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3 1 **CONTRASTING EFFECTS OF REINDEER GRAZING ON**  
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6 2 **CO<sub>2</sub>, CH<sub>4</sub> AND N<sub>2</sub>O FLUXES ORGINATING FROM THE**  
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9 3 **NORTHERN BOREAL FOREST FLOOR.**  
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25 24 Running Title: REINDEER GRAZING AND GREENHOUSE GAS FLUXES FROM FIELD  
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27 25 LAYER.  
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29 **Abstract**

30 Reindeer (*Rangifer tarandus* L.) are considered to be an important mammalian  
31 herbivore, strongly influencing Arctic lichen-dominated ecosystems. There is no wide  
32 knowledge about the effect of reindeer on greenhouse gas (GHG) fluxes in northern  
33 boreal forests. Ground vegetation plays an important role in absorbing nitrogen (N)  
34 and carbon dioxide (CO<sub>2</sub>) from the atmosphere. Lately, it has also been found to be a  
35 significant source of nitrous oxide (N<sub>2</sub>O) and a small source of methane (CH<sub>4</sub>). We  
36 investigated the influence of reindeer grazing on field layer GHG (CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O)  
37 fluxes, ground vegetation coverage and biomass, soil physical properties  
38 (temperature and moisture) in a northern boreal forests. At our study site, the  
39 reindeer-induced replacement of lichen by mosses had contrasting effects on the  
40 GHG fluxes originating from the field layer. Field layer CO<sub>2</sub> efflux was significantly  
41 higher in grazed areas. The field layer was a CH<sub>4</sub> sink in all areas, but grazed areas  
42 absorbed more CH<sub>4</sub> compared to non-grazed areas. While total N<sub>2</sub>O fluxes remained  
43 around zero in grazed areas, a small N<sub>2</sub>O sink occurred in non-grazed areas with  
44 lower moss biomass. Our results indicated that grazing by reindeer in northern boreal  
45 forests affects GHG fluxes from the forest field layer both positively and negatively,

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3 46 and these emissions largely depend on grazing-induced changes in vegetation  
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7 47 composition.  
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12  
13 49 **Keywords:** Reindeer grazing, boreal forests, greenhouse gases, mosses, lichens  
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For Peer Review

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3 51 **Introduction**  
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5 52 Over 35% (~415 000 km<sup>2</sup>) of the land area in Scandinavia is grazed by reindeer  
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9 53 (*Rangifer tarandus* L.) (Jernsletten & Klokov, 2002). They influence most of the  
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12 54 Scandinavian Arctic, subarctic and northern boreal regions, which are also expected  
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16 55 to be significantly influenced by the climate change (Anisimov *et al.*, 2007). Reindeer  
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18  
19 56 can strongly influence plant diversity and species composition, and consequently  
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21  
22 57 also the nutrient availability and carbon (C) balance of subarctic ecosystems (Väre *et*  
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26 58 *al.*, 1996; Köster *et al.*, 2013; Akujärvi *et al.*, 2014). The number of reindeer has  
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29 59 declined in northern boreal areas, except for Scandinavia where it has remained  
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32 60 stable or has even increased over the last decades (Vors & Boyce, 2009).  
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62 Reindeer grazing affects ground vegetation by trampling (Väre *et al.*, 1996;  
63 Suominen & Olofsson, 2000), selective grazing (Olofsson & Oksanen, 2002) and by  
64 returning nutrients in dung and urine (Barthelemy *et al.*, 2015). Van der Wal (2006)  
65 showed that in systems with low productivity, grazing changes the ground vegetation  
66 from lichen-rich to moss-rich. In more nutrient-rich systems, this change may be from  
67 moss-rich to graminoid-rich (Van der Wal, 2006). Heavy grazing increases the

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3 68 presence of bare patches of soil (Suominen & Olofsson, 2000), and grazing may  
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7 69 therefore warm up and dry the soil (Väre *et al.*, 1996). Temperature and soil moisture  
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10 70 are one of the key factors influencing activity of microbes, soil decomposition  
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13 71 processes and the greenhouse gas (GHG) exchange between the field layer and  
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17 72 atmosphere (Karhu *et al.*, 2014).

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21 73 Northern and high altitude soils have large soil organic matter (SOM) stores due to  
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24 74 the slowly decomposing recalcitrant plant litter (Hobbie *et al.*, 2002). Soils can be  
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26  
27 75 both source or/and sink of three GHGs: carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>) and  
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31 76 nitrous oxide (N<sub>2</sub>O) (Smith *et al.*, 2003). Effluxes of soil CO<sub>2</sub> can be originated from  
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34 77 heterotrophic respiration (decomposition of SOM), and autotrophic respiration of  
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37 78 roots and mycorrhizal fungi. Both of these are affected by the soil organic C  
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41 79 composition, concentration and soil physical properties (Kirschbaum 2000). The  
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44 80 CH<sub>4</sub> flux is a net effect of CH<sub>4</sub> methanogenesis and CH<sub>4</sub> oxidation (Le Mer & Roger,  
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47  
48 81 2001). N<sub>2</sub>O is a product or by-product in a series of microbial processes such as  
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50  
51 82 denitrification and nitrification (Wrage *et al.*, 2001). As both methanogenesis and  
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54 83 denitrification are anaerobic processes, CH<sub>4</sub> and N<sub>2</sub>O fluxes are sensitive to soil  
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58 84 moisture content changes.

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7 86 In addition, soil GHG fluxes are affected by cryptogams (lichens, bryophytes, etc.),  
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10 87 which can fix CO<sub>2</sub> from the atmosphere through photosynthesis and also fix nitrogen  
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12  
13 88 (N) in symbiosis with N-fixing bacteria (Elbert *et al.*, 2012). The role of cryptogams as  
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17 89 sources of N<sub>2</sub>O and CH<sub>4</sub> has also been demonstrated recently (Lenhart *et al.*, 2015).

