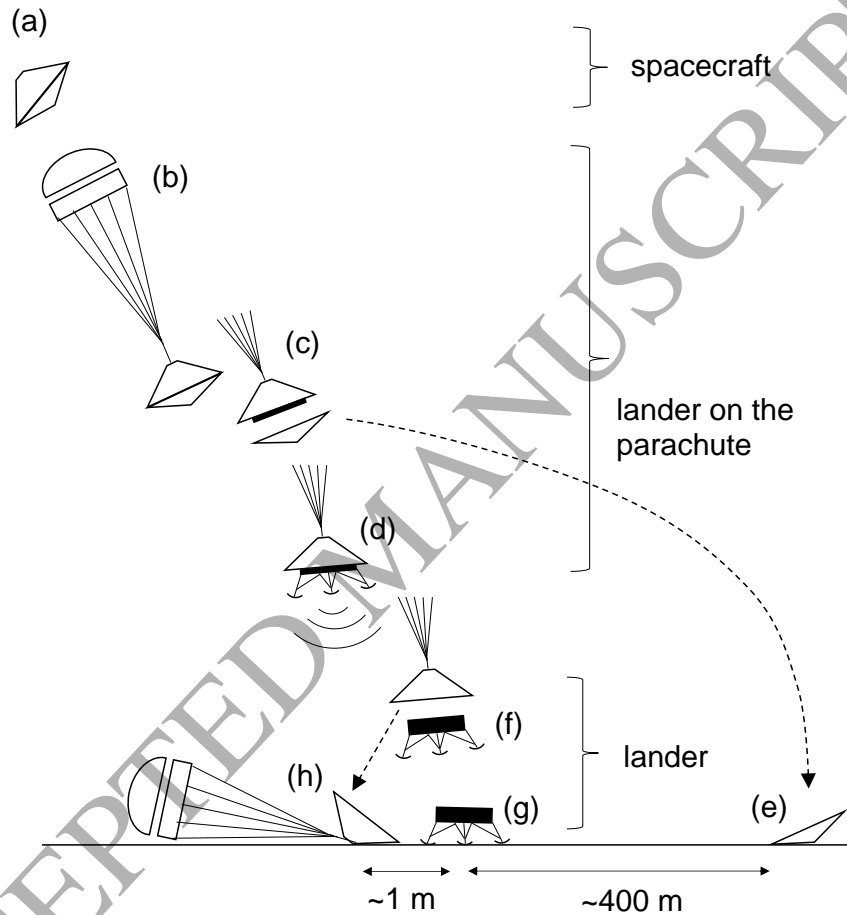
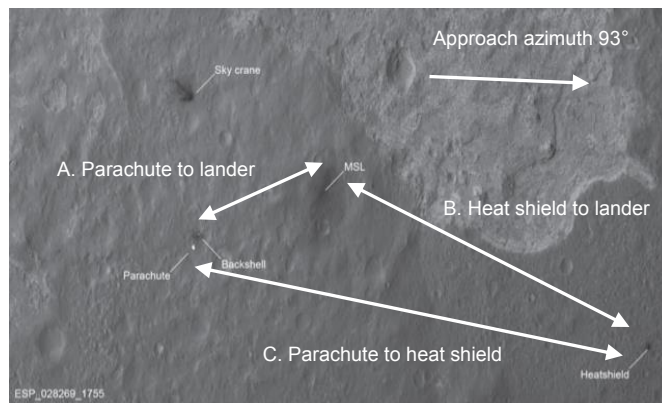


Figure 2: A landing architecture showing the sequence of events to reconfigure the lander during the descent and landing. The sequence are as follows (a) entry ~ 10 km altitude, (b) parachute deployment at ~ 10 km altitude, (c) heat shield jettison a few seconds after the parachute deployment, (d) leg deployment and radar activation, (e) impact of the heat shield on the surface (f) jettison of the backshell and parachute at 1-2 km altitude (g) touchdown on the surface (h) impact of backshell and parachute on the surface. The terminology shown for lander configurations used in discussion. (a) is the 'spacecraft', (b)-(d) is the 'lander on the parachute' and (f)-(g) is 'lander'.

(a) MSL and its EDLS hardware



(b) Distribution of hardware from Mars landers

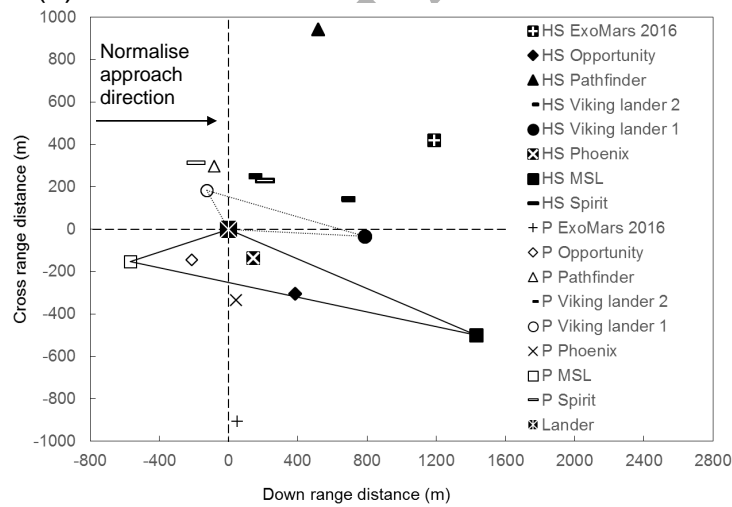


Figure 3: Locations on Mars of landed hardware from various missions. In (a) the EDL hardware at the landing site in Gale crater is shown. The distances used in section 3 The hardware for each mission is shown relative to the lander's touch down point. The dispersion of each mission's hardware on the surface is represented by the symbols in the legend. For VL-1 and Curiosity straight dotted and solid lines respectively are drawn between items of hardware to form triangles.

114 setting is an average value. In our work we use this setting as a comparison
 115 with our derived values of wind as it represents an average climate. We also
 116 use other scenarios, or setting, and these include a cold atmosphere with a low
 117 amount of dust present and a warm atmosphere which is dusty.

118 3. Method

119 3.1. Wind measurement procedure

120 Two steps were required to obtain the wind measurements. First the dis-
 121 tance between hardware components were obtained from the images. Then the
 122 trajectory model was run, varying the wind speeds and directions, until a match
 123 was found with the images.

124 The hardware component distances were calculated from JPEG2000 images
 125 (McEwen, 2007) available on the HiRISE website. The image viewing applica-
 126 tion HiView was used for this purpose. Table 2 lists the distance and orientation
 127 between the selected hardware components used to analyse the winds. These
 128 hardware components selected for analysis were based on those clearly identified
 129 in the images. Errors in measuring the distance between the hardware compo-
 130 nents was estimated to be at most one metre which translates to an uncertainty
 131 of about 2% when determining the wind speed.

Table 2: List of hardware components used for determining wind properties at the landing sites. For Beagle 2 'Pilot' and 'Ringsail' is shorthand for the pilot and ringsail parachutes respectively. For Spirit and Opportunity 'First bounce' refers to the first airbag impact on the surface. Uncertainty on the distance measurements are estimated to be of the order of one metre. The heading is an angle measured clockwise from the direction of the north pole. It is used to define the orientation of an object's path on the planet's surface, at any given point along its trajectory. The direction of travel is from the object's current surface location along the path defined by the heading. The heading is equivalent to azimuth as used in celestial mechanics. Note the meteorological convention for wind direction is in the opposite direction, i.e. the wind blows from the direction as defined by the azimuth angle.

Mission	From	To	Distance (m)	Heading (°)
Viking Lander 1	Lander	Backshell	245	282
Viking Lander 2	Lander	Backshell	422	342
Beagle 2	Pilot	Ringsail	106	323
Spirit	First bounce	Heat shield	373	72
Opportunity	First bounce	Heat shield	653	109
Phoenix	Lander	Backshell	331	173
Phoenix	Lander	Heat shield	195	45
Curiosity	Lander	Heat shield	1486	252
Curiosity	Lander	Backshell	569	115
Schiaparelli	Backshell	Heat shield	1773	40

132 Once the separation distances had been measured a Hill Climbing algorithm
 133 was used to control the fitting procedure. The algorithm automatically sampled
 134 a Gaussian probability distribution of the wind speeds, ran the trajectory model

135 with these wind speeds and compared the results from the model with the
 136 impact locations observed in an image. A single or double wind layer was used
 137 by the trajectory model depending on the information available, i.e. number of
 138 hardware components available. One layer was used if either the lander and the
 139 parachute or the lander and the heat shield had been identified in the images.
 140 Double layers were possible if both the parachute and heat shield were identified
 141 together with the lander. Each layer normally had constant wind speed and
 142 direction. More elaborate functions could in principle be used to represent the
 143 winds in the model (see section 3.2).

144 The Hill Climbing algorithm worked as follows. The known hardware and
 145 trajectory properties, listed in tables A5 and A6, were used first to initialise
 146 the model. Estimates for the u and v wind speeds could be provided, if known,
 147 to initialise the model and speed up the search otherwise the wind speed was
 148 set to zero. After the model had been initialised with the initial wind speeds
 149 and run once a preliminary fit parameter value was calculated comparing the
 150 impact locations in the model to those in the image.

151 The Hill Climbing algorithm then generated new wind speed estimates by
 152 sampling probability distributions, one distribution for each u and v wind speed
 153 component, centred around the initial estimates of the wind speeds. The model
 154 was then run for a second time and a new fit parameter calculated. If the new
 155 fit parameter improved on the previous one the algorithm then switched to sam-
 156 pling probability distributions centred around the latest wind speed estimates.
 157 If there was no improvement then the algorithm kept sampling the previous
 158 probability distributions centred on the old wind speed estimates. The process
 159 was repeated until the fit parameter reached a predetermined value.

160 Figure 4 shows an example of convergence towards a wind measurement
 161 using the Monte Carlo based Hill Climbing algorithm. The wind profile in this
 162 case has u and v components both equal to 5 m s^{-1} above 1000 m. Below this
 163 altitude the u and v components have values of 10 m s^{-1} . A best fit parameter
 164 is used by the hill climbing algorithm to search for a wind speed and direction.
 165 The best fit parameter for a double wind layer, i.e. if the parachute and heat
 166 shield have been identified, is as follows.

$$d_{bf} = ([x1_{hs} - x2_{hs}]^2 + [y1_{hs} - y2_{hs}]^2)^{0.5} + ([x1_{pc} - x2_{pc}]^2 + [y1_{pc} - y2_{pc}]^2)^{0.5} \quad (1)$$

167 where, for the model output, $x1_{hs}$ is the longitudinal distance between the
 168 lander and the heat shield, $y1_{hs}$ is that latitudinal distance between the lander
 169 and the heat shield, $x1_{pc}$ is the longitudinal distance between the lander and
 170 the parachute and $y1_{pc}$ is the latitudinal distance between the lander and the
 171 parachute. The variables followed by a '2' denote that these distances are from
 172 the images of the landing sites. A single wind layer would require at least one
 173 jettisoned hardware item to have been identified in the image.

174 The Hill Climbing algorithm was normally halted when equation 1 was less
 175 than one. This provided more than enough precision when calculating the wind
 176 speed, i.e. $<0.01 \text{ m s}^{-1}$ when tested against a reference wind profile with con-
 177 stant wind speed.

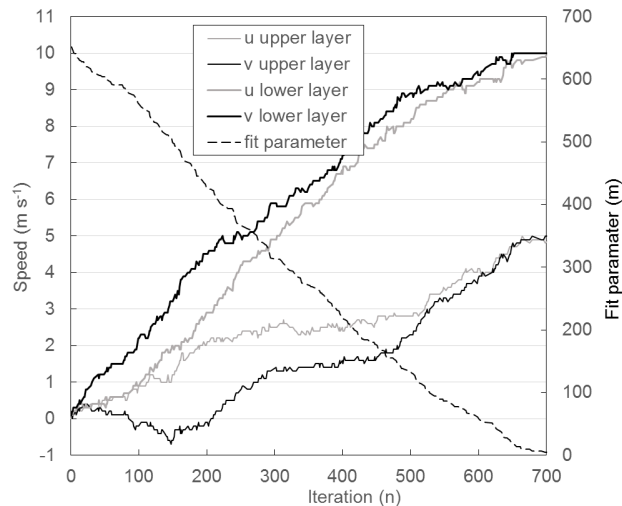


Figure 4: Progression of the fit, using a hill climbing algorithm, for a descending Phoenix-type lander that experiences winds.

178 Figure 5 shows an example of our approach applied to a notional reference
 179 Martian wind profile. This wind profile represents a typical wind profile on
 180 Mars that includes some variability with altitude. The wind blows from the
 181 west. Running the model with the reference wind profile produces a specific
 182 arrangement of impact points at the surface. These reference impact points can
 183 then be used to determine the wind speed and direction for one wind layer, if just
 184 the heat shield is used, or two layers if the heat shield and parachute are used.

