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INFLUENCE OF CAPILLARY HETEROGENEITY ON LEAKAGE OF CO₂ FROM A
BOREHOLE

A Thesis
presented in partial fulfillment of requirements
for the degree of Masters of Engineering Science
in the Department of Geology and Geological Engineering
The University of Mississippi

by

FRANK G. ROECKER

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ABSTRACT

We used a modified invasion percolation (MIP) model to examine the effect of capillary heterogeneity, buoyancy forces, and viscous forces on the surface area and saturation of a CO₂ plume leaking into a shallow aquifer. The purpose of this study is to provide a better understanding of how CO₂ migrates from a borehole, which is essential in implementing effective simulation and monitoring regimes to accurately detect CO₂ leakage from sequestration sites. The MIP model approach will simulate invasion of a light non-wetting fluid (e.g., CO₂) into a medium initially saturated with a dense wetting fluid (water). The style of capillary heterogeneity, the strength of buoyancy and viscous forces, and the size of the CO₂ source were systematically varied yielding 168 different simulation scenarios. We find that the interplay between capillary heterogeneity, buoyancy forces, and viscous forces controls the surface area and saturation of gaseous CO₂ leaking into an aquifer system. In unstructured systems with the absence of buoyancy and viscous forces, the CO₂ surface area and saturation are relatively large. In most cases, the CO₂ surface area decreases in weakly stratified systems, and as stratification increases, the CO₂ surface area increases and the CO₂ saturation decrease. Buoyancy forces stretch the invading CO₂ into a narrower structure, resulting in a higher surface area and a lower saturation. Our model implements weak viscous forces which cause CO₂ to pool around the leak source until, at a radial distance, either buoyancy or capillary forces begin to dominate CO₂ invasion. The dissolution rate of gaseous CO₂ into groundwater is proportional to the surface area of the CO₂ phase. Our study shows that variations in the style of capillary heterogeneity and

the strength of buoyancy and viscous forces can lead to large differences in CO₂ dissolution rates.

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I. INTRODUCTION

Carbon capture and sequestration in deep geologic formations is emerging as a possible way to reduce greenhouse gas concentrations. Proposed sequestration sites are commonly located in depleted oil and gas fields, saline aquifers, or deep un-mineable coal seams (*Ide et al., 2006; Folger, 2009, Birkholzer et al., 2009*). On closure, most CO₂ sequestration sites will contain numerous well bores that penetrate cap-rocks of reservoirs and CO₂ sequestration horizons. Degradation of well-bore casing, plugs, and cements can create pathways for buoyant CO₂ to migrate towards the surface (*Ide et al., 2006*). The consequences of CO₂ leakage from a storage reservoir include the contamination of shallow potable water sources and the direct or delayed release of CO₂ gas back into the atmosphere, negating the long term effects of CO₂ sequestration.

Many researchers (e.g., *Ide et al., 2006; Bachu and Celia, 2009; Kopp et al., 2010; Tao et al., 2011; Zhang and Bachu., 2011*) have identified leakage along well bores as the most likely pathway for CO₂ leakage from a storage reservoir. These workers have developed models that explore CO₂ leakage from abandoned wells [*Kopp et al., 2010; Nordbotten et al., 2005; Humez et al., 2011*], along the annular cement [*Bachu and Bennion, 2009; Viswanathan et al., 2008*], and through multiple geological layers and multiple leaky wells [*Nordbotten et al., 2009*]. Others have focused on how the well-bore cement reacts in a CO₂ saturated environment [*Carey et al., 2007; Carey et al., 2010; Huerta et al., 2011; Geloni et al., 2011*]. Upon leakage from a well bore, escaped CO₂ may come into contact with shallow aquifers and migrate through the porous medium away from the well bore (e.g., *Oldenburg et al., 2010*). CO₂ dissolution into

groundwater can lower pH, altering the aquifer geochemistry [Wang and Jaffe., 2004; Zheng et al., 2009; Kharaka et al., 2010; Keating et. al., 2010; Peter et al., 2011] and possibly triggering movement of trace metals [Apps et al., 2010; Peter et al., 2011].

Direct detection of CO₂ leakage in the shallow subsurface can be difficult due to large variations in natural background CO₂ fluxes (Lewicki et al., 2007). Other workers [Romanak et al, 2010; Romanak et al, 2012] have suggested that well-bore leaks into aquifers may be identified through long-term monitoring of groundwater. Measuring accurate amounts of dissolved CO₂ in groundwater is difficult, (Newell et al., 2008), but solute plumes containing newly dissolved minerals may be more readily identified (Benson and Cole, 2008). The character of solute plumes resulting from CO₂-driven alteration of aquifer materials will vary with the local CO₂ dissolution rate and the total mass of CO₂ within the aquifer, which control the evolution of aquifer pH near leaking boreholes. The dissolution rate of CO₂ into groundwater is limited by the surface area of the CO₂-groundwater interface, and the mass of CO₂ within the aquifer is a function of the CO₂-phase saturation. CO₂-phase saturation and surface area are directly influenced by the capillary heterogeneity within the aquifer, the magnitude of local buoyancy forces, and the influence of viscous forces arising from the displacement of groundwater by CO₂.

We examine the effect of capillary heterogeneity, buoyancy forces, and viscous forces on the surface area and saturation of a CO₂ plume in a shallow aquifer, resulting from a slow well-bore leak. The displacement of the denser wetting phase (e.g., water) by a lighter non-wetting phase (e.g., CO₂ gas) is simulated using a quasi-3D, Modified Invasion Percolation (MIP) model [e.g., Ioannidis et al., 1996; Glass et al., 2001; Holt, 2003] that includes capillary forces, buoyancy forces, and a first-order approximation of viscous forces. Over 8,400 simulations are

employed in a Monte Carlo approach used to quantify the influence of and uncertainty due to variations in capillary heterogeneity, buoyancy forces, and viscous forces on CO₂ surface area and saturation. The style of capillary heterogeneity is systematically varied from unstructured to stratified to mimic the range of capillary heterogeneity found in nature. Similarly, the strength of buoyancy and viscous forces are also varied, and the influence of the leak source size is evaluated. All simulations are non-dimensional and can be scaled using the gas-entry pressure of a porous medium.

We find that the interplay between capillary heterogeneity, buoyancy forces, and viscous forces controls the surface area and saturation of gaseous CO₂ derived from well-bore leaks in aquifer systems. In stratified systems, the CO₂ surface area increases and the CO₂ saturation decreases with increasing stratification and buoyancy forces, both of which focuses CO₂ invasion into narrow pathways. CO₂ surface area decreases and CO₂ saturation increases in the presence of viscous forces, which leads to pooling of CO₂ near the source. As the leak source size increases, CO₂ invades along more pathways near the source, increasing the CO₂ surface area and saturation.

