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MODELING OF HYDROGEOLOGIC CONDITIONS AND GROUND WATER QUALITY AT AN OIL WELL RESERVE PIT, RICHLAND COUNTY, MONTANA

by

Scott M. Payne

B.S., Northland College, 1985

Presented in Partial Fulfillment of the Requirements for the Degree of Masters of Science

UNIVERSITY OF MONTANA

1989

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august 15, 1989 Date

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MODELING OF HYDROGEOLOGIC CONDITIONS AND GROUNDWATER QUALITY AT AN OIL WELL RESERVE PIT IN RICHLAND COUNTY, MONTANA

Director: Dr. William W. Woessner WWW8-7-89

This study examined groundwater contamination in a shallow aquifer of the Yellowstone River Valley resulting from storage of oil-field brine and drilling fluid additives in a reserve pit. Results show brine seepage from the pit enters the shallow aquifer due to inadequate pit reclamation. Reserve pit mud is contaminated with ions and metals at concentrations one to three orders of magnitude greater than federal drinking water standards. Only certain ions and metals present in the pit mud reach the underlying shallow aquifer system at concentrations much less then in the pit mud.

Chloride concentrations up to 2800 mg/l are present in the shallow aquifer during spring and summer months as a result of the high water table intersecting the base of the reserve pit. In addition, precipitation generates wetting fronts which leach pit contaminants into the shallow aquifer. Chloride seepage declines in the fall through winter months as a result of a lowering water table and less recharge.

After entering the shallow aquifer, chloride migrates downgradient and enters a nearby return flow irrigation ditch. Maximum chloride concentrations measured in the ground water decrease from 2800 mg/l to approximately 600 mg/l at a distance of 350 feet (115 m) from the pit. High chloride concentrations are vertically limited to the upper five to six feet (2 m) of the 15 to 20 foot (6 m) thick sand and gravel aquifer. Lack of vertical mixing in the aquifer is probably a result of a high hydraulic conductivity or layering of strata in the horizontal direction.

Other constituents in the shallow aquifer detected at above background concentrations and attributed to brine seepage from the reserve pit are boron, lithium, barium, strontium, titanium, zinc, beryllium, calcium, magnesium, sodium, potassium, manganese, bicarbonate, sulfate, and nitrates.

Surface electromagnetic induction conductivity (EM) and resistivity were used to delineate the extent of brine contamination. Both methods produced similar results and delineated the brine contaminated area. EM was less cumbersome and time consuming to use compared to resistivity.

Transient groundwater flow modeling for one year of hydraulic head data using the PLASM 2-D numerical model successfully simulated the actual hydraulic heads. Solute transport modeling with Random-Walk successfully simulated chloride concentrations at some monitoring well locations but not at others. The simplicity of the groundwater flow model in conjunction with Random-Walk limitations yielded inaccurate simulated chloride concentrations.

ACKNOWLEDGMENTS

This project is the result of many individual's time and effort. I am deeply indebted for their assistance, knowledge and information. Residents living in the study area were helpful and cooperative throughout the study. I appreciated their patience and willingness to permit monitor well drilling and water quality sampling on their property.

The following organizations provided technical help and data: The Montana Bureau of Mines and Geology, Richland County, the Montana State University Agricultural Experiment Station in Sidney, the U.S. Department of Agriculture in Sidney, the U.S. Soil Conservation Service in Sidney, the U.S. Geological Survey in Helena and The U.S. Forest Service in Missoula. It is my opinion that the staff members of these public agencies provided outstanding assistance and were accessible and helpful in all of my requests.

A special thanks is due for Randy Skov, my field technician, who braved winter days to collect water quality data and water levels. Joe Donovan of the Montana Bureau of Mines and Geology also deserves a special thanks for his help in the field.

Key individuals in the project were my thesis committee members: Drs. Woessner, Thompson, Potts, and Sonderegger, and project supervisor Jon Reiten of the Montana Bureau of Mines and Geology.

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Chapter I

INTRODUCTION

Groundwater contamination from an oil well reserve pit was studied in a shallow aquifer of the Yellowstone River Valley, eastern Montana. Significant contamination by chloride and other ions and metals was documented. The source of contamination is a buried reserve pit containing extremely elevated concentrations of the ions and metals. Computer modeling was used to simulating groundwater flow and solute transport of chloride contamination through the shallow aquifer. The following discussion will familiarize the reader with a reserve pit. That discussion is followed by a description of oil production in the Williston basin, goals and objectives of this study, location and description of the study site, and a discussion of previous work done at the study site.

Drilling fluids and brine water are contained in a reserve pit during drilling of oil and gas wells in the oil producing areas of the United States. Pit dimensions vary, but average 150 to 200 feet long (46 to 61 m), 60 to 70 feet wide (18 to 21 m), and 8 to 12 feet deep (2 to 4 m). Most reserve pits are lined with a synthetic or clay liner to protect underlying groundwater from brine seepage, however, in some instances pit liners are not used depending on the geologic setting and location. Contamination of shallow aquifers by reserve pits has been documented in Montana,

Oklahoma, Alabama, North Dakota, Texas, Arkansas, Wyoming, Ohio, and Kansas (Beal, 1986; Beal and others, 1987; Collins, 1971; Dewey, 1984; Fryberger, 1975; Hicks, 1983a and 1983b; Knowels, 1965; Leonard, 1965; McMillion 1965; Murphy and Kehew, 1984; Murphy et. al., 1985, Novak and Yoram, 1988).

Techniques used to reclaim reserve pits usually involve removal of the brine water from the pit (disposed via injection well) and on site disposal of the more viscous drilling fluid or pit mud. On site mud disposal consists of digging trenches radially away form the reserve pit and pushing the pit mud into the trenches. Backhoes used to dig the trenches destroy the integrity of the existing pit liner. After the mud is pushed into the trenches, the site is resurfaced to ground level.

Surface recharge at a reclaimed reserve pit site may generate leachate and contaminate the underlying groundwater. In addition, high groundwater may come in contact with the bottom of the buried reserve pit or trenches and contaminate the groundwater.

Presently, drilling fluid wastes are designated as nonhazardous under the RCRA (Resource Conservation and Recovery Act). According to Kemblowski and Deeley (1987), this exemption is primarily based on the "...low toxicity of drilling fluid wastes". However, reserve pit studies in the last five to six years have shown that oil-field brine (Table 1) and drilling fluid additives (Appendix A) contain products

that are a threat to water, soil and vegetation (Hicks 1983a).

Oil production in the Williston Basin

As of 1982, there were over 4000 oil wells in the Williston Basin (Dewey, 1984). Chloride concentrations in brine and drilling fluids (Murphy and Kehew, 1984) are typically 300,000 ± 20,000 mg/l in north central North Dakota and 100,000 to 200,000 mg/l in eastern Montana (Dewey, 1984). Pit reclamation in the Williston Basin most often includes removal of brine water and on site mud disposal in excavated trenches (Dewey, 1984). Contamination of groundwater by brine seepage is reported at all reserve pit sites studied in the Williston Basin (Dewey, 1984; Murphy and Kehew, 1984; Hicks, 1983a and 1983b; and Beal, 1986). Dead or stunted vegetation has been documented at a number of these brine contaminated sites (Hicks, 1983a; Dewey, 1984). Many sites outside of Montana and North Dakota have also experienced dead or stunted vegetation in addition to contaminated drinking water and surface water in streams and rivers (Leonard, 1965; McMillion, 1965; Knowles, 1965).

The processes that leach reserve pit wastes into the shallow groundwater are poorly understood and regional impacts to eastern Montana groundwater are relatively unknown.

Goals and objectives

Goals of this study are: 1) to characterize the mechanism(s) controlling reserve pit waste migration into the shallow aquifer, and if the process(es) is continuous or

TABLE 1. Comparison of dissolved solids in seawater and oil field brine (USEPA, 1972).

Element	Scowater, mg/l. Oil-field brine, mg	
Sodium	10,600	12,000 to 150,000
Potassium	380	30 to 4,000
Lithium	0.2	1 to 50
Rubidium	0.12	0.1 to 7
Cesium	0.0005	0.01 to 3
Calcium	400	1,000 to 120,000
Mognesium	1,300	500 to 25,000
Strontium	8	5 to 5,000
Barium	0.03	0 to 1,000
Chlorine	19,000	20,000 to 250,000
Bromine	65	50 to 5,000
lodine	0.05	1 to 300

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episodic; 2) to determine if EM (electromagnetic induction conductivity) can be used to identify groundwater contaminated from brine seepage; 3) to estimate the aquifer parameters necessary to conduct a hydrogeologic model for the reserve pit area and predict contaminant dispersion behavior; 4) to characterize the chemistry of brine contaminated water and soil and 5) to suggest possible remediation measures. Objectives related to meeting the above goals consisted of measuring/determining the following parameters:

- The soil and groundwater conductance/resistivity using EM and resistivity surface geophysics;
- Groundwater and surface water quality;
- 3) The extent of brine contamination including related metal contamination;
- 4) The hydraulic properties of the unconfined aquifer;
- 5) The rate and direction of groundwater flow at the site;
- 6) The dispersion and dilution of contaminants in the unconfined aquifer;
- 7) Amount of chloride loaded into the groundwater;
- 8) Regional loading of chloride into the Yellowstone River;
- 9) The reclamation technique best suited to reverse soil contamination on site;

Location and description of study site

The Iverson site is located approximately 3 miles north of Sidney, Montana in Section 15 of Township 23 north, Range 59 east in Richland County (Figures 1 and 2). The site is





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Figure 2.Topography of Section 15, T.23N., R.59E. and location of study site (adapted from Dewey, 1984).

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situated on a terrace deposit on the western side of the Yellowstone River ("Crane Creek Gravel"), one mile (1.7 km) east of the western valley edge. The valley edge is composed of the Tongue River member of the Fort Union Formation which is 300 to 500 feet (120 m) higher in elevation than the valley bottom. The Yellowstone River Valley bottom is fairly level, except for a series of low terrace deposits.

At the study site an operating oil pumper and buried reserve pit are present (Figure 3). The reserve pit contained brine water and drilling mud, and was reclaimed in the summer of 1982 using the trenching method (Dewey, 1984). Brine water was removed from the pit and the drilling mud was disposed in both the reserve pit and trenches. According to Dewey (1984), initial groundwater contamination from brine seepage occurred within approximately 20 days after the pit was reclaimed. Wheat, corn, sugar beets, and alfalfa are grown on surrounding land and a return flow irrigation ditch abuts the west side of the oil pumper area. A number of small irrigation channels supply water for flood irrigation near the site.

<u>Climate</u>

The climate is semiarid, characterized by cold, dry winters, moderately hot and dry summers, and cool, dry falls (Slagle, 1984a). Winters are often interrupted by warming trends, with summers dominated by hot days and cool nights. January is generally the coldest month and July the warmest. Glendive, Montana has a 14.9° F average January temperature





and a 74.0° F average temperature for July, -10°C and 23.3°C respectively (Slagle, 1984). Average annual precipitation at Sidney is approximately 14 inches/year (35.6 cm) with about 65 percent of the precipitation falling from May through August, and June being the wettest month (Slagle, 1984a).

<u>Regional hydrogeology of Richland County, Montana</u>

Groundwater is present in eight major aquifer systems in Richland County from the Paleozoic Madison Group to the Quaternary alluvium. Aquifers present within 200 to 300 feet (61-91 m) of ground surface in Richland County are the Tongue River Member of the Tertiary Fort Union Formation and the Quaternary glacial, terrace and alluvial deposits. The water table in these aquifers usually reflects the land surface topography (Slagle, 1984a). Deeper aquifers, below 300 feet (91 m), have a regional flow direction towards the Yellowstone and Missouri Rivers.

Groundwater recharge generally in the occurs topographically high areas with discharge areas in the valley bottoms. Snow melt and precipitation are the major contributors of recharge in eastern Montana (Torry and Kohout, 1956), however, irrigation in Yellowstone River Vallev large quantities of water to the contributes shallow aquifer(s). Groundwater recharge is greatest in the spring when snow melt and precipitation are at their peak and irrigation begins (Torry and Kohout, 1956).

Prior to this study, no aquifer property data were

available for the Quaternary aquifer systems in Richland County, and only laboratory estimates of hydraulic conductivity were available for the Tongue River Member. Croft (1985) estimated hydraulic conductivities of the Tongue River claystone to be 0.15 to 0.099 ft/day (8.45 x 10^{-6} to 5.64 x 10^{-6} cm/sec).

A detailed discussion on the regional hydrogeology is in Appendix B.

Previous work

The Iverson site was previously studied by Michelle Dewey of the University of Montana from August 1982 through June 1983. Dewey determined that a shallow sand and gravel aquifer underlies the site, average depth to the groundwater is nine feet and groundwater flows towards the north.

Dewey's chemical analyses and surface resistivity surveys led her to conclude that: 1) brine is seeping into the shallow groundwater from the reserve pit which produces increased levels of sodium and chloride; 2) the reclamation techniques used to reclaim reserve pits are inadequate to inhibit brine migration; 3) surface electrical resistivity is a good indicator of possible brine contamination at oil well sites; and 4) contaminants may be moving as fast as 22 feet per day through the shallow aquifer.

Numerous other authors have studied reserve pit brine migration in the oil producing areas of the United States. These studies demonstrate that chloride and other ions often contaminate shallow aquifers and surface water systems comparable to the reserve pit contamination in the shallow aquifer of the Williston Basin (Powell and others, 1963; Knowles, 1965; Fryberger, 1972 and 1975; Todd and McNulty, 1976; Vander Leeden and others, 1975; Latta, 1963; Leonard, 1964; Bryson and others, 1966; Kreiger and Henderickson, 1960; Miller and others, 1977; McMillion, 1965; Scalf and others, 1975; Shaw, 1966; Pettyjohn, 1971, 1973, and 1975; McMillion, 1965; Payne, 1966; Miller; 1980; Baker and Brendecke, 1983).

Chapter II

METHODS

Analytical methods were selected to determine the movement and chemical composition of groundwater and surface water. Monthly soil moisture and groundwater quality sampling methods were selected to determine how and when contaminants enter the groundwater. Computer modeling methods were selected to simulate groundwater flow and solute transport in two dimensions.

Field methods - Geophysical techniques

Resistivity measurement

A Soiltest R-60 resistivity meter and Wenner electrode configuration with electrode a-spacings of 10, 20 and 30 feet (3.48 m, 6.1 m, and 9.57 m, respectively) were used to determine soil and groundwater resistivity. Figure 4 shows the location of resistivity stations. Standard operating procedures as described by Dejong and others (1979) and Soiltest (1976a and 1976b) were used to determine soil and groundwater resistivity.

EM measurement

A Geonics EM34-3 conductivity meter with 10, 20 and 40 metre coil spacings (32.8 ft, 65.6 ft, and 131.23 ft respectively) was used to determine soil and groundwater conductivity in both the horizontal and vertical EM modes. Location of EM stations are shown in Figure 5.

Use of two coil configurations (or modes) with the EM





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STATION MAP FOR 10 METRE EM



STATION MAP FOR 40 METRE EM



Figure 5. EM gird (10, 20 and 40 metre coil spacings).

equipment measures different depths of soil conductivity. Thehorizontal EM mode (vertical dipoles) measures soil conductivity to approximately 1.5 times the length of the coil spacing and is relatively insensitive to surface soil conductivity (McNeil, 1980c). The vertical EM mode (horizontal dipoles) measures soil conductivity from the surface to approximately .75 times the length of the coil spacing (McNeil, 1980c).

EM data were interpreted using standard procedures described by McNeil (1980a, 1980b, 1980c, 1985 and 1986).

Monitoring wells and lysimeters:

Figure 3 shows the location of monitoring wells and details well lysimeters. Appendix C and lysimeter The monitoring wells were located on the basis construction. of EM and resistivity results. Readings of large conductivity values and small resistivity values at geophysical survey points were interpreted to represent areas of reserve pit Monitoring wells were installed using the contamination. MBMG (Montana Bureau of Mines and Geology) Mobile-Drill 50 drill rig equipped with 6.5 inch OD (outer-diameter)/3.25 inch ID (inner-diameter) HSA (hollow-stem-auger) flights. One well, MW 15p, was installed using 10 inch OD/6.25 inch ID HSA flights.

The depth of bore holes varied from 10 to 37 feet (3 to 11 m), however, no monitoring wells were installed below 23 feet (7 m). All monitoring wells, except MW 15p, are 2-inch ID diameter schedule 20 PVC with either a 20 or 30 slot (0.020 or 0.030 inch) well screen. Well MW 15p is a 4.5 inch ID schedule 40 PVC well with a 20 slot (0.020 inch) well screen. The wells are completed with a natural sand and gravel (hole collapse) and a surface bentonite plug. Both single and nested well sites were installed. Nested well sites consisted of two to three monitoring wells completed at different depths a few feet apart. Each well was developed with a hand bailer or pumped with a small surface pump until clear water was obtained. Soil samples and logs were collected from auger cuttings and a geologic log was prepared for each hole (Appendix D).

Monitoring wells installed in the pit area were completed with two bentonite seals. A second bentonite seal (below the surface bentonite) was installed below the bottom of the reserve pit mud to inhibit vertical brine migration along the well casing.

Lysimeters were installed from 3.5 to 6 feet (1 to 2 m) below ground surface. Lysimeters are 1.5 inch diameter schedule 26 PVC with a ceramic cup at the lower end of the casing. The lysimeters are completed with a saturated silica flour surrounding the ceramic cup, natural fill above the silica flour, and a surface bentonite seal.

Water levels were measured with either a steel tape or electric M-scope from a top-of-casing datum surveyed by the United States Soil Conservation Service in Sidney, Montana. Water levels were measured monthly from July, 1987 to August 1988. Water level data were used to determine head at each well site and areal groundwater flow direction.

A stilling well was installed with a 30 day Stevens continuous water level recorder on the return flow irrigation ditch at site Ditch 1 (Figure 3). Two additional 30 day Stevens continuous recorders were installed on monitoring wells MW 1 and MW 3.

Aquifer testing:

The hydraulic conductivity of the shallow, unconfined sand and gravel aquifer was estimated from an aquifer test conducted at well MW 15p. The shallow aquifer was stressed using a 3 HP surface pump which sustained a constant discharge rate of 30 gpm (gallons/minute) for 4.5 hours. Discharge water was pumped through a pipe approximately 150 feet away to the southwest of the pumping well onto an alfalfa field. Drawdown and recovery were measured in the pumping well (MW 15p) and three monitor wells, 53 feet, 82 feet and 118 feet (16 m, 25 m and 36 m) from the pumping well (wells MW 6, MW 19s, and MW 20s, respectivley). Drawdown verses time data were first analyzed by assuming steady-state conditions The data were also evaluated assuming (Driscoll, 1986). transient conditions with delayed yield and partial well penetration (Boulton 1954a, 1954b, 1964, and Stallman, 1965).

<u>Water quality sampling:</u>

Measured field parameters include corrected SC (specific

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conductance), temperature, pH, and chloride. Chloride concentrations were measured using Quantabs (Appendix E), a small capillary tube that function as a chloride titrator. A single Quantab is lowered into a cup of water, after which water moves up the capillary passage on the Quantab to discolor the top. A value corresponding to the distance the sample water moved up the capillary passage is read off the side of the Quantab and converted into mg/l or percent chloride in solution (Appendix E). Field data were collected monthly and input into a computer data base.

Water samples for chemical analyses were collected on a quarterly basis. Sampling procedure included bailing water from monitoring wells until pH stabilized and rinsing sample bottles three times with sample water. Four samples were collected from each monitoring well site including one untreated, one filtered, and two filtered and acidified samples. Acidified samples were preserved with 2.5 milliliters of nitric acid per 500 ml of sample water. The filtered samples were pumped through a 0.45-micron filter. Decontamination of sampling equipment was accomplished with three rinses of deionized water.

Analyses for major ions and metals were done at the MBMG laboratory. Metals were also analyzed at the University of Montana Geology Department. Field quality control was checked with one deionized water sample filtered and preserved in the field, and analyzed at the MBMG laboratory. Sample water was

stored in a cooler over ice for delivery to the MBMG laboratory in Butte and the University of Montana Geology Department in Missoula.

Lysimeters were also sampled quarterly. Vacuum was drawn on lysimeters five to 30 days prior to sampling, after which, the soil moisture was sucked out using a vacuum pump. Sample bottles were rinsed with deionized water. The sample water was filtered in the field and sometimes acidified in the field if sufficient sample water was available. Sample bottles were stored in coolers over ice and delivered to the MBMG laboratory and the University of Montana Geology Department.

Laboratory methods - Chemical analyses:

Groundwater and soil moisture samples were analyzed at the MBMG laboratory for following ions and constituents including: calcium, magnesium, sodium, potassium, iron, manganese, silica, bicarbonate, carbonate, chloride, sulfate, nitrate, fluoride, phosphate, bromide, dissolved solids, sum of constituents, specific conductance, pH, hardness, alkalinity, Ryznar stability, Langlier saturation, and sodium absorption ratio. All MBMG laboratory procedures were EPA and USGS certified procedures (Reiten, 1988). Metals analyzed at the MBMG laboratory on an ICPS (inductively coupled plasma silver, aluminum, arsenic, boron, spectrometer) were: beryllium, cadmium, chromium, copper, lithium, molybdenum, barium, lead, nickel, strontium, titanium, vanadium, zinc, zirconium, and selenium.

Metal analyses at the University of Montana were also done on a ICPS. Groundwater, soil moisture, and soil samples were tested for: aluminum, arsenic, calcium, cadmium, copper, iron, potassium, magnesium, manganese, nickel, sodium, lead, phosphate, antimony, and titanium. Soil samples were acid extracted for total digestible metal content (Moore, 1988). Metal scans and sample preparation were accomplished using standard laboratory techniques done in accordance with USGS water analysis standards which included a 10 percent rerun of samples as duplicates (Moore, 1988).

Sieve analysis:

Sieve analyses were performed on three different samples collected from monitoring well cuttings. The samples were sieved at the United States Forest Service Materials Testing Lab in Missoula, Montana. Standard laboratory techniques were used to determine the grain size distribution.

Aquifer porosity was determined by saturating a volume of oven dried aquifer material and measuring the volume of water needed to fill the empty void spaces.

X-ray diffraction:

X-ray diffraction was used to identify the <2-micron size fraction clay mineralogy of mud samples taken from auger cuttings while drilling through the reserve pit. The samples were mixed with deionized water and centrifuged to separate the <2-micron size fraction. The <2-micron size fraction was put on glass slides as random, oriented, and glycol saturated clay particles for x-ray diffraction (Jennings and Thompson, 1986). Standard X-ray diffration techniques were used to identify the clay mineralogy of the pit mud.

<u>Computer modeling:</u>

The 2-dimensional PLASM groundwater flow model developed by Prickette and Lonnquist (1971) was used to model groundwater flow. Random-Walk, developed by Prickette, Naymik, and Lonnquist (1981), was used to model solute transport. PLASM is a finite difference model which can be used to simulate non-steady flow in a heterogeneous aquifer under unconfined and non-leaky, and leaky-confined conditions. PLASM estimates head using the alternating direction implicit numerical method to solve a set of finite difference equations. Input data will be discussed in detail in Chapter III.

Random-Walk is a 2-dimensional model that can simulate contaminant transport in an unconfined or confined aquifer (Prickette and others, 1981). The model takes into account groundwater velocity, convection, dispersion, and chemical reactions. The model simulates movement of a solute using a particles in a cell method, transported for a given time schedule. Particles are randomly dispersed according to model inputs by a statistical technique. The model has the ability to install sinks or sources, and can map solute as concentration or the number of particles for a given cell. The version of Random-Walk used in this study is linked with the PLASM groundwater flow model. Head data generated with PLASM is converted into a velocity vector field with a preprocessor for each simulated month. The velocity vector field is then linked to the Random-Walk model to move solute in accordance with the modeled potentiometric surface and dispersion coefficients.
Chapter III

RESULTS AND DISCUSSION

This chapter presents all data collected in this study. A short discussion will follow the data in each section. <u>Site hydrogeology</u>

Soil and fill (material used to cover the pumper area and roads) cover the top three feet (1 m) of land surface of the study area. From three to 26 feet (1 to 8 m) Quaternary sand and gravel of the oldest terrace deposit ("Crane Creek Gravel") of the Yellowstone River Valley is present. Included in the sand and gravel are local zones of well sorted sand, gravel, thin lenses of clay, layered silt and sand, layered clay and sand, and pit mud in the reserve pit area (see geologic cross sections in Appendix D). The depositional environment of the terrace deposit was apparently a fairly high energy system since coarse gravel was deposited with Terrace deposition apparently took place during an sand. interglacial time since till deposits are present stratigraphically above and below the terrace deposits about three miles (5 km) south of the study site (Prichard and Landis, 1975).