185 Notice how the single layer solution, in figure 5, is not some average of the
 186 winds acting on the heat shield above and below the parachute jettison altitude.
 187 For the single layer, with a wind speed of 10 m s^{-1} one may instinctively want
 188 to reduce the wind speed to obtain a fit. However this would not work because,
 189 above the parachute release altitude, a decrease in wind speed will increase
 190 the separation distance overall. See appendix section C for an example of this
 191 behaviour and its explanation.

192 In the model the winds are represented as u and v components as shown in
 193 figures 4 and 5. The results in the paper are presented using the meteorology
 194 convention for wind speed and direction. For example the values in figure 4, for
 195 the upper layer, translate into a wind speed of 7 m s^{-1} with a wind originating
 196 from an azimuth of 225° , i.e. blowing from the south-west. Note the convention
 197 used for calculating the location of hardware relative to the lander and for the
 198 heading of the lander in flight is in the opposite sense, i.e. towards rather than
 199 from the azimuth.

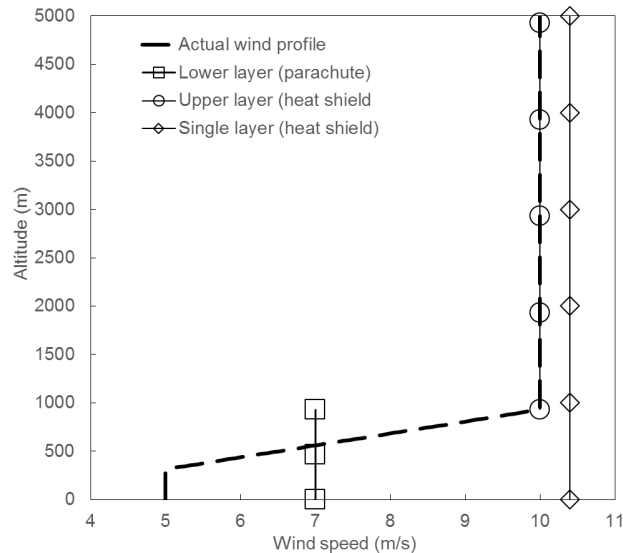


Figure 5: Wind layers as used in the results section fitted to a test profile. See the text for an explanation of the single and double wind layer solutions.

200 3.2. Fitting realistic wind profiles

201 In reality wind speed and direction, at any particular time on Mars, will likely
 202 vary with altitude (Petrosyan et al., 2011). An alternative approach to using
 203 wind layers with constant wind speed and direction is to vary these properties
 204 as some function of altitude in the model, e.g. using a polynomial or logarithmic
 205 function. This approach can then help with the interpretation with the results
 206 and identify possible trends.

207 As a minimum for an atmospheric layer of interest we use two linear functions
 208 connected at a mid-point to approximate a continuous profile and capture any
 209 swings in wind speed and direction. The formulation of the wind profile as a
 210 Piece-Wise Linear Function (PWLF) with two sections is as follows:

$$u = u_{top} + p_f \frac{1 - \frac{z - z_{bot}}{z_{top} - z_{bot}}}{1 - f_{mid}} \quad (2)$$

$$u = \frac{z - z_{bot}}{z_{top} - z_{bot}} \frac{u_{top} - u_{bot} + p_f}{f_{mid}} + u_{bot} \quad (3)$$

211 where u_{top} is the wind velocity at the top of the layer, p_f is a shape param-
 212 eter, z is the altitude, z_{top} is the altitude of the upper boundary, f_{mid} is the
 213 location of the point of the split in the layer and u_{bot} is the wind velocity at
 214 the bottom of the layer. If the mid-point is fixed then the only variable that is
 215 the shape parameter which makes it easy to incorporate into our Hill Climbing
 216 algorithm.

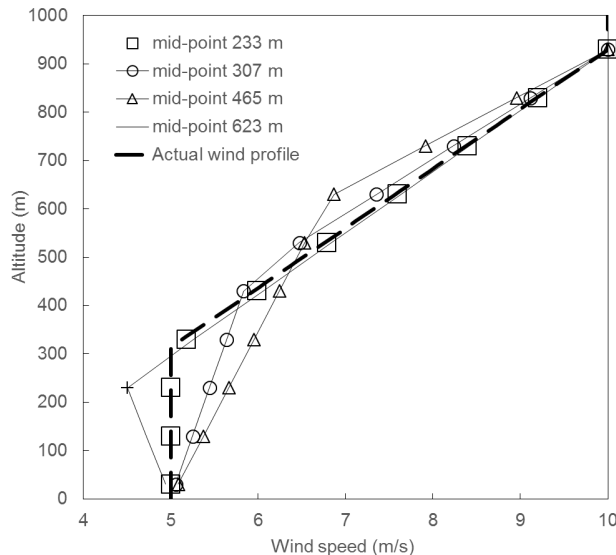


Figure 6: Reconstruction of the wind profile using a Piece-Wise Linear Function (PWLf). The point where each PWLF is connected is indicated in the legend.

217 Figure 6 shows the results of fitting a PWLF to a wind speed profile that
 218 varies in intensity with altitude. The mid-point for the fitted profiles are changed
 219 for each fitting attempt. As can be seen the algorithm can produce an exact
 220 solution for the wind profile if the mid-point is known. When using other points
 221 the profile is less accurate but the general trend can be discerned and aids
 222 interpretation when compared to GCM results or combined with previous wind
 223 measurements.

224 3.3. Measurement uncertainty

225 The vehicle's initial state, its aerodynamic properties during the descent and
 226 the jettisoned hardware aerodynamic properties are uncertain to one extent or
 227 another. If our trajectory model is run numerous times for a specific mission,
 228 each time with the same wind profile but sampling various probability distribu-
 229 tions of the initial state values or vehicle properties, a probability distribution
 230 of landing locations are produced, i.e. a landing ellipse. In our case we initialise
 231 the trajectory models from the parachute deployment altitude not from entry
 232 as this would introduce unnecessary complexity in the modelling.

233 To shrink the landing ellipses to a single point, i.e. the impact location of
 234 jettisoned hardware observed in an image, a range of wind profiles are required.
 235 These wind profiles, calculated by the model, then represent our wind measure-
 236 ment. Note the landing ellipses or uncertainties, apart for Beagle 2, use more
 237 tightly constrained initial conditions and vehicle properties than calculated be-
 238 fore entry as they use reconstructed values. See table A5 and A6 for trajectory

239 and vehicle parameters used in the model and table A7 for input uncertainties.

240 Table 3 lists the 3σ uncertainty on the separation distances between the heat
 241 shield and the lander for the Phoenix lander. See table A7 and the text in the
 242 appendix for the uncertainty on the model inputs for each mission. For the
 243 Phoenix case each model parameter was set their high and low values in turn,
 244 running the model with no winds, and noting the impact locations. As can be
 245 seen the uncertainty on the drag coefficient potentially accounts for more than
 246 half of the uncertainty (columns 3 and 4 in table 3) on the hardware locations.

Table 3: Sensitivity of hardware separation distance on initial conditions. Separation between the lander and heat shield without any uncertainty is 409 m. The error estimates for the initial state of the lander on the parachute are based on the 1σ values listed in the dataset of properties derived from IMU measurements (PHX-M-ASE-5-EDL-RDR-V1.0). The uncertainty on the drag coefficient is typical for parachutes deployed for Mars, e.g. see Cruz et al. (2003).

Model input	Input uncertainty	Heat shield displacement uncertainty (m)
α ($^\circ$)	+1 -1	+9 -9
v (m s^{-1})	+12 -12	+1 -1
z (m)	+700 -700	+2 -2
ρ (kg m^{-3})	+5% -5%	-8 +14
C_{Dp}	+6% -6%	-67 +69
C_{Dhs}	+6% -6%	+51 -54

247 The uncertainties on wind speed and direction are shown in figure 7 for
 248 the Phoenix case for a single wind layer from the surface to the parachute de-
 249 ployment altitude. These values were calculated as follows. A reference case
 250 was run with wind speed and direction set to some non-zero amount and the
 251 vehicle properties all set to their nominal values. This generated a set of refer-
 252 ence impact points. Next the vehicle's properties were set to their high and
 253 low 3σ values in turn and the search algorithm run to find the corresponding
 254 uncertainties on the winds speeds.

255 Figure 7 shows the dependence of the wind speed and direction uncertainties
 256 on wind speed for various model input uncertainties. Figure 7 (a) and (b) shows
 257 the wind uncertainties due to the uncertainty on the heat shield drag coefficient.
 258 In figure 7 (a) the wind speed uncertainty decreases with increasing speed. In
 259 figure 7 (b) the wind direction uncertainty totals 180° at a wind speed of 0 m s^{-1} .
 260 The directional uncertainty reduces to only a few degrees at 30 m s^{-1} . The
 261 large directional uncertainty at low wind speed is related to the displacement
 262 of a heat shield if its drag coefficient is changed. This displacement can be
 263 larger than the distance displaced by low speed winds. See Appendix D for
 264 a detailed discussion on how uncertainty in heat shield drag coefficient can

265 produce contradictory direction information.

266 Figure 7 (c) shows the uncertainty on wind speed due to other model input
 267 uncertainties. The uncertainties on initial altitude, velocity and flight path angle
 268 do not produce any uncertainty on wind direction so a figure demonstrating this
 269 is not shown. Figure 7 (d) shows the uncertainty on the wind speed determined
 270 from the jettisoned parachute due to uncertainty on its drag coefficient. The
 271 uncertainty on the initial conditions, parachute deployment condition, do not in
 272 general affect the displacement of the jettisoned parachute relative to the lander.
 273 By the time the parachute is jettisoned the lander has reached equilibrium with
 274 the atmosphere, i.e. it is descending vertically relative to the winds.

275 4. Low altitude (<12 km) winds

276 The wind profiles in this section consist of single layer constant value wind
 277 profiles for Viking lander 1 and 2, Beagle 2, Spirit and Opportunity. Double
 278 layer wind profiles, each layer with constant values, are included in this section
 279 for Phoenix and Curiosity. The wind profiles are derived from the technique
 280 described in subsections 3.1. and 3.2 Table A.5 and A.6 in the appendix contain
 281 information on the lander properties and details of the EDL events

282 4.1. Viking landers

283 The Viking Landers descended through the Martian atmosphere using an
 284 aeroshell fitted with a heat shield to protect the landers from the extreme ther-
 285 mal environment during the initial high speed entry. The aeroshell then slowed
 286 the lander down to a speed of just above Mach 1 (Cooley & Lewis, 1977) at
 287 an altitude of about 5.9 km above the landing site allowing a parachute to be
 288 deployed. A timer was used to trigger the release of the heat shield at an alti-
 289 tude of 4 km seven seconds after parachute deployment. At 1.5 km above the
 290 surface the backshell while still connected to the parachute, was jettisoned and
 291 the lander performed a powered descent to the surface using its rocket engines.

292 Only the location of the backshell relative to the lander was used for the
 293 analysis. The heat shield in the published HiRISE image has not been positively
 294 identified. For comparison, data from the MCD for different climate scenarios
 295 and the near-surface winds measured by the Viking landers over subsequent sols
 296 after the landing (Hess et al., 1977).

297 Seiff (1993a) have determined the winds between 5.5 and 1.5 km altitude
 298 for the Viking landers while they descended on the parachute. These winds are
 299 presented and discussed in section 5 together with our results.

300 For VL-1 the application of our method described in subsections 3.1 and 3.2
 301 yields a wind originating from south-east as shown in figure 7 (a) and a wind
 302 speed of about 3 m s^{-1} as shown in figure 7 (b). For VL-2 a wind blowing from
 303 the south and at a speed of about 3 m s^{-1} is determined using our method.
 304 The wind profiles determined from EDLS trajectory modelling appears to be in
 305 general agreement with the MCD wind profiles for a low dust and minimum solar
 306 activity climate scenario. They do not agree very well with wind measurements
 307 made by Seiff (1993a) an issue which is explored further in section 5.

