

PROCEEDINGS OF SPIE

SPIDigitalLibrary.org/conference-proceedings-of-spie

Use of statistical study methods for the analysis of the results of the imitation modeling of radiation transfer

M. A. Alekseenko, I. Yu. Gendrina

M. A. Alekseenko, I. Yu. Gendrina, "Use of statistical study methods for the analysis of the results of the imitation modeling of radiation transfer," Proc. SPIE 10466, 23rd International Symposium on Atmospheric and Ocean Optics: Atmospheric Physics, 104661P (30 November 2017); doi: 10.1117/12.2287848

SPIE.

Event: XXIII International Symposium, Atmospheric and Ocean Optics, Atmospheric Physics, 2017, Irkutsk, Russian Federation

USE OF STATISTICAL STUDY METHODS FOR THE ANALYSIS OF THE RESULTS OF THE IMITATION MODELING OF RADIATION TRANSFER.

M. A. Alekseenko¹, I. Yu. Gendrina²

ABSTRACT

Recently, due to the abundance of various types of observational data in the systems of vision through the atmosphere and the need for their processing, the use of various methods of statistical research in the study of such systems as correlation-regression analysis, dynamic series, variance analysis, etc. is actual. We have attempted to apply elements of correlation-regression analysis for the study and subsequent prediction of the patterns of radiation transfer in these systems same as in the construction of radiation models of the atmosphere. In this paper, we present some results of statistical processing of the results of numerical simulation of the characteristics of vision systems through the atmosphere obtained with the help of a special software package.¹

Vision system, angular brightness distribution, Monte Carlo method, regression.

Introduction

Vision system is understood as an observation scheme including the underlying surface, a "cloudy environment" (atmosphere), and an optical device that captures incoming radiation. To study radiation transfer in such systems, the theory of systems and the theory of radiation transfer are traditionally used.²

The model of the atmosphere includes the following characteristics: the total attenuation coefficient $\sigma(\lambda, \vec{r}) = \sigma_{sc}(\lambda, \vec{r}) + \sigma_a(\lambda, \vec{r})$, where σ_{sc} – the scattering coefficient, σ_a – the absorption coefficient; $g(\lambda, \mu, \vec{r})$ – aerosol phase function. Here $\vec{r} = (x, y, z)$ – radius-vector of the current point in space, $\mu(\vec{\omega}', \vec{\omega})$ – cosine of the scattering angle of radiation coming from direction $\vec{\omega}'$, in the direction $\vec{\omega}$, λ – is the wavelength of incident radiation.

The paper considers two models of the atmosphere:

1. vertically bounded plane-parallel layer-homogeneous aerosol-molecular;
2. vertically bounded plane-parallel aerosol-molecular, including overcast layer. For the cloud layer, different characteristics from those of the first model is assumed: coefficients of attenuation, absorption, scattering, and the aerosol phase function.

1. Problem Statement and methods

The geometric scheme of calculations is as follows: at the lower boundary of the atmosphere (on the underlying surface $z=0$) there is a point source of unit capacity. At the upper boundary of the atmosphere ($z=30\text{ km}$) there is an optical receiver that can receive scattered radiation coming from different directions (observation angles). The brightness of the scattered radiation is solution of the integro-differential transport equation³, which can be practically solved only by approximate or numerical methods.

One of the most universal methods for solving this problem is the simulation method, or the Monte Carlo method. The basis of the Monte Carlo method is the integral transport equation of the second kind with a generalized kernel for the density of particles' collisions³:

$$f(\vec{x}) = \int_X k(\vec{x}', \vec{x}) f(\vec{x}') d\vec{x}' + \psi(\vec{x}), \quad f = Kf + \psi$$

Here $\vec{x} = (\vec{r}, \vec{\omega})$ – is the point of the phase space of coordinates and directions, $\psi(\vec{x})$ – source function, K – integral operator with kernel $k(\vec{x}', \vec{x})$:

$$k(\vec{x}', \vec{x}) = \frac{\sigma_{sc}(\vec{r}) \cdot g(\mu) \exp(-\tau(\vec{r}', \vec{r})) \sigma(\vec{r})}{\sigma(\vec{r}') 2\pi |\vec{r} - \vec{r}'|^2} \cdot \delta\left(\vec{\omega} - \frac{\vec{r} - \vec{r}'}{|\vec{r} - \vec{r}'|}\right).$$

In this paper, one of the algorithms of the Monte Carlo method was used - the method of local estimation³.

