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Quantification of prediction uncertainty using imperfect subsurface models with model error estimation

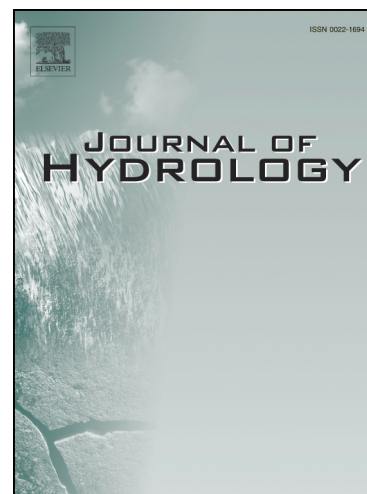
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1 Quantification of prediction uncertainty using imperfect subsurface
2 models with model error estimation

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7 **Abstract**

8 Subsurface reservoirs are far more heterogeneous and complex than the simulation models
9 in terms of scale, assumptions and description. In this work, we address the issue of prediction
10 reliability while calibrating imperfect/low-fidelity reservoir models. The main goal is to avoid
11 over-confident and inaccurate predictions by including a model for the bias terms (i.e. error-
12 model of a predefined form) during the history matching process. Our aim is to obtain unbiased
13 posterior distributions of the physical model parameters and thus improving the prediction
14 capacity of the calibrated low-fidelity reservoir models. We formulate the parameter estimation
15 problem as a joint estimation of the imperfect model parameters and the error-model parameters.
16 The structure of the error-model and the prior distributions of the error-model parameters are
17 evaluated before calibration through analysis of leading sources of the modeling errors. We
18 adopt a Bayesian framework for solving the inverse problem, where we utilize the ensemble
19 smoother with multiple data assimilation (ES-MDA) as a practical history matching algorithm.

20 We provide two test cases, where the impact of typical model errors originating from grid
21 coarsening/up-scaling and from utilizing an imperfect geological model description is investi-
22 gated. For both cases results from the ES-MDA update with and without accounting for model
23 error are compared in terms of estimated physical model parameters, quality of match to histori-
24 cal data and forecasting ability compared to held out data. The test results show that calibration

25 of the imperfect physical model without accounting for model errors results in extreme values
26 of the calibrated model parameters and a biased posterior distribution. With accounting for
27 modeling errors the posterior distribution of the model parameters is less biased (i.e. nearly un-
28 biased) and improved forecasting skills with higher prediction accuracy/reliability is observed.
29 Moreover, the consistency between the different runs of the ES-MDA is improved by including
30 the modeling error component. Although the examples in the paper consider the oil-water sys-
31 tem with permeabilities being parameters of the physical model, the developed methodology is
32 general and can be applied to typical ground water hydrology models.

33 1 Introduction

34 In subsurface reservoir modeling, various approximations are introduced at different stages of the
35 modeling process which in turn render most of the models to be imperfect and low-fidelity in nature.
36 However, these imperfect models are generally still useful for understanding the key physical inter-
37 actions within the subsurface regions of interest. The sources of approximations (a.k.a. modeling
38 errors) include: properties up-scaling (grid coarsening), discretization errors, imperfect reservoir
39 fluid properties, relative permeability, reservoir geology description/parameterization and approx-
40 imate representation of the complete complex subsurface fluid flow physics (e.g. black-oil model in
41 place of a compositional model or constant rock compressibility assumption).

42 In the context of error modeling, grid up-scaling has been widely studied within the reservoir
43 simulation community. In the published literature, there exist a number of efficient up-scaling
44 techniques (Durlofsky, 2003), aiming to obtain optimal upscaled properties. However, up-scaling
45 errors are not completely eliminated by most of these methods. Discretization errors also cannot
46 be eliminated (Ertekin et al., 2001), even after selecting an optimal grid size and utilizing adaptive
47 time stepping techniques. Additionally, numerical simulation using an optimal fine-grid could be
48 computationally prohibitive especially for tasks that typically requires many simulation runs, for
49 example history matching or robust optimization problems. Various techniques have been proposed
50 to address this computational bottleneck, for example reduced order modeling and proxy models
51 among many other techniques [c.f., Silva et al., 2007; Rammay and Abdulraheem, 2014; Cardoso

52 et al., 2009]. An alternative approach is to utilize an upscaled model instead of the fine scale model
53 for multi-query computationally demanding tasks (e.g. uncertainty quantification problems). In the
54 context of history matching, if the up-scaling errors are not modeled during the parameter inference
55 step, the posterior distributions of the model parameters is likely to be biased and this bias will
56 subsequently affect the future predictions of the engineering quantities of interest (e.g. oil, gas and
57 water rates/pressure). Omre et al. (2004) approximated the up-scaling and discretization errors
58 by computing samples or realizations of the error data using pairs of fine- and coarse-scale models.
59 The model errors due to up-scaling were then estimated using a multiple regression technique and
60 added to the coarse scale model predictions during the history matching process. Lodoen et al.
61 (2005) utilized a similar procedure on a different set of test cases while employing a more accurate
62 up-scaling procedure.

63 Accurate reservoir geology description is another challenging task due to various uncertainties
64 including: channel geometry, faults shape, facies proportion, stratigraphic and/or structural frame-
65 works. It is widely known that unrealistic geological models could be calibrated to match the histor-
66 ical data [c.f., Carter et al., 2006; Refsgaard et al., 2012]. However, these fitted models fail to provide
67 reliable predictions and could ultimately mislead the reservoir development plans [c.f., Carter et al.,
68 2006; Refsgaard et al., 2012]. Although considerable effort is often put into constructing geological
69 models that are as realistic as possible, it is very difficult to maintain this geological realism while
70 updating them to match the observed data (Sun and Durlofsky, 2017).

71 Accounting for model errors during the calibration process has attracted a large body of re-
72 search [c.f., Oliver and Alfonzo, 2018; Dreano et al., 2017; Josset et al., 2015], where various ap-
73 proaches have been developed to account for the model-error component during model calibra-
74 tion [c.f., Hansen et al., 2014; Evin et al., 2014; Reichert and Schuwirth, 2012]. These approaches
75 vary according to the different behavior and complexity of the modeled physical system. For ex-
76 ample in hydro-geophysical systems, Köpke et al. (2017) accounted for the model-error component
77 using orthonormal basis generated from an error dictionary which is continuously enriched during
78 the calibration process. The models of the bias or error component could be generally classified
79 as either input dependent (Giudice et al., 2013) or output dependent (Evin et al., 2014). Input

80 dependent model error formulation represents the error components as a function of the model
81 parameters. For example, a reservoir model-error can vary with permeability realizations or other
82 input model parameters. O’Sullivan and Christie (2005) utilized an input dependent formulation
83 for the model-error where the authors computed model-error realizations using the difference be-
84 tween a fine-grid and coarse-grid model outputs. During the calibration process, an interpolation
85 of the error component was performed to estimate a correction term to the coarse-grid model pre-
86 dictions. Lødøen and Tjelmeland (2010) used multiple linear regression algorithm to model errors,
87 where the residual part for the multiple regression was assumed to depend on the model input
88 parameters. The residual terms were modeled using a zero mean Gaussian Process. Giudice et al.
89 (2013) used an input dependent model-error representation to improve uncertainty estimation in
90 urban hydrological models. In that application, the model error variance was set to be depending
91 on the input of the rainfall term. Output dependent model error formulation represents the error
92 components as a function of the output of the physical model. For example, Evin et al. (2014)
93 utilized an output dependent formulation for the model-error heteroscedasticity as a function of
94 the simulated streamflow. In a realistic setting, where large models are utilized (i.e. millions of
95 input parameters), it is hard to relate the model errors to the high dimensional input parameter
96 space and output dependent or input/output (I/O) independent forms of the model-error were
97 proposed as an alternative approach that might have some advantages over the input-dependent
98 error-models (Giudice et al., 2013).

99 Model-error representation can also be classified as either an external bias description (EBD)
100 or an internal noise description (IND) (Giudice et al., 2015). EBD was developed using the back-
101 ground of statistical inference in a regression type framework (Giudice et al., 2015). In EBD, the
102 model-error term is added externally into the forward model (approximate or inadequate model)
103 output. In IND, the model-error is formulated as an additional term of the state space (Giudice
104 et al., 2015). This approach is also known as state space modelling or stochastic gray-box mod-
105 elling [c.f., Moradkhani et al., 2012; Kristensen et al., 2004]. Giudice et al. (2015) concluded that
106 EBD has some advantages over IND in terms of long-term predictions.

107 In this paper, we utilize an EBD and I/O independent error-model formulation as it is more

108 suitable for large scale models (e.g. subsurface oil and gas reservoir models) relying on black-box
109 simulators. Several typical sources of model errors are present in the test cases investigated in this
110 paper, including a coarse grid, less detailed geological representation (i.e. upscaling of different types
111 of geological features including variogram based and channelized geology), discretization errors and
112 slight change in well locations due to grid coarsening. The error-model formulation presented here,
113 assumes that the total modeling errors consists of two components: structural component and noise-
114 like component. We note that the structural component is often neglected in the bias correction
115 approaches developed in hydrological literature [c.f., Maier et al., 2014; Vrugt, 2016; White et al.,
116 2014]. In this study prior to model calibration, the structure of the model error is estimated
117 and represented using several basis functions, and the magnitude of the noise-like component is
118 quantified. During history matching, the weights of the basis functions are jointly calibrated with
119 the physical model parameters using data observed at well locations. The noise-like part of the
120 model error is also accounted for during the history matching process to avoid over-fitting of the
121 error model. We note that the presented formulation is general and can be applied to other sources
122 of modeling errors when dealing with low-fidelity physical models. The low-fidelity models are
123 generally used as efficient surrogate models for computationally demanding tasks [c.f., Asher et al.,
124 2015; Laloy et al., 2013].

125 For the Bayesian inversion, we use a particular type of iterative ensemble smoother ES-MDA (Em-
126 erick and Reynolds, 2013). The formulation of ES-MDA has some similarities with Kalman filtering
127 algorithms (Sun et al., 2016). However, ES-MDA assimilates data from different times simultane-
128 ously and the same set of data is assimilated multiple times with an inflated data noise covariance
129 matrix which is equivalent to annealing approaches (Stordal and Elsheikh, 2015). The rest of the
130 paper is organized as follows: In Section 2, we present some background on Bayesian inverse mod-
131 eling followed by the proposed error-model formulation. Following that, we present the case studies
132 in Section 3. The results of the case studies are discussed in Section 4 followed by the conclusions
133 of our work in Section 5.

134 2 Methodology

135 Bayesian inverse modeling is a generic inference framework that is widely adopted for calibration
 136 of reservoir models while accounting for different types/sources of uncertainties. In the Bayesian
 137 framework, the conditional probabilities $p(\mathbf{m}|\mathbf{d}_{obs})$ of the model parameters \mathbf{m} given the observa-
 138 tional data \mathbf{d}_{obs} (a.k.a. posterior distribution of the model parameters) is estimated using Bayes
 139 rule (Oliver et al., 2008):

$$p(\mathbf{m}|\mathbf{d}_{obs}) \propto p(\mathbf{d}_{obs}|\mathbf{m}) p(\mathbf{m}), \quad (1)$$

140 where \mathbf{m} is the model parameters vector of size N_m , \mathbf{d}_{obs} is the observations vector of size N_d , $p(\mathbf{m})$
 141 is the prior probability of the model parameters and $p(\mathbf{d}_{obs}|\mathbf{m})$ is the likelihood function of the data
 142 given a specific realization of the model parameters \mathbf{m} . It is common to assume a Gaussian prior:

$$p(\mathbf{m}) \propto \exp\left(-\frac{1}{2}(\mathbf{m} - \mathbf{m}_{pr})^\top \mathbf{C}_M^{-1}(\mathbf{m} - \mathbf{m}_{pr})\right), \quad (2)$$

143 where \mathbf{m}_{pr} is an N_m dimensional vector of the mean prior model parameters and \mathbf{C}_M is the
 144 covariance matrix of the prior model parameters. It is also common to assume that data noise is
 145 Gaussian, so that the likelihood function takes the form:

$$p(\mathbf{d}_{obs}|\mathbf{m}) \propto \exp\left(-\frac{1}{2}(\mathbf{d}_{obs} - \mathbf{d})^\top \mathbf{C}_D^{-1}(\mathbf{d}_{obs} - \mathbf{d})\right), \quad (3)$$

146 where \mathbf{d} is the simulated or predicted data vector using the model parameters \mathbf{m} and \mathbf{C}_D is the
 147 error/noise covariance matrix which is defined in Sect. 2.1 and 2.2 depending on the utilized history
 148 matching procedure. Using these definitions, Bayes' rule defined in Eq. 1 could be expanded as
 149 following:

$$p(\mathbf{m}|\mathbf{d}_{obs}) \propto \exp\left(-\frac{1}{2}\left((\mathbf{d}_{obs} - \mathbf{d})^\top \mathbf{C}_D^{-1}(\mathbf{d}_{obs} - \mathbf{d}) + (\mathbf{m} - \mathbf{m}_{pr})^\top \mathbf{C}_M^{-1}(\mathbf{m} - \mathbf{m}_{pr})\right)\right). \quad (4)$$