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20  
21 90 By affecting ground vegetation composition and soil temperatures reindeer grazing  
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24 91 could significantly impact the N<sub>2</sub>O and CH<sub>4</sub> fluxes from the forest floor (Porada *et al.*,  
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27 92 2017).  
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35 94 Over half of the subarctic area is covered by forest, but the number of studies  
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38 95 addressing the interactions between grazing and field layer GHG fluxes in northern  
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42 96 boreal forests is low (Olofsson *et al.*, 2010; Köster *et al.*, 2013; Köster *et al.*, 2015b).

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45 97 Previous studies on reindeer grazing have mostly focused on tundra areas and  
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49 98 grazing effect on CO<sub>2</sub> efflux, while the effects on other GHGs (CH<sub>4</sub> and N<sub>2</sub>O) have  
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52 99 not been studied. This study assesses the influence of reindeer on the fluxes of these  
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55 100 three GHGs in the field layer of a northern boreal forests. We compared grazed and  
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3 101 non-grazed areas and linked the GHG fluxes (CO<sub>2</sub> efflux from the field layer, the  
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6 102 proportion of root respiration amongst the CO<sub>2</sub> production, the CH<sub>4</sub> and N<sub>2</sub>O  
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10 103 exchange between the field layer and atmosphere) to vegetation biomass, plant  
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13 104 functional type and characteristics of soil (root biomass, soil temperature and  
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17 105 moisture).  
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25 107 We hypothesized that by reducing lichen biomass, reindeer grazing will increase soil  
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28 108 temperature and decrease soil moisture. We also hypothesized that these and the  
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31 109 following soil microbiological changes in the soil affect field layer CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O  
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34 110 fluxes: a) higher soil temperatures result in higher CO<sub>2</sub> emissions from the soil; b)  
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37 111 decreased soil moisture negatively affects CH<sub>4</sub> and N<sub>2</sub>O fluxes; c) differences in  
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41 112 ground vegetation characteristics between grazed and non-grazed areas influence  
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44 113 soil CH<sub>4</sub> and N<sub>2</sub>O fluxes, as mosses and lichens can both fix and emit these GHGs.  
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## 53 **Materials and methods**

### 54 55 56 57 116 *Study sites*

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3 117 Our study areas situate in the Värriö Strict Nature Reserve (67°46' N, 29°35' E) of  
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6 118 Finnish Lapland. The areas are situated close to the treeline (average altitude of the  
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10 119 study area was 300 m a.s.l.), in a northern boreal forest, in the zone of the last intact  
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13 120 subarctic Scots pine (*Pinus sylvestris* L.) forest landscapes in Fennoscandia  
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17 121 (Potapov *et al.*, 2008). The area is characterized by the sub-continental climate.  
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20 122 Snow cover persists for 200–225 days/year, and the average snow depth is  
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23 123 approximately 0.7 m (Köster *et al.*, 2015b). The annual growing season varies  
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27 124 between 105–120 days (Köster *et al.*, 2014; Köster *et al.*, 2015b). The soils belong to  
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30 125 haplic podzol (FAO, 1990) with no permafrost, and the pH of the soil is 4.4 (Köster *et*  
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33 126 *al.*, 2014). The study areas belong to the Pohjois-Salla reindeer herding district, with  
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37 127 approximately 2.2 reindeer per km<sup>2</sup> (Turunen *et al.*, 2016).  
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#### 45 *Experimental design*

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48 130 The experiment was established in a year 2013 as a split plot experiment with five  
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52 131 plots. Each plot was separated to the grazed and non-grazed subplot (400 m<sup>2</sup>). The  
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56 132 areas are separated by the fence running between Finland and Russia, preventing  
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3 133 Finnish domesticated reindeer to cross the border (Köster *et al.*, 2017). The official  
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7 134 border between Finland and Russia is a few hundred metres further east from the  
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10 135 reindeer fence. The subplots (n=10) were situated along this fence. The Finnish side  
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13 136 is mainly grazed during the wintertime, with an average density of 2.2 reindeer per  
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17 137 km<sup>2</sup>, while the non-grazed areas on the Russian side of the fence have been  
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20 138 excluded from heavy reindeer grazing for nearly a century (Stark *et al.*, 2003).

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28 140 *Biomass measurements and vegetation coverage*

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32 141 Stand characteristics and ground vegetation and root biomass measurements were  
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35 142 conducted in 2013, and the description of the methods used is published in Köster *et*  
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38 143 *al.* (2015b) and in Supporting experimental procedures: Methods S1. The ground  
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42 144 vegetation biomass measurements were repeated according to the same protocols in  
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45 145 2014 at the same time as GHG measurements. Four 0.5m x 0.5m quadrats placed  
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49 146 systematically inside each subplot were photographed and the ground vegetation  
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52 147 species composition, species coverage and possible area of bare ground were  
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55 148 visually estimated from these photos. As shown in Köster *et al.* (2015b), total ground

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3 149 vegetation and lichen biomasses were higher and moss biomasses lower in the  
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7 150 areas with no reindeer, while the dwarf shrub and grass biomasses were similar for  
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10 151 both areas (Table 1). The differences in ground vegetation and tree regeneration  
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13 152 caused differences in root biomasses between the non-grazed and grazed areas.  
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17 153 Both ground vegetation root biomasses and tree root biomasses were higher in non-  
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20 154 grazed areas (Table 1) (Köster *et al.*, 2015b).  
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28 156 *Chamber measurements of CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O.*  
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32 157 Soil GHG fluxes were measured using the static chamber method with 50 polyvinyl  
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35 158 chloride (PVC) collars (five collars per subplot; h = 0.05 m, Ø = 0.22 m) installed  
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38 159 permanently in the soil during the spring of 2013 one year before the flux  
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42 160 measurements. To avoid the leakage of air to the collar we installed collars in the  
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45 161 litter layer and sealed them with sand. The ground vegetation was left intact inside  
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49 162 the chamber in course of measurements. GHG measurements were performed  
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52 163 monthly (second week of each month) in June, July and August in 2014, with a  
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55 164 portable opaque chamber (h = 0.24 m, Ø = 0.22 m) (Pumpanen *et al.*, 2015).  
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3 165 CO<sub>2</sub> concentration in the chamber headspace was registered for a 5-minute chamber  
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6 166 deployment time at 5-sec intervals with a CO<sub>2</sub> probe (GMP343, Vaisala Oyj, Vantaa,  
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10 167 Finland); temperature and air humidity were recorded with a temperature and relative  
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13 168 humidity sensor (HM70, Vaisala Oyj, Vantaa, Finland) (Köster *et al.*, 2015a).

