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Motion and structure of atmospheric mesoscale baroclinic vortices: dry air and weak environmental shear

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A strongly tilted, nearly axisymmetric vortex in dry air with asymmetric diabatic heating is analyzed here by matched asymptotic expansions. The vortex is in gradient wind balance, with vortex Rossby numbers of order unity, and embedded in a quasi-geostrophic (QG) background wind with weak vertical shear. With wind speeds of 60 – 120 km/h, such vortices correspond to tropical storms or nascent hurricanes according to the Saffir-Simpson scale. For asymmetric heating, nonlinear coupling of the evolution equations for the vortex' tilt, its core structure, and its influence on the QG background is found. The theory compares well with the established linear theory of precessing quasi-modes of atmospheric vortices, and it corroborates the relationship between vortex tilt and asymmetric potential temperature and vertical velocity patterns as found by Jones (Q.J. Roy. Met. Soc., **121**, 1995) and Frank and Ritchie (Mon. Wea. Rev., **127**, 1999) in simulations of adiabatic tropical cyclones. A relation between the present theory and the local induction approximation for three dimensional slender vortex filaments is established.

1. Introduction

1.1. Related earlier work

There is a large body of literature, starting with Stewart (1943); Morikawa (1960) (see also Egger 1992; Reznik 1992; Reznik & Kizner 2007*a,b*, and references therein) that investigates the dynamics of concentrated atmospheric vortices in the limit of pointwise vorticity singularities in two-dimensional or multilayer shallow flow models. The related theories focus on the vortex motion but do not investigate the vortex core structure.

At the opposite end, the core structure of such vortices is often modelled using the Eliassen's balanced vortex model and extensions thereof (see, e.g., Eliassen 1952; Charney & Eliassen 1964; Schubert & Hack 1983; Emanuel 1991; Wirth & Dunkerton 2009, and references therein), which describes the evolution of an axisymmetric, upright vortex in gradient wind balance, i.e., in a regime where the radial pressure gradient is balanced to leading order by the centripetal and Coriolis accelerations. The developed theories provide no information regarding the vortex motion.

Asymptotic theories that resolve the vortex core structure and describe its coupling to the environmental flow and/or the motion of its centre were developed by Ling & Ting (1988) in the limit of geostrophic balance, by Smith & Ulrich (1990); Smith (1991) for two-dimensional flow on a β -plane, and by Reznik & Grimshaw (2001); McWilliams *et al.* (2003) in the context of shallow water theory. By restricting to two-dimensional shallow fluid models, these theories cannot address the vertical structure of the flow and, especially, a possible tilt of the vortex axis.

Shapiro & Montgomery (1993) develop a linearized asymptotic theory that is close, in terms of its principal aim, to the present work. It describes the evolution of three-dimensional balanced perturbations on an upright, axisymmetric vortex in gradient wind balance. It includes circumferential Fourier mode one perturbations that can be inter-

puted as a weak vortex tilt with horizontal displacements small compared with the vortex core radius. The authors also suggest extensions of their linearized perturbation analysis to include weakly nonlinear effects.

The dynamics of vortex tilt, and especially the resilience of a vortex to background vertical shear has been analysed by Reasor & Montgomery (2001); Reasor *et al.* (2004) both through theoretical analyses and numerical simulations based on the primitive flow equations. Their linear theory of vortex Rossby waves (VRWs) (see Montgomery & Kallenbach 1997) reveals a damping mechanism for the vortex tilt dynamics and allows them to explain the stabilization of an atmospheric vortex against vertical shear of the environmental flow. At the same time, these developments reveal weakly damped linear “quasi-modes” which, in a suitable regime of vortex Rossby numbers, can be interpreted as small-amplitude precessions of the entire vortex about its unperturbed vertical axis. Surprisingly, the theory – although based on linearizations – yields very good comparison with fully nonlinear simulations when relative deviations from the unperturbed state are of order unity, i.e., when the horizontal displacements of the vortex core are comparable to its characteristic diameter. In a first application the present analysis corroborates the existence of such large displacement precessions within the gradient wind regime and yields agreement of the predicted precession frequencies for vortex Rossby numbers of order unity. See also related numerical simulations by Jones (1995, 2004).

1.2. *Summary of the present approach and main results*

In this paper we consider concentrated nearly axisymmetric strongly tilted baroclinic vortices in gradient wind balance that are embedded in a quasi-geostrophic (QG) farfield flow with weak vertical shear. We restrict to the case of dry air with an ideal gas equation of state but include the principal effects of asymmetric diabatic sources in the vortex core.

Let $\mathbf{X}_0(t)$ denote the overall horizontal position of the vortex measured in terms of

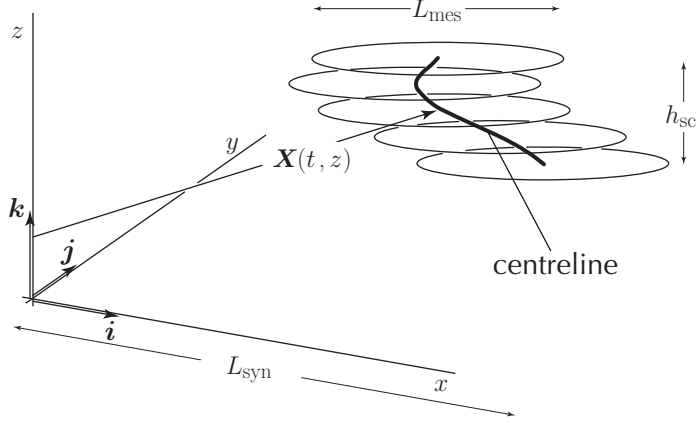


FIGURE 1. Tilted concentrated mesoscale vortex of characteristic height h_{sc} and typical core size L_{mes} embedded in a synoptic scale (L_{syn}) environment, where $h_{sc} \ll L_{mes} \ll L_{syn}$.

a synoptic length scale, L_{syn} , and let $\mathbf{X}^{(1)}(t, z) = \mathbf{i} X^{(1)}(t, z) + \mathbf{j} Y^{(1)}(t, z)$ describe the horizontal displacement of the vortex centre from that large-scale position at height z , measured in units of a typical vortex core diameter, L_{mes} . Then $\delta = L_{mes}/L_{syn} \ll 1$ for concentrated vortices, and the horizontal position of the vortex centre at height z measured in units of L_{syn} is (cf. Fig. 1)

$$\mathbf{X}(t, z) \equiv X(t, z)\mathbf{i} + Y(t, z)\mathbf{j} = \mathbf{X}_0(t) + \delta \mathbf{X}^{(1)}(t, z). \quad (1.1)$$

Thus we allow for horizontal displacements comparable to a typical vortex core size of $L_{mes} \sim 150$ km. This amounts to very strong tilt since the characteristic height of the vortex is comparable to $h_{sc} \sim 10$ km. With \mathbf{X}_0 independent of z we avoid unrealistically large tilt with horizontal displacements much larger than the vortex core size.

Let u_θ denote the primary, possibly baroclinic, circulation of the vortex with total strength $\Gamma(t, z)$. The gradient wind regime involves the dominant three term balance of the pressure gradient, centrifugal acceleration, and Coriolis terms and implies vortex Rossby numbers of order unity: With p, ρ , and f_0 denoting pressure, density, and the

Coriolis parameter, respectively, we have $\partial p/\partial r \sim \rho u_\theta^2/r \sim \rho f_0 u_\theta$ in the gradient wind regime, with r the horizontal distance from the centreline. Then, with L_{mes} of the order of a characteristic vortex radius, the comparison of the latter two terms yields an estimate for the vortex Rossby number, $\text{Ro}_v = u_\theta/f_0 L_{\text{mes}} = \mathcal{O}(1)$. Accordingly we let $\text{Ro}_v = \mathcal{O}(1)$ as $\delta \rightarrow 0$ in the subsequent asymptotic analysis.

Motivated by Ting and co-workers' approach to the dynamics of slender fluid vortices (see Ling & Ting 1988; Ting *et al.* 2007) we derive nonlinear coupled evolution equations for the vortex core structure, the vortex tilt, and the QG outer flow through matched asymptotic expansions of the compressible flow equations in three space dimensions.

Specifically, we consider in section 3 the QG farfield flow with an embedded tilted singular baroclinic line vortex. Given the vortex centreline from (1.1), the vortex' circulation, $\Gamma(t, z)$, and the atmospheric background state characterized by the density, $\rho_0(z)$, and the mean perturbation potential temperature stratification, $\Theta_1(z)$, we have the

Perturbed quasi-geostrophic farfield flow

$$\mathbf{u} = -\nabla_{\parallel}^{\perp} \psi, \quad (1.2a)$$

where ∇_{\parallel} is the horizontal gradient, and

$$\psi = \psi_0^{2d} + \psi_1^{2d} + \psi_2^*, \quad (1.2b)$$

is the QG streamfunction with the local vortex induced components

$$\psi_0^{2d} = -\frac{\Gamma}{2\pi} \ln r \quad \text{and} \quad \psi_1^{2d} = M_0 r^2 (\ln r - 1) + \delta M_1 r \ln r - \delta^2 M_2. \quad (1.2c)$$

Here $r(t, \mathbf{x}, z) = |\mathbf{x} - \mathbf{X}(t, z)|$ is the horizontal distance from the centreline and the coefficients M_i depend on $\rho_0(z)$, $\Theta_1(z)$, and $\Gamma(t, z)$ through eqs. (3.22) given below. The regularized stream function, ψ_2^* , satisfies

$$(\nabla_{\parallel}^2 + \mathcal{L}_z) \psi_2^* = -\mathcal{L}_z \psi_1^{2d} - q_r \quad \text{with} \quad \mathcal{L}_z[\cdot] = \frac{f_0^2}{\rho_0} \frac{\partial}{\partial z} \left(\frac{\rho_0}{d\Theta_1/dz} \frac{\partial \cdot}{\partial z} \right). \quad (1.2d)$$

The regular part of the potential vorticity, q_r , which includes planetary rotation, is advected with the horizontal flow, i.e.,

$$\left(\frac{\partial}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \right) q_r = 0. \quad (1.2e)$$

In section 4 we show that changes in the leading order axisymmetric circumferential flow velocity, $u_{\theta}(t, \hat{r}, z)$, in the vortex core are described by the

Core structure evolution equation

$$\frac{\partial u_{\theta}}{\partial t} + w_0 \frac{\partial u_{\theta}}{\partial z} + u_{r,00} \left(\frac{\partial u_{\theta}}{\partial \hat{r}} + \frac{u_{\theta}}{\hat{r}} + f_0 \right) = -u_{r,*} \left(\frac{u_{\theta}}{\hat{r}} + f_0 \right), \quad (1.3a)$$

where

$$w_0 = \frac{Q_{\Theta,0}}{d\Theta_1/dz}, \quad (1.3b)$$

$$u_{r,00} = -\frac{1}{\hat{r}} \int_0^{\hat{r}} \frac{r'}{\rho_0} \frac{\partial}{\partial z} \left(\rho_0 \frac{Q_{\Theta,0}}{d\Theta_1/dz} \right) dr', \quad (1.3c)$$

$$u_{r,*} = \frac{1}{2} \left[\frac{\partial X^{(1)}}{\partial z} \frac{Q_{\Theta,12}}{d\Theta_1/dz} + \frac{\partial Y^{(1)}}{\partial z} \frac{Q_{\Theta,11}}{d\Theta_1/dz} \right]. \quad (1.3d)$$

Here $Q_{\Theta,0}$ and $Q_{\Theta,1k}$ for $k = 1, 2$ are the axisymmetric and first circumferential Fourier mode amplitudes of the diabatic source term, $u_{r,00}$ and w_0 are the axisymmetric parts of the radial and vertical velocities that are directly induced by axisymmetric diabatic heating, and $\hat{r} = r/\delta$ is the stretched radial coordinate resolving the vortex core.

The terms on the left of (1.3a) are analogous to those found in the circumferential momentum balance of an axisymmetric system. The forcing on the right is due to nonlinear interactions of the vortex tilt with asymmetric diabatic heating. The forcing strength, represented by $u_{r,*}$, depends on the relative arrangement of the tilt direction and the diabatic heating asymmetries. Suitable arrangement can spin up or spin down the vortex. This is in agreement with results by Shapiro & Montgomery (1993) on balanced small

perturbations of an axisymmetric, vertically aligned vortex and extends them to vortices with leading order tilt.

In section 5 we derive the

Vortex centreline equations of motion

$$\frac{d\mathbf{X}_0}{dt} = \mathbf{u}_0, \quad (1.4a)$$

$$\frac{\partial \mathbf{X}^{(1)}}{\partial t} = \mathbf{u}_0^{(1/2)} + \mathbf{X}^{(1)} \cdot (\nabla_{\parallel} \mathbf{u}_0) - \ln \frac{1}{\delta} (\mathbf{k} \times \mathbf{M}_1) + (\mathbf{k} \times \Psi), \quad (1.4b)$$

where \mathbf{u}_0 and $\nabla_{\parallel} \mathbf{u}_0$ are the barotropic leading order regular part of the quasi-geostrophic background flow evaluated at $\mathbf{X}_0(t)$ and its horizontal gradient, respectively, $\mathbf{u}^{(1/2)}$ is a possibly baroclinic contribution to the background flow of order $\mathcal{O}(\delta)$ (the $(\cdot)^{(1/2)}$ -superscript will be explained shortly), \mathbf{M}_1 is related to M_1 from (1.2c), (3.22), and Ψ is a complex expression involving the vortex core structure, $u_{\theta}(t, \hat{r}, z)$, the asymmetric heating, and the vortex tilt (see eqs. (5.12), (5.10), and (4.33) below).

The main conclusions to be drawn from (1.4) are, first, that in the considered regime with relatively weak heating and tilt-induced vertical motions a vortex cannot sustain strong vertical shear of the background flow within which it is embedded. We must require the leading order background flow, \mathbf{u}_0 , to be barotropic, i.e., independent of z . Secondly, there is a self-induced motion of a tilted baroclinic vortex which induces horizontal displacements comparable to the vortex core size (terms involving \mathbf{M}_1 and Ψ). See section 6 for a comparison of predictions using (1.4) with established linear theories for small perturbations of an initially vertically aligned vortex by Reasor *et al.* (2004).

Two remarks of caution are in order: The present theory proceeds independently of whether we assume cyclonic or anti-cyclonic mean vortex rotation. It does assume, however, that the vortex is stable over the characteristic time scale considered and this may

exclude anticyclonic vortices with Rossby numbers larger than unity from consideration because of possible centrifugal instabilities (see, e.g., Afanasyev & Peltier 1998).

Also, the reader will have noticed the occurrence of the logarithm of the perturbation parameter in (1.4). This does not come as a surprise considering established theories for slender fluid vortices (see, e.g., Ting *et al.* 2007). Here we follow Ting and co-workers, who consider such logarithmic terms as of order unity relative to powers of the perturbation parameter. This seems appropriate here as $\delta > 1/10$ for realistic applications of the present theory (see section 7.3 for further discussion).

For conclusions, limitations of the theory, and an outlook see section 7.

2. Dimensionless governing equations and distinguished limits

The inviscid rotating compressible flow equations in the beta plane approximation,

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \mathbf{u} + w \frac{\partial \mathbf{u}}{\partial z} + \frac{1}{\rho} \nabla_{\parallel} p + f \mathbf{k} \times \mathbf{u} = 0, \quad (2.1a)$$

$$\frac{\partial w}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} w + w \frac{\partial w}{\partial z} + \frac{1}{\rho} \frac{\partial p}{\partial z} = -g, \quad (2.1b)$$

$$\frac{\partial \rho}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \rho + w \frac{\partial \rho}{\partial z} + \rho \nabla_{\parallel} \cdot \mathbf{u} + \rho \frac{\partial w}{\partial z} = 0, \quad (2.1c)$$

$$\frac{\partial \Theta}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \Theta + w \frac{\partial \Theta}{\partial z} = Q_{\Theta}, \quad (2.1d)$$

$$\Theta = \frac{p_{\text{ref}}}{\rho R} \left(\frac{p}{p_{\text{ref}}} \right)^{\frac{1}{\gamma}}, \quad (2.1e)$$

are our point of departure. Here $p, \rho, \Theta, \mathbf{u}, w$ are pressure, density, potential temperature, and the horizontal and vertical velocities, and γ is the specific heat ratio. The Coriolis parameter is $f = f_0 + \beta y$ where y denotes the local meridional coordinate and $\beta = \text{const.}$ The horizontal gradient is $\nabla_{\parallel} = \mathbf{i} \partial / \partial x + \mathbf{j} \partial / \partial y$ with x the zonal coordinate, and the three-dimensional gradient is $\nabla = \nabla_{\parallel} + \mathbf{k} \partial / \partial z$. Q_{Θ} is a generic diabatic source term.