150 Several algorithms could be used to generate samples from the posterior distribution of the
 151 model parameters (Oliver et al., 2008). Among those Markov Chain Monte Carlo (MCMC) is

152 an exact method for sampling. However, MCMC can be computationally expensive due to the
 153 large number of iterations needed to reach convergence and the sequential nature of the method.
 154 Ensemble-based methods have been widely used for calibrating subsurface flow models due to
 155 the computational feasibility and parallel nature of ensemble methods. In this study, we utilize
 156 the ensemble smoother with multiple data assimilation (ES-MDA) algorithm for the calibration
 157 step (Emerick and Reynolds, 2013). ES-MDA belongs to a class of iterative ensemble smoothing
 158 techniques that could be used to solve non-linear inverse problem iteratively with an inflated noise
 159 covariance matrix. The ES-MDA algorithm steps are summarized as follows:

- 160 – Select the number of iterations (number of data assimilation) N_a and the inflation factor α . A
 161 common choice of the inflation factor is to set it as a constant value for all iterations $\alpha = N_a$
- 162 – Initialize an ensemble of model parameters and perturb the observation data for each ensemble
 163 member using:

$$\mathbf{d}_{uc,j} = \mathbf{d}_{obs} + \sqrt{\alpha} \mathbf{C}_D^{1/2} \mathbf{z}_d, \quad (5)$$

164 where the subscript j is the ensemble member index $j = 1 \dots N_e$ and N_e is the ensemble size,
 165 $\mathbf{d}_{uc,j}$ is N_d dimensional vector of perturbed observation, \mathbf{z}_d is the N_d dimensional vector with
 166 standard Gaussian random variables as its components (i.e. $\mathbf{z}_d \sim \mathcal{N}(0, \mathbf{I}_{N_d, N_d})$).

- 167 – Update each ensemble member using,

$$\mathbf{m}_j^{(i+1)} = \mathbf{m}_j^{(i)} + \mathbf{C}_{MD} (\mathbf{C}_{DD} + \alpha \mathbf{C}_D)^{-1} (\mathbf{d}_{uc,j} - \mathbf{d}_j), \quad (6)$$

168 where the superscript i is the iteration index, \mathbf{C}_{DD} is the model output covariance matrix
 169 and \mathbf{C}_{MD} is the cross covariance matrix of model parameters and model predictions.

- 170 – Repeat the above steps for all iterations, from $i = 1$ to N_a

171 2.1 Procedures for history matching of reservoir models

172 In this study, two types of history matching procedures are investigated: history matching while
 173 neglecting model-discrepancy (i.e. standard history matching procedure) and joint history matching

174 of the model parameters and the parameters of an error/bias model. In this paper, we use the term
 175 model-discrepancy and model-error interchangeably. The standard history matching procedure
 176 relies on an implicit assumption that the model-errors are generally small and could be neglected
 177 (i.e. the simulation model is perfect). Mathematically, if an accurate/high-fidelity model is utilized,
 178 the observed data is formulated as (Giudice et al., 2013):

$$\mathbf{d}_{obs} = g(\mathbf{m}_{true}) + \boldsymbol{\epsilon}_d, \quad (7)$$

179 where $g(\cdot)$ is a nonlinear function representing the accurate/high-fidelity forward simulation model,
 180 \mathbf{m}_{true} is the true model parameters, $\boldsymbol{\epsilon}_d$ is the measurement errors which is usually assumed to follow
 181 a normal distribution $\mathcal{N}(0, \mathbf{C}_d)$ and \mathbf{C}_d is the measurement errors covariance matrix. In this study,
 182 uncorrelated measurement errors are considered, therefore the matrix \mathbf{C}_d is a diagonal matrix. In
 183 standard history matching (i.e. neglecting modeling errors), \mathbf{C}_D in Eqs. 3, 4, 5 and 6 is set to the
 184 covariance of measurement errors. Therefore,

$$\mathbf{C}_D = \mathbf{C}_d. \quad (8)$$

185 However, as noted in the introduction section, several approximations are commonly introduced
 186 in the computational model to simplify the simulation process (e.g. black-oil model versus compo-
 187 sitional flow), or to speed-up the simulations (coarsening of the simulation grid). During history
 188 matching if the model-error caused by these approximations is not accounted for, the obtained
 189 posterior distribution could be biased. In the case of utilizing an approximate/low-fidelity model,
 190 the observation data is related to the true model parameters \mathbf{m}_{true} as (Giudice et al., 2013):

$$\mathbf{d}_{obs} = \tilde{g}(\mathbf{m}_{true}) + \boldsymbol{\epsilon}_d + \boldsymbol{\epsilon}_m, \quad (9)$$

191 where $\boldsymbol{\epsilon}_m$ is the model-error and $\tilde{g}(\cdot)$ is a nonlinear function representing the imperfect (approximate/low-fidelity) simulation model. By subtracting Eq. 7 from 9, we obtain:

$$\boldsymbol{\epsilon}_m = g(\mathbf{m}_{true}) - \tilde{g}(\mathbf{m}_{true}). \quad (10)$$

193 In the following sub-section, we present a simple yet general parameterization of the model-error
 194 term ϵ_m .

195 2.2 Error-model formulation

196 In this study, EBD and I/O independent error-model approach is considered. As the model errors
 197 in our test cases were dominated by structured components, these errors are parameterized using
 198 smooth basis functions obtained by principle component analysis (PCA) method, which is an
 199 effective data-driven dimension reduction technique (Shlens, 2014; Kerschen et al., 2005). We rely
 200 on simulation output from pairs of models, accurate/high-fidelity versus approximate/low-fidelity,
 201 to obtain the basis functions and the prior statistics of the coefficients of the PCA basis functions.
 202 We acknowledge that this limits the applicability of the developed approach to the cases for which an
 203 accurate/high-fidelity model is available. However, we note that the accurate/high-fidelity model is
 204 only used to estimate the prior model-error statistics, and is not used during the calibration process.
 205 This is a notable difference between the presented framework and related studies by Josset et al.
 206 (2015) and Köpke et al. (2017).

207 In the current setting, the prior model-error realizations are estimated using:

$$\epsilon_{mr} = g(\mathbf{m}_r) - \tilde{g}(\mathbf{m}_r), \quad (11)$$

208 where r is the index of the prior realizations, i.e. $r = 1$ to N_r , N_r is the total number of realizations
 209 used to estimate the model-error statistics, ϵ_{mr} is an N_d dimensional vector of model-error for
 210 realization r . All prior model-error realizations are assembled into the matrix $\epsilon \in \mathbb{R}^{N_d \times N_r}$. The
 211 mean of the model-error prior is,

$$\bar{\epsilon}_m = \frac{1}{N_r} \sum_{r=1}^{N_r} (\epsilon_{mr}). \quad (12)$$

212 The covariance of the model-error prior is (Oliver et al., 2008),

$$\mathbf{C}_e = \frac{1}{N_r - 1} (\epsilon - \bar{\epsilon}_m \mathbf{I}_{N_r})(\epsilon - \bar{\epsilon}_m \mathbf{I}_{N_r})^\top, \quad (13)$$

213 where \mathbf{I}_{N_r} is an N_r dimensional row vector with all ones as its components. In this study, PCA is
 214 used to parametrize the prior model-error realizations and the weights of the obtained PCA basis
 215 vectors are jointly inferred with the model parameters during the history matching process. The
 216 basis functions are obtained by singular value decomposition (SVD) of the error covariance matrix
 217 \mathbf{C}_e (Oliver et al., 2008):

$$\mathbf{C}_e = \mathbf{U}\mathbf{\Sigma}\mathbf{V}^T, \quad (14)$$

218 where \mathbf{U} and \mathbf{V} are the orthonormal singular vectors (basis functions) and $\mathbf{\Sigma}$ is a diagonal matrix of
 219 the singular values. The error-model is formulated using the leading L singular vectors as following:

$$\hat{\epsilon}_{mr} = \mathbf{\Phi}\boldsymbol{\beta}_r + \bar{\epsilon}_m, \quad (15)$$

220 where $\mathbf{\Phi} \in \mathbb{R}^{N_d \times L}$ are the first L orthonormal singular vectors (basis functions) from \mathbf{U} and
 221 $\boldsymbol{\beta}_r \in \mathbb{R}^{L \times 1}$ are the coefficients of error-model for realization r .

222 The objective of the calibration process is then to find the posterior distribution of reservoir
 223 model parameters and the coefficients $\boldsymbol{\beta}$ of the PCA-based error-model. Since Bayesian inverse
 224 modelling require prior statistics of model parameters, therefore the prior statistics of the coefficients
 225 $\boldsymbol{\beta}$ should be estimated. Least square form of Eq. 15 is used to compute prior realizations of the
 226 coefficient vector $\boldsymbol{\beta}$ as following:

$$\boldsymbol{\beta}_r = (\mathbf{\Phi}^T \mathbf{\Phi})^{-1} \mathbf{\Phi}^T (\epsilon_{mr} - \bar{\epsilon}_m). \quad (16)$$

227 Since $\mathbf{\Phi}$ is an orthonormal matrix i.e. $\mathbf{\Phi}^T \approx \mathbf{\Phi}^{-1}$, therefore Eq. 16 can also be written as:

$$\boldsymbol{\beta}_r \approx \mathbf{\Phi}^T (\epsilon_{mr} - \bar{\epsilon}_m). \quad (17)$$

228 The prior statistic, such as the mean and covariance of $\boldsymbol{\beta}$ realizations are computed using:

$$\boldsymbol{\mu}_\beta = \frac{1}{N_r} \sum_{r=1}^{N_r} (\boldsymbol{\beta}_r), \quad (18)$$

229

$$\mathbf{C}_\beta = \frac{1}{N_r - 1} \sum_{r=1}^{N_r} (\boldsymbol{\beta}_r - \boldsymbol{\mu}_\beta)(\boldsymbol{\beta}_r - \boldsymbol{\mu}_\beta)^\top, \quad (19)$$

230 where $\boldsymbol{\mu}_\beta \in \mathbb{R}^{L \times 1}$ is the mean of $\boldsymbol{\beta}$ realizations and $\mathbf{C}_\beta \in \mathbb{R}^{L \times L}$ is the covariance of $\boldsymbol{\beta}$ realizations.

231 In this study, we only consider the diagonal terms of the matrix \mathbf{C}_β to generate prior samples of
232 error-model coefficients for the history matching purpose.

233 In order to avoid over-fitting the error-model, the number of coefficients of the PCA-based
234 error-model L should be limited to a small number. Therefore, the residual of error-model cannot
235 be neglected and need to be included in the inversion process. The residual of the least square fit
236 is defined as,

$$\boldsymbol{\zeta}_{mr} = \boldsymbol{\epsilon}_{mr} - \hat{\boldsymbol{\epsilon}}_{mr}. \quad (20)$$

237 All residual realizations are assembled into matrix $\boldsymbol{\zeta} \in \mathbb{R}^{N_d \times N_r}$. The covariance of the residual
238 from all error-model realizations is then estimated using:

$$\mathbf{C}_T = \frac{1}{N_r - 1} \boldsymbol{\zeta} \boldsymbol{\zeta}^\top, \quad (21)$$

239 where \mathbf{C}_T is denoted the error-model noise covariance. For history matching of imperfect-models,
240 the total error covariance matrix \mathbf{C}_D in Eqs. 3, 4, 5 and 6 contains both the measurement and
241 error-model noise components as following:

$$\mathbf{C}_D = \mathbf{C}_d + \mathbf{C}_T. \quad (22)$$

242 For simplicity, only the diagonal terms of the matrix \mathbf{C}_T is considered in this study. Conceptually,
243 \mathbf{C}_D is the total uncalibrated uncertainty, which includes both the measurement noise and the
244 model-error noise that is not captured by the truncated PCA-based error-model. In (Hansen et al.,
245 2014), modeling errors were considered as uncalibrated uncertainties using Gaussian distribution
246 (i.e. accounting for mean and covariance of errors). In their approach, the model error is accounted
247 for by replacing \mathbf{d}_{obs} with $\mathbf{d}_{obs} + \bar{\boldsymbol{\epsilon}}_m$ and replacing \mathbf{C}_d with $\mathbf{C}_d + \mathbf{C}_e$. However, this approach
248 would be inconsistent/inefficient for physical systems which exhibit highly complex statistics and

249 correlations of model errors that change significantly not only over the data space, but also as a
250 function of the input model parameters. (Köpke et al., 2017).