II. MODIFIED INVASION PERCOLATION MODEL

An MIP approach is used to simulate invasion of a light non-wetting fluid (gas) into a medium initially saturated with a dense wetting fluid (water) under the combined influence of capillary, buoyancy, and viscous forces. Compressibility and dissolution of CO₂ phase, trapping, and other dynamic effects, such as fragmentation and pulsation [e.g., *Wagner et al.*, 1997; *Vedvik et al.*, 1998; *Glass and Yarrington*, 2003], are neglected here, as our focus is on the broad factors affecting CO₂ dissolution rates and concentrations in aquifer systems. The current MIP differs from Invasion Percolation (IP) [e.g., *Wilkinson and Willemsen*, 1983] through the definition of macro, near-pore scale capillarity. Here, individual pore throats and necks are not considered. Instead, a near-pore scale block is defined and characterized by a local threshold spanning pressure (a local block-scale breakthrough pressure) that represents the behavior of the subscale network (*Holt et al.*, 2003). The model domain is discretized into an array of grid blocks with assigned spanning pressures. The invasion pressure for each block is determined by the sum of the spanning pressure, gravity forces, and viscous forces:

$$P_I^* = P_S^* + \Delta\rho g Z^* + \Delta P_v^* \quad (1)$$

where P_S^* is the capillary spanning pressure ($ML^{-1}T^{-1}$), $\Delta\rho = \rho_{\text{defender}} - \rho_{\text{invader}}$ is the difference in densities of the defending phase and invading phase (ML^{-3}), g is the acceleration due to gravity in the Z-direction (LT^{-2}), Z^* is the coordinate in the Z-direction (L), and $\Delta P_v^* = (P_v^{\text{defender}^*} - P_v^{\text{invader}^*})$ is the difference in viscous pressure ($ML^{-1}T^{-1}$). Dividing equation (1) by the average spanning pressure $\langle P_S^* \rangle$, our dimensionless invasion pressure becomes:

$$P_l = \frac{P_s^*}{\langle P_s^* \rangle} + B_o Z + \frac{(P_v^{d*} - P_v^{i*})}{\langle P_s^* \rangle} \quad (2)$$

where $\langle P_s^* \rangle$ is the average spanning pressure ($\text{ML}^{-1}\text{T}^{-1}$), the dimensionless Bond Number is defined as:

$$B_o = \frac{\Delta\rho g \Delta Z^*}{\langle P_s^* \rangle} \quad (3)$$

ΔZ^* is the grid block length in the Z-direction (L). The dimensionless Z-coordinate is defined as:

$$Z = \frac{Z^*}{\Delta Z^*} \quad (4)$$

The Bond Number reflects the strength of the buoyancy forces. For the simulations presented here, the Bond Number was systematically varied between 0, 0.1, and 1 to represent increasing buoyancy forces.

The viscous pressure in the defending fluid was set at zero for all simulations, to simulate static, non-flowing groundwater conditions. The viscous pressure in the invading fluid is calculated by a first-order approximation for steady flow from a point source in spherical coordinates with the reduced equation being:

$$\frac{P_v^{i*}}{\langle P_s^* \rangle} = VSN \left(\frac{1}{r} - \frac{1}{R} \right) \quad (5)$$

where VSN is the dimensionless viscous scaling number, r is the dimensionless grid-block radial coordinate (-), and R is the dimensionless radius of influence (-). An R value of 1000 was kept constant for all simulations. We selected VSN values of 0 and 250; a VSN value of 250 is equivalent to a CO_2 volumetric discharge rate of $0.03 \text{ m}^3/\text{d}$ in a typical sandstone and $3 \times 10^{-4} \text{ m}^3/\text{d}$ in a typical mudrock or shale.

Correlated spanning-pressure fields are generated using an FFT approach (e.g., Robin et al., 1993). First, a two-dimensional, correlated random field (512 by 512) with a standard-normal distribution is generated, then the value at each nodal location is mapped to an associated probability, and finally the nodal spanning pressure is determined from the nodal probability (P) using the van Genuchten (1980) pressure-saturation model:

$$P_s = \frac{1}{\alpha} \left\{ \left[\left(\frac{1}{P} \right)^{\frac{1}{1-N}} \right] - 1 \right\}^{\frac{1}{N}} \quad (8)$$

where P_s is the spanning pressure, α is a model parameter related to the air-entry pressure, N is a model parameter associated with the pore-size distribution. Alpha (α) values of 1 and N values of 1.5 were kept constant for all simulations. The style of capillary heterogeneity in each field is defined by a fixed correlation ratio, the ratio of the horizontal to vertical correlation lengths. Correlation ratios of 1, 10, 20, 30, 50, 100, 500, and 1000 were used to reflect increasing horizontal stratification. Values of 1, 10, and 50 are used as the dimensionless vertical correlation lengths (Figure 1). As the vertical correlation length increases, the horizontal correlation lengths are increased to maintain specified correlation ratios.

During simulations, non-wetting phase invasion proceeds from a source on the left edge of the domain towards the top or right edge domain boundary according to an IP algorithm that implements equation (2). We apply a network connectivity of 8 (communication is allowed with all surrounding blocks) to approximate three-dimensional behavior within a two-dimensional network (e.g., Holt et al., 2003). The IP algorithm sorts all the invadable nodes along the wetted surface and invades the block with the lowest invasion pressure. This process continues until the blocks along the active interface reach either the top or right edge boundaries, the percolation threshold. In the absence of buoyancy and viscous forces ($B_o = 0$ and $\Delta P_v^* = 0$), standard

percolation will exactly reproduce the drainage pressure-saturation curve given by equation (8) (Holt *et al.*, 2003). However, when $\Delta P_v^* \neq 0$ and $B_o \neq 0$, this is not the case.

Our model simulates leakage from a point source and a line source along the left edge of the domain. The model defines an initial wetted surface of CO₂ that acts as the leak source. The wetted surface is defined as all the possible nodes that are in contact with the growing cluster. The initial wetted surface for a point source allows invasion of CO₂ only into the bottom-most node on the left-hand side of the domain. This small source size is analogous to CO₂ escaping through a fracture in the well-bore casing or cement. For the larger leak source, the bottom 102 nodes (10% of the domain size) along the left-hand side of the domain are considered invadable at the start of the simulation. This arrangement simulates CO₂ leaking from a well bore through degraded annular cement. Studies have shown that well-bore cement can degrade after 200 years to allow potential fluid flow analogues to that through a silty sand (EPA, 2009).

Simulations including viscous forces were only conducted for a point source in random spanning pressure fields with a vertical correlation length of 1 grid block, as more complicated source sizes and styles of heterogeneity would lead to greater errors in our first-order approximation.

In the following discussion, the term “surface area” is defined as the surface area per unit volume of invaded.

III. RESULTS

We focus on the final CO₂ surface area (ratio of the CO₂-water surface area to CO₂ volume) and the final CO₂ saturation at percolation to either the right or top boundary. 8,400 simulations were run to evaluate the influence of capillary heterogeneity, buoyancy forces, and viscous forces on the CO₂ surface area and saturation. 168 different combinations were created by varying the correlation ratio, Bond Number, strength of first-order viscous forces, the Z-correlation length, and leak source size (Table 1). 50 simulations were run for each combination. CO₂ surface area and saturation results are summarized using the average and standard deviation for each 50-simulation set.

Figures 2 and 3 show the influence of varying styles of stratification on the evolution of the CO₂ phase in the absence of buoyancy and viscous forces for the case of a point source and a line source, respectively. Both source sizes show similar results. For unstructured media (Z-correlation length of 1 and correlation ratio of 1), the CO₂ surface area is fairly large (~1) as the CO₂ phase distribution is controlled by the random distribution of spanning pressures. The addition of structure, in the form of either vertical or horizontal correlation, decreases the surface area (~0.6), as invasion fills adjacent grid blocks, which have similar spanning pressures. As the degree of stratification increases, the CO₂ surface area increases, reflecting less vertical overlap in the CO₂ phase. As the correlation ratio becomes large, invasion tends to proceed similarly regardless of the vertical correlation length, and the average surface areas for all vertical correlation lengths become similar. The average surface area from a line source (Figure 3) (~1.5) is greater than that of a point source (Figure 2) (~1.2) at large correlation ratios, as the i

Invading CO₂ takes multiple pathways from the source. Average CO₂ saturation increases with increasing Z-correlation length, as regions of similar spanning pressure become completely filled with the invading CO₂. The average CO₂ saturation decreases with increasing correlation ratio, because CO₂ invasion becomes focused to fewer strata and the percolation threshold is reached in fewer invasion steps.