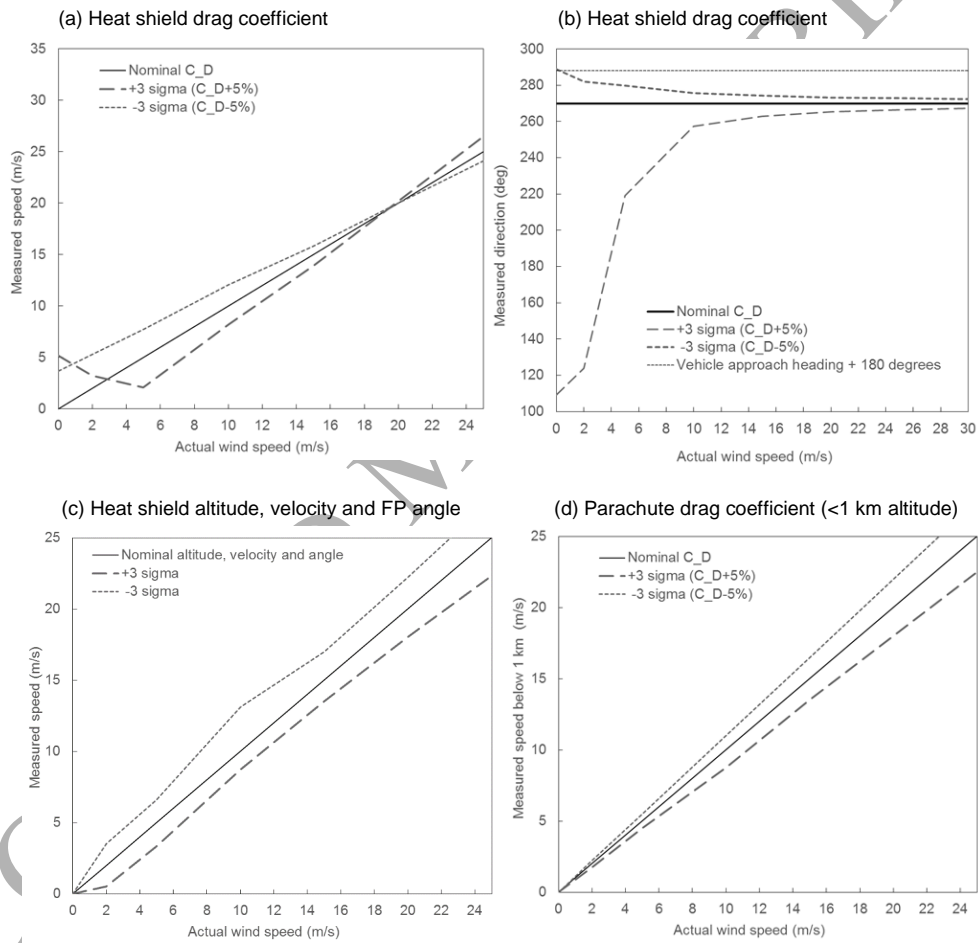


Figure 7: The uncertainty on the wind speed and direction measurements.

308 4.2. *Pathfinder, Beagle 2, Spirit and Opportunity*

309 During the final stage of landing Pathfinder (Spencer et al., 1999), Beagle 2
310 (Taylor, 2000), Spirit and Opportunity landers (Cheng et al., 2004) used landing
311 bags to bounce across the surface and absorb the impact velocity. For Spirit and
312 Opportunity images of the landing sites show the first impact of the air bags
313 shortly after the parachute had been jettisoned. Therefore it is straightforward
314 to determine the wind speed and direction from this impact point and the heat
315 shield impact point. For Pathfinder images of the landing site do not seem to
316 show the impact point of its airbags. A wind profile is therefore not provided for
317 Pathfinder. For Beagle 2, which did not use Rocket Assisted Descent (RAD),
318 the ring sail parachute (Fallon & Sinclair, 2003) would presumably impact the
319 surface close to where the airbags impacted the surface. This would make the
320 ring sail parachute imaged in the HiRISE image suitable for analysis together
321 with the backshell also in the image. A full list summarising the hardware
322 components and images used can be found in tables 1 and 2.

323 The wind speed and direction during the Beagle 2 landing was determined
324 to be 4 m s^{-1} blowing from the north-west. This result is representative of the
325 wind speed below 7.1 km after the first parachute had been deployed. The first
326 parachute was a relatively large pilot parachute aiding the deployment of the
327 main ringsail parachute by slowing the lander down from supersonic to subsonic
328 speed. See tables A.5 and A.6 for further details regarding Beagle 2's two-stage
329 parachute system.

330 For Spirit it was found that the it experienced a wind speed of 5 m s^{-1}
331 blowing from the west. For Opportunity, see figures 8 (e) and (f), the wind
332 speed was 8 m s^{-1} blowing approximately from the south.

333 4.3. *Phoenix*

334 The Phoenix lander used a similar EDLS to the Viking landers including an
335 aeroshell, parachute and powered descend for a soft touch down on the surface
336 (NASA, 2008; Witkowski et al., 2009; Szalai et al., 2011). The altitude of
337 parachute deployment for Phoenix was over twice the altitude of the Viking
338 landers (see table A.5). The jettisoning of the parachute and backshell, on the
339 other hand, was significantly lower at 930 m rather than 1.5 km for the Viking
340 landers. The hardware components from the Phoenix landing, i.e. the heat
341 shield, parachute and lander itself, are very clearly distinguished in the HiRISE
342 images. The heat shield was also imaged from the lander when on the surface.
343 With the heat shield and parachute available for the analysis it is possible to
344 obtain some information regarding the structure of the wind profile.

345 In the trajectory model the atmosphere of interest was divided into two
346 layers, with differing wind properties for each layer. The upper layer started
347 at the parachute deployment altitude of 13.3 km and ended at the parachute
348 jettison altitude of 930 m. The lower layer started at the parachute jettison
349 altitude and ended at the surface. The boundary between the layers was placed
350 at the parachute jettison altitude because the separation distance between the
351 parachute and lander is sensitive to winds below this altitude.

352 The wind speed and direction required to blow the hardware components to
 353 their impact locations are as follows. For the upper layer the wind speed is 21
 354 m s^{-1} and the direction is 316° , i.e. a north-westerly wind. For the lowest layer
 355 the wind speed is 10 m s^{-1} and the direction 360° , i.e. a northerly wind.

356 4.4. Curiosity

357 The Curiosity rover used a similar landing architecture to Viking and Phoenix
 358 in that parachutes and powered descent were employed to enable a soft landing.
 359 It was different in that the rover was delivered to the surface via a skycrane
 360 which hovered as it lowered the rover down on a tether (Martin et al., 2015).
 361 After the rover touched down the skycrane was detached and flew away to a safe
 362 distance. Another difference was that Curiosity performed a backshell avoidance
 363 manoeuvre (Martin et al., 2015) travelling about 300 m towards the north-east
 364 after jettisoning the backshell and parachute. Like Phoenix the hardware compo-
 365 nents in the HiRISE image are clearly identified on the surface allowing the
 366 determination of the atmosphere's structure at the time of landing (Fig. 3a).

367 To model the wind for the descent of Curiosity the atmosphere was divided
 368 into two layers. The upper layer started at parachute deployment (12.1 km) and
 369 ended with the jettison of the backshell and parachute at 1.7 km altitude. See
 370 table A.5 for more information on lander EDL events. The lower layer started
 371 from the backshell and parachute jettison altitude down to the surface. For the
 372 upper level the wind speed was calculated to be 4 m s^{-1} blowing from 200° , i.e.
 373 an south-westerly wind. For the lower level the wind speed was calculated to
 374 be 6 m s^{-1} blowing from 100° , i.e. a south-easterly wind.

375 4.5. Schiaparelli

376 Schiaparelli performed a nominal entry using its aeroshell and deployed its
 377 parachute at the correct altitude. It was then due to make a powered descent to
 378 the surface. The lander jettisoned its parachute at an altitude of 3.7 km (Ferri
 379 et al., 2017) which was higher than expected. The lander did not perform its
 380 powered descent correctly, the rockets only firing for a few seconds, resulting in a
 381 high speed impact with the surface and the end of the mission. For the analysis
 382 of Schiaparelli only the parachute and heat shield were used. The results are
 383 shown in figure 9 (e) and (f). The results are quite close to the MCD with a
 384 wind blowing at 5 m s^{-1} from the north east.

385 4.6. Key results collected together

386 Table 4 presents our wind measurements as a table together with the lati-
 387 tude, longitude, local time and solar longitude at the time of landing. Single
 388 wind layers with constant wind direction and speed were used for all landers
 389 except Phoenix and Curiosity where two wind layers, each with constant wind
 390 and speed were used.

391 Figure 8 shows the wind direction and speed for those winds aloft in the
 392 PBL as determined from trajectory modelling of the EDLS hardware compared

Table 4: Wind direction and speed. ADD MORE INFO.

Mission	alt. (km)	dir. (°)	speed (m s ⁻¹)	lat. (°)	long. (°)	LT	Ls (°)
Viking lander 1	0-1	129	3.0	22.5	312	1618	97
Viking lander 2	0-1	161	2.2	48	134	0949	121
Beagle 2	0-7	300	3.9	11.5	90	1235	322
Spirit	0-7	265	4.9	-14.6	175	1416	327
Opportunity	0-7	184	8.2	-2	354	1313	339
Phoenix	0-1	360	9.9	62	234	1600	77
Phoenix	1-12	316	21.2	62	234	1600	77
Curiosity	0-2	117	5.1	-5	137	1500	151
Curiosity	2-12	167	4.3	-5	137	1500	151
Schiaparelli	0-7	11	5.0	-2	354	1422	244

393 to the "standard atmosphere" obtained from the MCD. Some details regarding
 394 the MCD and its settings can be found in subsection 2.2.

395 In figure 8 it can be seen that the wind speed does not appear to agree
 396 very well with the MCD ($R^2=0.34$) while the wind direction agrees quite well
 397 ($R^2=0.86$). The EDLS derived wind speeds in figure 8 (a) appear to be weakly
 398 correlated with the MCD results. This could be at least partly due to the large
 399 uncertainty on knowing the aerodynamic properties of the jettisoned EDLS com-
 400 ponents. For about half of the cases the MCD appears to be in clear disagree-
 401 ment with the EDLS derived wind speeds and their associated uncertainties.
 402 The most notable is the wind speed for Phoenix between an altitude of 1 and
 403 12 km.

404 In figure 8 (b) the wind direction provided from the MCD appear to be in
 405 better agreement with the EDLS derived wind direction. Phoenix, represented
 406 by the circle and diamond near the top of the chart, has clearly different wind
 407 speeds to the MCD.

408 5. Winds in the PBL

409 In this section we discuss selected missions, i.e. VL-1, VL-2 and Phoenix,
 410 that allow us to probe the winds in the Martian PBL. The displacement of the
 411 Viking lander parachutes allowed us to measure the wind speed and direction
 412 below an altitude of 1.5 km. Coupled together with previously published wind
 413 profiles from 5.5 to 1.5 km this allows us to probe the wind conditions at these
 414 sights in some detail. For Phoenix the wind speeds are high compared to average
 415 PBL wind speeds suggesting some atmospheric disturbance .

416 5.1. The Viking landers

417 Wind profiles have been determined (Seiff, 1993a) from measurements made
 418 by the Viking landers as they descended by parachute (see figure 9) from an

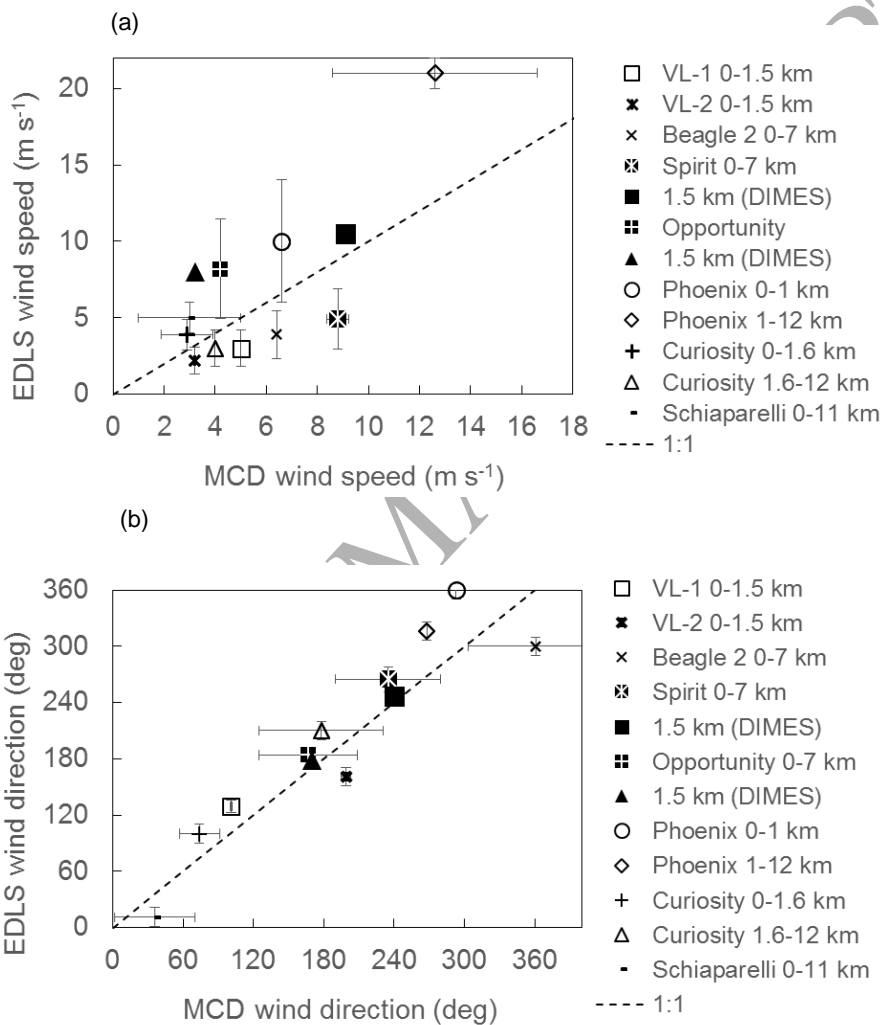


Figure 8: Wind speed and direction from EDLS and MCD compared. Vertical error bars take into account the uncertainty on the aerodynamic properties. The horizontal error bars represents the range of wind properties over the descent.

