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TIDAL ANALYSIS OF LONG SERIES

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Abstract

It is shown how M (\geq 5) sets of Fourier coefficients obtained from M successive Fast Fourier Transforms (FFT) of 2¹⁴ tapered hourly heights can be combined to obtain the harmonic constants of the clusters of the main astronomical and shallow-water constituents and their respective satellites. It is also shown how the clusters of the shallow-water constituents are formed.

1 - INTRODUCTION

Trustworthy long series of tidal hourly heights are now available from gauges spread all over the world. In addition to the 'mean' sea level obtained from these series, which is very important from the climatological point of view, tidal analyses of such data start to be worked out (AMIN, 1976; ZETLER *et al.*, 1985), in order to get rid of the perinodal factors f and angles u and to study the micro structure of the phenomenon.

If a series of tidal heights covering 18.61 Julian years is available, each main constituent and the respective satellite constituents can be handled separately, in order to obtain the harmonic constants, amplitude H and phase lag g, for all the cluster. The only exceptions are the neighbour satellite constituents for which the frequency difference is equal to 1/20,000 yr⁻¹, where 20,000 Julian years is the revolution period of the perihelium.

To arrive at such results AMIN (1976) used the CARTWRIGHT-TAYLER (1971) method whereas ZETLER et al. (1985) worked out Fourier computations for individual frequencies, ignoring the side band effects.

Since our harmonic analyses are based on the FFT, it is our purpose to show that this powerful tool can also be used to work out the analysis of such long series.

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The spectral refined method of analysis (FRANCO, 1981) is in current use to obtain accurate harmonic constants from records of 2^{14} or more hourly heights. Then it will be shown how the Fourier coefficients of $M \ge 5$ successive analyses of 2^{14} tapered hourly heights can be combined to compute the amplitudes and phases of the tidal constituents.

An interesting remark is that the proposed combination of the Fourier coefficients is based on the same principles of DOODSON (1928) analysis.

A numerical example will show that 21 diurnal constituents of close frequencies (six main constituents and their satellites) are handled together in a very well conditioned normal system, when M = 10.

2 — THEORETICAL BASIS

In order to make the reasoning clear let us consider the constituents π_1 , P_1 , S_1 , K_1 , ψ_1 and φ_1 which are handled together in the previously mentioned refined analysis. Table 1 shows the frequencies ω_j , in degrees per hour, and the virtual frequencies σ_j , in degrees per $2^{14} = 16,384$ hours, for the above mentioned constituents and their respective satellites.

TABLE 1

Clusters

	-		
Symbol	Cartwright numbers (*)	Angular Frequency °/h	Virtual Frequency º/16384 h
π1	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	14.9156583 14.9178647	298.1450958 334.2949181
P ₁	$ \begin{bmatrix} 1 & 1 & -2 & 0 & -2 & 0 & -1 \\ 1 & 1 & -2 & 0 & -1 & 0 & 1 \\ 1 & 1 & -2 & 0 & 0 & 0 & -1 \\ 1 & 1 & -2 & 2 & 0 & 0 & 1 \\ 1 & 1 & -2 & 2 & 1 & 0 & 1 \end{bmatrix} $	14.9545185 14.9567250 14.9589314 14.9682150 14.9704214	214.8317604 250.9815807 287.1314030 79.2348881 115.3847103
S ₁	11-1 0 0 0 2	15.000000	240.0000019
K,	$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	15.0295786 15.0366558 15.0388622 15.0410686 15.0432751 15.0454815	4.6152916 120.5689564 156.7187767 192.8685989 229.0184212 265.1682415
ψ_1	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	15.0821353 15.0843417	145.7050838 181.8549061
$arphi_1$	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	15.1139223 15.1161287 15.1232059 15.1254123 15.1276187	306.5023098 342.6521320 98.6057949 134.7556171 170.9054375

(*) Coefficients of 90° (phase constants)

The extreme harmonics of each 16,384 hourly data analysis, covering the constituents of table 1, are:

$$n_1 = 14.91565831 \times 16384/360 - 2 = 677$$

$$n_2 = 15.12761874 \times 16384/360 + 2 = 690$$
(2a)

so we have 14 Fourier lines.

According to FRANCO (1981) the Fourier coefficients a_n and b_n of any analysis can be expressed in terms of the tidal constituents by

where Q, R_j and r'_j are, respectively: the number of constituents, the amplitude and the phase of the jth constituent at 8192 hours, reckoned from the initial time, and

$$A_{nj} = (-1)^n \left\{ \frac{\sin\left[(\omega_n - \omega_j) (N+1)\Delta t/2\right]}{N \sin\left[(\omega_n - \omega_j)\Delta t/2\right]} - \frac{\sin\left[(\omega_n - \omega_j) (N+1)\Delta t/2 - \pi/N\right]}{2N \sin\left[(\omega_n - \omega_j)\Delta t/2 - \pi/N\right]} \right\}$$

$$\frac{\sin\left[\left(\omega_{n}-\omega_{j}\right)\left(N+1\right)\Delta t/2+\pi/N\right]}{2N\sin\left[\left(\omega_{n}-\omega_{j}\right)\Delta t/2+\pi/N\right]}\right\}$$
(2c)

In (2c) Δt represents the sampling interval ($\Delta t = 1$ hour), ω_n and ω_j are the Fourier and the jth constituent frequencies, respectively, and N is the number of points in the analysis (N = 16 384).