¹ National Research Tomsk State University, Tomsk, Russia, Post-Graduate, masha_af6@mail.ru

² National Research Tomsk State University, Tomsk, Russia, Ass. Prof., Ph.D., igendrina@bk.ru

Algorithm for local estimation consists in calculating following functional:

$$J(\Omega_i) = \int_{\Omega_i} \Phi(\vec{r}^*, \vec{\omega}^*) d\vec{\omega}^* = \int_X l_i(\vec{x}', \vec{x}^*) f(\vec{x}') d\vec{x}' = M \sum_{n=0}^N Q_n \cdot l_i(\vec{x}_n, \vec{x}^*), \quad (1)$$

$$l_i(\vec{x}, \vec{x}^*) = \frac{\exp(-\tau(\vec{r}, \vec{r}^*)) \cdot g(\mu^*)}{2\pi |\vec{r} - \vec{r}^*|^2} \Delta_i(\vec{s}^*). \quad (2)$$

Here $\vec{s}^* = \frac{\vec{r}^* - \vec{r}}{|\vec{r}^* - \vec{r}|}$, $\mu^* = (\vec{\omega}, \vec{s}^*)$, $\Delta_i(\vec{s})$ – is the indicator of region Ω_i . Φ – flux of particles at given

point \vec{x}^* . Q_n – weight of the particle, $f(\vec{x})$ – density of collisions.

2. Initial data

We will consider the process of radiative transfer through aerosol-molecular atmosphere, which comprises a layer overcast, by neglecting the reflection from underlying surface. We used the following data¹:

1. Wavelength (mkm) in transparent windows: 0.347; 0.530; 0.694; 0.860; 1,060; 3,390; 10.60.
2. Lower boundary of atmosphere 0 km above Earth's surface, upper boundary H of the atmosphere 30 km above the Earth's surface.
3. Optical thickness for a cloudless atmosphere are presented in Table 1.

Table 1. **Optical thickness of the cloudless atmosphere**

| Wavelength, mkm | Optical thickness |
|-----------------|-------------------|
| 0,347 | 0,228 |
| 0,53 | 0,158 |
| 0,694 | 0,124 |
| 0,86 | 0,098 |
| 1,06 | 0,092 |
| 3,39 | 0,067 |
| 10,6 | 0,041 |

4. Lower boundary of the cloud layer - 1 km above the Earth's surface, upper boundary - 2 km above the Earth's surface. The optical models of the cloud layer – “haze H” and “cloud C1”.⁴

5. In this work, we considered two models of sources of radiation: Lambertian and isotropic sources. In case of isotropic source, density of the initial areas looks like: $p(\vec{\omega}) = \frac{1}{2\pi}$. For Lambertian this value is defined as: $p(\vec{\omega}) = \frac{\mu}{\pi}$, where $\mu = \arccos \theta$, θ - zenith angle of initial direction.

3. Simulation results

Quantitative values brightness of scattered radiation for the cloudless atmosphere are presented in our previous publication.⁵ Similar values for various models of the cloud atmosphere are given in Tables 2,3.

Table 2 contains the results for the brightness of scattered radiation for the atmosphere with a cloud layer of the "Haze H" type.

Table 3 contains similar data for the atmosphere with a cloud layer of the "Cloud C1" type.

These types vary in value of the average cosine of scattering phase function:

$$\bar{\mu} = \frac{1}{2} \int_{-1}^1 \mu g(\mu) d\mu.$$

It is known that this parameter characterizes the elongation of aerosol phase function. For example in case $\lambda = 0,694$ mkm the average cosine amounts to 0,745 for type “Haze H” and 0,857 for type “Cloud C1”.

Table 2. Brightness of scattered radiation for the atmosphere with a cloud layer of the "Haze H" type, $W/mkm \cdot m^{2e}$

| Angles of reception, grad | Wavelength, mkm | | | | | | |
|---------------------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|
| | $\lambda=0,374$ | $\lambda=0,530$ | $\lambda=0,694$ | $\lambda=0,860$ | $\lambda=1,060$ | $\lambda=3,390$ | $\lambda=10,60$ |
| 4,5 | 2,24E-05 | 1,67E-05 | 1,38E-05 | 1,13E-05 | 1,19E-05 | 5,55E-06 | 3,51E-06 |
| 13,5 | 1,09E-06 | 7,65E-07 | 6,60E-07 | 5,30E-07 | 4,56E-07 | 2,09E-07 | 1,43E-07 |
| 22,5 | 2,23E-07 | 1,83E-07 | 1,55E-07 | 1,39E-07 | 1,12E-07 | 4,84E-08 | 3,51E-08 |
| 31,5 | 8,44E-08 | 5,63E-08 | 5,54E-08 | 4,73E-08 | 3,85E-08 | 1,95E-08 | 1,32E-08 |
| 40,5 | 3,49E-08 | 2,73E-08 | 2,81E-08 | 6,65E-08 | 4,33E-08 | 8,70E-09 | 5,38E-09 |
| 49,5 | 1,86E-08 | 5,15E-08 | 1,54E-08 | 1,18E-08 | 9,30E-09 | 4,23E-09 | 4,59E-09 |
| 58,5 | 1,30E-08 | 1,10E-08 | 1,71E-08 | 7,75E-09 | 6,64E-09 | 3,58E-09 | 3,57E-09 |
| 67,5 | 1,94E-08 | 1,12E-08 | 7,95E-09 | 2,11E-08 | 1,73E-08 | 8,16E-08 | 6,37E-09 |
| 76,5 | 1,58E-07 | 2,13E-08 | 4,96E-09 | 3,68E-09 | 2,90E-09 | 1,99E-09 | 1,18E-09 |
| 85,5 | 1,20E-08 | 6,37E-09 | 1,23E-09 | 1,03E-09 | 1,00E-09 | 5,13E-11 | 1,16E-11 |