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16 169 The CH<sub>4</sub> and N<sub>2</sub>O flux measurements were performed using the same collars and the  
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20 170 same chamber as in the CO<sub>2</sub> efflux measurements by collecting air samples from the  
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23 171 chamber with a polypropylene syringe (50-ml) (BD Plastipak 60, BOC Ohmeda,  
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26 172 Helsingborg, Sweden). The samples were collected 0, 1, 5, 10 and 15 min after the  
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30 173 chamber was installed, and immediately stored into glass vials (Soda glass (12 ml)  
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33 174 Labco Exetainer®, Labco Limited, UK). We analysed the samples by an Agilent Gas  
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36 175 Chromatograph model 7890A (GC, Agilent Technologies, USA). The linear  
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40 176 regression between the measurement time and GHG concentration change within  
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43 177 the chamber was used to calculate the fluxes of CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O. Outliers from the  
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45  
46 178 CH<sub>4</sub> and N<sub>2</sub>O concentrations measured during each chamber deployment were  
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50 179 filtered using Grubbs' test for outliers (Stefansky, 1972).

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53 180 During the GHG measurements, soil moisture was determined with a ML2 sensor  
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56 181 (Delta-T Devices, UK) and soil temperature (top 10 cm) with a P 300w temperature  
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3 182 probe (Dostmann Electronic GmbH, Germany). In addition, soil temperature was  
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7 183 measured continuously at every subplot with iButton temperature sensors (Maxim  
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10 184 Integrated, San Jose, California, U.S.A.) and soil moisture at two subplots with an  
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13 185 ML2x soil moisture sensor (Delta-T Devices Ltd., Cambridge, England) installed  
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17 186 under the organic layer and connected to a data logger.  
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22 188 *The proportion of root respiration*  
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26 189 Five soil cores ( $h = 0.15$  m,  $\varnothing = 0.05$  m) were collected from each subplot for the  
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29 190 incubation study and stored at  $+4$  °C for two weeks. Soil cores were incubated in  
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33 191 cylinders ( $h = 0.3$  m,  $\varnothing = 0.075$ ), and  $\text{CO}_2$  exchange was followed with  $\text{CO}_2$  and  
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36 192 water vapour analysers (Licor LI-840 and Licor LI-7000, Li-Cor Inc., Lincoln,  
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39 193 Nebraska). The system temperature was constantly  $+12$  °C, and differences of  $\text{CO}_2$   
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43 194 concentrations were measured from the incoming and outgoing air. The proportion of  
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45  
46 195 root respiration was calculated from one time point (the beginning of June) by  
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48  
49 196 comparing the field layer  $\text{CO}_2$  efflux measurements (corrected to the rates occurring  
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53 197 at  $+12$  °C) in the field to the lab incubation measurements. Field layer  $\text{CO}_2$  efflux was  
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3 198 used as total respiration (heterotrophic + autotrophic respiration). Lab incubation  
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7 199 measurements provided us with the heterotrophic respiration values.  
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14 201 *Data analysis*  
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18 202 Normality check for the data was done with the Shapiro–Wilk test ( $P > 0.05$ ) and  
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21 203 logarithm transformations were calculated for the measured CO<sub>2</sub> fluxes. The other  
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24  
25 204 parameters did not need transformation. We used the linear mixed effects model  
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27  
28 205 (PROC MIXED) for testing the effects of various factors (grazed or non-grazed area,  
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31 206 subplot, collar location, measurement time, root biomass, various ground vegetation  
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35 207 biomasses and coverage, etc.) on soil GHG emissions as follows:  
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38 208 
$$Y_{p,s} = \beta_0 + \beta_1 A + b_p + b_{C*P} + b_T + \epsilon_S \quad (1)$$
  
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41 209 
$$b_p \sim N(0, \delta_p^2) \quad (2)$$
  
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44 210 
$$b_{C*P} \sim N(0, \delta_{C*P}^2) \quad (3)$$
  
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47 211 
$$b_T \sim N(0, \delta_T^2) \quad (4)$$
  
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$$\epsilon_S \sim N(0, \delta_S^2) \quad (5)$$
  
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3 214 Where  $\beta_0$  is the intercept and  $\beta_1$  is the regression slope.  $b_P$  is the random intercept of  
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5  
6 215 the regression for plots,  $b_{C*P}$  is the random intercept for collars within a plot,  $b_T$  is the  
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10 216 random intercept for measurement times and  $\epsilon_S$  is the sample residual.  $\delta_{C*P}^2$  and  $\delta_T^2$   
11  
12  
13 217 are the variances between the plots ( $p$ ) and collars ( $s$ ) within the plots. We assumed  
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16  
17 218 unstructured covariance between all random factors. The subplots were treated as  
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19  
20 219 random factors and the three GHG measurements conducted in June, July and  
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23 220 August at each collar were used as repeated random factors. We treated grazing  
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26  
27 221 (grazed or non-grazed) as a fixed factor.  
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29  
30 222 A similar approach was used for all three GHGs. The first estimation of the model  
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32  
33 223 was done using all existing co-variates: soil moisture, soil temperature, root biomass,  
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37 224 lichen biomass, moss biomass, shrub biomass, lichen coverage, moss coverage,  
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39  
40 225 area (grazed or non-grazed). We therefore dropped the non-significant terms one by  
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43  
44 226 one and separate model based Akaike values (AIC) were calculated (Table S1).  
45  
46  
47 227 Factor area (grazed or non-grazed) was forced to stay in each tested model. The  
48  
49  
50 228 best model to explain factor effects on different GHGs was selected based on the  
51  
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54 229 lowest AIC value. For quantifying the fit of the linear mixed effects model, we  
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56  
57 230 calculated the coefficient of determination ( $r^2$ ):  
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6 232  $r^2 = 1 - \left(\frac{V(m)}{V(r)}\right)$  (6)  
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13 234 Where  $V(m)$  is the sum of squared residuals from the model and  $V(r)$  is the sum of  
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17 235 squared residuals from the intercept-only model (Nakagawa & Schielzeth 2013).  
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20 236 The Pearson correlation coefficients were calculated between the fluxes and  
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22  
23 237 explanatory effects (same co-variables as in model) to evaluate the direction of  
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26  
27 238 different effects (Table S2).  
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30 239 Tukey's HSD post hoc test was employed to compare the grazing effects on different  
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33 240 months. All statistical analyses compared grazed and non-grazed areas, and the  
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37 241 significance level was  $P = 0.05$ . SAS version 9.3 (SAS Institute Inc., Cary, NC, USA)  
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40 242 was used for statistical analyses.  
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## 48 244 **Results**