Si r'_{i} corresponds to the first sub series, then we have for each sub series:

$$\frac{a_n(k)}{b_n(k)} = \sum_{j=1}^{Q} A_{nj} R_j \begin{cases} \cos(r'_j + \sigma_j k) \\ \sin(r'_j + \sigma_j k) \end{cases}$$
(2d)

where

$$k = 0, 1, 2, ..., M - 1$$
 (2e)

for the M analyses.

If we make

$$m = (M - 1)/2$$
 (2f)

then

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$$\frac{a_n(k)}{b_n(k)} = \sum_{j=1}^{Q} A_{nj} R_j \left\{ \begin{array}{c} \cos \\ \sin \end{array} \right\} \left[r'_j + \sigma_j m + \sigma_j \left(k - m \right) \right]$$
(2g)

hence by making

$$\mathbf{r}_{i} = \mathbf{r}_{i}^{\prime} + \sigma_{i} \mathbf{m} \tag{2h}$$

we have

$$\begin{cases} a_n(k) \\ b_n(k) \end{cases} = \sum_{j=1}^{Q} A_{nj} R_j \begin{cases} \cos r_j \cos \sigma_j(k-m) - \sin r_j \sin \sigma_j(k-m) \\ \sin r_j \cos \sigma_j(k-m) + \cos r_j \sin \sigma_j(k-m) \end{cases}$$
(2*i*)

It is easy to grasp that for values of k symmetrical with respect to m, $\cos \sigma_j(k-m)$ have the same value and the same sign whereas $\sin \sigma_j(k-m)$ have the same value but opposite sign. Consequently, we can write:

$$a_n(k) + a_n(2m-k) = \sum_{j=1}^Q 2A_{nj} \cos \sigma_j(k-m) R_j \cos r_j$$
 (2j)

$$b_n(k) - b_n(2m-k) = \sum_{j=1}^Q 2A_{nj} \sin \sigma_j(k-m) R_j \cos r_j$$
 (2k)

$$b_n(k) + b_n(2m-k) = \sum_{j=1}^Q 2A_{nj} \cos \sigma_j(k-m) R_j \sin r_j$$
 (21)

$$-a_n(k) + a_n(2m-k) = \sum_{j=1}^Q 2A_{nj} \sin \sigma_j(k-m) R_j \sin r_j \qquad (2m)$$

where k is expressed by (2e).

Let us check our formulation for M = 10. In this case there will be five equations like (2j) and another five like (2k) for each *n*. Consequently, for $677 \le n \le 690$ there will be $14 \times 10 = 140$ equations to obtain 21 unknowns $R_j \cos r_j$. Since the coefficients of $R_j \cos r_j$ in (2j) and (2k) are the same as those of $R_j \sin r_j$ in (2l) and (2m), then the redundant matrix will be the same for both $R_j \cos r_i$ and $R_i \sin r_j$. Obviously, the normal matrix will be the same also.

In order to evaluate how well conditioned would be the matrix of the normal system resulting from (2j) and (2k) or (2l) and (2m) its elements relative to the set of constituents given in table 1 are shown in table 2. The inverse matrix was also computed to examine the lateral contamination of the side bands (table 3).

The fact that the matrix of the normal system and its inverse are well conditioned is not surprising, since $2^{14} \times 10 = 163840$ hours correspond to 18.69 Julian years, while 18.61 Julian years is the nodal cycle.

	0	0	0	0	0	0	0	0	ŝ	4	4	9-	9	2.	1594	4554	-1857	1779	131	- 126	29987
	0	0	0	0	0	0	0	0	ů.	4	و	9	- 7 -	6	4554	- 1334	1779	1871	-126	29992	· 126
ę	0	0	0	0	0	0	0	0	4	ę	9	- 7	6	-10	-1334	968	1871	2172	29996	- 126	131
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	0	0	0	0	0	- 1	-	9	32	538	1606	4549	-1338	971	-126	30000	1586	1261	968	-1334	4554
¢,	0	0	0	0	0	1	-	-7	- 50	1606	4549	-1338	971	-888	30000	-126	1260	-1099	- 1334	4554	1594
	0	0	- 4	S	ŝ	9	7-7	1601	-2169	154	-140	131	-126	30000	-888	971	11	6,	· 10	6	- 7
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	0	0	9	- 7	6	6-	11	696	1873	-126	30000	-126	131	- 140	4549	1606	2	9	9	9	4
	0	0	2	6	- 10	11	-13	-886	-2174	30000	-126	131	-140	154	1606	- 538	9	ÿ	9-	4	4-
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S	9	2	1601	4552	-1336	1260	-1586	30000	1580	-886	696	1336	4552	1601	- 7	9	-	-	0	0	0
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ፈ	4548	-1337	131	-126	30000	-1871	1779	-1336	6	-10	6	2-	9	Ϋ́	0	0	0	0	0	0	0
	-1338	971	-126	30000	126	1779	-1857	4552	80	6	2	9	Ϋ́	ŝ	0	0	0	0	0	0	0
	971	-887	30000	-126	131	-1857	2165	1601	2	2-	9	Ŷ	ഗ	-4	0	0	0	0	0	0	0
	-126	30000	-887	971	-1337	-83	49	7	1	0	0	0	0	0	0	0	0	0	0	0	0
# ₁	29998	-126	971	-1338	4548	49	-32	9	- i	•	0	0	0	0	0	0	0	0	0	0	0