A sieve analysis on the sand and gravel material (Appendix D) shows an even distribution of particles, from fine sand to coarse gravel. This material is the dominant aquifer material encountered in the monitoring well bore holes. A sieve analyses on a sample collected from eight to 20 feet (2.5 to 6 m) at monitoring well MW 12 showed that 85% of the sample retained was sand size. A sieve analyses on a sample collected from 14 to 33 feet (4.5 to 10.5 m) at monitoring well MW 13 showed 75% of the sample retained was gravel size. Results of the sieve analyses indicate that the aquifer is relatively coarse grained and little silt and clay are present.

The Tongue River Member of the Fort Union Formation underlies the terrace deposits at approximately 26 to 28 feet (8 m) below ground surface. At the study site the Tongue River Member consists of a dense, low permeability grayish clay with sandy intervals.

The shallow aquifer is restricted to the sand and gravel terrace deposits. The shallow aquifer has a seven to 10 foot (3 m) unsaturated zone and a saturated thickness of about 17 feet (5 m). Below the shallow aquifer, clay of the Tongue River Member separates the shallow unconfined aquifer and the deeper confined aquifer. The upper clay of the Tongue River Member has a relatively low hydraulic conductivity, estimated to be 0.15 to 0.099 ft/day (8.45 X 10⁻⁶ to 5.64 X 10⁻⁶ cm/sec) (Croft, 1985), which probably inhibits groundwater movement between the shallow and deep aquifers. Domestic wells in the study are completed in water bearing coal and sandy intervals of the Tongue River Member 35 to 45 feet (13 m) below land surface (as determined from domestic well logs).

In general, groundwater flow direction (Figure 6 and

Appendix F) is towards the northeast with some variations in flow direction along the return flow irrigation ditch. The ditch is recharged by groundwater in the spring, summer and fall and the water table is in contact with the ditch since it is above the ditch water level and a visible seepage face is present in the ditch channel. Estimated ditch flow in the summer months is approximately three to four cfs.

In the winter, groundwater flow (Appendix F) is parallel to the ditch, and water in the ditch apparently leaks into the underlying shallow aquifer since the measured ditch water levels are above the water table. The ditch bottom is comprised of mud and silt, and permeability of these sediments is estimated to be relatively low. The low permeability and small winter ditch flow (approximately .5 cfs) indicate that the ditch is most likely leaking water to the shallow aquifer and forming a small groundwater mound.

The Yellowstone Irrigation District shuts down the irrigation canals in the winter months, yet flow is present in the return flow ditch. The source of winter ditch water is probably from upgradient reaches where the water table intersects the ditch causing groundwater to discharge into the ditch.

Precipitation and recharge

Precipitation and free water evaporation data collected at the Montana State University Agricultural Experiment Station in Sidney about two miles from the study area are in



Figure 6. Groundwater flow map (spring and summer).

Figure 7 and Appendix G. Precipitation is lower in the spring and summer of 1988 compared to the spring and summer of 1987. In addition, evaporation is greater in 1988 compared to 1987. The higher evaporation and lower precipitation in 1988 indicate there was less surface recharge available to move pit contaminants compared to 1987. Since the reserve pit area is not flood irrigated, precipitation is the only surface recharge factor for leaching brine to the water table.

Annual water level fluctuation at well MW 3 is about four feet in the 1987-1988 year (Figure 8). The water table rise and fall is coincident with the start of spring flood irrigation in May and irrigation shutdown in November surrounding the site. The crops grown on the surrounding farmland (alfalfa, corn and sugar beets) require 18 to 24 inches (46 to 61 cm) of water during the growing season (SCS, 1987), which is significantly greater than the average annual precipitation of 14 inches/year (36 cm). The irrigation water used to supplement the crops is an important source of recharge to the shallow aquifer.

The peak 1988 water table position (Figure 8) matches the peak 1987 water table position, yet precipitation is less and free water evaporation is greater in 1988 compared to 1987. This suggests that more flood irrigation water was applied in 1988 compared to 1987 to counter drought conditions.

All monitoring wells and ditch locations have a similar annual water level fluctuation (Appendix H), indicating water



Figure 7. Annual precipitation and evaporation from June 1987 to June 1988 (Montana State University Agricultural Experiment Station Sidney, Montana).

CENTIMETERS



Figure 8. Annual water table fluctuation from August 1987 to July 1988.

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level response to recharge is similar across the entire site.

Hydraulic conductivity and specific yield

One constant discharge aquifer test was used to estimate K (hydraulic conductivity) using both Boulton's partial penetration and delayed yield equations (Lohman, 1978). Delayed yield was evident in the pumping well (MW 15p) and one monitoring well (MW 6). The delayed yield method was employed to estimate K at these locations. Delayed yield was not evident in the other monitoring wells (MW 19s and MW 20s) and the partial penetration method was used to estimate K at these locations.

One foot of drawdown was measured at the end of the 4.5 hour aquifer test at the pumping well MW 15p (pumping 30 gpm), and only a few tenths of a foot drawdown were measured at the monitoring wells (Appendix I). Limited drawdown made K interpretation difficult. Hydraulic conductivity estimates ranged from 1,500 gal/day/ft² (7 x 10^{-2} cm/sec) to 25,000 gal/day/ft² (1.0 cm/sec) depending on the well location and whether the delayed yield or the partial penetration method was employed. K is probably close to 4200 gal/day/ft² (.2 cm/sec) since the most accurate simulated drawdown in an aquifer test was modeled using this value (to be discussed later). Assuming a 17.5 feet thick aquifer, transmissivity is 73,500 gallons/day/ft.

Hydraulic conductivity was also estimated by using the Theim steady-state equations (Driscoll, 1986) since drawdown

at the end of the aquifer test was small. Theim K estimates averaged one order of magnitude below K estimates calculated with the delayed yield and partial penetration methods. Transient K calculations probably represent a better estimate of K since steady-state conditions were not present at the end of the stress period.

A representative Sy (specific yield) for an unconfined sand and gravel aquifer was not calculated with the above methods. All calculated estimates were a number of orders of magnitude too low. Specific yield was estimated from 13 documented Sy values for sand and gravel aquifers (Turcan, 1963). Specific yield values ranged from 0.2 to 0.35, and .2 was chosen since it is a conservative estimate. Porosity of a repacked aquifer sample was determined to be .25 in the laboratory. To calculate groundwater velocity at the pit area sites 7s) Ι used the following (wells MW 3 and MW relationship:

V = KI where V = velocity (ft/day)
7.48 n K = hydraulic conductivity
(4200 gal/day/ft²)
I = gradient (.0026 to .001
depending on which month)
n = porosity (.25)
7.48 = conversion factor for
gallons to cubic feet

and calculated the following monthly groundwater velocities:

87'	July:	5.6 ft/day	(1.70 m/day)
	August:	5.3 ft/day	(1.62 m/day)
	September:	4.7 ft/day	(1.43 m/day)
	October:	4.5 ft/day	(1.37 m/day)
	November:	3.2 ft/day	(.97 m/day)
	December:	2.9 ft/day	(.88 m/day)

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88'	January:	2.9	ft/day	(.88 m/day)
	February:	2.8	ft/day	(.85 m/day)
	March:	2.6	ft/day	(.79 m/day)
	April:	2.4	ft/day	(.73 m/day)
	May:	5.6	ft/day	(1.70 m/day)
	June:	5.9	ft/day	(1.79 m/day)

Soil chemistry

Vadose zone soil analyses results are in Table 2 and Appendix K. Table 2 data represents the total acid digestible metal content for sodium, calcium and lead concentrations in samples collected from three reserve pit locations and three samples collected downgradient of the pit. Only sodium, calcium and lead are presented in the Table 2 since the other ions tested are at background concentrations. Chloride was not analyzed.

Table 2. Soil analyses (* less than detection)

<u>Sample site</u>	Depth	<u> </u>	<u>% Ca</u>	pom Po
PITMUD 1	6-8 ft.	1.9	6.4	30.0
PITMUD 2	6-8 ft.	2.1	7.7	44.0
MW 3	5-7 ft.	1.6	3.8	21.9
MW 4	6-12 ft.	2.3	4.6	16.8
MW 7	2-7 ft.	0.8	2.6	*
MW 7 (DUPLICATE)	2-7 ft.	0.9	2.9	*
FT. UNION	28 ft.	0.7	7.8	*

Sodium, calcium, and lead results are elevated in pit samples PITMUD 1, PITMUD 2, and MW 3, and downgradient of the site at MW 4. Analyses indicate the pit mud results are approximately 50 to 75% greater than background analyses results (MW 7 and FT.UNION), with the exception of calcium in the Fort Union sample. MW 4 sample site is approximately 100 feet north of the pit area. The apparent soil contamination at monitoring site MW 4 is probably from pit mud being pushed into a reclamation trenches extending from the reserve pit.

Groundwater and vadose water chemistry

Field measurements

Specific conductance (SC) and Quantab data (chloride concentrations) are in Figures 9 and 10, and Appendix J. Figure 9 shows Quantab and SC results for nested well site MW 4. The highest chloride and SC levels are present in upper four feet of the aquifer (MW 4s). Intermediate chloride and SC levels are present from four to eight feet below the water table (MW 4m). Background or near background chloride and SC levels are present in the lower half of the aquifer from eight feet below the water table to the top of the Fort Union Formation (MW 4d). Figure 9 also shows MW 4s chloride and SC levels in the summer of 1988 are less compared to chloride and SC levels in the summer of 1987.

Figure 10 shows Quantab and SC levels for monitoring wells MW 7s and MW 4m. Both wells are completed in the upper half of the shallow aquifer where brine contamination concentrates. Chloride and SC levels are significantly less in MW 7s than in MW 4m, and the peak timing of chloride and SC levels are approximately 2.5 to three months behind in MW 7s (which is down gradient of the MW 4 site). Contaminant loading in the spring and summer of 1988 appears to be less compared to the summer of 1987. Lower chloride and SC levels in 1988 are probably a result of less surface recharge and in



Figure 9. Annual chloride concentration and specific conductance at MW 4 nest.



Figure 10. Annual chloride concentration and specific conductance at MW 4m and MW 7s.

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turn less brine seepage into the underlying aquifer.

Field data indicate the brine contaminants tend to disperse near the top of the shallow aquifer. The lack of vertical contaminant migration suggests that K may be greater in the horizontal direction compared to the vertical direction and/or an upward flow component is keeping the pit contaminants near the top of the aquifer.

An approximation of groundwater velocity between the two well sites in Figure 10 indicates that contaminants move through the aquifer at approximately 3.3 ft/day to 4 ft/day. This was estimated by evaluating the time required for a chloride or SC peak at MW 4s to be measured at MW 7s. Figure 10 show a 2.5 to three month lag time between chloride and SC peaks which approximates the above velocities since the two sites are about 250 feet (80 m) apart.

Contaminants apparently enter the shallow aquifer by seepage from surface recharge and from contact between the water table and the pit mud. Monitoring well logs (Appendix D) show the that pit mud extends to a depth of about nine feet (3 m) below land surface at the pit area. The water table fluctuates from seven to ten feet (2 to 3 m) below land surface at MW 3. This demonstrates that the groundwater intersects the pit material. The water table apparently saturates the pit mud and in turn pit contaminants are released into the aquifer.

In 1987, pit wastes probably entered the shallow aquifer

as a result of both seepage caused from precipitation and groundwater saturating the pit mud. Wetting fronts were less frequent in 1988 due to drought conditions, which apparently caused the lower chloride concentrations in the groundwater.

Laboratory analyses

Tables 3 and 4 show suspect brine contaminants in the shallow aquifer. Groundwater contamination is indicated by annual changes in ion and metal concentrations above background concentrations. This change is characterized by high spring and summer concentrations, lower fall concentrations and winter very low or background concentrations. Background concentrations are present under the pit area in the winter months indicating contaminant loading temporarily stops. Up gradient control wells (MW 2 and MW 21) have lower concentrations compared to wells in and downgradient of the reserve pit year round.

Pit soil moisture analyses at lysimeter L3 indicate the pit material is contaminated with all metals and ions tested (except titanium and zinc) at concentrations sometimes a number of orders of magnitude above those in the groundwater.

Chloride and sodium are the two most elevated constituents detected in the reserve pit mud and groundwater. Other ions and metals were detected in the pit mud and groundwater but to a lesser degree compared to chloride and sodium. Constituents detected above the recommended federal drinking water standards are nitrates (>10 mg/l), Manganese

Table 3. Ions in the shallow aquifer attributed to brine seepage (dissolved ions reported as mg/l)*.

		LOCATION				
PARAMETER	R DATE	SITE 2/21	SITE 3	SITE 4	MW 7s	L3
Ca	07/22/87	145	357	226	119	9510
	10/22/87	111	113	172	~ 109	
	04/23/89	102	96.3	83.2	86.4	8350
	06/28/88	105	147	175	84.9	7310
No	07/22/87	31.2	90.7	50.6	53.6	1270
-	10/22/87	34	35.B	44.5	53.2	
	04/23/88	32.4	37.4	42.6	42	1260
	06/28/88	37.2	3B.7	43	43.3	1300
Na	07/22/87	63.5	1436	681	196	6470
	10/22/87	85.2	229	462	211	
•	04/23/88	91.8	105	143	162	88400
	06/28/88	89.1	324	476	146	73500
		••••			• • •	
ĸ	07/22/87	3.9	25.4	28.4	8.2	1040
••	10/22/87	4.5	8.1	18.7	8.8	••••
	04/23/88	3.3	4.5	0.6	6.3	1020
	06/28/88	5	14.4	19.2	7.3	885
	VV/ 10/ VV	•	• • • •			
C1	07/22/87	31.9	2650	1060	217	14500
	10/22/97	19.7	221	678	274	
	A/23/8R	21.1	24.4	28.2	70.7	156000
	14/28/8R	23.2	A11	680	40.6	130000
	VU/ 10/ 00	T41 T	7.4.4	000		
Ko	07/22/87	C.001	5.01	0.29	0.003	42
	10/22/87	6.001	0.4	0.085	<.001	-
	04/23/89	0.006	0.28	0.022	0.015	35.9
	04/28/88	0.001	0.077	0.001	0.002	30.9
	V0/20/00	*****	•••		••••	••••
มกถุง	07/22/87	319	532	589	476	220.8
11000	10/22/07	179	191	539	478	
	A4/23/89	874	470	501	506	332
	NE 120 100	700	444	508	448	268.4
	V0/20/00	JOL	111	370	476	
604	17 177 107	203	254	253	243	694
204	10/00/07	107	200	200	217	
	10/22/0/	207	200	227	267	1110
	01/23/00	210	233	210	207	223
	10112122	207	293	100	1.37	***
N	07/22/87	22.7	2.61	9.14	4.03	25
н	10/22/97	5,89	3.85	3.21	2.22	
	A4/27/00	4.34	1.55	0.97	1.24	<.04
	V4120100	11	10.6	3.28	2.75	35
	<u>40170100</u>		····		 41 •	
		XSAA	316A V#		11 V (

*See also Br Appendix L

Table 4. Metals in the shallow aquifer attributed to brine seepage (dissolved metals reported as ug/l).

		LOCATION						
PARAMETE	R DATE	NW 21s	MW 3	HW 3s	HW 4s	HW 16s	MW 20s	L3
8	07/21/87		1030					
	10/22/87		220				570	
	04/23/88		140					38000
	06/28/88	120		400	820	600		38000
Li	07/21/87		370					
	10/22/87		71				160	
	04/23/88		20					13100
	06/28/88	20		130	190	100		10000
Ba	07/21/87		530					
	10/22/B7						92	
	04/23/88		60					320
	06/28/89	1		110	100			270
Sr	07/21/87	t	7300					
	10/22/87		1020				2510	
	04/23/89	1	990					241000
	05/29/88	630		2660	2780	940		207000
Ti	07/21/97	,	29					
	10/22/87	1	3				5	
	04/23/98	l	4	•				(1
	06/28/88	4		6	8	5		{}
Zn	07/21/87	ŗ	14					
	10/22/87		4				<3	
	04/23/88	1	<3					880
	06/28/88	3 (3		<4	3	(3		430
Be	07/21/87	r	21					
	10/22/87	1						
	04/23/88	3	(1					26
	06/28/88	3		1				21

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(>0.05 mg/l) and chloride (>250 mg/l). Nitrates and Manganese only exceed these standards in the spring and summer at the pit area; chloride is generally above the recommended drinking water standards year round.

Ions and metals at elevated concentrations in the pit mud but not detected in the groundwater are probably absorbed, bonded, precipitated and/or absorbed by the clay minerals of the pit material. An x-ray diffraction pattern of the pit mud (Appendix M) shows the drilling mud is rich in smectite (or bentonite, which is the common drilling mud used drill oil wells). Smectite is a 2:1 layer silicate clay mineral that will absorb hydrated ions in the interlayer spaces of the 2:1 structure (depending on the size and charge of the element or compound).

The ions that do not seep into the groundwater are identified by comparing the pit mud soil moisture analyses with the groundwater analyses. They are: nickel, vanadium, zirconium, aluminum, silver, cadmium, iron, chromium, copper, molybdenum, and lead (Appendix K and L). These are retained in the pit mud.

Figures 11 through 14 are chloride plume maps. Most contaminant loading occurs in the spring and summer as shown by high chloride concentrations originating from the pit area in Figures 11 and 12. Loading wanes in the fall and chloride concentrations drop (Figure 13). In the winter, the water table continues to lower and little surface recharge is

CI CONCENTRATION mg/I FOR JULY, 1987





CI CONCENTRATION mg/I FOR OCT., 1987





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CI CONCENTRATION mg/I: APRIL, 1988





CI CONCENTRATION mg/I: JUNE, 1988





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available. Background chloride concentrations are present in the groundwater under the reserve pit at this time (Figure 14). In the spring, chloride concentrations increase at the pit area from the high water table intersecting the bottom of the pit and increased precipitation leaching pit contaminants into the groundwater.

Shallow groundwater discharges to the irrigation ditch in the spring, summer and fall since the ditch acts as a groundwater sink, removing chloride from the aquifer. In late winter, the ditch is influent and a small groundwater divide probably forms under the ditch causing chloride migration to be parallel to the ditch.

Figures 11 through 14 and Tables 3 and 4 show that contaminants load into the shallow aquifer in a annual cyclic fashion. In addition, Figures 11 through 14 show that the chloride concentrations decline dramatically over a short distance from the source (350 feet or 115 m), from 2800 mg/l under the pit to 600 mg/l at MW 16s.

Data plots for Figures 11 through 14 are in Appendix N, which also contains SC plume maps and data for the same locations and seasons.

Michelle Dewey of the University of Montana studied the same reserve pit site from 1982 to 1983. Dewey completed five monitoring wells in the shallow aquifer and collected water quality samples for major ion analysis. Chloride concentrations down gradient of the reserve pit at Dewey's wells 3 and 5 (Appendix O) have similar chloride concentrations as down gradient monitoring wells sampled in this study. Background chloride concentrations at Dewey's well 1 (Appendix O) also has similar chloride concentrations as background wells sampled in this study (with the exception of one anomalous peak in April, 1983).

A comparison of Dewey's data and data collected in this study indicate that brine seepage has not decreased significantly in the past four years, suggesting the pit will be a long term contaminant source. However, comparison of Dewey's data with this study is awkward since her wells were located at different points and completed differently. The comparison does show the pit mud on site may load brine wastes into the shallow aquifer for the next tens of years.

Yellowstone River chemistry

Data provide by the United States Geological Survey shows no significant increase in chloride, sodium, sulfate, calcium, potassium, and magnesium concentrations occur in the chemistry of the Yellowstone River between Terry, Montana and Sidney, Montana (Slagle, 1983, 1984 and 1988) (Appendix O). Terry is approximately 90 miles (155 km) up stream of Sidney on the Yellowstone River. Many reserve pits are located along this stretch of the Yellowstone River in both valley fluvial deposits and the nearby Tertiary deposits. Brine from these reserve pits may seep into the underlying aquifers and in turn increase chloride and other ion concentrations in the

Yellowstone River (assuming groundwater velocity, time, dispersion and discharge will move the contaminants to the river). Chemical data collected from the Yellowstone River indicate dilution of brine contaminants in the groundwater or in the Yellowstone River are too great to measure an increase in concentration of ions analyzed or no increase occurs.

Apparent Conductivity and resistivity

EM

Results of the EM 10 m horizontal and vertical mode surveys are in Figures 15 and 16. The EM 20 m and 40 m surveys show similar conductivities to those measured in the EM 10 m survey (Appendix P). All EM data were collected in July, 1987. Figures 15 and 16 show soil conductivity data measured near the reserve pit and do not include the conductance data measured over the entire EM 10 m grid in Figure 5. EM 10 m conductance data collected away from the pit were background readings and are plotted in Appendix P.

In Figure 15, the pit area is delineated by conductance values ranging from 40 to 110 mmhos/m compared to background readings of 20 to 25 mmhos/m. Only a limited area around the pit shows a high conductivity. Figure 16 accurately defines the pit area with conductivity values ranging from 90 to 160 mmhos/m, and appears to delineate a conductive zone (30 to 60 mmhos/m) in the shape of a plume downgradient of the pit, similar to the chloride plume map in Figure 14.

Water quality data for July, 1987 indicate that brine



Figure 15. Contour map and data plot of EM 10m horizontal mode survey (CI = 15 mmhos/m) (July 1987).

22 26 24 25 22 23 21

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Figure 16. Contour map and data plot of EM 10m vertical mode survey (CI = 20 mmhos/m) (July 1987).

contamination was present downgradient of reserve pit. The presence of brine in the groundwater was apparently significant enough that the EM 10 m vertical mode could delineate the contaminated area. The EM survey may also have delineated soil contamination. This is indicated from lysimeter data at L2 which shows contaminated soil moisture approximately four to six feet below ground surface, 100 feet (30 m) west of the pit area. Soil chemistry data at the MW 4 site indicates that contaminated pit mud is present from approximately six to 12 feet (4 m) below land surface, 100 feet (30 m) north of the pit area.

The contrast between the EM data in Figures 15 and 16 is probably caused by a difference in EM depth penetration. The horizontal mode (Figure 15) is less influenced by the surface and near surface conductivity, probing soil approximately 1.5 times the coil spacing. The vertical EM mode (Figure 16) examines conductivity from the ground surface down to about .75 times the coil spacing. The shallower EM penetration mode) did a job delineating brine (vertical better contamination because the water table and reserve pit are within seven to nine feet of the ground surface. The "deeper" soil was not as affected by the brine contamination making the vertical EM mode a better tool to evaluate brine contamination at this site.

A comparison of EM conductivity values with chloride and SC levels collected in July 1987 show the EM highs are located in the same area as chloride and SC highs (Appendix M). A correlation of EM versus chloride or SC could not be determined with these data. Monthly EM data, collected during water quality sampling, is needed to quantitatively determine groundwater quality from EM conductivity.

Results indicate that EM is a valuable tool for locating brine contamination. Highly conductive pit mud and brine water are delineated with the equipment.

Resistivity

Resistivity 10, 20 and 30 foot a-spacing data and plots support conductivity data collected with EM (Appendix N). Low resistivity values are present in the pit area extending northward in the same direction as groundwater flow away from the pit area. Operation of resistivity equipment was significantly more time consuming and cumbersome to use compared to the EM equipment.

Resistivity data collected in this study are very similar to Dewey's resistivity work in 1983, suggesting that contaminant concentrations have not significantly changed since the first resistivity surveys in 1983.

Groundwater and solute transport modeling

Flow modeling

Figure 17 displays the model area. Figure 18 displays the node location, boundary conditions, and site landmarks. K was set at 4200 gal/day/ft² for each node except those representing the irrigation ditch. The bottom of the ditch







is composed of mud and silt, and consequently is assumed to be less transmissive then the underlying sand and gravel. The actual K of the ditch bottom is unknown but probably is a number of orders of magnitude lower then the sand and gravel aquifer. Since PLASM is two dimensional model it interprets K to penetrate the entire aquifer thickness. To account for the lower K in the ditch, K was lowered in all ditch nodes to 420 gal/day/ft² which approximates the actual ditch K by averaging K in the ditch bottom and the aquifer. No other K zones were modeled. Specific yield was set at .2 for all nodes.

October, 1987 head data were used to calibrated a steadystate flow model simulation. Ten time steps were used, with a model error tolerance of .1 feet. Constant head data were varied until simulated head were within .1 to .2 feet of actual head data (see October data Appendix R and Appendix S).

Modeling groundwater flow alone, a wide range of K inputs could be used to simulate head within the .1 to .2 error tolerance using the same constant head data. To determine a more precise estimate of K, the aquifer test was simulated. In order to match the actual and simulated aquifer test drawdown, K had to be limited to approximately 3500 to 4500 gallons/day/ft² and Sy equal to .2 (Appendix T).

Simulated drawdown at the pumping well is off by approximately one foot. This is caused by the model removing water from a nodal area instead of discrete well diameter, which in turn caused the simulated drawdown at the pumping node to be less then actual well drawdown.

After achieving a steady-state simulation by adjusting constant head data and simulating the aquifer test, the same model and different constant head data were used to simulate 12 months of transient head data. Each month was modeled by reading a new constant head file (Appendix S) every thirty days and adjusting the constant head files until head was simulated within a .2 to .3 foot error tolerance.