251 3 Case Studies

252 In this section, we present the details of the case studies. The dimension of the subsurface reservoir
253 is 7500 ft \times 7500 ft \times 20 ft in the x, y and z directions, respectively. Incompressible two-phase
254 porous media flow of oil and water is considered. The initial reservoir pressure is 5000 psi and
255 the reservoir has uniform porosity of 20%. The reservoir contains one injector well (I1) and three
256 production wells (P1, P2, P3) and is simulated using a 2D grid. We utilize the Matlab Reservoir
257 Simulation Tool-box (MRST) (Lie, 2016) for the forward model simulations. Corey's power law
258 model is used to represent relative permeabilities. Parameter values for the Corey's model and fluid
259 properties are listed in Appendix A. The gravitational and capillary pressure effects are neglected.
260 The production wells are operated under constant bottom hole pressure constraint of 4500 psi and
261 the injector well is operated under constant injection rate constraint with varying control values as
262 shown in Fig. 1(a). The wells open/shut schedule is shown in Fig. 1(b). Figure 1 also shows end
263 of historical period (i.e. 2 years). In the historical period, the flow rates at the production wells
264 and the bottom hole pressure of the injector well are used as the historical data for the calibration
265 process. We also note that one of the production wells (P3) is only used in prediction phase in
266 order to assess predictions from calibrated models on wells drilled in future development plans.

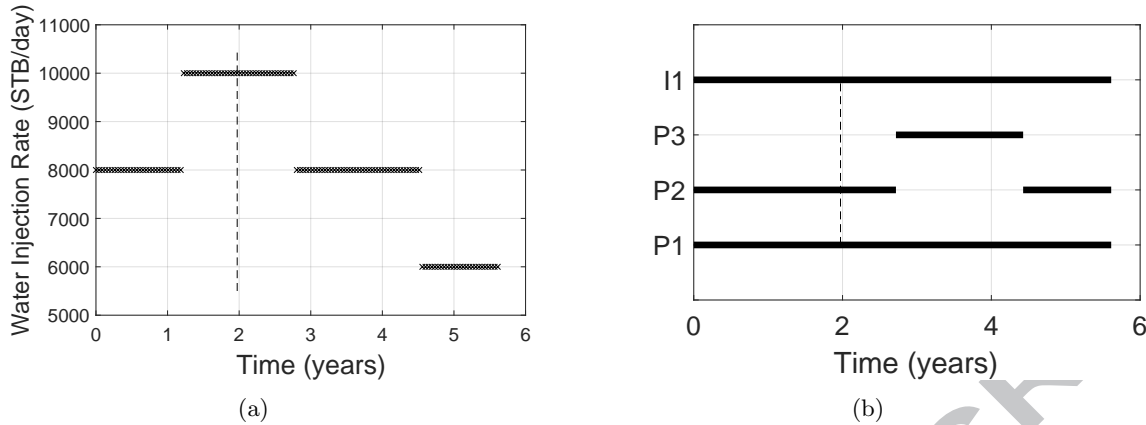


Figure 1: Injection well control rates and wells open/shut schedule. Dashed black lines show end of historical period. Part (a) shows the water injection rate of injector well. Part (b) shows wells open/shut schedule. In part (b) solid back lines indicate the time periods when a well is open to flow.

267 3.1 Case 1: Coarse scale model

268 In the first case study, fine-grid/high-fidelity model uses a 2D grid with 75×75 cells. The distributed
 269 log-permeability fields are modeled as multivariate Gaussian with exponential covariance function:

$$\mathbf{c} = \sigma^2 \exp\left(-3\left(\frac{\mathbf{s}}{r_a}\right)^\gamma\right), \quad (23)$$

270 where \mathbf{s} is the lag distance and r_a, σ^2, γ are the correlation range, variance and exponent respectively
 271 (which are 35 cells, 1 and 1 respectively in this test case). The log-permeability field $\ln(\mathbf{K})$ is
 272 parameterized using PCA and only two leading basis functions are retained:

$$\ln(\mathbf{K}) = \overline{\ln(\mathbf{K})} + \sum_{b=1}^{N_w} w_b \psi_b, \quad (24)$$

273 where $\overline{\ln(\mathbf{K})}$ is the mean log-permeability (equal to 4 in this test case), b is index of the basis weight
 274 w and basis function ψ and $N_w = 2$. Figure 2(b) shows the leading two principal basis functions
 275 obtained by singular value decomposition of covariance of log-permeability fields (Eq. 23). Figure
 276 2(a) shows the prior distribution of weights obtained by projecting the log-permeability fields into
 277 the PCA-basis functions.

278 Figure 3(a) shows the reference fine scale log-permeability field. The fine scale reference log-

279 permeability field is generated by the leading two-PCA basis functions and reference basis weights
280 are shown as the red vertical lines in Fig. 2. The coarse-grid/low-fidelity reservoir models contain
281 only 5×5 grid blocks. The coarsened version of the reference fine model is shown in Fig. 3(b),
282 in which harmonic averaging is used to up-scale the log-permeability field. The observed data are
283 generated by the fine scale model using reference log-permeability field (Fig. 3(a)) with the addition
284 of measurement noise of 2% of the reference solution. We note that except permeabilities, the rest
285 of the static and dynamic properties (i.e. porosity, relative permeability, viscosity, and density, well
286 controls and schedule) of the coarse scale model are the same as the fine scale model.

287 Two different procedures of history matching the coarse scale model were considered. In the
288 case of neglecting model-discrepancy, PCA basis weights \mathbf{w} of the log-permeabilities are calibrated
289 (i.e. $\mathbf{m}_j = \mathbf{w}$ in Eq. 6). In the case of joint inversion with error-model, the estimated parameters is
290 the combined vector of the log-permeability PCA weights and the error-model coefficients (i.e. $\mathbf{m}_j =$
291 $[\mathbf{w}; \boldsymbol{\beta}]$ in Eq. 6).

292 Prior statistics of the model-discrepancy were estimated using Eq. 11. One hundred fine scale
293 permeability realizations were generated through Eq. 24 by sampling the prior distribution of
294 the PCA-basis weights and a corresponding number of coarse scale permeability realizations were
295 obtained using harmonic up-scaling. Forward runs were then performed for both the coarse and
296 fine scale models to obtain the error realizations using Eq. 11. A smaller number of realizations
297 could be used to evaluate the model-discrepancy statistics. In that case, special care should be
298 taken to select a representative set of prior realizations to cover the respective statistic. Figure 4
299 shows the prior statistics of the model-discrepancy in the simulated well production data (bottom
300 hole pressure of the injector well and flow rates of the producers).

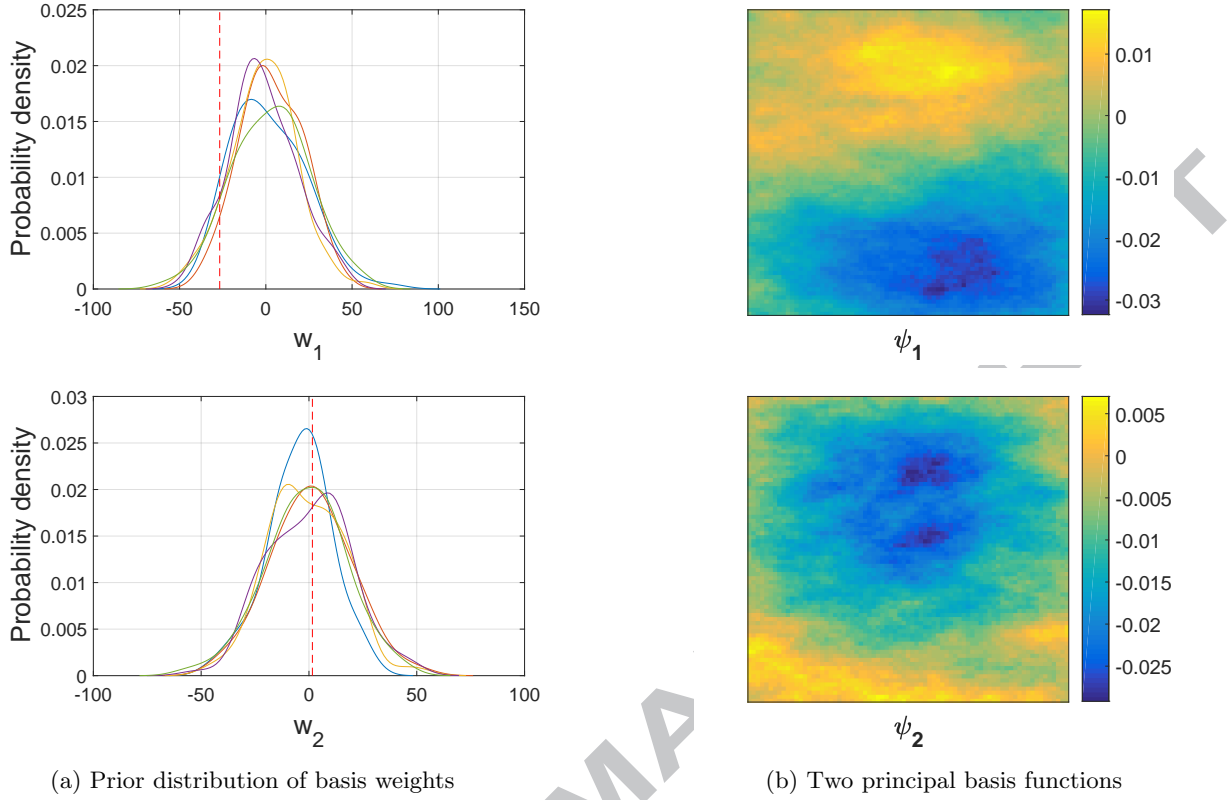


Figure 2: Prior distribution of basis weights from five ensembles and two principle basis functions for log-permeability. Red dashed lines show reference solution and five prior ensembles distribution are shown by five different colors in part (a).

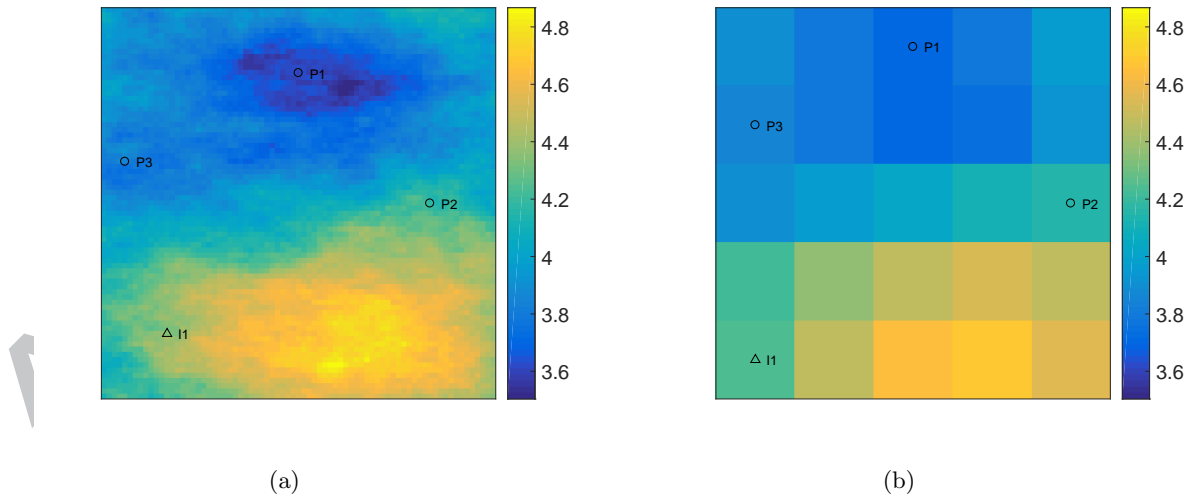


Figure 3: The fine scale (75×75) reference log-permeability (a) and the corresponding up-scaled log-permeability (5×5) using harmonic average (b).