Table 3. Brightness of scattered radiation for the atmosphere with a cloud layer of the "Cloud C1" type, isotropic source, $W/mkm \cdot m^2$

| Angles of reception, grad | $\lambda = 0,374$ mkm | $\lambda = 0,53$ mkm | $\lambda = 0,694$ mkm | $\lambda = 0,86$ mkm | $\lambda = 1,06$ mkm | $\lambda = 3,39$ mkm | $\lambda = 10,6$ mkm |
|---------------------------|-----------------------|----------------------|-----------------------|----------------------|----------------------|----------------------|----------------------|
| 4,5 | 4,24E-07 | 5,01E-07 | 5,45E-07 | 5,69E-07 | 5,18E-07 | 4,91E-07 | 3,38E-07 |
| 13,5 | 1,32E-07 | 1,36E-07 | 1,01E-07 | 9,47E-08 | 8,07E-08 | 3,59E-08 | 4,21E-08 |
| 22,5 | 5,06E-08 | 4,30E-08 | 4,23E-08 | 3,50E-08 | 2,85E-08 | 8,39E-09 | 9,36E-09 |
| 31,5 | 2,29E-08 | 1,84E-08 | 1,20E-08 | 1,19E-08 | 7,86E-09 | 3,73E-09 | 2,66E-09 |
| 40,5 | 7,00E-09 | 8,43E-09 | 6,45E-09 | 5,16E-09 | 3,70E-09 | 6,67E-10 | 9,83E-10 |
| 49,5 | 3,29E-09 | 2,44E-09 | 3,10E-09 | 2,69E-09 | 1,97E-09 | 5,41E-10 | 4,98E-10 |
| 58,5 | 1,84E-09 | 1,59E-09 | 1,36E-09 | 1,39E-09 | 1,19E-09 | 3,49E-10 | 2,76E-10 |
| 67,5 | 1,17E-09 | 1,13E-09 | 7,77E-10 | 8,25E-10 | 7,02E-10 | 2,17E-10 | 1,37E-10 |
| 76,5 | 8,77E-10 | 9,73E-10 | 7,45E-10 | 5,32E-10 | 5,11E-10 | 9,01E-11 | 3,96E-11 |
| 85,5 | 1,70E-10 | 5,77E-10 | 8,95E-11 | 1,78E-10 | 1,12E-10 | 1,08E-11 | 5,50E-12 |

4. Statistical analysis of simulation results

To establish functional relationship between the angular distribution of brightness and optical parameters, regression analysis was used, which is widely used to restore aerosol and cloud characteristics, and also to assess their effect on climate.⁶ The regression equation for the angular distribution of brightness in the aerosol-molecular and cloud

atmosphere relatively wavelength of the incident radiation was obtained in the form $y = \frac{b_1}{x} + b_0$ for all reception angles.

Regression coefficients for the cloudless atmosphere are given in Table 4. The regression coefficients for the cloudy atmosphere are given in Table 5. The tables also shown coefficients of determination R^2 . This coefficient indicates the proportion of the total variation in the dependent variable y due to variability of x and characterizes the total quality of regression. The statistical significance of determination coefficient can be confirmed with the help of Fisher

statistics: $F = \frac{R^2}{1-R^2} \cdot (n-2)$. Here n is the number of observations. The value F is compared to value $F_{\alpha; k_1, k_2}$ from Fisher table. Here α - the given level of significance, $k_1 = 1$ and $k_2 = n - 2$.