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52 245 *Vegetation structure and environmental factors*  
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3 246 Grazing did not create patches of bare soil in our study areas (Table 1). Lichen  
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7 247 coverage was on average 77% higher in the areas where reindeer were not present  
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10 248 ( $P < 0.01$ ), and grazing resulted in 94% loss of lichen cover (Table 1). The effect of  
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13 249 reindeer grazing was the opposite for moss coverage: 66% of the ground was  
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17 250 covered with mosses in the grazed areas, whereas mosses covered only 22% of the  
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20 251 ground in the non-grazed areas ( $P = 0.0042$ ) (Table 1). Similar patterns were  
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23 252 observed when biomasses of various vegetation types (lichens, mosses, dwarf  
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27 253 shrubs, grass) were compared between grazed and non-grazed areas (Table 1)  
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30 254 (Köster *et al.*, 2015b).  
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38 256 Average soil temperatures during the measuring period were +10.2 °C in grazed  
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41 257 areas and +9.4 °C in the non-grazed areas ( $P = 0.2527$ ). Coldest soil temperatures  
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45 258 were measured in June and the warmest in July (Figure 1), and a significant  
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48 259 difference in temperature between the areas was observed only in July ( $P = 0.0218$ ),  
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51 260 the temperatures being 0.9 °C warmer in grazed areas.  
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3 261 Soil was driest in July and wettest in June (Figure 1) in both areas. In August, soil  
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6 262 water content in the grazed area was lower compared to the non-grazed area ( $P <$   
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10 263 0.0001). A negative correlation was detected between soil water content and soil  
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13 264 temperature (Table S2:  $R = -0.835$ ,  $P < 0.0001$ ).  
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20 266 *Field layer CO<sub>2</sub> efflux*  
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24 267 The highest field layer CO<sub>2</sub> effluxes were measured in July and August and the  
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27 268 lowest in June (Figure 2), and there was a positive correlation between soil  
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30 269 temperature and field layer CO<sub>2</sub> emission ( $R = 0.44$ ,  $P < 0.0001$ ). In general, CO<sub>2</sub>  
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33 270 effluxes were higher in grazed areas, but the differences were statistically significant  
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37 271 only in June and August (Figure 2).  
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40 272 Based on the linear mixed model effect, soil temperature ( $P < 0.0001$ ) and lichen  
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43 273 coverage ( $P = 0.0411$ ) were the most important factors affecting the CO<sub>2</sub> efflux  
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47 274 (Table 2). The effect of grazing on the CO<sub>2</sub> efflux from the field layer was not  
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50 275 statistically significant ( $P = 0.1566$ ) (Table 2). The CO<sub>2</sub> efflux increased with warmer  
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53 276 soil temperatures (Table S2:  $R = 0.665$ ,  $P < 0.0001$ ) and decreased with higher  
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3 277 lichen coverage (Table S2:  $R = -0.223$ ,  $P = 0.0565$ ). Approximately 50–79% of the  
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7 278 total field layer CO<sub>2</sub> efflux originated from root and rhizosphere respiration, but no  
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10 279 significant difference was observed between grazed and non-grazed areas (Table 1).

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17 281 *Soil CH<sub>4</sub> flux*

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21 282 Soils in our study areas were CH<sub>4</sub> sinks throughout the entire measuring period, and  
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24 283 the sink was significantly higher ( $P = 0.0108$ ) in grazed areas (Figure 3). The grazed  
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28 284 areas were larger sinks for CH<sub>4</sub> in June and August ( $P < 0.0001$ ), but there was no  
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31 285 difference between the areas detected in July (Figure 3). Based on the linear mixed  
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34 286 effect model, soil temperature ( $P = 0.0009$ ) and moss coverage ( $P = 0.0042$ ) were  
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38 287 the most important factors affecting the CH<sub>4</sub> flux (Table 2). Higher moss coverage in  
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41 288 grazed areas resulted in a higher CH<sub>4</sub> sink (Table S2:  $R = -0.406$ ,  $P = 0.0003$ ), but at  
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44 289 the same time warmer soils decreased the field layer CH<sub>4</sub> sink (Table S2:  $R = 0.330$ ,  
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48 290  $P = 0.0041$ ). The effects of grazing and soil moisture on the CH<sub>4</sub> flux were not  
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51 291 statistically significant (Table 2).

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293 *Soil N<sub>2</sub>O flux*

294 A difference ( $P = 0.0116$ ) in N<sub>2</sub>O fluxes was observed between the areas grazed and  
295 non-grazed. The non-grazed areas acted as N<sub>2</sub>O sinks throughout the entire  
296 measuring period, while the trend was not as clear in the grazed areas (Figure 4). In  
297 June and August, we observed no significant differences between the treatments,  
298 whereas in July the non-grazed areas were larger sinks for N<sub>2</sub>O. Moss biomass was  
299 the only factor affecting the N<sub>2</sub>O fluxes ( $P = 0.0494$ ) (Table 2), while N<sub>2</sub>O  
300 consumption (sink) was larger with lower moss biomass (Table S2:  $R = 0.528$ ,  $P <$   
301  $0.0001$ ).