An error tolerance of .2 to .3 feet was chosen after considering factors that might account for inaccuracies in the actual head data. These are: human error, survey elevations and accuracy of well location in the model.

Simulated monthly groundwater flow from July 1987 to June 1988 are in Appendix R. The accuracy of each monthly flow field is depicted in Figures 19 through 21 and Appendix T. Figure 19 is a regression plot of the actual and simulated head from both the transient (top diagram) and steady-state simulations (bottom diagram). Both regressions show a good correlation between the actual and simulated head.

Figure 20 compares the actual and simulated head in monitoring wells MW 3 and MW 7. This figure shows a good correlation between actual and simulated head for one year of transient modeling.

Figure 21 is a variance plot of simulated head data. This figure shows that 85% of the simulated head are within



Figure 19. Regression plot of actual and simulated head.



MW 3 and MW 7s.




.2 feet of actual head, and greater than 95% of simulated head data are within .3 of actual head data. In addition, the PLASM water balance accounted for 99.9% of all water moving through the model over the duration of the transient simulation.

The modeled hydrogeologic parameters and boundary conditions comprise a relatively simple model in an aquifer that is not homogeneous and isotropic. Modeling head data is not an important factor by itself since monthly groundwater flow is approximate with available water level data. However, modeling different hydrogeologic parameters and boundary conditions provide a better understanding of the hydrogeologic system and were needed to generate flow fields for solute transport modeling.

Solute transport

A preprocessor was used to convert simulated head data from PLASM into twelve groundwater flow velocity fields. They represented flow for each simulated month from July 1987 to June 1988. Chloride transport modeling was initiated in May 1988 since most of 1987 chloride is off site (via the return flow ditch), and a new chloride plume is present at the pit area. The drawback to this method is that chloride was simulated forward in time through July 1988 and then modeled with previous chloride concentration data (August 1987 through April 1988) in the model. This reasoning assumed that the previous year (1987) chloride data (and head distributions) were 1988 data.

Brine contamination is represented by loading particles into the model, where each particle represents 15 pounds of chloride. Most particles were input to the model during groundwater highs. Particle input was decreased during groundwater declines (Appendix U). In a few months, particle input deviated from the initial approach to increase or decrease chloride concentrations to improve the match of simulated and actual chloride concentrations.

Model input parameters consisted of: porosity (.25), longitudinal dispersivity (20 feet), transverse dispersivity (1 foot), and retardation (= 1; none). Particles were removed from the model if they migrated within a ditch node area since the ditch acts as a groundwater sink during most of the year. Figures 22 and 23 show simulated chloride and actual concentrations for four monitoring well sites. Appendix V contains simulated and actual chloride concentrations at each monitoring well site.

The comparison of actual chloride concentration in monitoring wells with the simulated chloride concentrations were done by comparing the highest chloride level for a specific date at each well site with the simulated value for that location. The chloride concentration at the well sites were assumed to represent the chemistry of the entire saturated thickness of the shallow aquifer (for this reason, only well site identification is presented in Figures 22 and

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Figure 22. Comparison of actual verses simulated chioride concentrations at well sites MW 7 and MW 3.



Figure 23. Comparison of actual verses simulated chloride concentrations at well sites MW 16 and MW 19.

23 and no deep or shallow well designations are used to identify well sites).

Figure 22 shows a good correlation between the simulated chloride concentrations at monitoring well sites MW 7 and MW These data indicate simulated chloride concentrations are 3. fairly close to the actual chloride concentrations. However, in Figure 23 the timing of actual versus simulated chloride concentrations at MW 16 and MW 19 is off by two to three months in both diagrams. In addition, the bottom diagram of Figure 23 shows that the simulated chloride concentrations are two to three times greater compared to the actual chloride concentrations. These data indicate the model did an inadequate job simulating chloride concentration through time at these sites. I explain the inaccurate simulated chloride concentrations in a number of ways:

- 1) I modeled my site in two dimensions when in reality it a three dimensional problem. Random-Walk assumes chloride concentration is evenly distributed throughout the entire aquifer thickness which was determined to not be true according to water quality data.
- 2) The flow field was calibrated within .2 to .3 feet error tolerance of actual groundwater elevations which may be insufficient to accurately move particles through the solute transport model.
- 3) Quantab chloride results were sometimes 10% to 50% different from the laboratory chloride concentrations.
- 4) The nodal scale in the model may be too large to accurately simulate head without creating a "blocky" velocity field to move particles in Random-walk.
- 5) Modeling 1988 chloride concentrations with 1987 groundwater flow data may poorly represent the

actual migration and chloride concentrations.

- 6) The flow model may be too simplistic in describing the geologic system. Actual aquifer geology is complected with variable K in three dimensions.
- 7) Solute modeling input parameters may be incorrect. Annual Chloride load

To calculate annual chloride loading, a cross section extending from MW 5 site, MW 4 site and MW 19s/MW 6 site was used to represent the shallow aquifer. The cross section was divided into blocks or zones vertically and horizontally throughout the saturated zone (Appendix W). Volumes of monthly chloride input from the reserve pit were calculated knowing the volume of groundwater moving through the cross sectional area and chloride concentrations at nested well sites. Monthly chloride data from Quantabs and laboratory determine chloride analyses used to chloride were concentrations. If no data were available, a monthly chloride trend to estimate chloride concentration was used concentrations for a well site.

Monthly chloride loading is in Table 5. Total annual chloride load from July 1987 to June 1988 is calculated to be 3,137 kg/yr from the pit, which is 3.7 times that of background chloride load (852.3 kg/yr chloride). Some error in the load calculations probably occur since about 35% of the chloride concentration data had to be estimated. Other factors which may affect the load accuracy are estimates of groundwater velocity and the area of contamination and

	<u>Pit Input</u>	Background	Simulated
July:	732.45	100.25	2503.9
August:	822.46	94.9	1564.9
September:	450.86	81.43	251.7
October:	242.06	80.56	149.7
November:	113.75	55.44	183.7
December:	70.56	51.9	45.0
January:	56.28	51.9	45.0
February:	40.71	45.28	44.5
March:	47.01	46.55	45.0
April:	44.66	41.58	45.0
May:	210.19	100.25	2245.3
June:	306.01	102.22	898.1
TOTAL:	3137	852.3	8021.9
lbs/yr	6915	1879	17658

Table 5. Annual chloride load in groundwater (kilograms).

chloride analyses/Quantab results.

The total simulated chloride load (Random-Walk simulation) appears to be high compared to the calculated chloride load. However, the model is only two dimensional and must assume the total saturated thickness of the simulated aquifer has the same chloride concentration. Chemical data show that only about half of the shallow aquifer is contaminated with chloride. This required the simulated load be twice as high as the field data in order to reproduce the measured values. The simulated annual chloride is therefore reasonable considering that the lower half of the aquifer is concentrations and generally background chloride at approximately 60% more chloride was input into the simulation compared to the calculated chloride load. The simulated chloride load is greater in the initial and ending months in comparison to the calculated monthly chloride load. This may be a results of improperly loading chloride in the model.

Elevated chloride concentrations in the ditch water were not detected by chemical analyses. The ditch water varied from approximately 20 to 35 mg/l chloride depending on the ditch flow rate. Chloride data from the up gradient ditch control site did not show any relative increase in chloride concentrations that could be attributed to brine seepage. In some instances the up gradient chloride concentrations were higher by a few mg/l than the down gradient control site where the brine enters the ditch.

Inability to measure the chloride input into the ditch is probably due to the relatively small amount of chloride loaded into the ditch from the reserve pit. This is demonstrated by calculating the chloride concentration which enters the ditch via groundwater (500 mg/l) and the groundwater flux into the ditch (two gpm or 7.57 l/minute). Using these values, 0.057 kg of chloride per minute enter the ditch which produces an increase in chloride concentration of This small .64 mg/l at an average flow rate of 3.5 cfs. amount of chloride loaded into the ditch is not detectable in chemical analyses since the analytical methods used to determine chloride concentration were not precise enough to measure chloride differences of .64 mg/l.

In the winter, when ditch flow is much less, the ditch is loses water to the shallow aquifer. For this reason chloride is not detected in the winter months.

The pit will continue to load chloride into the shallow aquifer system until the brine contaminants are purged from the pit material. The time necessary to remove available chloride is difficult to determine since controlling factors, including future climate condition, percolation rates and the amount of chloride available in the pit area, are poorly understood or unknown. However, for illustrative purposes an estimate will be attempted.

Parameters used in the attempt include assuming a pit volume (200 feet long by 70 feet wide by 12 feet deep or 61 m by 21 m by 4 m respectively) (Dewey, 1984) and a soil porosity (.25) since open void spaces will not contain chloride. Multiplying these values yields a hypothetical pit volume of 126,000 ft³ or 3,565,800 l of soil. Assuming all chloride in the soil volume loads into the shallow aquifer system, the required time to purge the chloride can be calculated knowing annual chloride load.

Soil moisture analyses at lysimeter L-3 (located in the pit material) indicate the potential amount of removable chloride in the pit material. The largest chloride concentration at L-3 (156,000 mg/l) was used to calculate the volume of removable chloride (where $3,565,800 \ l \ x \ 156,000 \ mg/l$ = 5.5626 x 10^{11} mg or 55,626 Kg of available chloride in the pit material). This value represents the volume of chloride that would take the longest to purge from the pit material.

Assuming the 1987-1988 calculated loading values are the average future chloride loading values, approximately 3,000 kg of chloride are loaded into the shallow aquifer per year. This value is divided into 55,626 kg total available chloride which equals 18.5 years or about 20 years until most of the chloride is removed from the pit from 1988.

The 20 year estimate is probably conservative since it represents a chloride loading pattern similar to the 1987-1988 year for the next 20 years. In reality more time will probably be required to purge all chloride from the pit materials. The "less easily removed" chloride (i.e., small concentrations held up by soil and chemical reactions) will require more leaching action and/or groundwater contact to be removed. This indicates chloride concentrations may be measurable in the shallow groundwater for more than 20 years.

The annual chloride load from the numerous reserve pit sites in the region is probably dependant on the oil-field brine encountered during drilling and the volume and type of drilling additives used during drilling. In addition, not all reserve pit sites may seep brine contaminant into the underlying aquifer. Without knowing the site specific information it is difficult to determine what regional chloride loading impacts will occur in the Yellowstone River Valley as a result of brine seepage from reserve pits.

Reclamation practices and remediation

This and other studies document that reserve pit

reclamation techniques are inadequate. Synthetic or clay liners used to protect the underlying groundwater are intended to inhibit brine seepage. However, when pit liners are ripped or punctured as a result of reclamation, a pathway for brine seepage into the groundwater is created. New cost effective reclamation methods are needed to stop brine from entering the groundwater.

A technique used in some areas is replacement of permanent on site reserve pits with mobil reserve pits. Mobil pits may be more cost effective depending on disposal costs (Reiten, 1988). Drilling fluids can be disposed of by injection wells and solid wastes can be buried at a drilling mud disposal site (properly designed to handle the wastes).

Pit solidification with fly ash has been used for several years in states other than Montana to stop brine seepage from reserve pits. The fly ash causes the mud to be fairly impermeable, limiting brine seepage into the underlying groundwater. This process may be effective, however, it needs to be tested at reserve pits located in the Williston Basin.

Other methods include pumping or evaporating the liquid fraction off the reserve pits and excavating the solid fraction. Brine water can be injected back into saline aquifers and the solid fraction shipped to a central disposal site for burial.

Remediation

Remediation at reserve pit sites with excessive brine

contamination would require one or more of the following:

- The reserve pit and trenches be cleared of mud to eliminate the source.
- 2) A cap over the pit and trenches be installed to limit wetting fronts and slow brine migration (would not affect mud below the water table).
- 3) installation of brine removal pump(s) in severe cases of groundwater contamination.

These are expensive measures and not the only solution to localized brine contamination. Most reserve pit sites are probably similar to the Iverson site and will require little or no remediation. However, highly contaminated sites which threaten vegetation and water supplies need to be cleaned up. Improved pit reclamation techniques could significantly decrease the need to curb brine seepage from reserve pits.

CONCLUSIONS AND RECOMMENDATIONS

Conclusions

- 1) Groundwater quality analyses indicated elevated concentrations of calcium, magnesium, sodium, potassium, chloride, manganese, bicarbonate, sulfate, nitrate, boron, lithium, barium, strontium, titanium, zinc. beryllium and possibly bromide are present in the shallow aquifer at the Iverson site. These contaminants apparently enter the shallow aquifer as a result of wetting fronts moving through the contaminated pit mud and the high water table intersecting the base of the reserve pit.
- 2) Analyses of pit mud show that high concentrations of all constituents detected in the groundwater are in the pit mud in addition to nickel, vanadium, zirconium, aluminum, silver, cadmium, iron, chromium, copper, molybdenum, and lead. These ions were at concentrations a number of orders of magnitude above background levels detected in soil water and groundwater analyses.
- 3) Nitrates and Mn concentrations in the groundwater exceed federal drinking water standards near the pit area in the spring and summer months. Chloride concentrations exceed recommended federal drinking water concentrations year around at the pit area and down gradient of the pit.

- 4) Rising groundwater at the pit area in the spring and summer months, accompanied by greater precipitation, generate high concentrations of brine contaminants in the groundwater. Lowering of the water table in the fall and winter months, accompanied with less precipitation, slows the contaminant loading processes producing lower concentrations at those times.
- 5) High horizontal hydraulic conductivity limits vertical contaminant migration to approximately the upper five to ten feet of the shallow aquifer. Dispersion and dilution appear to lower chloride concentrations by a factor of five at a point 350 feet (115 m) down gradient of the pit.
- 6) The reserve pit will be a groundwater contaminant source for at least 20 years and possibly more.
- 7) Local groundwater flow direction in the spring, summer and fall is partly controlled by the ditch. Groundwater flows is towards the ditch during these months. During winter, groundwater flow is approximately parallel to the ditch. Winter ditch levels are slightly above that of the groundwater, which produces a small groundwater divide below the ditch.

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- 8) No water supplies are threatened by brine contamination at this site since groundwater discharge to the ditch removes chloride and other pit contaminants from the shallow aquifer.
- 9) Cumulative groundwater contamination impacts are difficult to measure. However, broad scale regional brine loading impacts to the quality of the Yellowstone insignificant. River appear to be Local brine contamination from oil well sites could have serious local effects on water supplies and vegetation depending geology, groundwater flow direction the and on contaminant loading rate. The potential of these effects should be considered when siting and reclaiming reserve pits.
- 10) EM is a valuable tool to delineate brine contamination and plume migration at oil well sites. EM equipment can be used to survey a site quickly and easily to determine the extent of brine contamination. Resistivity delineates brine contamination as accurately as EM but it is cumbersome and time-consuming to use.
- 11) Groundwater flow and solute transport modeling are important tools to improve the understanding and characterization of the hydrogeologic system, aquifer

parameters, and contaminant migration.

- 12) Flow modeling accurately simulated flow direction, head and groundwater velocities.
- 13) The accuracy of simulated chloride concentrations were limited by the simplicity of the groundwater flow model, Random-Walk limitations of dispersion and dilution, boundary conditions and model inputs. Results show chloride concentrations were accurately simulated at some locations and inadequately simulated at other locations.
- 14) Current reclamation practices in Montana are inadequate to prevent brine contamination in shallow alluvial aquifers.

Recommendations

Individuals pursuing research involving reserve pit contamination should consider the following topics:

1) Determine whether organic contamination is a potential threat to groundwater quality at oil well sites drilled with oil based mud or mud containing diesel fuel additives. Volatile, semi-volatile and long chain organic compounds are potential contaminants not examined in current reserve pit literature.

- 2) Perform EM work needed to located contaminated oil well sites in eastern Montana, and correlate chloride concentrations in the groundwater with EM conductance values.
- 3) Determine cost effective measures that can be implemented to replace current reclamation techniques to stop or reduce brine contamination.

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APPENDIX A

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FUNCTION AND GENERAL PURPOSE OF DRILLING FLUIDS ADDITIVES

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Function and general purpose of drilling fluid additives (adapted from Murphy and Kewhew, 1985)

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Function	General Purpose	Common Additives	
Weighting Naterial	Control formation pressure, check caving, facilitate pulling dry pipe, and well completion operations	Barite, lead compounds, iron oxides	
Viscosifier	Viscosity huilders for fluids, for a high viscosity-solids relationship	Bentonite, attapulgite clays, all colloids, fibrous asbestos	
Thianer Dispersant	Modify relationship between the viscosity and percentage of solids, vary gel strength, deflocculant	Tannins (Quebracho), poly- phosphates, lignitic materials	
Filtrate Reducer	Cut the loss of the drilling fluid's liquid phase into the formation	Bentonite Clays, sodium car- boxymethyl cellulose (CNC), pregelatinized starch, various lignosulfonates	
Lost Circulation Haterial	Primary function is to plug the zone of lows	Valnut shells, shredded cello- phane flakes, thixotropic ce- ment, shredded cane fiber, pig hair, chicken feathers, etc.	
Alkalinity, pH Control	Control the degree of acidity or alkalinity of a fluid	Lime, caustic soda, bicarbon- ate of soda	
Emulsifier	Create a heterogeneous mixture of two liquids	lignosulfonates, mud deter- gent, petroleum sulfonate	
Surfactaat	. Used to the degree of emulsifi- cation, aggregation, dispersion, interfacial tension, foaming, and defoaming (surface active agent)	Include additives used under emulsifier foamers, detoamers, and flocculators	
Corrosion Inhibitor	Materials attempt to decrease the presence of such corrosive compounds as oxygen, carbon dioxide, and hydrogen sulfide	Copper carbonate, sodium chromate, chromate-zinc selu- tions, chrome lignosulfonates, organic acids and amine poly- mers, sodium arsenite	
Defoamer	Reduce foaming action especially in sult-water-based mods	Long chain alcohols, silicones, sulfonated	
Fosmer	Surfactants which foam in the presence of water and thus permit air or gas drilling in formations producing water	Organic sodium and sulfonates, alkyl benzene sulfonates	
Flocculants	Used commonly for increases in gel strength	Salt, hydrated lime, gypsub, sodium tetraphosphates	
Bactericides	Reduce hactoria count	Starch preservative, paraform- aldehyde, caustic soda, lime, sodium pentachloraphenate	
Lubricants	Reduce torque and increase horse- power at the bit by reducing the coefficient of friction	Graphite powler, soaps, Cer- tain oils	
Calcium Remover	Prevent and avercome the contame ination effects of analydrite and gypSum	Canstie soda (N.OH), sada ash, bicarbonate of soda, barium carbonate	
Shale Control Inhibitors	Used to control caving by swelling or hydroos disintegration	Gypsom, sodium silicate, cal- cium lignosolfonates, lime, salt	

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APPENDIX B

DISCUSSION ON THE REGIONAL HYDROGEOLOGY

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The site is located on the western flank of the Williston Basin (Exhibit 1) in Richland County, Montana which is a deep structural depression containing 14,000 to 15,000 feet (4500 m) of sedimentary rock overlying a gneissic precambrian basement (Carlson, 1985).

As of 1982 oil removed from the Williston basin in Montana had totaled 21,923,760 barrels from 1,360 wells (Dewey, 1984). Major oil producing zones 9,000 to 13,000 feet (2700 to 4000 m) below ground surface include the Paleozoic Mission Canyon Limestone and Red River formation (Prichard and Landis, 1975). The Spearfish Evaporite sequence is above the oil producing zones and marks the Permian boundary (Carlson, 1985).

The Regional surficial geology encompasses the Tertiary Fort Union Formation to the Quaternary alluvial and fluvial deposits. Exhibit 2 shows a geologic cross section of the area and Exhibit 3 describes the Mesozoic through the Quaternary stratigraphy. The Fort Union formation has four members, however, only the Tongue River member is significant to this study. The Tongue River Member is a light grayish to brown layered sand, silt, and clay deposit with .25 to five feet thick coal beds and clinker zones. Gypsum layers and iron nodules are found in many of the silt and clay Most beds are massive with local crossbedding, and layers. weathered to a yellowish buff color. Cementation is weak in most rock layers making the Tongue River Member soft. The depositional environment was probably similar to a terrestrial flood plain, marsh, or swamp setting (Prichard and Landis, 1975). The Tongue River Member generally yields approximately 10 gpm of water from

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Exhibit 1. Location of Richland County Montana, Montana in the Williston basin shown by hatched lines (adapted from Dewey, 1984).

Cast Tel



(diagrammatic)

Exhibit 2.

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Exhibit 3.

Stratigraphic column for Richland County, Montana (from Slagle, 1981)

wells, most of which is derived from coal beds. Many wells are completed in the Tongue River Member since the water tends to be softer then other near surface water supplies and wells are generally not deeper then 75 feet.

The Yellowstone River deposited terrace gravel upon the Fort Union formation. These gravel were previously named the "Flaxville Gravel" which has now been abandoned. The oldest terrace deposits as reported by Howard (1960) are Miocene or Pliocene and may have a surface younger than Pliocene. The gravel cap the Fort Union highs and are composed of coarse gravel and cobbles in a coarse to fine sand matrix, and in some instances, the deposits are weakly cemented with calcite.

The Yellowstone River cut into the oldest terrace deposits and formed the Missouri Plateau peneplain, on which a younger terrace gravel were deposited (previously named the Cartwright Gravel). These gravel deposits are early Pleistocene age laid down by the ancestral Missouri and Yellowstone Rivers (Howard, 1960) and found on topographically high Fort Union banks as erosional remnants of a once extensive gravel blanket. These terrace deposits ("Cartwright Gravel") are lithologically similar to the older terrace deposits ("Flaxville Gravel"), and may in fact be reworked (Alden, 1932). Both of these deposits terrace deposits ("Cartwright and Flaxville") generally do not yield more than 10 gpm and the quality of the water is often poor due to a high TDS and hardness.

Another period of erosion and aggradation period deposited a younger terrace gravel sequence (previously named the Crane Creek

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Gravel). These gravel exhibit a similar lithology and texture as the two older terrace deposits and hence, thought to be in part derived from these deposits (Prichard and Landis, 1975). The major difference between the two older terrace gravel and the younger terrace gravel ("Crane Creek Gravel") is that the younger it is constrained to the present Yellowstone drainage, and is not preserved on plateaus and broad interstream uplands of the ancestral Missouri and Yellowstone Rivers. Howard (1960) correlates these gravel as pre-Wisconsin Pleistocene or Yarmonth Interglaciation. These younger terrace deposits produce up to 1,000 gpm in the major river valleys and are recharged by irrigation and channel seepage in addition to precipitation.

These younger terrace deposits ("Crane Creek Gravel") are infrequently covered by Glacial till, but do show evidence of continental glaciation after deposition. Glacial drift is early Wisconsin (?) according to Howard (1960) and Iowan or Illinoian (?) according to Alden (1932). Glacial deposits include ground moraines, stratified drift, till, melt-water and diversion-channel deposits, eskers, and kames (Prichard and Landis, 1975). The exact number of ice advances and glacial interims is disputed and not discussed in this text (see Alden, 1932; Carlson, 1985; Howard, 1960; and Prichard and Landis, 1975). Glacial deposits are reported to produce 20 gpm of water from wells but are seldom utilized.

Alluvium in the Yellowstone River Valley is mapped as two older flood plain terraces and alluvium deposited within the river channel. Alluvial deposits are predominately Holocene age

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deposited after an erosional period, which topographically has changed little since deposition of the older alluvium (Howard, 1960). Exhibit 3 is a stratigraphic column of units discussed above which also reviews water-yielding characteristics of each geologic unit. Alluvial deposits generally yield 10 gpm.