Table 4. Coefficients of the regression equation for the cloudless atmosphere

| Angles of reception, grad | Coefficient b_0 | Coefficient b_1 | Coefficient of determination, R^2 |
|---------------------------|-------------------|-------------------|-------------------------------------|
| 4,5 | 1,19E-05 | 3,25E-05 | 0,993 |
| 13,5 | 2,31E-07 | 5,05E-07 | 0,976 |
| 22,5 | 6,98E-08 | 1,19E-07 | 0,927 |
| 31,5 | 3,92E-08 | 4,04E-08 | 0,784 |
| 40,5 | 2,44E-08 | 1,51E-08 | 0,570 |
| 49,5 | 1,39E-08 | 6,51E-09* | 0,482 |
| 58,5 | 6,82E-09 | 3,64E-09 | 0,609 |
| 67,5 | 2,74E-09 | 2,35E-09 | 0,771 |
| 76,5 | 5,21E-10 | 1,23E-09 | 0,907 |
| 85,5 | 3,00E-11 | 1,76E-10 | 0,929 |

Table 5 Coefficients of the regression equation for the cloud atmosphere

| Angles of reception, grad | Coefficient b_0 | Coefficient b_1 | Coefficient of determination, R^2 |
|---------------------------|-------------------|-------------------|-------------------------------------|
| 4,5 | 3,52E-06 | 7,12E-06 | 0,985 |
| 13,5 | 1,08E-07 | 3,63E-07 | 0,997 |
| 22,5 | 3,63E-08 | 7,53E-08 | 0,973 |
| 31,5 | 1,26E-08 | 2,66E-08 | 0,976 |
| 58,5 | 3,68E-09* | 4,34E-09 | 0,598 |
| 76,5 | -3,34E-08* | 5,04E-09 | 0,606 |
| 85,5 | -2,37E-09* | 4,50E-09 | 0,809 |

We would like to present also the brightness dependence relatively the upper boundary of cloud. The regression equation in this case was obtained in a simple form: $y = b_1x + b_0$ for all reception angles. Regression coefficients and determination coefficients for the case of isotropic source, wavelength of $\lambda = 0,347$ mkm, cloud of the "Haze H" model are given in Table 6.

Table 6. Coefficients of the regression equation for the cloud atmosphere. Model «Haze H»

| Angles of reception, grad | Coefficient b_0 | Coefficient b_1 | Coefficient of determination, R^2 |
|---------------------------|-------------------|-------------------|-------------------------------------|
| 4,5 | 5,14E-07 | -3,07E-08 | 0,990 |
| 13,5 | 1,53E-07 | -7,70E-09 | 0,975 |
| 22,5 | 5,92E-08 | -2,33E-09 | 0,851 |
| 31,5 | 2,65E-08 | -6,54E-10* | 0,654 |
| 40,5 | 7,36E-09 | -1,75E-10 | 0,946 |
| 49,5 | 3,60E-09 | -6,36E-11 | 0,625 |
| 58,5 | 2,11E-09 | -5,13E-11 | 0,840 |
| 67,5 | 1,25E-09 | -2,96E-11 | 0,886 |

Note. * - insignificant coefficient.

Conclusions

Statistical estimation of regression equations for significance was carried out on the basis of the F -test and estimation of the determination coefficient. With 90% confidence, it can be argued that the considered dependence is statistically significant.

The analysis shows that between obtained angular distributions of intensity and wavelength λ in transparent windows for both cloudless and cloudy for the atmosphere, there is a link that can be with a good degree of accuracy to describe hyperbolic regression equation.

Between the brightness and the upper boundary of cloud, there is a link that can be with a degree of accuracy to describe linear regression equation.

References

- [1] Gendrina I.Yu., Kvach A.S., “The Monte Carlo method for determining the vision system characteristics”, J. of International Scientific Publication: Education Alternatives, 11(1), 236 – 244 (2013).
- [2] Zuev V.E., Belov V.V., Veretennikov V.V., [Systems with applications in scattering media], Publishing House “Spectrum” Institute of Atmospheric Optics of the Siberian Branch of the Russian Academy of Sciences, 157-168 (1997).
- [3] Marchuk G.I., Mikhailiov G.A., Nazaraliev M.A., Darbinjan, Kargin B.A., Elepov B.S., [The Monte Carlo method in atmospheric optics], Nauka, Novosibirsk, 4-95 (1976).
- [4] Deirmendjian D., [Electromagnetic scattering on spherical polydispersions], American Elsevier Pub. Co., 1-290 (1969).
- [5] Gendrina I.Yu., Alekseenko M. A., “The regression analysis of statistical simulation results in vision systems through the atmosphere”, Izvestija-vuzov-fizika, 58 (8/2), 294-296 (2015).
- [6] Khayer M. M. et al., “Evaluation of a 5-Year Cloud and Radiative Property Dataset Derived from GOES-8 Data Over the Southern Great Plains”, Twelfth ARM Science Team Meeting Proceedings, 1-14 (2002).