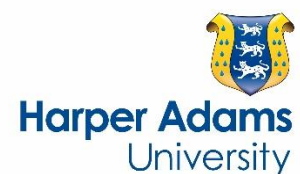


Biochar effects on methane emissions from soils: a meta-analysis

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1 **Biochar effects on methane emissions from soils: A meta-analysis**

2

3 Running title: Meta-analysis of biochar effects on CH₄ flux

4

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22 **Keywords**

23 Biochar; methane; soil; meta-analysis; standardised mean difference; greenhouse gas.

24

25 Type of Paper: Research Review

26 **Abstract**

27 Methane (CH₄) emissions have increased by more than 150% since 1750, with
28 agriculture being the major source. Further increases are predicted as permafrost
29 regions start thawing, and rice and ruminant animal production expand. Biochar is
30 posited to increase crop productivity while mitigating climate change by sequestering
31 carbon in soils and by influencing greenhouse gas fluxes. There is a growing
32 understanding of biochar effects on carbon dioxide and nitrous oxide fluxes from soil.
33 However, little is known regarding the effects on net methane exchange, with single
34 studies often reporting contradictory results. Here we aim to reconcile the disparate
35 effects of biochar application to soil in agricultural systems on CH₄ fluxes into a single
36 interpretive framework by quantitative meta-analysis.

37 This study shows that biochar has the potential to mitigate CH₄ emissions from
38 soils, particularly from flooded (i.e. paddy) fields (Hedge's $d = -0.87$) and/or acidic
39 soils (Hedge's $d = -1.56$) where periods of flooding are part of the management regime.
40 Conversely, addition of biochar to soils that do not have periods of flooding (Hedge's
41 $d = 0.65$), in particular when neutral or alkaline (Hedge's $d = 1.17$ and 0.44 ,
42 respectively), may have the potential to decrease the CH₄ sink strength of those soils.
43 Global methane fluxes are net positive as rice cultivation is a much larger source of
44 CH₄ than the sink contribution of upland soils. Therefore, this meta-study reveals that
45 biochar use may have the potential to reduce atmospheric CH₄ emissions from
46 agricultural flooded soils on a global scale.

47

48 **1. Introduction**

49 Agriculture accounts for 10–12% of total global anthropogenic emissions of
50 greenhouse gases (GHGs), which includes 50% of global methane (CH₄) emissions
51 (Smith et al., 2007). Methane emissions have increased by 151% since 1750 (IPCC,
52 2007), and are currently increasing at a rate of 0.003 μmol mol⁻¹ year⁻¹ (Butenhoff and
53 Khalil, 2007; Bloom et al., 2010). Further increases are projected due to the growing
54 demand for food, particularly animal protein, which could require *ca.* 70 million ha of
55 additional land to fall under agricultural production (Alexandratos and Bruinsma,
56 2012).

57 Methane is primarily produced in water-logged anoxic soils by methanogenic
58 archaea via methanogenesis (Conrad, 2007). Conversely, well-aerated upland soils are
59 biological sinks for atmospheric CH₄ (Boone et al., 1993; Dunfield, 2007). Soil CH₄
60 uptake is driven by microbial oxidation of CH₄ by methanotrophs from groups
61 including α- and γ-proteobacteria, a group of obligate aerobic bacteria some of which
62 feed solely on CH₄ and others, along with genera such as *Methylocella* and
63 *Methylocapsa*, that are facultive methanotrophs (Pratscher et al., 2011; Knief, 2015).
64 Generally, both processes – methanogenesis and methanotrophy – can occur
65 simultaneously in micro-sites within the soil, or can be stratified with CH₄ production
66 occurring in more highly anoxic depths, and CH₄ consumption occurring in overlaying
67 oxic soil horizons. Here, the soil acts as a net source or sink depending on which is the
68 overriding process (Hiltbrunner et al., 2012). However, these two processes can
69 dynamically interact (Kammann et al., 2009) with CH₄ consumption functioning as a
70 "biofilter" process that can ameliorate CH₄ emissions in various ecosystems, including
71 rice paddies and landfill cover soils.

72 One of the main attractions underlying the biochar concept is the combination
73 of soil carbon (C) sequestration with soil fertility (crop yield) increases (Glaser et al.,
74 2001; Lehmann, 2007). Initial research efforts have focused on biochar's recalcitrance
75 as a potential means to sequester C in soils (Lehmann et al., 2006; Nguyen and
76 Lehmann, 2009; Gurwick et al., 2013) while concurrently increasing crop yields
77 (Jeffery et al., 2011). It has also been shown to mitigate nitrous oxide (N₂O) emissions
78 from agricultural soils (meta-analysis: Cayuela et al., 2014). The interactions between
79 biochar and GHG fluxes such as carbon dioxide (CO₂) and N₂O, and the associated
80 mechanisms, are becoming better understood (Cayuela et al., 2013; Maestrini et al.,
81 2014; Cayuela et al., 2015; Obia et al., 2015; Sagrilo et al., 2015). However, there is
82 still a paucity of information on CH₄ flux effects beyond the single study scale, which
83 often report contradictory results.

84 Biochar has been shown to increase (Zhang et al., 2010; Spokas et al., 2011),
85 decrease (Feng et al., 2012; Dong et al., 2013; Reddy et al., 2014), or have no significant
86 effect (Kammann et al., 2012) on CH₄ emissions from soils. Mechanisms are usually
87 only assumed or hypothesised and remain unclear. Meta-analysis is a useful tool for
88 comparing results across studies to reveal common response patterns. It facilitates
89 extrapolation of results and formulation of mechanistic hypotheses (e.g. within the
90 same soil conditions; or with the same biochar types) and thus increases the robustness
91 of extrapolations and predictions across systems.

92 The mechanisms by which biochar may affect soil CH₄ fluxes include sorption
93 of CH₄ to biochar's surfaces (Yaghooubio et al., 2014), and soil aeration by biochar
94 addition, which may increase diffusive CH₄ uptake (van Zwieten et al., 2010; Karhu et
95 al., 2011), as microbial CH₄ oxidation in upland soils is mostly substrate-limited (Castro
96 et al., 1994). However, in anoxic environments, the labile C pool of biochar may

97 function as methanogenic substrate, promoting CH₄ production (Wang et al., 2012).
98 Biochar has also been shown to promote methanotrophic CH₄ consumption at
99 oxic/anoxic interfaces in anoxic environments, lowering CH₄ emissions via the
100 “biofilter” function of CH₄ consumption (Feng et al., 2012; Reddy et al., 2014).