TABLE 2

Normal matrix - Values multiplied by 1000

TABLE3

Inverted matrix - Values multiplied by 10000

347	2	-11	16	-53	-5	5	-4	0	0	0	0	1	0	0	0	0	0	0	0	0
-2	335	10	-11	16	3	~3	2	0	0	0	0	0	0	0	0	0	0	0	0	0
-11	10	338	1	3	22	-26	-21	1	0	1	-1	3	1	0	0	0	0	0	0	0
16	-11	1	345	6	-19	19	-53	2	- 1	2	-2	8	3	0	1	0	0	0	0	0
-53	16	3	6	345	21	-20	15	-1	0	-1	1	-2	-1	0	0	0	0	0	0	0
~5	3	22	~19	21	338	~.3	-12	0	0	0	- 1	2	1	0	0	0	0	0	0	0
5	-3	- 26	19	-20	-3	338	16	0	1	1	1	-3	1	0	0	0	0	0	0	0
4	2	-21	~53	15	-12	16	356	-15	9	-11	15	-53	-20	3	-3	0	0	0	- 1	0
0	0	1	2	-1	0	0	15	341	24	-21	20	-19	25	5	-4	0	0	0	-1	0
0	0	0	-1	0	0	1	9	24	336	2	-1	0	-1	-19	6	1	-1	- 1	3	0
0	0	1	2	- 1	0	- 1	- 11	-21	2	344	0	2	0	-54	-19	1	- 1	- 2	7	6
0	0	- 1	-2	1	-1	1	15	20	- 1	0	344	-5	1	16	-54	-3	2	2	5	7
1	0	3	8	- 2	2	3	-53	-19	0	2	- 5	344	2	-12	16	1	· 1	· 1	3	-2
0	0	1	3	- 1	1	1	20	25	- 1	0	1	2	337	10	-11	-1	1	1	2	1
0	0	0	0	0	0	0	3	5	-19	-54	16	-12	10	355	0	- 12	10	14	-53	-20
0	0	0	1	0	0	0	-3	-4	6	- 19	-54	16	11	0	354	14	-10	~10	14	52
0	0	0	0	0	0	0	0	0	1	1	3	1	- 1	-12	14	338	- 3	20	-18	19
0	0	0	0	0	0	0	0	0	- 1	- 1	2	- 1	1	10	- 10	- 3	338	24	19	-19
0	0	0	0	0	0	0	0	0	· 1	-2	2	- 1	1	14	- 10	20	- 24	337	-4	2
0	0	0	0	0	0	0	- 1	- 1	3	7	- 5	3	-2	53	14	-18	19	-4	344	0
0	0	0	0	0	0	0	0	0	0	6	7	- 2	1	-20	-52	19	-19	2	0	345

As a result of the analysis, the amplitude $H_j = R_j$ and the phase lag $g_j = (V_{o} - r)_j$ are computed, where V_o is the astronomical argument. No nodal correction is necessary.

3 - COMMENTS ON THE PROPOSED THEORY

Expressions in section 2 were derived for N = 16384. However (2c) is a general formula which can be used for any value of N. Thus the same theory holds to combine M FFT analyses of 2^{ρ} samples for any p. Hence the method can be adapted to combine M short series of 2^{ρ} samples in order to use a micro computer of small capacity for conventional FFT analyses of $M \times 2^{\rho}$ samples.

Expressions (2j) to (2m) can also be used to compute the Fourier spectrum of the long series. Let us consider the long series of 163840 hours and compute the Fourier harmonics of such series between which are contained the extreme frequencies of table 1. All we need to obtain these harmonics is to use 163840 instead of 16384 in (2a). Then we find 6786 and 6887. The respective Fourier frequencies are:

$$\omega_1 = 360 \times 6786/163840 = 14.9106445^{\circ}/h$$
(3a)

$$\omega_2 = 360 \times 6887/163840 = 15.1325684^{\circ}/h$$

The left hand side of (2j) to (2m) can be taken as the 140 known terms of a redundant system where the Fourier coefficients $R_n \cos r_n$ and $R_n \sin r_n$ corres-

ponding to the long series are the 102 unknowns. It is easy to understand that the values of σ_i to be used in (2*j*) to (2*m*) are given by

$$\sigma_{j} = [\omega_{1} + p(\omega_{2} - \omega_{1})/101] \, 16384 \tag{3b}$$

where

$$p = 0, 1, 2, ..., 101$$
 (3c)

The values obtained from (3b) must always be reduced to the first circle. The resolution of the frequency band will be $(\omega_2 - \omega_1)/101 = 0.00219727^{\circ}/h$. And finally the Fourier coefficients $R_n \cos r_n$ and $R_n \sin r_n$ could be used to work out the tidal harmonic analysis of the long series as in the usual method.