419 altitude of 5.5 to 1.5 km. We used two PWLFs in our model, one for the u
 420 wind speed component and another for the v wind component, to represent the
 421 winds below 1.5 km altitude. A Hill Climbing algorithm (see section 2.2) was
 422 used to vary the PWLF shape parameters until a fit between the model and
 423 observed location of the parachute was found. The PWLFs were bounded by
 424 wind measurements at 1.5 km altitude and at the surface (Seiff, 1993a). The
 425 altitude of the PWLF mid-point was set manually to three different heights to
 426 provide three pairs of u and v PWLFs. These are converted into wind speed
 427 and direction in figure 9.

428 For VL-1 Seiff (1993a) suggested that high wind speeds during the parachute
 429 descent, see figure 9 (b), was likely caused by cold air advection from the south-
 430 east. Figures 9 (a) and (b) possibly support this interpretation, i.e. south-
 431 easterly wind. However in figure 9 (a) the wind swings around so it is blowing
 432 from the north quite soon after the jettison of the parachute. This is different
 433 from the MCD and perhaps suggests a more local disturbance in the atmosphere
 434 below an altitude of 1.5 km.

435 Seiff (1993a); Haberle et al. (1993) suggested local topography modified the
 436 large-scale slope wind as revealed by the winds measured for VL-2, e.g. see
 437 wind profile above 1.5 km altitude in figures 9 (c) and (d). Fitting a PWLF
 438 to the Viking data does not seem to support or refute this interpretation from
 439 Seiff (1993a); Haberle et al. (1993) but what is interesting is the structure in the
 440 wind profile in figure 9 (c) seems to support the prediction that the boundary
 441 layer at the VL-2 site was only 500 m in height at the landing time of 0949 LT.
 442 There is a small rotation from about 130° to 180° in the wind direction below
 443 this altitude.

444 5.2. The Phoenix case

445 Out of all the wind measurements, in section 4, Phoenix had the largest
 446 difference in wind speed when compared to the MCD. To investigate non-
 447 modelling, i.e. physical, causes of the mis-match we added realistic PWLF
 448 wind profiles to our trajectory model. See section 3.2 for more information on
 449 using PWLF wind profiles.

450 The PWLFs were arranged as follows. Two PWLFs, one for each u and
 451 v component of wind speed, were used in the top wind layer and two similar
 452 PWLFs were used in the bottom layer. The upper point of the PWLF repre-
 453 senting the u component of wind speed in the bottom layer was set equal to
 454 the lower point of the u component PWLF in the top layer. The v component
 455 PWLFs in the top and bottom layers were also vertically connected in this way.
 456 Altogether the trajectory model used four PWLFs.

457 Ten PWLF parameters were varied in the model to achieve a fit. These
 458 parameters were varied by the Hill Climbing algorithm described in section 3.1.
 459 Each parameter was sampled from ten independent Gaussian distributions. Four
 460 of the parameters were two u and v wind speed pairs, one pair for the upper
 461 points of two top layer PLWFs, and the other pair for the lower points of the
 462 two bottom layer PLWFs. An additional two parameters, a u and v pair, were
 463 used to vertically connect the PWLFs. The remaining four parameters were for

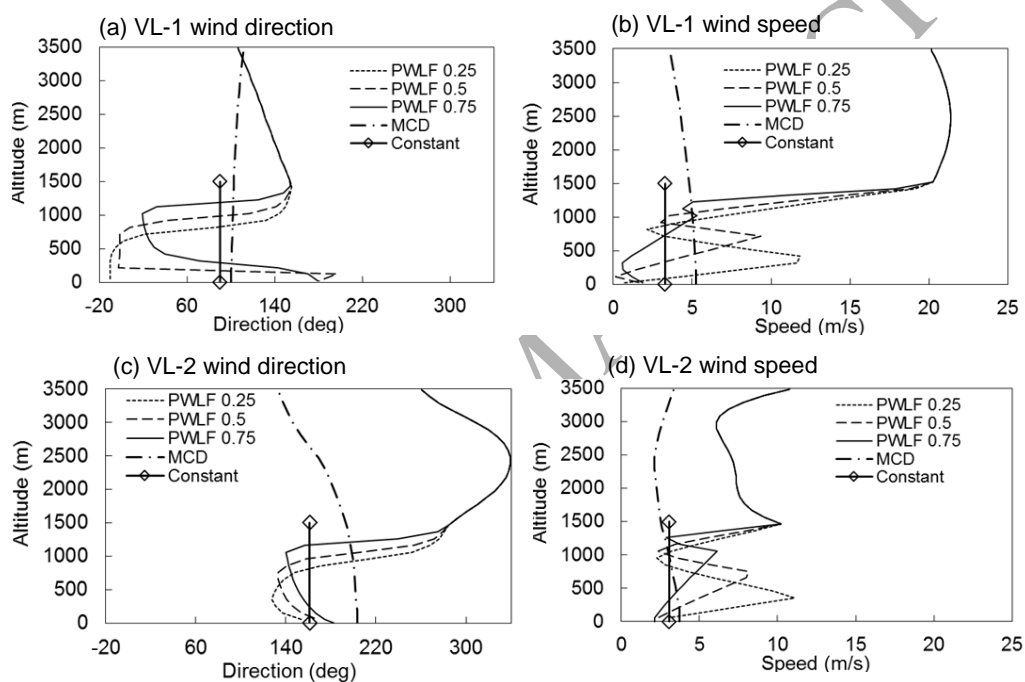


Figure 9: Wind speed and direction from EDLS and MCD compared for the Viking landers. A piece-wise linear function described in section 3 is used to help highlight any possible structural detail. The numbers 0.25, 0.5 and 0.75 in the legend refer to a scaling factors used to vary the mid-point of the PwLF relative to the lower boundary, i.e. 1.5 km, of the wind profiles provided by Seiff (1993a) and presented in this figure as the constant curves.

464 controlling the shape of the four PLWFs. The altitudes of the points connecting
465 the upper and lower PLWFs could also be varied but were set manually. Their
466 values can be found in the caption for figure 10 presented as a fraction of the
467 parachute deployment altitude.

468 The results in figure 10 for PWLF A are in general agreement with the results
469 in section 4. In figure 10 it can be seen that the wind speed rapidly decreases
470 to a minimum at about 2 km above the surface. The wind direction then
471 abruptly swings round from a north-westerly to a north-easterly and the wind
472 speed rapidly increases towards the surface. Figure 11 shows our trajectory
473 simulations for wind blown lander and the jettisoned hardware components
474 superimposed on an image of the Phoenix landing site. It can be seen the
475 heat shield was only mildly deflected from the incoming path of Phoenix while
476 the parachute had been strongly deflected. It should be noted that while the
477 heat shield was not deflected to the south it did experience a strong tail wind,
478 together with the lander on the parachute, which reduced its separation from
479 the lander. See appendix D for an explanation of this behaviour as it may
480 not immediately obvious how a tail wind could reduce the separation distance
481 between the heat shield and the lander.

482 The wind speed and direction profiles from the three MCD climate scenarios
483 are included in figure 10 for reference. It can be seen that none of the scenarios
484 can produce the high wind speeds especially at higher altitude. The upper alti-
485 tude wind speeds we measured are similar to measurements made by Karatekin
486 & Asmar (2011) who estimated a wind speed of 30 m s^{-1} during the parachute
487 descent. These Doppler measurements were made along the line of site between
488 Earth and Phoenix, which approximately lined up with the approach heading
489 of the lander to its landing site, i.e. from the north-west to the south-east. Lift
490 forces, which we neglected in our model, caused by a possibly deformed heat
491 shield (Desai et al., 2011) are not thought to have a significant contribution.
492 The heat shield was most likely spinning during the descent (Desai et al., 2011)
493 cancelling out any lift forces.

494 A magnified image of the Phoenix parachute is shown in figure 12. A strong
495 northerly wind near or at the surface, as suggested in our modelling results, is
496 apparent. The roundness of parachute canopy of the Phoenix lander in figure 12
497 (d) suggests the canopy was inflated by an intense horizontal wind as it came to
498 rest on the surface. The orientation of the parachute relative to the backshell,
499 to which it is attached, is consistent with a wind blowing from the north. In
500 contrast to the round canopy of Phoenix images of the parachute canopies for
501 Spirit, Opportunity and Curiosity in figures 12 (a) to (c) suggests their canopies
502 collapsed before reaching the surface. This would indicate the winds at these
503 sites were not so strong at landing.

504 The shape of the Curiosity parachute canopy was observed to change spo-
505 radically due to strong surface winds ([https://www.uahirise.org/releases/msl-
506 chute.php](https://www.uahirise.org/releases/msl-chute.php)). The first change was observed in an image acquired 3.5 months after
507 the landing. Previous images were acquired after 12 days and about one month
508 after landing showing no change in canopy shape. Images of the Phoenix canopy
509 were acquired 11 hours, 22 hours and 55 sols (see MRO image PSP_009290_2485)

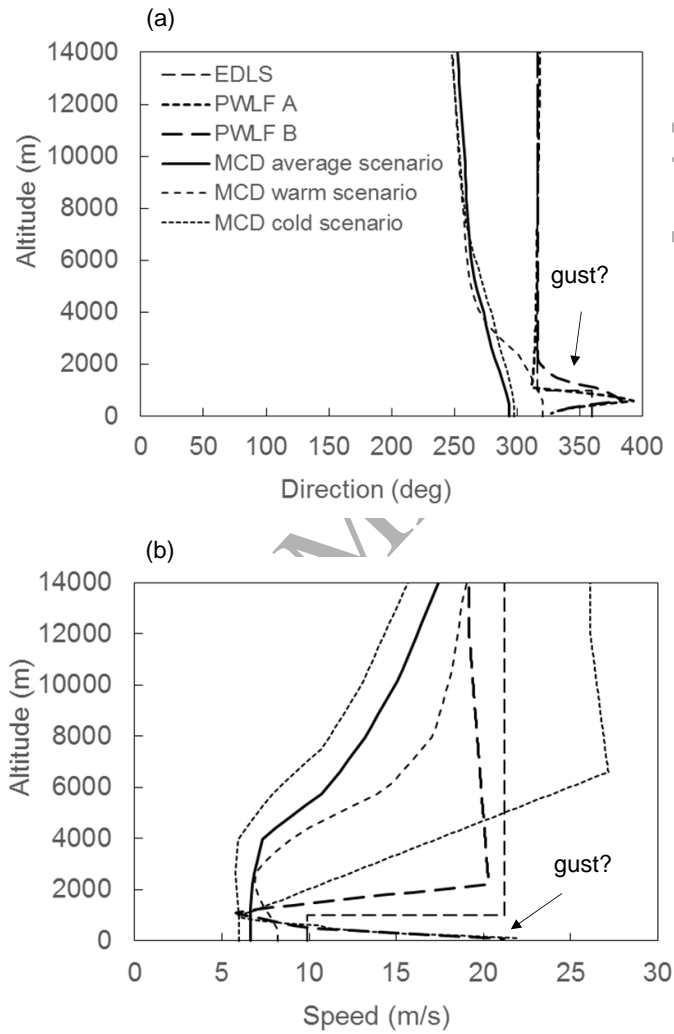


Figure 10: Wind speed and direction derived from the scatter of EDLS hardware on the surface compared to MCD standard atmosphere for Phoenix landing site. A PwLF (section 3.3) was used with four sections. The locations of these fixed points were at 13, 2, 1, 0.5 km for PwLF A and at 13, 6, 1, 0.5 km for PwLF B.