101 A recent work has also included meta-analysis of CH₄ emissions in response to
102 biochar application as part of a wider analysis (Song et al., 2016). However, the method
103 applied in their analysis does not allow the inclusion of negative fluxes (i.e. all CH₄
104 sinks) and thus was restricted in the conclusions that could be drawn. Here, we present
105 the first comprehensive meta-analytical investigation of the effects of biochar
106 application to soil in agricultural systems on CH₄ emissions drawing on studies with a
107 global distribution.

108

109 **2. Material and Methods**

110 *2.1. Data collection and categorisation*

111 The keywords “biochar” AND “methane” OR “CH₄” were entered into the search
112 engines of Scopus, Web of Science and Google Scholar to identify relevant studies for
113 inclusion in the meta-analysis. This led to identification of 62 studies, to a cut-off date
114 of 31st December 2014. Studies were vetted using inclusion criteria consisting of
115 studies: (i) using a randomised design; (ii) using replicated samples per treatment; and
116 (iii) containing a “treatment” and “control” such that the treatment was the same as the
117 control in all aspects apart from the inclusion of biochar. Only cumulative net CH₄
118 fluxes were included. Where only daily or seasonal fluxes were reported, corresponding
119 authors were contacted to ask for data on cumulative fluxes and means. When these
120 data were provided the studies were included; otherwise they were excluded. Of the

121 total studies, 42 met the inclusion criteria (Table S1), from which 189 pairwise
122 comparisons were extracted.

123 Data were collected from tables presented in manuscripts where possible, or
124 from figures using Plot Digitizer 2.6.6 (Huwaldt, 2015) or Web Plot Digitizer (Rohatgi,
125 2016), or from authors directly. Error bars were usually present in the form of standard
126 errors; standard deviations were back calculated from these when necessary. When no
127 measure of variance was available, corresponding authors were contacted to obtain such
128 information. Categorical information concerning biochar, soil and environmental
129 properties was also collected from manuscripts and recorded as auxiliary variables.
130 These can be found in the full database which is available in supplementary
131 information.

132 Auxiliary variables were grouped to facilitate cross-comparisons between
133 studies using the same groupings as Cayuela et al. (2014). These variables related to
134 *soil pH* grouped to <6 and 6-8 and >8 representing the optimum pH range for
135 methanogenesis and methanotrophy; biochar *feedstock*, grouped as Manure - manures
136 or manure-based materials from poultry, pig or cattle), Wood - oak, pine, willow,
137 sycamore and unidentified wood mixtures, Herbaceous - greenwaste, bamboo, maize
138 stover, straws; Biosolids – sewage sludge from water treatment plants and
139 lignocellulosic wastes - including rice husk, nuts shells, paper mill waste; *pyrolysis*
140 *temperature*, grouped as <450°C 450 – 600 and >600°C, *H:C_{org}*, grouped as <0.3 0.3-
141 0.5 >0.5; Brunauer, Emmett and Teller (BET) surface area ($\text{m}^2 \text{g}^{-1}$), grouped as <100,
142 100-500 and >500; water regime, water regime, grouped as Flooded (paddy soils and
143 studies conducted under continuous waterlogged conditions), Cycles (paddy soils
144 involving flooding-drying in which CH₄ emissions were measured during both the wet
145 and dry periods) and Non-Flooded (studies where flooding was not part of the

146 experimental setting); and N and Phosphorus (P) fertilization, grouped by rate for N,
147 $\leq 120 \text{ kg N ha}^{-1}$ and $>120 \text{ kg N ha}^{-1}$, and as P or No P if P fertilizer was applied or not
148 for P respectively.

149

150 2.2. *Meta-analytical metric*

151 Soils can function as both CH₄ sinks (negative values, uptake, consumption) and
152 sources (positive values, emissions). The notation of the flux direction follows the
153 convention by biogeochemists and takes the view from the atmosphere that gains or
154 loses the gas in question. Since it is not possible to take a logarithm of a negative
155 number, this precludes the use of the response ratio (calculated as the natural log of the
156 experimental mean over the control mean) as a metric for comparison between studies,
157 which is considered the preferred metric for ecological studies (Hedges et al., 1999).
158 Here we utilise the standardised mean difference metric “Hedge’s d ” for analysis
159 (Equation 1; Hedges and Olkin, 1985). This is a less biased indicator than “Hedge’s g ”
160 (Equation 2; Hedges, 1981; Hedges and Olkin, 1985). Note that this is a different
161 standardised mean difference metric to “Cohen’s d ” which was developed for
162 behavioural science (Cohen, 1988); Hedge’s d is less biased by small sample sizes
163 (Hedges and Olkin, 1985) and was used as this was the case for most studies included
164 in this meta-analysis.

165

166 Equation 1

$$d = \left(1 - \frac{3}{4(n-2)-1}\right) g$$

167

168 Where n is the total sample size on which g is based, and g is Hedge’s g as calculated
169 by Equation 2

170

171 Equation 2

$$g = \frac{\bar{x}_1 - \bar{x}_2}{s}$$

172

173 Where \bar{x}_1 and \bar{x}_2 are the experimental and control means and s is the pooled standard
174 deviation.

175 Here, experimental treatment refers to the treatment with biochar – controls are
176 samples that are the same in all aspects, including any other amendment, without
177 addition of biochar. A categorical random effects model was applied to d , with means
178 weighted by the inverse of the variance. Confidence intervals (95%; CIs) were
179 generated by bootstrapping (9999 iterations). To obtain a standardised mean effect size,
180 the effect size was then divided by an estimate of the standard deviation of the effect
181 sizes (Hedges and Olkin, 1985). Input data were arranged in Microsoft Excel 2010.
182 Calculations were performed using Metawin Version 2 statistical software (Rosenberg
183 et al., 2000).