APPENDIX C

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WELL CONSTRUCTION

WELL

LYSMETER



WELL CONSTRUCTION INFORMATION

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			SCREEN	TOTAL	DEPTH OF
WELL #	ELEVATION	CASING	SLOT/LENGTH	LENGTH	PENETRATION
HW L	1922.1'	2"PVC	#20/3.0'	21.75'	20.0'
MW 2	1920.14'	2"PVC	120/5.0'	20.0'	19.0'
MN 3	1920.14'	2"PVC	\$20/5.0'	20.0	18.17',
NW Js	1920.89'	2 PVC	\$30/2.25'	14.0	11.4"
MW 4s	1921.09'	2*?VC	#30/2.0'	14.0'	12.33'
MW 4a	1919.74'	2"PVC	\$30/2.0*	16.0'	14.17'
MN 4d	1922.24'	2"PVC	#30/2.0'	22.0'	19.75'
MW 5	1919.59'	2"PVC	130/5.0'	20.0'	18.08'
MW 5s	1919.34'	2"PVC	\$20/3.0°	13.5'	11.5'
MN 6	1921.14'	2"PVC	\$30/5.0'	22.0'	19.0*
MW 75	1919.84'	2"PVC	#30/1.5*	15.3'	13.8'
MW 70	1919.69'	2"PVC	130/2.0'	24.0'	22.17'
MH 8	1915.72'	2"PVC	#20/3.0*	22.0'	20.5*(?)
MW 9	1920.48'	2"PVC	#30/2.0*	22.2'	19.62'
NW 10	1916.67'	2"PVC	#30/3.07	23.0'	20.17'
MW 11	1919.74'	2"PVC	\$30/3.0'	23.6'	21.75'
MW 12	1917.77'	2"PVC	#20/3.0'	22.0'	20.0'
MW 13	1919.09*	2"PVC	\$20/3.0'	22.0'	19.5'
MW 14	1921.547	2*PVC	\$20/5.07	23.0'	21.33'
MW 15p	1919.84'	4.5"PVC	#20/14.0*	21.1'	19.6'
MW 165	1919.87'	2*PVC	\$20/6.0°	14.0'	12.0'
MW 17s	1918.18'	2*PVC	\$20/3.0°	15.0'	12.6'
NW 18s	1917.75'	2"PVC	120/3.0'	13.07	11.0'
MW 195	1920.74'	2"PVC	\$20/6.0'	17.0'	14.9'
NW 20s	1917.96'	2*PVC	#20/3.5'	13.5'	12.9'
NW 21s	1920.64'	2"PVC	#2073.07	15.0'	12.4

LYSINETER CONSTRUCTION INFORMATION

			TOTAL	DEPTH OF		
LYSIM.	Ħ	CASING	DEPTH	PENETRATION		
L2		1.5"PVC	5.17'	3.591		
L3		1.5*PVC	6.83'	6.41'		
L3s		1.5"PVC	4.25'	2.92'		
L350		(NONFUNCT)	(ONAL)			
L3SS		INONFUNCT	IONAL)			
19		1.5"PVC	5.42'	3.84'	(DESTROYED	10/87)

APPENDIX D

MONITOR WELL LOGS, SIEVE ANALYSES AND SITE GEOLOGY

LITHOLOGY KEY

SO	50 - Blackish tan/dark brown soil, fine to coarse sands with 5 to 15% silt and clay, and 5 to 20% gravel, aoist, moderate to poorly sorted, rounded and Y subangular.
so	SO - Slackish tan to green soll, fine grained sand with 25 to 35% silt and i clay, moist, mocerately well sorted, subrounded to subangular, a few gravel bits and clay lenses.
FL	* FL - Blackish red fill soil and debris, coarse to fine sand, gravel and silt= with 5 to 10% clay (?), v. poorly sorted, subrounded gravel and cobbles, 5% coarse clinker bits, trace wood and plastic debris.
FL	FL - Blackish brown and grayish fill soil, sedium to v. fine same with 10% coarse grains and 20 to 30% silts and clays, 5 to 15% gravel, poorly sorted, subrounded to subangular, 5% clinker bits.
SD	SD - Tan brown to green clean sand, medium to fine grained with 3 to 5% gravel, fairly well to well sorted, subrounded to subangular, 9tz rich with 5 to 10% dologite and clinker, 2% mica, and trace coal bits.
SD .	SD - Tan brown fairly clean sand, sedium grained with 10 to 202 pea gravel, fine and coarse sand, moderately sorted, subrounded to subangular sand and subrounded gravel, Otz rich, 52 clinker, dolomite, agates, and coal bits.
, , SD&GR , , , ,	SD46R - Brownish sand and gravel, fine to coarse sand with .5 to 3 inch gravel, 5Z cobbles, 2 to 10% silt and clay, v. poorly sorted, subangular sand and subrounded gravel, 9tz rich with siltstone, granite, gneiss, dolomite, sandstone, and acate clasts, compon aquifer material for site.
20000000000000000000000000000000000000	GR - Brownish clean gravel, 1 to 2 inch gravel with 10 to 201 scaller and larger, 5 to 10% cobbles, 2 to 5% fines, moderately to poorly scrtad, subrounded to subangular, Qtz rich with siltstone, granite. gneiss, dologite, sandatone, and acate clasts, channel cut deposit (?).
CL	CL - Dark green to gray pit aud, silty/clayey fine sand with a few clinker bits and pee gravel, 30% silt and 20% clay, acderately sorted, rounded to subrounded clasts, acist, cily saell (diesel?).
SL&CL	SLUCL - Gray to dark green clayey silt, trace gravel, moderately sorted. rounded to subrounded clasts, v. moist, oily smell near pit area-possibly a trench dug from the reserve pit(?), more tan colored away from the pit.
CL/SD	CL/SD - Brown interlayered clay and sand with 20% gravel. v. poorly sorted sands and fairly sorted clays, subrounded. Qtz rich, trace clinker bits.
SL/SD	SL/SD - Brown interlayered silt and sand with 15 to 20% gravel, poorly sorted sands and fairly sorted silts, subrounded, Qtz rich.
SL/GR	SL/GR - Gray to Brown clayey-silty sand with 5 to 102 pea gravel, SUX silt and clay, wedium to commense sand, moderate to poorly sorted, moist, till(?).
FU	FU - Fort Union bedrock, thick layers of gray silty clay and black coal beds, 20% fine sandy lenses and trace pea gravel in silty clay, fairly well sorted, rounded to subangular, v. moist to wet, v. dense "siltstome".
ETMAR Trilobites STRATCOL

Scale 1: 120



ETMAR TELLOUILUS STRATCOL

Scale 1: 120

Feet MONITOR WELL 45 MONITOR WELL 4m MONITOR WELL 4d 0 FL FL FL FL SL&CL SL&CL 10 SD&GR SD&GR SD&GR 20 FU

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ETHAR Trilobites STRATCOL

Scale 1: 120



ETHAR TRIIGBILDE STRATCOL

Scale -1: 120

Feet	MONITOR WELL 5	MONITOR WELL S	S MONITOR WELL 6
	FL FL SL/SD	FL FL SL/SD SD&GR	SO SD&GR

ETMAR Trilobites STRATCOL



ETMAR TELLOBILES STRATCOL

Scale 1: 120



ETMAR THLIGBLEGS STRATCOL

Scale 1: 120



ETMAR THIIGBLEDD STRATCOL

Scale 1: 120

F	feet	MONITOR	WELL 15p	MONITOR	WELL 16s	MONITOR	WELL 17s
	O		80		SD		SO SD&GR
-	10		SD&GR		SD&GR SD		SD&GR SD SD&GR SD
	20						

ETMAR Trilobites STRATCOL

120 Scale 1:

F	eet	MONITOR	WELL 18s	MONITOR	WELL 195	MONITOR	WELL 20s
Γ	0		SO		50		50
			SD&GR SD				SD
	10	<u>(1999-1997)</u>	SD&GR		SD&GR		SD&GR

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	SAND SI	AMPLE	(MW12)				
SIEVE A	NALYSIS OF	SOIL	FOREST				
			NU. SAND NO.				
AND	AGGREGATE		PROJECT IVERSON				
USDA FOR	EST SERVICE REG	ON 1	LOCATION RICHLAND, CO.				
SIEVE SIZE	WEIGHT RETAINED	*/+ RETAINED	*/* PASSING SPEC. GRADATION				
3" (75.0mm)	0	0	100				
2" (50,0mm)	<u> </u>	0	100				
1-1/2"(37.5mm)	0	0	100				
1" (25.0mm)	0	0	100				
3/4" (19.0 mm)	0	0	100				
1/2" (12.5 mm)	51.5g	4.43	95.57				
3/8" (9.5 mm)	12.0 g	1.03	94.54				
#4 (4.75mm)	44.0 g	3.78	G 90.76				
PAN	1055.0g	90.75					
TOTAL	1162.0 g	WEIGHT	OF				
Sieve Loss or Gain	-3 g	ORIGINA	L SAMPLE				
SIEVE SIZE	Veight Retained X G	*/• RETAINED	% PASSING SPEC. GRADATION				
*8 (2.36 mm)	13.6	3.15	87.6/				
* 10 (2,00mm)			1. A. A.				
#16 (1.18 mm)	13.8	3.20	84.41				
[#] 30 (.600mm)	20.1	4.66	79.75				
₹40 (425 mm)							
₹50 (.300mm)	39.9	9.25	70.5				
[#] 100 (.150 mm)	139.0	32.24	3.8.26				
*200 (.075mm)	109.4	25.37	12.89				
Total Minus *200	E 55.5	12.87	MINUS *40 MATERIAL IS:				
TOTAL	H 391.3		PLASTIC				
Sieve Loss or Gain	н-а .8		NON-PLASTIC				
TOTAL DRY WEIG	HT BEFORE WASHING A	390.5	REMARKS				
TOTAL DRY WEIG	HTAFTER WASHING, VING B	347.6					
MINUS #200 WA	SHED OUT (A-B) C	42.9					
MINUS #200 FRO	OM SIEVING PAN D	12.6					
TOTAL MINUS	*200 (C+D) E	55.5					
MINUS #4 CORR	RECTION FACTOR. $\frac{G}{H} = $. 23194					
FORM NO.	OPERATOR	DATE	CHECKED BY DATE				
51-7100 MTL-2	2/80 50	11/87					

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AND	AGGREGATE	-	PROJECT IVE	ESON
USDA FOR	EST SERVICE REGI		OCATION ZIC	HLAND, CO.
SIEVE SIZE	WEIGHT RETAINED	% RETAINED	% PASSING	SPEC. GRADATION
3" (75.0mm)	0	0	100	-
2" (50.0mm)	0	0	100	
1-1/2"(37.5mm)	157.0 g	Z.57	97.43	
1" (25.0mm)	104.0 g	1.70	95.73	
3/4" (19.0 mm)	Z64.0 g	4.31	91.42	
1/2" (12.5 mm)	444.5 g	7.26	84.16	
3/8" (9.5 mm)	Z87.5 g	4.70	79.46	
#4 (4.75mm)	621.09	10.15	G 69.31	
PAN	42.42.0 g	69.31	J	
TOTAL	6120.0 g	WEIGHT ()F	
Sieve Loss or Gain	-19	ORIGINAL	SAMPLE	<u>6/2/.0 g</u>
SIEVE SIZE	Veight Retained X G	*/. RETAINED	% PASSING	SPEC. GRADATION
[₩] 8 (2.36mm)	55.4 g	6.77	62.54	
*10 (2.00mm)	· · ·			·
[#] 16 (1.18 mm)	67.49	8.23	54.31	
¥30 (.600mm)	95.19	11.61	42.7	
₩40 (.425 mm)	0	·		
¥50 (.300mm)	<u>84.1 g</u>	10.27	32.43	-
*100 (.150 mm)	<u>144.9</u> g	17.70	14.73	
*200 (.075mm)	<u>59.9 g</u>	7.31	7.42	
Total Minus ⁷⁶ 200	E 60.7 g	7.4/	MINUS *4	0 MATERIAL IS:
TOTAL	H 567.59	4	PLASTIC	
Sieve Loss of Gain	[<u>H-A — . 9 g</u>		NON-PL	
TOTAL DRY WEIG	HT BEFORE WASHING A	568.4	RE	MARKS
TOTAL DRY WEIG	HTAFTER WASHING, B	513.8		
MINUS #200 WA	SHED OUT (A-B) C	54.6		
MINUS #200 FR	OM SIEVING PAN D	6.1	1	
TOTAL MINUS	₩200 (C+D) E	60.7		
MINUS 74 CORR	RECTION FACTOR. $\frac{G}{H} = $.12213		
FORM NO. 51-7100 MTI-2	OPERATOR	DATE	CHECKED BY	DATE

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	GRAVE	C (M	W 13) SAMPL
SIEVE A	NALYSIS OF	SOII F	OREST -	
		F	IELD NO. GRAVEL	
AND	AGGREGATE	F	ROJECT IVE	RSON
USDA FORI	EST SERVICE REGI	ON 1	OCATION RICH	VLAND, CO.
SIEVE SIZE	WEIGHT RETAINED	% RETAINED	*/. PASSING	SPEC. GRADATION
3" (75.0mm)	0	0	100	
2" (50.0mm)	0	0	100	
1-1/2"(37.5mm)	253.5 g	12.03	87.97	
1" (25.0mm)	305.5 g	14.49	73.48	
3/4" (19.0 mm)	361.5 g	17.15	56.33	
1/2" (12.5 mm)	464.0 g	22.01	34.32	
3/8" (9.5 mm)	254.0 g	12.05	22.27	
∜ 4 (4.75mm)	132.5 g	6.29	G 15.98	
PAN	337.5 g	16.01		
TOTAL	Z108.0g	WEIGHT C)F	
Sieve Loss or Gain	- 11.5 g	ORIGINAL	SAMPLE	Z119.5 g
^{**} 8 (2.36 mm) ^{**} 10 (2.00 mm)	<u>435 д</u>	2.05	13.93	
#16 (1.18mm)	41.7 a	1.97	11.96	
*30 (.600mm)	51.1 g	2.41	9.55	
#40 (425 mm)	d ²			
₹50 (.300mm)	39.5 a	1.86	7.68	
#100 (.150 mm)	52.6 d	Z.48	5.21	
*200 (.075 mm)	42.0 g	1.98	3.23	
Total Minus *200	E 68.5 g	3.23	MINUS #40	MATERIAL IS:
TOTAL	H 338.99		PLASTIC	
Sieve Loss or Gain	H-A - 1.4 g	J	NON-PLA	ISTIC 🔀
TOTAL DRY WEIG	HT BEFORE WASHING A	337.5	REI	MARKS
TOTAL DRY WEIG BEFORE SIE	HTAFTER WASHING, VING B	Z76.8		
MINUS *200 WA	SHED OUT (A-B) C	60.7		
MINUS #200 FR	OM SIEVING PAN D	7.8		
TOTAL MINUS	#200 (C+D) E	68.5		
MINUS #4 CORF	RECTION FACTOR. $\frac{G}{H} = \frac{1}{2}$	047153		
FORM NO. 51-7100 MTL-2	2/80 SP	DATE 11/87	CHECKED BY	DATE

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CROSS SECTION B-B'

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APPENDIX E

QUANTAB DISCUSSION

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QUANTAB[®] CHLORIDE TITRATORS



GENERAL

INTRODUCTION DUANTAB Chiloride Trirators are convenient measuring devices for sait (chiloride) in aqueous solutions or dilute abueous extractions of solids in the approximate ranges from 01% (60 com CIT) to soluration of NACI QUANTAB Trirators are selfacting and provide recorded objective results which are days to read undimercipit into meaningful answers where monitoring sait concentration is important.

PSCOUCT	APPHOXIMATE
DENTIFICATION	TATION RANGE
OUANTAB Chickide	51 m to 0.08% NaCi
Titrator No. 1175	60 to 460 ppm, CI =
QUANTAB Chioride	053 to 0.5% NaCI-
Titrator No. 1176	300 to 3 600 ppm CI-
	Owith delation/extraction
	procedure up to
	6 91 NaCl or
	36 000 ppm CI-)
QUANTAB Chlorice	0 6% 10 6 4% NaCI-
Titrator No. 1177	3 600 to 38,400 pcm CI-
	:Up to saturation
	with dilation 1

APPLICATIONS The QUANTAB method for measuring chiprides employs a direct interaction of chloride with super dichromate. Dichromate is a frequently employed indicator for hitration of chloride with silver sats.

QUANTAB Chloride Tittators may be used whenever sait: chloride) measurement for control is important.

Listed below are three industries in which OUANTAB is being used successfully:

1. Chemical 2. Food 3 Petroleum

Additional use or application information is available

UDON request PREPARATION OF TEST SOLUTIONS

Good results with QUANTAB Chloride Titrators may be achieved by following these instructions carefully

For Aqueous Samples

Measure directly by using the appropriate OLANTAB Chipride Trator and carofully following the sectors headed DIRECTIONS. Aduebus samples containing more sait that the OLANTAB upper limit carb be measuured with QUANTAB Chipride Titrators by cilluling the sample to be tested before performing the measurement To obtain the correct value of salt content before diffuon multiply result on OUANTAB calibration table by the dilution factor.

For Solid or Semi-Solid Samples. Use Dilution, Estraction Procedure Below

- Mix or grind a representative portion of solid or semi-solid product thus dividing the product into small particles to insure extraction of sall.
- 2 Weigh 10 grams of finely-divided product and place in a suitable container
- 1137AD

3 Add 90 milliboiiing water Stir mixture vigorously for 30 seconds, then wait one minute and stir another 30 seconds to obtain a good extraction of sait from the sample. (Dilution factor is 10.)

4 Foid filter paper circle in half times, open into cone-shaped cup and place cup into extraction solution to collect a few drops of filtrate solution inside cup before performing test.

DIRECTIONS

- Place lower end of QUANTAB in the solution to be tested. (Immersion of entire QUANTAB will pretrigger completion signal.)
- Allow test solution to saturate column. This is accomplished two minutes after the yellow test completion signal across the top of the column begins to turn dark blue. This usually requires about 20 minutes.
- Results may be read from 30 seconds to 5 minutes after complete signal color change occurs (Strip excess fluid out open end to make test result permanent).

INTERPRETATION OF REACTION

- Record QUANTAB reasing to the rearest on-shalf pivision on the numbered scale at the <u>to</u> of the white color change.
- Convert QUANTAB reading to percent sait or com chloride (mg. chloride per litter) using the calibration table on the revorse side.
- If sample has been diduted, multiply result on calibration table by didution factor to obtain salt content of sample.

TECHNICAL

CriEMISTRY: QUANTAB Chloride Titrators consist of a thin, chemically inert plastic strip Laminated within the strip is a capiliary column impregnated with silver dictoromate. When QUANTAB Chloride Titrators are placed in abueous solutions fluid will reain the column by capillary action and continue to progress as long as chloride solution enters the column. The reaction of silver dichromate with chloride (salt, produces a white color change in the capillary column. When the capillary column is completely saturated a molaturesensitive signal across the top of the column turns cark loue and the chloride concentration has then been massured as silver chloride. The reaction that taxes place in this measurement can de explained by the chemical equation.

The length of the white color change in the sabiliary column is proportional to phonde concentration.

SPECIFICITY: For the application described QUANTAB Chichde Titrators are specific for chlorides Bromides, lodides, Sulfides, strong pucks and strong bases can react with QUANTAB Chicr ba Titrators, however, they are not present in most samples to be tested in sufficient amounts to affect test results. Nitrite and nitrate have no effect on the fest

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PERFORMANCE TIPS

Gond Results may be expected using the OUANTAB according to the directions. Occasionally unusual results may occur The following ideas will help resolve problems which may acise

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Low Results	Caused by incomprete extraction of salt, prugged wick, variable product. Resplyed by more comprete mixing, filtering, pin hole in signal fyellow area, and recarioration.
High Results.	Caused by variable product or over easy extraction. Resolved by recalibration, or using 10% more water for extraction.
Slow Reaction or No Completion.	Caused by sample plugging, sealed off signal, or compressed wick. Resolved by filtering and or larger dilution, pin hole in signal, not water
Inaccurate High Readings:	Readings above 7 on the QUANTAB scale will cover a broad, sometimes unacceptable spread. Always read results below 7.0 by choosing appropriate QUANTAB or diluting.
Variable Results	Non-representative samples or variable product. Be sure sample and mixture are correct. Recoupling batch.

If technical questions arise, white or call Venture Systems. Ames Division, Miles Laboratories. Inc., 1127 Myrtle Street, P.O. Box 70 Elkhart, Indiana 45515. U.S.A. Telephone, 2191 264-6534

CALIBRATION TABLE

Be sure control number for table corresponds to bottle label control number This table is made from representative samples of the batch. Small variations from table may be noted within any batch. For greater accuracy recalibration by user is suggested.

	QUAR	TAB CHLO		CONTROL NO. 0192119					
QUANTAS Reading	NaCi in solu % NaCi i pi m	ition as cm CIT g CIT/L	OUANTAB Reading	NaCl in solu % NaCl p m	ntion as pm_CI™ g_CI™, L	QUANTAB Reading	NaCt in solut 3 NaCt - po mg	ion as m Cim p Cim /L	
1.0	υ.025	143	4.Ŭ	Ú.189	1130	7.0	0.61	3630	
1.2	0.032	190	4.2	0.21	1230	7.2	0.65	3910	
1.4	0.039	233	4.4	0.22	1340	7.4	0.70	4220	
1.6	0.046	279	4.6	0.24	1450				
1.8	0.055	32 9	4.8	0.26	1570				
2.0	0.064	384	5.0	0.28	1700				
2.2	0.074	441	5.2	0.31	1840		11/15/	79	
2.4	0.084	501	5.4	0.33	1980				
2.6	0.094	564	5.6	0.36	2140				
2.8	0.105	630	5.8	0.38	2310				
3.0	0.117	700	6.0	0.41	2490				
3.2	0.130	780	6.2	0.45	2680				
3.4	0.144	860	6.4	0.48	2900				
3.6	0.158	950	6.6	0.52	3130				
3.8	0.173	1040	6.8	0.56	3370				



Ames Division, Miles Laboratories, Inc. PO Box 70, Elkhart, Indiana 46515

C 1979 Miles Laboratories. Inc.

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APPENDIX F

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FALL AND WINTER GROUND WATER FLOW MAP



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APPENDIX G

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CLIMATE DATA

CLIMATE	DATA FOR	THE SIDNE	Y AREA C	MONTA	WA ST.	ATE UI	UIVERS	ITY EXT	ELINIE	1.17- SZA
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DATE	AV.MAX T		AV.MIN T		AVER6. T		INCH PPT	(ca)	INCH EVP	(ca)
Jun-87	84.6	29.22221	53	11.65665	68.9	20.49999	4.02	10.21075	7.71	19.58332
Jul-87	B5.9	29.94444	56.7	13.72222	71.3	21.83333	2.87	7.289770	E.14	15.59553
Auo-87	80.5	26.94444	51.4	10.77777	6 5	18.99999	0.97	2.453790	5.11	12.97934
Sep-87	77	24.99999	43.3	6.277777	60.2	15.65556	0.92	2.336790	3.59	9.118563
Oct-87	60.3	15,72222	31	-0.55555	45.7	7.511110	0.1	0.253998	2.63	5.680173
Nov-87	50.2	10.11111	24.2	-4.33333	37.2	2.888888	0.21	0.533397		
Dec-87	37.3	2.944444	15.1	-9.39898	25.2	-3.22222	0.05	0.126999		
ปรก-88	25.1	-3,83333	-1.1	-18,3898	12	-11.1111	0.55	1.422394		
Feb-88	32.9	0.499999	5.8	-14.5555	19.3	-7.05555	0.49	1.244595		
Mar-88	51.4	10.77777	22.8	-5.11111	37.1	2.833333	0.58	1.473194		
86- reA	67.5	19.77777	28.6	-1.88888	48.1	8.944443	0.17	0.431798	4.97	12.6238
May-88	82.1	27.93333	46.4	7.999999	64.2	17.88888	0.79	2.006591	8.24	20.9295
Jun-28	94.7	34.83332	61.4	16.33333	78.1	25.61110	1.27	3.225787	10.44	26.5176

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APPENDIX H

ANNUAL WATER TABLE FLUCTUATION



COMPARISON OF DITCH WL AND WELL WL FROM 7/11/87 TO 8/12/88



APPENDIX I

AQUIFER TEST DRAWDOWN AND RECOVERY





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Full Logarithmic, 3 × 5

"Page 122 omitted in numbering."













APPENDIX J FIELD DATA

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12	((36	(35							34		(50	
13	53	53		111			54	54			(50	
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ASHCRAFT	(32	(36										
SCHNIGT												
NEULER	93	79										
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DUANTAB RESULTS OF OP# CI IN WATER AT THE IVERSON SITE. RICHLAND, CD.

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1	1600	1597	1597	1553						1085		743		
2	925	1145	1061	1137	1005							999	82	3
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17				1457	1479	1345		1367	1293	1302	745	907	81	1
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SCHM101			723	536	487									
NEULER	2	1525	1534											
SHEELY		814	952	1044	990					931		670		
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DITCH E			677	1521	1413	1681	1142		1012	B 51		750	1	

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CONDUCTANCE OF WATER IN unhosics AT THE IVERSON SITE, RICHLAND, CO. (COPPECTED TO 25°C)

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TEMPERATURE OF WATER IN CENTIGRADE AT THE IVERSOM SITE. RICHLAND, CD.