Table 1 lists general characteristics of the near-tropical atmosphere (latitude $\phi =$

Gravitational acceleration	g	=	9.81	m s^{-2}
Coriolis parameter ($\phi = 30^\circ \text{ N}$)	f_0	=	$7.3 \cdot 10^{-5}$	s^{-1}
$(df/dy)_0$ ($\phi = 30^\circ \text{ N}$)	β	=	$2.0 \cdot 10^{-11}$	$\text{m}^{-1} \text{s}^{-1}$
Pressure	p_{ref}	=	10^5	Pa
Temperature	T_{ref}	=	300	K
Brunt-Väisälä frequency	N	=	10^{-2}	s^{-1}
Dry gas constant	R	=	287	$\text{m s}^{-2} \text{K}^{-1}$
Isentropic exponent	γ	=	1.4	

TABLE 1. Characteristic atmospheric flow parameters

Density	ρ_{ref}	=	$\frac{p_{\text{ref}}}{RT_{\text{ref}}}$	\sim	1.16	kg m^{-3}
Horizontal velocity	u_{ref}	=	$\frac{\tan \phi N^2}{\pi/2 f_0^2} \beta h_{\text{sc}}^2$	\sim	10	m s^{-1}
Vertical velocity	w_{ref}	=	$\frac{h_{\text{sc}}}{L_{\text{syn}}} u_{\text{ref}}$	\sim	0.1	m s^{-1}
Horizontal distance	L_{syn}	=	$\frac{N}{f_0} h_{\text{sc}}$	\sim	1200	km
Vertical distance	h_{sc}	=	$\frac{p_{\text{ref}}}{g \rho_{\text{ref}}}$	\sim	8.8	km
Time	t_{ref}	=	$\frac{L_{\text{syn}}}{u_{\text{ref}}}$	\sim	$1.2 \cdot 10^5$	s

TABLE 2. Derived reference values for non-dimensionalization of (2.1)

30° N) which we combine in Table 2 to obtain reference values for non-dimensionalization,

$$\begin{aligned}
 p^* &= \frac{p}{p_{\text{ref}}}, & \rho^* &= \frac{\rho}{\rho_{\text{ref}}}, & \mathbf{u}^* &= \frac{\mathbf{u}}{u_{\text{ref}}}, & w^* &= \frac{w}{u_{\text{ref}}} \frac{c_1 L_{\text{syn}}}{h_{\text{sc}}}, \\
 t^* &= \frac{t u_{\text{ref}}}{c_1 L_{\text{syn}}}, & \mathbf{x}^* &= \frac{\mathbf{x}}{c_1 L_{\text{syn}}}, & z^* &= \frac{z}{h_{\text{sc}}}.
 \end{aligned}
 \tag{2.2}$$

Here u_{ref} is an estimate of the thermal wind shear, $\mathbf{x} = \mathbf{i}x + \mathbf{j}y$, and c_1 is a constant chosen later to eliminate unessential free parameters from the dimensionless equations.

The Brunt-Väisälä frequency, $N = \sqrt{(g/\overline{\Theta})(d\overline{\Theta}/dz)}$, with the overbar denoting a horizontal average, characterizes the vertical variation, $\Delta\Theta \sim T_{\text{ref}}(N^2 h_{\text{sc}}/g)$, of the potential

temperature. An appropriate non-dimensionalization of Θ then reads

$$\Theta^* = \frac{\Theta}{T_{\text{ref}}} = 1 + c_2 \frac{\Delta\Theta}{T_{\text{ref}}} \left(\Theta_1(z^*) + \tilde{\Theta} \right), \quad (c_2 = \mathcal{O}(1) \text{ to be determined below}). \quad (2.3)$$

In deriving dimensionless equations, we introduce the Mach, Froude, Rossby, and Damköhler numbers, the aspect ratio, ℓ , and the dimensionless Coriolis β -parameter,

$$\begin{aligned} \text{M} &= \frac{u_{\text{ref}}}{\sqrt{RT_{\text{ref}}}} \approx 3.4 \cdot 10^{-2}, & \text{Fr} &= \frac{u_{\text{ref}}}{Nh_{\text{sc}}} \approx 1.1 \cdot 10^{-1}, & \beta^* &= \frac{\beta L_{\text{syn}}}{f_0}, \\ \text{Ro} &= \frac{u_{\text{ref}}}{f_0 L_{\text{syn}}} \approx 1.1 \cdot 10^{-1}, & \ell &= \frac{h_{\text{sc}}}{L_{\text{syn}}} \approx 7.3 \cdot 10^{-3}, & \text{Da} &= \frac{Q_{\Theta_{\text{ref}}}}{T_{\text{ref}} t_{\text{ref}}}. \end{aligned} \quad (2.4)$$

Note that $\beta^* \equiv (\pi/2) \text{Fr} / \tan \phi$. The first five of these parameters are then replaced through a set of distinguished asymptotic limits that are in line with Pedlosky's small Rossby number derivation of the QG model (see Pedlosky 1987; Klein 2010),

$$c_2 \frac{\Delta\Theta}{T_{\text{ref}}} = \varepsilon, \quad \frac{1}{\text{M}^2} = \frac{1}{\varepsilon^3}, \quad \frac{1}{\text{Ro}} = \frac{\hat{f}_0}{\varepsilon}, \quad \beta^* = \hat{\beta} \varepsilon, \quad \frac{c_1}{\ell} = \frac{1}{\varepsilon^2}, \quad (2.5)$$

where $(\hat{f}_0, \hat{\beta}) = \mathcal{O}(1)$ as $\varepsilon \rightarrow 0$. The constants c_1, c_2 are included in these distinguished limits in hindsight in such a way that they cancel from the dimensionless equations (2.6) below. For the scaling of the Damköhler number, Da, see (4.13) below.

Dropping the $*$ and $\hat{}$ indicators for brevity, we obtain the dimensionless equations

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \mathbf{u} + w \frac{\partial \mathbf{u}}{\partial z} + \frac{1}{\varepsilon^3} \frac{1}{\rho} \nabla_{\perp} p + \frac{1}{\varepsilon} (f_0 + \varepsilon \beta y) \mathbf{k} \times \mathbf{u} = 0, \quad (2.6a)$$

$$\frac{\partial w}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} w + w \frac{\partial w}{\partial z} + \frac{1}{\varepsilon^5} \frac{1}{\rho} \frac{\partial p}{\partial z} = -\frac{1}{\varepsilon^5}, \quad (2.6b)$$

$$\frac{\partial \rho}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \rho + w \frac{\partial \rho}{\partial z} + \rho \nabla_{\perp} \cdot \mathbf{u} + \rho \frac{\partial w}{\partial z} = 0, \quad (2.6c)$$

$$\frac{\partial \tilde{\Theta}}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \tilde{\Theta} + w \frac{\partial \tilde{\Theta}}{\partial z} + w \frac{d\Theta_1}{dz} = \frac{\text{Da}}{\varepsilon} Q_{\Theta}, \quad (2.6d)$$

$$\varepsilon \left(\tilde{\Theta} + \Theta_1 \right) = \frac{p^{\frac{1}{\gamma}}}{\rho} - 1. \quad (2.6e)$$

For the reader familiar with Klein's work, we remark that the atmospheric flow parameters in table 1 are essentially equivalent to those used in the general modelling framework in (Klein 2010), while the reference quantities for nondimensionalization in

table 2 have been constructed from these to fit the present application and streamline the subsequent developments.

3. Quasi-geostrophic farfield flow

This section describes the asymptotic expansions and the key results for the farfield including its separation into a smooth and a singular vortex induced part.

3.1. Expansion scheme for the quasi-geostrophic (QG) farfield flow

Here we rederive the QG theory following (Pedlosky 1987) except for one modification: While the classical expansions proceed in integer powers of ε , we allow for intermediate level perturbations of order $\mathcal{O}(\varepsilon^{i+1/2})$ for integers i . The latter appear because the synoptic and vortex core scales differ by $\mathcal{O}(\varepsilon^{1/2})$ (see section 4.1.1) and they are needed for consistently matched asymptotic solutions. For simplicity we neglect diabatic effects in the outer flow. The analysis then proceeds with the expansion scheme

$$p = p_0 + \varepsilon p_1 + \varepsilon^2 \left(p^{(2)} + p_2 \right) + \varepsilon^{5/2} \left(p^{(5/2)} + p_{5/2} \right) + \mathcal{O}(\varepsilon^{5/2}) \quad (3.1a)$$

$$\rho = \rho_0 + \varepsilon \rho_1 + \varepsilon^2 \left(\rho^{(2)} + \rho_2 \right) + \varepsilon^{5/2} \left(\rho^{(5/2)} + \rho_{5/2} \right) + \mathcal{O}(\varepsilon^{5/2}) \quad (3.1b)$$

$$\Theta = 1 + \varepsilon \Theta_1 + \varepsilon^2 \left(\Theta^{(2)} + \Theta_2 \right) + \varepsilon^{5/2} \left(\Theta^{(5/2)} + \Theta_{5/2} \right) + \mathcal{O}(\varepsilon^{5/2}) \quad (3.1c)$$

$$\mathbf{u} = \mathbf{u}^{(0)} + \varepsilon^{1/2} \mathbf{u}^{(1/2)} + \varepsilon \mathbf{u}^{(1)} + \varepsilon^{3/2} \mathbf{u}^{(3/2)} + \mathcal{O}(\varepsilon^{3/2}) \quad (3.1d)$$

$$w = \varepsilon w^{(1)} + \varepsilon^{3/2} w^{(3/2)} + \mathcal{O}(\varepsilon^{3/2}) \quad (3.1e)$$

where $(p_i, \rho_i, \Theta_i)(z)$ and $(\mathbf{u}^{(i)}, w^{(i)}, p^{(i)}, \rho^{(i)}, \Theta^{(i)})(t, \mathbf{x}, z)$ represent the mean background state and the flow variability, respectively.

From the scalings of the pressure gradient and Coriolis terms in the horizontal momentum balance from (2.6a) it follows that, under leading-order geostrophic balance, deviations of the pressure from the background state will arise first at order $\mathcal{O}(\varepsilon^2)$. The

vertical momentum equation in (2.6b) is dominated by hydrostatic balance up to at least five orders in ε and, as a consequence, deviations of density from the background state follow the pressure scaling and also start at $\mathcal{O}(\varepsilon^2)$. This explains the expansion schemes in (3.1a)–(3.1c). In (3.1e) we have used that the vertical velocity in a QG flow is by one order in the Rossby number smaller than expected on the basis of the aspect ratio scaling because the leading-order horizontal divergence vanishes under geostrophic balance (see, e.g., Pedlosky 1987).

To streamline the notation, we let

$$\delta \equiv \varepsilon^{1/2} \quad (3.2)$$

from here on and use the following abbreviations in the rest of this section,

$$\begin{pmatrix} \mathbf{u} \\ \tilde{\mathbf{u}} \\ w \\ \pi \\ \tilde{\Theta} \end{pmatrix} \equiv \begin{pmatrix} \mathbf{u}^{(0)} + \delta \mathbf{u}^{(1/2)} \\ \mathbf{u}^{(1)} + \delta \mathbf{u}^{(3/2)} \\ w^{(1)} + \delta w^{(3/2)} \\ (p^{(2)} + \delta p^{(5/2)}) / \rho_0 \\ \Theta^{(2)} + \delta \Theta^{(5/2)} \end{pmatrix}. \quad (3.3)$$

3.2. Quasi-geostrophic (QG) theory

Using (3.3), the combined leading and first order results in δ from the momentum equations, the mass balance, and the potential temperature transport equation read

Geostrophic Balance

$$f_0 \mathbf{k} \times \mathbf{u} + \nabla_{\parallel} \pi = 0, \quad (3.4a)$$

Hydrostatic Balance

$$\frac{\partial \pi}{\partial z} = \tilde{\Theta}, \quad (3.4b)$$

Anelastic Constraint

$$\rho_0 \nabla_{\parallel} \cdot \tilde{\mathbf{u}} + \frac{\partial}{\partial z} (\rho_0 w) = 0, \quad (3.4c)$$

Potential Temperature Transport

$$\frac{\partial \tilde{\Theta}}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \tilde{\Theta} + w \frac{d\Theta_1}{dz} = 0. \quad (3.4d)$$

For later reference, we conclude from (3.4a) that

$$\nabla_{\parallel} \cdot \mathbf{u} = 0 \quad \text{and} \quad \mathbf{u} = -\mathbf{k} \times \nabla_{\parallel} \psi, \quad (3.5)$$

where the streamfunction, ψ , for the horizontal flow is defined as

$$\psi \equiv -\pi/f_0. \quad (3.6)$$

An additional equation for the divergence of the perturbation velocity, $\nabla_{\parallel} \cdot \tilde{\mathbf{u}}$, appearing in (3.4c) is obtained from the curl of the next-order horizontal momentum equation,

Vorticity Transport

$$\left(\frac{\partial}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \right) (\zeta + \beta y) + f_0 \nabla_{\parallel} \cdot \tilde{\mathbf{u}} = 0, \quad (3.7)$$

with the relative vertical vorticity

$$\zeta = \mathbf{k} \cdot (\nabla_{\parallel} \times \mathbf{u}). \quad (3.8)$$

Eqs. (3.4)–(3.8) are the QG model for the unknowns $(\mathbf{u}, w, \pi, \tilde{\Theta}, [\nabla_{\parallel} \cdot \tilde{\mathbf{u}}])(t, \mathbf{x}, z)$ defined in (3.3) given the background state through $\rho_0(z)$ and $(d\Theta_1/dz)(z)$. The essence of the system is revealed through its classical formulation involving an advection equation

$$\left(\frac{\partial}{\partial t} + \mathbf{u} \cdot \nabla_{\parallel} \right) q = 0 \quad (3.9)$$

for the QG potential vorticity

$$q = \zeta + \beta y - \frac{f_0}{\rho_0} \frac{\partial}{\partial z} \left(\frac{\rho_0}{\Theta'_1} \tilde{\Theta} \right) \quad \text{where} \quad \Theta'_1 = \frac{d\Theta_1}{dz}. \quad (3.10)$$

To verify these equations, one eliminates $[\nabla_{\parallel} \cdot \tilde{\mathbf{u}}]$ from (3.7) using (3.4c), and then eliminates w from the remaining equation using (3.4d).

Using that $\zeta = -\nabla_{\parallel}^2 \psi$ due to (3.8), (3.6), (3.5), and the geostrophic balance from (3.4a), eq. (3.10) becomes an elliptic equation for ψ given the potential vorticity, viz.

$$(\nabla_{\parallel}^2 + \mathcal{L}_z) \psi = -q_{\text{rel}}, \quad (3.11)$$

where

$$\mathcal{L}_z[\cdot] = \frac{f_0^2}{\rho_0} \frac{\partial}{\partial z} \left(\frac{\rho_0}{\Theta_1'} \frac{\partial \cdot}{\partial z} \right) \quad \text{and} \quad q_{\text{rel}} = q - \beta y. \quad (3.12)$$

3.3. Singular vortex theory

3.3.1. Regular–singular farfield flow decomposition

Consider the relative potential vorticity, q_{rel} , to be composed of a line singularity, q_s , located on the vortex centreline, and a regular contribution, q_r , so that

$$q_{\text{rel}} = q_s + q_r, \quad (3.13)$$

where

$$\begin{aligned} q_s(t, \mathbf{x}, z) &= \int_0^H \Gamma(t, z') \delta_3 \left([\mathbf{x} + \mathbf{k}z] - [\mathbf{X}(t, z') + \mathbf{k}z'] \right) dz' \\ &= \Gamma(t, z) \delta_2(\mathbf{x} - \mathbf{X}(t, z)) \end{aligned} \quad (3.14)$$

Here $\delta_k(\cdot)$ denote the k -dimensional delta distributions, and

$$\Gamma(t, z) = \Gamma^{(0)}(t, z) + \delta \Gamma^{(1)}(t, z) + \mathcal{O}(\delta), \quad (3.15a)$$

$$\mathbf{X}(t, z) = \mathbf{X}_0(t) + \delta \mathbf{X}^{(1)}(t, z) + \mathcal{O}(\delta), \quad (3.15b)$$

are the total vortex circulation and the horizontal position of its centreline, respectively.