184 The interpretation of the standardised mean effect size differs from the response
185 ratio as it cannot be expressed as a percent change in response of an experimental
186 treatment compared to a control. Rather, it is equivalent to a Z-score and as such
187 represents the number of standard deviations that the standardised mean of the
188 experimental treatment is from the standardised mean of the control. The effect of a
189 response variable can be considered significant if the 95% CI does not intersect the
190 standardised control mean (i.e. Z-value = 0). Groupings of auxiliary variables are
191 considered significantly different if their 95% CIs do not overlap.

192

193 2.3. *Interpretation of standardised mean effect size*

194 There is no rigorously applied framework for interpretation of standardised
195 means in terms of “effect sizes” because, unlike response ratios, they are probabilistic.
196 That is, they describe the probability that a sample drawn from the control treatments
197 would fall between the experimental mean and the control mean, assuming a normal
198 distribution. By convention, a large effect is indicated by $d > 0.8$, a moderate effect by
199 $d = 0.2 - 0.8$, and a small effect by $d = 0 - 0.2$ (Cohen, 1988; Gurevitch et al., 1992).
200 However, it is generally acknowledged that these terms are relative and likely
201 dependent on research area and methods (Hedges, 1981). A key point is that, using this
202 metric, an effect size of (for example) 0.2 for a category does not equate to an effect
203 size of 0.2 for categories in independent analyses presented in this paper, in absolute
204 terms. Only categories within individual analyses, as differentiated by the horizontal
205 dotted bars in Fig. 1 and 2, can be compared relatively (i.e. only within each category
206 does an effect size of 0.4 equate to twice the size of 0.2; comparisons between figures
207 are qualitative only). Further, small effect sizes (~ 0.2) may indicate significant changes
208 in cumulative GHG fluxes, in absolute terms, particularly if effects persist over the long
209 lifetime of biochar. Data are presented in two figures to allow use of different scale x-
210 axis only and does not represent any fundamental difference in analyses.

211 Interpretation of effect sizes here is further confounded by the CH₄ sink/source
212 flux direction in soils. A positive effect size implies a shift to the right on a scale going
213 from strong net sink (i.e. negative flux values) to strong net source (i.e. positive flux
214 values). However, it does not necessarily mean a change has occurred in the net
215 sink/source status of the soil. Rather, it signifies that either the net sink strength has
216 decreased, the soil has switched from sink to source, or that the net source strength has
217 increased – and vice versa for negative effect sizes.

218

219 *2.4. Control of biases*

220 We tested the effects of publication bias using the Fail-safe N technique (Orwin,
221 1983; Rosenthal and Rosnow, 1991). A weakness with meta-analyses of experimental
222 studies is that several experimental treatments are often compared to a single (identical)
223 control in a published study. This artificially increases the number of replicate pairs and
224 violates the assumption of independence that the effect size metric is based upon; the
225 controls are necessarily counted repeatedly in pairwise control versus experimental
226 treatment comparisons. Means of controlling for this bias (Borenstein et al., 2009;
227 Aguilera et al., 2013) often show little effect (van Groenigen et al., 2006; Gattinger et
228 al., 2012; Abalos et al., 2014; Skinner et al., 2014). Therefore, we here report results
229 from the analysis on the level of single comparisons.

230

231

232 3. Results

233 Figure 1 shows the effect of biochar application to soils under different
234 irrigation regimes. Biochar addition to *Flooded* soils (as part of their management
235 practice) significantly increased in CH₄ sink strength / reduced source strength
236 compared to *Flooded* soils without biochar application (Hedge's $d = -0.87$). Studies
237 reporting biochar additions to *Non-Flooded* soils showed an overall moderate but
238 significant decrease in the CH₄ sink strength / increase in source strength (Hedge's $d =$
239 0.65). Experiments in which irrigation was applied as *Cycles* of flooding and draining
240 did not show a significant response to biochar application.

241 Biochar application to acidic soils (i.e. with a pH <6) resulted in the strongest
242 effect size, causing a statistically significant increase in CH₄ sink strength / decrease in
243 source strength following biochar application (Hedge's $d = -1.56$; Fig. 1). Conversely,
244 addition of biochar to soils within the neutral pH range (i.e. 6-8) showed a statistically
245 significant decrease in CH₄ sink strength / increase in source strength (Hedge's $d =$
246 1.17). Application of biochar to soils with a pH greater than 8 did not show a
247 statistically significant response to biochar application.

248 Biochar effects on CH₄ flux interact with N fertilizer rate (Fig. 1). Application
249 of N fertilizers at rates less than 120 kg ha⁻¹ caused a strong and statistically significant
250 increase in CH₄ sink strength / decrease in source strength in the presence of biochar
251 (Hedge's $d = -3.1$). Applications of N fertilizer at higher rates showed no interaction
252 with biochar on soil CH₄ fluxes.

253 Biochars produced at high temperatures caused a statistically significant
254 increase in CH₄ sink strength / reduction in source strength following application to
255 soils (Hedge's $d = -1.3$; Fig. 1). Mid-temperature biochars (450-600°C) led to

256 significant reductions in CH₄ sink strength / increased source strength when applied to
257 soil (Hedge's $d = 0.67$).

258 In terms of interactions with feedstock source, biochar produced from *biosolids*
259 led to a statistically significant increase in sink strength / reduction in source strength
260 (Hedge's $d = -6.03$; Fig. 2). When produced from *Lignocellulosic waste*, biochar
261 significantly decreased the CH₄ sink strength / increased the source strength (Hedge's
262 $d = 0.74$). No other feedstock showed statistically significant effects on CH₄ fluxes.

263 No significant effects or differences between sub-groups were found for the
264 category BET Surface Area; however, there was an apparent trend whereby increased
265 BET surface area resulted in increasing sink strength / decreased source strength (Fig.
266 2).

267

268 **4. Discussion**

269 Using standardised mean differences as the meta-analysis metric precludes making
270 firm conclusions in terms of changes in CH₄ sink/source functioning. A statistical
271 approach based on measurements of net CH₄ fluxes alone does not enable
272 differentiation between changes in methanogenesis or methanotrophy. However, it can
273 identify the effect of biochar on the direction of net CH₄ fluxes (i.e. changes in overall
274 sink/source strength). It also allows identification of the key management practices and
275 soil and biochar properties which likely underlie the observed effects, and since the
276 "usual" CH₄ flux direction in flooded wetland or aerated upland soils is known in
277 general, the results provide first general insights into associated factors that need further
278 investigation.