Figures 4 and 5 display how intermediate buoyancy forces ($B_O = 0.1$) affect CO₂ phase invasion from a point and a line force respectively. Both source sizes show similar results. For unstructured media (Z-correlation lengths of 1 and a correlation ratio of 1), the CO₂ surface area is fairly large (~1.3) and CO₂ saturation is low (~0.02), as the CO₂ phase is being stretched vertically into a narrow structure by buoyancy forces. When the vertical correlation ratio increases from 1 to 10 in un-stratified media, CO₂ fills adjacent grid blocks with similar spanning pressures, decreasing CO₂ surface area (~0.6) and slightly increasing CO₂ saturation (~0.03). When the vertical correlation length reaches 50, however, stronger vertical correlation of spanning pressures allows buoyancy forces to dominate the local capillary heterogeneity, leading to slightly higher surface areas.

For both leak source sizes (Figures 4 and 5), buoyancy forces become less influential over the CO₂ plume as the medium becomes more stratified, and the percolation threshold is reached on the right boundary. Instead of a narrow, vertical structure, the final CO₂ phase distribution reflects trapping beneath horizontal capillary barriers of higher spanning pressure. When the vertical correlation length is 1, buoyancy forces are strongest, leading to vertical stacking of CO₂-invaded layers, lower CO₂ surface areas, and higher CO₂ saturations. Weakly stratified systems (correlation ratio of 10) show lower surface areas (~0.7) and higher saturations (~0.05) than the unstructured systems, as CO₂ fills more vertically adjacent gridblocks. As

stratification increases, CO₂ invasion proceeds along an increasingly horizontal pathway, increasing CO₂ surface area and decreasing CO₂ saturation. At large correlation ratios, the average surface area from a line source (Figure 5) (~1.5) is greater than that from a point source (Figure 4) (~1.2), because CO₂ can take multiple pathways from the source.

Figures 6 and 7 show how strong buoyancy forces ($B_o = 1$) affect CO₂ invasion from a point source and from a line source, respectively. Both source sizes show similar results. In unstructured media (Z -correlation lengths of 1 and a correlation ratio of 1), the CO₂ surface area is very large (~2.2) and the CO₂ saturation is very low (~0.004), because the CO₂ phase is being stretched vertically into a very narrow structure. In weakly stratified systems, buoyancy forces remain strong relative to capillary forces, and vertical stacking of CO₂-invaded layers causes lower surface areas (~0.8 for $Z = 1$) and slightly higher saturations (~0.025 for $Z = 1$). This process is particularly important for simulations run with a vertical correlation length of 1, where strong buoyancy forces lead to percolation at the upper boundary and increased vertical stacking to correlation ratios of 30 and 50. As the correlation ratios increase, horizontal capillary forces dominate the system, vertical stacking is reduced, and the percolation threshold is quickly reached at the right boundary; under these conditions, CO₂ surface areas increase, and CO₂ saturations decrease. The CO₂ surface areas for line sources are generally slightly larger than those observed from point sources.

The influence of viscous forces on CO₂ invasion with increasing levels of stratification and increasing buoyancy forces from a point source is shown in Figures 8. Here, results for $VSN = 250$ and a vertical correlation length of 1 are presented. Viscous forces lead to pooling of CO₂ around the source, which decreases CO₂ surface area and increases CO₂ saturation when compared with the results for simulations where viscous forces are absent (e.g., Figures 2, 3, and

6). The effect of CO₂ pooling is most apparent in strongly stratified systems, where the CO₂ surface area is greatly reduced (~ 0.7) when compared with simulations without viscous forces. For example, in Figure 2, where CO₂ invasion is only controlled by spanning pressures, CO₂ surface areas were around 1.2 in strongly stratified systems (correlation ratio of 1000). As buoyancy forces were increased, the CO₂ surface area increases to ~ 1.2 and ~ 1.5 at Bond Numbers of 0.1 (Figure 4) and 1 (Figure 6), respectively.

The overall trends observed in the surface area and saturation plots in Figure 8 are similar to those reported above. Here, the average CO₂ surface area in unstructured media increases as buoyancy forces are increased, while the average CO₂ saturation decreases. Increasing stratification initially leads to vertical stacking of CO₂ saturated layers, reducing the surface area. As stratification is increased further, the CO₂ surface area increases slightly, but does not reach the high values observed in the absence of viscous forces. When buoyancy forces are absent, CO₂ saturation is very large (~ 0.2), and then decreases as stratification increases. In the presence of buoyancy forces, saturations peak (~ 0.08) in weakly stratified media as CO₂ is vertically stacked in overlapping layers. As stratification increases, CO₂ invasion becomes more strongly focused into horizontal layers, reducing saturations (~ 0.02).

The coefficient of variation (C_v) of the CO₂ surface area (Figure 9) reflects the relative uncertainty in the average CO₂ surface area. The pattern of the surface area C_v versus correlation ratio or stratification strength is varies mainly with the vertical correlation length. When the vertical correlation length is 1 (Figure 9a), the C_v generally increases as stratification becomes stronger. Here, the unstructured systems are more closely ergodic, and the systems become less ergodic (more variability in invasion pathways) as the correlation ratio increases. When the vertical correlation length is 10 (Figure 9b), the surface area C_v initially increases and then

decreases, as stratification becomes stronger, reflecting a shift from mainly upward to dominantly horizontal invasion patterns. When the vertical correlation length is 50 (Figure 9c), the surface area C_v generally decreases with increasing stratification, as CO₂ invasion proceeds along increasingly narrow, horizontal pathways. The surface area C_v during CO₂ invasion with viscous forces into media with a vertical correlation length of 1 is similar to that observed in Figure 9a, indicating that the addition of viscous forces does not lead to significantly different CO₂ invasion patterns between simulations.

In general, the uncertainty increases when B_o increases from 0 to 0.1, as CO₂ invasion pathways are pulled upward by buoyancy forces. A further increase in B_o from 0.1 to 1, generally leads to a decrease in the uncertainty, as invasion is focused into narrower pathways.

Our results show that the interplay between capillary heterogeneity, buoyancy forces, and viscous force controls the dissolution rate and aqueous concentrations of CO₂ derived from well-bore leaks in aquifer systems. The dissolution rate of gaseous CO₂ into groundwater is proportional to the surface area of the CO₂ phase. Our study shows that variations in the style of capillary heterogeneity and the strength of buoyancy and viscous forces can lead to differences in CO₂ dissolution rates of well over an order of magnitude (factor of ~17 in this study).

Depending on the scenario, the CO₂ surface area increases by a factor between 3 and 9, as the degree of stratification increases. As buoyancy forces increase (from $B_o = 0$ to $B_o = 1$), the CO₂ surface area typically increases by a factor of ~2. The addition of weak viscous forces lowers the CO₂ surface area by a factor of 1.1 in unstructured media and a factor of 2 in well stratified media. Changing from a point to a line source can increase CO₂ surface areas by a factor of 1.3.