Let us now consider the way of finding the confidence intervals.

Since we use spectral methods to analyse tidal records from 15 days to 2 years, tables of residual amplitudes for the Fourier frequency bands of interest are part of the output. Inspection of these tables indicate possible gaps in the constituents choice. Thus the variance of each frequency band is the sum of the half squared residual amplitudes $(\Delta R^2_n/2)$ between the extreme values of *n*. However, such tables of residual amplitudes should not give any useful information if the method of analysis proposed here is used. Hence we took advantage of an old technique to compute the variance (WRIGHT and HAYFORD, 1906).

If P is the redundant matrix, $\{L\}$ the vector of the known terms and $\{X\}$ the vector of the unknowns, then we can write:

Residual variance =
$$-\{X\}^T(P^T\{L\}) + \{L\}^T\{L\}$$
 (3d)

where the symbol T represents the transposed matrix.

Then the confidence intervals are computed in the same way as indicated by FRANCO (1981).

4 — THE SHALLOW-WATER CONSTITUENTS

As early as 1921 Doodson succeeded in separating the main constituents, derived by Darwin in 1883, from the respective satellite constituents. From such clusters one can derive accurate values of f and u. CARTWRIGHT-TAYLER (1971) improved the results found by Doodson through the Fourier analysis of the equilibrium tide. Another analysis by CARTWRIGHT-EDDEN (1973) corrected the Cartwright-Tayler tables. Thus the clusters of the astronomical constituents are well established and can be used in any 18.61 Julian years' analysis.

However, many shallow-water constituents exist in the frequency bands of the purely astronomical tides and it would be impossible to include such shallowwater constituents in the very long period analysis without separating them from their ignored satellite constituents. Thus it was necessary to provide the clusters of these constituents, in order to make the analysis feasible. The full development of the basic theory can be found in FRANCO (1981) and, in an abridged form, in FRANCO (1980). Hence, only the fundamental expression of the tidal height $\zeta(t)$ progressing in a channel will be written:

$$\zeta(t) = \sum_{j=Q}^{Q} a_{j} \exp[i\omega_{j}(t-x/c_{0})] + \sum_{n=1}^{5} \left\{ \sum_{j=Q}^{Q} \sum_{k=Q}^{Q} \dots \sum_{p=Q}^{Q} Z'_{nj}(a_{j}a_{k}\dots a_{p}) \exp[i(\omega_{j}+\omega_{k}+\dots+\omega_{p})(t-x/c_{0})] (4a) \right\}$$

where Q, R_j , r_j and ω_j are as previously defined; *n* is the order of the constituent; *x* is the distance of any section of the channel to its entrance; c_o is the tidal wave celerity at the wave points $\zeta = 0$; Z'_{nj} is the general complex coefficient representing the topographical conditions; *t* is the time and

$$a_{j} = \delta_{j} \mathbf{R}_{j} \exp(-ir_{j}) \quad \delta_{j} = \begin{cases} 1/2 \text{ for } j = 0\\ 0 \quad \text{for } j \neq 0 \end{cases}$$
(4b)

The following convention must be observed:

$$\begin{array}{lll} r_j < 0 & \omega_j < 0 & \text{for} & j < 0 \\ r_j = 0 & \omega_j = 0 & \text{for} & j = 0 \\ r_i > 0 & \omega_i > 0 & \text{for} & j > 0 \end{array}$$

Expression (4a) shows that Z'_{nj} depends on the jth constituent and on the order *n* of the shallow-water term.

Expression (4a) can be considerably simplified, in order to handle the shallow-water constituents individually. In fact, the amplitude of any shallow-water constituent can be expressed by

$$\left[(n+1)!/(k_1!k_2!...) \right] \left(2^n R_j R_k...R_p \right)$$
(4*d*)

for the even species and by

$$[n!n/(k_1!k_2!...)] (2 \ "R_iR_k...R_p)$$
(4e)

for the odd species. k_1 , k_2 , etc., are the number of times that each astronomical constituent enters the combination to form the shallow-water one. If R_j , R_k , etc., are the equilibrium amplitudes of the purely astronomical constituents, then (4*d*) and (4*e*) will give the relative amplitudes of the shallow-water constituents considered.