510 after landing showing no change in canopy shape. It appears unlikely that a
511 surface gust could have altered the shape of the parachute canopy in the first
512 11 hours of landing although it is not beyond the realms of possibility.

513 Wind gusts near the surface of Mars could be quite common and may present
514 an unseen risk to landing spacecraft. Dust devils at least appear to quite com-
515 mon on Mars (Spiga et al., 2016; Chapman et al., 2017; Lorenz & Jackson, 2016).
516 A dust devil or large turbulent eddy may be responsible for the high wind speed
517 near the surface in figure 10 (b). The Phoenix landed in afternoon when the
518 atmosphere is at its most turbulent. At least one lander, Spirit, has experienced
519 a Martian gusts during landing (Cheng et al., 2004).

520 For Phoenix the probability of encountering a convective vortex during land-
521 ing can be easily calculated. Ellehoj et al. (2010) identified 502 pressure drops in
522 the Phoenix pressure data that were associated with convective vortices passing
523 over the lander. Ellehoj et al. (2010) determined that during the first sols of
524 the mission the frequency of passing convective vortices in the afternoon was
525 around 0.2 per hour. This would translate to a 1 in 300 chance Phoenix would
526 have encountered a convective vortex during the final minute of its descent as
527 it descended in the afternoon.

528 Characterisation of the winds aloft in the PBL properly, in terms of its
529 gustiness, with the appropriate instruments and modelling would probably be
530 a prudent exercise given the apparent variations in the intensity of the Martian
531 winds. Such measurements would benefit future Mars landings and the pro-
532 tection of prelanded assets from wind blown components jettisoned during the
533 descent (Paton, 2017).

534 6. Conclusions

535 We have determined in situ for the first time the wind properties aloft in
536 the PBL at a variety of landing sites over the surface of Mars. The resulting
537 wind information spans a period of four decades or 21 Mars years. We used a
538 trajectory model to analyse the orbital images of jettisoned hardware and de-
539 termined the wind speed and direction that could account for their distribution
540 on the surface. In addition to orbital images of successful landed missions we
541 were able to extract useful information from images of Beagle 2 and Schiaparelli
542 hardware on the surface. To interpret our results comparisons were made to the
543 atmospheric winds from the state of the art Mars Climate Database (MCD).

544 We found the wind speed in the Martian PBL to be $<8 \text{ m s}^{-1}$. One exception
545 was Phoenix landing site where wind speeds $>20 \text{ m s}^{-1}$ were required to explain
546 the displacement of the heat shield and a wind gust close to the surface was
547 required to explain the displacement of the parachute. Further analysis suggests
548 the wind speed dropped below the tolerance of the EDLS just long enough for
549 Phoenix to be successfully released from the parachute.

550 Our results suggest that the wind direction on Mars generally correlates
551 over a period of at least 21 Martian years, i.e. over the period of landings on
552 Mars, with the MCD. On the other hand the wind speeds are not in such good

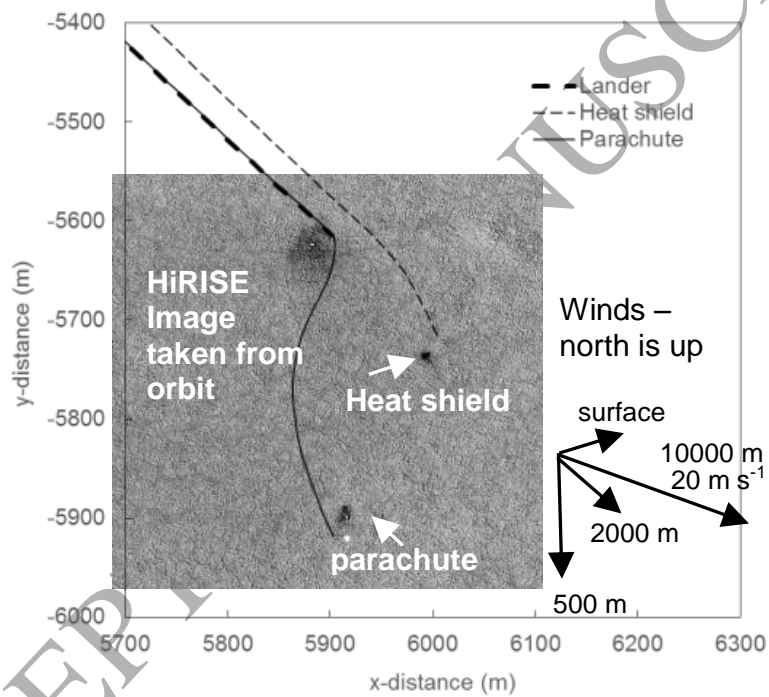


Figure 11: Ground track of Phoenix and jettisoned EDLS components.

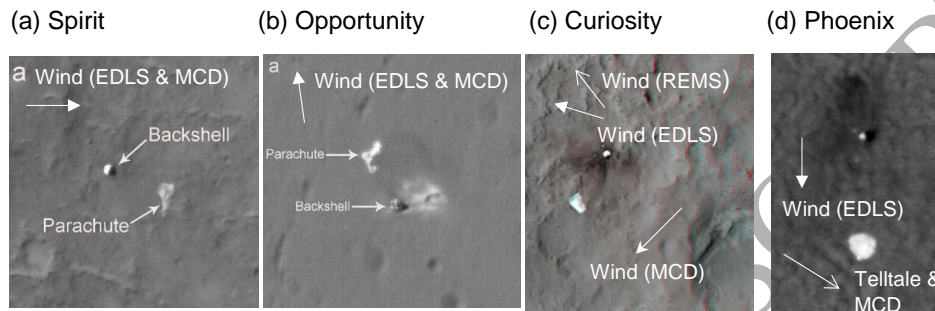


Figure 12: Parachute direction and canopy appearance.

553 agreement, which may be because wind speed is more sensitive to the thermal
 554 and mechanical turbulence in the PBL. The apparent high winds at the Phoenix
 555 site, turbulence during the landing of Viking lander 1, together with a gust
 556 previously reported experienced by the Spirit rover, suggest underlying large
 557 eddies in the winds aloft in the PBL will need to be more fully characterised in
 558 the future.

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561 References

- 562 Bauske, R. (2004). Dependence of the Beagle2 Trajectory on the Mars Atmo-
 563 sphere. In *18th International Symposium on Space Flight Dynamics* (p. 577).
 564 volume 548 of *ESA Special Publication*.
- 565 Chamberlain, T. E., Cole, H. L., Dutton, R. G., Greene, G. C., & Tillman,
 566 J. E. (1976). Atmospheric measurements on Mars - The Viking meteorology
 567 experiment. *Bulletin of the American Meteorological Society*, *57*, 1094–1104.
- 568 Chapman, R. M., Lewis, S. R., Balme, M., & Steele, L. J. (2017). Diurnal
 569 variation in martian dust devil activity. *Icarus*, *292*, 154–167.
- 570 Cheng, Y., Goguen, J., Johnson, A., Leger, C., Matthies, L., Martin, M. S.,
 571 & Wilson, R. (2004). The Mars Exploration Rovers Descent Image Motion
 572 Estimation System. *IEEE Intelligent Systems*, *19*, 13–21.

- 573 Cooley, C. G., & Lewis, J. G. (1977). *Viking lander system primary mission*
574 *performance report*. Martin Marietta.
- 575 Cruz, J. R., Mineck, R. E., Keller, D. F., & Bobskill, M. V. (2003). Wind Tunnel
576 Testing of Various Disk-Gap-Band Parachutes. In *17th AIAA Aerodynamics*
577 *Decelerator Systems Technology Conference, 19-22 May 2003, Phoenix, AZ,*
578 *United States*.
- 579 Cruz, J. R., Way, D. W., Shidner, J. D., Davis, J. L., Adams, D. S., & Kipp,
580 D. M. (2014). Reconstruction of the Mars Science Laboratory Parachute
581 Performance. *Journal of Spacecraft and Rockets*, *51*, 1185–1196.
- 582 Day, M., & Kocurek, G. (2016). Observations of an aeolian landscape: From
583 surface to orbit in Gale Crater. *Icarus*, *280*, 37–71.
- 584 Desai, P. N., & Knocke, P. C. (2007). Mars exploration rovers entry, descent,
585 and landing trajectory analysis. *Journal of the Astronautical Sciences*, *55*,
586 311–323.
- 587 Desai, P. N., Prince, J. L., Queen, E. M., Schoenenberger, M., Cruz, J. R., &
588 Grover, M. R. (2011). Entry, Descent, and Landing Performance of the Mars
589 Phoenix Lander. *Journal of Spacecraft and Rockets*, *48*, 798–808.
- 590 Dutta, S., & Braun, R. D. (2014). Statistical Entry, Descent, and Landing Per-
591 formance Reconstruction of the Mars Science Laboratory. *Journal of Space-*
592 *craft and Rockets*, *51*, 1048–1061.
- 593 Ellehoj, M. D., Gunnlaugsson, H. P., Taylor, P. A., Kahanpää, H., Bean, K. M.,
594 Cantor, B. A., Gheynani, B. T., Drube, L., Fisher, D., Harri, A.-M., Holstein-
595 Rathlou, C., Lemmon, M. T., Madsen, M. B., Malin, M. C., Polkko, J.,
596 Smith, P. H., Tamppari, L. K., Weng, W., & Whiteway, J. (2010). Convective
597 vortices and dust devils at the Phoenix Mars mission landing site. *Journal of*
598 *Geophysical Research (Planets)*, *115*, E00E16. doi:10.1029/2009JE003413.
- 599 Fallon, E., & Sinclair, R. (2003). Design and Development of the Main Parachute
600 for the Beagle 2 Mars Lander. In *17th AIAA Aerodynamic Decelerator Sys-*
601 *tems Technology Conference and Seminar. Monterey, California..*
- 602 Ferri, F., Karatekin, O., Aboudan, A., VanHove, B., Colombatti, G., Bettanini,
603 C., Debei, S., Gerbal, N., Lewis, S., & Forget, F. (2017). Atmospheric Mars
604 entry and landing investigation and analysis (AMELIA) by ExoMars 2016
605 Sciaparelli entry descent module. In *Sixth International Workshop on the-*
606 *Mars Atmosphere: Modelling and Observations, 17-20 Jan 2017, Granada,*
607 *Spain, CNES*.
- 608 Gunnlaugsson, H. P., Holstein-Rathlou, C., Merrison, J. P., Knak Jensen, S.,
609 Lange, C. F., Larsen, S. E., Madsen, M. B., Nørnberg, P., Bechtold, H.,
610 Hald, E., Iversen, J. J., Lange, P., Lykkegaard, F., Rander, F., Lemmon,
611 M., Renno, N., Taylor, P., & Smith, P. (2008). Telltale wind indicator for

- 612 the Mars Phoenix lander. *Journal of Geophysical Research (Planets)*, 113,
613 E00A04.
- 614 Haberle, R. M., Houben, H. C., Hertenstein, R., & Herdtle, T. (1993). A
615 boundary-layer model for Mars - Comparison with Viking lander and entry
616 data. *Journal of Atmospheric Sciences*, 50, 1544–1559.
- 617 Hess, S. L., Henry, R. M., Leovy, C. B., Tillman, J. E., & Ryan, J. A. (1977).
618 Meteorological results from the surface of Mars - Viking 1 and 2. *Journal of*
619 *Geophysical Research*, 82, 4559–4574.
- 620 Holstein-Rathlou, C., Maue, A., & Withers, P. (2016). Atmospheric studies
621 from the Mars Science Laboratory Entry, Descent and Landing atmospheric
622 structure reconstruction. *Planetary and Space Science*, 120, 15–23.
- 623 Jackson, D. W. T., Bourke, M. C., & Smyth, T. A. G. (2015). The dune effect
624 on sand-transporting winds on Mars. *Nature Communications*, 6, 8796.
- 625 Justus, C. G., Duvall, A., & Johnson, D. L. (2004). Mars-GRAM validation
626 with Mars global surveyor data. *Advances in Space Research*, 34, 1673–1676.
- 627 Karatekin, ., & Asmar, S. R. (2011). Entry trajectory reconstruction us-
628 ing Phoenix radio link: Deducing Martian wind velocities. In *IPPW-8,*
629 *Portsmouth, Virginia, 6-10 June 2011*.
- 630 Kass, D. M., Kleinböhl, A., Schofield, J. T., McCleese, D. J., Mischna, M. A.,
631 & McS Team (2008). MCS Views of the Northern Polar Atmosphere During
632 Phoenix Approach. In *Third International Workshop on The Mars Atmo-*
633 *sphere: Modeling and Observations*. volume 1447 of *LPI Contributions*.
- 634 Kauhanen, J., Siili, T., Jrvenoja, S., & Savijrvi, H. (2008). The mars lim-
635 ited area model and simulations of atmospheric circulations for the phoenix
636 landing area and season of operation. *Journal of Geophysical Research:*
637 *Planets*, 113, n/a–n/a. URL: <http://dx.doi.org/10.1029/2007JE003011>.
638 doi:10.1029/2007JE003011. E00A14.
- 639 Kutty, P. (2014). *Reconstruction and uncertainty quantification of entry, de-*
640 *scend and landing trajectories using vehicle aerodynamics*. Georgia Institute of
641 Technology, Masters Thesis.
- 642 Liu, Z. Y.-C., & Zimbelman, J. R. (2015). Recent near-surface wind directions
643 inferred from mapping sand ripples on Martian dunes. *Icarus*, 261, 169–181.
- 644 Lorenz, R. D., & Jackson, B. K. (2016). Dust Devil Populations and Statistics.
645 *Space Science Review*, 203, 277–297.
- 646 Martin, M. S., Mendeck, G. F., Brugarolas, P. B., Singh, G., Serricchio, F.,
647 Lee, S. W., Wong, E. C., & Essmiller, J. C. (2015). In-flight experience of
648 the Mars Science Laboratory Guidance, Navigation, and Control system for
649 Entry, Descent, and Landing. *CEAS Space Journal*, 7, 119–142.