279 Application of biochar to soil produced a range of effects on CH₄ fluxes across
280 studies, as expected. In most instances the "Grand Mean" (i.e. the mean response of all

281 studies combined) was not significantly different to the control. This result is most
282 likely due to contrasting responses (i.e. positive and negative effects on net CH₄ fluxes)
283 cancelling each other out when studies assigned to all functional categories were
284 combined. The data are unlikely to be significantly affected by publication bias, as
285 studies finding either a positive or negative result are equally publishable, and most
286 studies also investigated other factors such as N₂O fluxes and/or yield response. These
287 have been shown to have a positive response to biochar application (Cayuela et al.,
288 2014; Jeffery et al., 2011). As such, studies also investigating these metrics would have
289 an increased chance of publication of the “associated” CH₄ flux results.

290

291 *4.1. Irrigation management*

292 Biochar addition to soils that were flooded as part of their management practice
293 significantly increased CH₄ sink strength / reduced source strength compared to their
294 controls. Methanogenesis is an exclusively anaerobic process (Thauer, 1998). Here
295 (*Flooded*; Fig. 1), the change in CH₄ flux would likely equate to reduced net CH₄
296 emissions from flooded paddy soils, indicating that either the production decreased or
297 methanotrophy in the rhizosphere increased through influencing the
298 methanogenic/methanotrophic ratio of soils. Feng et al. (2012) reported that biochar
299 decreased the ratio of methanogenic archaea to methanotrophic bacteria. In flooded
300 soils, CH₄ consumption occurs at the aerated root interface where most CH₄ produced
301 in the surrounding anoxic sediment usually enters the aerenchymatic root-shoot rice
302 tissue, leaving the soil via this plant ‘chimney’. Thus, increased CH₄ oxidation at the
303 “biofilter” anoxic/oxic interface may explain the apparent CH₄ efflux mitigation
304 potential of biochar application to flooded soils observed in our meta-analysis. Other
305 studies in paddy soils (Liu et al., 2011; Singla et al., 2014) found no significant effects

306 on methanogenic archaeal diversity between biochar treated and non-treated soils (but
307 did not investigate CH₄ oxidizer communities).

308 Non-flooded (i.e. predominantly oxic) upland soils are an important sink for
309 CH₄ and are considered to contribute to approximately 15% of global CH₄ oxidation
310 (Powlson et al., 1997). Figure 1 suggests that biochar application may decrease net CH₄
311 oxidation by such soils. As intensively managed agricultural soils are relatively poor
312 sinks of CH₄, it is likely that the decrease in net CH₄ efflux from *Flooded* soils more
313 than counteracts any decrease in the net uptake from *Non-flooded* soils. Therefore,
314 biochar use in rice agriculture may contribute to reducing the C footprint of rice
315 production, which is usually worse than for example that of wheat production due to
316 the CH₄ emission burden.

317 Experiments in which irrigation was applied as *Cycles* of flooding and draining
318 did not show a significant response to biochar application. However, considerably
319 fewer pairwise comparisons contributed to this category: 14 compared to 56 for
320 *Flooded* and 85 for *Non-Flooded*. As such, there is reduced confidence in this result
321 evidenced by the relatively large error bars (Fig. 1).

322 All studies included in this analysis were conducted in managed systems: either
323 in the field or in controlled laboratory or greenhouse experiments. Currently, there is
324 no work in the published literature that has investigated biochar effects when applied
325 to natural wetlands, such as marshes, bogs and swamps, which can be significant
326 sources of CH₄ emission (Bubier and Moore, 1994). This represents an unknown area
327 of biochar research that may grow in importance as novel biochar applications are
328 sought and potentially the biochar load of these systems increases due to biochar
329 transport over time through waterways following erosion events (Jaffé et al., 2013).

330

331 4.2. Soil pH

332 Soil pH is one of the main environmental parameters that affects both
333 methanogenesis and methanotrophy (Hanson and Hanson, 1996; Semrau et al., 2010).
334 The optimum pH range of most methanogens ranges from 6 to 8 (Garcia et al., 2000),
335 thereby overlapping with the optimum pH range for methanotrophy, which also extends
336 to more acidic conditions (Le Mer and Roger, 2001; Semrau et al., 2010). Biochar
337 generally has a higher pH than the soil to which it is applied, thereby providing a liming
338 effect (Chidumayo, 1994; Yamato et al., 2006; Jeffery et al., 2011); the pairwise
339 comparisons of this meta-analysis have an average pH of 6.2 for soil and 9.6 for biochar.
340 As the optimum pH range for both methanogenesis and methanotrophy is similar, it
341 may be expected that raising the soil pH to within the optimum range would affect both
342 processes equally. However, we observed a significant increase in CH₄ sink strength /
343 decrease in source strength for acidic soils (Fig. 1). A potential explanation is that the
344 size and/or structure of methanotrophic communities may be more sensitive to rising
345 soil pH than that of methanogens. Experiments quantifying, for example, the
346 *mcrA/pmoA* ratios of soils are required to identify the cause underlying this observed
347 effect.

348 Another possible explanation for the large CH₄ mitigating effect of biochar in
349 acidic soils is related to Al³⁺ toxicity. Soils with a low pH are associated with increased
350 Al³⁺ solubility, which is highly toxic for methanotrophic bacteria (Tamai et al., 2007).
351 By increasing soil pH, biochar may reduce Al³⁺ release from cation exchange sites in
352 the soil, thereby reducing toxicity levels for methanotrophs. A further analysis of initial
353 soil pH effects on CH₄ fluxes, utilising a cut off at a pH of 5, the threshold above which
354 Al³⁺ availability strongly decreases, provides further evidence for this explanation (Fig.
355 S1). Biochar applied to soils with a pH <5 showed a significant increase in sink strength

356 / reduction in source strength compared to soils with a pH >5. When biochar was
357 applied to soils already above this threshold, no significant effect on CH₄ flux was
358 observed. This hypothesis is in line with the literature on this topic (reviewed in
359 Dunfield 2007; e.g. Sitaula & Bakken 2001) but more empirical studies are required to
360 confirm or reject this hypothesised mechanism.