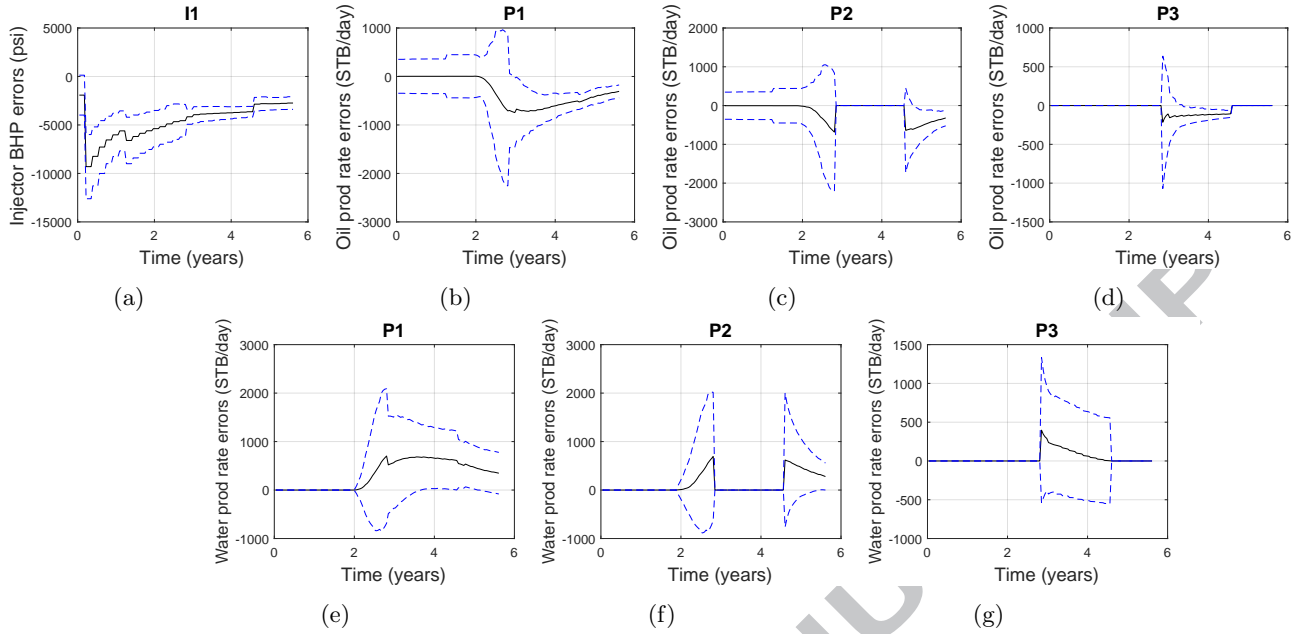


Figure 4: Prior model-error statistics of all wells for Case 1. Black lines show mean model errors, dashed blue lines show the 95% confidence interval (mean plus and minus two standard deviations) of model errors.

3.2 Case 2: Up-scaled imperfect geology model

Geologists commonly try to build geologically realistic prior models. However, maintaining the geological realism during the history-matching process is quite challenging (Sun and Durlofsky, 2017). For example, multipoint statistics (MPS) is widely used to represent channelized geological patterns. Geologically consistent history matching using MPS prior is still a subject of active research (Chen et al., 2016). Sometimes the predictability of the history matched MPS models may not be satisfactory, often due to limitation of the available history matching methods in handling this type of non-Gaussian models (Chen et al., 2016). In this study we do not aim for obtaining calibrated models that are consistent with the channelized geological feature, instead we focus on improving predictability of the calibrated coarse models by including the error-model.

For this case, the permeability fields are based on a similar test case presented in (Chen et al., 2016). Figure 5(a) shows the reference fine scale log-permeability with channelized features and Fig. 5(b) shows the corresponding up-scaled log-permeability field in which the channelized features have been lost due to harmonic averaging. The reference and prior fine scale channelized log-

315 permeability fields are generated using a two-facies training image with the direct sampling version
 316 of MPS (Mariethoz and Caers, 2014). The observed data are generated by the fine scale model
 317 using reference log-permeability field (Fig. 5(a)) with the addition of measurement noise of 2%.

318 Similar to the first test case, one hundred realizations of the model-discrepancy were obtained
 319 using Eq. 11 by running the fine scale simulation using the MPS permeability images of size 75×75
 320 grid blocks and the corresponding up-scaled permeability field with 5×5 grid blocks. Figure 6
 321 shows the prior statistics of the model-discrepancy in the simulated well production data.

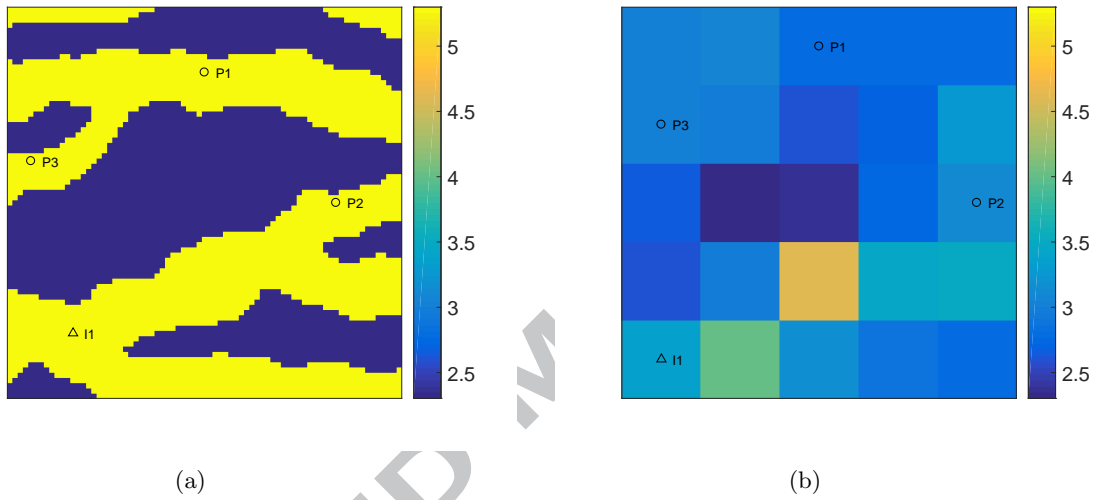


Figure 5: The fine scale (75×75) reference log-permeability with channelized features (a) and the corresponding up-scaled log-permeability (5×5) using harmonic average (b).

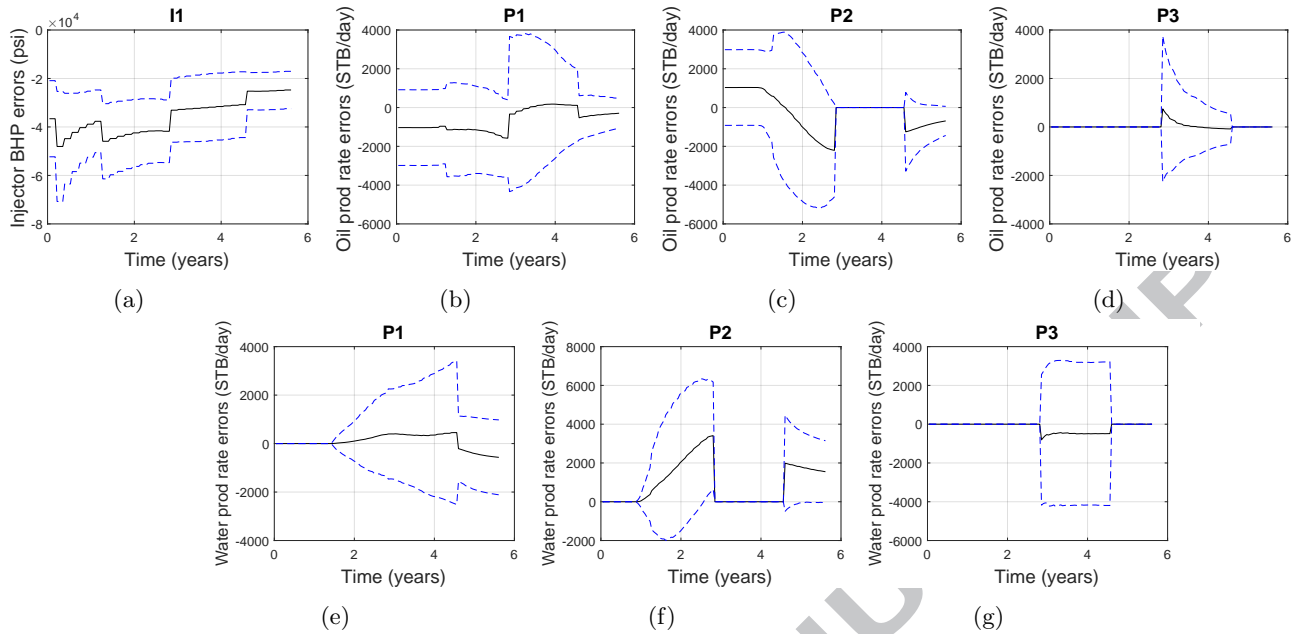


Figure 6: Prior model-error statistics of all wells for Case 2. Black lines show mean model errors, dashed blue lines show the 95% confidence interval (mean plus and minus two standard deviations) of model errors.

322 4 Results and Discussion

323 In this section, we present history matching results for test Case 1 and 2 with and without ac-
 324 counting for the model-discrepancy. All the ensemble-based history matching results are presented
 325 for multiple runs (five independent ensembles) in order to investigate the consistency and relia-
 326 bility of the parameter estimation process. Each ensemble run consists of 100 ensemble members
 327 and the measurement errors are assumed to be 2% of the observation data. We utilized the ES-
 328 MDA algorithm with eight iterations ($\alpha = 8$) for calibration. In the case of joint inversion with
 329 error-model, two PCA components were retained per each output time series to parametrize the
 330 model-discrepancy. Since we have seven output time series, the total number of error-model pa-
 331 rameters is 14.

332 The calibrated models are evaluated using three different forecasting metrics to assess the quality
 333 of the estimated parameters and the capacity of the calibrated models in making future predictions.
 334 The utilized forecasting metrics are: coverage probability (CP), mean continuous ranked probability
 335 score (CRPS) and mean square error (MSE). CP indicates the fraction of the actual data that lie

336 within the 95% confidence interval of the estimation. A value of 0.95 for CP indicates a consistent
 337 estimation of uncertainty and values below 0.95 indicate underestimation of uncertainty. Mean
 338 CRPS quantifies both accuracy and precision (Hersbach, 2000) and higher values of CRPS indicate
 339 a less accurate results. MSE is widely used as a metric for parameter estimation problems. However,
 340 MSE measures the quality of data-fitting and is not enough to provide a probabilistic assessment
 341 of the estimation and prediction from an ensemble of models. In this study, we observed that a
 342 combination of MSE, CP and CRPS provides a good assessment of the quality for the probabilistic
 343 forecast (Skauvold and Eidsvik, 2018). The mathematical formulations of the three forecasting
 344 metrics are listed in Appendix B.

345 4.1 Case 1 Results

346 In this test case, the log-permeability is calibrated in terms of the PCA-basis weights, i.e. w_1, w_2
 347 as detailed in the problem description. Figure 7 shows the posterior distributions of the basis
 348 weights for the five runs, for both cases of neglecting and accounting for model error. The results
 349 presented in Fig. 7(a) show that the posterior distribution from the inversion (neglecting model-
 350 error) procedure are biased and the estimated basis weights do not capture the reference weights.
 351 In contrast, the posterior distributions obtained by the joint inversion procedure are less biased
 352 (i.e. nearly unbiased) and successfully cover the true model parameters as shown in Fig. 7(b).

353 Figure 8 shows the mean and standard deviation of the posterior $\ln(\mathbf{K})$ for test case 1. In
 354 Fig. 8(a) the mean of posterior log-permeability field obtained from five different runs are shown
 355 for the inversion procedure. This posterior mean is clearly different from the coarse scale reference
 356 log-permeability field shown in Fig. 3(b) due to the bias in the inferred posterior distributions.
 357 Figure 8(b) shows the mean posterior log-permeability fields obtained by the joint inversion pro-
 358 cedure. These fields are quite similar to the coarse scale reference log-permeability field. We also
 359 observe that the posterior standard deviations are quite low for the inversion (neglecting model-
 360 error) procedure, which could be a sign of over-fitting the data. In contrast for the joint inversion
 361 procedure, the standard deviations of the posterior fields are much larger due to accounting for the
 362 model error and including the error-model noise covariance matrix \mathbf{C}_T in Eq. 22.

363 Figures 9 and 10 show the oil and water production rates of the different production wells
364 and the bottom hole pressure of the injector well. The 50th percentile p_{50} and 95% confidence
365 interval (the shaded region) are obtained by combining results from all five runs. In part (a) of
366 these figures, the results for the inversion (neglecting model-discrepancy) are presented and the
367 results of the joint-inversion with error-model are shown in part (b). For the inversion (neglecting
368 model-error) procedure, the results are mixed where the data are matched for some cases and not
369 matched for others. For example in Fig. 9(a), the data match is quite good for the wells P1 and
370 P2. However in Fig. 10(a) the data match for well I1 is not as good. Moreover, the predictions
371 for all wells are inaccurate for the inversion (neglecting model-error) procedure. Furthermore,
372 the prediction envelop is really narrow, resulting in invisible confidence interval in the plots, which
373 shows over-confidence in the inaccurate predictions. In comparison, better matches and predictions
374 are obtained by the joint inversion procedure as shown in Figs. 9(b) and 10(b).