To construct a perturbation expansion of ψ that explicitly represents the local QG flow response to the vortex singularity we first rewrite (3.11) as

$$(\nabla_{\parallel}^2 + \mathcal{L}_z) \psi = -q_s - q_r. \quad (3.16)$$

Next, introducing local polar coordinates via

$$\mathbf{x} - \mathbf{X}(t, z) = r \mathbf{e}_r = r (\mathbf{i} \cos \theta + \mathbf{j} \sin \theta), \quad (3.17)$$

with \mathbf{e}_r a radial, and \mathbf{i}, \mathbf{j} cartesian horizontal unit vectors, and observing that the quasi two-dimensional potential vortex streamfunction

$$\psi_0^{2d}(t, r, z) = -\frac{\Gamma(t, z)}{2\pi} \ln r \quad \text{solves} \quad \nabla_{\parallel}^2 \psi_0^{2d} = -q_s \quad (3.18)$$

we let, in a first step,

$$\psi(t, \mathbf{x}, z) = \psi_0^{2d}(t, r, z) + \psi_1^*(t, \mathbf{x}, z). \quad (3.19)$$

Then, ψ_1^* obviously satisfies

$$(\nabla_n^2 + \mathcal{L}_z) \psi_1^* = -\mathcal{L}_z \psi_0^{2d} - q_r. \quad (3.20)$$

Taking into account the expansion of the centreline in (3.15b) and the coordinate transformation in (3.17) we express the first term on the right of this equation as

$$-\mathcal{L}_z \psi_0^{2d} = 4M_0(t, z) \ln r + \frac{\delta}{r} 2M_1(t, \theta, z) + \frac{\delta^2}{r^2} 4M_2(t, \theta, z), \quad (3.21)$$

where

$$M_0 = \frac{1}{8\pi} \frac{f_0^2}{\rho_0} \frac{\partial}{\partial z} \left(\frac{\rho_0}{\Theta_1'} \frac{\partial \Gamma}{\partial z} \right), \quad (3.22a)$$

$$M_1 = -\frac{f_0^2}{4\pi\rho_0\Gamma} \frac{\partial}{\partial z} \left(\frac{\rho_0\Gamma^2}{\Theta_1'} \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right) \cdot \mathbf{e}_r \equiv -\mathbf{M}_1 \cdot \mathbf{e}_r, \quad (3.22b)$$

$$M_2 = \frac{f_0^2\Gamma}{8\pi\Theta_1'} \left(\left[\mathbf{e}_\theta \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right]^2 - \left[\mathbf{e}_r \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right]^2 \right). \quad (3.22c)$$

The source term in (3.20) is still singular, yet the singularity is much weaker than the original Dirac distribution in (3.14). Also, the singular part of this source term as given in (3.21) involves the first and second circumferential Fourier modes through the terms $M_1(t, \theta, z)$ and $M_2(t, \theta, z)$, respectively. For later reference we extract again the dominant singular behavior of the stream function by introducing

$$\psi_1^{2d} = M_0 r^2 (\ln r - 1) + \delta M_1 r \ln r - \delta^2 M_2 \quad (3.23)$$

which solves

$$\nabla_n^2 \psi_1^{2d} = -\mathcal{L}_z \psi_0^{2d}(t, r, z). \quad (3.24)$$

Collecting results we have

$$\psi = -\frac{\Gamma}{2\pi} \ln r + M_0 r^2 (\ln r - 1) + \delta M_1 r \ln r - \delta^2 M_2 + \psi_2^* \quad (3.25)$$

where the regularized stream function, ψ_2^* , satisfies

$$(\nabla_0^2 + \mathcal{L}_z) \psi_2^* = -\mathcal{L}_z \psi_1^{2d} - q_r. \quad (3.26)$$

With (3.25) we have an explicit representation of the singular behavior of the stream-function near the vortex centreline ($r \rightarrow 0$), including its expansion in terms of δ . In App. B we discuss the behavior of ψ_2^* at $r = 0$ and *inter alia* show that it is continuous.

4. Vortex Core Structure

4.1. Expansion scheme for the meso-scale vortex core flow in gradient wind balance

4.1.1. Pressure, velocity, and spacial scaling

If a given total vortex circulation, Γ , is realized on the synoptic scale with velocities of order u_{ref} , i.e., $\Gamma = \mathcal{O}(2\pi L_{\text{syn}} u_{\text{ref}})$, then it is realized with wind speeds of order u_{ref}/δ on the smaller meso scale $L_{\text{mes}} = \mathcal{O}(\delta L_{\text{syn}})$. At the same time, if the synoptic scale Rossby number is $\text{Ro} = u_{\text{ref}}/f_0 L_{\text{syn}} = \mathcal{O}(\delta^2)$ (see (2.5)) then the vortex Rossby number becomes $\text{Ro}_{\text{mes}} = u_{\text{mes}}/f_0 L_{\text{mes}} = \mathcal{O}(1)$ and the vortex is in approximate gradient wind balance (see section 1.2). This sets the length, velocity, and pressure scales for the vortex core,

$$L_{\text{mes}} = \delta L_{\text{syn}}, \quad \delta p_{\text{mes}} = \mathcal{O}(\delta^4 p_{\text{ref}}) \equiv \mathcal{O}(\varepsilon^2 p_{\text{ref}}), \quad u_{\text{mes}} = \mathcal{O}(\delta^{-1} u_{\text{ref}}). \quad (4.1)$$

Given that $u_{\text{ref}} \sim 10 \text{ ms}^{-1}$ (see Table 2) and $\varepsilon \sim 0.1$ or $\delta = \sqrt{\varepsilon} \sim 1/3$ (see (2.4), (2.5)) these scalings are compatible with tropical storms and nascent hurricanes, which feature velocities of $\mathcal{O}(30 \text{ ms}^{-1})$. Motivated by changes of storm intensities being observed over a day or two ($t_{\text{ref}} \sim 30 \text{ h}$ in Table 2), we assume the core structure to evolve on the synoptic time scale, i.e., on the same time scale as the quasi-geostrophic outer flow.

4.1.2. Co-moving coordinate system

For the meso scale analysis we introduce vortex centred horizontal coordinates so that

$$\mathbf{x} = \mathbf{X}(t, z) + \delta \hat{\mathbf{x}} = \mathbf{X}_0(t) + \delta \left(\mathbf{X}^{(1)}(t, z) + \hat{\mathbf{x}} \right), \quad (4.2)$$

The relative horizontal offset, $\hat{\mathbf{x}}$, resolves the vortex core diameter scale as estimated in (4.1). To efficiently deal with the near-axisymmetry of the vortex we use the polar coordinates from (3.17) with an appropriate rescaling of the radius, i.e.,

$$\hat{\mathbf{x}} = \hat{x} \mathbf{i} + \hat{y} \mathbf{j} \quad \text{where} \quad \begin{cases} \hat{x} = \hat{r} \cos \theta; & \mathbf{i} = \mathbf{e}_r \cos \theta - \mathbf{e}_\theta \sin \theta \\ \hat{y} = \hat{r} \sin \theta; & \mathbf{j} = \mathbf{e}_r \sin \theta + \mathbf{e}_\theta \cos \theta \end{cases} \quad (4.3)$$

with \mathbf{e}_r and \mathbf{e}_θ the radial and circumferential unit vectors, respectively. For later reference, here is the polar coordinate representation of the vortex tilt,

$$\mathbf{X}^{(1)} = \left(X^{(1)} \cos \theta + Y^{(1)} \sin \theta \right) \mathbf{e}_r + \left(-X^{(1)} \sin \theta + Y^{(1)} \cos \theta \right) \mathbf{e}_\theta. \quad (4.4)$$

The horizontal velocity is decomposed into the vortex' motion plus the relative velocity,

$$\mathbf{u} = \frac{\partial \mathbf{X}}{\partial t} + (u_r \mathbf{e}_r + u_\theta \mathbf{e}_\theta), \quad (4.5)$$

and the transformation rules for derivatives now read

$$\nabla_{\parallel} = \delta^{-1} \left(\mathbf{e}_r \frac{\partial}{\partial \hat{r}} + \mathbf{e}_\theta \frac{1}{\hat{r}} \frac{\partial}{\partial \theta} \right) \equiv \delta^{-1} \hat{\nabla}, \quad (4.6)$$

$$\frac{\partial}{\partial z} \Big|_{t,x,y} = \frac{\partial}{\partial z} \Big|_{t,\hat{r},\theta} - \delta^{-1} \frac{\partial \mathbf{X}}{\partial z} \cdot \hat{\nabla}, \quad (4.7)$$

$$\frac{\partial}{\partial t} \Big|_{x,y,z} = \frac{\partial}{\partial t} \Big|_{\hat{r},\theta,z} - \delta^{-1} \mathbf{V} \cdot \hat{\nabla}. \quad (4.8)$$

For later reference here is our notation for the Fourier expansion of functions of θ ,

$$F(\theta) = F_0 + \sum_n (F_{n1} \sin(n\theta) + F_{n2} \cos(n\theta)) \quad (4.9)$$

4.1.3. *Vortex core expansion scheme*

Up to order $\mathcal{O}(1)$ the circumferential velocity is assumed to be axisymmetric, so that

$$u_\theta(t, \mathbf{x}, z; \varepsilon) = \delta^{-1} u_\theta^{(0)}(t, \hat{r}, z) + u_\theta^{(1)}(t, \hat{r}, z) + \delta u_\theta^{(2)}(t, \hat{r}, \theta, z) + \mathcal{O}(\delta), \quad (4.10a)$$

$$u_r(t, \mathbf{x}, z; \varepsilon) = \delta u_r^{(2)}(t, \hat{r}, \theta, z) + \mathcal{O}(\delta). \quad (4.10b)$$

That the first-order velocity must be axisymmetric can be derived instead of having to be assumed (Mikusky 2007) (E. Päsche's maiden name). Asymmetries of order $\mathcal{O}(\delta u_{\text{ref}})$ in the horizontal velocity distributed over the core size length scale, L_{mes} , induce horizontal divergences of order $\delta u_{\text{ref}}/L_{\text{mes}} = u_{\text{ref}}/L_{\text{syn}}$, see (4.1). The flow field is essentially anelastic as we will see below, and from the anelastic divergence constraint follow dimensionless vertical velocities of order $w/u_{\text{ref}} = \mathcal{O}(h_{\text{sc}}/L_{\text{syn}})$, and this implies, given the vertical velocity scaling in (2.2),

$$w(t, \mathbf{x}, z; \varepsilon) = \hat{w}^{(1)}(t, \hat{r}, \theta, z) + \mathcal{O}(1). \quad (4.11)$$

Expansions for the thermodynamic variables are anticipated as follows: The estimates in (4.1) settle the lowest order at which horizontal pressure variations can occur. All lower order pressure terms will be horizontally homogeneous. Moreover, they will have to match the corresponding terms in the outer QG solution since any differences would be equilibrated rapidly by fast internal waves. The outer pressure expansion does not include half-order scalings below $\varepsilon^2 \equiv \delta^4$, and this is why such terms must also be absent in the core. The pressure perturbations turn out to be hydrostatic in the regime considered so that half order scalings cannot appear in the Θ and ρ expansion at the respective orders

either. These considerations lead to

$$p = p_0 + \delta^2 p_1 + \delta^4 \left(\widehat{p}^{(4)} + \widehat{p}_4 \right) + \delta^5 \left(\widehat{p}^{(5)} + \widehat{p}_5 \right) + \mathcal{O}(\delta^5), \quad (4.12a)$$

$$\rho = \rho_0 + \delta^2 \rho_1 + \delta^4 \left(\widehat{\rho}^{(4)} + \widehat{\rho}_4 \right) + \delta^5 \left(\widehat{\rho}^{(5)} + \widehat{\rho}_5 \right) + \mathcal{O}(\delta^5), \quad (4.12b)$$

$$\Theta = 1 + \delta^2 \Theta_1 + \delta^4 \left(\widehat{\Theta}^{(4)} + \widehat{\Theta}_4 \right) + \delta^5 \left(\widehat{\Theta}^{(5)} + \widehat{\Theta}_5 \right) + \mathcal{O}(\delta^5). \quad (4.12c)$$

Here $(p_0, p_1, \rho_0, \rho_1, \Theta_1)$ represent the background state as in (3.1), $(\widehat{p}_i, \widehat{\rho}_i, \widehat{\Theta}_i)(t, z)$, are higher-order horizontal mean values, and $(\widehat{p}^{(i)}, \widehat{\rho}^{(i)}, \widehat{\Theta}^{(i)})(t, \widehat{r}, \theta, z)$ are the perturbation variables of interest in this section. With the diabatic source terms scaling as

$$\frac{\text{Da}}{\varepsilon} Q_\Theta = \delta^2 Q_\Theta^{(2)} + \mathcal{O}(\delta^5), \quad (4.13)$$

they will modify the adiabatic dynamics but will not change the balances fundamentally (see discussion below (4.19)). This also settles the asymptotic scaling of the Damköhler number as announced in section 2 (below (2.5)).

4.2. Asymptotic equation hierarchy for the vortex core

The governing equations transformed to the co-moving coordinates are provided in App. A. Inserting the vortex core expansion scheme from section 4.1.3 we obtain

horizontal momentum balance

$\mathcal{O}(\delta^{-3})$:

$$-\frac{(u_\theta^{(0)})^2}{\widehat{r}} + \frac{1}{\rho_0} \frac{\partial \widehat{p}^{(4)}}{\partial \widehat{r}} - f_0 u_\theta^{(0)} = 0 \quad (4.14a)$$

$$\frac{\partial \widehat{p}^{(4)}}{\partial \theta} = 0 \quad (4.14b)$$

$\mathcal{O}(\delta^{-2})$:

$$-\frac{2u_\theta^{(0)} u_\theta^{(1)}}{\widehat{r}} + \frac{1}{\rho_0} \frac{\partial \widehat{p}^{(5)}}{\partial \widehat{r}} - f_0 \left(u_\theta^{(1)} + \mathbf{e}_\theta \cdot \frac{d\mathbf{X}_0}{dt} \right) = 0 \quad (4.15a)$$

$$\frac{1}{\rho_0 \widehat{r}} \frac{\partial \widehat{p}^{(5)}}{\partial \theta} + f_0 \mathbf{e}_r \cdot \frac{d\mathbf{X}_0}{dt} = 0 \quad (4.15b)$$

$\mathcal{O}(\delta^{-1})$:

$$\begin{aligned} & \frac{u_\theta^{(0)}}{\hat{r}} \frac{\partial u_r^{(2)}}{\partial \theta} - \frac{2u_\theta^{(0)} u_\theta^{(2)}}{\hat{r}} - \frac{(u_\theta^{(1)})^2}{\hat{r}} + \mathbf{e}_\theta \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \frac{\hat{w}^{(1)} u_\theta^{(0)}}{\hat{r}} \\ & + \frac{1}{\rho_0} \frac{\partial \hat{p}^{(6)}}{\partial \hat{r}} - \frac{\rho_1}{\rho_0^2} \frac{\partial \hat{p}^{(4)}}{\partial \hat{r}} - f_0 \left(u_\theta^{(2)} + \mathbf{e}_\theta \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial t} \right) - \beta Y_C^{(0)} u_\theta^{(0)} = 0 \end{aligned} \quad (4.16a)$$

$$\begin{aligned} & \frac{\partial u_\theta^{(0)}}{\partial t} + \hat{w}^{(1)} \frac{\partial u_\theta^{(0)}}{\partial z} + u_r^{(2)} \left(\frac{\partial u_\theta^{(0)}}{\partial \hat{r}} + \frac{u_\theta^{(0)}}{\hat{r}} \right) + \frac{u_\theta^{(0)}}{\hat{r}} \frac{\partial u_\theta^{(2)}}{\partial \theta} \\ & - \hat{w}^{(1)} \mathbf{e}_r \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \frac{\partial u_\theta^{(0)}}{\partial \hat{r}} + \frac{1}{\rho_0 \hat{r}} \frac{\partial \hat{p}^{(6)}}{\partial \theta} + f_0 \left(u_r^{(2)} + \mathbf{e}_r \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial t} \right) = 0 \end{aligned} \quad (4.16b)$$

The flow is hydrostatic to very high order with the first non-trivial vertical balance

vertical momentum balance

 $\mathcal{O}(\delta^{-6})$:

$$\frac{\partial \hat{p}^{(4)}}{\partial z} - \frac{\partial \mathbf{X}^{(1)}}{\partial z} \cdot \hat{\nabla}_\parallel \hat{p}^{(4)} = -\hat{\rho}^{(4)}. \quad (4.17)$$

The leading two contributions to the flow velocity are horizontal and axisymmetric according to (4.10), (4.11), and therefore divergence free. For the second order velocity field we find an anelastic divergence constraint that is modified by the vortex' tilt,

mass continuity

 $\mathcal{O}(1)$:

$$\frac{\rho_0}{\hat{r}} \left(\frac{\partial}{\partial \hat{r}} \left(\hat{r} u_r^{(2)} \right) + \frac{\partial u_\theta^{(2)}}{\partial \theta} \right) + \frac{\partial}{\partial z} \left(\rho_0 \hat{w}^{(1)} \right) - \frac{\partial \mathbf{X}^{(1)}}{\partial z} \cdot \hat{\nabla}_\parallel (\rho_0 \hat{w}^{(1)}) = 0. \quad (4.18)$$

The first non-trivial potential temperature transport equation reads

potential temperature

 $\mathcal{O}(\delta^2)$:

$$\frac{u_\theta^{(0)}}{\hat{r}} \frac{\partial \hat{\Theta}^{(4)}}{\partial \theta} + \hat{w}^{(1)} \frac{d\Theta_1}{dz} = Q_\Theta^{(2)}. \quad (4.19)$$

For any $\hat{w}^{(i)}$ with $i < 1$ we would have obtained $w^{(i)} d\Theta_1/dz = Q_\Theta^{(i+1)}$, so that these orders of w would be induced solely by diabatic sources. As a consequence, the lowest-order vertical velocity that remains in the diabatic case is $\hat{w}^{(1)}$ as anticipated in the expansion scheme (4.11). Stronger diabatic effects will be considered in future work.