The CO₂ saturation is a metric for the dissolved CO₂ concentration in groundwater, assuming all gaseous CO₂ is dissolved into groundwater, and the geochemical impact of the CO₂

well-bore leak, as high CO₂ saturations will eventually lead to high dissolved CO₂ concentrations. Increasing stratification decreases CO₂ saturation by a factor of between 8 and 60, depending upon the scenario. The addition of strong buoyancy forces ($B_o = 1$) decreases CO₂ saturation by a factor of 23 to 35 in unstructured media. Strong buoyancy forces reduce saturations by a factor of 1.2 to 3.6 in weakly-stratified media, and increase saturations by a factor of 1.1 to 2.5 in well-stratified media. Saturations increase by factors between 1.2 and 6, when weak viscous forces are added. The presence of a line source, rather than a point source, can increase CO₂ saturations by up to a factor of 4.

IV. SUMMARY AND DISCUSSION

We used a quasi-3D, MIP model to examine the effect of capillary heterogeneity, buoyancy forces, and first-order viscous forces on the surface area (per unit volume) and saturation of a CO₂ plume in a shallow aquifer, resulting from a slow well-bore leak. Compressibility and dissolution of CO₂ phase, trapping, and other dynamic effects, such as fragmentation and pulsation, were neglected. The style of capillary heterogeneity, the strength of buoyancy and viscous forces, and the size of the CO₂ source were systematically varied yielding 168 different scenarios. Each scenario was then simulated 50 times and the CO₂ surface area and saturation was determined. The style of capillary heterogeneity was systematically varied from unstructured to stratified to mimic the range of capillary heterogeneity found in nature.

We find that the CO₂ surface area and saturation is governed by the interplay between capillary heterogeneity, buoyancy forces, and viscous forces. In stratified systems, the CO₂ surface area increases and the CO₂ saturation decreases with increasing stratification and buoyancy forces, both of which focuses CO₂ invasion into narrow pathways. Viscous forces, which lead to pooling of the CO₂ near the source, lead to lower CO₂ surface area and higher CO₂ saturation. As the leak source size increases, CO₂ invades along more pathways near the source, increasing the CO₂ surface area and saturation. Variations in the style of capillary heterogeneity and the strength of buoyancy and viscous forces can lead to differences in CO₂ dissolution rates of well over an order of magnitude, and CO₂ saturations can vary by nearly two orders of magnitude.

Our results provide critical insight into the processes affecting CO₂ dissolution rates and aqueous CO₂ concentrations resulting from well-bore leaks into groundwater aquifers and can inform continuum-based modeling of well-bore leak processes and geochemical reactions within the aquifer. Knowledge of the style of capillary heterogeneity and the strength of buoyancy and viscous forces is essential for determining apparent diffusion coefficients and dissolution rates for CO₂ gas into groundwater. The capillary pressure-saturation relationships used in most continuum-based models should be modified to include the impact of stratification, the directionality of the invasion process, and the strength of local buoyancy and viscous forces. Prior to modeling CO₂ dissolution processes in aquifer materials, site-specific geologic data (core descriptions, samples, etc.) are required to identify and characterize the extent of capillary heterogeneity.

Slow leaks from boreholes may be difficult to detect in well-stratified geologic materials, as high CO₂ dissolution rates and low CO₂ saturations may make geochemical changes in the aquifer difficult to identify. In unstructured (e.g., sedimentary media affected by bioturbation or pedogenesis) and weakly stratified materials, CO₂ well-bore leaks may be more readily identifiable from geochemical changes in groundwater, as CO₂ saturations will be higher.

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List of Appendices











Appendix: A

Table 1.

Simulation variables and values

Viscous Scaling Number (VSN)	250
Leak Source Sizes	1, 102
Vertical Correlation Lengths	1, 10, 50
Correlation Ratios	1, 10, 20, 30, 50, 100, 500, 1000
Bond Number	0, 0.1, 1

Table 2. Summary of simulation results as changes to the CO₂ surface area and saturation. Values are factors of increase or decrease.

Increase	<u>Surface Area</u>	<u>Saturation</u>
Stratification	 1.5	 8-60
Vertical Correlation Length	 3 – 4	 1.3 – 1.6
Source Size	 1.2 – 1.7	 1.2 – 4
Buoyancy Forces	 1.3 – 3.7	 6.5 – 35
Viscous Forces	 1.1 – 1.7	 1.7 – 6.2

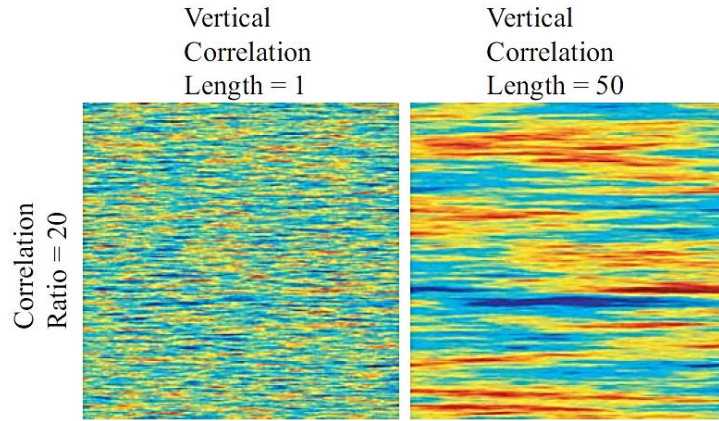


Figure 1. Random fields with (a) a horizontal correlation length of 20 and a vertical correlation length of 1, and (b) a horizontal correlation length of 100 and a vertical correlation length of 50. The horizontal correlation length will vary to maintain correlation ratios.