- 650 McEwen, A. (2007). Mars Reconnaissance Orbiter High Resolution Imaging
651 Science Experiment. *MRO-M-HIRISE-2-EDR-V1.0, NASA Planetary Data*
652 *System*, .
- 653 Merrifield, J. (2015). Beagle 2 discovered. In *UK Space Conference. Liverpool,*
654 *UK.*
- 655 Michaels, T. I., & Rafkin, S. C. R. (2008). Meteorological predictions for candi-
656 date 2007 Phoenix Mars Lander sites using the Mars Regional Atmospheric
657 Modeling System (MRAMS). *Journal of Geophysical Research (Planets)*, *113*,
658 E00A07.
- 659 Millour, E., Forget, F., Gonzalez-Galindo, F., Spiga, A., Lebonnois, S., Lewis,
660 S. R., Montabone, L., Read, P. L., Lopez-Valverde, M. A., Gilli, G., Lefvre,
661 F., Montmessin, F., Desjean, M.-C., & Huot, J.-P. (2009). The mars climate
662 database (version 4.3). In *SAE Technical Paper*. SAE International.
- 663 Millour, E., Forget, F., & Lewis, S. R. (2015). *Mars Climate Database v5.2 user*
664 *manual*. Technical Report.
- 665 Montmessin, F. (2014). The MARs Boundary Layer Lidar experiment (MAR-
666 BLL): Winds at last! In *EGU General Assembly Conference Abstracts* (p.
667 16617). volume 16 of *EGU General Assembly Conference Abstracts*.
- 668 Montmessin, F. (2017). The MARs Boundary Layer Lidar experiment (MAR-
669 BLL): Winds at last! . In *Exomars Atmospheric Science and Missions Work-*
670 *shop, Saariselka, Finland, 26-30 March*.
- 671 Moores, J. E., Lemmon, M. T., Rafkin, S. C. R., Francis, R., Pla-Garcia,
672 J., de la Torre Juárez, M., Bean, K., Kass, D., Haberle, R., Newman, C.,
673 Mischna, M., Vasavada, A., Rennó, N., Bell, J., Calef, F., Cantor, B., Mc-
674 connochie, T. H., Harri, A.-M., Genzer, M., Wong, M., Smith, M. D., Javier
675 Martín-Torres, F., Zorzano, M.-P., Kemppinen, O., & McCullough, E. (2015).
676 Atmospheric movies acquired at the Mars Science Laboratory landing site:
677 Cloud morphology, frequency and significance to the Gale Crater water cycle
678 and Phoenix mission results. *Advances in Space Research*, *55*, 2217–2238.
679 doi:10.1016/j.asr.2015.02.007.
- 680 Moores, J. E., Lemmon, M. T., Smith, P. H., Komguem, L., & Whiteway, J. A.
681 (2010). Atmospheric dynamics at the Phoenix landing site as seen by the
682 Surface Stereo Imager. *Journal of Geophysical Research*, *115*, E00E08.
- 683 Moores, J. E., Schieber, J., Kling, A. M., Haberle, R. M., Moore, C. A., Ander-
684 son, M. S., Katz, I., Yavrouian, A., Malin, M. C., Olson, T., Rafkin, S. C. R.,
685 Lemmon, M. T., Sullivan, R. J., Comeaux, K., & Vasavada, A. R. (2016).
686 Transient atmospheric effects of the landing of the Mars Science Laboratory
687 rover: The emission and dissipation of dust and carbazic acid. *Advances in*
688 *Space Research*, *58*, 1066–1092. doi:10.1016/j.asr.2016.05.051.

- 689 NASA (2008). *Phoenix Landing: Mission to the Martian Polar North*. NASA
690 Press kit.
- 691 Parteli, E. J. R., Durán, O., Tsoar, H., Schwämmle, V., & Herrmann, H. J.
692 (2009). Dune formation under bimodal winds. *Proceedings of the National*
693 *Academy of Science*, *106*, 22085–22089.
- 694 Paton, M. (2017). Protection of surface assets on Mars from wind blown jetti-
695 sioned spacecraft components. *Acta Astronautica*, *136*, 395–406.
- 696 Petrosyan, A., Galperin, B., Larsen, S. E., Lewis, S. R., Määttänen, A., Read,
697 P. L., Renno, N., Rogberg, L. P. H. T., Savijärvi, H., Siili, T., Spiga, A.,
698 Toigo, A., & Vázquez, L. (2011). The Martian Atmospheric Boundary Layer.
699 *Reviews of Geophysics*, *49*, RG3005.
- 700 Savijärvi, H. (1995). Mars boundary layer modeling: Diurnal moisture cycle
701 and soil properties at the Viking Lander 1 Site. *Icarus*, *117*, 120–127.
- 702 Savijärvi, H. (1999). A model study of the atmospheric boundary layer in the
703 Mars Pathfinder lander conditions. *Quarterly Journal of the Royal Meteorological*
704 *Society*, *125*, 483–493.
- 705 Savijärvi, H. (2012). Mechanisms of the diurnal cycle in the atmospheric bound-
706 ary layer of Mars. *Quarterly Journal of the Royal Meteorological Society*, *138*,
707 552–560. doi:10.1002/qj.930.
- 708 Savijärvi, H., Crisp, D., & Harri, A.-M. (2005). Effects of CO₂ and dust on
709 present-day solar radiation and climate on Mars. *Quarterly Journal of the*
710 *Royal Meteorological Society*, *131*, 2907–2922.
- 711 Savijärvi, H., & Kauhanen, J. (2008). Surface and boundary-layer modelling
712 for the Mars Exploration Rover sites. *Quarterly Journal of the Royal Meteorological*
713 *Society*, *134*, 635–641.
- 714 Savijärvi, H., & Määttänen, A. (2010). Boundary-layer simulations for the Mars
715 Phoenix lander site. *Quarterly Journal of the Royal Meteorological Society*,
716 *136*, 1497–1505.
- 717 Savijärvi, H., Määttänen, A., Kauhanen, J., & Harri, A.-M. (2004). Mars
718 Pathfinder: New data and new model simulations. *Quarterly Journal of the*
719 *Royal Meteorological Society*, *130*, 669–683.
- 720 Seiff, A. (1993a). Mars atmospheric winds indicated by motion of the Viking
721 landers during parachute descent. *Journal of Geophysical Research*, *98*, 7461–
722 7474.
- 723 Seiff, A. (1993b). Mars atmospheric winds indicated by motion of the Viking
724 landers during parachute descent. *Journal of Geophysical Research*, *98*, 7461–
725 7474.

- 726 Seiff, A., & Kirk, D. B. (1977). Structure of the atmosphere of Mars in summer
727 at mid-latitudes. *Journal of Geophysical Research*, *82*, 4364–4378.
- 728 Seiff, A., Tillman, J. E., Murphy, J. R., Schofield, J. T., Crisp, D., Barnes,
729 J. R., LaBaw, C., Mahoney, C., Mihalov, J. D., Wilson, G. R., & Haberle, R.
730 (1997). The atmosphere structure and meteorology instrument on the Mars
731 Pathfinder lander. *Journal of Geophysical Research*, *102*, 4045–4056.
- 732 Smith, D. E., Zuber, M. T., Frey, H. V., Garvin, J. B., Head, J. W., Muhleman,
733 D. O., Pettengill, G. H., Phillips, R. J., Solomon, S. C., Zwally, H. J., Banerdt,
734 W. B., Duxbury, T. C., Golombek, M. P., Lemoine, F. G., Neumann, G. A.,
735 Rowlands, D. D., Aharonson, O., Ford, P. G., Ivanov, A. B., Johnson, C. L.,
736 McGovern, P. J., Abshire, J. B., Afzal, R. S., & Sun, X. (2001). Mars Orbiter
737 Laser Altimeter: Experiment summary after the first year of global mapping
738 of Mars. *Journal of Geophysical Research*, *106*, 23689–23722.
- 739 Spencer, D. A., Blanchard, R. C., Braun, R. D., Kallemeyn, P. H., & Thurman,
740 S. W. (1999). Mars pathfinder entry, descent, and landing reconstruction.
741 *Journal of Spacecraft and Rockets*, *36*, 357366.
- 742 Spiga, A., Barth, E., Gu, Z., Hoffmann, F., Ito, J., Jemmett-Smith, B., Klose,
743 M., Nishizawa, S., Raasch, S., Rafkin, S., Takemi, T., Tyler, D., & Wei,
744 W. (2016). Large-Eddy Simulations of Dust Devils and Convective Vortices.
745 *Space Science Review*, *203*, 245–275.
- 746 Spiga, A., & Lewis, S. R. (2010). Martian mesoscale and microscale wind vari-
747 ability of relevance for dust lifting. *International Journal of Mars Science
748 and Exploration*, *5*, 146–158.
- 749 Stanzel, C., Pätzold, M., Williams, D. A., Whelley, P. L., Greeley, R., Neukum,
750 G., & the HRSC Co-Investigator Team (2008). Dust devil speeds, directions
751 of motion and general characteristics observed by the Mars Express High
752 Resolution Stereo Camera. *Icarus*, *197*, 39–51.
- 753 Szalai, C., Thoma, B., Lee, W., Maki, J., Wilcoxson, W. H., Venkatapathy, E.,
754 & White, T. (2011). Mars exploration rover heatshield observation campaign.
755 *42nd AIAA Thermophysics Conference*, . doi:10.2514/6.2011-3956.
- 756 Tamppari, L. K., Bass, D., Cantor, B., Daubar, I., Dickinson, C., Fisher, D.,
757 Fujii, K., Gunnlaugsson, H. P., Hudson, T. L., Kass, D., Kleinböhl, A.,
758 Konguem, L., Lemmon, M. T., Mellon, M., Moores, J., Pankine, A., Pathak,
759 J., Searls, M., Seelos, F., Smith, M. D., Smrekar, S., Taylor, P., Holstein-
760 Rathlou, C., Weng, W., Whiteway, J., & Wolff, M. (2010). Phoenix and
761 MRO coordinated atmospheric measurements. *Journal of Geophysical Re-
762 search (Planets)*, *115*, E00E17. doi:10.1029/2009JE003415.
- 763 Taylor, A. P. (2000). Design Optimization of the Beagle II Mars Lander Airbags
764 Through Explicit Finite Element Analysis An Update . In *6th International
765 LS-DYNA Conference, April 9-11, 2000, Dearborn, Michigan*.

- 766 Way, D. W. (2013). Preliminary assessment of the Mars Science Laboratory
767 entry, descent, and landing simulation. In *Proceedings of the 2013 IEEE*
768 *Aerospace Conference* (p. 343).
- 769 Withers, P., & Smith, M. D. (2006). Atmospheric entry profiles from the Mars
770 Exploration Rovers Spirit and Opportunity. *Icarus*, *185*, 133–142.
- 771 Witkowski, A., Kandis, M., & Adams, D. S. (2009). Mars 2007 scout phoenix
772 parachute decelerator system program overview. *2007 IEEE Aerospace Con-*
773 *ference, Decelerator Systems Technology Conference and Seminar 4 - 7 May*
774 *2009, Seattle, Washington, .*

775 **Appendix A. Vehicle properties, EDL trajectories and model input** 776 **uncertainties**

777 Tables A.5, A.6 and A.7 contain the vehicle properties, reconstructed EDL
778 properties and the model input uncertainties respectively. The vehicle and re-
779 constructed EDL properties can be found in Cooley & Lewis (1977); Cheng
780 et al. (2004); Desai & Knocke (2007); Desai et al. (2011); Way (2013); Martin
781 et al. (2015). The model input uncertainties listed in table A7 are based on in-
782 formation found in Seiff & Kirk (1977); Seiff (1993b); Cruz et al. (2003); Bauske
783 (2004); Withers & Smith (2006); Dutta & Braun (2014); Kutty (2014); Ferri
784 et al. (2017). The atmospheric density used in the model and listed in table A6
785 are derived from measurements made during landings by Viking, MER, Phoenix,
786 Curiosity and Schiaparelli can be found in Seiff & Kirk (1977); Withers & Smith
787 (2006); Dutta & Braun (2014); Holstein-Rathlou et al. (2016); Ferri et al. (2017).
788 For Beagle 2 the atmospheric density was obtained from atmospheric models
789 which all tended to converge at low altitude making the modelling uncertainty
790 small (Bauske, 2004).