The equation of state relates the perturbations of the thermodynamic variables through equation of state

$\mathcal{O}(\delta^4)$:

$$\hat{\rho}^{(4)} = \rho_0 \left(\frac{\hat{p}^{(4)}}{\gamma p_0} - \hat{\Theta}^{(4)} \right). \quad (4.20)$$

4.3. Temporal evolution of the vortex structure

The goal here is to derive from (4.14)–(4.20) an evolution equation for the leading order circumferential velocity, $u_\theta^{(0)}$, given the vortex tilt, $\partial \mathbf{X}^{(1)}/\partial z$, and the diabatic source term, $Q_\Theta^{(2)}$. Consider to this end the second order circumferential momentum balance from (4.16b). Averaging in θ taking into account that $u_\theta^{(0)}$ is axisymmetric we obtain

$$\frac{\partial u_\theta^{(0)}}{\partial t} + \hat{w}_0^{(1)} \frac{\partial u_\theta^{(0)}}{\partial z} + u_{r,0}^{(2)} \left(\frac{\partial u_\theta^{(0)}}{\partial \hat{r}} + \frac{u_\theta^{(0)}}{\hat{r}} + f_0 \right) - u_{r,*}^{(2)} \frac{\partial u_\theta^{(0)}}{\partial \hat{r}} = 0 \quad (4.21)$$

where, using Fourier decomposition, (4.9),

$$u_{r,*}^{(2)} = \left(\hat{w}^{(1)} \mathbf{e}_r \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right)_0 = \frac{1}{2} \left[\hat{w}_{12} \frac{\partial X^{(1)}}{\partial z} + \hat{w}_{11} \frac{\partial Y^{(1)}}{\partial z} \right]. \quad (4.22)$$

Expressions for $\hat{w}_0^{(1)}$ and $\hat{w}_{11}, \hat{w}_{12}$ follow from the Fourier decomposition of the potential temperature transport equation in (4.19), viz.

$$\hat{w}_0^{(1)} \frac{d\Theta_1}{dz} = Q_{\Theta,0}^{(2)}, \quad (4.23a)$$

$$\hat{w}_{1k}^{(1)} \frac{d\Theta_1}{dz} = Q_{\Theta,1k}^{(2)} + (-1)^k \frac{u_\theta^{(0)}}{\hat{r}} \hat{\Theta}_{1k^*}^{(4)} \quad \left(k \in \{1, 2\}, \quad k^* = 3 - k \right). \quad (4.23b)$$

Now we have $\hat{\Theta}_{1k}^{(4)} = -\hat{\rho}_{1k}^{(4)}/\rho_0$ from the equation of state, (4.20), because $\hat{p}^{(4)}$ is axisymmetric according to (4.14b), and hence $\hat{p}_{1k}^{(4)} \equiv 0$. By extracting the first Fourier modes of the vertical momentum balance in (4.17) we find

$$\hat{\Theta}_{11}^{(4)} = -\frac{\hat{\rho}_{11}^{(4)}}{\rho_0} = -\frac{1}{\rho_0} \frac{\partial Y^{(1)}}{\partial z} \frac{\partial \hat{p}^{(4)}}{\partial \hat{r}}, \quad \hat{\Theta}_{12}^{(4)} = -\frac{\hat{\rho}_{12}^{(4)}}{\rho_0} = -\frac{1}{\rho_0} \frac{\partial X^{(1)}}{\partial z} \frac{\partial \hat{p}^{(4)}}{\partial \hat{r}}. \quad (4.24)$$

Using the gradient wind balance, (4.14a), to eliminate $\partial\widehat{p}^{(4)}/\partial\widehat{r}$, and going back to (4.23b)

we obtain the desired expressions for the $\widehat{w}_{1k}^{(1)}$ in terms of $u_\theta^{(0)}$, $\partial\mathbf{X}^{(1)}/\partial z$, and $Q_\Theta^{(2)}$,

$$\widehat{w}_{1k}^{(1)} = \frac{1}{\Theta'_1} \left[Q_{\Theta,1k}^{(4)} + (-1)^k \frac{\partial X_k^{(1)}}{\partial z} \frac{u_\theta^{(0)}}{\widehat{r}} \left(\frac{(u_\theta^{(0)})^2}{\widehat{r}} + f_0 u_\theta^{(0)} \right) \right] \quad (k = 1, 2), \quad (4.25)$$

where $X_1^{(1)} \equiv X^{(1)}$ and $X_2^{(1)} \equiv Y^{(1)}$. Inserted into (4.22) we find the expression for $u_{r,*}^{(2)}$

(called $u_{r,*}$ there for simplicity of notation) announced in (1.3d).

To find a corresponding expression for $u_{r,0}^{(2)}$ (see the third term in (4.21)), consider the circumferential average of mass continuity, (4.18). A brief calculation yields first

$$\frac{\partial(\widehat{r} u_{r,0}^{(2)})}{\partial\widehat{r}} = -\frac{\widehat{r}}{\rho_0} \frac{\partial(\rho_0 \widehat{w}_0^{(1)})}{\partial z} + \frac{1}{2} \left[\frac{\partial X^{(1)}}{\partial z} \frac{\partial(\widehat{r} \widehat{w}_{12}^{(1)})}{\partial\widehat{r}} + \frac{\partial Y^{(1)}}{\partial z} \frac{\partial(\widehat{r} \widehat{w}_{11}^{(1)})}{\partial\widehat{r}} \right]. \quad (4.26)$$

Then, exploiting (4.25), we integrate in \widehat{r} requiring that $u_{r,0}^{(2)}$ be finite at $r = 0$ to find

$$u_{r,0}^{(2)} = -\frac{1}{\widehat{r}} \int_0^{\widehat{r}} \frac{r'}{\rho_0} \frac{\partial}{\partial z} \left(\rho_0 \frac{Q_0^{(3)}}{\Theta'_1} \right) dr' + u_{r,*}^{(2)} \equiv u_{r,00}^{(2)} + u_{r,*}^{(2)}. \quad (4.27)$$

With (4.21)–(4.27) now yield the core structure evolution equation announced in (1.3a).

4.4. First Fourier modes of the second-order horizontal flow

Here we study the second order horizontal momentum and mass balances in (4.16) and (4.18) to derive expressions for the first θ -Fourier modes of $u_r^{(2)}$, $u_\theta^{(2)}$. These will be of particular interest in matching the inner and outer solutions in section 5 as their large- \widehat{r} behaviour determines the evolution of the vortex centerline perturbation, $\mathbf{X}^{(1)}$.

The non-symmetric components of $u_r^{(2)}$ and $u_\theta^{(2)}$ are represented by a velocity potential, $\phi^{(2)}$, (divergent part) and a streamfunction, $\psi^{(2)}$, (non divergent part) so that

$$u_r^{(2)} = \frac{1}{\widehat{r}} \frac{\partial\psi^{(2)}}{\partial\theta} + \frac{\partial\phi^{(2)}}{\partial\widehat{r}}, \quad u_\theta^{(2)} = -\frac{\partial\psi^{(2)}}{\partial\widehat{r}} + \frac{1}{\widehat{r}} \frac{\partial\phi^{(2)}}{\partial\theta}. \quad (4.28)$$

After projection onto the first Fourier modes, mass continuity (4.18) yields,

$$\left(\frac{\partial^2}{\partial\widehat{r}^2} + \frac{1}{\widehat{r}} \frac{\partial}{\partial\widehat{r}} - \frac{1}{\widehat{r}^2} \right) \phi_{1k}^{(2)} = -\frac{1}{\rho_0} \frac{\partial(\rho_0 \widehat{w}_{1k}^{(1)})}{\partial z} + \frac{1}{2} \left(\mathbf{e}_r \cdot \frac{\partial\mathbf{X}^{(1)}}{\partial z} \right)_k \frac{\partial\widehat{w}_0^{(1)}}{\partial\widehat{r}} \equiv R_{1k}. \quad (4.29)$$

Knowing that the quasi-geostrophic outer flow is horizontally non divergent we let

$$\phi_{1k}^{(2)} = -\hat{r} \int_{\hat{r}}^{\infty} \frac{1}{\bar{r}^3} \left[\int_0^{\bar{r}} \bar{r}^2 R_{1k} d\bar{r} \right] d\bar{r}. \quad (4.30)$$

The integral in (4.30) converges and yields a vanishing divergent part of the second order flow as $\hat{r} \rightarrow \infty$ provided $\hat{w}_0^{(1)}$ and $\hat{w}_{1k}^{(1)}$ decay sufficiently rapidly for large \hat{r} . Considering (4.23a) and (4.25), this is the case if (i) $Q_{\Theta,0}^{(4)}$ and $Q_{\Theta,1k}^{(4)}$ decay sufficiently rapidly, and (ii) $u_{\theta}^{(0)} = \Gamma/(2\pi\hat{r}) (1 + o(1))$ as $\hat{r} \rightarrow \infty$. For the diabatic source terms we take the liberty to assume such rapid decay, while the required large- \hat{r} behaviour of $u_{\theta}^{(0)}$ follows from our assumption that the vortex has a well-defined total circulation, $\Gamma(t, z)$.

With the divergent part of the flow vanishing for large \hat{r} , the solutions for the streamfunction $\psi^{(2)}$ will be central in determining the evolution of the vortex tilt. Elimination of $\hat{p}^{(4)}$ and $\hat{p}^{(6)}$ from (4.16a) and (4.16b) by cross-differentiation, extraction of the first Fourier modes from the resulting equations, and inserting (4.28) yields

$$\left(\frac{\partial^2}{\partial \hat{r}^2} + \frac{1}{\hat{r}} \frac{\partial}{\partial \hat{r}} - \left[\frac{\zeta_{\hat{r}}^{(0)}}{u_{\theta}^{(0)}} + \frac{1}{\hat{r}^2} \right] \right) \psi_{1k}^{(2)} = \frac{(-1)^{k^*}}{u_{\theta}^{(0)}} \left(\mathcal{H} + \mathcal{I} + \mathcal{J} + \mathcal{Q} \right)_{1k^*}, \quad (4.31)$$

where $k^* = 3 - k$,

$$\zeta^{(0)} = \frac{\partial u_{\theta}^{(0)}}{\partial \hat{r}} + \frac{u_{\theta}^{(0)}}{\hat{r}}, \quad (4.32)$$

is the vorticity of the leading-order axisymmetric circumferential flow, and

$$\mathcal{H}_{1k} = \frac{\partial}{\partial \hat{r}} \left(\hat{r} \hat{w}_{1k}^{(1)} \frac{\partial u_{\theta}^{(0)}}{\partial z} \right), \quad (4.33a)$$

$$\mathcal{I}_{1k} = -\hat{r} [\zeta^{(0)} + f_0] \frac{1}{\rho_0} \frac{\partial (\rho_0 \hat{w}_{1k}^{(1)})}{\partial z}, \quad (4.33b)$$

$$\mathcal{J}_{1k} = \frac{\partial \phi_{1k}^{(2)}}{\partial \hat{r}} \left(\hat{r} \frac{\partial \zeta^{(0)}}{\partial \hat{r}} \right), \quad (4.33c)$$

$$\mathcal{Q}_{1k} = \left(\mathbf{e}_r \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right)_k \left(\hat{w}_0^{(1)} \frac{u_{\theta}^{(0)}}{\hat{r}} - \frac{\partial}{\partial r} \left(\hat{r} \hat{w}_0^{(1)} \frac{\partial u_{\theta}^{(0)}}{\partial r} \right) \right). \quad (4.33d)$$

Particular solutions with homogeneous boundary conditions

$$\psi_{1k}^{(2)} = \frac{\partial \psi_{1k}^{(2)}}{\partial \hat{r}} = 0 \quad \text{at} \quad \hat{r} = 0 \quad (4.34)$$

read

$$\psi_{1k}^{(2)} = (-1)^{k*} \hat{r} \left[\frac{u}{\hat{r}} \int_0^{\hat{r}} \frac{1}{\bar{r} u^2} \int_0^{\bar{r}} r' [\mathcal{H} + \mathcal{I} + \mathcal{J} + \mathcal{Q}]_{1k*}(r') dr' d\bar{r} \right]. \quad (4.35)$$

This is also the final solution because one of the two linearly independent homogeneous solutions produces singular velocities near \hat{r} , the other represents a singular dipole flow for large \hat{r} which is incompatible with the outer QG flow (see App. C.1 for details).

We have restricted this section to those results from the core structure analysis that are needed to derive our main results, i.e., the core structure evolution equation (see section 4.3), and the evolution equation for the vortex tilt (see section 5 below). For more details of the core structure solution, such as the higher harmonics of $\hat{p}^{(4)}, \hat{\rho}^{(4)}, \hat{\Theta}^{(4)}, \hat{w}^{(1)}$, an evaluation of the first order horizontal momentum balances in (4.15), and the higher harmonics of $\phi^{(2)}$ and $\psi^{(2)}$ see Appendix C.

5. Matching and the Vortex Centreline Velocity

5.1. Eckhaus' notation for the matching procedure

Here we employ van Dyke's procedure for the asymptotic matching of the farfield and vortex core solutions (van Dyke 1964) using Eckhaus' notation (Eckhaus 1979): Let $\Phi_\delta(x)$ be defined for $x \in D \subset \mathbb{R}^n$ and $0 < \delta \ll 1$. When Φ_δ solves a boundary layer type problem, there is one subdomain $D_0 \subset D$ within which Φ_δ has a regular expansion

$$\left. \begin{aligned} \Phi_\delta(x) &= \sum_{n=0}^m \delta^n \Phi^{(n)}(x) + \mathcal{O}(\delta^m) \\ &\equiv \left(\mathbf{E}^{(m)} \Phi_\delta \right) (x) + \mathcal{O}(\delta^m) \end{aligned} \right\} \quad \text{for} \quad x \in D_0, \quad (5.1)$$

and another subset $S_\delta \subset D$ such that Φ_δ when expressed in boundary layer coordinates,

$$\Phi_\delta(x) \equiv \hat{\Phi}_\delta(\hat{x}) \quad \text{where} \quad \hat{x} = x/\delta, \quad (5.2)$$

has an expansion

$$\left. \begin{aligned} \widehat{\Phi}_\delta(\widehat{x}) &= \sum_{n=m_0}^m \delta^n \widehat{\Phi}^{(n)}(\widehat{x}) + \mathcal{O}(\delta^m) \\ &\equiv \left(\widehat{\mathbf{E}}^{(m)} \widehat{\Phi}_\delta \right)(\widehat{x}) + \mathcal{O}(\delta^m) \end{aligned} \right\} \quad \text{for } \widehat{x} \in \widehat{S}. \quad (5.3)$$

Here $\widehat{x} \in \widehat{S}$ for some δ -independent $\widehat{S} \subset \mathbb{R}^n$ implies $x \in S_\delta$. In boundary layer type problems the size of S_δ shrinks to zero as $\delta \rightarrow 0$. In (5.1), (5.3), $\mathbf{E}^{(m)}$ is the $(m+1)$ -term expansion of functions of x in D_0 , while $\widehat{\mathbf{E}}^{(m)}$ is the expansion of functions of \widehat{x} in \widehat{S} ; note that $\widehat{\mathbf{E}}^{(m)}$ allows for a different scaling of the leading order term than $\mathbf{E}^{(m)}$.