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303 **Discussion**

304 We studied the impact of reindeer grazing on field layer GHG (CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O)  
305 fluxes in a northern boreal coniferous forest, and linked these fluxes to grazer-  
306 induced changes in ground vegetation coverage and biomass along with soil  
307 temperature and soil moisture. Areas grazed by reindeer consumed more CH<sub>4</sub> in  
308 June and August, and there were significantly higher CO<sub>2</sub> effluxes from these areas

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3 309 at the same time. Non-grazed areas were a sink for N<sub>2</sub>O. Our study indicates that  
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7 310 grazing caused changes in soil temperature and lichen coverage affect field layer  
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10 311 CO<sub>2</sub> emissions. Concurrently, grazing-induced changes in moss coverage, soil  
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13 312 moisture content and temperature affected the CH<sub>4</sub> fluxes, and changes in moss  
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17 313 biomass affected the N<sub>2</sub>O fluxes.  
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23 315 We observed that field layer CO<sub>2</sub> effluxes were generally higher in areas grazed by  
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27 316 reindeer. As hypothesized, the linear mixed effect model showed that the important  
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30 317 factor explaining the CO<sub>2</sub> efflux was the soil temperature (Table 2). Together with soil  
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33 318 temperature lichen coverage also significantly influenced the CO<sub>2</sub> efflux as there was  
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37 319 a negative correlation between soil CO<sub>2</sub> efflux and lichen coverage (Table 2). Our  
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40 320 study detected significant grazing effects on the CO<sub>2</sub> efflux only in June and August  
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43 321 (Figure 2), and the reason for that could be that the CO<sub>2</sub> measurements in July were  
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47 322 performed directly after a few rainy days, and the low temperature and high soil  
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50 323 moisture possibly influencing the CO<sub>2</sub> efflux (Figure 1).  
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53 324 Previous studies from boreal pine forests have showed grazing to have no effect  
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57 325 (Köster *et al.*, 2015b) or only a minor effect (Köster *et al.*, 2017) on the field layer CO<sub>2</sub>  
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3 326 efflux. Stark *et al.* (2003) documented lower soil respiration and microbial activity in  
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7 327 the grazed areas in the northern boreal forest soils. These outcomes were explained  
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10 328 by decreased activity of microbes and Scots pine root biomass. As our study deals  
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13 329 with field layer respiration that encompasses both soil respiration and the respiration  
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17 330 by vegetation, also the changed species composition and the altered environmental  
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20 331 conditions induced by it may explain the difference.  
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23 332 The results of our study revealed also that grazed areas were stronger CH<sub>4</sub> sinks (i.e.  
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27 333 increased CH<sub>4</sub> uptake) in June and August. The results of the mixed effect model  
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30 334 analysis indicated that soil temperature and moss coverage together with soil  
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33 335 moisture explained most of the CH<sub>4</sub> fluxes (Table 2).  
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37 336 Based on the model, colder and drier soils acted as stronger CH<sub>4</sub> sinks. Another  
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41 337 explanation for higher CH<sub>4</sub> uptake in the grazed areas was the grazing induced  
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44 338 differences in ground vegetation, as the field layer CH<sub>4</sub> sink correlated positively with  
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48 339 the moss coverage. Lichens and bryophytes contain methane-producing  
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51 340 cyanobacteria (Lenhart *et al.*, 2015). Higher moss coverage (and smaller lichen  
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54 341 coverage) is likely to have increased the capacity of the ground layer to act as a CH<sub>4</sub>  
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3 342 sink. Our results therefore suggest, as hypothesized, that changes in the CH<sub>4</sub> flux  
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7 343 depend on the changes in ground layer vegetation.  
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14 345 The forest field layer was a clear N<sub>2</sub>O sink in non-grazed areas, but the trends were  
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18 346 not as clear in grazed areas. Our hypothesis that differences in ground vegetation  
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21 347 characteristics might influence the N<sub>2</sub>O fluxes was confirmed by our model analysis,  
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24 348 which showed moss biomass to be the only factor which affected the N<sub>2</sub>O fluxes in  
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28 349 our study (Table 2). Lichens and bryophytes are one of the main N<sub>2</sub>O sources in  
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31 350 Arctic regions (Zhuang *et al.*, 2012), where the N<sub>2</sub>O release also depends on  
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34 351 temperature (Lenhart *et al.*, 2015). Non-grazed areas have less mosses, which  
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38 352 probably emit less N<sub>2</sub>O. As the soil acts as an N<sub>2</sub>O sink regardless of grazing, the  
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41 353 higher moss biomass on the grazed area likely negated this and turned the field layer  
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44 354 to be smaller N<sub>2</sub>O sink. However, other reasons may exist behind the difference  
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48 355 between grazed and non-grazed areas, e.g. temperature and soil moisture content.  
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51 356 Supporting the observations by Olofsson *et al.* (2010), soil temperatures in our study  
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54 357 were slightly higher in grazed areas (although no statistically significant difference  
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3 358 was observed in our case). Soil warming goes along with changes in species  
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7 359 composition in our study, where lichens were replaced with mosses, and these  
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10 360 darker surfaces absorb more solar radiation.  
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17 362 Our previous study (Köster *et al.*, 2015b) documented significant effect of grazing on  
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20 363 the biomass of lichens and mosses, but no effect was observed on dwarf shrub and  
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23 364 grass biomasses. Lichens are the main winter forage for reindeer, and similarly to  
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27 365 other studies (Akujärvi *et al.*, 2014; Susiluoto *et al.*, 2008; Olofsson *et al.*, 2010;  
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30 366 Köster *et al.*, 2013), reindeer grazing reduced the lichen biomass by more than 90%  
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34 367 in our study areas. One additional important finding in our study was that grazing by  
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37 368 reindeer did not create patches of bare soil, which was shown to be the case in  
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40 369 studies conducted in tundra areas (Suominen and Olafson, 2000; Väre *et al.*, 1996).  
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44 370 In our study areas, mosses were occupying the areas grazed by reindeer and bare  
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47 371 patches of soil were uncommon. Also, previous studies (Köster *et al.*, 2013; Väre *et*  
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50 372 *al.*, 1996) have showed mosses to quickly invade grazed areas, probably due to  
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54 373 decreased competition by lichens. Thus, through the promotion of mosses and the  
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3 374 reduction in tree regeneration (Köster *et al.*, 2013), grazing changes the vegetation  
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7 375 composition and its functions.  
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13 377 Grazing by reindeer increases nutrient cycling (Olofsson & Oksanen, 2002) by  
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16 378 changing the litter quality and abiotic conditions (temperature and moisture) relevant  
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20 379 for litter decomposition (Olofsson & Oksanen, 2002; Stark & Väisänen, 2014). It has  
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23 380 been found that reindeer grazing in areas grazed during summer has a strong  
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26 381 positive effect on ecosystem nutrient cycling, as dung and urine deposits provides  
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30 382 resources that are highly decomposable (Olofsson *et al.*, 2004; Barthelemy *et al.*,  
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33 383 2015), and have positive effect on net N mineralization (Stark *et al.*, 2000) and GHG  
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36 384 fluxes (Turunen *et al.* 2013). Still, our study areas are located in the winter ranges,  
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40 385 where dung and urine depositions are partly wasted away with snowmelt, and  
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43 386 thereby, they do not affect microbial activity in a similar manner as in the summer  
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46 387 grazing areas.  
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50 388 Grazing may also affect belowground C dynamics through changes in ground  
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53 389 vegetation. A substantial proportion of assimilated C is allocated to both tree and  
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57 390 belowground plant root biomass and/or root-associated mycorrhizal fungi, which are  
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3 391 an important contributors to soil respiration. Moreover, the recently assimilated C  
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7 392 may also increase SOM decomposition by providing energy for decomposing  
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10 393 microorganisms and possibly affect GHGs (Högberg et al., 2002). Ground vegetation  
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13 394 root biomasses were slightly higher in non-grazed areas, but the CO<sub>2</sub> efflux was  
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17 395 concurrently higher in the grazed areas, thus changes in root and rhizosphere  
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20 396 respiration of ground vegetation did not explain the higher CO<sub>2</sub> efflux in the grazed  
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23 397 areas. The amount of litter in the grazed area could also be expected to be smaller  
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27 398 when ground vegetation biomass was reduced due to grazing, and this would result  
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30 399 in smaller decomposition rates. However, the main litter input in forested areas  
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33 400 comes from trees, but as shown by Köster *et al.* (2015b), no differences were  
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37 401 observed in the aboveground biomasses between the treatments.