In order to make clear the use of (4d) and (4e) let us consider the constituent $3M\overline{2S}_2$. One can see that M_2 enters three times (positively) and S_2 twice (negatively) to form the constituent, thus n = (3 + 2) - 1 = 4 which

indicates that $3M\overline{2S}_2$ is a 4th order constituent. In addition, the equilibrium amplitudes of M_2 and S_2 are, respectively, 0.90809 and 0.42248. Consequently, the relative amplitude of $3M\overline{2S}_2$ will be

$$\left[(4+1)!/(3!2!) \right] (2^{-4} \times 0.90809^3 \times 0.42248^2) = 0.08354$$

Let us now consider the odd species constituent $2MNO_7$ to which M_2 contributes twice whereas N_2 and O_1 contribute once each. Since the equilibrium amplitudes of M_2 , N_2 , and O_1 are, respectively, 0.90809, 0.17386 and 0.37694 and n = (2 + 1 + 1) - 1 = 3 then the amplitude of $2MNO_7$ will be:

$$[(3!\times3)/2!)](2^{3}\times0.90809^{2}\times0.17386\times0.37694)=0.06080.$$

It is possible to connect our solution to the DOODSON-WARBURG power law (1941). Let us ignore the topographical coefficient and express the tidal height in the following complex form:

$$\zeta(t) = \sum_{j=Q}^{Q} a_j \exp(i\omega_j t)$$
(4*f*)

where a_i is given by (4b). If three purely astronomical constituents are the only ones taken into account, then the tidal height at t = 0 will be:

$$\zeta(0) = a_1 + a_1 + a_2 + a_2 + a_3 + a_3 \qquad (4g)$$

Thus, according to (4b) and (4c) we have:

$$A = a_{1} + a_{11} = \frac{1}{2} R_{1} [exp(-ir_{1}) + exp(ir_{1})]$$

$$B = a_{2} + a_{12} = \frac{1}{2} R_{2} [exp(-ir_{2}) + exp(ir_{2})] \qquad (4h)$$

$$C = a_{3} + a_{13} = \frac{1}{2} R_{3} [exp(-ir_{3}) + exp(ir_{3})]$$

The Doodson power law indicates that the relative height of the third degree will be:

$$\zeta(0) = (A + B + C)^{3}$$

= A³ + B³ + C³
+ 3(A²B + AB² + AC² + A²C + B²C + BC²) + 6ABC (4*i*)

Using the complex expressions (4h) in (4i), it is possible to derive all the shallowwater constituents resulting from the combination of the purely astronomical constituents A, B and C. The powers 3, 2, 1 and 0 of A, B and C indicate the number of times A, B and C enter the combination, then the numerical coefficients appearing in (4i) can be evaluated by the expression:

$$(n+1)!/(k_1!k_2!k_3!)$$
 (4j)

where n = 2 and

$$k_1 = 3$$
 $k_2 = 0$ $k_3 = 0$ for A^3 , B^3 and C^3
 $k_1 = 2$ $k_2 = 1$ $k_3 = 0$ for A^2B , AB^2 , etc.
 $k_1 = 1$ $k_2 = 1$ $k_3 = 1$ for ABC

Note that

$$k_1 + k_2 + k_3 = n + 1$$

and that (4j) is valid for any value of n.

We can derive from (4*h*):

$$A^{3} = \frac{1}{4} R_{1}^{3} (\cos 3r_{1} + 3\cos r_{1})$$
(4k)

$$A^{2}B = \frac{1}{4} R_{1}^{2} R_{2} \left[\cos(2r_{1} - r_{2}) + \cos(2r_{1} + r_{2}) + 2\cos r_{2} \right]$$
(41)

$$ABC = \frac{1}{4} R_1 R_2 R_3 \left[\cos(r_1 + r_2 + r_3) + \cos(r_1 - r_2 + r_3) + \cos(r_1 - r_2 - r_3) + \cos(r_1 - r_2 - r_3) \right]$$
(4m)

An interesting remark is that A^3 contains the shallow-water constituent $\frac{3}{4}R_1^3\cos r_1$, in phase with the purely astronomical A, and A^2B contains the shallow-water constituent $\frac{2}{4}R_1^2R_2\cos r_2$, with the same phase as B. These shallow-water disturbing effects on the astronomical constituents were also detected by FRANCO (1977).

The products of the coefficients of (4k) to (4m) by (4j) are the relative amplitudes of the shallow-water constituents. Fraction 1/4 appearing in these expressions can be generalized by writing $1/2^n$, where *n* is the order of the shallow-water constituent in our development (FRANCO, 1981), for the shallowwater even species. Consequently, each shallow-water constituent can be obtained individually. In fact, B³, C³, B²A, C²A, etc., can be formed exactly in the same way as A³ and A²B.

If (4) is compared to the coefficient of $2 \ {}^{n}R_{j}R_{k} \dots R_{p}$ of (4d) one can see that they are identical, thus the relative amplitudes of the even species derived by FRANCO (1981) follow exactly the Doodson power law. As far as the odd species

are concerned expression (4e) shows that the resulting amplitudes are n/(n + 1) times the power law amplitudes. In fact, our results would be the same if the topographical coefficients were multiplied by n/(n + 1). However, we do not use this factor in order to maintain the connection between theory and practice.