361

362 4.3. N Fertilizer

363 Figure 1 suggests that when biochar is applied with <120 t ha⁻¹ N fertilizer, it
364 can reduce CH₄ fluxes, while it has no effect when applied with >120 t ha⁻¹ N. However,
365 the effect of N fertilizer type and application rate on CH₄ flux is, in general, highly
366 controversial. In soils where methanotroph N supply is not limiting to growth and
367 activity, it is generally expected that the addition of NH₄⁺-containing or delivering
368 fertilizers will lead to decreased CH₄ oxidation due to competitive exclusion of CH₄ at
369 binding sites by NH₄⁺ (Bédard and Knowles, 1989; Sylvia et al., 2005). However, this
370 effect is rate dependent; smaller amounts of N tend to stimulate CH₄ uptake while larger
371 amounts tend to inhibit uptake into the soil (Aronson and Helilker, 2010). Despite this
372 general rule, in severely N-limiting environments, the addition of an N source, even
373 NH₄⁺ which may also competitively inhibit CH₄ oxidation (Bedard and Knowles 1989;
374 Gullede et al., 1997), can lead to an increase in CH₄ oxidation due to an increase in
375 methanotrophic biomass (Bodelier et al., 2000; Nazaries et al., 2013). The switch
376 between stimulation and inhibition of CH₄ uptake has been reported to occur at between
377 100 kg N ha⁻¹ (Aronson and Helilker, 2010) and 140 kg N ha⁻¹ (Banger et al., 2012).
378 As such, we set the threshold for our analysis to the mid-point between these studies -
379 120 kg N ha⁻¹ (Fig. 1). This analysis shows that when N is applied above this threshold
380 there is no significant difference between the experimental treatments (with biochar)

381 and the controls (without biochar). When biochar is applied with levels of N below the
382 threshold a significant difference is observed between the experimental treatments
383 (with biochar) and the controls (without biochar) with increased sink strength/ reduced
384 source strength being observed following biochar application with N fertilization rates
385 below 120 k N ha⁻¹. The mechanism for this response pattern when biochar is applied
386 with low N rates remains unclear and warrants further investigation

387

388

389

390 *4.4. Pyrolysis temperature*

391 Biochars produced at high temperatures caused a statistically significant
392 increase in CH₄ sink strength / reduction in source strength following application to
393 soils (Fig. 1). High temperature biochars are characterised by fewer labile compounds
394 remaining on the surface of biochar particles, and so introduce less microbial substrate
395 than lower temperature biochars when applied to soil (Brunn et al., 2011).

396 Reduced H:C_{org} ratios in high temperature biochars indicate increased
397 aromaticity, which is associated with the reducing effect of biochar on N₂O emissions
398 (Cayuela et al., 2015). However, we did not find any relationship between H:C_{org} and
399 CH₄ fluxes from soil (Fig. S2).

400 Mid-temperature biochars (450-600°C) led to significant reductions in CH₄ sink
401 strength / increased source strength when applied to soil (Fig. 1). The majority (73%)
402 of the studies that used mid-temperature biochar were performed on non-flooded soils.
403 This means that there is a confounding effect: it may be that the effect observed here is
404 due to either biochar properties or soil water management - it is not possible to
405 distinguish between the two with the analysis used for this study.

406

407 4.5. Feedstocks

408 In general, the feedstock from which biochar was produced did not lead to
409 significantly different effect on CH₄ flux, with the exception of *biosolids* (Fig. 2). The
410 effect size for biochar produced from biosolids is remarkably large (Hedges, 1981), as
411 are the associated confidence intervals. This may be exacerbated by the low number of
412 pairwise comparisons on which the statistic is based; all of the four pairwise
413 comparisons were drawn from one study (Khan et al., 2013). The biochar used for this
414 study was produced from sewage sludge (here grouped as *Biosolids*; according to
415 Cayuela et al., 2014) and was applied to very acidic soil (i.e. pH = 4.02). Possible
416 mechanisms, as discussed above, include potential changes in the size and/or structure
417 of methanotrophic communities, or potentially reduced Al³⁺ toxicity effects. In
418 addition, the effect may also be partly due to the high sulphur content of this feedstock
419 (5.3% dry weight). This hypothesis is consistent with previous results that showed
420 decreased CH₄ emissions when ammonium sulphate was used as a fertilizer compared
421 to urea (Bufogle et al., 1998).

422 Biochar produced from *Lignocellulosic waste* led to a significantly decreased
423 CH₄ sink strength / increased source strength. The mechanism underlying this effect
424 remains unclear and warrants further research.

425

426 4.6. BET Surface Area

427 Biochar production temperature and the Brunauer, Emmett and Teller (BET)
428 surface area of biochars have been shown to be positively correlated (Ronsse et al.,
429 2013; Kambo and Dutta, 2015). This suggests that adsorption of CH₄ to the surface of
430 biochars (Sadasivam and Reddy, 2014) may also be responsible for the reduced flux in

431 the high temperature biochars (Fig. 1). However, this characteristic is often not reported
432 in biochar studies, which hinders investigation of this potential mechanism. It appears
433 that there is a trend whereby increased BET surface area results in decreased CH₄ flux
434 (Fig. 2). However, the data are highly variable in the highest category (>500) with little
435 confidence in the mean value owing to the low number of pairwise comparisons on
436 which this statistic is based (n = 3). More studies using high surface area biochars, or
437 systematically varying BET, are needed to investigate the importance of CH₄ or
438 inhibitory N adsorption onto biochar as a mechanism underlying observed reductions
439 in CH₄ fluxes.

440

441 **5. Conclusions**

442 Evidence presented in this study shows that biochar does have the potential to
443 mitigate CH₄ emissions from soil, particularly from paddy fields and/or acidic soils that
444 use periods of flooding as part of their management regime. However, addition of
445 biochar to neutral or alkaline soils that do not have periods of flooding, may have the
446 potential to decrease the CH₄ sink strength of those soils. These results indicate that soil
447 and biochar properties, as well as management conditions, must be considered to
448 maximise biochar's potential to mitigate CH₄ emissions and minimise trade-offs.