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Well (DATE													
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4.	1 12.B	10.8	13.9	1	10.5	9.1					79	10.0	11 9	1.511
4	17.8	11.8	12.9		10.5	9	11.1	11.5	7.1	7			17.5	
	5 11.1	12.9	13.8		10	9.1	,,,,	11.5	, ,,,, ,, ,	•	91		11.1	
5	,				10.5	9.7						11.1	12 7	14.9
	11.7	12.4	12.5		11	9.1					7.1		13.2	
1	5				10.5	9		11	5	9.9	8.9	11.1	13.1	15.1
20	•)				10	9.1	10.1	10.4	5.9	6.9	7.9			
19	9				11	9.1	10.9		6.9	7.5	7.9	10.1	15.2	16.9
1	5				19.5	9.9							15.9	
7	K 11.7	12.9	12.0	1	11	9.9	9,9	8.1	6.9	7	9.1	10.9	17.9	17
7.	4	17.4	11.1		10.5	9.1		•••	•••	•	6.9	••••	11.9	
1	- 7				11.5	B.2	8.7		5	7	6.9	15.1	17.1	18.1
11	Â				12	9				-	6.5	13.1	15.9	
	R 14.4	13.9	12.9	•	11	10			6	7	7	11	18.9	
	2 22 2	13.1	13.9	,	11.5	10			•		, i		17.9	
54	0 13.9	13.9	13.0	1	11.8	9							15.1	
1	1 12.2	12.1	12.9	•	11.1	9.9					6.9			
	2 11 1	12.7	17.9	1	10.5	9.1					R		14.9	
		14.0	16.1		11.5	8.9			8.9	6.1	-		15.9	
1	4 127	16.5	15.0	ł	13	11			•••		. 6.9		14.9	
ACHERAE	7 J£+A T	11 9	12 5			•••					••••		••••	
CCUNTR.	י ז		15 0	l.	13	11.1								
MENIL EI	י ס	14.2	13.0		••	••••								
CUCCI	n V	14.0	14.5	k	17.4						6.8		15.1	
E DMAS	r N	15 0	10.V		14.0	••••								
EDULONA	n; M	13.0												
PUTTUR	ħ	74 0	10 1		11.1		,							
1. 1.	ν 7	17.7	20.0	1	11.1									
L. 17.		10.0	20.0											
L3	5 7	10.1	17 0			ε 3								
L.) ຄ	10./	11.0	,		4.2								
LI DITCU I	0 7	10.9			9 6			,		7			15.1	21.1
DITCH 1	\$	15./			d.J		1. J.I. 1. J.			, 1	17.5		27.1	
BIICH I			12.1		4		2.3	ا الب	•	د			4.7.1	

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Well #	DATE													
	07/11/87	07/21/87	08/17/87	09/25/97	10/20/87	11/22/87	12/19/97	01/23/88	02/21/89	03/20/88	04/23/89	05/24/88 04	/26/98	08/12/88
1			1.1		6.9						7.1			
			2.3			6.89								
21					1	6.7		7.1			7.5			
5)		5.8		1.2				6.4	6.9	7.1			
55	i				5.8	6.1								
45	i		4.5		6.8	-								
40			6.7		1.3	1					7.3			
4a			4.8		7.3	6.4	7.49	7.1	5.8	7.2	7.3			
5	i		6		7.1	7.2			7.2		7.3			
5s	i				1	6.8					1.2			
6			6.3		7.1	7.1					7.1			
16)				7.3	7		6.8	5.9	7.4	7.4			
20					7.3	7	7,41		6.7	6.6	7.1			
19)				7.1	6.9	6.6		6.B	6.9	7.1			
15	;				7.2	7.2								
75	5		4.9		7.2	7.2	6.9	7.1	6.8	6.7	7.1			
78			5.5		7.4	7.5					7.9			
17	1				7.1	7.1	6.9		6.9	7	7.2			
19	1				7.2	7.3					7.4			
8)		5.3		7.2	6.7			7.1	7.1	7.3			
9	}		7.5		7.15	6.9					7.3			
10)		7.3		7.2	5.8								
11			5.5		7.2	5.9					7.2			
12	2		5.5		7.2	7					7			
13	5		5.5		7.2	6.9			7.3	7.2				
14	ł		5.9		7.1	6.9					7.3			
ASHCRAFT	1		7.?											
SCHMIDT	[
HEULER	5		7											
SHEELY	1	8	6.8		7.5						6.8			
LDHAX	4	8												
Sherman	4													
CHITWOOD)		7.9											
- L:	2	6.9	6.4		6.7									
L3:	5	<6.0	6.6											
13	3	9.5	6.1							•				
LS	3	7.6	7.6											
DITCH 13	5		5.5		8.1	7.9	7.9	1		8.1				
DITCH 1	N/A		8		7.9	6.81	1.7	7.5	i	6.8	8			

H OF WATER AT THE IVERSON SITE, RICHLAND, CO.

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HELL	MP ELEVA.	NEW NP D	ATE			:					1								
LGCATIO	W IN FL.	ELEV. 0	7/11/67 6	17/21/187 (0 18/11/80	19/30/87	10/20/87	11/21/87	12/21/87 (01/24/88 0	2/20/88 (3/19/98	0 81/22/88 0	5/15/88 0	5/21/89 0	6/05/88 ()6/26/68 (91/15/88 (8/12/88
	1922.1		8.82	8.60	8.58	9.58	10.35	10.87	11.26	11.57	11.8	11.96	12.14		6.96		8.8	8.37	
	NL ELEV.		1913.29	1913.5	1913.52	1912.52	22.11.91	1911.23	1910.84	1910.53	1910.3	1910.14	1909.95		1913.14		1913.3	1913.73	
e 4	1920.14		9.57	10.11	9.76	11.10	11.94	12.41	12.66	12.85	12.97	13.15	13.19	10.49	9.85	9.56	9.73	9.31	9.3
	M. ELEV.		1916.57	1910.03	1910.38	1909.: 4	1908.2	1907.73	1907.48	1907.29	1907.17	1906.99	1906.95	1909.65	1910.29	1910.58	1910.41	1910.83	1910.34
21	1920.64						12.46	12.92	13.18	13.37		13.6	13.7	II	10.37	10.07	10.2	9.77	
	NL ELEV.						1908.18	1907.72	1907.46	1907.27		1907.04	1906.94	1909.64	1910.27	1910.57	1910.44	1910.87	
~	1920.14		9.65	10.20	9.95	11.15	12.01	12.48	12.73				13.22		9.96		9.85	9.38	9.87
	HL ELEV.		1910.49	196.94	1910.18	1908.99	1998.13	1997.66	1907.41				1906.92		1910.18		1910.29	1910.76	1910.27
35	1920.89						12.82	13.29	13.52	13.59	13.81	ORY	DRY	11.39	10.75	10.5	10.67	10.21	10.66
	HL ELEV.						1998.07	1907.61	1907.37	1907.2	1907.09			1909.5	1910.14	1910.39	1910.22	1910.6B	1910.23
Ŧ	1921.09		9.75	11.33	10.96	12.31	13.12	DRY	DRY	DRY	DRY	CRY	DRY	11.71	1.11	10.91	11.13	10.61	11.96
	HL ELEY.		1910.34	1909.76	1910.13	1908.78	1907.97							1909.38	1909.99	1910.18	1909.95	1910.48	1910.03
₽ŧ	1920.24			10.46	10.13	11.48	12.29	11.12	12.95	13.1	13.23	13.33	13.44	10.29	19.27	10.08	10.3	9.76	
	AL ELEV.		1910.33	1909.76	1910.11	1908.76	1907.95	1907.52	1907.29	1907.14	1907.01	1905.91	1906.8	1909.35	1909.99	1910.16	1909.94	1910.48	
4 B	1919.74		9.41	16.9	5.5	10.93	11.75	12.18	12.42	12.57	12.7	12.79	12.91	10.35	9.72	9°24	51.9	9.23	
	HL ELEV.		1910.33	1909.8	1910.15	1909.81	1907.99	1907.55	1907.32	1907.17	1907.04	1906.95	1906.83	1909.33	1910.02	1910.2	1909.99	1910.51	
ŝ	1919.59		8.99	9.92	9.59	10.90	11.71	11.13	12.35	12.5	12.64	12.72	12.83	10.29	5.7	9.6	9.82	9.22	
	M. ELEV.		1910.61	19.90.91	1919.01	1908.59	1907.89	1997.46	1907.24	1907.09	1906.95	1905.97	1906.76	1909.3	1909.89	1909.99	1909.77	1910.37	
55	1919.34						11.53	115	12.17	12.34	12.45	12.53	12.64	10.11	9.5	9.44	9.65	9.05	5.5
	N ELEY.						19.7091	1907.39	1907.17	1907	1905.29	1905.61	1906.7	1909.23	1909.84	1909.9	1909.69	1910.29	1909.91
ą	1921.14		10.81	11.59	11.24	12.58	13.38	13.8	14.03	14.17	14.31	14.39	14.5	12.06	11.44	11.22	11.47	10.93	
	N ELEV.		1910.33	1909.55	1909.9	1908.55	1907.76	1907.34	11.70%1	1905.97	1906.83	1906.75	1906.64	1909.08	1969.7	1909.92	1909.57	1910.21	
16	1919.87						12.17	12.57	12.79	13.04	13.05	13.14	13.25	10.8	10.22	10.15	10.36	9.71	10.21
	N. ELEV.						1907.7	1907.3	1907.08	1906.83	1906.83	1906.73	1906.62	1909.07	1909.65	1909.72	1909.51	1910.16	1909.66
20	1917.96						10.24	10.71	19.87	11.1	11.14	11.22	11.34 0	ESTROYED					
	NL ELEV.						1907.72	1907.25	1907.09	1906.86	1905.82	1906.74	1906.62						
6	1920.74						13.06	13.47	13.7	13.85	13.96	14.04	14.16	11.74	11.14	Ξ	11.19	10.42	11.05
	HL ELEV.						1907.69	1907.27	1907.04	1906.89	1906.78	1906.7	1906.58	6061	1909.6	1909.74	1909.55	1910.12	1909.69
15	1919.94						12.98	12.51	12.77	12.39	13.03	13.09	13.21	10.75	10.09	9.87	10.11	9.6	
	H EEV.						1907.76	1907.33	1907.12	1906.96	1906.81	1906.75	1906.63	1909.09	1909.75	1909.97	1909.73	1910.24	
75	1919.84		9.56	10.53	10.19	11.48	12.25	12.65	12.86	11	13.14	13.21	13.31	10.96	10.4	10.35	10.49	9.81	10.34
	M ELEV.		1910.19	1909.31	1909.65	1908.36	1907.58	1907.15	86.3091	1905.84	1906.7	1906.63	1906.53	1908.88	1909.44	1909.49	1909.35	1910.63	2.9091
P2	1919.69		9.77	10.30	9.98	11.28	12.67	12.46	12.69	12.61	12.94	13.01	19.12	10.79	10.2	10.1	10.3	51.9	
	NL ELEV.		1910.12	1909.39	17.9091	1908.41	1907.62	1907.23	1907.01	1906.88	1906.75	1906.68	1906.57	1908.91	1909.49	1909.59	1909.39	1909.96	
17	1918.18						10.86	11.22	11.44	11.56	11.68	11.75	11.85	9.58	9.14	9.21	9.26		
	NL ELEV.						1907.32	1906.95	1906.74	1906.62	1906.5	1996.43	1906.33	1908.5	1909.04	1909.11	1908.92		
18	1917.75						10.5	10.64	11.09	11.18	11.29	11.36	11.46	9.44	8.98	8.86	9.8	8.04	
	W ELEV.						1907.25	1906.91	1905.66	1906.57	1906.46	1906.39	1906.29	1908.31	1908.77	1909.97	1908.95	1909.71	
8	1915.78	(1915.72)	1.54	8.82	1.99		8.79	9.42	9.27	9.38	ò" tò	9.55	9.61	9.03	7.47 0	ESTROYED			
	W. ELEV.		1909.24	1907.95	1908.79		1906.93	1906.31	1996.45	1906.34	1906.23	1906.17	1906.11	1907.69	1908.25				

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HELL	MP ELEVA.	NEN NP	DATE					•											
LOCATION	m et.	ELEV.	07/11/97 1	18/12/10	18/11/80	09/30/87	10/20/87	11/21/97	12/21/87 (01/24/88 0	2/29/88 (3/19/88 (M/23/98 0	5/15/98 0	5/21/89 6	88/50/%	06/26/09	07/16/08 0	8/12/88
6	1920.48		12.07	12.63	12.46	13.39	14.04	14.3	14.44	14.53	14.64	14.68	14.73	13.51	13.25	12.8	12.8	12.45	
	HL ELEV.		1908.41	1907.85	1908.02	1907.09	1906.44	1906.19	1906.04	59.2091	1905.84	1905.8	1905.75	1906.97	1907.23	1907.68	1907.68	1908.03	
10	1916.57		9.22	9.92	08.9	10.89	11.65	11.85	11.94	11.67	12.05	12.1	12.07	11.06	10.89	10.84	10.63	10.11	
	AL ELEY.		1907.45	52°906i	1905.87	1905.78	1905.02	1994.82	1904.73	1905	1904.51	1904.57	1904.6	1905.61	1905.79	1905.83	1906.04	1906.56	
11	1919.74		9.42	10.36	9.98	11.29	12.06	12.49	12.7	12.82	12.95	13.04	13.16	10.96	10.34	10.16	10.16	9.45	
	HL ELEV.		1910.32	1909.38	1909.76	1908.45	1907.68	1907.25	1907.04	1906.92	1906.79	1906.7	1906.58	1908.88	1909.4	1909.58	1909.58	1910.29	
ដ	1917.77		7.12	8.09	7.74	9.04	9.85	10.28	10.51	10.69	10.8	10.89	Ξ	9.41	1.8.1	7.71	1.91	1.37	
	NL ELEV.		1910.65	1909.68	1910.03	1908.73	1907.92	1907.49	1907.26	1907.09	1906.97	1906.88	1906.77	1909.36	1909.95	1910.06	1909.86	1910.4	
5	1919, 00		7.59	10.35	10.69	11.28	12.01	12.37	12.56	12.65	12.79	12.86	12.95	Ξ	10.57	10.41	10.29	9.54	
	NL ELEV.		1911.5	1908.74	1909	1907.91	1907.68	1906.72	1906.53	1906.44	1906.3	1906.23	1906.14	1908.09	1908.52	1908.68	1908.8	1909.55	
1	1921.54		10.87	11.58	11.20	12.57	13.42	13.89	14.14	14.32		14.52	14.65	12.03	11.34	Ξ	11.2	10.75	
	HL ELEV.		1910.67	96.9091	1910.34	1908.57	1908.12	1907.66	1907.4	1907.22		1907.02	1906.89	1909.51	1910.2	1910.54	1910.34	1910.79	
ASHCRAFT	1919.37				11.5		12.97	13.28	11.43	13.55	13.64	13.46	13.75						
	ML ELEV.				1907.87		1906.4	906°.09	1905.94	1995.32	1905.73	1905.71	1905.52						
SCHINLDT	1917.45			5.53	9.25	10.12	10.61	10.95	1.11	11.12	11.24	11.29	11.41	10.42	9.89				
	ML ELEV.			1907.92	1908.2	1997.33	1906.32	1906.5	1906.35	1906.33	1906.21	1906.15	1906.04	1907.03	1907.57				
MEULER	1918.47				9.83	10.73	11.22	11.57	11.72	11.74	11.83	11.96	12.11		10.24	9.5		9.52	
	NE ELEV.				1908.54	1907.74	1907.25	1906.9	1906.75	1906.73	1906.64	1906.51	1506.36		1908.23	1908.97		1908.95	
YJI334KS	1917.99				3.15	01	10.45	10.81	10.98	10.9	11.61	11.14	11.41		11.21	9.19	9.5	8.54	
	ML ELEV.				1908.54	1907.99	1907.54	1907.18	1907.01	1907.09	1906.58	1996.85	1906.59		1906.79	1908.81	1908.49	1909.45	
SHEFFIRM	1918.49			19.9	9.58		DRY												
	HL ELEY.			1508.88	1908.91														
CHI TYDOD	1918.43			9.76	8.08		7.9	8.92	8.74	18.9	10.75	9.82	9.53	6.91	10.5	25.41	6.14	7.66	
	NL ELEV.			1908.67	1910.35		1910.53	19.901	1909.69	1908.62	1907.68	1908.61	1909.68	1911.52	567.93	1892.92	1912.29	1910.77	
BITCHIJ	1907.81						XK	1.11	1.15	FROZE	FROZE	1.24	0.89	0.65	0.53	0.67	0.68	XI	XX
	1908.36						1.51	XX	XX	FROZE	58026	1.79	X X	II	II	XX	XI	XX	XX
	1913.98						XX	XX	XX	FR01E	5902E	II	XX	XX	XX	XX	11	5.36	5.33
	WL ELEV.						1905.85	1906.7	1996.65			1996.57	1906.93	1908.16	1906.22	1909.14	1908.13	1908.62	908.65
390 i 35	1913.51				6.75	7.75	8.31	9.45	8.46	8.63	8.65	8.6	8.16	7.96	1.19	7.95	7.84	6.86	
	ML ELEV.				1906.75	1905.76	1905.2	1905.06	1905.05	1994.88	1904.86	1904.91	1905.35	1905.55	1905.72	1905.56	1905.67	1996.65	
011CH 1	1919.3	(1918.2)			5.56	Ĩ	XX	IX	XX	Ţ	XX	X	ΪK	XX	II	X	X	II	
	1915.63				11	3.55	3.92	3.98	3.99	1.2	4.15	3.85	3.62	3.5	2.95	3.65	3.72	3.12	
	IN ELEV.				1912.74	1912.08	1911.71	1911.45	1911.64	1911.39	1911.48	1911.78	1912.01	1912.13	1912.69	1912.18	19.1191	1912.51	

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APPENDIX K

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METAL ANALYSES (UNIVERSITY OF MONTANA ICP)

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DISSOLVED METALS

									(NE	GAT	IVE	IND	HCA1	231	LESS	: 7/	YAN	DEC	EL TION
	HETAL ANAL	YSES FOR	WATER SA	WFLES AT	THE IVERS	DH STIE. I	RICHLAND,	CO.			-								LIMIT).
		ppa	3 94	pen	PP	\$pe	pps.	ppe	2 04	p pa		99 8	P # b	ppa	. ppa	PP#	pps.	pp =	~
100.	8812	#1 • • • •	A6	54 134 804	64	Ca.	Fe	h.	lig	In	Na	H1	h	56	P	51	h	2m	
NW 2	08/20/08	4.414	4. 4 72	124.708	-0.001	0.001	9 ,015	3.409	29.799	0,001	\$4.500	9.00I	-0.424	0.961	0.035	9.J7v	0,001	0.007	
- P 11 3	16/22/87	0.000	6.663	104.300	0.001	8.000	4.414	4 504	15 504	A 401	211 144		-0.011			** ***		4 441	
F	04/23/08	0.029	9.906	97.500	-0.001	-8.0ut	8.010	3. 600	37.000	0.300	110.000	6.004	-8.617	9.071 9.071	A.673	10.710	-0.001	6.091	
1 60	06/20/88	0.020	4. 608	124.204	6.000	0.001	0.010	12.706	32.990	0.074	353.900	ð. #04	-0.007	6.689	0.042	12.360	6_0H3	8.891	
- 10UP.1	10/22/87	0.110	0.017	112.900	0.090	0.002	0.030	11.600	35, 194	0. 107	218.200	0.000	0.018	0.090	9.130	12.710	0.002	9.905	
•																			
MI 4 Est	19/22/07	9.044	0.098	143.500	000	0.011	0.020	18, 3%	45.900	4.979	154,700	0.007	-0.006	0.993	9, 488	12.410	0. vet	0.9%	
(8)	04/23/98	8.040	0.005	84.600	-9.001	0.000	0.670	4.600	42.000	0.023	140.900	0.004	-0, 041	0.612	0.114	9.510	9.040	0.001	
15)	667.58788	0.020	0.029	151.200	-0.44)	e.003	4.434	16.900	37.400	8.004	443.902	0.v01	-0.019	0.975	9.118	12.919	0.092	ú.002	
NV Sa	10/22/07	0.030	0, D21	145.000	0.000	-0.001	0.910	4.200	54.490	4.001	144.200	0.001	-0.081	0,104	0.141	13.060	0.603	0.992	
	04/23/08	0.020	9.019	116.300	E.0%V	6.000	0.010	5.300	59,109	0.011	137.299	0.005	0.005	9.115	0.218	10.599	-9.991	0.003	
	AP1581.00	V. V.30	V. 911	113.400	8.008	0.304	4.014	2.100	43.600	4.402	125.400	•.094	0.001	0.971	0.152	11.450	9,997	a' ĉaj	
Int 2.	16/22/82	0.846	0.072	197,290	-8.001	-4.401	6 616		51 500		204 506	Cak 0				10 714	104.0		
	-94/23/88	4.050	9.929	81.504	0.010	4.000	0.020	4. 700	40.500	6.002	154.700	9.04	0.007	10.000	6.108	8.990	-0.001	-9.001	
	96/28/08	0.130	0.010	91.200	0.0V1	6.001	0.060	4, 500	49.400	0.006	137.500	6.002	6.014	4.084	9.196	10.440	6.992	0.092	
- (SUP. 1	04/23/88	-9.419	9.014	81.000	-0.001	9.000	9.010	5.100	44. 300	9.001	158.900	0.002	-0.012	0.000	0.100	9.170	-9, 191	-9.961	
•																			
HW 0	04/23/8E	0.130	0.015	88.990	0.009	-0,001	0.050	1.500	44.700	4.665	122.709	0.006	6,074	4.196	9.182	10.020	- 9 _0ài	6.000	
	96/20/99	9.07 9	8.017	55.400	4.199	-0.991	8.949	9.509	41.200	4.177	143.000	0.003	0.072	9.976	0.184	9.700	0.003	0.007	
													_						
NH 145	30/22/07	0.029	9,013	179,800	9.011	-0.002	0.016	12.100	108.500	0.001	370.209	0.005	-0.004	0.172	9.269	14,399	9,007	0,005	
	#4/23/88 AL/19/00	0.0/0	0.912	94 565	9.907 8.000	W.913	4.020	6.409 to 160	53.70¥	4.476	261.299	0.005	9.940	9.134	9.177	12,180	0.097	9.9/5	
	487 287 00	* .e4 v	4.041	147.364	0.070	-41441	4.9ZV	14.749	12.900	4.444	317.100	A' 061	9. VV4	W. 143	4. 244	131854	A*445	A * AA1	
MM 87a	14/77/03	4 074		\$7.600	8.000	-8 965		6 100	49 300	0.061	144 200	6.061	-6 60)	6 161	# 163	10.376	5 061	0 00A	
	64/23/88	0.010	0.016	83.509	0.901	-8,901	0.070	5.000	45.269	6.007	177.400	4.045	0.001	0.084	4. 178	1.490	0.000	9.001	
	04/28/88	8,040	0.011	62.700	0.001	-0.002	0.020	4.100	35,500	0.002	87.800	0.001	0.008	9.087	9.071	19.470	9.092	0.002	
M 18s	19/22/87	0.040	0.017	92.300	9,001	-0.002	9.910	8.109	53.600	0.00t	183, 600	8.003	0.006	0.121	8.135	9.430	Ø. 1994	9.990	
	04/23/08	0.070	0.016	Bú. 299	0.002	-0.901	9.099	5.400	47.809	0.007	161.400	6.002	0.005	0.102	9.197	9,199	0_002	8, ú01	
	96/28/89	0.099	-0.901	61.809	0.001	-9.994	0.010	5.600	41.200	0.99L	135.200	-0.001	-0.000	4.079	0.093	9.270	0,002	0.001	;
76 17s	10/22/87	-0.030	0.019	48.200	-9.001	0.000	0.060	3.400	45,300	9,090	141.500	-4.992	-0.021	0.083	9,117	11.940	0.093	9.090	
	94/23/88	-9.530	0.019	42.894	-0.002	V. 001	0.000	1.300	68, 790 74 844	9.306	133.200	9.91/	-0.013	9.083	9.147	7.180	9,000	#.U91	
	V6/20/W0	-4.4/4	0.007	101.000	-0.002	-0.001		3.144	76.600	6.641	113-344	0.000	-0.410	4.961	V. 111	11.749	4.442	4.001	
NH 26s	10/22/87	0.000	0.008	160.700	-0.001	0.000	0.010	14.500	50.400	â. vôi	448.500	6.003	-0.004	6.603	0.163	12.376	0.001	-0.001	
	64/23/88	0.010	0.008	149.200	0.000	-0.001	9.010	4.000	35.400	0.001	239.700	0.002	-0.001	0.053	9.149	9.738	-0.003	0.000	
🚰 🖬 216	10/22/87	0.000	0.911	168. 500	e.000	-0.001	ê. 916	6, 400	23.299	0.001	\$1.500	0.003	-0.005	0.048	0.057	11.950	0.091	9.965	
[94/23/88	0.070	0.010	101.700	U. 00 I	0.001	÷.139	1.700	32.000	4.014	69.200	0.095	0.004	0.039	9.081	7.410	0.0#Z	0.004	
L	45/29/86	0.030	9,007	95,960	0,000	-0.002	Q.018	5, 500	\$3,300	0.001	80.000	0.005	-0.003	0.045	9.059	9.970	0.001	-0.00?	
% (\$0P.)	10/22/07	-9.040	6.098	107.400	9.000	-0.002	Q. (H)U	1.200	33.608	0.001	92.600	-0.001	-9.913	9.041	9.057	11.020	0,041	9.903	
										A 484					A 175	18 704			
E MJIIG	10/22/8/	-9.010	4 .410	B1.849	4.601	-4.001	0.014	1.194	\$Y.800	4.4 54	122.400	4.002	-0.00/	A.113	0.173	10.300	4. 052	01013	
-	14/19/41	-8 104	6 447				-0 -04	-1 164	6 766	-	0 BOA	-0.667	40.057	-8.474	6.469	-0.080	-0.962	-0,004	
- WILK IS	A4/31/88	-7.//V	9.997 8.669	55.764	0,664	8,661	6, 100	2.164	54.946	6_61E	102.104	-0,001	-4,017	6.041	0.441	4,080	0.004		
1	AL178/88	-0.030	b. 514	48.761	8.661	-1.005	-9.618	-0.100	30,700	0.034	74.800	-0.003	-0.025	0.039	9.097	6.778	0.003	-0.004	
La (BUP.)	10/22/0?	-0.050	9.010	63. 109	0.600	-0.003	0.000	5,790	67.100	0.039	154.500	0.003	-0.917	4.114	0.127	10.060	0,043	-0.004	
				•••••															
13	04/23/89	77.50	1.60	\$470.00	-1.05	0.00	368.50	4870.00	1300.00	37.35	37445.00	1.70	19,05	-5.20	15.55	47.00	-0.10	0.15	
	06/25/86	23.00	4.89	7305.00	0.55	-1.10	312.50	1180.00	1249.00	32.60	35085.00	°. 80	3,45	5.30	15.30	14.00	-0.19	-0.35	
12	10/22/87	0,974	-0.001	-0.100	0.004	9.003	0.150	B. 500	-0,200	0.075	-0.300	9.040	0,115	1.131	t. 746	25.420	-0.044	0.028	
	06/28/98	0.750	0.013	-9.100	0.001	0.255	0.15ú	1.000	-0.260	0.035	-0.300	0.033	0.129	1.077	9.800	27.000	-0.065	9.403	
												-		OTP	12	Flov	EQA	BCE)	
	TETAL ANAL	LYSES FOR	SEDIMENT	SAMPLES	AT THE IV	ERSON SIT	E, RICIULA	WD, ED.	GTUTAL DI	GESIABLE	NUMES D	STALMENL							
		1	1	્ર	1	<u>י</u>	1	p pa	\$ \$\$	e pe	PP4	ppe.	ppe L-	ppe t-	. PP B	ppe e	1×	978 7_	
LUC.	PEFTH Later 44	#1 1	1 110	F0	,	Fa	51 10 114	-34 AAA	1 0	19 18 214	1 44 1	197.304	ta Adm. Et	36.604		۳ ۲۵ فر	1293.9984	49,044	
PITHON T	- 079 FL. - 6-8 44	3.//J 1 AA3	9.3/7	7.191	3.76 7	1.09) 9 117	301 330 35, 141	-16 ul4	1.049	20.770	1.233	411.517	106.104	44.044	\$5.055	0.041	1423.323	5851	
100 3	5-1-1	4.171	3.847	2.107	2.191	6. 193 1. Lta	37.7++	- 10.910 - 30.471	1.494	14.944	6.942	434.461	84.461	21.912	21.912	0.649	1573,765	54.574	
2 Ni 4	4-12 (1.	4.318	4.599	2,907	2.077	2.2.4	29.690	-24.729	1.060	40.752	1.482	459.059	156.291	16.815	45.500	8.069	2288.724	91.998	
1 m 7	2-1 (1.	5,360	2.634	3.504	2.055	0.824	25.935	-36.3/4	1.111	15.285	0.750	612.157	31.762	-30.76	-14.998	9.059	1547,295	57.965	
- He 7 (6)/	1 2-7 41.	3.729	2.910	4.104	2.733	0.872	32.010	-17.864	1.290	17.469	1.058	740.844	32.754	-13.070	29,777	9.064	1777.568	67.593	
FT. UNLEUM	20 ft.	3.755	7.823	1.020	1.042	6.757	30.525	-21.991	0.8%	13.041	3.21/	435.341	173.271	-20,764	97,561	0.059	1443.396	44.391	
				-		•			<u> </u>		The	-	3.4.4.4		• • - •				
		(r	NEG.	ATIU	モル	VDIC	4TE?	CES	S TH	AAN	PE	1267	00	(K	^ IT }	•			