It is remarkable indeed that such law was derived by Doodson through a very simple geometrical construction based on the tidal wave celerity, developed into series and reduced to the first order term.

When n = 1 it is possible to keep a single coefficient for all the species. In addition, when n = 2 an algebraic manipulation permits to use (4d) only for the even an odd species.

The Doodson power law does not work for n = 1, to derive the amplitudes of the shallow-water long period constituents.

Notwithstanding its generality, the theory was used to compute the relative amplitudes of the main shallow-water constituents, through the combination of some main Darwin constituents. However, according to the modern harmonic developments, these constituents were shown to be the main lines of clusters where the minor constituents are considered as satellites of the main constituents. Thus it becomes evident that the shallow-water effects should be represented by the combination of the whole clusters of the main constituents. Then a program was written to generate the clusters of the shallow-water constituents. A sample of a shallow-water cluster is given in table 4.

TABLE 4

The $3M\overline{2S}_2$ cluster

Symbol	Cartwright	Relative	Frequency
	numbers (*)	amplitude	Degrees/Hour
3M2S ₂	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	0.00002 0.00048 0.00936 0.08354 0.00038 0.00015 0.00005	26.943029 26.947900 26.950106 26.952313 26.954519 26.961596 26.963803

(*) Coefficients of 90° (phase constants).

5 - COMMENTS ON THE RESULTS

Tables in the Annex contain the harmonic constants resulting from an analysis of 163840 hourly heights of Cananéia (25°01.0'S, 47°55.5'W, Brazil), with the solution of six different systems, and from a refined analysis of 16384 tidal hourly heights, both starting on 1 January 1955.

The first important thing to be observed in these tables is that with only four exceptions all the main constituents rejected have been so in both analyses.

In fact, all the clusters of $2S\overline{K}_2$ and $2MS\overline{T}_4$, including the main constituents, were statistically rejected in the 18.69 years' analysis. It is interesting to note that many constituents not flagged in both analyses have amplitudes much less than 1 cm.

The long series analysis showed that several constituents appear very stable in spite of their small amplitudes. Thus two important conclusions can be drawn:

- a) the constituents, including the shallow-water ones, were judiciously chosen;
- b) the refined method resisted to a very severe test.

In order to give a clear idea of the close agreement between theory and practice, table 5 was drawn. Column 1 of that table shows the most important satellite constituents, identified with the coefficient of N' (360° minus the longitude of the ascending node) after the symbol of the main constituent of the cluster. The

TABLE 5

Comparison between practice and theory

Symbol	Equilibrium relation	H(I)	H(A) entimetr	+/- res)	DG (de	+/- grees)
K 1(-1)	0.01983	0.13	0.16	0.15	5.65	63.25
K1(+1)	0.13556	0.85	0.80	0.15	5.39	10.46
M2(-1)	0.03733	1.34	1.34	0.24	-7.58	10.33
*MKS2(+1)	0.30231	0.14	0.14		-27.63	
K2(+1)	0.29796	2.19	2.24	0.24	-0.62	6.21
M3(-1)	0.05640	0.41	0.26	0.24	3.84	72.65
M4(-1)	0.07465	0.50	0.44	0.13	-7.34	16.56
MS4(-1)	0.03509	0.12	0.18	0.15	-43.21	58.60
MK4(+1)	0.30008	0.29	0.30	0.15	-5.79	30.64

second column gives the theoretical ratio between the amplitude of each satellite and the amplitude of the main constituent in the respective cluster; the third column is the inferred amplitude H(I) of each satellite; the fourth column gives the amplitudes H(A) found through the long series analysis; the sixth column is the difference between the phase lag of each satellite and the phase lag of the main constituent in the respective cluster, both phase lags being given by the long series analysis. The fifth and seventh columns are the confidence intervals of amplitudes and phases, respectively, in the long series analysis.

The rejected satellite $MKS_2(+1)$ was included because its amplitude and phase are in good agreement with theory.

Notwithstanding the objections arisen against FRANCO's theory (1980, 1981) on the shallow-water constituents generation, because friction was not taken into account, some fairly good practical confirmation resulted from the present analysis. In fact, table 5 shows how the most important satellite shallow-water constituents, as referred to the main constituent in the cluster, follow closely the relative amplitudes established by the theory.

A last comment must be done on the third diurnal cluster, which has a

large contribution of constituent M_3 surrounded by several overlapping clusters of shallow-water constituents. In fact, the satellites of MM'_3 ($M_2 + M_1$), NK'_3 ($N_2 + K_1$), LO_3 and KQ_3 have exactly the same frequency but different amplitudes. Then we summed up the amplitudes of the satellites with equal frequencies and built up a compound cluster, the amplitudes of which guided us in the choice of the constituents to include in the analysis, which gave results confirming the relative importance of the constituents. The small effects of the second order shallow-water constituents were ignored.