375 Figure 11 shows the forecasting metrics (CP for the estimated log-permeabilities and mean
376 CRPS, MSE, CP for the well data in history matching and prediction periods) for individual
377 ensemble (E1 to E5) and for results from all five ensembles assembled together (denoted as “All” in
378 the figure). These metrics provide a good assessment on the consistency, reliability and accuracy of
379 forecasting capacity of the calibrated models. Figures 11(a) and 11(b) show the coverage probability
380 of reservoir model parameters ($\ln(\mathbf{K})$), well data for both the history matching period and the
381 forecasting period. Both the inversion (neglecting model-error) and the joint inversion results are
382 shown. In Fig. 11(a) CP of $\ln(\mathbf{K})$ is zero for each individual ensemble (E1 to E5) as well as for
383 all five ensembles combined meaning that none of the ensemble captures the true log-permeability
384 using the inversion (neglecting model-error) procedure. In Fig. 11(b) CP of $\ln(\mathbf{K})$ is one for E1
385 to E4 as well as for the combined ensemble that means that four out of five runs managed to
386 enclose the reference log-permeabilities completely when using the joint inversion procedure. For
387 the inversion (neglecting model-error) procedure, the CP lies between 0.06–0.11 for the historical
388 data and lies between 0.01–0.013 for the validation data (prediction), as shown in Fig. 11(a). For
389 the joint inversion procedure, these values of CP of are increased to be between 0.29–0.31 and 0.77–
390 0.82, respectively as shown in Fig. 11(b). Although the value of CP equal to one (higher than the

391 correct value of 0.95) for the estimated permeability field when the joint inversion procedure is used
392 clearly indicates that the uncertainty of the permeability field is overestimated, the overall results
393 still show reasonable improvement from the joint inversion procedure compared to the standard
394 inversion (neglecting model-error) procedure.

395 Figures 11(c) and 11(d) show the mean CRPS of history matching and prediction periods of the
396 well data for the inversion (neglecting model-error) and the joint inversion procedures, respectively.
397 Figure 11(c) shows that the mean CRPS lies between 117–118 and 466–469 for history matching and
398 prediction periods, respectively. The results for the inversion (neglecting model-error) is unreliable
399 and inaccurate due to the biased posterior distributions for all the different runs. Figure 11(d) shows
400 that using joint inversion with error-model, the mean CRPS lies between 72–73 and 170–187 for
401 history matching and prediction periods of the well data, respectively. A significant improvement
402 in terms of reliability and accuracy is observed, by incorporating the error-model in the inversion
403 process. Figures 11(e) and 11(f) show the MSE of the individual runs and the combined ensemble
404 of all runs. With the joint inversion procedure, lower MSE values are obtained for both the history
405 matching and the prediction periods (indicated by subscript “h” and “p” respectively in the plot).

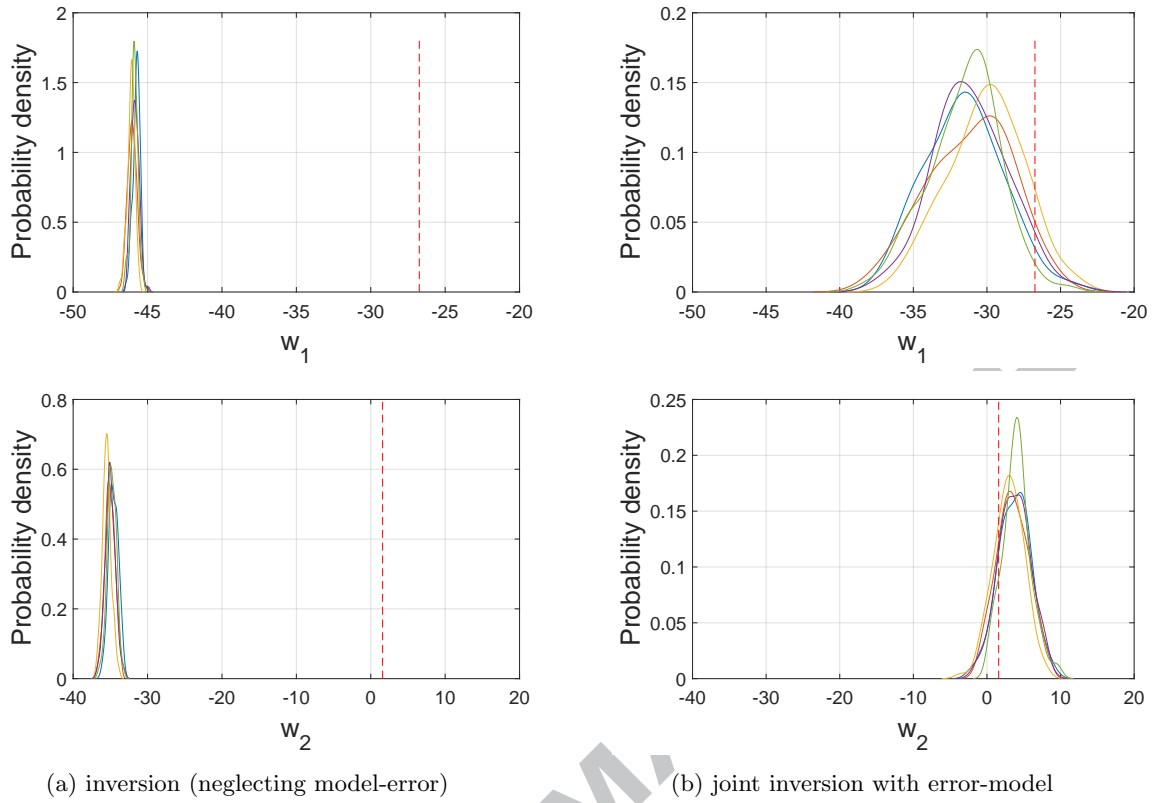


Figure 7: Posterior distribution of two PCA basis weights (of $\ln(\mathbf{K})$) obtained after history matching for coarse scale model case. Dashed red lines show the reference solution and the posterior distribution of the five ensembles are shown by five different colors.

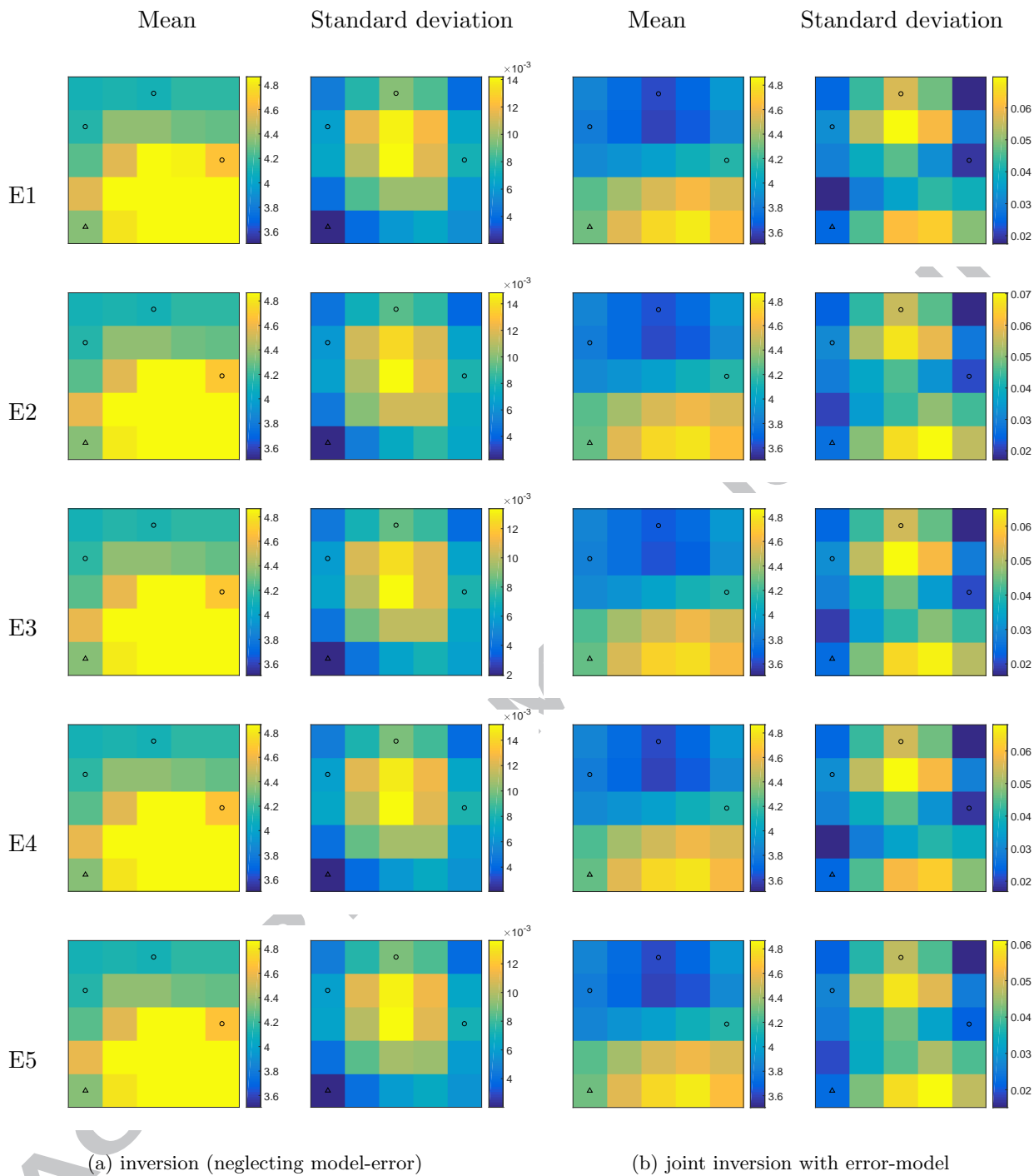


Figure 8: Mean and standard deviation of $\ln(\mathbf{K})$ posterior ensembles obtained after history matching of two PCA basis weights for coarse scale model case.

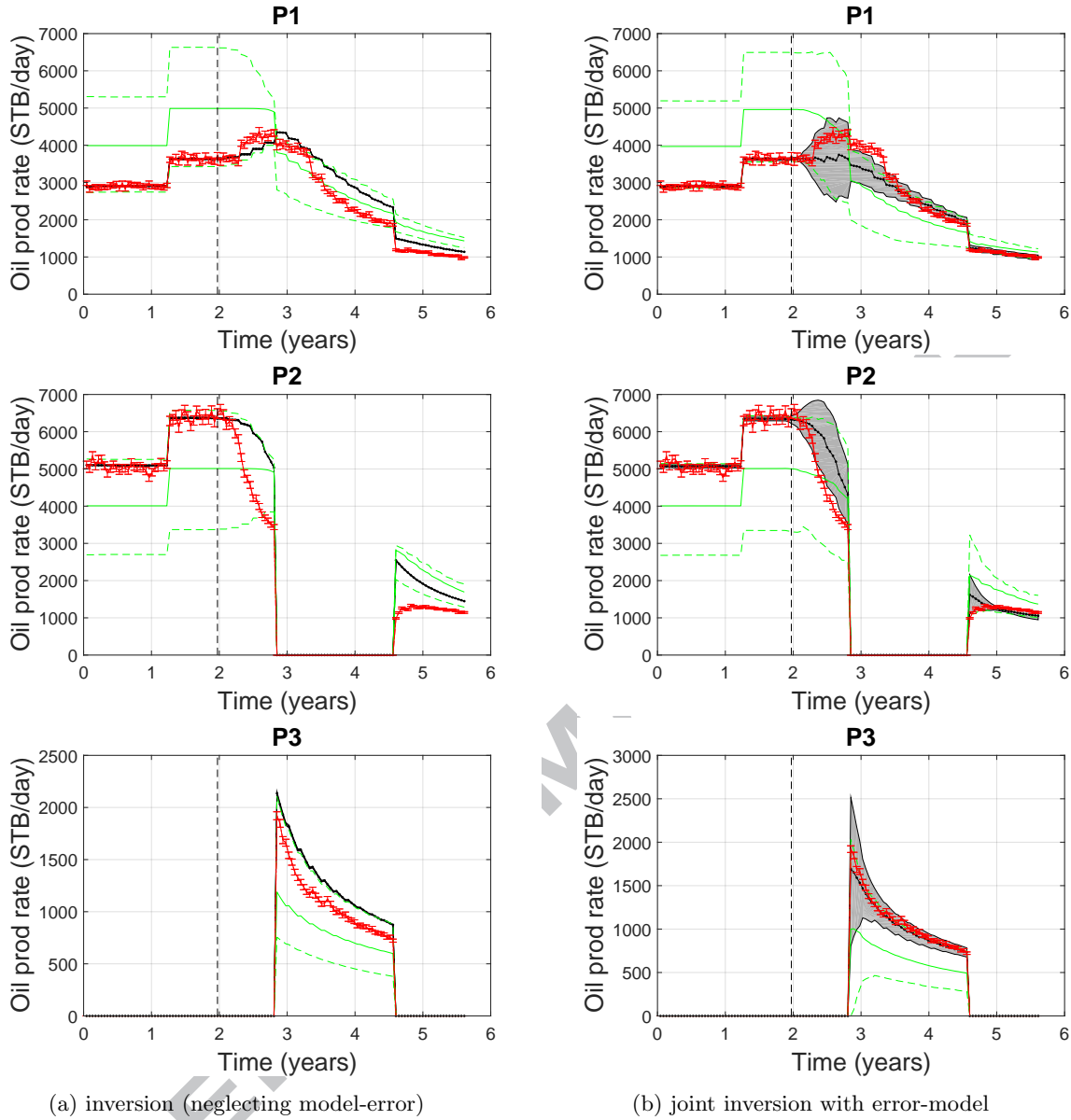


Figure 9: Prior and posterior of oil production data obtained from all ensembles for coarse scale model case. Red lines show observation data and bar on red lines shows measurement error. Dashed black lines show end of historical period. Solid green lines show 50th percentile p_{50} of prior distribution, dashed green lines show 95% confidence interval of prior distribution. Solid black lines show p_{50} of posterior distribution, gray shaded area shows 95% confidence interval of posterior distribution.