791 **Appendix B. Model**

792 The coordinate system is based on a rotating equatorial frame of reference
793 with its origin at the centre of Mars as illustrated in figure 13. The frame of
794 reference is stationary in inertial space i.e. it moves around the sun with Mars
795 and its vertical axis is the same as the rotation axis of Mars. The three main
796 spatial coordinates are longitude (θ), latitude (ϕ) and radial distance from the
797 centre (r). The heading (h) is the direction of travel relative to the vector
798 pointing north from the current position. The radial velocity is positive when
799 r is increasing and likewise the longitudinal and latitudinal velocity is positive
800 as θ and ϕ increases.

801 The equations of free motion without any lift or drag from interactions with
802 the atmospheric flow are as follows:

$$\frac{\partial^2 r}{\partial t^2} = r \left(\frac{\partial \theta}{\partial t} \right)^2 (\cos \phi)^2 + r \frac{\partial \phi}{\partial t} - \frac{GM}{r^2} \quad (\text{B.1})$$

Table A.5: Vehicle and EDLS component properties.

Parameter	VL-1	VL-2	PF	B2	MER 1	MER 2	PHX	MSL	SCH
Aeroshell									
LoD	0.18	0.18	0	0	0	0	0.06	0.24	0
$m_{entry}(kg)$	977	977	585	69	827	832	602	3100	577
C_D	1.55	1.55	1.5	1.5	1.5	1.5	1.6	1.55	1.55
LoD	0.18	0.18	0	0	0	0	0.06	0.24	0
d_{aero} (m)	3.5	3.5	2.65	0.9	2.65	2.6	2.65	4.5	2.4
A_{aero} (m ²)	15	15	8	1	8	8	9	25	7
β_{aero} (kg m ⁻²)	66	66	71	72	100	101	65	126	82
Parachute deployed									
$m_{noseshield}$ (kg)	740	740	521	59	740	754	511	2718	497
d_{chute} (m)	16	16	12.74	3.2	14.1	14.1	11.73	19.7	12
C_D	0.67	0.67	0.44	0.6	0.6	0.6	0.6	0.6	0.6
A_{chute} (m ²)	135	135	56	4	69	69	49	139	45
$\beta_{withshield}$ (kg m ⁻²)	7	7	10	17	9	9	9	17	10
$\beta_{noseshield}$ (kg m ⁻²)	5	5	9	15	8	8	8	15	9
Lander									
m_{lander} (kg)	673	673	464	50	540	545	401	2369	477
T (kN)	2.1	2.1	16.7	0	22.4	22.4	3.6	24.9	3.6
T/W (N kg ⁻¹)	3	3	36	0	41.5	41	9	10.5	7.5
Heat shield									
m_{shield} (kg)	237	237	64.4	9.78	78	78	62	382	80
C_D	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0
d_{shield} (m)	3.5	3.5	2.65	0.9	2.65	2.65	2.65	4.5	2.4
A_{shield} (m ²)	15	15	8	1	8	8	8.5	25	7
β_{shield} (kg m ⁻²)	25	25	12	15	14	14	11	24	18
Backshell									
m_{back} (kg)	67	67	57	9.5	209	209	110	349	20
β_{back}	0.5	0.5	1.0	2.4	3.0	3.0	2.2	2.5	0.4
Ringsail									
m_{ring}				1					
D_{ring}				10					
C_D				0.8					
A_{ring}				62.8					
β_{ring}				0.02					
β_{wl}				1.25					

Table A.6: Trajectory parameters values required for modelling and verification of the simulations.

Parameter	VL-1	VL-2	PF	B2	MER 1	MER 2	PHX	MSL	SCH
Atmosphere parameters									
ρ_{surf} (kg m ⁻³)	0.017	0.018	n/a	0.019	0.015	0.013	0.022	0.014	0.014
h_{atm} (km)	10.0	10.0	10.0	10.0	10.0	10.0	10.0	10.0	10.0
Entry									
z_{entry} (km)	242	242	207	120	130	128	130	125	121.5
v_{entry} (km s ⁻¹)	4.7	4.7	7.4	5.6	5.6	5.7	5.7	5.8	5.7
γ_{entry} (deg)	17	17.1	16.9	16.6	11.5	11.5	12.5	15.5	16
α_{entry} (deg)	56	43	255	86	79	86.5	79	94	118
θ_{entry} (deg)	298	117	344	83	162	341	200.5	127	343
ϕ_{entry} (deg)	13	37	23	11	-18	-3	69	-4	4
Parachute deployment									
z_{chute} (km)	5.9	5.9	9.40	7.1	7.5	7.5	13.3	12.0	11.0
v_{chute} (m s ⁻¹)	238	237	377	360	411	430	384	492	478
γ_{chute} (deg)	-53.4	-50.8	-25.5	-25.4	-28.1	-26.7	-27.5	-22.4	-26.6
α_{chute} (deg)	56	55	253	88	75	86	108	93	119
Heat shield jettison									
z_{shield} (km)	4.5	4.6	7.7	0.3	5.7	5.7	11	11	5
t_{shield} (s)	7	7	20	20	20	20	15	20	20
v_{shield} (m s ⁻¹)	115	115	100	100	100	100	123	100	100
Backshell and parachute jettison									
z_{lander} (km)	1.46	1.44	0.02	0.03	0.02	0.02	0.93	1.66	1.20
v_{lander} (m s ⁻¹)	100	100	100	100	100	100	100	100	100
θ_{surf} (deg)	312	134	326	90	175	354	234	137	354
ϕ_{surf} (deg)	22.5	48	19	11.5	-14.6	-2	62	-5	-2
Beagle 2 main parachute									
z_{ring} (km)				2.6					
v_{ring} (m s ⁻¹)				96					

Table A.7: Uncertainty (σ) on model input parameters.

	VL-1	VL-2	B2	MER 1	MER 2	PHX	MSL	SCH
Angle (deg)	1	1	1	1	1	1	1	1
Velocity (m s ⁻¹)	1	1	1	1	1	12	0.1	1
Altitude (m)	100	100	2000	1700	1800	238	75	100
Density (%)	0.5	0.5	5.0	5.0	5.0	5.0	5.0	5.0
Heat shield drag coefficient (%)	5	5	5	5	5	5	5	5
Parachute drag coefficient (%)	5	5	5	5	5	5	5	5

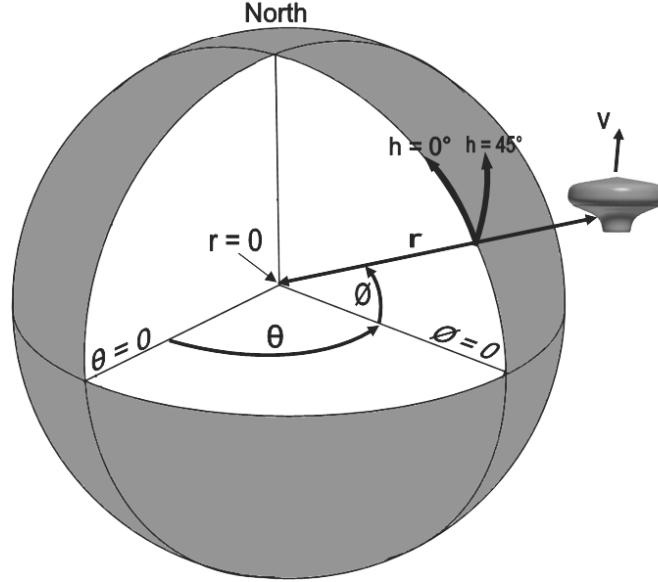


Figure B.13: Coordinate system used for trajectory modelling showing the important parameters used in the model and for interpreting the results.

$$\frac{\partial^2 \theta}{\partial t^2} = 2 \frac{\partial \theta}{\partial t} \frac{\partial \phi}{\partial t} \frac{\sin \phi}{\cos \theta} - \frac{2}{r} \frac{\partial r}{\partial t} \frac{\partial \theta}{\partial t} \quad (\text{B.2})$$

$$\frac{\partial^2 \phi}{\partial t^2} = - \left(\frac{\partial \theta}{\partial t} \right)^2 \cos \phi \sin \phi - \frac{2}{r} \frac{\partial r}{\partial t} \frac{\partial \phi}{\partial t} \quad (\text{B.3})$$

803 Where r is the radial distance from the centre of Mars, θ is the longitude,
 804 ϕ is the latitude, γ is the flight path angle, h is heading, D is the drag force, L
 805 is the lift force, M is the mass of Mars, G is the gravitational constant and t is
 806 time.

807 Figure B13 and B14 show the essential parameters for modeling the motion
 808 of the vehicle through the atmosphere. The flight path angle (γ) is the angle
 809 between the velocity vector (direction of motion) and the horizon. The angle of
 810 attack (α) is the angle between the long axis, in this case the direction the nose
 811 is pointing, and the velocity vector. Normally the angle of attack will remain
 812 constant for extended periods of time while the flight path angle will vary as
 813 aerodynamic forces effect its motion.

814 The equations used to calculate the lift and drag forces are:

$$L = 0.5 C_L \rho A v^2 \quad (\text{B.4})$$

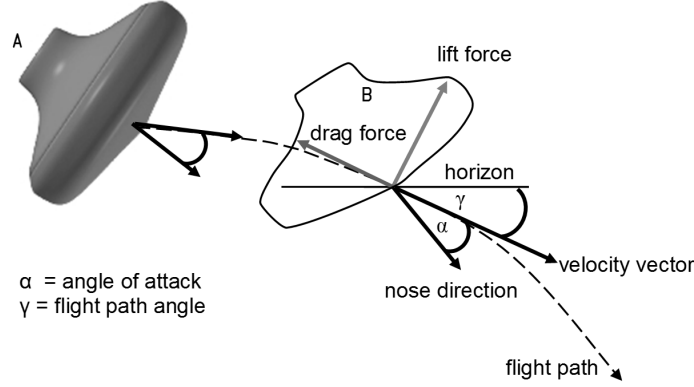


Figure B.14: Essential parameters for calculating the trajectory of a lifting body through the atmosphere.

$$D = 0.5C_D\rho Av^2 \quad (\text{B.5})$$

where C_L is the lift coefficient, C_D is the drag coefficient, ρ is the density of the local atmosphere, A is the reference area and v is the velocity. The lift and drag coefficients can be measured directly in a wind tunnel or calculated using a model. These coefficients describe the airflow around the vehicle and are dependent on the angle of attack. The equations of motion for drag are as follows.

$$\frac{\partial^2 r}{\partial t^2} = -\frac{D}{m} \sin(\gamma) \quad (\text{B.6})$$

$$\frac{\partial^2 \theta}{\partial t^2} = -\frac{D \cos(\gamma) \sin(h)}{m r \cos(\phi)} \quad (\text{B.7})$$

$$\frac{\partial^2 \phi}{\partial t^2} = -\frac{D \cos(\gamma) \cos(h)}{m r} \quad (\text{B.8})$$

Similar equations are used to calculate lift but with a positive instead of a negative sign in front. The local atmospheric density can be calculated with varying level of complexity depending on the accuracy required. The atmosphere is defined as follows:

$$\rho = \exp(-z/H) \quad (\text{B.9})$$

where H is the scale height. The force from the winds includes that from the rotation of the atmosphere. The bulk of the atmosphere is assumed to rotate at the same speed as the surface when there are no winds. For the zonal case the wind is as follows.

$$v_x = v_{surf} + v_\theta \quad (\text{B.10})$$

829 where v_{surf} is the speed of the surface and v_θ is the zonal wind relative to
830 the bulk rotational motion of the atmosphere.

831 This would mean that a spacecraft dropped vertically into the planet's at-
832 mosphere at the equator would be decelerated horizontally even though the drag
833 is assumed to act through the direction of motion. This is because the velocity
834 used in the drag equations is as follows.

$$v = (v_x + v_y + v_z)^2 \quad (\text{B.11})$$

835 where v_x is the longitudinal velocity, v_y is the vertical velocity and v_z is the
836 latitudinal velocity. This approach makes sense for entry vehicles as they are
837 designed to be dynamically stable, i.e. the centre of mass is in front of the centre
838 of pressure, as the vehicle would reorientate itself to line up with the relative
839 velocity vector of the wind.