Loosely speaking, van Dyke's matching principle requires "the near-field behaviour of the outer solution to match the farfield behavior of the inner solution" when both are expressed in the inner coordinates. Let \mathbf{T} denote the transformation from (5.2), i.e.,

$$(\mathbf{T} \Phi_\delta)(\widehat{x}) \equiv \widehat{\Phi}_\delta(\widehat{x}) = \Phi_\delta(\mathbf{T}^{-1}(\widehat{x})) \quad \text{with} \quad \mathbf{T}x = \widehat{x} = x/\delta. \quad (5.4)$$

Then, given the expansion depth m_o of the outer expansion, the matching principle reads

$$\widehat{\mathbf{E}}^{(m_i)} \mathbf{T} \mathbf{E}^{(m_o)} \Phi_\delta = \widehat{\mathbf{E}}^{(m_i)} \mathbf{T} \mathbf{E}^{(m_o)} \mathbf{T}^{-1} \widehat{\mathbf{E}}^{(m_i)} \mathbf{T} \Phi_\delta \quad (5.5)$$

for the smallest possible inner expansion depth m_i . Part of the matching procedure is to determine values for m_o and m_i such that the resulting system of equations for the asymptotic inner and outer representations of Φ_δ is closed and sufficiently accurate.

This simplified description of the matching principle suffices the present purposes (for a more comprehensive exposition of the technique, see (Eckhaus 1979)).

5.2. The centreline equation of motion

Matching the horizontal velocities in this section will yield the equation of motion for the vortex centreline in analogy with the analysis of three-dimensional slender vortex filaments (see Ting *et al.* 2007, and references therein). For the horizontal velocities, we keep terms up to and including $\mathcal{O}(\delta)$, so that $m_o = 1$ in (5.5).

5.2.1. *The nearfield behavior of the outer solution*

Based on (3.25) and recalling that ψ_2^* represents the non-singular part of the outer flow, we find the nearfield behaviour of the horizontal velocity in the outer solution

$$\begin{aligned}
\widehat{\mathbf{E}}^{(1)} \text{TE}^{(1)} \mathbf{u} &= \mathbf{u}_0^{(0)} + \delta \left(\mathbf{u}_0^{(1/2)} + \mathbf{X}^{(1)} \cdot (\nabla_{\parallel} \mathbf{u})_0^{(0)} \right) \\
&+ \left(\frac{1}{\delta} \frac{\Gamma}{2\pi \widehat{r}} - \delta M_0 \widehat{r} (2 \ln \widehat{r} - 1 + \ln \delta) \right) \mathbf{e}_\theta \\
&+ \delta \left(\widehat{r} \mathbf{e}_r \cdot (\nabla_{\parallel} \mathbf{u})_0^{(0)} \right), \\
&+ \delta \left(\left[\ln \widehat{r} - \ln \frac{1}{\delta} \right] \mathbf{k} \times \mathbf{M}_1 + (\mathbf{M}_1 \cdot \mathbf{e}_r) \mathbf{e}_\theta \right),
\end{aligned} \tag{5.6}$$

where \mathbf{u}_0 and $(\nabla_{\parallel} \mathbf{u})_0$ denote the regular part of the outer flow velocity and its horizontal gradient evaluated at (t, \mathbf{X}_0, z) , and M_0 and \mathbf{M}_1 were defined in (3.22). At this result we arrive by (i) calculating $-\mathbf{k} \times \nabla_{\parallel} \psi_s$ from the outer solution in (3.25), and (ii) taking into account that ψ_2^* is sufficiently smooth up to order $\mathcal{O}(\delta)$ (see App. B) so that the associated velocity has a two-term Taylor expansion near the centreline and becomes a contribution to the regular part of the outer velocity, \mathbf{u}_0 .

For later reference, we note that $\widehat{r} \mathbf{e}_r \cdot (\nabla_{\parallel} \mathbf{u})_0^{(0)}$ in (5.6) can be expressed in terms of a local quadratic stream function, and the local vorticity, ζ_0^* , (see App. D)

$$\widehat{r} \mathbf{e}_r \cdot (\nabla_{\parallel} \mathbf{u})_0^{(0)} = \widehat{r} \zeta_0^* \mathbf{e}_\theta - \mathbf{k} \times \widehat{\nabla}_{\parallel} \left(\widehat{r}^2 \left(\psi_{21}^* \sin(2\theta) + \psi_{22}^* \cos(2\theta) \right) \right), \tag{5.7}$$

with constants ψ_{2k}^* for $k \in \{1, 2\}$.

5.2.2. *The farfield behavior of the inner solution*

We recall from (4.5), (3.15b), (4.10), and (4.28) that the absolute horizontal velocity within the vortex core up to and including terms of order $\mathcal{O}(\delta)$ reads

$$\begin{aligned}
\widehat{\mathbf{E}}^{(1)} \mathbf{u} &= \frac{d\mathbf{X}_0}{dt} + \delta \frac{\partial \mathbf{X}^{(1)}}{\partial t} \\
&+ \frac{1}{\delta} \left(u_\theta^{(0)} + \delta u_\theta^{(1)} + \delta^2 u_{\theta,0}^{(2)} \right) \mathbf{e}_\theta \\
&+ \delta \left(\nabla_{\parallel} \phi^{(2)} - \mathbf{k} \times \nabla_{\parallel} \psi^{(2)} \right).
\end{aligned} \tag{5.8}$$

For the divergent part of the flow we have, by construction of $\phi^{(2)}$ in (4.30) that

$$\mathbf{E}^{(1)} \mathbf{T}^{-1} \left[\nabla_{\parallel} \phi^{(1)} \right] = 0. \quad (5.9)$$

For the farfield behaviour of $-\mathbf{k} \times \nabla_{\parallel} \psi^{(2)}$ in (5.8), consider the expression for $\psi_{1k}^{(2)}$ in (4.35). With de L'Hôpital's rule for limits of type “ ∞/∞ ”, the term in brackets obeys

$$\lim_{r \rightarrow \infty} \left[\frac{u}{\hat{r}} \int_0^{\hat{r}} \frac{1}{\hat{r} u^2} \int_0^{\bar{r}} r' \mathcal{F}(r') dr' d\bar{r} \right] = \frac{\pi}{\Gamma} \int_0^{\infty} r \mathcal{F}(r) dr \equiv L[\mathcal{F}], \quad (5.10)$$

provided $[u, \hat{r}^2 du/d\hat{r}] = [\Gamma/2\pi\hat{r}, -\Gamma/2\pi\hat{r}^2] (1 + o(1))$ as $\hat{r} \rightarrow \infty$, and the integral exists.

The integral converges for $\mathcal{F} \in \{\mathcal{H}_{1k}, \mathcal{J}_{1k}\}$ (see (4.33), (4.35)), and it does for $\mathcal{F} = \mathcal{Q}_{1k}$ if the diabatic source term vanishes sufficiently rapidly for large \hat{r} (see App. E). For $\mathcal{F} = \mathcal{I}_{1k}$, extraction of its dominant part and application of L'Hôpital's rules reveals a logarithmic divergence for large \hat{r} (see App. E). Combining all results we find

$$\hat{\mathbf{E}}^{(1)} \mathbf{T} \mathbf{E}^{(1)} \mathbf{T}^{-1} \left[\tilde{\psi}_1^{(1)} \right] = \left(-\hat{r} \ln \hat{r} \mathbf{M}_1(\tau, z) + \hat{r} \Psi(\tau, z) \right) \cdot \mathbf{e}_r(\theta), \quad (5.11)$$

where $\tilde{\psi}_1^{(1)} = \psi_{11}^{(1)} \sin \theta + \psi_{12}^{(1)} \cos \theta$,

$$\Psi = L \left[\mathcal{H} + \tilde{\mathcal{I}} + \mathcal{J} + \mathcal{Q} \right] \quad (5.12)$$

with $\mathcal{K} = -i \mathcal{K}_{11} + j \mathcal{K}_{12}$ for $\mathcal{K} \in \{\mathcal{H}, \tilde{\mathcal{I}}, \mathcal{J}, \mathcal{Q}\}$ and $L[\cdot]$ defined in (5.10). Here

$$\tilde{\mathcal{I}}_{1k} = \mathcal{I}_{1k} + H(\hat{r} - 1) I_{1k} / \hat{r}^2 \quad (5.13)$$

is the integrable part of \mathcal{I}_{1k} in that $\int_0^{\infty} r \tilde{\mathcal{I}}_{1k} dr$ exists, $H(\cdot)$ is the Heaviside step function, and I_{1k} is defined by $\mathbf{M}_1 = (\pi/\Gamma) (-i I_{11} + j I_{12})$ with \mathbf{M}_1 from (3.22b).

Collecting (5.9), (5.11), we find the farfield behavior of the inner horizontal velocity,

$$\begin{aligned} \hat{\mathbf{E}}^{(1)} \mathbf{T} \mathbf{E}^{(1)} \mathbf{T}^{-1} \hat{\mathbf{E}}^{(1)} \mathbf{T} \mathbf{u} &= \frac{d\mathbf{X}_0}{dt} + \delta \frac{\partial \mathbf{X}^{(1)}}{\partial t} \\ &+ \hat{\mathbf{E}}^{(1)} \mathbf{T} \mathbf{E}^{(1)} \mathbf{T}^{-1} \left[\frac{1}{\delta} \left(u_{\theta}^{(0)} + \delta u_{\theta}^{(1)} + \delta^2 u_{\theta,0}^{(2)} \right) \right] \mathbf{e}_{\theta} \\ &- \hat{\mathbf{E}}^{(1)} \mathbf{T} \mathbf{E}^{(1)} \mathbf{T}^{-1} \left[\delta \mathbf{k} \times \hat{\nabla} \tilde{\psi}_2^{(2)} \right] \\ &+ \delta \left(\ln \hat{r} \mathbf{k} \times \mathbf{M}_1 + (\mathbf{M}_1 \cdot \mathbf{e}_r) \mathbf{e}_{\theta} - \mathbf{k} \times \Psi \right), \end{aligned} \quad (5.14)$$

where $\tilde{\psi}_2^{(2)} = \psi_{21}^{(2)} \sin(2\theta) + \psi_{22}^{(2)} \cos(2\theta)$.

5.2.3. Matching

Here we match the large- \hat{r} inner solution in (5.14) with the nearfield outer solution in (5.6). Comparing the second line in (5.14) with the second line plus the first term of the third according to (5.7) in (5.6) we find

$$\left. \begin{aligned} u_\theta^{(0)} &= \frac{\Gamma^{(0)}}{2\pi\hat{r}} (1 + \mathcal{O}(1)) \\ u_\theta^{(1)} &= \frac{\Gamma^{(1/2)}}{2\pi\hat{r}} (1 + \mathcal{O}(1)) \\ u_\theta^{(2)} &= \hat{r} (\zeta_0^* + M_0 (2 \ln \hat{r} + 1 + \ln \delta)) (1 + \mathcal{O}(1)) \end{aligned} \right\} \text{ as } \hat{r} \rightarrow \infty. \quad (5.15)$$

The results for $u_\theta^{(0)}, u_\theta^{(1)}$ were expected from the theory of point vortices, while $u_\theta^{(2)}$ involves an additional influence of baroclinicity, $\partial\Gamma/\partial z$, through M_0 (see (3.22)).

The remainder of the third line of (5.6) according to (5.7) and the third line of (5.14) provide the large- \hat{r} behaviour of the Fourier modes, $\psi_{2k}^{(1)}$, of the perturbation stream function in the vortex core,

$$\hat{\mathbf{E}}^{(1)} \mathbf{T} \mathbf{E}^{(1)} \mathbf{T}^{-1} \psi_{2k}^{(1)} = r^2 \psi_{2k}^* (1 + \mathcal{O}(1)) \quad \text{as } r \rightarrow \infty. \quad (5.16)$$

On one hand, the ψ_{2k}^* are related to the irrotational strain of the regular part of the outer solution in \mathbf{X}_0 (see App. D). On the other hand, they represent the homogeneous solutions to the equations for the second Fourier modes of the stream function of the core flow which are thus determined (see the discussion before (C 12) and (C 13) in App. C).

Comparing the remaining terms in (5.6) and (5.14) we find the centreline velocity,

$$\frac{d\mathbf{X}_0}{dt} = \mathbf{u}_0^{(0)}, \quad (5.17a)$$

$$\frac{\partial \mathbf{X}^{(1)}}{\partial t} = \mathbf{u}_0^{(1/2)} + \mathbf{X}^{(1)} \cdot (\nabla_{\mathbf{u}} \mathbf{u})_0^{(0)} - \ln \frac{1}{\delta} (\mathbf{k} \times \mathbf{M}_1) + (\mathbf{k} \times \mathbf{\Psi}). \quad (5.17b)$$

6. Vortex Rossby waves (VRWs) and quasi-modes

6.1. VRWs and quasi-modes in the literature

In a series of papers, Montgomery and colleagues discuss the roles of vortex Rossby wave (VRW) damping and the long-lived precession of so-called quasi-modes in stabilizing an atmospheric vortex against vertical shear of the background flow (see Shapiro & Montgomery 1993; Montgomery & Kallenbach 1997; McWilliams *et al.* 2003; Reasor & Montgomery 2001; Reasor *et al.* 2004, and references therein). VRWs are analogous to the classical mid-latitude Rossby waves, except that in VRWs meridional advection of planetary rotation is replaced with radial advection of the axisymmetric mean vorticity of the vortex. Quasi-modes arise in these linear theories as the characteristic flow patterns associated with the low-frequency continuous spectrum of the linearized perturbation operators. Here we discuss why vortex Rossby wave dynamics on the vortex core cannot be captured by the present asymptotic theory whereas it does provide a nonlinear theory for the quasi-modes. Note, however, that VRWs in the outer flow are captured as part of the QG outer solution!

If a vortex features a long-range decay of vorticity away from its core, then a critical radial layer may exist within which circumferential advection by the mean vortex and the phase propagation of VRWs coincide and this resonance induces wave damping. Of particular interest in the present context are VRWs of circumferential Fourier mode one as these may be interpreted as oscillatory lateral displacements of the entire vortex and thus correspond roughly with vortex centreline motions. Why do VRWs (Montgomery & Kallenbach 1997) on the vortex core and the associated damping mechanism not emerge from the present asymptotic analysis? This is understood by comparing the characteristic time scales. The circumferential phase speed of VRWs is comparable with the circumferential advection time scale of the mean vortex, $T_{\text{VRW}} \sim L_{\text{mes}}/u_{\text{max}}$, where L_{mes} is

the radial location of the maximum circumferential mean velocity, u_{\max} . In contrast, the vortex centreline motions analyzed in this paper develop on longer time scales of order $L_{\text{syn}}/u_{\text{ref}} = \delta^{-2} T_{\text{VRW}}$. In the language of matched asymptotics for concentrated vortices, the VRWs thus correspond to the fast time solutions discussed by Ling & Ting (1988) (see also Ting *et al.* 2007), whereas the present centreline equations of motion correspond to slow time solutions.