449 This meta-analysis highlights the importance of reporting key functional
450 characteristics of biochar properties. Biochar pH has been shown to be highly pertinent
451 for predicting response of some ecosystem functions to biochar application, in both this
452 current study and previous studies (Jeffery et al., 2011; Sagrilo et al., 2015). Other
453 functional characteristics (or proxies thereof) such as the molar H:C_{org} ratio are
454 becoming more recognised as effective predictors (Cayuella et al., 2015). Here we show
455 that BET surface area may be an important functional characteristic in terms or

456 predicting CH₄ flux mitigation potential of biochar. However, insufficient numbers of
457 experiments have reported the characteristic to draw firm conclusions. It is vital that
458 biochar researchers characterise and report functional characteristics of their biochars
459 wherever possible.

460 Finally, it is apparent that trade-offs are inevitable and clear goals are necessary
461 before effective advice can be offered to land managers and policy makers (Jeffery et
462 al., 2015). For example, low temperature, slow pyrolysis maximises biochar production
463 (Sohi et al., 2010) and thereby also C sequestration potential. However, evidence
464 presented in this study shows that high temperature biochars are more effective at
465 mitigating CH₄ emissions (the same applies for N₂O, Cayuela et al., 2015). Which one
466 has the greatest potential to mitigate climate change thus remains to be determined and
467 will require life cycle assessment approaches. However, market forces are likely to
468 make the former more attractive until the full environmental costs of production are
469 included as part of agricultural products.

470

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481

482

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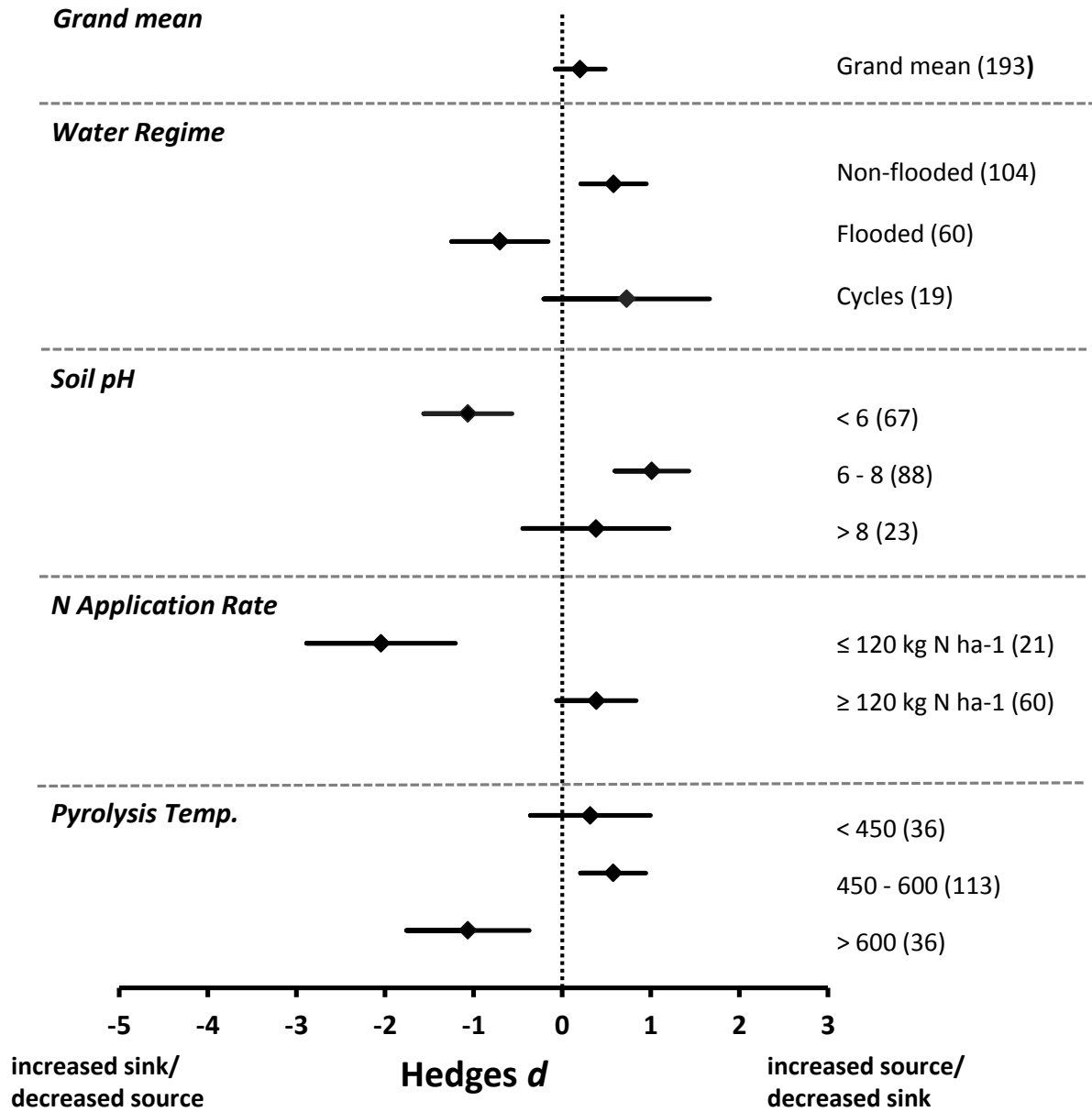
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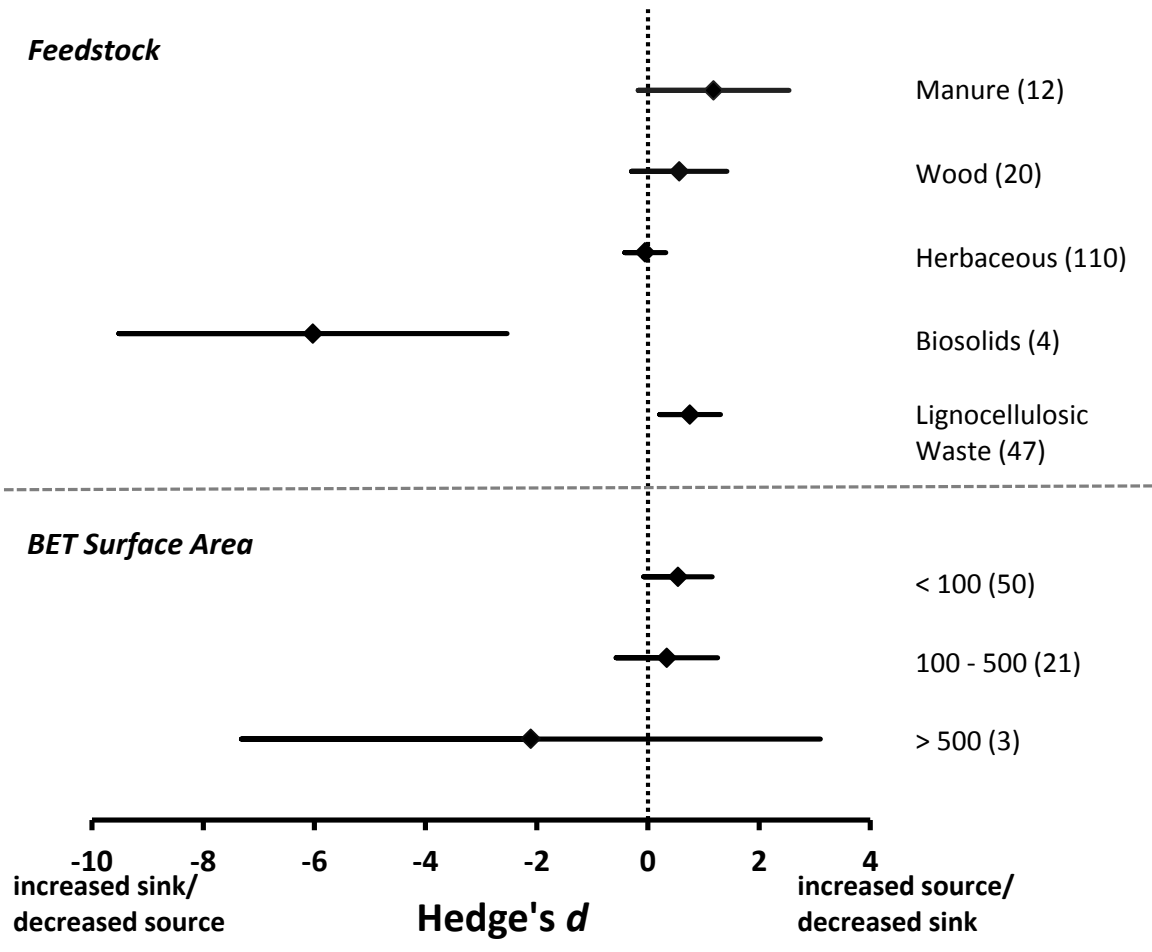
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Figure 1. A forest plot of Hedge's d calculated from published literature grouped by experimental water regime, soil pH pre-biochar amendment, N fertilizer application rate and biochar pyrolysis temperature. Points show means, bars show 95% confidence intervals. The numbers in parentheses indicate the number of pairwise comparisons on which the statistic is based. (For an explanation of the Hedge's d metric see text).