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43 403 Our study showed that grazing significantly reduced lichen coverage and biomass.

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46 404 The biomass and coverage of mosses concurrently increased significantly in grazed

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49 405 areas. Our results indicated that grazing-induced changes in vegetation composition

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52 406 both positively and negatively affect GHG fluxes from the forest field layer in a

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55 407 northern boreal forests. Our results showed clearly higher CO<sub>2</sub> emissions from

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3 408 grazed areas, and these were affected by soil temperature and lichen coverage. Our  
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7 409 results also clearly demonstrated that CH<sub>4</sub> and N<sub>2</sub>O fluxes were dependent on shifts  
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10 410 in aboveground vegetation communities. Soil temperature and moss coverage  
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13 411 together with soil moisture explained most of the CH<sub>4</sub> fluxes, and grazed areas  
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16 412 consumed more CH<sub>4</sub>. Non-grazed areas with lower moss biomass were concurrently  
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20 413 a small N<sub>2</sub>O sink. The results presented in our study are collected from a single  
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23 414 experiment with similar site characteristics (same soil, tree species, grazing  
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26 415 practices, etc.), and thus it would be difficult to extrapolate our results to entire  
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30 416 forested areas in Arctic regions. Overall, our results show for the first time that  
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33 417 besides influencing the CO<sub>2</sub> fluxes from the northern boreal forest floor, reindeer  
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36 418 grazing has an impact on the fluxes of other GHGs. Thus, reindeer may have a  
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40 419 significant role in determining the GHG balance of northern boreal forests.  
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17 428 Supporting Information:  
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21 429 Supporting experimental procedures: Methods S1  
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25 430 Supporting tables: Table S1, Table S2  
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561 Table 1. Average ground vegetation biomasses ( $\text{kg m}^{-2}$ ), root biomasses ( $\text{kg m}^{-2}$ ),  
 562 number of trees per hectare (referred from Köster et al., 2015b). Average vegetation  
 563 coverages (%), patches of bare soil (%) and values for root respiration (%) in grazed  
 564 and non-grazed areas. The given values are the mean and SE in parentheses.

	Grazed	Non-grazed
Total ground vegetation biomass ( $\text{kg m}^{-2}$ )	0.61 ( $\pm 0.09$ )	0.94 ( $\pm 0.07$ )*
Ground vegetation root biomass ( $\text{kg m}^{-2}$ )	0.52 ( $\pm 0.02$ )	0.66 ( $\pm 0.02$ )*
Tree root biomass ( $\text{kg m}^{-2}$ )	2.2 ( $\pm 0.12$ )	2.5 ( $\pm 0.2$ )
Nr of trees per ha ( $h < 1.3$ )	3892 ( $\pm 839$ )	7133 ( $\pm 233$ )*
Nr of trees per ha ( $h > 1.3$ )	583 ( $\pm 79$ )	558 ( $\pm 33$ )
Lichen biomass ( $\text{kg m}^{-2}$ )	0.04 ( $\pm 0.02$ )	0.45 ( $\pm 0.09$ )*
Moss biomass ( $\text{kg m}^{-2}$ )	0.36 ( $\pm 0.07$ )	0.28 ( $\pm 0.04$ )*
Dwarf shrub biomass ( $\text{kg m}^{-2}$ )	0.21 ( $\pm 0.02$ )	0.21 ( $\pm 0.03$ )
Grass biomass ( $\text{kg m}^{-2}$ )	0.01 ( $\pm 0.01$ )	0.01 ( $\pm 0.01$ )
Lichen coverage (%)	4.2 ( $\pm 0.9$ )	76.7 ( $\pm 8.6$ )*
Moss coverage (%)	65.9 ( $\pm 12.3$ )	22.2 ( $\pm 8.8$ )*
Dwarf shrub coverage (%)	49.9 ( $\pm 7.9$ )	38.3 ( $\pm 6.4$ )
Patches of bare soil (%)	< 1	< 1
Grass coverage (%)	< 1	< 1
Root respiration (%)	65.6 ( $\pm 4.3$ )	61.1 ( $\pm 4.9$ )

565 \* $P \leq 0.01$

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567 Table 2. Parameter values of the linear mixed effect models that best explained the  
 568 GHG fluxes. The tested models and their selection based on AIC criteria are shown  
 569 in Table S1.