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APPENDIX L

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BASIC ION AND METAL ANALYSES (MBMG LAB)

WATER ANALYSES RESULTS FOR THE LVERSON SITE, RICHLAND. CC. (DI3SOCUED)

	ATTCH		64.				2				16				0.01				0.00				0.05		
	NTCH 13	N2.2	6.3 2	62.2	53.2	14.7	2	40.7	33.2	97.2	161	121	81.5	5.2	01	7.2	6.3	0.072	(.002	0.04	0.006	0.022	0.039	0.017	0.037
	1	9510		9350	7310	1270		1260	1300	8170		38600	73560	1040		1020	683	399		358	329	¥		35.9	30.9
	NN 70c	Į	167	109			20.2	35.3			137	282			51	8.8			<.002	0.099			(,001	0,006	
	MN 194		98.9	88	110		44.4	45.2	\$		135	E	116		6.7	5.4	6.7		(,002	0.28	. 002		(.001	0.065	0.001
	MW 18<		94.5	75.9	23		53.2	44.3	43,9		111	149	MEI		1.7	5.5	7.3		(.002	0.044	(.002		(.001	0.009	(.001
	MN 174		91.6	79.3	66.6		50.3	42.3	38.5		162	161	69.A		6.7	5.1	5.9		(,002	0.051	<.002		100')	0.016	0.001
	XX 16c		185	118	98.8		112	65.5	17.6		370	257	327		13.5	8.6	=		(.002	0.023	(,002		0.001	0.02	100")
	EL MU	71.5				45.6				HI				6.5				0.007				0.25			
	8	61		62.4	60.7	45.2		47.4	47.5	156		147	161	5.7		5.2	7.1	(,002		0.35	0.005	0.5		0.21	0.19
	61 7s	119	109	86.4	64.9	53.4	53.2	4	43.3	195	211	162	146	8.2	8.8	6.3	7.3	0.005	(.002	0.066	100")	0.003	(.001	0.015	0.002
	9 NK	94.9				45.8				112				ŝ				(, 002				1067)			
	N 55		148	115	119		55.6	99	46.7		145	136	130		b.4	6.1	5.6		<, (102	0.016	<002		(,001	0.015	0.001
	M4 44	80.7				54.9				192				6.9				(,0 02				0.053			
 7	4 24 24			83.2				42.5				143				0.6				0.014				0.022	
	NN 45	226	172		5/1	50.6	44.5		43	661	462		476	28.4	18.7		19.2	0.004	(.002		0.002	0.29	0.085		0.001
•	AN 35				147				38.7				H 22				14.4				(,002				0.077
	194 3 194 3	357	113	96.3		90.7	35.8	37.4		1136	229	19 5		25.4	e. 1	5. 1		<. 002	(.002	0.11		5.01	0.4	0.23	
	M 215		Ш	102	501		34	1. 1.	37.2		85.2	91.8	89.1		4.5	3.3	ŝ		(.002	0.012	< . 002		(.001	ý. Čý,	100.0
	14	145				31.2				63.5				3.8				(.602				(, 001			
ATTON	9A/DC	5				.,				! ")				Э				0.003				(*001			
LDC	PARAMETER DATE	PPA Ca 07/22/87	10/22/97	04/23/89	06/28/88	ppa Ng 07/22/87	10/22/01	04/23/88	06/28/58	122:22:42 bos Na 07/22:427	18/22/01	04/23/88	05/28/89	pam K 07/22/87	19/22/81	04/23/83	06/23/36	pon Fe 07/22/87	10/22/87	04/23/88	06/29/88	ppm Mn 07/22/87	10/22/87	04/23/98	06/28/89

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							191	SSOCU	(0)												
PARAMETER DATE per F 07/22/97	0a/0C (.1	아이 쫖 수	MA 215	MM 3 0.2	2) X2	NN 45 0.3	4 Ri	6.6 0	1 12 12	8 HH 6 0.6	K¥ 75 0.6	浅 8 0.7	MH 13 0.7	MH 165	571 MM	NU 18s	MN 195	HN 205	10	11CH 13 0.4	DITCH 1
10/22/67			0.1	0.6		0.6			0.4		0.7	-		9.4	0.6	0.6	0.6	0.6	;	0.6	0.6
04/23/68			9.6	0.7			0.8		0.5		0.3			0.5	0.4	0.8	0.8	0.4	¢4	0.3	
06/28/88			9.6		0.5	0.6			0.5		0.8	1.1		0.5	0.7	0.7	0.7		8	0.2	
ppe Br 07/22/87	5	5		6.1		2.5		0.4		5	0.5	5	0.2						110	3	
10/22/87			5	5		0.6			0.1		0.4			0.4	0.2	0.1	0.1	0.7		5	3
04/23/89			:.)	5					5		3	:-		(.)		(.1	5	Э	0.2	5	
68/82/90			! ";			Ţ			5		3	17)		('')	5	(')	3		. .	! '}	
pea P 07/22/87	(1)	5		(1		0.2)		5	(.1	5	(.1						.1	3	
10/22/97			:>	::		(.)					(.1			1.>	Э	!' }	Э	5			5
04/22/38			1">						:1		:->	5		('1		1:>	3	(.)	3	1.2	
06/29/88			! ")		12	3			(.1		:			;;}	С	3	(.)			5	
05 07/22/81	0.3	693.72		5113.98		2629.98		969.16		744.43	1100.22	786.92								624.3	
10/22/87			688.85	1086.67		1899.51			1039.72		1095.28			1950.49	915.58	992.82	843.58	1837.07		955.73	966.39
04/23/99			706.62	760			828		954		196	÷03		1285	868	846	827	1146		715	
06/29/38			692.07		1435.81	1935.42			882.77		823.84	789.89		1459.43	596.69	757.88	782.35			513.15	
SUN 9155 07/22/87	ė.4	870.6		5383.92		2928.83		1223.36		973.27	1341.73	1022.35								762.31	
10/22/97			911.09	1332.35		2172.99			1293.42		1337.82			2250	1155.53	1225.36 1	080.03	2100.92	-	1209.93	1213.49
04/23/68			928.9	366			1079		1199		1164	1047		1580	1158	1102	1077	1407		865	
83/82/90			682°¢		1659.57	2193.18			1123.79		1051.15	1021.77		1717.7	778.34	16.93	987.95			645.89	
LAB SC 07/22/37	9.5	1178.9		845.8		5059.1		1602.2		1187.1	1860.9	1276.1	1277.2						113502	1037.3	
10/22/87	_		é£11	1261		2194			1726		1942			3534	1555	1681	1421	3362		1223	1584
04/23/88			1125	1165			1269		1612		1119	1244		2192	1440	1345	1242	1794	100900	1544	
06/23/88			10.67		2445	3313			1432		1317	1213		2413	934.9	1179	1234			796.3	

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	I HOLLO	22.3				181				14.6				30.9				342				2.93	
	DITCH 13 M.4	22.8	5°3	15.7	212	201	299	261.6	G	0	0	Ģ	ц.,	33.8	21.4	17.7	2249	340	305	17	E.3	7	0.67
	5		7.5	1.1	220.8		332	269.4	0		0	0	14500		156000	13000	069		1110	223	Я		, 04 36
	MH 205	26.6	23.8			520	115			•	0			646	220			234	262			5. 1	2.3
	561 MN	32	22.6	27.6		965	261	402		0	0	¢		87.9	49.4	56.1		Ē	260	218		3.51	1.75
	th 185	21.5	19.7	21.8		178	202	418		0	•	o		146	11.1	30.2		244	255	240		2.76	1.15
	NN 175	27.8	21	24		11	512	356		0	0	•		119	71.3	20.8		227	265	112		7.53	1.25
	NN 165	32.1	29	31.9		610	583	509		Û	0	¢		626	268	262		295	247	259		5	5.25
	NW 13 21.3				423				0				51.8				124				0.23		
	71 8 21.5		20.4	22.6	191		486	427	¢		3	¢	31		31.4	29.2	236		245	235	0.75		0.39 0.39
	NW 75 24.3	23.9	21.8	23.5	476	478	506	8\$\$	6	0	Ģ	¢	217	224	70.7	60.6	242	53	267	234	4.03	2.7.2	2.75
ê	MU 6 23.8				151				0				37.4				210				2.77		
SULE	25 22	29.6	23.9	27.3		9 7	493	475		Ð	¢	•		186	130	83.7		213	822	122	:		8.3 4.9e
325/G	KN 4d 19.9				501				0				1117				250				0.15		
\mathbf{O}	4 R		22.1				201				•				29:5 29:5				251			:	0.47
	NN 45 31.1	5.15		28.1	583	539		508	¢	0		¢	1060	678		989	11	223		260	=:	17.5	3.28
	NU 35			29.3				441				•				111				243			10.6
	E NH 10	78.5	21.8		225	181	170		0	c	¢		2660		÷.5		756	208	222		2,51	3.83	5
	MN 215	35.4	22.3	23.9		438	824	382		0	0	6		19.7	21.1	23.2		187	213	209	1	2.89	1. 34
	NN 2 20.5				349				Û				31.9				202				72.7		
	0.1 0.1				0.2				6				0.1								(.02		
	PARAMETER DATE DPm Si02 07/22/97	10/22/87	04/22/88	06/23/88	ppn HC03 07/22/87	10/22/87	04/23/86	06/29/38	ppm CO3 07/22/87	10/22/87	04/23/88	06/23/98	ppm C1 07/22/87	10/22/87	04/23/98	06/28/88	ppn 504 07/22/87	18/22/01	04/23/08	96/23/99	28/22/20 N #dd	18/22/01	06/28/88

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DITCH 1		8.35				467.67				123.77				-0.003		
DITCH 13	B.06	8.3	9.04	1.92	298.13	461.05	222	269.49	223,08	410.9	245	214.55	0.604	0.035	0.032	-2.179
5	5,25		5.49	ŧ.%	8973.79		26036	3603.87	181.09		272.29	220.13				
MH 20s		1.32	7.48		FI	623.62	115	C-1		426.48	422			0.094	4.253	
195 195		7.34	7.52	1.91		429.7	404	439.31		382.19	101	332.16		0.502	2.098	-2.13
NN 185		1.32	11.71	8.01		455.18	372	362.97		392.04	114	342.83		6.749	2.919	-1.551
MI 175		3.68	1.1	1.83		435.76	376	324.76		387.94	420	293.62		9.361	1.571	-1, 198
MK 165		7.64	7.9	7.89		922.93	564	555.6		500.3	478	417.46		0.117	-0.049	-0-51
NN 13	1.71												112			
8 MH	1.71		7.43	7.89	339.34		351	347.07	380.55		399	374.81	758		1.675	-2.747
st kh	7.56	1.47	11.1	1.62	517.76	491.14	369	390.21	390.4	392.04	415	367.43	560	-0.387	2.934	-0.56
9 MA	1.74				400.5				369.89				0.145			
NN 5s		7.5	11.1	1.85		598.4	526	189.36		410	396	389.58		0.367	1.63	-0.732
94 MA	7.69				127.47				110.9				-,977			
NN 4m			7.59				383				ŧ				0.995	
MM 45	1.17	7.01		7.5	172.59	512.64		613.96	463.08	442.07		416.64	413	0.204		-0.592
MU 35				7,48				526.34				361.69				-0.823
5 M	7.39	7.54	7.51		1264.75	429.51	1 62		434.33	396.96	385		0.498	0,091	2.663	
NJ 215		7.48	7.29	7.85		\$17.11	398.05	415.3		359.23	359.23	313.3		-0.315	2.525	-0.78
NU 2	7.52				490.48				286.23				019			
0A/0C	5.65				0				0.16				0.055			
DATE	11/22/97	0/22/87	14/23/68	6/28/88	122187	0/22/87	4/23/68	16/28/88	7/22/87	0/22/87	14/23/88	6/28/68	11/22/87	10/22/87	M123/68	06/23/98
PARANE TER	LAP OH (÷	Ŧ	HARDNESS O		÷	9	ALKALIN 0	-	ç		10N 9AL 0		~	~

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	NETAL A	NALYSES OF LOCATION	WATER SAM	PLES AT	THE IVERSO	N SITE, RI	ICHLAND,	CO. (DISSOCUED)
PARAMETE	R DAT	E WW 21s	MN 3	MW 35	MW 45	MW 16s	MW 20s	L3
ppb Al	07/21/8	7	<30					
	10/22/8	7	< 30				<30	
	04/23/8	8	<30					80
	06/28/8	8 <30		<30	〈 30	(30_		150
ppb Ag	07/21/8	7	<2					
	10/22/8	7	<2				<2	
	04/23/8	6	<2					44
	06/28/8	8 <2		<2	₹2	<2		94
pab 8	07/21/8	7	1030					
	10/22/8	7	220				570	
	04/23/8	8	140					38000
	06/28/8	8 120		400	820	600		38000
ppb Cd	07/21/8	7	<2					
	10/22/8	7	<2				<2	
	04/23/9	8	<2					16
	06/28/8	8 (2		<2	2	<2		30
pob Cr	07/21/8	17	(2					
•	10/22/8	7	<2				<2	
	04/23/8	8	〈2					14
	06/28/8	8 <2		<2	<2	<2		12
oob Cu	07/21/8	7	<2					
	10/22/8	7	4				<2	
	04/23/8	8	<2					1920
	06/28/8	B <2		<2	11	<2		890
ppb Li	07/21/8	17	370					
	10/22/8	17	71				160	
	04/23/8	18	20					13100
	06/28/8	8 20		130) 190	100		10000
peb No	07/21/8	7	<20					
	10/22/5	7	{20				<20	
	04/23/8	8	<20					50
	06/28/8	8 (20		(20) (20	<20		20
pob Ba	07/21/9	17	5 30					
	10/22/8	17					92	
	04/23/8	8	6 0					320
	06/28/9	19		110) 100			270
pob Pb	07/21/8	17	₹40				<40	
	10/22/8	7						
	04/23/8	8	(40					350
	06/28/9	8		(40) (4 0			270

CON	τ.								
PARAME	TER DATE	MW 21s	周報 3	HW 35	MW 4c	MW 14c	MH 20c	17	
pob Ni	07/21/87		(10			114 103	HM 243	LJ	
•	10/22/B7		(10				210		
	04/23/88		(10				10		
	06/28/88	(10	` * v	710	/10	/ 1 A		160	
				715	/10	CIU		150	
oob Sr	07/21/87		7300						
1 12	10/22/87		1020				00.00		
	04/23/89		1020				2510		
	04/28/89	430	110	7/10	2204			241000	
	00,20,00	000		100V	2730	740		207900	
onh Ti	07/21/87		23						
,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,	10/22/B7		3				Ę		
	04/23/88		a d				له	/ 1	
	04/28/88	4	••	4	0	E.			
		•		U	9	5		71	
pob V	07/21/87		α						
••	10/22/87		d				71		
	04/23/88		a				14	51	
	96/28/88	<1	••	(1	1	23		15 15	
				•	11	11		10	
pob Zn	07/21/87		14						
• •	10/22/87		4				13		
	04/23/88		(3					020	
	06/28/88	(3		٤.	3	(3		430	
				••	· ·			AAA	
pob Ir	07/21/87		(4						
••	10/22/87		(3				24		
	04/23/88		(4					66	
	06/28/88	<4		<4	(4	{4		66	
pob As	07/21/87		0.5						
-	10/22/87						0.5		
	04/23/88		0.7					0.3	
	06/28/99			(.1				2	
								-	
pob Se	07/21/87								
•••	10/22/87						1		
	04/23/88		1.2				•	0.6	
	06/28/88							2.2	
ppb Be	07/21/87		21						
	10/22/87								
	04/23/88		(1					26	
	06/28/88		• •	1				21	
				*				**	

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APPENDIX M

CLAY MINERALOGY



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APPENDIX N

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CHLORIDE DATA AND SC DATA/PLOTS



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APPENDIX O

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YELLOWSTONE RIVER CHEMISTRY AND DEWEY'S DATA

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CHLORIDE CONCENTRATION IN opm 10/15/82 01/15/83 02/15/83 03/15/83 04/15/83 06/15/83 WELL 3 560 600 600 321 220 150 WELL 5 445 264 192 161 214 395 WELL 1 33.5 27.8 19.1 20.7 106 25 SULFATE CONCENTRATION IN ppm	
LHLUMIDE CUNLENTRATION IN Ope 10/15/82 01/15/83 02/15/83 03/15/83 04/15/83 06/15/83 WELL 3 560 600 600 321 220 150 WELL 5 445 264 192 161 214 395 NELL 1 33.5 27.8 19.1 20.7 106 25 SULFATE CONCENTRATION IN ppa	
10/13/82 01/15/83 02/15/83 03/15/83 04/15/83 06/15/83 NELL 3 560 600 600 321 220 150 WELL 5 445 264 192 161 214 395 NELL 1 33.5 27.8 19.1 20.7 106 25 SULFATE CONCENTRATION IN ppa	
WELL 3 560 600 600 321 220 150 WELL 5 445 264 192 161 214 395 WELL 1 33.5 27.8 19.1 20.7 106 25 SULFATE CONCENTRATION IN ppa	
WELL 5 445 264 192 161 214 395 WELL 1 33.5 27.8 19.1 20.7 106 25 SULFATE CONCENTRATION IN ppm	
NELL 1 33.5 27.8 19.1 20.7 106 25 SULFATE CONCENTRATION IN ppm	
SULFATE CONCENTRATION IN ppa	
SULFAIL CUNCENIKAIIUN IN PPa	
WELL 3 263 252 260 254 260 264	
NELL D 234 223 231 240 249 247	
NELL 1 217 217 225 225 240 237	
SPECIFIC CONDUCTANCE IN weber (or DE C	
00/10/00 A0/04/00 A0/04/00 A0/A4/00 A0/A0/00 AA/00/00 A4/0	74/07
V0/10/02 V0/24/02 V7/V1/02 V7/V4/02 V7/V7/02 10/25/02 V1/3	1600
WELL 5 2217 1/70 1002 1034 20/3	1322
WELL 0 2000 5362 2339	2333
WEEL 1 1125 1135 487 1128 1345 1251	1213
(DASSOCUES IONS) CHEMICAL PARAMETERS AUGUST, 1982 (DISSOCUES IONS))
PARAMETER WELL 1 WELL 2 NELL 3 DITCH BL	-
Ca 85.3 93.1 124.8 56.9	
Ma 38.1 31.6 70 35.5	
Na 114 102.4 285.7 84.2	
K 6.8 6.4 11.4 5.1	
HC03 436.B 424.6 510 305	
C03 0 0 0 0	
CI 14.5 13.5 358 17.5	
S84 230 195 250 190	
pH 7.77 7.69 7.78 8.21	
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2	
pH 7.77 7.69 7.78 8.21 TDS 925 855.6 1619 695.2 CONDUCT. 1139 1085 2312 911	
pH 7.77 7.69 7.78 8.21 TDS 925 855.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289	
pH7.777.697.788.21TDS925856.61619695.2CONDUCT.113910852312911HARDNESS370363600289ALKALIN.358349418250	
pH7.777.697.788.21TDS925856.61619695.2CONDUCT.113910852312911HARDNESS370363600289ALKALIN.358348418250SAR2.62.35.12.2	
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 358 348 418 250 SAR 2.6 2.3 5.1 2.2 (DDB) CHEMICAL PARAMIERS FOR SEPT. 1982 (DISSOLVEN ICT)	(2Mc
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 358 348 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMITERS FOR SEPT. 1982 (Dissolves) 100 PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L	(240
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 348 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMIERE FOR SEPT. 1982 (Dissources ic) PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 283 ALKALIN. 35B 349 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMIERS FOR SEPT. 1982 (Dissources ic) PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mo 35.04 40.8 49.5 60.1 29.5	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 866.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 348 418 250 SAR 2.6 2.3 5.1 2.2 (ppe) CHEMICAL PARAMIERS FOR SEPT. 1982 (Dissocutes ic) PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 81 Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 855.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 358 348 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMIERS FOR SEPT. 1982 (Dissources) 100 PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 348 418 250 SAR 2.6 2.3 5.1 2.2 (ppe) CHEMICAL PARAMTERS FOR SEPT. 1982 (Dissources ic) PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8 NF03 392.8 395.3 475.8 283.1 293	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 349 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMIERS FOR SEPT. 1982 CDISSOLUERS IC PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 81 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0	(ZMC
pH 7.77 7.69 7.78 8.21 TDS 925 866.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 348 418 250 SAR 2.6 2.3 5.1 2.2 (ppe) CHEMICAL PARAMTERS FOR SEPT. 1982 (Dissocures) ic PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0 C03 0 0 0 0 17.5	(ZMC
pH 7.77 7.69 7.78 8.21 TDS 925 866.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 358 348 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMIERS FOR SEPT. 1982 (DisSoures let some second s	נאכ)
pH 7.77 7.69 7.78 8.21 TDS 925 866.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 348 418 250 SAR 2.6 2.3 5.1 2.2 (ppm) CHEMICAL PARAMTERS FOR SEPT. 1982 (Dissocutes is PARAMETER WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0 0 C1 29 14.8 200 450 17.5 504 S04 225 235 25	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 855.6 1619 695.2 CONDUCT. 1139 1085 2312 911 MARDNESS 370 363 600 283 ALKALIN. 358 349 418 250 SAR 2.6 2.3 5.1 2.2 (ppe) CHEMICAL PARAMTERS FOR SEPT. 1982 (DisSocular for secolar for for secolar for secolar for secolar for secolar for secolar for secolar for for secolar for secolar for secolar for secolar for secolar for secolar for for secolar for secolar for secolar for secolar for secolar for secolar for seco	נאכ)
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 358 349 418 250 SAR 2.6 2.3 5.1 2.2 (ppe) CHEMICAL PARAMTERS FOR SEPT. 1982 (Dissocutes ic) PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0 C1 29 14.8 200 450 17.5 S04 225 235 250	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 866.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 35B 348 418 250 SAR 2.6 2.3 5.1 2.2 (pps) CHEMICAL PARAMYERS FOR SEPT. 1982 (DisSocures ic PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0 0 C1 29 14.8 200 450 17.5 504 C03 0 0	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 856.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 283 ALKALIN. 358 348 418 250 SAR 2.6 2.3 5.1 2.2 (ppm) CHEMICAL PARAMIERS FOR SEPT. 1982 (DisSocures ic) PARAMETER WELL 1 WELL 2 WELL 3 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 81 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0 C1 29 14.8 200 450 17.5 S04 225 235 250 225	(2MC
pH 7.77 7.69 7.78 8.21 TDS 925 866.6 1619 695.2 CONDUCT. 1139 1085 2312 911 HARDNESS 370 363 600 289 ALKALIN. 358 349 418 250 SAR 2.6 2.3 5.1 2.2 (ppm) CHEMICAL PARAMYERS FOR SEPT. 1982 (Dissocutes for sept) PARAMETER WELL 1 WELL 2 WELL 3 USEL 5 PARAMETER WELL 1 WELL 2 WELL 5 DITCH 8L Ca 89.2 67.2 89.6 156.8 58.9 Mg 35.04 40.8 49.5 60.1 29.5 Na 113.6 111.2 241.9 261 B1 K 7.4 8 10.1 13.4 5.8 HC03 392.8 395.3 475.8 283.1 293 C03 0 0 0 0 0 0 C1 29 14.8 200 <	(2MC