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(See Annex, pages 154-158)

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Harmonic constants (H, G) of Cananeia (Brazil, 25° 1.0' S, 47° 55.5' W), resulting from analyses of spans of 18.69 years and 16384 hours

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					1		
hours +	17.27	0.42	27.86	23.11	45.25	28.11	0.88
(degrees) 16384 G	96.88	93.80	269.61	162.15	221.71	101.99	94.48
Phase lags (years + .	21.62	65.94 0.38 0.38	***** 39.71	30.24		21.35	62.94 68.54 0.61
18.69 G	87.44 109.21 85.35	5.63 79.21 102.23 94.65 106.48 99.24 98.39 98.39 49.08	252.03 232.87	121.65 299.76 170.78 198.41 299.31 195.71	70.30 176.40 258.78 93.19	126.24	65.71 63.54 95.68 119.87 65.24
hours +	0.26	0.26	0.26	0.26	0.33	0.34	0.34
es (cm) 16384 H	0.89	35.89	0.56	0.65	0.47	0.72	22.27
Amplitud years + _	0.24	0.24 0.24 0.24	0.24	0.24		0.24	0.24 0.24 0.24
18.69 H	0.21 0.65 0.09	0.14 0.26 1.34 35.78 0.22 0.17 0.17 0.17 0.17	0.02 0.38	0.07 0.07 0.148 0.14 0.14 0.08	0.08 0.07 0.22 0.24	0.66	0.27 0.26 22.71 0.22 0.15
Angular frequency (degrees/h)	28.9408311 28.9430375 28.9452440	28.9726142 28.9796914 28.9818978 28.9811042 28.9863106 28.983170 28.9933879 28.9933879 28.9955943	29.0229645 29.0251709	29.0618287 29.0640351 29.0662415 29.0684479 29.0706543 29.0755252	29.9134499 29.9156563 29.9178627 29.9200691	29.9589333	29.9955872 29.9977936 30.0000000 30.0022064 30.0041128
Cartwright numbers and phase constants	2 0 - 1 0 - 1 1 2 2 0 - 1 0 0 1 0 2 0 - 1 0 1 0	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	2 0 1 0 -1 -1 2 2 0 1 0 0 -1 0	2 0 2 0 -2 0 0 2 0 2 0 1 0 2 2 0 2 0 0 0 0 2 0 2 0 1 0 2 2 0 2 0 0 0 2 0 2 0 0 2 0 2 0 0 2 0 2 0	2 2-4 0-2 0 0 2 2-4 0-1 0 0 2 2-4 0 0 0 0 2 2-4 0 1 0 2	2 2 - 3 0 0 1 0	2 2-2 0-2 0 0 2 2-2 0-1 0 0 2 2-2 0 0 0 0 2 2-2 0 1 0 0 2 2-2 0 1 0 0 2 2-2 0 2 0 2
Constituents	27* 28 MTS2 29*	30* S-KO2 31 32 33 M2 33 M2 34* S-KO2 35* S-KO2 37* 38*	39* 40 MST2	41* 42* 43 MKS2 44* 46*	47* 48* 49* 2SK2 50*	51 T2	52 S.KP2 53 54 55 55* S.KP2 56* S.KP2 56* S.KP2

TIDAL ANALYSIS OF LONG SERIES

hours +		41.97		2.63				·		*****					2.95							*****			
(degrees) 16384 G		116.01		84.72						85.40	2				236.85							241.88			
Phase lags years + -	****	25.07	*****	1.88	0.21				*****	50.83		*****	****	72.65	1.93	*****	17.84			*****	*****	****			
18.69 G	231.11	119.33 222.29	108.46	86.30 86.00	104.56		16.36	253.44	118.39	100.91	88.57	124 20	279.27	236.02	239.86	239.29	84.83	118./4 189.83	357.30	58.74	96.86	29.76	203.61 35.60	00.00	77.077
hours +		0.34		0.33						*****					0.33							****			
des (cm) 16384 H		0.51		7.24						0.26					b.4 6							0.10			
Amplitu years +	•••••	0.24	****	0.24 0.24				*****	*****	0.24	****	****	*****	0.24	0.24		0.24	*****	*****	****	*****	****		*****	
18.69 H	0.02	0.57 0.04	0.08	7.36 2.24	0.23	900	0.05	0.21	0.07	0.31	0.07	0.06	0.09	0.26	07.1	0.18	0.80	0.10	0.03	0.05	0.03	0.11	0.08 0.06	0.07	
Angular frequency (degrees/h)	30.0092837	30.0410667 30.0432770	30.0799309	30.0821373 30.0843437	30.0865501	43 389758A	43.3849644	43.3942481	43.3964545	43.3986609	43.4079445	43.4671017	43.4693081	43.4739499 43.4761563	0001014.04	43.4785918	43.480/982 43 4830046	43.4852110	57.8794519	57.8816583	57.8838648	57.8860711	57.8953548	57.9249354	
Cartwright numbers and phase constants	$2 \ 2 \ -2 \ 2 \ 0 \ 0 \ 0$	2 2-1 0 0-1 2 2 2-1 0 1 1 2	2 2 0 0 -1 0 2	2 2 0 0 1 0 0	2 2 0 0 2 0 0	3 0+2-1-3 0 1	3 0 -2 -1 -2 0 1	3 0 - 2 1 - 2 0 1	3 0 - 2 1 - 1 0 - 1	3 0 - 2 1 0 0 - 1	3 0 - 2 3 0 0 1	3 0 0 - 1 - 2 0 - 1	3 0 0 - 1 - 1 0 - 1	3000000		3001-10-1	3001101	3 0 0 1 2 0 1	4 0 - 2 0 - 3 0 2	4 0-2 0-2 0 0			$\begin{array}{cccccccccccccccccccccccccccccccccccc$	4 0 -1 0 -1 1 2	
Constituents	57*	58 R2 59*	60* 61 K2	62 62	63*	64*	65*	66*	67" 60 505	68 SQ3 60*	60	70* S-NK3	71° 5-NK3	73 M3		75 NK3	76*	-77	78* 70*	13	81* 2MCKA	80#	83 *	84*	