	CO <sub>2</sub>	CH <sub>4</sub>	N <sub>2</sub> O
	Model 7	Model 6	Model 10
r <sup>2</sup>	0.6517	0.245	0.249
AIC	70.3	-1686.3	-1827.5
Intercept	-2.248	-0.000015	-4.4 x 10 <sup>-7</sup>
(stderr, P)	(0.245, P = 0.0116)	(4.4 x 10 <sup>-6</sup> , P = 0.0729)	(7.9 x 10 <sup>-8</sup> , P < 0.0001)
Soil temperature	0.167	7.8 x 10 <sup>-7</sup>	
(stderr, P)	(0.016, P < 0.0001)	(2.2 x 10 <sup>-7</sup> , P = 0.009)	
Lichen coverage	-0.005		
(stderr, P)	(0.002, P = 0.0167)		
Soil moisture		1.3 x 10 <sup>-7</sup>	
(stderr, P)		(8.2 x 10 <sup>-8</sup> , P = 0.0945)	
Moss coverage		-2.8 x 10 <sup>-8</sup>	
(stderr, P)		(9.7 x 10 <sup>-9</sup> , P = 0.0042)	
Moss biomass			1.3 x 10 <sup>-6</sup>
(stderr, P)			(2.5 x 10 <sup>-7</sup> , P < 0.0001)
Area (grazed or non-grazed)	0.259	-7.1 x 10 <sup>-7</sup>	
(stderr, P)	(0.181, P = 0.1566)	(7.7 x 10 <sup>-7</sup> , P = 0.3480)	

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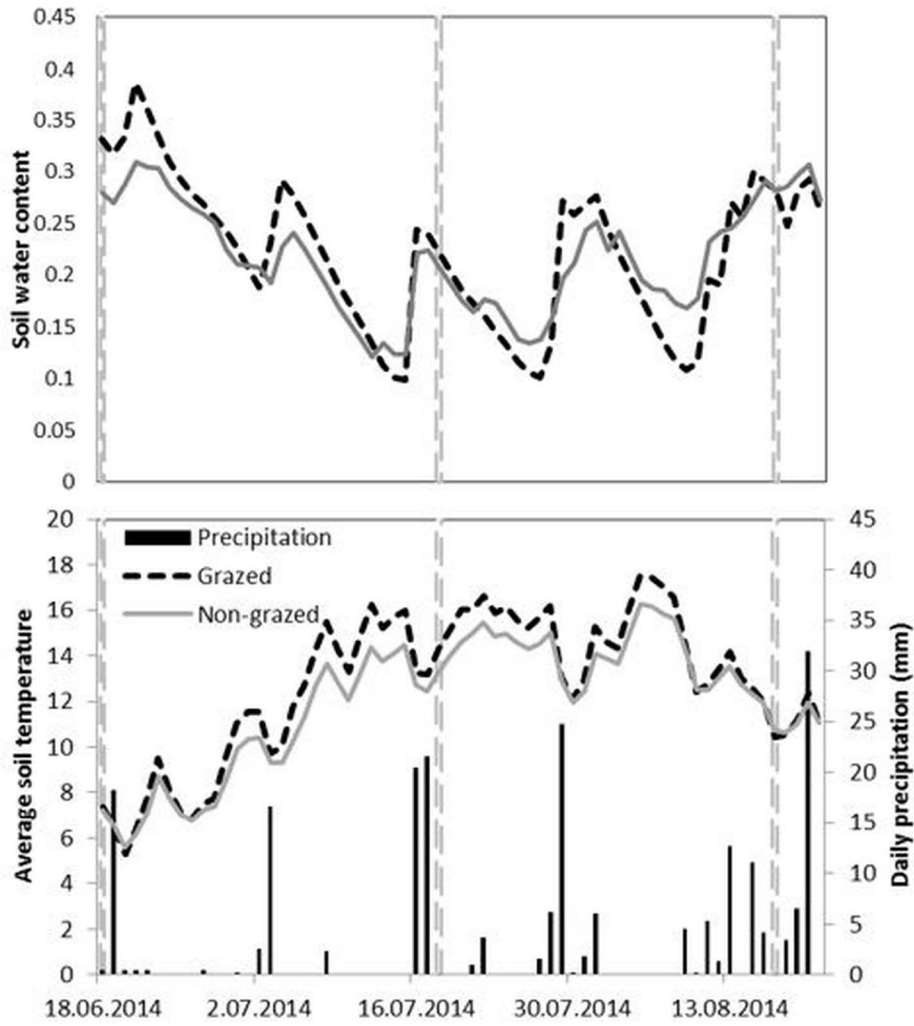


Figure 1. Soil water content ( $\text{m}^3 \text{m}^{-3}$ ) and average soil temperature ( $^{\circ}\text{C}$ ) and precipitation (mm) during the measuring period (June–August). Dotted vertical lines represent the GHG flux measurement days.