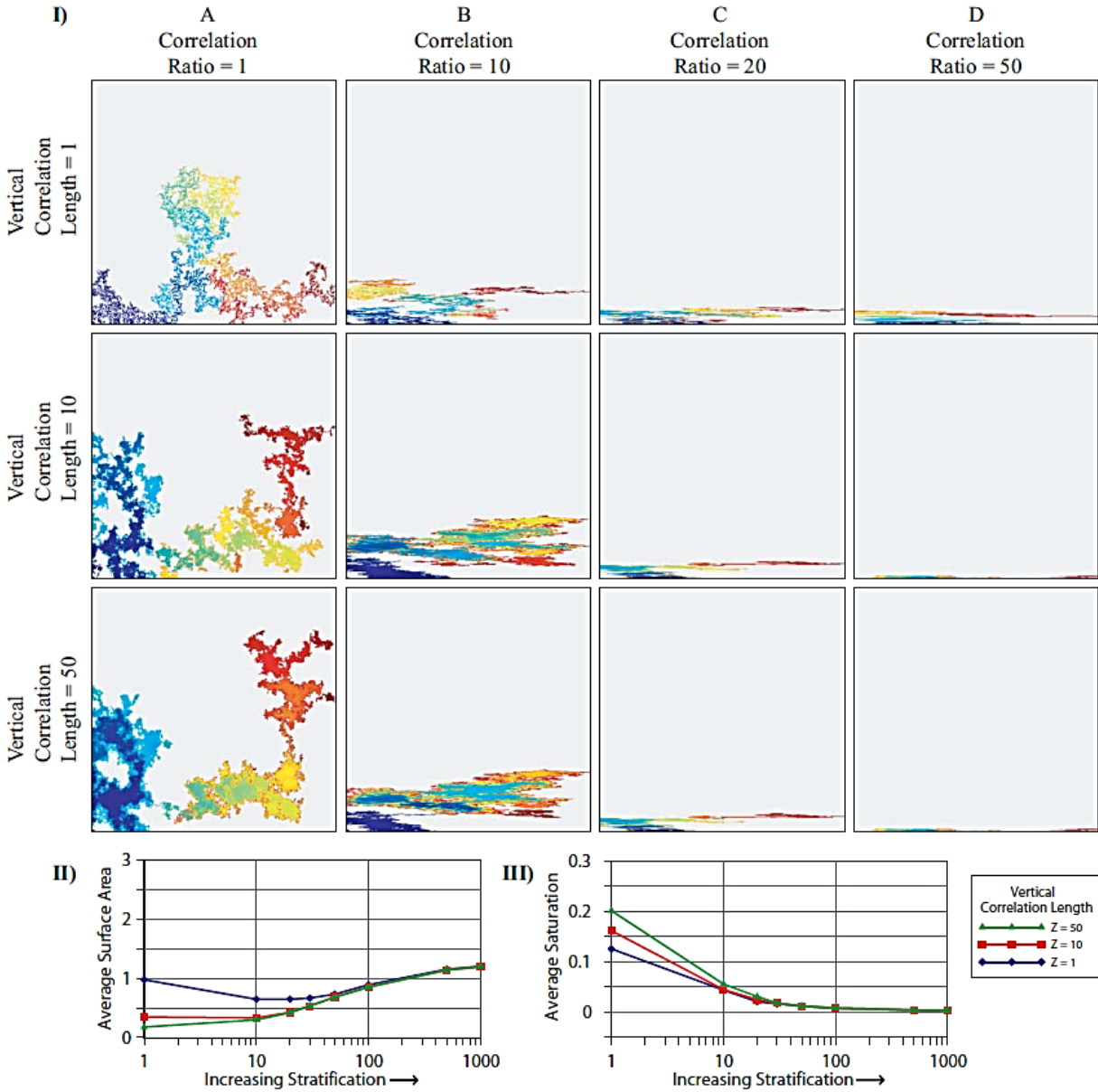


Figure 2. CO₂ invasion history (I) from a point source in the absence of buoyancy forces and viscous forces in (a) unstructured media, (b) weakly stratified media, (c) more stratified media, and (d) highly stratified media. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

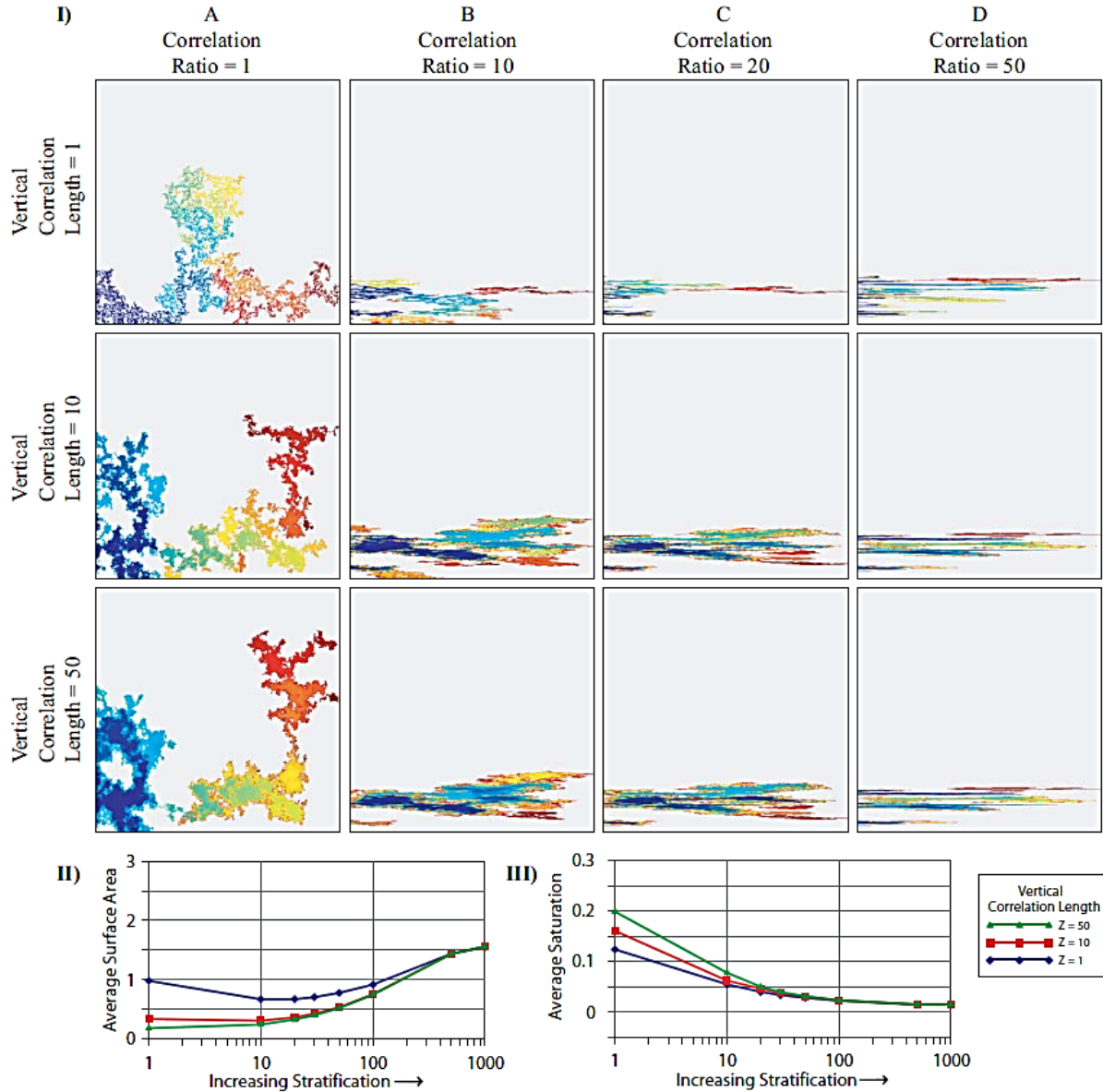


Figure 3. CO₂ invasion history (I) from a line source with a Bond Number = 0 in (a) unstructured media, (b) weakly stratified media, (c) more stratified media, and (d) highly stratified media. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