Reasor *et al.* (2004), studying dry flow dynamics, find that vortex Rossby wave damping is weak in the typical flow regime for Hurricane development. This is demonstrated through numerical simulations based on the non-hydrostatic primitive equations for initially upright barotropic vortices with Gaussian vorticity distribution and maximum circumferential mean velocities of $u_{\max} = 10\text{...}40 \text{ ms}^{-1}$ at a radius of $L_{\text{mes}} = 100 \text{ km}$ corresponding with vortex Rossby numbers $\text{Ro}_{\text{mes}} = u_{\max}/f_0 L_{\text{mes}} = 3.2\text{...}12.7$. The internal Rossby radii of the background state are $L_{\text{syn}} = NH/\pi f_0 = 630\text{...}1250 \text{ km}$, so the vortices are concentrated with $L_{\text{mes}}/L_{\text{syn}} \sim 0.08\text{...}0.16$. The stability of the background stratification is characterized by $N/f_0 = \pi L_{\text{syn}}/H = 200\text{...}400$, where N is the (constant) Brunt-Väisälä frequency, and $H = 10 \text{ km}$ is the height of the flow domain. The stabilizing vortex Rossby wave damping is weak in these cases for two reasons: (i) the critical radius lies far outside the vortex core in a region with very small remaining radial vorticity gradient, and (ii) the vortex turnover time at the critical radius is very long because of the slow remaining circumferential velocity and because of the large length of the circumference. With weak VRW damping, these vortices support long-time precessions of quasi-modes with circumferential Fourier mode one, and these correspond roughly with horizontal displacements of the entire vortex structure as considered here.

Reasor *et al.* propose a heuristic model to describe these nearly undamped quasi-modes motivated by previous linearized analyses of Schecter *et al.* (2002); Schecter &

Montgomery (2003) in the context of the quasi-geostrophic and the asymmetric balanced model of Shapiro & Montgomery (1993). The heuristic model states that potential vorticity perturbations, q' , due to the precessing quasi-mode should obey

$$q' = -\frac{F_{q,s}(r, z)}{\omega_p} [\sin(\lambda - \omega_p t) - \sin(\lambda)] , \quad (6.1)$$

where λ denotes the circumferential angle relative to the background vertical shear direction, $F_{q,s}(r, z)$ is a forcing function that depends on the structure of the unperturbed axisymmetric vortex, and ω_p is the angular frequency of precession.

There are indications in the literature that VRW damping may not be weak at all in moist vortex flows (Schechter & Montgomery 2007). However, an in-depth discussion of moist flows is beyond the scope of the present paper.

6.2. Quasi-modes on the vortex core in the present theory

Here we demonstrate that, for vortices in gradient wind balance, i.e., for vortex Rossby numbers $\text{Ro}_{\text{mes}} = u_{\text{max}}/f_0 L_{\text{mes}} = O(1)$, the present theory provides an explanation of such quasi-modes as the net result of the vortex centreline dynamics. The first-order centreline evolution equation in (5.17b) may be rewritten as (see App. F)

$$\frac{\partial \mathbf{X}^{(1)}}{\partial t} = \mathbf{u}_0^{(1/2)} + \mathbf{X}^{(1)} \cdot (\nabla_{\parallel} \mathbf{u})_0^{(0)} - \mathbf{k} \times \left(A \frac{\partial^2 \mathbf{X}^{(1/2)}}{\partial z^2} + B \frac{\partial \mathbf{X}^{(1/2)}}{\partial z} \right) \quad (6.2)$$

where

$$A = \ln \frac{1}{\delta} \frac{\Gamma}{4\pi} \frac{f_0^2}{\Theta_1'} - \tilde{A}, \quad B = \ln \frac{1}{\delta} \frac{\Gamma}{4\pi} \frac{f_0^2}{\Theta_1'} \frac{\partial}{\partial z} \left(\ln \left[\frac{\rho_0 \Gamma^2}{\Theta_1'} \right] \right) - \tilde{B} \quad (6.3)$$

and where $\tilde{A}(t, z)$, $\tilde{B}(t, z)$ are integral expressions involving the leading order vortex core structure as detailed in App. F, eqs. (F6)–(F8). In the absence of diabatic forcing, the core structure becomes time independent as discussed in section 4.3, so that all coefficients in (6.2) depend on z only. Following Reasor *et al.* (2004) in assuming, in addition, a barotropic leading-order vortex and constant (logarithmic) background stratifica-

tion, corresponding to $\Theta'_1 = \text{const}$, the equations further simplify. Finally, since Reasor *et al.* (2004) use models based on the Boussinesq approximation for the development of their vortex alignment theory, we may consider here the simplified cases of exponential, $\rho_0(z) = \rho_\infty \exp(-z/H_\rho)$ or constant, $\rho_0(z) \equiv \rho_\infty$, background density. In both cases, all coefficients in (6.2) become constant, i.e., independent of z .

Reasor *et al.* (2004) assume forcing of an initially vertically aligned barotropic vortex by horizontally homogeneous easterly vertically sheared winds, so that

$$\mathbf{u}_0^{(1/2)} = \left(U \cos(\pi z/H), 0 \right)^T, \quad (6.4)$$

and $(\nabla_{\parallel} \mathbf{u})_0^{(0)} \equiv 0$ in (6.2), where H is the height of the top boundary of flow domain.

At the bottom and top of the domain, they assume constant potential temperature boundary conditions which, considering the potential temperature transport equation and the fact that the potential temperature is horizontally homogeneous in the farfield, corresponds to rigid lid conditions, i.e., to $w = 0$ at $z = 0$ and $z = H$. Recalling eqs. (4.25) for the first Fourier modes of the vertical velocity in the vortex core and considering the adiabatic case like Reasor *et al.* (2004), we find that rigid lid conditions require

$$\frac{\partial \mathbf{X}^{(1)}}{\partial z} = 0 \quad \text{at} \quad z = 0 \quad \text{and} \quad z = H. \quad (6.5)$$

This provides the top and bottom boundary conditions for solutions of (6.2).

With these simplifications, the equations in (6.2) are linear in $\mathbf{X}^{(1)}$, so that solutions may be constructed as superpositions of a particular solution $\mathbf{X}_p^{(1)}(z)$ satisfying

$$A \frac{d^2 \mathbf{X}_p^{(1)}}{dz^2} + B \frac{d\mathbf{X}_p^{(1)}}{dz} = -\mathbf{k} \times \mathbf{u}_0^{(1/2)} \quad (6.6)$$

and suitable homogeneous solutions which satisfy

$$\frac{\partial \mathbf{X}_h^{(1)}}{\partial t} = -\mathbf{k} \times \left(A \frac{\partial^2 \mathbf{X}_h^{(1)}}{\partial z^2} + B \frac{\partial \mathbf{X}_h^{(1)}}{\partial z} \right) \quad (6.7)$$

and will guarantee compliance with the initial conditions.

Particular solutions satisfying (6.6) for $\mathbf{u}_0^{(1/2)}(z)$ from (6.4) are of the form

$$\mathbf{X}_p^{(1)}(z) = \left(X_0, a \cos(\pi z/H) + b \sin(\pi z/H) + Y_1 \exp(-Bz/A) + Y_0 \right)^T \quad (6.8)$$

where

$$a = U \left(\frac{A\pi^2}{H^2} + \frac{B^2}{A} \right)^{-1}, \quad b = -\frac{BH}{A\pi} a. \quad (6.9)$$

$\mathbf{X}_p^{(1)}$ satisfies the boundary conditions from (6.5) provided $b = Y_1 = 0$, so that

$$\mathbf{X}_p^{(1)}(z) = \left(X_0, a \cos(\pi z/H) + Y_0 \right)^T. \quad (6.10)$$

This finding corroborates and extends earlier results in the literature: for a concentrated atmospheric vortex embedded in a background vertical shear there is an equilibrium configuration involving a down-shear left tilt of the vortex centre (see Reasor *et al.* 2004, and references therein). We extend this in that we allow for horizontal displacements of the vortex center comparable to the vortex core size whereas earlier studies relied on linearization and small displacement assumptions.

Real-valued homogeneous eigenmode solutions from (6.7) satisfying the boundary conditions in (6.5) read

$$\mathbf{X}_{h,k}^{(1)}(t, z) = e^{i\omega_k t} e^{-\mu z} \left(e^{i\nu_k z} - \frac{\mu - i\nu_k}{\mu + i\nu_k} e^{-i\nu_k z} \right) \left[\mathbf{X}_0^{(1)} + i\mathbf{k} \times \mathbf{X}_0^{(1)} \right] + c.c., \quad (6.11)$$

where *c.c.* indicates the complex conjugate, and

$$\omega_k = \frac{B^2}{4A} + \frac{A\pi^2}{H^2} k^2, \quad \mu = \frac{B}{2A}, \quad \nu_k = \frac{\pi}{H} k \quad (k = 1, 2, \dots). \quad (6.12)$$

For the Boussinesq case, $\rho_0 \equiv \rho_\infty = \text{const}$, we have $\mu = B = 0$ and the eigenmode solutions further simplify. In particular, the gravest mode with $k = 1$ becomes

$$\mathbf{X}_{h,1}^{(1)}(t, z) = \cos(\pi z/H) \left(\cos(\omega t) \mathbf{X}_0^{(1)} + \sin(\omega t) \mathbf{k} \times \mathbf{X}_0^{(1)} \right). \quad (6.13)$$

For an initially upright vortex perturbed by the shear from (6.4) we find,

$$\mathbf{X}^{(1)}(t, z) = \frac{U}{\omega} \cos(\pi z/H) \begin{pmatrix} \sin(\omega t) \\ 1 - \cos(\omega t) \end{pmatrix} \quad \text{where} \quad \omega = \frac{A\pi^2}{H^2}. \quad (6.14)$$

A quantitative comparison with Reasor *et al.* (2004) reveals that the precession frequency extracted from simulations of vortices with Gaussian distribution of potential vorticity agrees well with the present theory for their first case in their section 5. This case involves $u_{\max} = 10$ m/s, $L_{\text{mes}} = 100$ km, $H = 10$ km, $L_{\text{syn}} = NH/\pi f_0 = 1250$ km, and a vortex Rossby number $\text{Ro}_{\text{mes}} = u_{\max}/L_{\text{mes}}f_0 = 3.2$. While Reasor *et al.* report a ratio of the vortex turn-over time at the radius of maximum wind speed and the vortex precession time of $\omega^* \equiv \omega L_{\text{mes}}/u_{\max} = 0.053$, we have agreement with $\omega^* = 0.059$ on the basis of (6.14), the $\sim 10\%$ difference is compatible with a higher order error in $\delta = \sqrt{L_{\text{mes}}/L_{\text{syn}}} \sim 0.09$ for this case.

For the more intense vortex with $\text{Ro}_{\text{mes}} = 12.7$, however, the discrepancy is substantial with $\omega^* \approx 0.2$ in (Reasor *et al.* 2004) and $\omega^* \approx 0.5$ from the present theory. Such a deviation is not surprising because the regime of validity of the present theory is the gradient wind regime which, by definition, has $\text{Ro}_{\text{mes}} = O(1)$ as $\varepsilon \rightarrow 0$. With actual values of $\varepsilon \sim 1/8$, the stronger vortex lies outside this regime since $\text{Ro}_{\text{mes}} = 12.7 \sim 1.5 \varepsilon^{-1}$. Another indication that our theory does not apply to this case comes from the asymptotic scaling. We should have $\omega^* = O(\varepsilon)$ as $\varepsilon \rightarrow 0$, whereas the precession and vortex turn-over times become comparable for the stronger vortex.

7. Discussion

Here we provide an additional discussion of the core structure evolution equation in (4.21) and the centreline equations of motion in (5.17), and we summarize several limitations of the theory together with possible remedies.

7.1. The centreline equations of motion

7.1.1. General properties

The left side of the leading order centreline equation of motion, (5.17a), does not depend on z . This implies the regular part of the quasi-geostrophic background flow on the right to be barotropic to leading order, and is consistent with results in the literature stating that (nearly) adiabatic vortices cannot sustain a strong horizontal shear (see, e.g., Jones 1995, 2000; Reasor & Montgomery 2001; Reasor *et al.* 2004, and references therein). It is also consistent with the observation that strong, concentrated atmospheric vortices develop predominantly in the sub-tropics, where the Coriolis effect and thermal wind shear are reduced in comparison with the middle latitudes. We recall, however, from the previous discussions that the vortex even with the assumed weak shear will undergo horizontal displacements comparable to the vortex core size and will thus be strongly tilted. Given the implied weak shear of the background flow, one may surmise that the outer flow could then have been described by a simpler than the quasi-geostrophic theory in the first place. We leave an exploration of this possibility for future work.

The first two terms on the right in (5.17b) describe the first order advection of the vortex centre by the background flow. The remaining terms capture the self-induced vortex motion. The definitions of \mathbf{M}_1 and Ψ from (3.22b), (5.12), together with (4.33), (4.23a), (4.25), and (4.30), reveal that vertical tilt and baroclinicity, asymmetries of the diabatic sources, and the background density stratification interact nonlinearly to produce this deviation of the vortex centre motion from simple advection by the background wind.

7.1.2. Local induction

In the special case of a barotropic vortex, $\partial\Gamma/\partial z \equiv 0$, in an atmosphere with constant density, $d\rho_0/dz \equiv 0$, and constant stratification, $d^2\Theta_1/dz^2 \equiv 0$, the leading logarithmic

term in (5.17b) is proportional to $-(\Gamma/4\pi)\kappa\mathbf{b}$ with κ the centreline curvature and \mathbf{b} its binormal unit vector. To see this, consider the definition of \mathbf{M}_1 in (3.22b) and use that $\partial^2\mathbf{X}^{(1/2)}/\partial z^2 = \kappa\mathbf{n} + \mathcal{O}(\delta)$ and $\mathbf{k} = \boldsymbol{\tau} + \mathcal{O}(1)$. At a first glance, this result appears to be in line with the theory of slender vortex filaments (Callegari & Ting 1978), where it emerges as the leading-order term in the centreline equation of motion. Keeping only this leading term one has the classical “local induction approximation (LIA)” (see also Ricca 1991; Ting *et al.* 2007, and references therein). There is a discrepancy, however, in the sign of the term. Whereas a three-dimensional filament moves in the direction of the binormal unit vector for positive circulation, a hurricane-like vortex precesses in the opposite direction according to this term.

To obtain a physical interpretation in the present context, consider eqs. (3.20)–(3.22b), which provide the expressions for \mathbf{M}_1 and emerge from the analysis of the QG solutions in the immediate vicinity of the singular vortex. The term in question involving \mathbf{M}_1 is part of the first term in (3.20), i.e., $-\mathcal{L}_z\psi_0^{2d}$. Recalling the derivations leading to this point, we find that this term embodies the net effect on the stream function of bending a potential vortex, represented by ψ_0^{2d} , to follow a curved centreline.

Close inspection of the detailed derivation for three-dimensional vortex filaments in (Callegari & Ting 1978) reveals that the corresponding term is the sum of two terms in the asymptotic evaluation of the Biot-Savart kernel which, in analogy with the present situation, provides the farfield behavior in the immediate vicinity of the singular vortex line. To verify this, see the first equation on p. 172 in (Callegari & Ting 1978) and trace back the derivation of the second term on the right. Now, the first term in the slender filament theory is in agreement with the present result for atmospheric vortices and has an interpretation analogous to that given in the previous paragraph: it represents the local net effect of bending a potential vortex along a curved centreline. This effect is

present independently of whether or not the vortex is long-stretched in the axial direction, like the classical slender vortex, or whether it is a flow in a shallow fluid layer, like the atmospheric vortices considered here. In contrast, the second term in the slender vortex theory, which is twice as large with opposite sign, does not appear here. This term is the result of nonlocal induction from vorticity located far way along the vortex filament axis when measured in units of the core diameter. This is also why this term is absent here: due to the extreme “pan cake-like” aspect ratio of the atmospheric vortex its axial extent is small compared to the vortex core diameter and the far-field induction effect does not materialize.

7.2. Vortex core structure

7.2.1. The adiabatic lifting mechanism

Jones (1995, 2000) numerically simulates an adiabatic atmospheric vortex in a background shear. After 6 hours she observes and discusses a constant “relationship between the direction of the vortex tilt and the orientations of the potential temperature perturbation and vertical-velocity pattern”. This relationship corresponds with the tilt–potential temperature asymmetry described in (4.23b) and (4.25) above. In the adiabatic setting, a wavenumber-one pattern of ascent and descent with a phase shift of 90° relative to the potential temperature pattern is found which is reflected in (4.23b) (note the index shift on the Θ modes). These vertical velocity asymmetries induce the lowering and raising of Θ surfaces called “adiabatic lifting mechanism” by Frank & Ritchie (1999).