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	LDCA	15/51	IVS4	EUPN		5/5A	11/SA	BURN		5/54	YS/T	BURN		5/SX	1:51	NADS		5/SX	1/54	FUKN		5/SA	1/57	196 9		1SV			/S/ /S/	Ξ'n2
L HC	TION OC	IDNEY	1883	S E	1105	TONEY	ERRY	۲ ۲	SULF	10HEY	ERRY	5	כארנ	TONEY	ERPY	20	1104	1 DAEY	CERY	19 CX	NAG!	SEDNEY	IERPY	€ CX	as	SIGNEY	IEANY VS CK	412	SIDNEY TERRY	HS LF.
felde teol TE	1-35 M	:	:	6.5	UN (pps)	6t	2	200	ATE (pon)	19(-	96 i	ē11	(neq) MUL	6	5	3	IDI WILES!	4.7	3.7	ê.]	41 MILS3	ដ	51	83	(add) SGI	476		EAN FI DN	11100	1.1
-	51-10	2	입	6.7		13	3	200	~	300	206	1£0	_	ت	5	9.	(89	3.5	5.4	7.5	(8 5	51	ដ	F.		69¥		IN CES	9990 16600	
	Dec - 75	=	5	6.9		69	22	061		210	240	130		59	22	F		3.4	ы. 5	13		ž	2			506			12696 6760	ſ
	Jan-76	=	13	5.5		Ç,	ŝ	200		190	216	460		90 90	19	2		3.8	3.ċ	7,8		ន	23	f :		476			10800 9730 5-1	147
	Feb-76	Ξ	Ξ			2	29	5		220	972 139	140		ដ	Ĵ,	5		3.6	J. 5	6.2		23	2	F .		23			00 0 01 00011	3
	Nar-76	81	2	4.6		ŝ	13	120		022	2eé	200		63	. 9	8			1.1	7.9		38	23	3		1/5			13200 10660	-
	Àpr - 76	6	=	b. 1		6 8	5	210		246	973 270	540		6 6	6	61		5.5	1.1			5	2	i.		515			12100 15700	
	Nav-76	5.2	2	5.4		2	2	230		63	ŝ	530		36	2	39		2.1	2.4	9.3		Ξ	2	5		111			37600 32500 A. e	170
	Jun-76	5.6	6.E	5.5		12	×	200		140	96	4 29		88	2	85		2.3	2.3	8.S		::	=	83		274			33300 42106	0.F
	9/-19 f	6. J	÷.4	7.3		R	8	230		6	ž	1 90		2	8	\$		2.6	2.3	9.8		21	2	2		151			13800 26506 0 77	
	Aug-76	8.4	E. 1	9.6		5	=	220		346	136	961		46	4	6 1		3.5	3.2	9.6		2	2	"		356			6140 9590 978	A.10
	5ep-76	6.6	۴.۲	6.8		5	4	226		130	150	430		1 8	41	8		3.7	3.2	9.5		7	£	£.		353			9300 9154	10.0
,	9ct-76	:	Ξ	è.4		59 59	5	200		220	210	9 <u>0</u>		61	51	ė.		3.9	3.5	1.1		5	23	2		217			9590 10000	9
I	Mov-76	2	Û	8.3		13	5	279		260	ñ	530		11	55	vis		3.9	3.7	ġ.		97	R	001		579			7600 9360	-
•	Dec-76	E	=	b.2		23	53	140		922	210	390		66	53	£		3.7	3.5	b.b		22	23	2		955			7606 8200	
	11-nst	=	e :			19	: ?			100	260			36	76			1.1	3.9			2	29			244			9800 3930	
	Feb-77	57	2	6.4		50	64	19 (236	216	350		65	96	5		3.6	3.6	6.B		z	5	19		506			7400 8230 1 9	
	fiar-77	23	# 3	5.5		6	1	180		360	38 2	25		53	5 9	5		1.1	-	5.6		Ş	8	5		194			6700 5960 7 5	•••
	Aor-77	16	1	7.1		11	6	260		350	580	919		59	7 9	11		17	1 .1	6.¢		2	5	5		280			6300 6710 2 1	1.3
	Nav-77)()	12	4.4		() 9	5	240		210	196	520		5	5	3		3.6	3.5	9.4		2	2	8		462			11100	1•1
	lt-nut	6.]	6.8	6.9		4	ន	266		021	911	916 216		4	22	5		2.9	2.5	8,9		2	1	8)		317			10500 15800	6°3
	li-Int	=	8.	7.6		1	2	987		220	260	260		5	ដ	5		14	1.9	=		79	2	6		513			3636 5356	U.J
	Aeg-77	12	Ξ			6	5	i		250	ù 1 2			з,	5			ŝ	1			ដ	22			578			4000 4100	
	Sep-77	91	2	7.3		ž	99	2		210	220	661		ន	61	7 7		1.2	-	8.5 1		2	12	"		115			5170 5320	C. 67

APPENDIX P

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EM 10,20, AND 40 m DATA/PLOTS





EM 10V DATA PLOT



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10 METRE EM - HORIZONTIAL COIL, CI=20 mmhos/m

10 METRE EM - VERTICAL COIL, CI=10 mmhos/m



EM 20H DATA PLOT



EM 20V DATA PLOT



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20 METRE EM - HORIZONTIAL COIL, CI=15 mmhos/m



20 METRE EM - VERTICAL COIL, CI=5 mmhos/m





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(SMALL #'S ARE CI)



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10m EM-HORZ.(mmhos/m) PLOTTED with SC (umhos/cm)

(SMALL #'S APE SC)

10m EM-VERT.(mmhos/m) PLOTTED with SC (mmhos/cm) • 1500



APPENDIX Q

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RESISTIVITY DATA/PLOTS

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SAME SCALE AS 10 FL DATA

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APPENDIX R

JULY, 87 TO JUNE, 88 SIMULATED FLOW FIELDS



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DECEMBER









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APPENDIX S

· ...

CONSTANT HEAD DATA FOR MODELING

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	HO	INTHLY C	CONSTANT HEA	NO LINPUT FOR FLOW	NODEL						
NONTH	ì	j	K HEAD	MONTH	ì	j	K HEAD	MONTH	i	j	K HEAO
7	1	1	1908.65	9	i	1	1908.93	9	1	i	1907.9
7	2	1	1908.6	9	2	1	1908.88	9	2	1	1907.85
7	3	1	1908.44	8	3	- t	1908.75	9		;	1907 R
7	4	1	1908.3	ĥ	i		1908 45	, o	J J	i	1907 7
7	5	1	1908 49	9	, i		1000.00	,	т с		1007 10
7	4		1909 5				1700.73	1		1	1707.77
,	7		1000 51	о А	5	1	1908.76	4	0	1	1907.79
,	0		1700.JI	8		1	1400-19	9	1	1	1907.79
<u>,</u>	0		1708.31	8	8	1	1908.76	9	8	1	1907.82
,	,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,	1	1408.51	B	9	i	1908.76	9	9	1	1907.84
	10	1	1908.52	B	10	1	1908.77	9	19	1	1907.85
,	31	1	1908.52	8	11	1	1908.77	ዮ	11	1	1907.86
7	12	1	1908.52	8	12	1	1908.77	9	12	1	1907.87
7	13	1	1908.52	8	13	1	1998.77	9	13	1	1907.80
7	14	1	1908.53	9	14	1	1908.78	9	14	1	1907.89
7	15	1	1908.53	8	35	1	1908.78	9	15	1	1907.89
7	16	t	1908.53	8	16	1	1908.78	s	16	1	1907.9
7	17	1	1908.53	8	17	ī	1908.78	9	17	1	1907.9
7	19	1	1908.54	, ,	18		1908.29	, q	18	1	1907.9
7	1	24	1911.65	Ř		24	1911 95	, 0	1	25	1010 15
;	;	24	1911 64		3	21	1011 04	,	,	21	1010-13
,	ь т	24	1011 64	0	2	21	1011 04	7	4	24	1719.14
<u>'</u>		1	1711.09	8	3	24	1411.44	4	3	24	1910.14
1	•	24	1411.63	8	4	24	1911.93	9	4	24	1919.13
7	5	24	1911.63	8	5	24	1911.93	9	5	24	1910.13
7	6	24	1911.63	8	6	24	1911.93	9	ģ	24	1910.13
1	7	24	1911.53	9	7	24	1911.93	9	7	24	1910.13
7	8	24	1911.63	8	8	24	1911.93	9	8	24	1910.13
7	9	24	1911.5	8	9	24	1911.8	9	9	24	1910.08
7	10	24	1911.2	8	10	24	1911.4	9	10	24	1910.06
7	11	24	1911	8	11	28	1911.1	9		24	1909.15
7	12	24	1911 05	9	12	24	1911 32	e	17	24	1909 95
,	17	24	1011.00	0	17	27	1711.01	,	14	11	1010 00
,	13	27	4731+4	a 0	1.5	24	1711.0	7	13	24	1710.00
,	14	24	1711.34	3	14	29	1711.75	1	11	29	1410.08
1	15	24	1911.34	8	15	24	1911.74	Ŷ	15	24	1910.09
/	15	24	1411.35	8	16	24	1911.75	9	16	24	1710.1
7	17	24	1911.35	9	17	24	1911.75	9	17	24	1710.11
7	18	24	1911.39	8	18	24	1911.78	9	18	24	1919.13
7	4	1	1909.3	9	4	1	1908.55	9	4	i	1907.7
7	6	2	1998.55	8	6	2	1909	9	6	2	1997.75
7	9	3	1908.75	8	9	3	1909.2	9	9	3	1707.8
7	9	4	1909.82	8	9	4	1909.28	9	9	4	1907.83
7	9	5	1908.87	8	9	5	1909.36	9	9	5	1907.86
,	ģ	-	1908.9	9	9	Å	1909.4	ç	9	6	1907.9
, ,	é	7	19/00 97	9	ġ	7	1909 41			,	1907.97
7	7	ó	1000 04	D 0	, a	, D	1000 10	, G	, a	, 0	1907 97
,	7	0	1700.T3 1040.00	6	1	0 ^	1003 65	7	5	0 0	1260
1	¥ -	¥	1700.77	8	4	¥ •^	1000 50	7	7	7	1000 A.
/	4	10	1909.02	8	4	16	1909.38	7	7	10	1405.04
7	9	11	1909.07	9	9	11	1909.63	9	9	11	1408.64
7	9	12	1907.13	8	9	17	1909.69	9	ų	12	1708.13
7	9	13	1909.25	6	9	13	1909.75	9	ò	13	1908.19
7	9	14	1909.39	8	9	14	1909.79	9	9	14	1908.21
7	9	15	1909.5	8	9	15	1909.82	9	9	15	1908.24
7	9	16	1909.6	8	9	16	1909.86	9	9	15	1908.27
7	9	17	1909.7	9	9	17	1909.9	9	9	17	1908.34
7	á	19	1909 9	<u> </u>	á	19	1909.95	9	9	18	1909.4
	, a	10	1906 9	D	, 0	10	1000 20	, 0		10	1908 45
	Ţ Ċ	17	1797+7		1	17	1010 03	, 0	, o	30	1000 5
•	7	20	1719	- U	1	29	1717-04	,	י ה	2V 01	1000 55
1	4	21	1910.1	B	4	21	1410.00	¥	, , , , , , , , , , , , , , , , , , ,	21	1000.13
7	9	22	1910.22	8	9	22	1910.11	9	9	22	1408.03
7	9	23	1910.4	8	Ģ	23	1910.16	9	9	13	1908.8
7	11	24	1911	8	11	24	1911.1	9	11	24	1909.15

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YONTH	:	:	W 187 AD	M011711							* 11540
10			K DEHU	nunin	1	1	K HEAD	PUNTH	1	,	K HERU
10	1	1	1906./8	11	1	ĩ	1906.3	12	1	1	1906.23
10	2	1	1906.75	11	2	1	1906.29	12	2	1	1906.22
10	3	1	1906.73	11	2	1	1906.28	12	3	1	1906.08
10	4	1	1906.53	11	4	1	1905.98	12	4	1	1905.94
10	5	1	1906.73	11	5	1	1906.29	12	5	1	1905.08
10	6	1	1906 74	11			1001 70	12	ž		1001 0
10	7		1004 75		7	1	1740-17	14			1001 9
10	,		1100.73	11	1	1	1946.29	14			1708.2
10	8	1	1906.75	11	8	1	1906.29	12	8	1	1906.2
10	4	1	1906.75	11	9	1	1906.29	12	9	្រ	1906.21
10	19	1	1906.75	11	10	1	1906.3	12	10	1	1906.21
10	11	1	1905.76	11	11	1	1906.31	12	11	1	1906.22
10	12	1	1906.76	-11	12	1	1906.33	12	12	1	1906.22
10	13	1	1906.76	11	13	1	1906.35	12	13	1	1996.23
10	14	1	1906.77	11	14		1904 37	12	14	i	1205 25
10	15	i	1906 17	11	15		1001 30	17	15		1904 77
10	14		1004 70	41	13	:	1001 41	12	1.		100.21
10	17		1790.70	11	10	1	1706.41	12	10	,	1790.27
10	17	1	1790.77	11	17	Ļ	1905.43	12	<i>U</i>	1	1906.31
10	18	1	1906.8	i !	18	1	1906.45	12	19	1	1906.33
10	1	24	1907.37	11	1	24	1908.7	12	1	24	1908.3
10	2	24	1909.36	11	2	24	1908.69	12	2	24	1908.29
10	2	24	1909.36	11	3	24	1908.69	12	3	24	1908.29
10	4	24	1909.35	11		4	1908 48	17	i	24	1008 70
10	Ś	24	1909 35		Ę	34	1000 10	12	, ,	74	1000 70
10	,	24	1000 35	**	5	29	1708.00	14		29	1703.27
10	•	24	1909.35	11	6	24	1408.93	12	b	25	1409.29
19	1	24	1909.35	11	1	24	1909.69	12	1	24	1908.20
10	8	24	1909.35	11	8	24	1908.68	12	B	24	1908.28
10	9	24	1909.3	11	9	24	1909.63	12	ò	24	1908.26
10	10	24	1909.2	11	10	24	1908.53	12	10	24	1908.24
10	11	24	1908.65	11	11	74	1908.19	12	11	74	1997.84
10	12	74	1909 23	11	17	24	1000 7	17	17	74	1000 34
44	17	27	1000 35	11	12	27	1000.7	12	17	74	100,10
10	15		1107.20	11	13	24	1908.72	12	13	24	1908.27
10	14	24	1909.25	11	14	24	1908.74	12	14	24	1998.28
10	15	24	1909.26	11	15	24	1909.75	12	15	24	1908.3
10	15	24	1909.27	11	16	24	1908.78	12	16	24	1908.32
10	17	21	1909.28	11	17	24	1908.8	12	17	24	1905.34
10	18	24	1909.3	11	18	24	1908.82	12	18	24	1908.36
10	4	1	1906.53	11	4	- 1	1905.88	12	4	1	1905.74
10	ì	;	1964 41		i i	•	10AL AL	12		,	1906 14
10		,	1002 20	11		, , , , , , , , , , , , , , , , , , ,	1001 22	12	e e		1001 74
10	,	3	1700.07	11	7	3	1798.22	14	7	3	1700.24
10	9	4	1906.77	11	9	4	1906.3	12	y y		1906.31
10	9	5	1906.85	1!	9	5	1906.38	12	9	5	1905.4
10	9	6	1905.97	11	9	6	1906.5	12	9	6	1906.45
10	9	7	1907.09	11	9	7	1906.62	12	9	7	1905.56
10	9	B	1907.21	11	9	8	1906.74	12	9	8	1906.64
10	9		1907 33	11	ė	9	1706.86	12	9	9	1906.78
10	ó	10	1007 45	11	á	10	1001 00	17	ġ	10	1904 97
10	,	19	1707.13	**	,	17	1007 1	10	á		100/ 05
10	Y	11	1997.57	11	4	11	190/.1	12	7	11	1700.75
10	9	12	1907.69	11	9	12	1997.19	12	9	12	1907.03
10	9	13	1907.78	11	9	13	1907.26	12	9	13	1907.1
10	9	14	1907.84	11	9	14	1907.34	12	9	14	1907.17
10	9	15	1907.87	11	9	15	1907.4	12	ģ	15	1907.24
10	٩	16	1967 9	11		14	1907 41	17	9	16	1907.31
14	, a	17	1007 07	**			1007 5		, o	17	1997 77
10	7	17	170/.7/	11	7		1007 4	14	,	1/	1007 44
10	Y _	18	1106.62	11	¥.	19	110/.5	12	7	18	170/.94
10	9	19	1908.13	11	9	19	1907.66	12	9	19	1401.2
10	9	20	1908.22	11	9	20	1907.73	12	9	20	1907.55
10	9	21	1908.32	11	9	21	1907.75	12	ģ	21	1907.61
10	9	22	1908.43	11	Ŷ	22	1907.96	12	9	22	1907.65
10	9	23	1908.55	11	, Q	27	1907.99	17	9	21	1707 75
10	, ,	24	19(19 44			4.0 17.4	1000 10	17	11	24	1907 04
10		41	4100.00	11	11	<u> </u>	1100.17			11	

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MONTH	i	j	K HEAD	NONTH	i	i	K HEAD	MCNTH	i	i	K HEAD
1	1	1	1906.16	2	1	i	1905.91	3	1	i	1905.69
1	2	1	1906.15	2	2	i	1905.89	1	2	1	1905.69
1	3	1	1906.03	2	3	1	1905.88	3	3	i	1705.69
1	4	1	1905.9	2	4	i	1965.76	1	i	1	1905.82
1	5	1	1906.03	2	5	1	1905.92	1	5	1	1905.79
1	6	1	1906.15	2	- 6	i	1905.93	3	Ĩ,	i	1905.79
1	1	1	1906.15	2	ĩ	1	1905 94	ג ז	,		1905 79
1	8	1	1906.15	2	P		1905 94	J T	á	1	1905 9
1	9	1	1906.15	,	, ,		1005 05	ت ۲		;	1005.0
1	10	i	1906.16	2	10		1965 05	J 1	10	-	1005 01
i	11	÷	1906.17	2	11		1005 04	J 7	10		1703.01
ī	12	í	1906.19	2	11		1703.70	3	11	1	1703.01
- i	13		1906 19	2	11	1	170J.70 1005 01	3	17	1	1945 03
1	14	;	1905 2	1 1	10		1703.77 1005 00	۲ ۲	51 51	1	1703.92
1	15	;	1906 22	÷ ,	15	1	1903.90	3	19	1	1705.82
i	14	;	1004.24		13	1	1903.99	з Т	10	1	1402.87
1	17		1004 94	۲ ۲	10	I	1906	3	10	1	1462.87
1	10		1006.20		17		1796.01	3	17	1	1905.84
	10		1700.40	2	18	1	1905.02	3	18	1	1905.84
	3	24	1708.13	2	1	24	190/.88	3	1	24	1907.41
1	2	21	1908.14	2	2	24	1907.87	2	2	24	1907.41
1	<u>ن</u>	24	1908.14	2	3	24	1907.86	3	3	24	1907.41
1	+	24	1708.13	2	4	24	1997.85	2	4	24	1907.41
1	5	24	1908.13	2	5	24	1907.94	3	5	21	1907.41
1	6	- 24	1908.13	2	6	24	1907.83	3	٤	24	1907.41
1	7.	. 24	1908.13	2	1	24	1907.92	3	1	24	1907.41
1	9	24	1908.12	2	8	24	1907.81	3	8	24	1907.4
1	9	- 24	1908.1	2	9	24	1907.8	3	9	24	1907.4
1	10	24	1908.05	2	10	24	1907.8	3	10	24	1907.4
1	11	24	1907.91	2	11	24	1907.7	3	11	24	1907.44
1	12	24	1908.07	2	12	24	1907.87	3	12	24	1907.4
1	13	24	1908.08	2	13	24	1907.88	3	13	24	1907.4
1	14	24	1908.1	2	14	24	1907.89	3	14	24	1907.4
1	15	24	1908.12	2	15	24	1907.9	3	15	24	1907.41
i	16	24	1908.14	- 2	16	24	1907.91	3	16	74	1907.41
	17	24	1908.14	2	17	24	1907.97	t	17	74	1907 41
	19	74	1000 10	2	19	24	1967.91	्र र	10	74	1907 47
4	10		1005 0	2		1	1005 74	र १		1	1965 67
			1703.7	2	1		1743.70	J T		, ,	1964 67
1	0	4	1700.13	1		1 7	1703.70	Т	0 0	4	1004 94
1	y 0		1906.27				1708.13	л т	7	3	1706.24
1	9	2	1409.23		· · ·		1706.20	Т	7		1795.33
3	4	Ş	\$905.45	2	4	2	1906.33	2	7	3	1708.37
1	9	6	1905.51	2	4	6	1906.4		7		1700.40
1	9	1	1905.6	2	9	1	1905.45	3	4		1908.51
1	9	8	1906.65	2	9	8	1906.5	3	4	8	1905.56
1	9	9	1906.7	2	9	9	1906.55	2		9	1906.51
1	9	10	1906.76	2	9	10	1905.6	3	9	19	1906.66
1	9	11	1906.84	2	9	11	1906.69	3	à	11	1966.69
1	9	12	1906.92	2	9	12	1906.75	3	9	12	1906.75
1	9	13	1906.98	2	9	13	1906.8	3	ò	13	1906.79
1	9	14	1907.04	2	9	14	1906.84	3	à	- 14	1906.84
1	9	15	1907.09	2	9	15	1905.89	3	9	15	1906.08
i	9	16	1907.12	2	9	16	1906.95	3	9	16	1906.92
•	9	17	1907.19	2	ç	17	1907	3	9	17	1905.96
1	á	10	1907 70	3		10	1947 01	3	ę	18	1907
•	7	10	1907 14	4 1	,	10	1907 04	t		19	1907.04
1	0	56	1067 41	1		17 2A	1967 13	ג ז	q	20	1907.08
4	7	20	1707.91	4	т 6	20	1047 3		á	71	1967 13
1	7	21	170/.01	1	7	11	110/.2	37	1	11	1003 17
1	7	22	1907.39	2	9	22	1907-23	3	7		1007.1/
1	y 	23	1907.71	2	. 9	23	1401*24	2		23	1707.20
1	11	24	1997.91	2	11	- 24	1997.7	3	11	24	1201-47