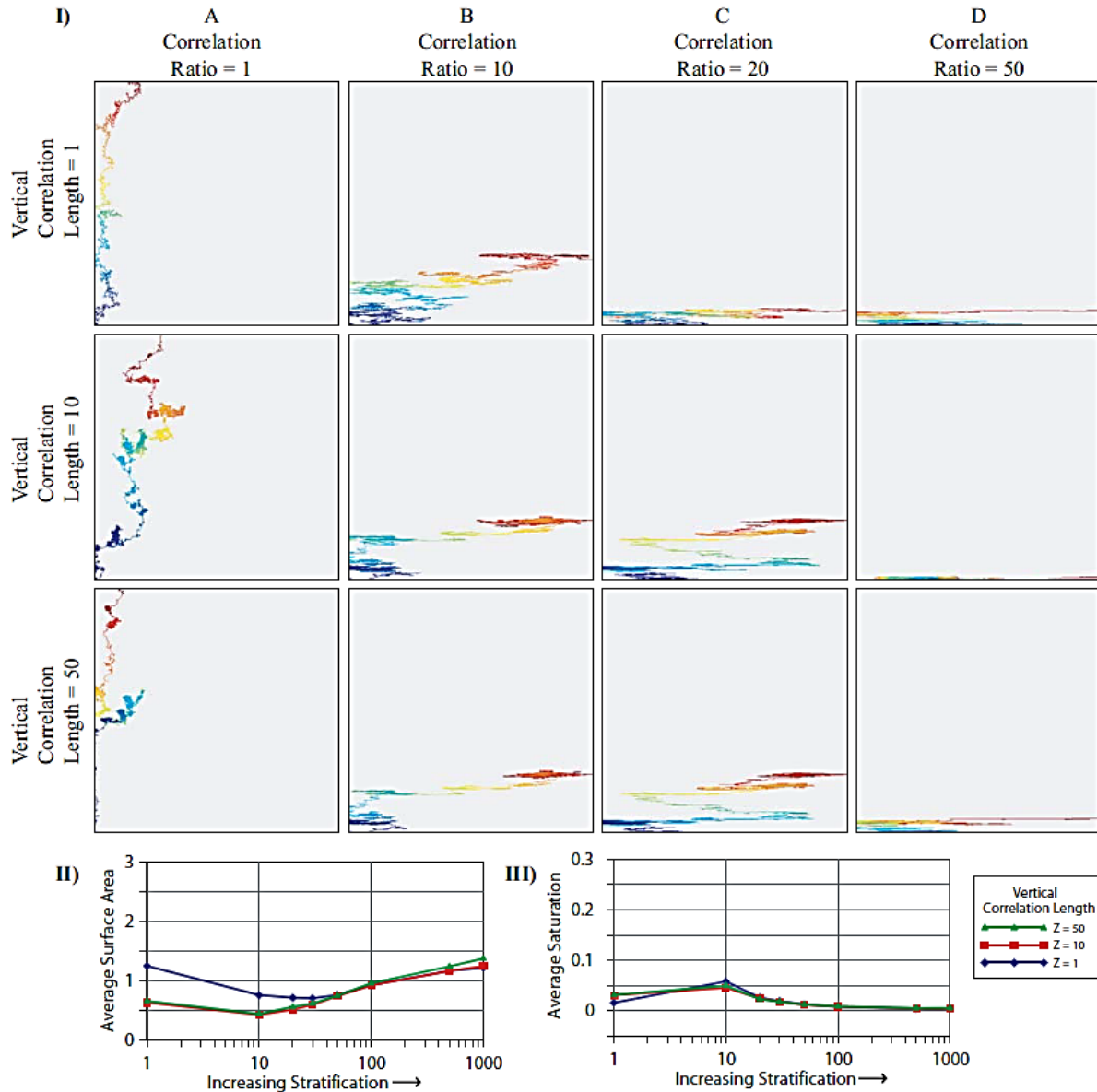


Figure 4. CO₂ invasion history (I) from a point source with a Bond Number = 0.1 and no viscous forces in (a) unstructured media, (b) weakly stratified media, (c) more stratified media, and (d) highly stratified media. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

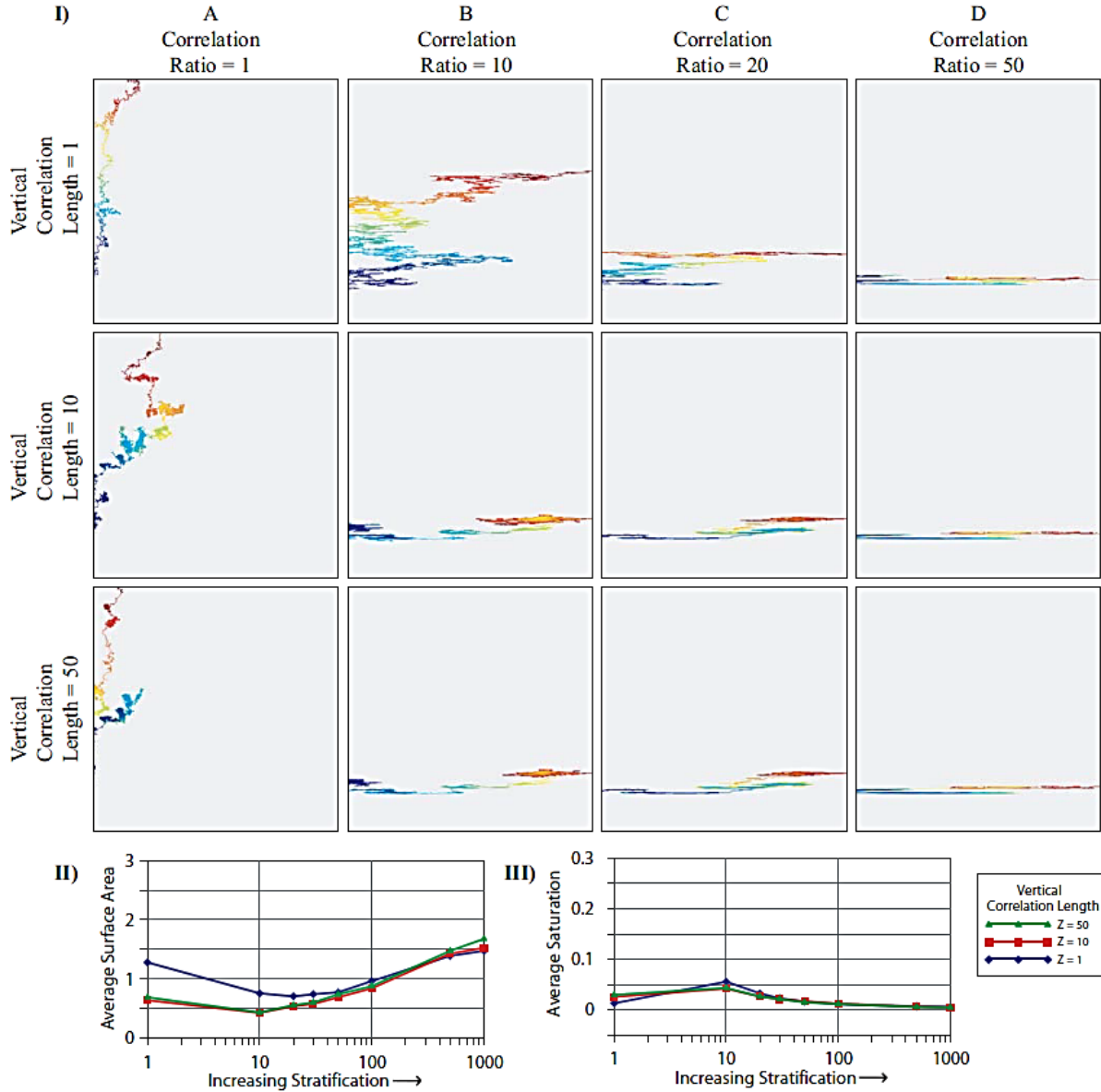


Figure 5. CO₂ invasion history (I) from a line source with a Bond Number = 0.1 and no viscous forces in (a) unstructured media, (b) weakly stratified media, (c) more stratified media, and (d) highly stratified media. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