840 The equations are solved numerically integrating with time.

841 Appendix C. Displacement of jettisoned hardware by wind layers

842 The characteristic response of the Viking lander on the parachute to a step
843 change in the horizontal wind speed was shown by Seiff (1993a) to be charac-
844 terised as an exponential acquisition of the wind speed by the lander. After
845 correcting for a dimensionally inconsistent exponent in paper's equation A7 we
846 have the following.

$$\frac{v}{v_w} = 1 - e^{-\frac{g}{v_{eq}} t} \quad (\text{C.1})$$

847 where v is the horizontal velocity of the object, v_w is the horizontal wind ve-
848 locity, g is the acceleration due to gravity, v_{eq} is the equilibrium descent velocity,
849 i.e. terminal velocity, and t is the time. Seiff (1993a) define the characteristic
850 response time of the object, which is the time to acquire $1 - e^{-1}$ or 63%, of the
851 step change in the horizontal wind speed, as follows.

$$\tau = \frac{v_{eq}}{g} \quad (\text{C.2})$$

852 Figure C15 shows the displacement of a descending lander and its jettisoned
853 hardware calculated using our model. In figure C15 (a) and (b) a two kilometre
854 thick wind layer is located with its mid-point located at an altitude of 7000 m.
855 Throughout the wind layer the wind is blowing at a speed of 5 m s^{-1} from the
856 west. Outside the wind layer the wind speed is set to zero.

857 In figure C15 (b) the heat shield takes a longer time than the lander on the
858 parachute to respond to the changes in the wind speed. Equation C2 can be used
859 to compare the response times of the two objects. In our model the heat shield
860 is descending at approximately 76 m s^{-1} while the lander on the parachute is
861 descending at approximately 60 m s^{-1} . The characteristic response time are

862 then 20 and 16 s for the heat shield and lander on the parachute respectively.
863 These times are consistent with the acquisition of wind speeds by the heat shield
864 and lander on the parachute as can be seen in figure C15 (b).

865 In figure C15 (c) and (d) a wind layer has been placed close to the surface
866 in our model. The free flying lander, having jettisoned the parachute, descends
867 through the layer using its rocket motors and is not affected by the winds. In
868 contrast the parachute is significantly displaced from the lander by the winds
869 as it has a low ballistic coefficient and hence a low descent velocity. In our
870 model the jettisoned parachute descends with a terminal velocity of 27 m s^{-1}
871 corresponding to a response time of about 7 s which is consistent with figure
872 C15 (d). The heat shield is only weakly deflected from its vertical patch by the
873 wind layer due to its large ballistic coefficient and resulting high descent speed.

874 Some important conclusions regarding the analysis of displaced hardware
875 can be drawn from the calculations in this section. The response times of the
876 jettisoned hardware suggest a full trajectory model is required for measuring
877 the wind speeds rather than relying on approximate calculation that assume
878 the objects instantaneously acquire the wind speed. For example close to the
879 surface, where the object may not have time to acquire the wind speed, a wind
880 layer will produce different separation distances to a layer of the same thickness
881 but at a higher altitude.

882 It is somewhat obvious that a jettisoned parachute will be increasingly dis-
883 placed from the lander by stronger winds during its descent, i.e. for higher winds
884 speeds but constant wind direction. When considering the jettisoned heat shield
885 it may not be immediately obvious that a stronger wind will decrease the sep-
886 aration distance between the lander and heat shield. For example see figure
887 C15 (a) for the case where there is a tail wind. This results in a decrease in
888 separation distance compared to the no wind case, i.e. $A < B$. The lander on the
889 lander on the parachute is blown closer to the heat shield.

890 **Appendix D. Impact locations for a jettisoned heat shield**

891 After the parachute has been deployed the lander begins to rapidly turn
892 towards the vertical. At some time after the parachute has been deployed and
893 before the lander has completed its turn into a vertical descent the heat shield is
894 jettisoned. See table A5 for jettison angles and timer settings for each lander. If
895 there are no winds a jettisoned heat shield will impact the surface at some point
896 downrange from the lander due to its high ballistic coefficient and its horizontal
897 velocity component upon being jettisoned. For a vehicle approaching from the
898 north-west as with the Phoenix case the impact point will be south-west of the
899 lander.

900 The displaced heat shield impact point due to winds will be directly related
901 to the magnitude and direction of the winds during the descent. In figure C16
902 shows the displacement of a heat shield with a nominal drag coefficient by winds
903 from different directions. The displacement of the parachute due to winds is also
904 shown for comparison.

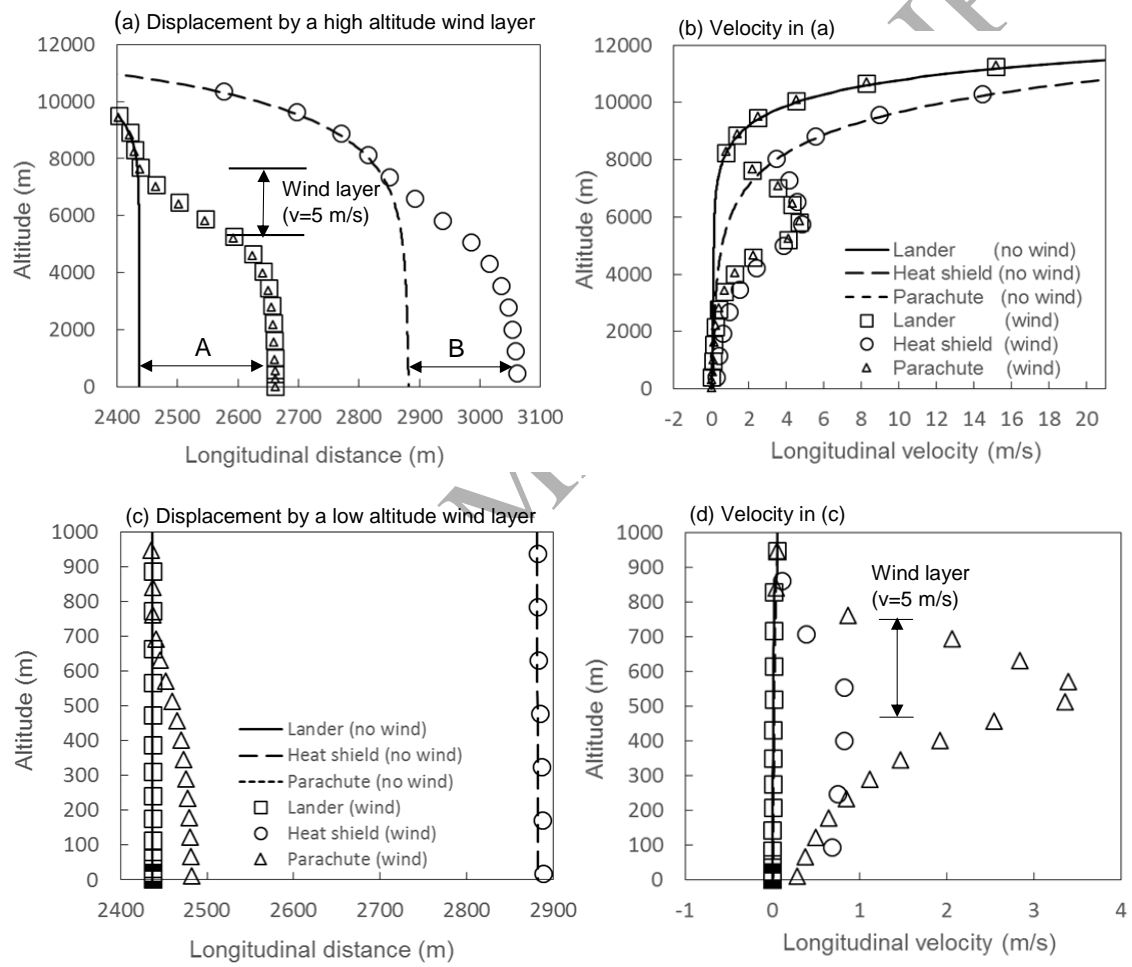


Figure C.15: Wind layers and their effect on the trajectory.

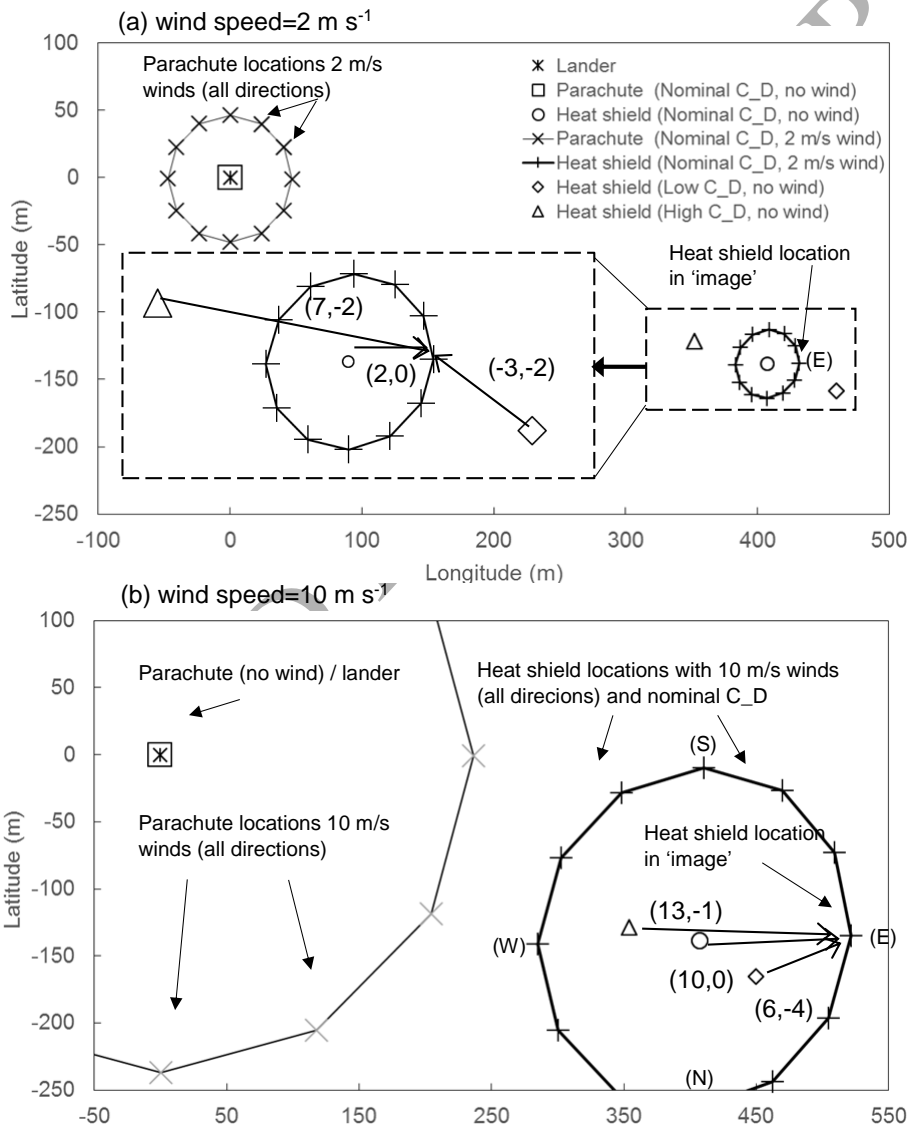


Figure D.16: Predicted impact locations the heat shield due to winds and uncertainty on its drag coefficient.

905 As the impact points of the heat shield are directly related to the winds
906 it is possible to visualise, in terms of location, the influence of the heat shield
907 uncertainty on the measurements. Figure C16 shows the impact locations of
908 the heat shield for high and low 3 sigma drag coefficients. If we assume one of
909 the impact locations due to winds as a reference case then it can be seen that a
910 model, initialised with a high or low 3 sigma uncertainty, will produce larger
911 uncertainties in direction (see arrows in figure C16) when weaker winds are
912 blowing, i.e. the model will calculate nearly opposite wind directions for low
913 and high 3 sigma uncertainties (also see figure 7).

914 If the winds are weak then a heat shield with a low drag coefficient will
915 overshoot the observed heat shield impact location while a heat shield with a
916 high drag coefficient will undershoot this point. When this situation arises the
917 model will calculate winds in opposite directions eventually reaching a total
918 uncertainty on the wind direction of 180 degrees. Wind direction uncertainty
919 can be seen to reverse increases to more than 90° below a wind speed of 4 m
920 s^{-1} .