7.2.2. The vortex core evolution equation

Whereas the left side of the vortex core structure evolution equation (1.3a) describes advection of circumferential momentum in the meridional plane of the vortex, the right side represents nonlinear interactions of vertical velocity asymmetries with the vortex tilt

(see definition of $u_{r,*}$ in (1.3d) as derived in (4.22)). Such an effect has also been discussed by Shapiro & Montgomery (1993) in the context of linear and weakly nonlinear theories for balanced perturbations of a strong vertically aligned axisymmetric vortex. Collecting all contributing terms we find that $\partial u_\theta / \partial t \equiv 0$ in the absence of diabatic heating even when the vortex has a tilt and associated asymmetric vertical velocity fluctuations. In this case, these asymmetries and the tilt have a particular phase relationship (see section 7.2.1) that annihilates $u_{r,*}$. This annihilation will not occur in general when diabatic effects are active, and in that case a nonlinear coupling of the vortex core structure evolution with the vortex tilt dynamics arises and is mediated by the first circumferential Fourier modes of the asymmetric heating (see eqs. (5.12), (4.33), and the relation between the asymmetric vertical velocities, \hat{w}_{1k} , and the asymmetric heating in (4.23)). Note that these results on the local self-induced motion due to processes in the vortex core will complement any strengthening, weakening, or alignment processes due to vortex Rossby waves in the QG farfield (see Montgomery & Kallenbach 1997; Reasor & Montgomery 2001; Schecter *et al.* 2002, and item (b) in Sec. 7.3.), and these processes can be captured here through solutions of the regularized outer flow problem in (1.2).

7.3. *Limitations of the present theory and outlook*

The theory presented here is based on a number of simplifying assumptions that limit its range of applications and deserve scrutiny in future research work:

(a) Large Rossby numbers and the inner core of a hurricane: In this paper we have extended asymptotic theories for concentrated quasi-geostrophic vortices to the gradient wind regime, thereby moving from small to moderate vortex Rossby numbers, and covering “tropical storms” and “incipient hurricanes”. Mature hurricanes, however, feature large vortex Rossby numbers with their inner cores in approximate cyclostrophic balance, where pressure gradients and centripetal accelerations dominate the Coriolis effect.

Describing such a vortex asymptotically may require a three layer analysis involving a cyclostrophic inner core, a first outer layer in gradient wind balance, and matching to the geostrophic far field. In such a setting, new dynamical effects will arise due to the fact that the vortex turn-over time is no longer separated asymptotically from the vortex Rossby and buoyancy wave eigenfrequencies, so that a Rossby-Inertia-Buoyancy (RIB) wave instability may occur (see Plougonven & Zeitlin 2002; Schecter & Montgomery 2004, 2006, and references therein).

(b) Long-range tails of vorticity: We assume here that the leading order farfield behavior at every vertical level is consistent with that of a point vortex. Since turbulent friction and vertical momentum transport due to moist convection will couple the vertical layers, this assumption will have to be modified in a theory for cyclones in moist air. With such an extension one can then address the effect of larger scale farfield tails of potential vorticity. Combining the VRW theory for a QG vortex to describe the outer flow from (Schecter *et al.* 2002) with the present quasi-stationary theory for the vortex core in gradient wind balance appears to be promising in this context.

(c) Fast processes – vortex stabilization: One central assumption adopted here is that the characteristic time scale of all processes considered is commensurate with the synoptic time scale, i.e., with the characteristic time scale of the large-scale motion of the vortex. This precludes the description of internal waves, vortex Rossby waves, and advection on the vortex core, all of which act on much shorter time scales. Following Ting & Ling (see, e.g., 1983); Ting *et al.* (see, e.g., 2007) one can incorporate such faster processes systematically through multiple scales asymptotics in time.

One important effect in this context is the resilience against vertical background shear induced by vortex Rossby wave damping (see, e.g., Reasor *et al.* 2004, and references therein). Note that VRWs on the vortex core have a systematically shorter time scale

than those on the QG outer flow, so the present statement does not contradict the related remark on VRWs under item (b) above.

(d) The β -drift: The reader might be missing an explicit representation of the beta drift (see, e.g., Smith & Ulrich 1990; Smith 1991; Reznik 1992; Reznik & Kizner 2007b) in (5.17), which is the technical term for the self-induced motion of a vortex due to the large scale transport of the horizontally inhomogeneous planetary rotation. According to Reznik & Kizner (2007b) the order of magnitude of the β -drift velocity is $U_\beta \sim \beta L_{\text{syn}}^2$, where L_{syn} is the Rossby deformation radius. In nondimensional terms using our notation from (2.4), (2.5), this means $U_\beta/u_{\text{ref}} = (\beta L_{\text{syn}}/f_0)(f_0 L_{\text{syn}}/u_{\text{ref}}) = \beta^*/\text{Ro} = \widehat{\beta}/\widehat{f}_0 = \mathcal{O}(1)$. Thus, the β -drift will contribute at leading-order to the vortex motion. As the effect is due, however, to the advection of planetary rotation over distances comparable to the Rossby deformation radius, L_{syn} , it is an effect that materializes implicitly in the outer quasi-geostrophic flow as part of $\mathbf{u}_0^{(0)}$ in (5.17a). Within the vortex core of size $L_{\text{mes}} \sim \delta L_{\text{syn}}$ the effect can merely generate a dimensionless higher order contribution of order $\beta L_{\text{mes}}^2/u_{\text{ref}} \sim \mathcal{O}(\delta^2)$, which is why it does not appear in the vortex core dependent terms \mathbf{M}_1, Ψ in (5.17b).

An explicit representation of the development of the β -drift is given by Reznik and co-workers in (Reznik 1992; Reznik & Kizner 2007a,b; Reznik & Grimshaw 2001) who consider the initial value problem of a concentrated vortex being placed in an unperturbed QG background flow on the β -plane. They find that the rearrangement of planetary rotation by the farfield potential vortex generates two large-scale counter-rotating gyres which in turn induce an effective mean flow component in the vicinity of the vortex. This self-induced flow shifts the entire vortex in the westward direction, and this is called the β -drift. We note that the outer flow description in the present theory in (1.2) includes explicitly the perturbation of the farfield by the presence of the vortex through the stream

function contributions ψ_1^{2d} and ψ_1^{2d} , and we expect future solutions of these equations to reproduce the β -drift as a contribution to the leading order centreline velocity.

Reznik & Grimshaw (2001) analyse the evolution of a singular vortex based on the shallow water equations. In addition to the quasi-geostrophic part of the β -effects they reveal ageostrophic dynamics on the faster inertia-gravity wave time scale. Capturing these faster-scale processes in the present theory would require introduction of multiple time scale asymptotics.

(e) Stronger diabatic source terms and moist processes: Preliminary work by Mikusky (2007) (E. Päschke’s maiden name) indicates that the strength of diabatic source terms in the presence of moisture can be an order of magnitude larger in δ than assumed here. The core structure evolution equation in (1.3a) with the strength of the source term from (1.3d) reveals that with such stronger diabatically induced vertical velocities weaker tilts will suffice to drive changes of the vortex core structure on the same time scale. Also, as mentioned already in section 6, the presence of moisture can enhance the damping effect of vortex Rossby waves (Schecter & Montgomery 2007), thus increasing the vortex’ resilience against vertical shear.

(f) Boundary layer effects: Clearly, we have thus far neglected boundary layer effects, while the established literature (see, e.g., Emanuel 2003; Smith & Montgomery 2010) reports that boundary layer processes will have to be included as soon as moist processes are considered, as these feed and partially determine the vortex’ secondary circulation.

(g) Logarithmic terms: It is beyond the scope of the present paper to either attempt an “analysis beyond all orders” following Keller & Ward (1996) for the logarithmic terms appearing in the asymptotic matching in section 5, or to exploit the intermediate asymptotic ordering induced by the logarithms as done for slender fluid vortices by Klein & Majda (1991) (see also Ting *et al.* 2007).

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Appendix A. Governing equations in the co-moving coordinates

Transforming (2.6) to the vortex-centred coordinates from section 4.1.2 using (4.6) – (4.8) we using, $\mathbf{U} \equiv \partial \mathbf{X} / \partial t$, and

$$\nabla_{||} \equiv \frac{1}{\delta} \widehat{\nabla} \equiv \frac{1}{\delta} \left(\mathbf{e}_r \frac{\partial}{\partial \widehat{r}} + \mathbf{e}_\theta \frac{1}{\widehat{r}} \frac{\partial}{\partial \theta} \right) \quad \text{and} \quad \frac{\partial}{\partial z} \Big|_{t, \mathbf{x}} \equiv \frac{\partial}{\partial z} \Big|_{t, \widehat{\mathbf{x}}} - \frac{1}{\delta} \frac{\partial \mathbf{X}}{\partial z} \cdot \widehat{\nabla} \quad (\text{A } 1)$$

we find

$$\begin{aligned} \frac{\partial(\mathbf{U} + \mathbf{u}_{\text{rel}})}{\partial t} + \frac{1}{\delta} \mathbf{u}_{\text{rel}} \cdot \widehat{\nabla} \mathbf{u}_{\text{rel}} + w \left[\frac{\partial}{\partial z} - \frac{1}{\delta} \frac{\partial \mathbf{X}}{\partial z} \cdot \widehat{\nabla} \right] (\mathbf{U} + \mathbf{u}_{\text{rel}}) \\ + \frac{1}{\delta^7} \frac{1}{\rho} \widehat{\nabla} p + \frac{1}{\delta^2} f \mathbf{k} \times (\mathbf{U} + \mathbf{u}_{\text{rel}}) = 0, \end{aligned} \quad (\text{A } 2a)$$

$$\begin{aligned} \frac{\partial w}{\partial t} + \frac{1}{\delta} \mathbf{u}_{\text{rel}} \cdot \widehat{\nabla} w + w \left[\frac{\partial}{\partial z} - \frac{1}{\delta} \frac{\partial \mathbf{X}}{\partial z} \cdot \widehat{\nabla} \right] w \\ + \frac{1}{\delta^{10}} \left(\frac{1}{\rho} \left[\frac{\partial}{\partial z} - \frac{1}{\delta} \frac{\partial \mathbf{X}}{\partial z} \cdot \widehat{\nabla} \right] p + 1 \right) = 0, \end{aligned} \quad (\text{A } 2b)$$

$$\frac{\partial \rho}{\partial t} + \frac{1}{\delta} \widehat{\nabla} \cdot (\rho \mathbf{u}_{\text{rel}}) + \left[\frac{\partial}{\partial z} - \frac{1}{\delta} \frac{\partial \mathbf{X}}{\partial z} \cdot \widehat{\nabla} \right] (\rho w) = 0, \quad (\text{A } 2c)$$

$$\frac{\partial \Theta}{\partial t} + \frac{1}{\delta} \mathbf{u}_{\text{rel}} \cdot \widehat{\nabla} \Theta + w \left[\frac{\partial}{\partial z} - \frac{1}{\delta} \frac{\partial \mathbf{X}}{\partial z} \cdot \widehat{\nabla} \right] \Theta = Q_\Theta, \quad (\text{A } 2d)$$

where $\Theta = p^{\frac{1}{\gamma}} / \rho$.

Appendix B. The regular part of the outer flow streamfunction

To show that under (3.26) ψ_2^* will be continuous at $r = 0$, consider a repetition of the procedure from section 3.3.1. Let ψ_2^{2d} solve

$$\nabla_{||}^2 \psi_2^{2d} = -\mathcal{L}_z \psi_1^{2d}. \quad (\text{B } 1)$$

Applying \mathcal{L}_z to (3.23) the most singular behavior in r is generated by the second term on the right. When the two derivatives in \mathcal{L}_z meet the factor $r \ln r$, the following term arises,

$$\frac{\delta^3}{r} \frac{f_0^2}{\Theta_1'} M_1 \left(\mathbf{e}_r \cdot \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right)^2 \quad (\text{B } 2)$$

The corresponding contribution to ψ_2^{2d} will be linear in r . All other contributions will be at least as smooth near $r = 0$ as $r^2 \ln r$. As a consequence, the source term in the equation for the remainder, i.e.,

$$(\nabla_n^2 + \mathcal{L}_z) \psi_3^* = -\mathcal{L}_z \psi_2^{2d} - q_r. \quad (\text{B } 3)$$

is no longer singular at $r = 0$ and ψ_3^* will be smooth and thus lending itself for accurate numerical solution.

Appendix C. Further details of the core solution

C.1. Homogeneous solutions to the second order core streamfunction equation

There are two linearly independent homogeneous solutions to (4.31). Because $\zeta_{\hat{r}}^{(0)} \rightarrow 0$ as $\hat{r} \rightarrow 0$ and $\hat{r} \rightarrow \infty$, the operator on the left of (4.31) reduces to $\partial_{\hat{r}\hat{r}} + (1/\hat{r})\partial_{\hat{r}} - 1/\hat{r}^2$, with homogeneous solutions \hat{r} and \hat{r}^{-1} , in these limits. One of the homogeneous solutions to (4.31) is of the form $C_{1k} u_{\theta}^{(0)}$ with $C_{1k} = \text{const}$. It is proportional to \hat{r} near $\hat{r} = 0$ and to \hat{r}^{-1} as $\hat{r} \rightarrow \infty$. The other homogeneous solution is singular at $\hat{r} = 0$ and must therefore be discarded.

C.2. Higher harmonics of the leading order core solutions

Here we provide some considerations that were not listed in the main text to not distract from the principal line of thought.

Since $\hat{p}^{(4)}$ is axisymmetric (see (4.14b)), the vertical momentum balance in (4.17)

implies that the circumferential average is trivially hydrostatic, i.e.,

$$\widehat{\rho}_0^{(4)} = -\frac{\partial \widehat{p}^{(4)}}{\partial z}, \quad (\text{C } 1)$$

and that all higher than the first circumferential harmonics of $\widehat{\rho}^{(4)}$ and – via the equation of state (4.20) – of $\widehat{\Theta}^{(4)}$ vanish identically, i.e.,

$$\widehat{\Theta}_{lk}^{(4)} = -\frac{\widehat{\rho}_{lk}^{(4)}}{\rho_0} = 0 \quad \text{for} \quad l \geq 2; \quad k = 1, 2. \quad (\text{C } 2)$$

Thus, $\widehat{\Theta}^{(4)}$ has no higher than the first circumferential harmonics. This allows us subsequently to obtain closed expressions for the harmonics of the second order horizontal flow despite the presence of the nonlinear advection terms, which would otherwise couple all Fourier modes through triad interactions and render the equation system unclosed.

All higher harmonics $\widehat{w}_{lk}^{(1)}$ for $l \geq 2$ obey equations such as (4.23b) without the terms involving $\widehat{\Theta}_{lk}^{(4)}$ because these vanish identically (see (C 2)). Thus, in an adiabatic vortex $\widehat{w}_{lk}^{(1)} = 0$ for $l \geq 2$ and $k \in \{1, 2\}$. Since we wish to allow only for source terms Q_Θ that do not principally alter the solution structure relative to the adiabatic case, we assume

$$Q_{\Theta, lk}^{(2)} = 0 \quad \text{and hence} \quad \widehat{w}_{lk}^{(1)} = 0 \quad \text{for} \quad (l \geq 2; \quad k = 1, 2). \quad (\text{C } 3)$$

C.3. First order horizontal momentum balance

The radial momentum balance at $\mathcal{O}(\delta^{-2})$, (4.15a), yields higher order gradient wind balance for $\widehat{p}_0^{(5)}$ after circumferential averaging,

$$\frac{1}{\rho_0} \frac{\partial \widehat{p}_0^{(5)}}{\partial \widehat{r}} - \frac{2u_\theta^{(0)} u_\theta^{(1)}}{\widehat{r}} - f_0 u_\theta^{(1)} = 0. \quad (\text{C } 4)$$

The first Fourier modes of (4.15) reveal the mode one components of $\widehat{p}^{(5)}$ to be in geostrophic balance with the Coriolis force due to the leading-order vortex motion, i.e.,

$$\widehat{p}_{11}^{(5)} = -\widehat{r} \rho_0 f_0 \frac{dX_0}{dt}, \quad \widehat{p}_{12}^{(5)} = \widehat{r} \rho_0 f_0 \frac{dY_0}{dt}. \quad (\text{C } 5)$$

All higher harmonics of $\widehat{p}^{(5)}$ are found to vanish identically.