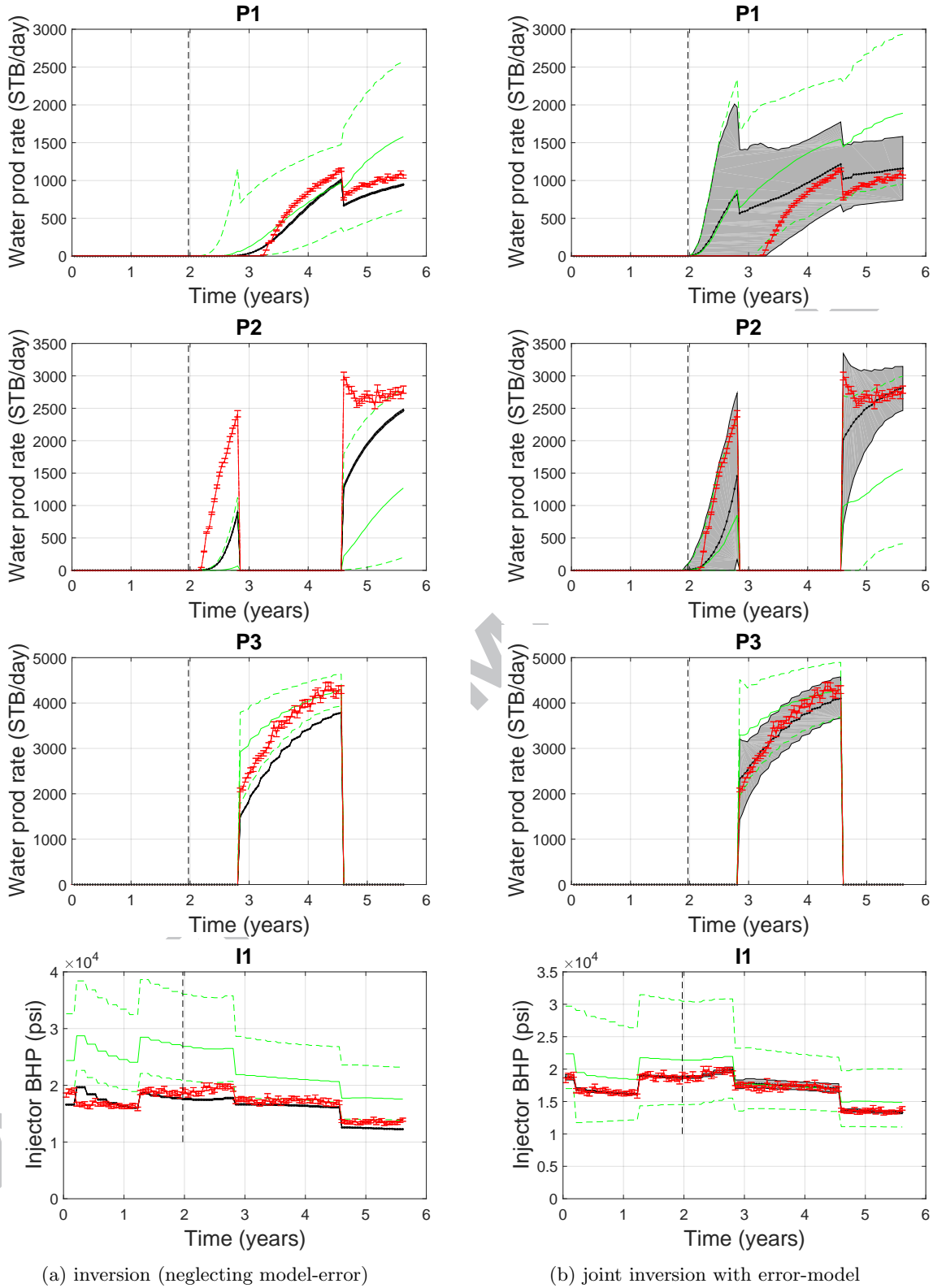


Figure 10: Prior and posterior of water production and injection pressure data for coarse model case. The explanation of colors and lines are the same as in Fig. 9.

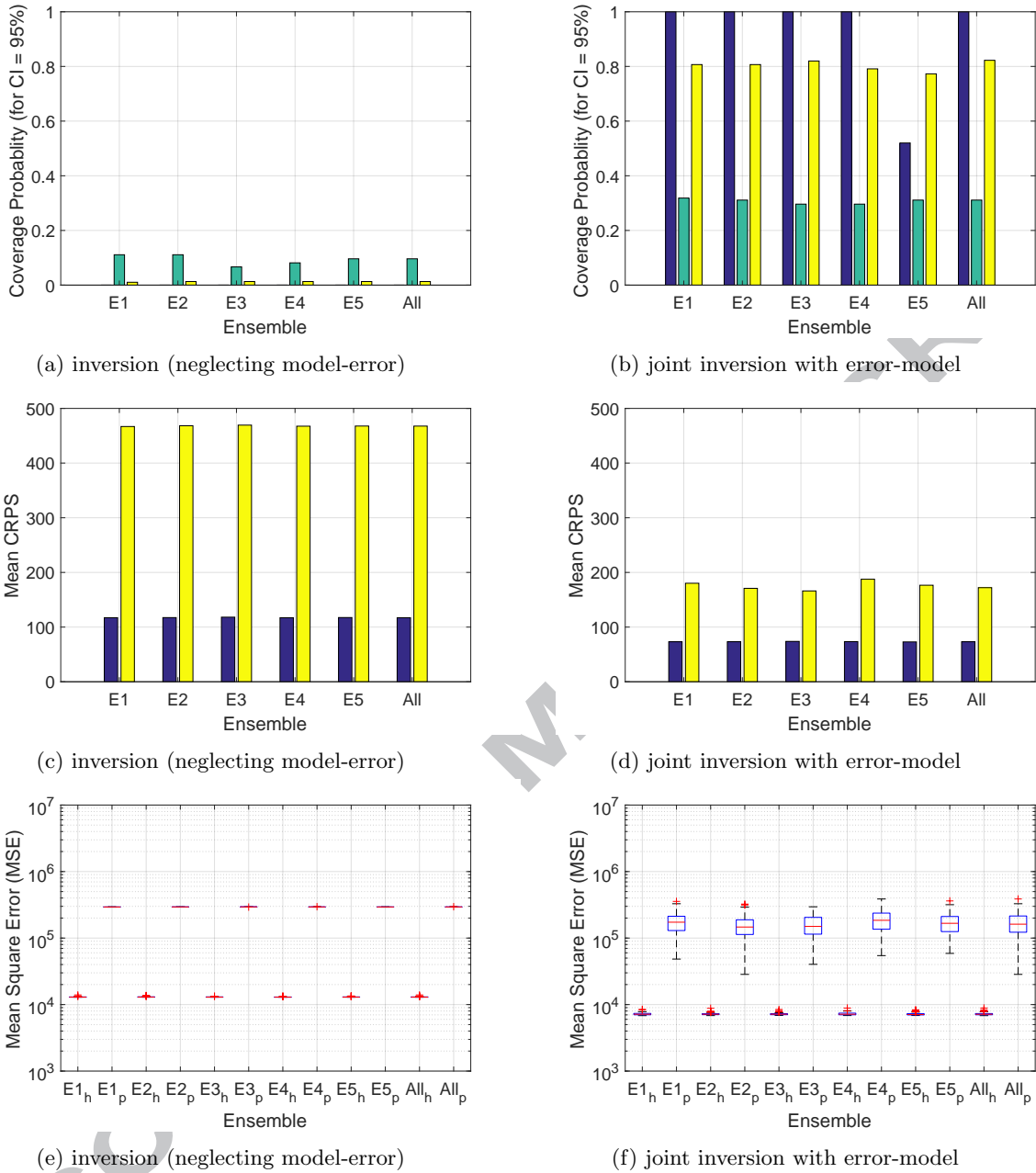


Figure 11: Forecasting metrics of coarse scale model case. In part (a) and (b) blue bars show the CP of true log-permeabilities, green bars show the CP of the historical data and yellow bars show the CP of prediction. In part (c) and (d) blue bars show the mean CRPS of the historical data and yellow bars show the mean CRPS of prediction. In part (e) and (f) box plots of MSE of the simulated well data from each ensemble are shown, subscript h and p are used for history and prediction respectively. On each box, the central red line indicates the median, and the bottom and top blue edges of the box indicate the 25th and 75th percentiles, respectively. The whiskers represent extreme data points without outliers, and '+' symbol represents outliers (more than 1.5 times of interquartile range).

406 **4.2 Case 2 Results**

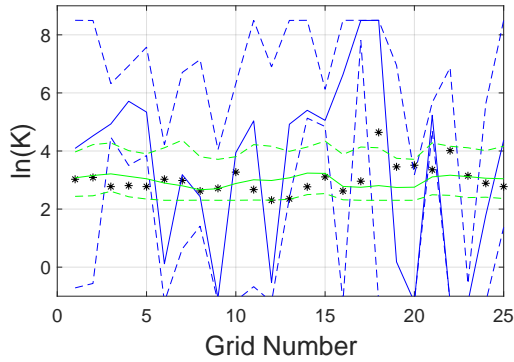
407 In this test case, log-permeability of every grid cell is calibrated using both inversion (neglecting
408 model-error) and joint inversion procedures. Figure 12 shows prior and posterior distributions
409 obtained by combining realizations from all five runs for both history matching procedures. Figure
410 12(a) shows that the posterior distribution of $\ln(\mathbf{K})$ is biased where the estimated log-permeabilities
411 show extreme values and do not capture the reference log-permeabilities. Figure 12(b) shows that
412 by accounting for model error, relatively small changes have been made to the physical parameters
413 (log-permeability in this case), and the mean of the posterior distribution of $\ln(\mathbf{K})$ remains smooth
414 after data assimilation.

415 Figure 13 shows the mean and standard deviation of $\ln(\mathbf{K})$ posterior ensembles as maps. In
416 Fig. 13(a) the posterior mean log-permeability field obtained from five different runs are shown
417 for the inversion (neglecting model-error) procedure. This posterior mean is clearly different from
418 the reference coarse log-permeability field as shown in Fig. 5(b). Moreover, in Fig. 13(a) the mean
419 log-permeability field of every ensemble is different from each other, which is an indication of
420 convergence of every ensemble to different non-unique local peak (mode) of the biased posterior.
421 The mean of log-permeability ensemble from the joint inversion procedure is shown in Fig. 13(b)
422 where no extreme features are observed. Similar to the observation from Case 1, the standard
423 deviation is higher for the joint inversion compared to the inversion (neglecting model-error).

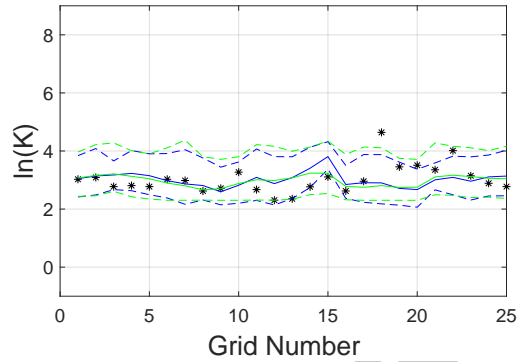
424 Figures 14 and 15 show the oil and water production rates of the different production wells
425 and the bottom hole pressure of the injector well. The 50th percentile p_{50} and 95% confidence
426 interval (the shaded region) are obtained by combining results from all five ensembles. In part (a)
427 of these figures, the results for the inversion (neglecting model-discrepancy) are presented and the
428 results of the joint-inversion with error-model are shown in part (b). For the inversion (neglecting
429 model-error) procedure, only historical data at some wells are matched. For example Figs. 14(a)
430 and 15(a) show that the data match is good for wells I1 and P1, however the data match of well
431 P2 is not good. Moreover the future prediction from the estimated model parameters using the
432 inversion (neglecting model-error) procedure is quite poor. A good example is the prediction of
433 water production rate of well P1: the models predicted early water breakthrough between year 2 to

3, while the actual water breakthrough is after year 5 because P1 is separated from the injector by a low permeability region (see Fig. 5(a)). In comparison, better matches and predictions are obtained by the joint inversion procedure as shown in Figs. 14(b) and 15(b) for individual ensembles as well as all ensembles combined together. The prediction from the combined multiple ensembles may seem good for some well data for the case with the inversion (neglecting model-error) procedure, for example BHP pressure of well I1 in Fig. 15(a), even though the prediction from each individual ensemble is not good. This is often due to the fact that different ensemble converges to different local peaks (modes) of the biased posterior and the combined prediction from these multiple biased final ensembles happen to enclose the validation data.