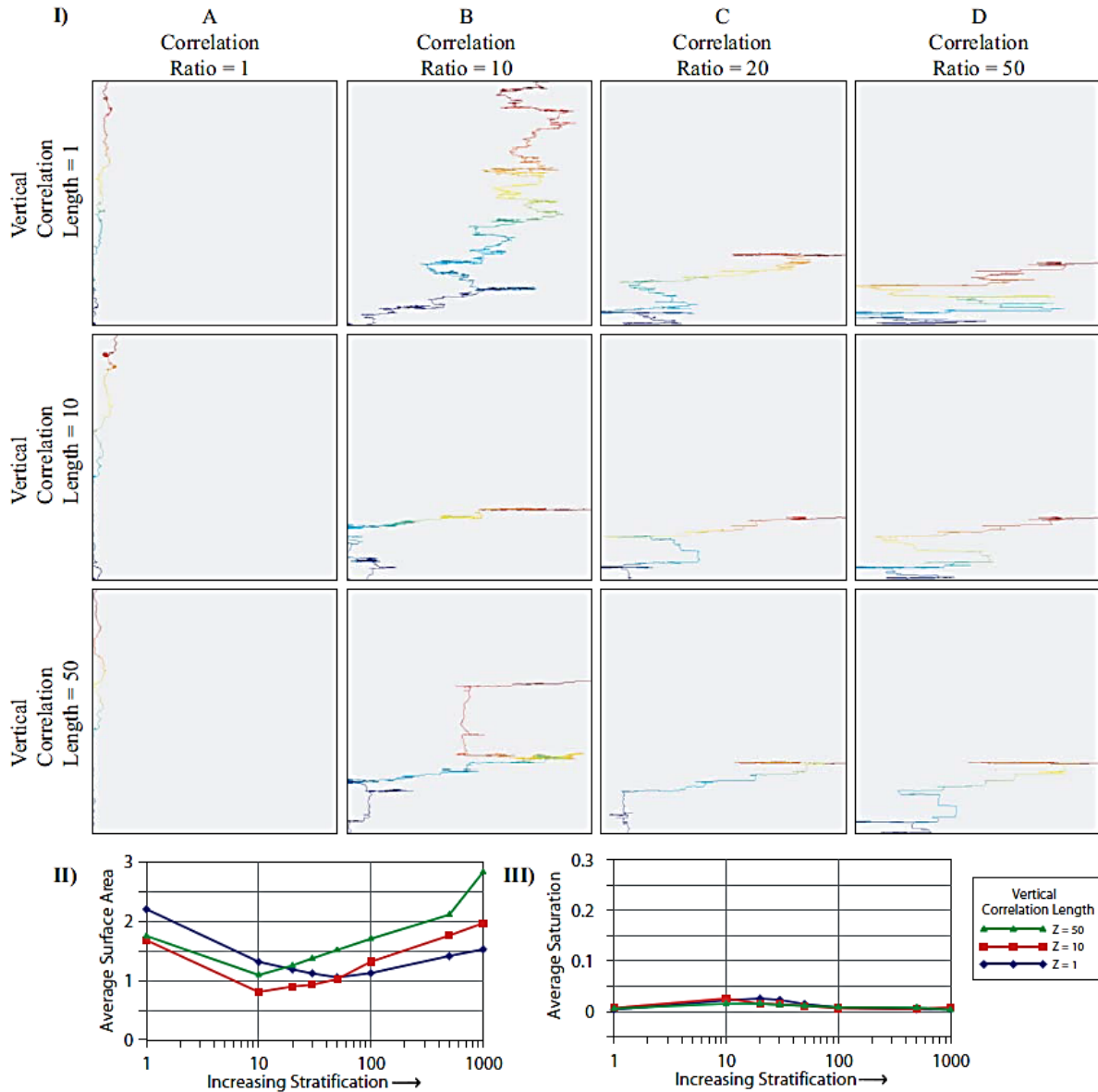


Figure 6. CO₂ invasion history (I) from a point source with a Bond Number = 1 and no viscous forces in (a) unstructured media, (b) weakly stratified media, (c) more stratified media, and (d) highly stratified media. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

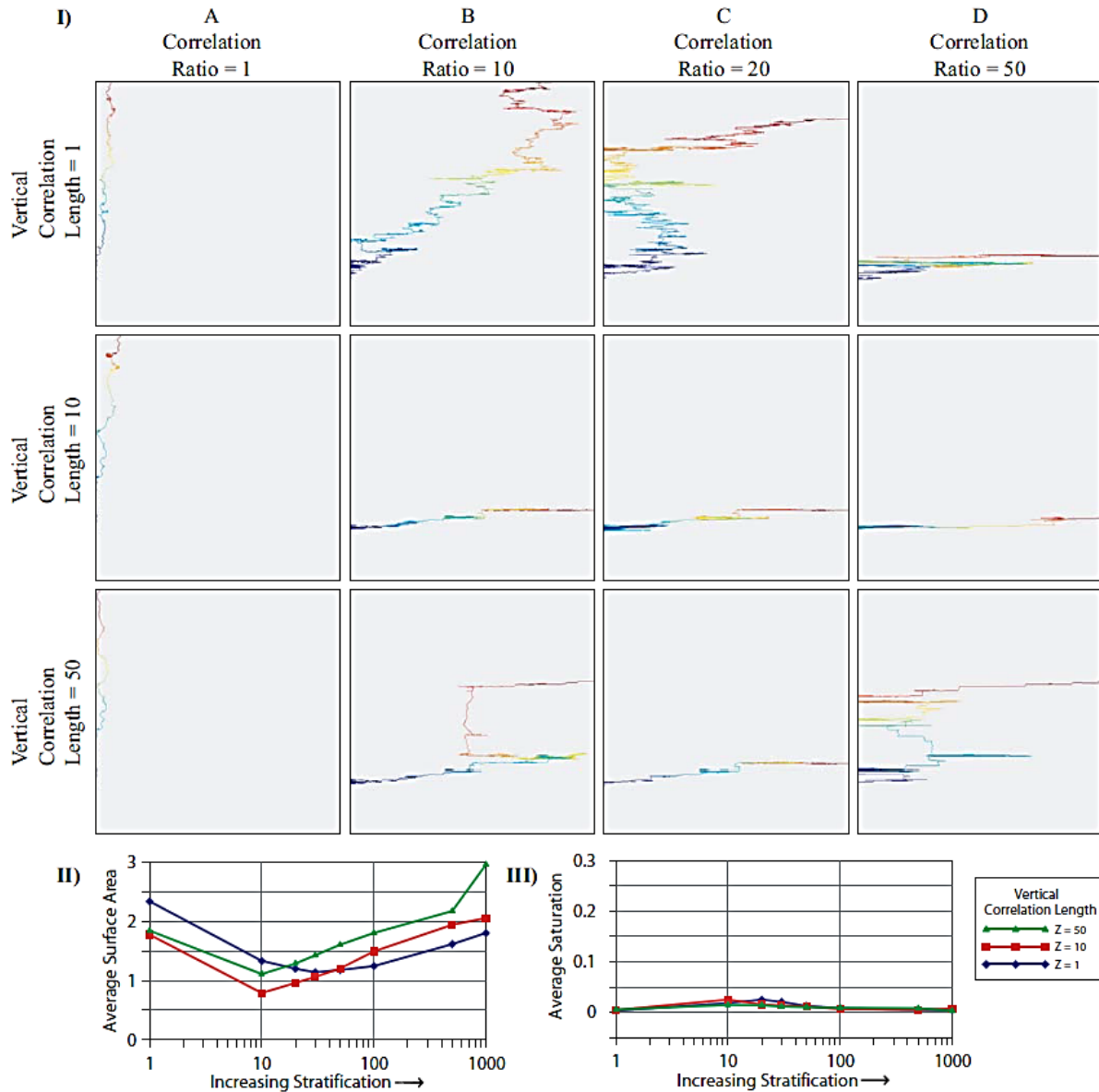


Figure 7. CO₂ invasion history (I) from a line source with a Bond Number = 1 and no viscous forces in (a) unstructured media, (b) weakly stratified media, (c) more stratified media, and (d) highly stratified media. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