Figure 2. A forest plot of Hedge's d calculated from published literature grouped by biochar feedstock type and BET (Brunauer, Emmett and Teller) surface area. Points show means, bars show 95% confidence intervals. The numbers in parentheses indicate the number of pairwise comparisons on which the statistic is based. (For an explanation of the Hedge's d metric see text).





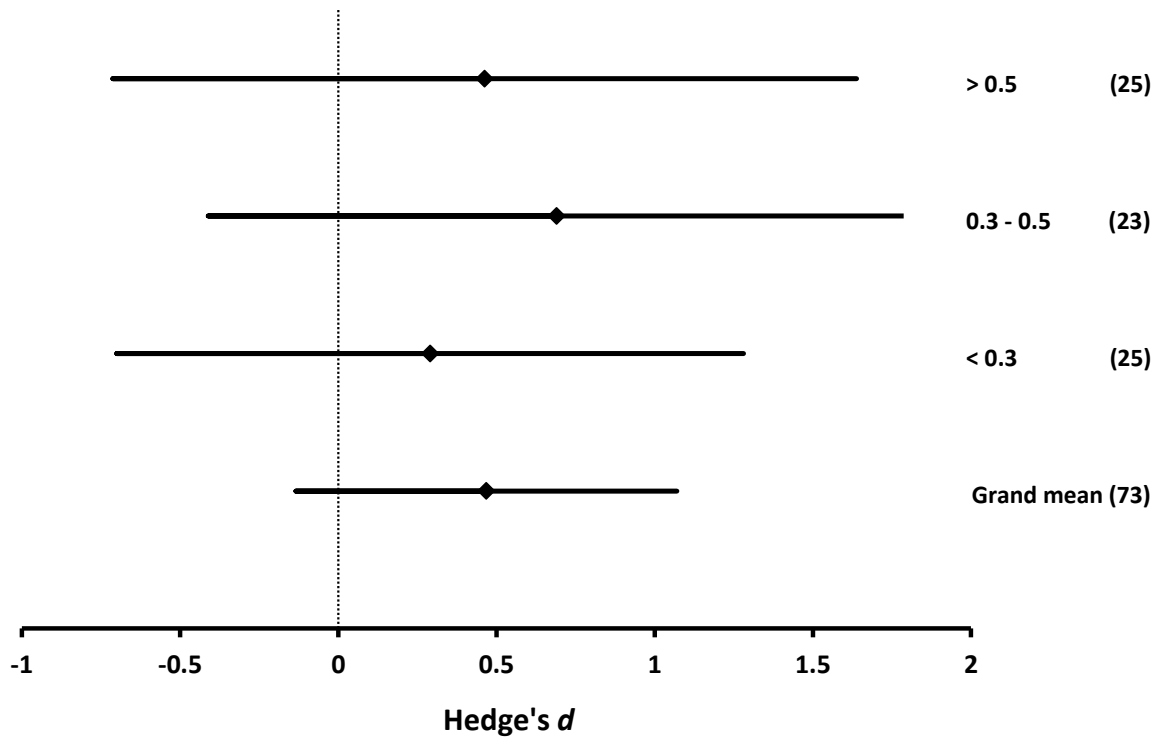


Figure S1. A forest plot of Hedge's d calculated from published literature grouped by pre-biochar amendment soil pH. The pH 5 threshold is applied to investigate the potential effects of aluminium bioavailability/toxicity. Points show means, bars show 95% confidence intervals. The numbers in parentheses indicate the number of pairwise comparisons on which the statistic is based. (For an explanation of the Hedge's d metric see text).