Figure 16 shows the forecasting metrics (CP for the estimated log-permeabilities and mean CRPS, MSE, CP for the well data in history matching and prediction periods) for the individual ensembles and for results from all five ensembles assembled together. Figures 16(a) and 16(b) show the coverage probability of reservoir model parameters ($\ln(\mathbf{K})$), well data for both the history matching period and the forecast period. In Fig. 16(a) CP of $\ln(\mathbf{K})$ is between 0–0.12 for ensembles (E1 to E5), however CP of $\ln(\mathbf{K})$ is 0.6 for the combined ensembles. This relatively high coverage from the combined ensemble is due to the overshooting of $\ln(\mathbf{K})$ values and the different final estimation from each individual ensemble as shown in Fig. 12(a) and Fig. 13(a) respectively. With the joint inversion procedure, the coverage probability is improved for all three quantities investigated (log-permeabilities, historical data, future prediction). Similarly, the mean CRPS and MSE measures also show significant improvement by accounting for model error using the joint inversion procedure (2nd and 3rd row of Fig. 16). In addition, based on all three forecasting measures, the results from multiple ensemble runs using the joint inversion procedure are very consistent, which indicates the statistical consistency of the proposed procedure.



(a) inversion (neglecting model-error)



(b) joint inversion with error-model

Figure 12: Prior and posterior distribution of $\ln(\mathbf{K})$ obtained after history matching for up-scaled imperfect geology model case using five ensembles. In both part (a) and (b), green and blue lines show the prior and posterior distribution respectively. Solid green and blue line show the $p50$ prior and posterior respectively. Dashed green and blue lines show the 95% confidence interval of prior and posterior respectively. Black asterisks show the reference solution.

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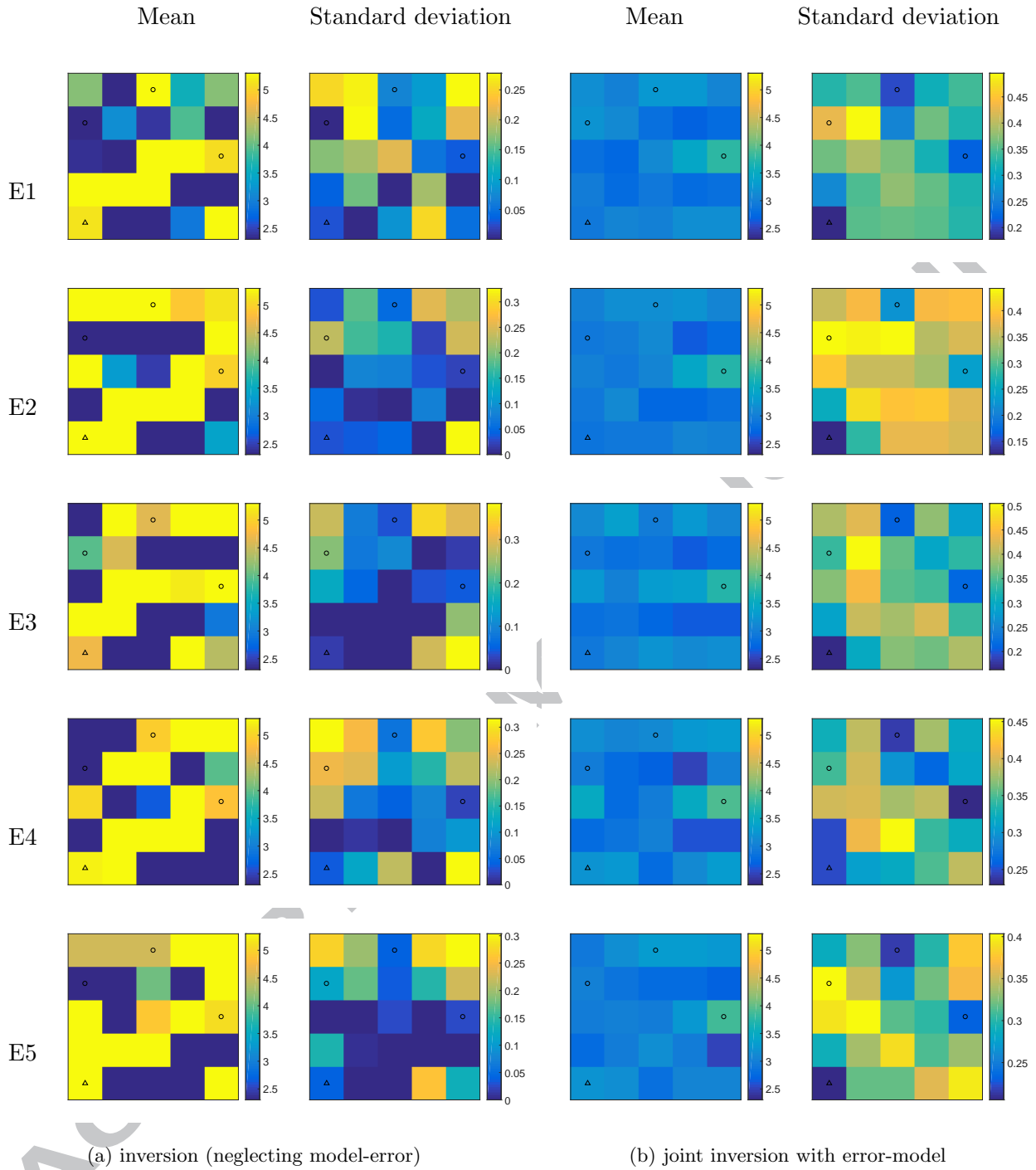


Figure 13: Mean and standard deviation of $\ln(\mathbf{K})$ posterior ensembles obtained after history matching of all grids log-permeabilities for up-scaled imperfect geology model case.

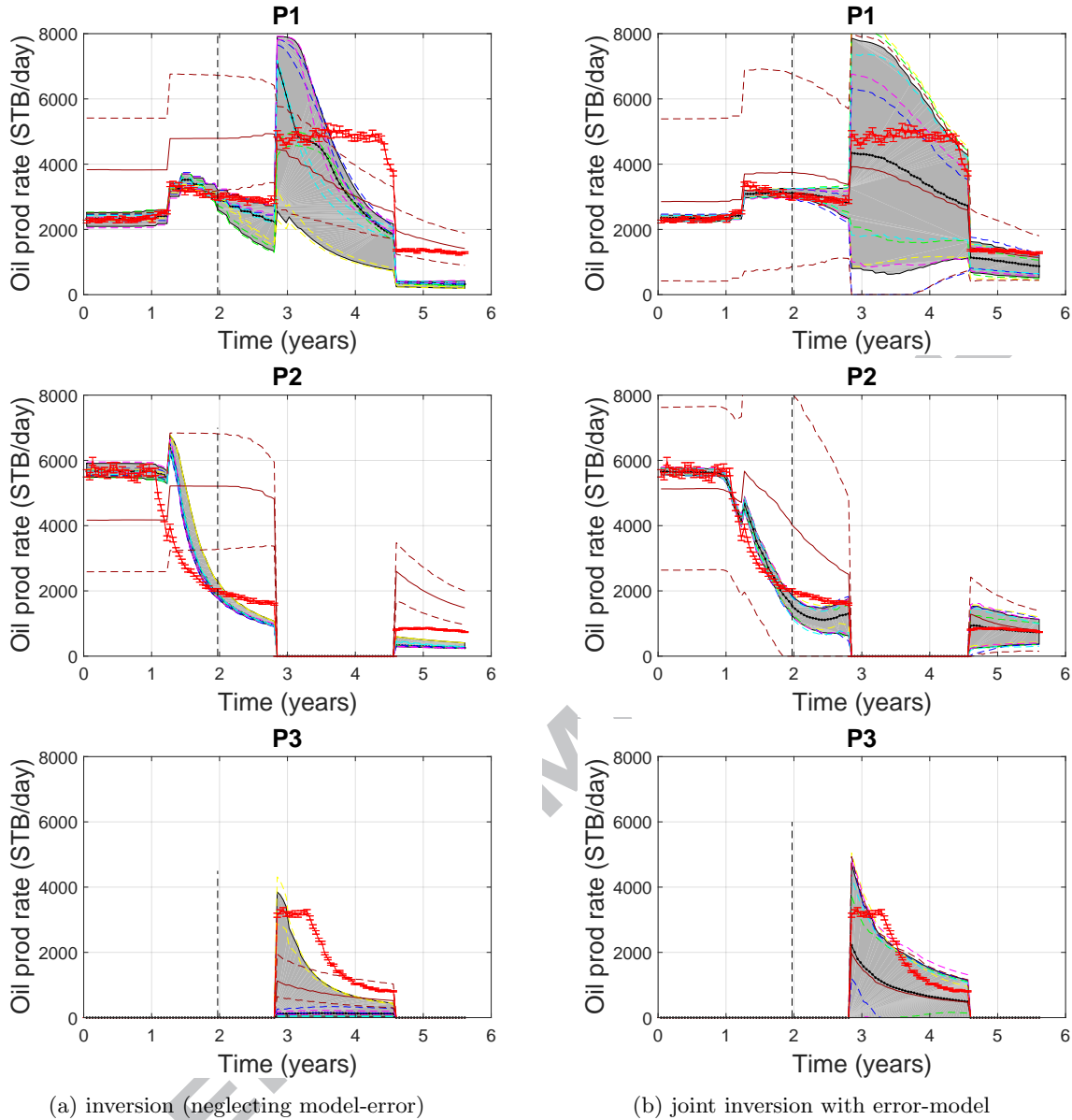
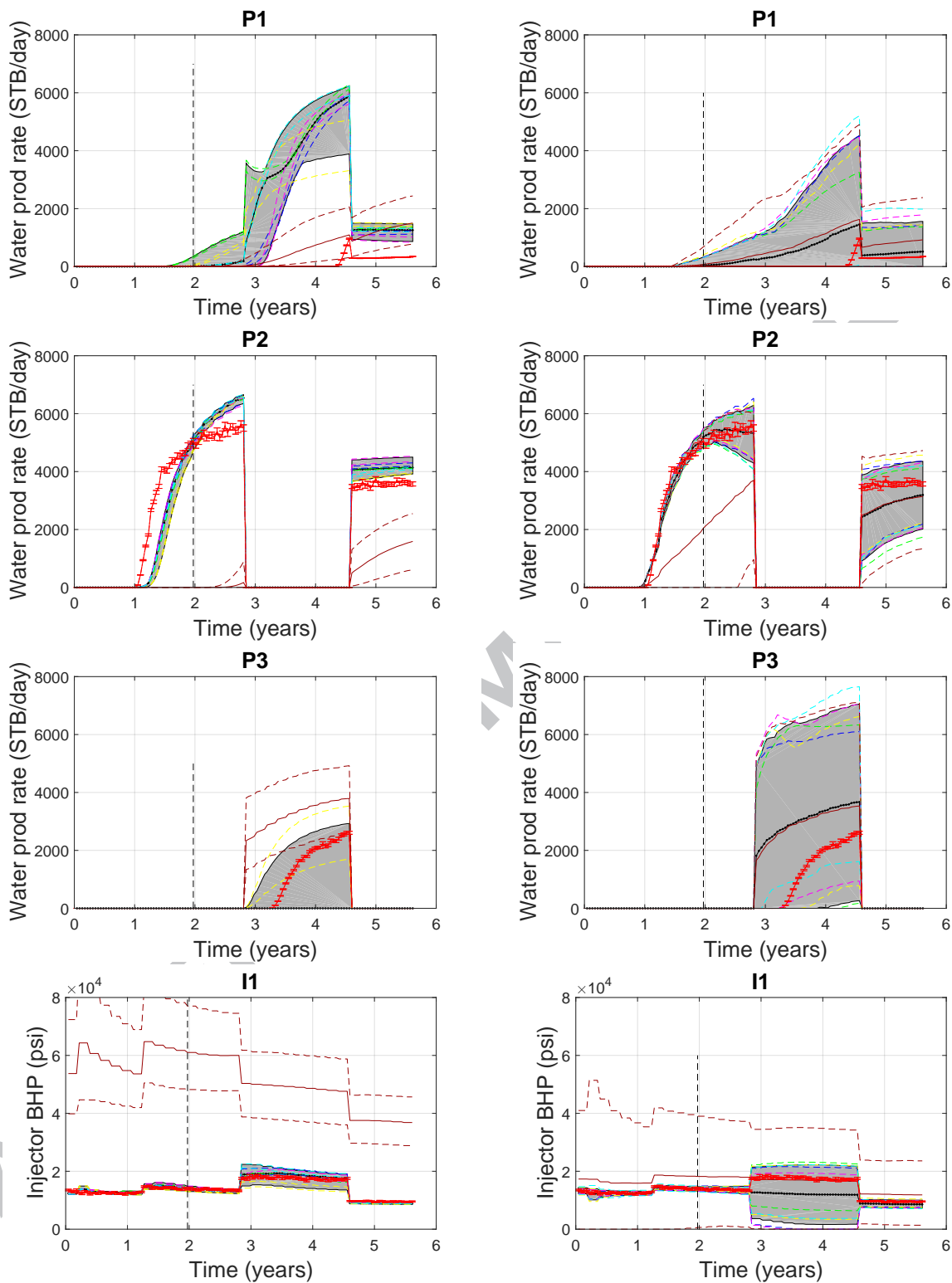