C.4. Higher harmonics of the second order horizontal flow

Here we first define

$$\mathcal{D}_j = \left(\frac{\partial^2}{\partial \hat{r}^2} + \frac{1}{\hat{r}} \frac{\partial}{\partial \hat{r}} - \frac{j^2}{\hat{r}^2} \right), \quad (\text{C } 6)$$

and then, in analogy with the derivations in section 4.4, arrive at

$$4 \mathcal{D}_2 \left[\phi_{21}^{(2)} \right] = \frac{\partial X^{(1)}}{\partial z} \left[\frac{\partial}{\partial \hat{r}} + \frac{1}{\hat{r}} \right] \hat{w}_{11}^{(1)} + \frac{\partial Y^{(1)}}{\partial z} \left[\frac{\partial}{\partial \hat{r}} + \frac{1}{\hat{r}} \right] \hat{w}_{12}^{(1)} \quad (\text{C } 7a)$$

$$4 \mathcal{D}_2 \left[\phi_{22}^{(2)} \right] = \frac{\partial X^{(1)}}{\partial z} \left[\frac{\partial}{\partial \hat{r}} + \frac{1}{\hat{r}} \right] \hat{w}_{12}^{(1)} + \frac{\partial Y^{(1)}}{\partial z} \left[-\frac{\partial}{\partial \hat{r}} + \frac{1}{\hat{r}} \right] \hat{w}_{11}^{(1)} \quad (\text{C } 7b)$$

and

$$\mathcal{D}_j \left[\phi_{jk}^{(2)} \right] = 0, \quad \text{for } j \geq 3; \quad k = 1, 2. \quad (\text{C } 8)$$

The second harmonics, $\phi_{2k}^{(2)}$, can be solved for analytically given the asymmetries of $\hat{w}^{(1)}$ and the components of the tilt (Bronstein & Semendjajew 1979). Just as for the first harmonics, these solutions should be determined under the constraint that the divergent part of the horizontal flow has to vanish at large \hat{r} consistent with the non-divergent outer flow. For higher harmonics mass conservation is guaranteed if $\phi_{jk}^{(2)} \equiv 0$ for $j \geq 3; k = 1, 2$.

In analyzing the second and higher harmonics of the streamfunction we find

$$\left(\frac{\partial^2}{\partial \hat{r}^2} + \frac{1}{\hat{r}} \frac{\partial}{\partial \hat{r}} - \left[\frac{\zeta_{\hat{r}}^{(0)}}{u_{\theta}^{(0)}} + \frac{4}{\hat{r}^2} \right] \right) \psi_{22}^{(2)} = \frac{(-1)^{k^*}}{u_{\theta}^{(0)}} \left(\mathcal{H} + \mathcal{I} + \mathcal{J} + \mathcal{Q} \right)_{2k^*}, \quad (\text{C } 9)$$

where $k^* = 3 - k$, and

$$\mathcal{H}_{2k} = \frac{1}{2} \frac{\partial}{\partial \hat{r}} \left(\hat{r} \hat{w}_{2k}^{(1)} \frac{\partial u_{\theta}^{(0)}}{\partial z} \right) \quad (\text{C } 10a)$$

$$\mathcal{I}_{2k} = \frac{1}{2} [\hat{r}(\zeta^{(0)} + f_0)] \mathcal{D}_2 \phi_{2k}^{(2)} \quad (\text{C } 10b)$$

$$\mathcal{J}_{2k} = \frac{1}{2} \frac{\partial \phi_{2k}^{(2)}}{\partial \hat{r}} \left(r \frac{\partial \zeta^{(0)}}{\partial \hat{r}} \right) \quad (\text{C } 10c)$$

$$\mathcal{Q}_{2k} = \left[-\frac{1}{4} \frac{\partial}{\partial \hat{r}} \left(\hat{r} \frac{\partial u_{\theta}^{(0)}}{\partial \hat{r}} \left[\cdot \right] \right) + \frac{u_{\theta}^{(0)}}{2\hat{r}} \right] \mathcal{X}_{2k} \quad (\text{C } 10d)$$

with

$$\mathcal{X}_{21} = \left(\frac{\partial X^{(1)}}{\partial z} \hat{w}_{11}^{(1)} + \frac{\partial Y^{(1)}}{\partial z} \hat{w}_{12}^{(1)} \right), \quad \mathcal{X}_{22} = \left(\frac{\partial X^{(1)}}{\partial z} \hat{w}_{12}^{(1)} - \frac{\partial Y^{(1)}}{\partial z} \hat{w}_{11}^{(1)} \right). \quad (\text{C } 11)$$

The second harmonics of the horizontal velocity depend on the first Fourier modes and the vortex tilt which do not depend on the second harmonics in return. Thus we will not need the second and higher harmonics to arrive at our main results, the evolution equations for the leading order core structure and the vortex tilt. Yet, we will need the large- \hat{r} behaviour of the second harmonics in constructing consistently matched asymptotic solutions for the vortex core and outer flow up to and including terms of order $\mathcal{O}(\delta)$.

In preparation for this step we consider the large- \hat{r} behaviour of homogeneous solutions to (C 9). Since the derivative of the leading order vorticity, $\zeta_{\hat{r}}^{(0)}$, in (C 9) vanishes for large \hat{r} by assumption, these homogeneous solutions satisfy

$$\left(\frac{\partial^2}{\partial \hat{r}^2} + \frac{1}{\hat{r}} \frac{\partial}{\partial \hat{r}} - \frac{4}{\hat{r}^2} \right) \psi_{2k,h}^{(2)} = \mathcal{O}(1) \quad \text{as} \quad \hat{r} \rightarrow \infty, \quad (\text{C } 12)$$

and one finds that

$$\psi_{2k,h}^{(2)} = \psi_{2k}^* \hat{r}^2 (1 + \mathcal{O}(1)) \quad \text{as} \quad \hat{r} \rightarrow \infty \quad (\text{C } 13)$$

with $\psi_{2k}^* = \text{const}$ is consistent with this behaviour. The amplitudes ψ_{2k}^* of the corresponding homogeneous contributions to $\psi_{2k}^{(2)}$ are determined in the course of matching the inner and outer solutions in section 5.

The higher harmonics of the streamfunction satisfy

$$\left(\frac{\partial^2}{\partial \hat{r}^2} + \frac{1}{\hat{r}} \frac{\partial}{\partial \hat{r}} - \left[\frac{\zeta_{\hat{r}}^{(0)}}{u_{\theta}^{(0)}} + \frac{j^2}{\hat{r}^2} \right] \right) \psi_{jk}^{(2)} = 0, \quad \text{for} \quad j \geq 3; \quad k = 1, 2 \quad (\text{C } 14)$$

and these equations are trivially solved by $\psi_{jk}^{(2)} \equiv 0$.

Appendix D. Linearization of the outer flow near the centreline

As the outer flow regular velocity is divergence free its local linearization can be expressed in terms of a local quadratic stream function,

$$r \mathbf{e}_r \cdot (\nabla_{\parallel} \mathbf{u}^*)_0 = -\mathbf{k} \times \nabla_{\parallel} \left(-\frac{r^2 \zeta_0^*}{2} + \psi_2^* \right). \quad (\text{D } 1)$$

where

$$\psi_2^*(r, \theta) = r^2 \left(\psi_{21}^* \sin(2\theta) + \psi_{22}^* \cos(2\theta) \right). \quad (\text{D } 2)$$

Here ζ_0^* is the vorticity of the regular part of the flow at \mathbf{X}_0 , and ψ_{21}^*, ψ_{22}^* define the components of the symmetric part of the velocity gradient, i.e., of the local irrotational strain tensor. A straightforward calculation shows that

$$\left((\nabla_{\parallel} \mathbf{u}^*)_0 \right) = - \begin{pmatrix} -\psi_{21}^* & \psi_{22}^* + \zeta_0^* \\ \psi_{22}^* - \zeta_0^* & \psi_{21}^* \end{pmatrix}. \quad (\text{D } 3)$$

The representation in (D 1), (D 2) implies (5.7).

Appendix E. Large- \hat{r} behavior of the operator in (5.10), (5.11)

Consider the solution operator for the first Fourier mode stream function of the inner solution in (4.35), i.e.,

$$\mathcal{L}_u [\mathcal{F}] = \frac{u}{\hat{r}} \int_0^{\hat{r}} \frac{1}{\bar{r} u^2} \int_0^{\bar{r}} r' \mathcal{F}(r') dr' d\bar{r}. \quad (\text{E } 1)$$

Assume that $u = (\Gamma/2\pi\hat{r}) (1 + \mathcal{O}(1))$ and $\hat{r}^2 du/d\hat{r} = -(\Gamma/2\pi\hat{r}^2) (1 + \mathcal{O}(1))$ as $\hat{r} \rightarrow \infty$.

Following de L'Hôpital's rule we have

$$\lim_{\hat{r} \rightarrow \infty} \mathcal{L}_u [f] (\hat{r}) = \lim_{\hat{r} \rightarrow \infty} \frac{\int_0^{\hat{r}} r' f(r') dr'}{\hat{r} u^2 \left(\frac{1}{u} - \frac{\hat{r} u'}{u^2} \right)} = \frac{\pi}{\Gamma} \int_0^{\infty} r f(r) dr \quad (\text{E } 2)$$

if the limits on the right exists. For the four source terms of the equations for the $\psi_{1k}^{(1)}$ in (4.35) we now need to investigate the convergence of this last integral.

For \mathcal{H}_{1k} , defined in (4.33), we recall the solution for $w_{1k}^{(1)}$ in (4.25). Since we consider the diabatic source term $Q_{\Theta}^{(2)}$ as given externally here, we may assume it to decay sufficiently rapidly so that any time it occurs in the integrand in $\mathcal{L}_u [f]$, the inner integral will converge. Then, the term in $w_{1k}^{(1)}$ with the slowest decay is proportional to $u_{\theta}^{(0)2}/r \sim r^{-3}$,

since $u_\theta^{(0)} \sim r^{-1}$ for large r . This implies $\widehat{r}\mathcal{H}_{1k} \sim \widehat{r}^{-3}$, and last integral in (E2) will converge for $f = \mathcal{H}_{1k}$. Analogously, $\widehat{r}\mathcal{Q}_{1k} \sim \widehat{r}^{-4}$ and the integral converges.

For the convergence of the integral with $f = \mathcal{J}_{1k}$ we have to assume in addition that $\partial^2 u_\theta^{(0)}/\partial\widehat{r}^2 = \Gamma/(\pi\widehat{r}^3)(1 + \mathcal{O}(1))$. In that case, since $\phi_{1k}^{(1)} \rightarrow 0$ under the conditions formulated below (4.30) (checked via two-fold application of de L'Hospital's rule) we also have $\partial\phi_{1k}^{(1)}/\partial r \rightarrow 0$ and $r\mathcal{J}_{1k} = \mathcal{O}(r^{-2})$ as $r \rightarrow \infty$, and the integral converges.

Finally with $w_{1k}^{(1)} \sim \widehat{r}^{-3}$, the term with slowest decay in \mathcal{I}_{1k} from (4.33) is the one involving the factor $\widehat{r}f_0$, and we have $r\mathcal{I}_{1k} \sim \widehat{r}^{-1}$ for large \widehat{r} and the last integral in (E2) diverges. More precisely,

$$\widehat{r}\mathcal{I}_{1k} = -\frac{I_{1k}}{\widehat{r}}(1 + \mathcal{O}(1)) \quad \text{as} \quad \widehat{r} \rightarrow \infty, \quad (\text{E3})$$

where

$$I_{1k} = \frac{f_0^2}{4\pi^2\rho_0} \frac{\partial}{\partial z} \left(\frac{\rho_0\Gamma^2}{\Theta_1'} (-1)^k \frac{\partial X_k^{(1)}}{\partial z} \right). \quad (\text{E4})$$

By adding and subtracting this dominant behavior we find a decomposition of $\mathcal{L}_u[\mathcal{I}_{1k}](\widehat{r})$ into a logarithmically diverging term plus a finite contribution, viz.

$$\mathcal{L}_u[\mathcal{I}_{1k}](\widehat{r}) = -\frac{\pi I_{1k}}{\Gamma} \ln \widehat{r} + L[\widetilde{\mathcal{I}}_{1k}] + \mathcal{O}(1) \quad (\widehat{r} \rightarrow \infty) \quad (\text{E5})$$

where

$$\widetilde{\mathcal{I}}_{1k} = \mathcal{I}_{1k} + H(\widehat{r} - 1)I_{1k}/\widehat{r}^2 \quad (\text{E6})$$

is the regularized version of \mathcal{I}_{1k} with $H(\cdot)$ the Heaviside step function.

Appendix F. Centreline motion for the barotropic, adiabatic case

We start from the first-order centreline equation of motion in (5.17b), viz.

$$\frac{\partial \mathbf{X}^{(1)}}{\partial t}(t, z) = \mathbf{u}_0^{(1/2)} + \mathbf{X}^{(1)} \cdot (\nabla_{\mathbf{r}} \mathbf{u})_0^{(0)} - \ln \frac{1}{\delta} (\mathbf{k} \times \mathbf{M}_1) + (\mathbf{k} \times \boldsymbol{\Psi}) \quad (\text{F1})$$

where

$$\mathbf{M}_1 = \frac{f_0^2}{8\pi\rho_0\Gamma} \frac{\partial}{\partial z} \left(\frac{\rho_0\Gamma^2}{\Theta'_1} \frac{\partial \mathbf{X}^{(1)}}{\partial z} \right), \quad (\text{F } 2a)$$

$$\Psi = \left(-i L \left[\mathcal{H}_{11} + \tilde{\mathcal{I}}_{11} + \mathcal{J}_{11} \right] + j L \left[\mathcal{H}_{12} + \tilde{\mathcal{I}}_{12} + \mathcal{J}_{12} \right] \right), \quad (\text{F } 2b)$$

furthermore,

$$L \left[\mathcal{H}_{1k} + \tilde{\mathcal{I}}_{1k} + \mathcal{J}_{1k} \right] = R_1 [W T_k] + R_2 \left[\frac{1}{\rho_0} \frac{\partial}{\partial z} (\rho_0 W T_k), \frac{1}{\rho_0} \frac{\partial}{\partial z} (\rho_0 \widehat{W} T_k) \right]. \quad (\text{F } 3)$$

Here we have used (4.25) for the diabatic case, so that

$$w_{1k}^{(1)} = W T_k \quad \text{with} \quad W = \frac{1}{\Theta'_1} \frac{u_\theta^{(0)2}}{r} \left(\frac{u_\theta^{(0)}}{r} + f_0 \right) \quad (\text{F } 4)$$

and where we have defined the abbreviations

$$\widehat{W} = \frac{\Gamma^2}{4\pi^2} \frac{f_0^2}{\Theta'_1}, \quad T_1 = -\frac{\partial X_C^{(1)}}{\partial z}, \quad T_2 = +\frac{\partial Y_C^{(1)}}{\partial z}, \quad (\text{F } 5)$$

and

$$R_1[W] = \frac{\pi}{\Gamma} \int_0^\infty r \frac{\partial}{\partial r} \left(r W \frac{\partial u_\theta^{(0)}}{\partial z} \right) dr, \quad (\text{F } 6a)$$

$$R_2[W, \widehat{W}] = \frac{\pi}{\Gamma} \int_0^\infty \left(-r^2 [\zeta^{(0)} + f_0] W + \frac{H(r-1)}{r} \widehat{W} \right) dr \quad (\text{F } 6b)$$

$$\begin{aligned} & - \frac{\pi}{\Gamma} \int_0^\infty \left(r^2 \frac{\partial \zeta^{(0)}}{\partial r} \right) \frac{\partial}{\partial r} \left(r \int_0^r \frac{1}{\bar{r}^3} \left[\int_0^{\bar{r}} \bar{r}^2 W d\bar{r} \right] d\bar{r} \right) dr \\ & + \frac{\pi}{\Gamma} \int_0^\infty \left(r^2 \frac{\partial \zeta^{(0)}}{\partial r} \right) dr \int_0^\infty \frac{1}{\bar{r}^3} \left[\int_0^{\bar{r}} \bar{r}^2 W d\bar{r} \right] d\bar{r} \end{aligned} \quad (\text{F } 6c)$$

Finally, we find

$$\Psi = \tilde{A} \frac{\partial^2 \mathbf{X}^{(1)}}{\partial z^2} + \tilde{B} \frac{\partial \mathbf{X}^{(1)}}{\partial z} \quad (\text{F } 7)$$

where

$$\tilde{A} = R_2 [W, \widehat{W}], \quad \text{and} \quad \tilde{B} = R_1 [W] + \frac{1}{\rho_0} \frac{\partial}{\partial z} (\rho_0 R_2 [W, \widehat{W}]) \quad (\text{F } 8)$$

Collecting all results from this section we obtain the equation of motion for $\mathbf{X}^{(1)}$ as announced in (6.2), (6.3).

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