Figure S2. A forest plot of Hedge's d calculated from published literature grouped by H:C_{org}. Points show means, bars show 95% confidence intervals. The numbers in parentheses indicate the number of pairwise comparisons on which the statistic is based. (For an explanation of the Hedge's d metric see text).

Table S1. A list of the studies included in the meta-analysis database.

<i>Reference</i>	<i>Country</i>	<i>Soil pH</i>	<i>Water regime</i>	<i>BET surface area</i>	<i>Pyrolysis temperature</i>	<i>H:Corg</i>	<i>Biochar feedstock</i>
Castaldi et al., 2011	Italy	<6	Non-flooded	-	450-600	-	Wood
Dong et al., 2013	China	<6	Flooded	<100	>600	0.3-0.5, >0.5	Herbaceous
Feng et al., 2012	China	6-8	Flooded	-	<450, 450-600	-	Herbaceous
Fungo et al., 2014	Kenya	6-8	Non-flooded	-	<450, 450-600	0.3-0.5, >0.5	Herbaceous
Khan et al., 2013	China	<6	Flooded	<100	450-600	-	Biosolids
Liu et al., 2011	China	<6	Flooded	-	>600	0.3-0.5, >0.5	Herbaceous
Scheer et al., 2011	Australia	<6	Non-flooded	-	450-600	>0.5	Herbaceous
Singla and Inubushi, 2014	Japan	<6	Flooded	-	<450	-	Manure
Spokas et al., 2009	USA	6-8	Flooded	<100	450-600	0.3-0.5	Lignocellulosic
Wang et al., 2012	China	6-8	Flooded	-	450-600	-	Lignocellulosic
Wu et al., 2013	Canada	<6	Non-flooded	-	450-600	0.3-0.5	Herbaceous
Xie et al., 2013	China	6-8	Cycles	-	<450	-	Herbaceous
Zhang et al., 2012	China	6-8	Cycles	<100	450-600	-	Herbaceous
Zhang et al., 2012	China	>8	Non-flooded	<100	450-600	-	Herbaceous
Zhang et al., 2010	China	6-8	Flooded	<100	450-600	-	Herbaceous
Zheng et al., 2012	USA	6-8, >8	Non-flooded	100-500	450-600	<0.3	Lignocellulosic
Liu et al., 2014	China	<6	Flooded	-	450-600	-	Herbaceous
Pandey et al., 2014	Vietnam	-	Cycles	-	-	-	Herbaceous
Schimmelpfennig et al., 2014	Germany	<6	Non-flooded	>500	>600	<0.3	Herbaceous
Shen et al., 2014	China	<6	Flooded	-	-	-	Herbaceous
Singla et al., 2014	Japan	6-8	Flooded	-	<450	>0.5	Manure
Zhao et al., 2014	China	6-8	Flooded	-	450-600	-	Herbaceous
Zhang et al., 2014	Canada	6-8	Non-flooded	-	>600	-	Lignocellulosic

Zhang et al., 2013	China	6-8	Cycles	<100	450-600	-	Herbaceous
Jia et al., 2012	China	<6	Non-flooded	-	<450	-	Herbaceous
Ali et al., 2013	Bangladesh	6-8	Cycles	-	450-600	-	Lignocellulosic
Spokas et al., 2013	USA	6-8	Non-flooded	<100	450-600	<0.3, 0.3-0.5	Lignocellulosic
Angst et al., 2014	USA	6-8	Non-flooded	<100	450-600	>0.5	Lignocellulosic
Case et al., 2014	UK	6-8	Non-flooded	-	<450	-	Lignocellulosic
Li et al., 2013	China	<6, >8	-	<100	<450, 450-600	-	Herbaceous
Ly et al., 2014	Cambodia	<6	Flooded	-	450-600	-	Herbaceous
Stewart et al., 2013	USA	>8	Non-flooded	100-500	450-600	<0.3	Lignocellulosic
Watanabe et al., 2014	Japan	-	Non-flooded	-	>600	-	Herbaceous
Karhu et al., 2011	Finland	-	Non-flooded	<100	<450	-	Lignocellulosic
Troy et al., 2013	Ireland	6-8	Non-flooded	-	>600	-	Manure
Mukherjee et al., 2014	USA	6-8	Non-flooded	100-500	>600	-	Wood
Thomazini et al., 2015	USA	<6, 6-8	Non-flooded	<100	450-600	-	Wood
Vu et al., 2015	Vietnam	<6	Cycles	-	-	-	Herbaceous
Zhang et al., 2015	China	6-8	Cycles	<100	450-600	-	Herbaceous
Li et al., 2015	China	<6	Non-flooded	<100	<450	-	Herbaceous
Lin et al., 2015	China	>8	Non-flooded	-	<450	-	Herbaceous
Yoo et al., 2015	Korea	<6, 6-8	Flooded	-	<450, >600	0.3-0.5, >0.5	Herbaceous
Yu et al., 2013	China	<6	Non-flooded, Flooded	-	-	>0.5	Manure

Supplementary Table S2. Between-group heterogeneity (Q_b), within-group heterogeneity (Q_w) and total heterogeneity (Q_t).

	Q_b	Q_w	Q_t
<i>Water regime</i>	16.55***	342.32***	358.88***
<i>Soil pH</i>	40.92***	324.99***	365.92***
<i>N application rate</i>	27.76***	117.67**	145.44***
<i>Pyrolysis temperature</i>	17.98***	349.95***	367.93***
<i>Feedstock</i>	41.46***	348.31***	389.77***
<i>BET Surface Area</i>	5.41	148.12***	153.54***
<i>Soil pH - cut off at pH 5</i>	6.45*	350.41***	356.86***
<i>H:C_{org} molar ratio</i>	0.31	180.76***	181.07***
<i>Soil texture</i>	7.76*	238.73***	246.49***

* $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$