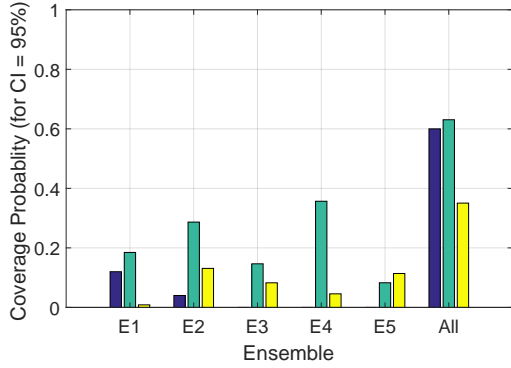
Figure 14: Prior and posterior of oil production data for up-scaled imperfect geological model case. Red lines show observation data and bar on red lines shows measurement error. Dashed black lines show end of historical period. Solid brown lines show 50th percentile p_{50} of the prior distribution, dashed brown lines show 95% confidence interval of prior distribution. Solid black lines show p_{50} posterior distribution obtained from all ensembles. Shaded gray area show 95% confidence interval of posterior distribution obtained from all ensembles. Dashed blue, green, yellow, magenta and cyan lines show 95% confidence interval of posterior distribution of individual ensembles.



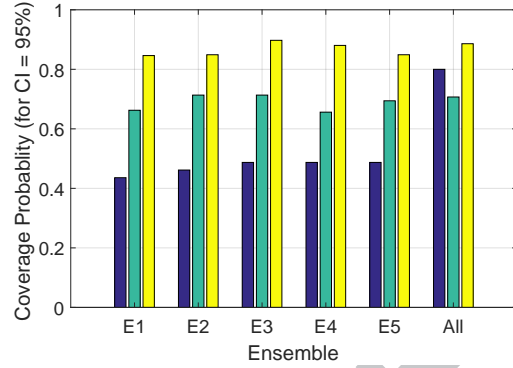
(a) inversion (neglecting model-error)

(b) joint inversion with error-model

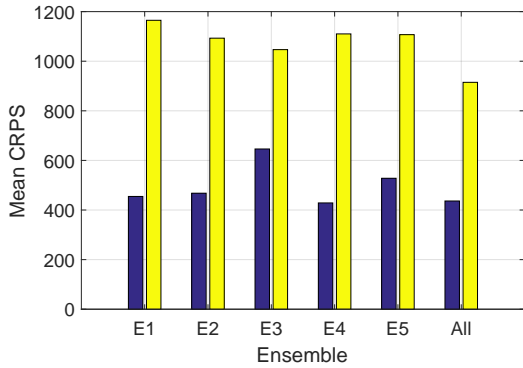
Figure 15: Prior and posterior of water production and injection pressure data for up-scaled imperfect geological model case. The explanation of colors and lines are the same as in Fig. 14.



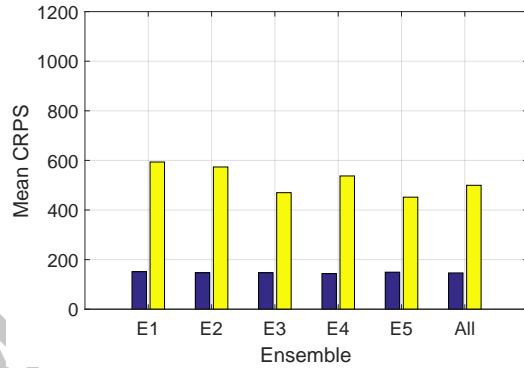
(a) inversion (neglecting model-error)



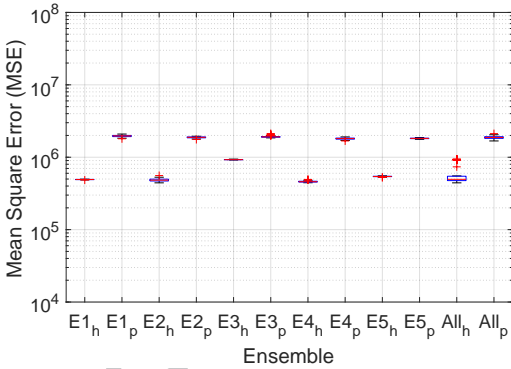
(b) joint inversion with error-model



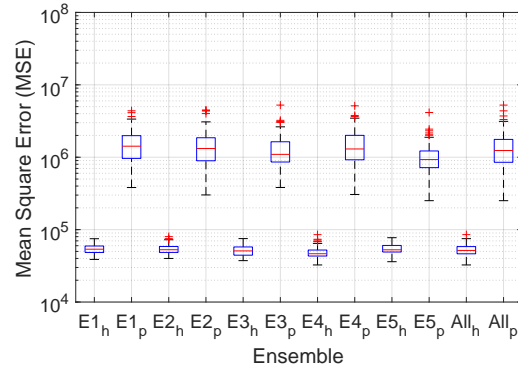
(c) inversion (neglecting model-error)



(d) joint inversion with error-model



(e) inversion (neglecting model-error)



(f) joint inversion with error-model

Figure 16: Forecasting metrics of up-scaled imperfect geology model case. In part (a) and (b) blue bars show the CP of true log-permeabilities, green bars show the CP of the historical data and yellow bars show the CP of prediction. In part (c) and (d) blue bars show the mean CRPS of the historical data and yellow bars show the mean CRPS of prediction. In part (e) and (f) box plots of MSE of the simulated well data from each ensemble are shown, subscript h and p are used for history and prediction respectively. On each box, the central red line indicates the median, and the bottom and top blue edges of the box indicate the 25th and 75th percentiles, respectively. The whiskers represent extreme data points without outliers, and '+' symbol represents outliers (more than 1.5 times of interquartile range).

457 5 Conclusions

458 In this paper, a generic procedure for history matching of imperfect/low-fidelity reservoir models
459 has been developed where, we formulate the history matching problem as a joint inversion of
460 reservoir model parameters and an error model parameters. We used principal component analysis
461 to parameterize the error model, where the PCA basis function and prior statistics of the PCA basis
462 weights were obtained using pairs of accurate and inaccurate models. We note that the accurate
463 model is only used for defining the prior model-error statistics and during history matching only
464 the imperfect/low-fidelity model is used.

465 We evaluated the proposed history matching procedure on low-fidelity models with modeling
466 errors due to aggressive grid coarsening/up-scaling of the permeability field obtained from two-point
467 statistics and low-fidelity models where the main source of error is the grid coarsening/up-scaling
468 of a channelized geology. Detailed comparison were performed against standard history matching
469 (inversion while neglecting model error). The obtained results show that the estimated model
470 parameters are biased using standard history matching procedure in the presence of large modeling
471 errors. Consequently the calibrated low-fidelity model predictions are unreliable and generally
472 inaccurate. Utilizing the developed joint inversion procedure results in significant improvements
473 in terms of the quality of the estimated parameters, the matching capacity to historical data
474 and prediction accuracy/reliability of the calibrated low-fidelity models. This is attributed to a
475 reduction (and in some cases elimination) of the bias in the estimated posterior distribution of
476 the model parameters when we included a flexible error-model terms in the inversion process.
477 The numerical test cases were assessed using three forecasting metrics and it was observed that
478 the consistency of ensemble-based history matching technique was also improved by including the
479 error-model terms in the inversion procedure. We argue that this observed consistency might be
480 due to the elimination of multiple biased peaks (modes) in the posterior distribution of the model
481 parameters once the error modeling terms are included in the formulations.

482 The proposed framework is generally flexible and could be applied to large scale models as
483 the error-model formulation is I/O independent and the prior error-model parameters could be
484 estimated before the history matching step. However, for general error-modeling an accurate model

485 may be missing or the sources of the modeling errors could be unknown. It is also possible that the
 486 fine/high-fidelity model (which is assumed to be perfect) is also biased. In these cases, the proposed
 487 methodology can only improve the parameter estimation and the prediction up to the fine/high-
 488 fidelity model accuracy. Addressing the effects of unknown modeling errors without relying on an
 489 accurate (high-fidelity)/approximate (low-fidelity) model pairs is the subject of our future work.

490 6 Acknowledgments

491 The first author thanks Total E&P, UK for the financial support. The authors acknowledge Total
 492 S.A. for authorizing the publication of this paper.

493 Appendix A: Reservoir properties

494 Corey model in form of power law is used to generate relative permeability data for the reservoir
 495 model. Mathematically Corey model in form of power law is written as follows.

$$k_{rw} = (\hat{S}_w)^{n_w} k_w^0. \quad (\text{A.1})$$

$$k_{ro} = (1 - \hat{S}_w)^{n_o} k_o^0. \quad (\text{A.2})$$

$$\hat{S}_w = \frac{S_w - S_{wc}}{1 - S_{or} - S_{wc}}. \quad (\text{A.3})$$

496 The notations of above equations are described in MRST manual Lie (2016). The fluid data
 497 and core relative permeability model parameters used in the reservoir model are shown in the
 498 Table A.1.

Table A.1: Reservoir fluid data and Corey relative permeability model parameters

Fluid properties		Corey relative permeability model parameters			
water viscosity	0.5 cp	S_{or}	0.2	k_o^0	1
oil viscosity	1 cp	S_{wc}	0.2	k_w^0	1
water density	1000 kg/m ³	n_w	2		
oil density	700 kg/m ³	n_o	2		

499 Appendix B: Forecasting metrics

500 B.1: Mean Square Error (MSE)

501 The Mean Square Error (MSE) is obtained using,

$$MSE = \frac{1}{N_d} \sum_{n=1}^{N_d} (d_n - d_{obs,n})^2, \quad (B.1)$$

502 where n is index of observation or model prediction at corresponding time.

503 B.2: Coverage Probability (CP)

$$CP = \frac{N_{CI}}{N_t}. \quad (B.2)$$

504 N_{CI} = Number of samples, parameters or observations in Confidence Interval

505 N_t = Total number of samples, parameters or observations

506 B.3: Continuous Ranked Probability Score (CRPS)

507 The details of CRPS for ensemble prediction system were described by Hans Herbach (2000) Hers-
508 bach (2000). In this section summary of CRPS is explained.

509 Mathematically CRPS can be defined as,

$$CRPS = \int_{-\infty}^{\infty} [p(x) - H(x - x_{obs})]^2 dx, \quad (B.3)$$

510 where $p(x) = \int_{-\infty}^x \rho(y) dy$ Cumulative distribution of quantity of interest, $H(x - x_{obs}) =$ Heaviside

511 function (Step function) i.e.

$$H(x) = \begin{cases} 0 & \text{if } x < 0 \\ 1 & \text{if } x \geq 0 \end{cases}$$

512 For an ensemble system with N_e realizations, the CRPS can be written as follows,

$$CRPS = \sum_{i=0}^{N_e} c_i. \quad (\text{B.4})$$

$$c_i = \alpha_i p_i^2 + \beta_i (1 - p_i)^2. \quad (\text{B.5})$$

513 where $p_i = P(x) = i/N_e$, for $x_i < x < x_{i+1}$ (Cumulative distribution is a piece wise constant
514 function).

$$\alpha_i = \begin{cases} 0 & \text{if } x_{obs} < x_i \\ x_{obs} - x_i & \text{if } x_i < x_{obs} < x_{i+1} \\ x_{i+1} - x_i & \text{if } x_{obs} > x_{i+1} \\ x_{obs} - x_{N_e} & \text{if } x_{obs} > x_{N_e} \\ 0 & \text{if } x_{obs} < x_1 \end{cases}$$

$$\beta_i = \begin{cases} x_{i+1} - x_i & \text{if } x_{obs} < x_i \\ x_{i+1} - x_{obs} & \text{if } x_i < x_{obs} < x_{i+1} \\ 0 & \text{if } x_{obs} > x_{i+1} \\ 0 & \text{if } x_{obs} > x_{N_e} \\ x_1 - x_{obs} & \text{if } x_{obs} < x_1 \end{cases}$$

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