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hours		64.29	I.I.	33.49	**	*	5.90	20.60
(degrees) 16384 C	2	260.73	211.43	144.55	308.46	57.83	319.90	323.34
Phase lags years	 	27.96	16.56 1.07		71.49		58.60 2.58	8.90 30.64
18.69	כ	308.11 291.16	59.82 239.50 218.38 211.04 25.95 25.95 119.22	343.56 134.36	120.01 46.91 40.13 41.22 322.63 358.42	169.81 11.54 63.64 308.87	287.81 2.55 319.34 137.72 240.22	273.45 10.07 319.08 324.87
hours	ł	0.13	0.13	0.13	ł		0.37	0.37
es (cm) 16384	Ę	0.15	6.98	0.24	0.11	0.11	3.64	1.06
Amplitud years	 	0.13	0.13	***	0.13		0.15 0.15	0.15 0.15
18.69	E	0.27 0.08	0.07 0.02 0.44 6.71 0.12 0.09 0.07	0.03 0.09	0.05 0.04 0.13 0.07 0.02 0.09	0.03 0.08 0.06 0.06	0.05 0.18 3.36 0.07 0.05	0.04 0.05 0.30
Angular frequency	(degrees/h)	57.9271417 57.9293482	57.9615892 57.9637956 57.9660020 57.9682084 57.9752857 57.9774921 57.9776985	58.0070687 58.0092751	58.0459329 58.0481393 58.0503457 58.0525521 58.0547585 58.0596294	58.0892060 58.0914124 58.0936188 58.0958252	58.9796914 58.9818978 58.9841042 58.9933879 58.9955943	59.0618287 59.0640351 59.0662415 59.0684479
Cartwright numbers and	phase constants	4 0 1 0 0 1 0 4 0 1 0 1 1 0	4 0 0 -3 0 2 4 0 0 -2 0 0 4 0 0 -1 0 2 4 0 0 0 -1 0 2 4 0 0 2 -1 0 2 4 0 0 2 1 0 2 4 0 0 2 1 0 2 4 0 2 1 0 0 0	4 0 1 0 1 1 2 4 0 1 0 0 1 0	4 0 2 0 -2 0 0 4 0 2 0 -1 0 2 4 0 2 0 0 0 0 4 0 2 0 1 0 0 4 0 2 0 1 0 0 4 0 2 0 2 0 0 4 0 2 0 2 0 0 4 0 2 0 2 0 0	4 0 3 01 -1 2 4 0 3 0 01 0 4 0 3 0 1 -1 0 4 0 3 0 1 -1 0 4 0 3 0 2 -1 0	4 2 2 0 2 0 0 4 2 2 0 1 0 2 4 2 2 0 1 0 2 4 2 2 0 0 0 0 4 2 2 2 0 0 0 4 2 2 2 1 0 0	4 2 0 2 0 0 4 2 0 1 0 2 4 2 0 0 1 0 2 4 2 0 0 1 0 0 4 2 0 1 0 0 0
Constituents		85 2MTS4 86*	87* 888 90 M4 91* 93*	94* 95* 2MST4	96* 97* 98 2MKS4 99* 100* 101*	102* 103* 2MKT4 104* 105*	106* 107 108 MS4 109* 110*	111* 112* 113 MK4 114

TIDAL ANALYSIS OF LONG SERIES

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Symbol S

Phase lags (degrees) 18.69 years 16.384 hours G +- G +--

Amplitudes (cm) ears 16384 hours +- H + .

18.69 years H + -

Angular frequency (degrees/h)

Cartwright numbers and phase constants

Constituents

302.59 72.76 219.53

.....

0.05 0.02 0.07

59.0706543 59.0755252 59.0777316

115* 116* 117*