156x181mm (300 x 300 DPI)



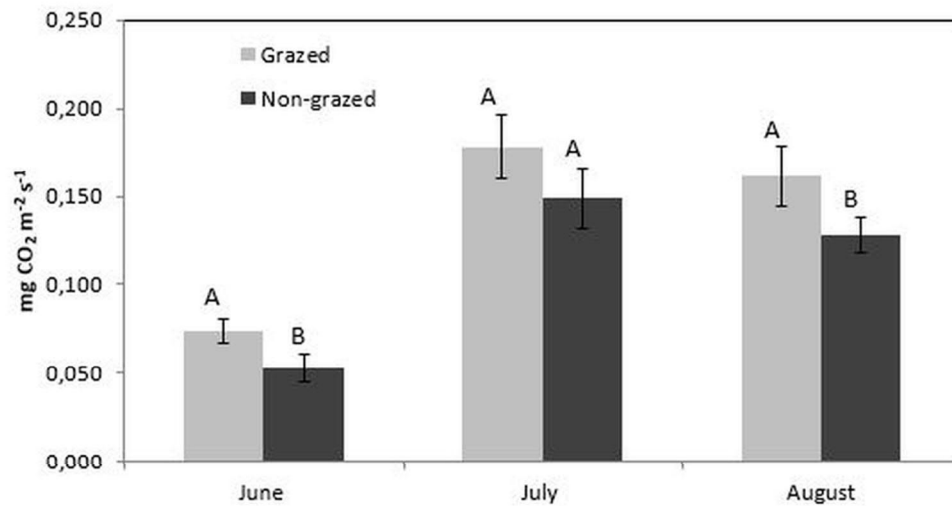


Figure 2. Average field layer CO<sub>2</sub> effluxes (mg CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>) throughout the growing season in areas grazed by reindeer and in non-grazed areas. Error bars represent SE. Statistically significant differences between the grazed and non-grazed areas during separate months are denoted by letters above the bars.

84x49mm (300 x 300 DPI)

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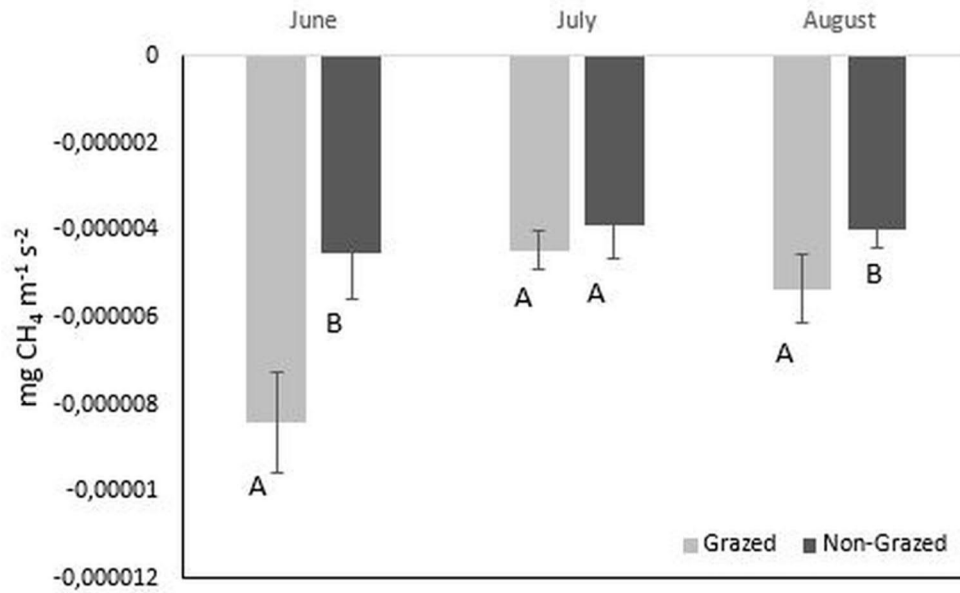


Figure 3. Average field layer methane fluxes ( $\text{mg CH}_4 \text{ m}^{-2} \text{ s}^{-1}$ ) throughout the growing season in areas grazed and not grazed by reindeer. Error bars represent SE. Statistically significant differences between the grazed and non-grazed areas during separate months are denoted by letters above the bars.

74x44mm (300 x 300 DPI)

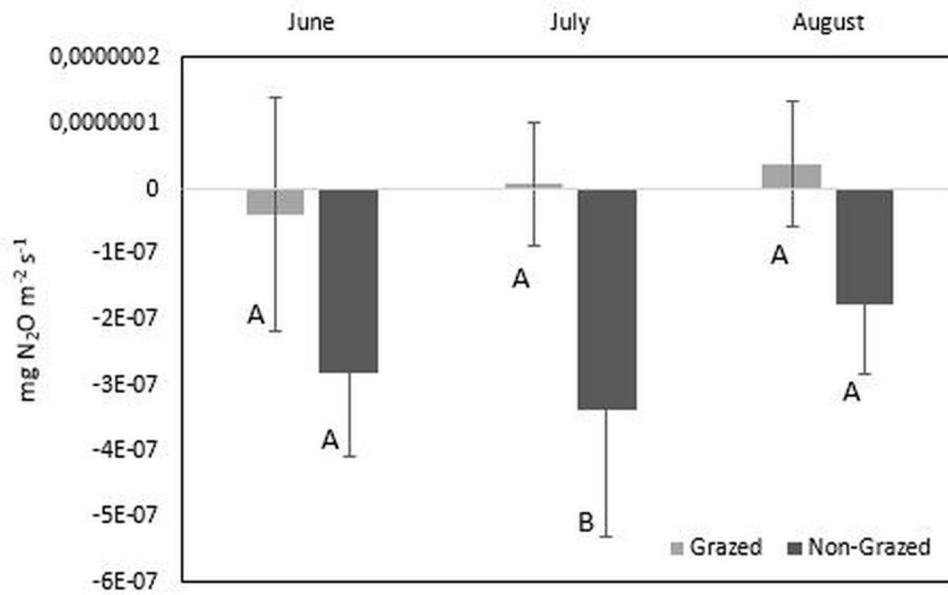


Figure 4. Average field layer fluxes of nitrous oxide ( $\text{mg N}_2\text{O m}^{-2} \text{s}^{-1}$ ) throughout the growing season in areas grazed by reindeer and in non-grazed areas. Error bars represent SE. Statistically significant differences between grazed and non-grazed areas during separate months are denoted by letters above the bars.

75x45mm (300 x 300 DPI)