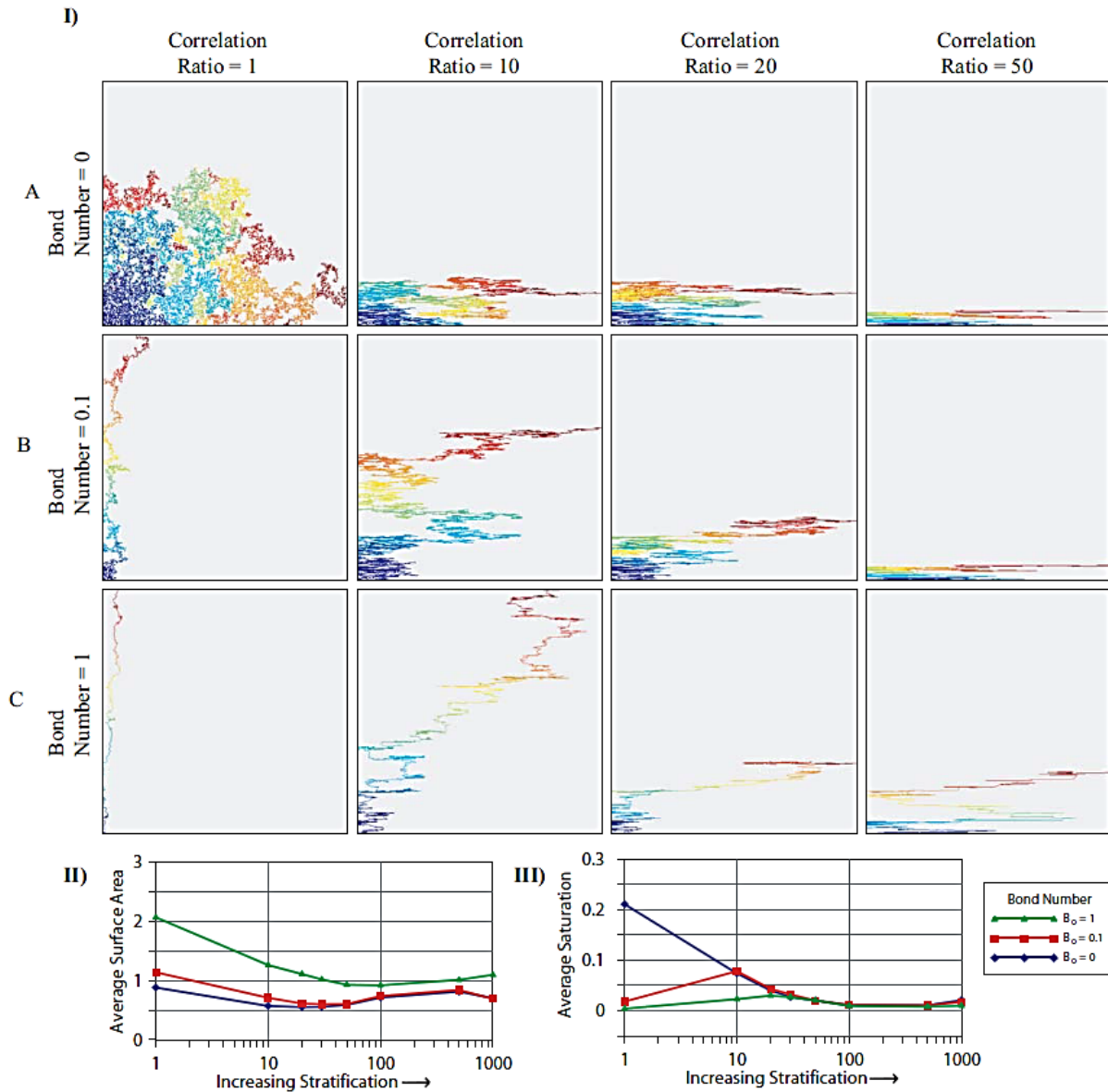


Figure 8. CO₂ invasion history (I) from a point source with VSN = 250 with Bon Numbers of (A) 0, (B) 0.1, (C) 1. Invasion pictures for correlation ratios of 100, 500, and 1000 are not shown because the invasion phase only populates a few layers near the bottom of the domain. CO₂ invasion progress from early to late is reflected by the progression of dark to light colors. Grid blocks saturated with the wetted phase are shown in gray. Plots displaying the (II) average surface area, and (III) average saturation for these simulations versus increasing stratification. The average saturation is plotted between -0.5 to 0.3 to clearly illustrate results close to zero. No simulations produced saturations less than zero.

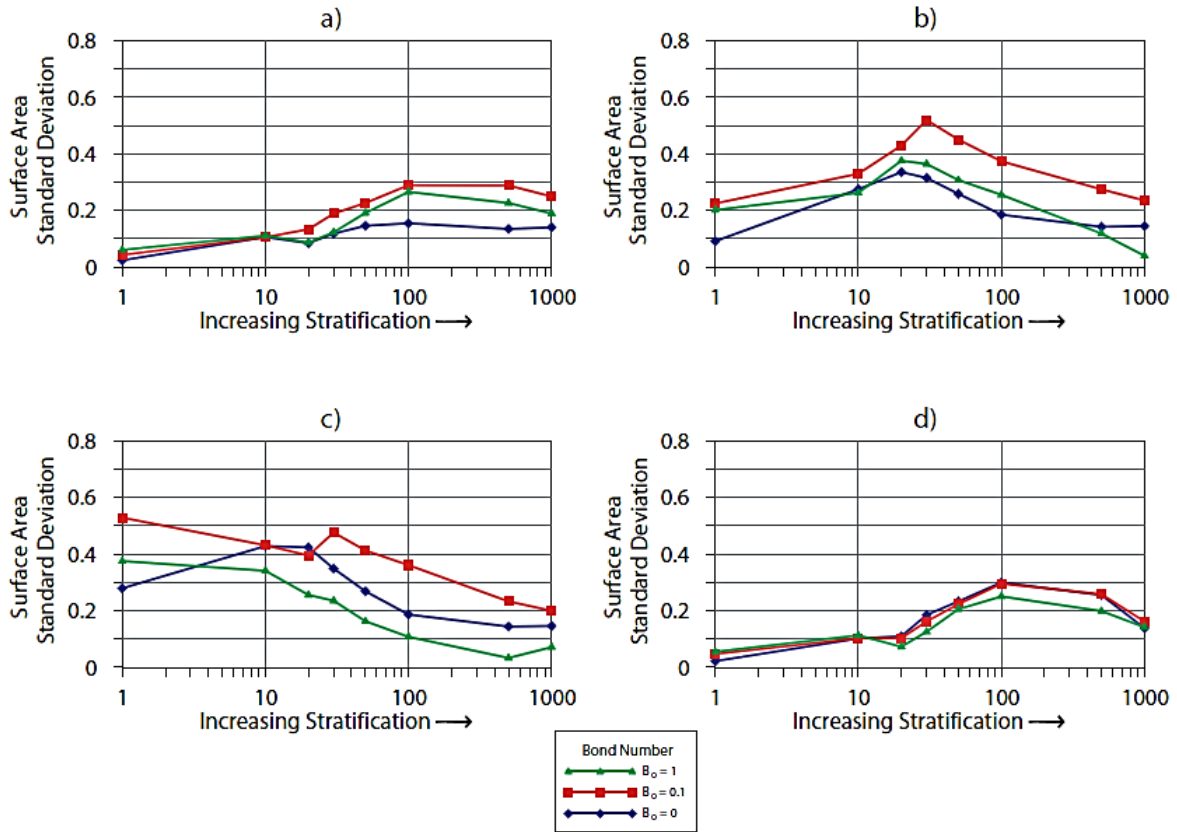


Figure 9. Plots displaying the coefficient of variation (C_v) of the CO_2 surface area for correlation lengths of (a) 1, (b) 10, and (c) 50 from a line source. Plot (d) displaying C_v of the CO_2 surface area for a correlation of 1 with a $\text{VSN} = 250$ invading from a point source.

Appendix: B

The appendix for this paper is not attached. It includes a copy of the program code, user manuals, and all program files.

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