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HONTH	i	j	K HEAD	MONTH	i	i	X HEAD	NONTH	i	i	r head
4	1	1	1906	5	1	í	1908.54	6	1	í	1908.54
4	2	1	1906	5	2	i	1908.45	6	2	i	1908.45
4	3	1	1906	5	3	1	1908.4	6	3	i	1908.25
4	4	1	1906	S		1	1907. 9R	6		i	1907.89
4	5	1	1906	5	5	1	1908.52	6	5		1908.45
4	6	1	1905	Ś	6	ī	1909.53	с к	, ,	•	1909 44
4	7	L	1905	5	7	i	1209 54	5	7		1908 47
4	8	1	1906	5	Ŕ	i	1908 55	5	, 4		1909 4B
4	9 '	1	1906	5	Ģ	i	1968 57	, L	ä	1	1000 5
4	10	1	1906	5	ŁÓ	1	1998.6	6	10		1908.53
4	11	1	1905	5	n	1	1908.62	6	11	÷	1968.55
4	12	1	1906	5	12	1	1908-62		12		1908.55
4	13	1	1906	5	13	i	1908.62	6	13		1908.55
4	14	1	1905	5	14	1	1908.61	Ň	14	i	1908.54
4	15	1	1906	5	15	1	1908.59	6	15	i	1909.52
4	16	1	1905	5	16	1	1968.57	6	16	1	1908.5
4	17	1	1906	5	17	i	1908.57	6	17	1	1908.5
4	18	1	1906	S	18	1	1998.55	6	18	1	1909.48
4	1	24	1908	5	1	24	1911.92	- -	1	74	1911 97
4	2	24	1908	5	2	24	1911.79	5	,	24	1911 94
4	3	24	1908	5	3	24	1911.79	1	i	78	1051 94
4	, i	24	1908	5		24	1911 77	, ,	1	24	1011 07
4	5	24	1908	5	5	26	1911 77	6	Ś	24	1011 07
i i	Å	24	1909	5	, i	24	1911 76			21	1011 01
i i	,	24	1908	5	7	24	1911 74	5	2	54	4711.74
i	8	24	1908	5	R	28	1911 74		á	27 74	1011 01
í	ů,	Ĵa	1908	Š	ő	24	1911 7	6		24	1011 05
, i	10	24	1908	्र	10	51 78	1911 4	5	10	53 74	1711-0J 1011 75
4	11	24	10/10		10	21	1016 05	5	10	21	1711./J
4	17	24	1908	J K	12	24	1011 LL	6		14	1011 01
	11	24	1909	5	12	74	1911.00	6	14	24	1016 97
	14	34	1962	्र इ	1.0	74	1011 7	6	3.5	14	1011 05
	15	54	1900	ں ج	15	24	1911 7	B 4	17	24	1711-73
	15	74	1700	J 5	15	24	1015 22	8	1.0	27 74	1011 01
	17	27	1960	J 4	10	27	1711.00	5	17	74	1011 0
	10	24	1009	J 5	10	44 74	1911 45	0 4	19	24	1711+7
7	10	1	1901	J 5	10	27	1711.03	5	15	1	1711+7
1	1	2	1904 73	J K		· ·	1000 4	6			1000 20
7	0	4 T	1046 4		0 0		1000.7	÷	0 6		1000.20
1	,	ž	1001 80	J E	0		1000.72	u 4			1000 70
	7		1004 57	J	1	, ,	1000.10	10 6	,	, i	1008 00
T	7	لت د	1001 10		1	J	1060 14	a 4	, 9	J 1	1100.72
	7	<i>о</i> т	1001 0	J	7	а т	1960 10	0 4	7	1	1000.00
	4		1900.0	J	,	6	1707.17	6	7		1703.17
+	7	8	1700.72	2	y ~	8	1000 7-	a ,	Y A	5	1940 4/
2	7	4	100/ 1/		Y	7	1707.31	a .	7	1	1000 11
	9	10	1706.10	5	, ,	10	1707.30	0 (7	10	1797.11
•	9	11	1907.28	5	Y	11	1404.4	6	7	11	1707.13
4	Ŷ	12	1997.4	3	Y	12	1909.44	8	y 9	12	1909.19
4	9	13	1907.47	5	9	15	1909.47	6	4	15	1909.22
4	9	14	1907.56	5	9	14	1909.5	6	y y	15	1909.25
4	9	15	1907.58	5	9	15	1909.53	6	à	15	1909.26
4	9	16	1907.61	5	9	16	1909.58	6	9	16	1909.33
4	9	17	1907.68	5	9	17	1909.63	6	9	17	1909.38
4	9	18	1907.78	5	9	18	1909.66	6	9	18	1909.41
4	9	19	1907.84	5	9	19	1909.7	6	9	19	1909.45
4	9	20	1907.93	5	9	20	1909.73	6	Ģ	20	1909.43
4	9	21	1908.03	5	9	21	1909.76	6	ģ	21	1909.51
4	9	22	1908.14	5	9	22	1709.95	6	ò	22	1909.55
4	9	23	1908.23	5	9	23	1910.1	6	9	23	1909.58
4	11	24	1908.37	5	11	24	1910.85	6	11	24	1907.64

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APPENDIX T

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FLOW MODELING STATISTICS

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	COMPARISO	N OF DRAW	down of a	ICT. VS. 9	ih. At
	TWO DIFFE	RENT PUMP	ING NODES	AND THO	K VALUES
		PUMP=	PUMP=	PUNP =	PUMP=
DISTANCE		14,11	13,10	13,10	14.11
AMAY	ACTUAL s	LOW K	LOW K	HI K	HI K
0	1906.58	1907.24	1907.28	1907.4	1907.41
35	1907.41	1907.42	1907.41	1907.45	1907.46
82	1907.46	1907:44	1907.44	1907.48	1907.48
119	1907.465	1907.47	1907.46	1907.49	1907.49
136	1907.47	1907.48	1907.4B	1907.495	1907.49
185	1907.475	1907.48	1907.48	1907.5	1907.495
192	1907.46	1907.49	1907.48		1907.495
200	1907.48	1907.49	1907.48		1907.5
260	1907.48	1907.49	1907.47		

DATA FOR MODEL RUN WITH K = 11000 AND T = 192500.

PUMP	ING NODE 1	13,10	PUMPING NODE 14,11					
DISTANCE	DRAWDOWN	WE ELV.	DISTANCE	DRANDOWN	WL ELV.			
0	0.1	1907.4	0	0.09	1907.41			
35	0.04	1907.46	35	0.04	1907.46			
82	0.02	1907.48	82	0.02	1907.48			
118	0.01	1907.49	118	0.01	1907.49			
136	0.005	1907.495	136	0.01	1907.49			
185	0	1907.5	185	0.005	1907.495			
			192	0.005	1907.495			
			200	0	1907.5			

DATA FOR HODLE RUN WITH K = 4400 AND T = 77000.

PUMPII	NS NODE 13	5,10	PUMP:	ING NODE I	14,11
DISTANCE	DRAWDOWN	WE ELV.	DISTANCE	DRAWDOWN	WL ELV.
0	0.22	1907.28	0	0.26	1907.24
35	0.09	1907.41	40	0.08	1907.42
82	0.06	1907.44	80	0.06	1907.44
118	0.04	1907.45	120	0.03	1907.47
136	0.02	1907.4B	160	0.02	1907.48
185	0.02	1907.48	200	0.02	1907.48
192	0.02	1907.48	240	0.01	1907.49
200	0.02	1907.48	290	0.01	1907.49
260	0.03	1907.47	320	0.01	1907.49
ACTUAL	AQ. TEST	DATA			

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	. пат тъчт	WIN1
DISTANCE	DRAWDOWN	WL ELV.
0	0.92	1906.58
35	0.09	1907.41
82	0.04	1907.46
118	0.035	1907.465
136	0.03	1907.47
185	0.025	1907.475
192	0.04	1907.46
200	0.02	1907.48
260	0.02	1907.48

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STEADY-STATE STATISTICS (OCT.)

		HE	90					
NEL	L	ACTUAL	SIMULATED	DIFF.		WELL	ACUTAL	SIMULATED
2		1908.2	1908.5	-0.3		3	1908.99	1908.93
3		1908.1	1908.26	-0.16		4	1908.78	1908.56
4		1907.93	1907.91	0.02		5	1908.69	1908.48
5		1907.85	1907.85	0		7	1908.38	1908.27
6		1907.76	1907.7	0.06		6	1908.56	1908.47
7		1907.6	1907.5	0.1		14	1908.57	1908.8
8		1906.93	1906.81	0.12	007	2	1908.2	1908.5
11	l	1907.68	1907.75	-0.07		3	1908.1	1908.26
12	2	1907.92	1907.97	-0.05		4	1907.93	1907.91
1.	5	1907.08	1907.05	0.03		5	1907.B5	1907.85
14	4	1908.12	1908.05	0.07		-	1907.76	1907.7
15	5	1907.76	1907.74	0.02		7	1907.6	1907.5
1	5	1907.7	1907.7	0		Ŕ	1906.93	1906 81
Ē	7	1907.32	1907.24	0.08		11	1907.68	1907 75
18	3	1907.2	1907.13	0.07		12	1907 97	1907 97
10	- 9	1907.68	1907.64	0.0A		17	1907 00	1007 05
25	,)	1907.72	1907 7	0.07		14	1000 10	1900 05
2	v I	1000 10	1969 5	-0 30		15	1700.12	1700.03
4	1	1700.10	1700.3	-0.32		11	170/./0	1707.74
		TOANC	ICHT STATIS	TICC		10	170/./	170/6/
		ENNER I LI	ICHI JIMIII Tar	1112		1/	1707.32	1907.29
			CIMUM ATEA	01 <i>CE</i>		18	1907.2	1997.13
11.11		HUDIHL	SINULAIED	UIFF.		17	1907.68	1907.64
JUL	8	1908.56	1908.68	-0.1		20	1907.72	1907.7
	13	1908.74	1908.98	-0.24		21	1908.18	1908.5
	11	1909.38	1909.6	-0.22	NOV	8	1906.31	1906.37
	12	1909.68	1909.74	-0.06		13	1906.72	1906.56
	2	1910.11	1910.42	-0.31		18	1906.91	1906.65
	3	1909.94	1910.12	-0.19		11	1907.25	1907.23
	- 4	1909.9	1909.66	0.24		12	1907.49	1907.47
	5	1909.65	1909.6	0.05		2	1907.73	1908
	7	1909.35	1909.27	0.08		21	1907.72	1908.02
	6	1909.55	1909.5	0.05		3	1907.64	1907.77
	14	1909.96	1909.9	0.06		4	1907.54	1907.41
AUG	8	1908.79	1908.98	-0.19		5	1907.42	1907.37
	13	1909	1909.41	-0.41		16	1907.3	1907.21
	11	1909.76	1909.94	-0.19		20	1907.25	1907.23
	12	1910.03	1910.04	-0.01		7	1907.19	1907.04
	,	1910.39	1910.59	-0.21		19	1967.27	1907.18
	रे	1010 10	1010 30	-0.21			1907 34	1907 24
	Å	1010 17	1010.01	0.21 A 17		15	1067 37	1907 28
	र इ.	1/10/10	1900 05	0 0L		10	1007 14	1007 50
	J	1010.01	1707,73	-0.00		19	17V/.00	170/.30
	1	1000 0	1710.67	-0.01	ħ <i>₽</i> =	17	1700.70	1700./7
	0	1707.7	1404.81	0.05	DEC	8	1708.90	1798.3
	14	1710.34	1410.23	9.11		15	1906-02	1700.31
327	12	1907.81	1907.99	-0.18		18	1906.65	1405.57
	11	1708.45	1908.45	0		11	1907.04	1907
	12	1908.23	1908.46	-9.23		12	1907.26	1907.3
	2	1707.04	1909.07	-0.03		2	1907.48	1907.71

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DIFF.

0.06

0.22

0.11

0.09

-0.23

-0.3

-0.16

0.02

0.06

0.1

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-0.07

-0.05

0.03

0.07

0.02

0.08

0.07

0.04

0.02

-0.32

-0.06

0.16

0.26

0.02

-0.27

-0.3

-0.13

0.13

0.05

0.09

0.02

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0.1

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0.08

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0.15

0.07

-0.04

-0.23

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	WELL	ACUTAL	SIMULATED	DIFF.		WELL	ACIITAI	SIMULATED	OIFF.
	21	1907.46	1907.71	-0.25		17	1906.88	1906.89	0
	3	1907.39	1907.51	-0.12		2	1904.99	1907.11	-0.12
	4	1907.3	1907.21	0.09		21	1907.04	1907.09	-0.05
	5	1907.2	1907.19	0.01		4	1906.93	1906.76	0.17
	16	1907.08	1907.04	0.04		5	1905.84	1906.78	0.06
	20	1907.09	1907.06	0.03		16	1906.73	1906.68	0.05
	7	1907	1906.9	0.1		20	1905.74	1906.67	0.07
	19	1907.04	1907.01	0.03		7	1906.45	1906.66	-0.01
	6	1907.11	1907.06	0.05		19	1906.7	1906.62	0.08
	15	1907.12	1907.09	0.03		6	1906.75	1906.64	0.11
	14	1907.4	1907.33	0.07		15	1906.75	1906.66	0.09
	17	1906.64	1906.6B	-0.04		14	1907.02	1906.79	0.23
JAN	8	1906.34	1906.27	0.07		17	1906.43	1905.43	0
	13	1906.44	1906.51	-0.07	APR	B	1906.11	1905.7	0.41
	18	1906.57	1906.58	-0.01		13	1906.14	1906.17	-0.03
	11	1906.92	1906.98	-0.06		18	1905.29	1906.25	0.04
	12	1907.09	1907.17	-0.08		11	1906.58	1906.54	0.04
	2	1907.29	1907.59	-0.3		12	1906.77	1906.76	0.01
	21	1907.27	1907.58	-0.31		2	1906.95	1907.07	-0.12
	3	1907.2	1907.39	-0.19		21	1906.94	1907.05	-0.11
	4	1907.15	1907.1	0.05		3	1906.92	1906.89	0.04
	5	1907.05	1907.08	-0.03		14	1906.89	1906.68	0.21
	16	1907.83	1907.94	-0.11		5	1905.73	1905.64	0.09
	20	1906.86	1906.97	-0.11		Ĩ	1906. Bt	1906.63	0.18
	7	1906.86	1906.83	0.03		16	1906.62	1906.52	0.1
	19	1906.89	1906.92	-0.03		19	1906.59	1906.46	0.17
	6	1906.97	1906.97	0		5	1906 61	1906 49	0 15
	15	1906.96	1907	-0.04		15	1904 43	1906.51	0 12
	14	1907.22	1907.22	0		7	1904.55	1904.4	0.15
	17	1906.62	1906.65	-0.03		17	77 4091	1906 23	0.15
FFR	R	1906.23	1906.08	0.15	MAY	ġ	1908 25	1908.7	-0.45
	13	1906.3	1906.35	-0.05	11771	13	1909.52	1909.08	-0 55
	19	1906 46	1906 47	0.03		19	1905 77	1909 19	-0 41
	11	1004 70	1901.75	0.03		11	1000 4	1909 49	-0.28
	12	1906 97	1966.96	0.01		17	1909 94	1909 74	0.20
		1907 17	1907 34	-0.17		2	1010 20	1910 45	-0.16
	र्म र	1007-17	1707.34	-V.I/ -0.00		21	1010 97	1916 53	-0.26
	ي د	1907.00	1001 0	0.07		- 7 1	1910-14	1910.00	-0.10
	۲ ج	107.01 1064 0	1001 00	0.12			1710-10	1910.23	0.07
	14	1001 07	1700.00 40AL 7L	0.02		тт Т	1000 04	1000 73	0.10
	20	1700.03	1700.70	0.07		ت ۸	1010	10/0 0	0.32
	10	1001 75	1001 15	0.00		12	1000 72	1000 50	0.04
	10	100.7J	1700,0J	4.1 A AE		10	100101	1000 50	0.00
	17	1004 00	1700.73			17	1000 7	1000 15	0.02
	0 16	1006 04	1790.17	V+13 A A4		10	1000 75	17V7.0J (0AD 10	0.0J 0.01
	10	1770.01	1700.0	V.VI 0.00		د د۱	1707./3	1000 #1	0.00
MAG	1/ . 0	1700.3	1708,48	0.02		/	1747.45	17V7.90 1000 AC	V (()
nnn	. 0	1706.1/	1703.77	V.18		1/	1707.28	1909.02	V.Z6
	10	1790.23	1706,38	-0.15	JUN	13	1408.8	1908.92	-0.12
	18	1706.39	1706.44	-0.05		19	1908.79	1909.01	-0.72
	11	1792.7	1706.69	0.01		11	1909.58	1909.55	0.03

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A100 1	* * • • • • • •		
WELL	ACUTAL	SIMULATED	QIFF.
12	1909.86	1909.57	0.29
2	1910.41	1910.34	0.07
21	1910.44	1910.49	-0.05
3	1910.25	1910.15	0.1
14	1910.34	1910.01	0.33
5	1909.74	1909.57	0.17
4	1909.95	1909.69	0.27
16	1909.51	1909.44	0.07
19	1909.55	1909.47	0.08
6	1909.67	1909.55	0.12
15	1909.73	1909.58	0.15
7	1909.37	1909.31	0.06
17	1908.92	1909.12	-0.2

STEADY-STATE RESULTS Regression Dutput: Constant 0 Std Err of Y Est 0.126713 R Squared 0.926028 No. of Observations 18 Degrees of Freedom 17

X Coefficient(s)1.000020 Std Err of Coef.0.000015

TRANSIENT RESULTS	
Regression Output	::
Constant	0
Std Err of Y Est	0.171654
R Squared	0.993229
No. of Observations	187
Degrees of Freedom	195

X Coefficient(s)1.000001 Std Err of Caef.0.000006

APPENDIX U

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RANDOM-WALK LOADING

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APPENDIX V

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SIMULATED C1 PLOTS AND ACTUAL VERSES SIMULATED CHLORIDE DATA

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) L E	' M.				خت	0	D	/H Y Z	>		PAF		LES	= 0	
		· • - ·	н г								_ <u>k</u>	ELL S	SITE (LOCAT	ION	
			С(נסד	[NA	TES	5 A	RE	IN	F	EET				
24		+		+		+	-	+	+		+	•†-	+	+	+	
23		+	~		8	+	-	+	+		+	+	4	+	+	
	i	+	_`*¶	へ		+	•	+	+		+	+	+	+	+	
21 ;		+		1		-	-	+	+		+	+	+	+	+	
20 i	i I	+		•		÷.	-	÷	+-		+	+	+	+	+	
17 1	i	+		7	17	÷	-	+	+		+	+	+	+	+	
10 1	1			I		+	7	+ L	+		+	+	+	+	+	
				7				T	Ŧ		Ŧ	Ŧ	-4.	Ŧ	+	
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14 i		+		•	16	: ⁺ 6		⊢ ·	9		+	C.	9 +	+	+	
15 :		+	T.	t	10	/+ *	-	+	+		+-	+ ***	₩ + 15	+	+	
12 3		+	Ĕ	+		1		+	+	4		+ •	10 +	+	+	
		+	ā	1		+ •	ح	+	+		+	+	+	+	+	
		+		t		+	J.	1 -	+		+	+	+	+	+	
7 i - i		+		+		+	-	† - 	+		+ 1		2 +	+	+	
ម : _ :		+		Ť		+•	-	+	+		+ (21/2	2 +	+	+	
7		+		1		+	-	+	+		+PI	Γ +	+	+	+	
6 i	i			-		-+-	•	+	-		+	+	+	+	+	
		 8		 9					·		1					
		8				1 0		 1 1	1 2		1 3	1 4	1 5	1 6	1 7	
ACCU Conc	INUL: ENTI	8 NTED RATIO	TIM N M	 9 E = AP 1	3ù IN FI	1 0 043	 (S	1 1 P	1 2 'ARTI		1 3 = 3	1 4 328	1 5	1 6	1 7	
ACCU Conc 	INULI ENTI	8 ITED RATIC	TIM DN M	57 E = AP 1 DIN(30 IN FI ATES	1 O DAY PM ARE	rs In F	1 1 P EET	1 2 'ARTI		1 3 = 3	1 4 328	1 5	1 6	1 7	
ACCU CONC 	 INUL: ENT: 	8 NTED NATIO	TIM DN M DOR	= = AP] DIN(+	30 IN FI ATES +	1 O DAT PM ARE +	rs In F	1 1 FEET	1 2 'ARTI	 CLES 	1 3 = ;	1 4 320 +	1 5 7	1 6	1 7	
ACCU CONC 16 15	INULI ENTI 	8 NTED RATIC I	TIMI DN M DOORI + +	 	30 IN FI ATES + +	1 O DAY PM ARE +	rS IN F +	1 1 FEET +	1 2 ARTI + +	CLE5	1 3 = ; +	1 4 328	1 5	1 6	1 7	
ACCU CONC 16 15 14	///9L4 ENTI ENTI	8 NTED RATIO TATIO	TIM DN M DOOR + + +	=== == == == == == == == == ==	30 IN FI ATES + + 127	1 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	(S IN F + +	1 1 EET + +	1 2 ARTIC + + +	CLE5	1 3 = 3 + + +	1 4 328 + + +	1 5 7	1 6	1 7	
ACCU CONC 16 15 14 13		8 TTED RATIO + + +	TIM DN M CDORI + + + +	==== ===== ===== ===== ===== ===== ==== ==== ==== ==== ==== ==== ==== ==== ==== == ==	30 IN FI ATES + + 127 295	1 O PM ARE + + 95 473	(S IN F + + + 126	1 1 EET + + 42	1 2 ARTIC + + + +	 + + + +	1 3 = 5 + + + + +	1 4 328 + + + +	1 5	1 6	1 7	
ACCU CONC 16 15 14 13 12		8 NTED RATIO RATIO + + + + + + + +	TIMI DN 11 Coori + + + + +	E = AP 1 DINK + + + +	30 IN FI ATES + + 127 295 +	1 O FM ARE + + 75 473 785	/S IN F + + 126 563	1 1 EET + + + 42 +	1 2 ARTI + + + +	CLE5	1 3 + + + + + +	1 4 328 + + + + + +	1 5	1 6	1 7	
ACCU CONC 16 15 14 13 12 11		8 11ED 13 11 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	TINI DN 111 CDORI + + + + + + + + + +	E = AP 1 DIN/ + + + + + + + +	30 IN FI ATES + 127 295 + +	1 0 FM ARE + + 95 473 785 625	(5 1H F + + 126 563 1029	1 1 EET + + 42 + 247	1 2 ARTI + + + + +	CLES + + + + + + +	1 3 + + + + + + + +	1 4 328 + + + + + + +		1 6	17	
ACCU CONC 16 15 14 13 12 11 10		B NTED RATIO + + + + + + + + + + +	TINI DN 11 CDORI + + + + + + + + + +	E = P E = AP DINK + + + + + + + + + +	30 IN FI ATES + + 127 295 + + + +	1 O DAY PM ARE + + 75 473 785 6251 +)	(5 IN F + 126 563 1029	1 FEET + + 42 + 247 743	1 2 ARTIC + + + + + +	 + + + + + + + + + +	1 3 = : + + + + + + + + + + +	1 4 328 + + + + + + + + +	1 5	1 6	1 7	
ACCU CONC 16 15 14 13 12 11 10 9		8 NTED RATIO RATIO + + + + + + + + + + + +	TIMI DN M COORI + + + + + + + + + +	==== ===== ===== ===== ===== ===== ===== ===== ==== ==== ==== ==== ==== ==== ==== == =	30 IN FI ATES + + 127 295 + + + + +	1 O M ARE + + 75 473 785 625 + +	(S IN F + + 126 563 1029 1335 4631	1 1 FEET + + 42 + 247 743 1127	1 2 ARTIC + + + + 1 5	 + + + + + + + + + + + +	1 3 + + + + + + + + + + + + +	1 4 328 + + + + + + + + + +	15	1 6	17	
ACCU CONC 16 15 14 13 12 11 10 3 9		B TED TATIO	TIMI DN M CDORI + + + + + + + + + + + + +	57 E = 1 DINK + + + + + + + + + + + + + +	30 IN FI ATES + + 127 295 + + + + + +	1 O FM ARE + + 75 473 785 6251 +) +) +	(S IN F + 126 563 1029 1335 4631 31	1 P EET + + 42 + 247 743 1127 654	1 2 ARTI + + + + + 1 5 +	CLES + + + + + + + + + + + +	1 3 + + + + + + + + + + + + +	1 4 328 + + + + + + + + + + + +	1 5	1 6	17	
ACCU CONC 16 15 14 13 12 11 10 9 9 7		B TED TED TATJO + + + + + + + + + + + + +	TIMI DN 11 CDORI + + + + + + + + + + + + + +	57 E = 1 DIN + + + + + + + + + + + + + +	30 IN FI ATES + 127 295 + + + + + + +	1 O EAN M ARE + + 95 473 785 625 + + + + +	(S 1) F + + 126 563 1029 1335 4631 31 +	1 P EET + + 247 743 1127 654 41	1 2 ARTI + + + + + 1 5 + 15	CLES + + + + + + + + + + + + + +	1 3 + + + + + + + + + + + + + +	1 4 328 + + + + + + + + + + + + + +	1 5	1 6	17	
ACCU CONC 16 15 14 13 12 11 10 9 7 6		B TED TED TATJE + + + + + + + + + + + + +	TIMI DN 11 CDORI + + + + + + + + + + + + + + + +	57 E = 1 DINK + + + + + + + + + + + + + +	30 IN FI ATES + 127 295 + + + + + + + +	1 O EAN M ARE + + 95 473 785 625 + + + + + +	(5 1) F + + 126 563 1029 1335 4631 31 + +	1 P EET + + 247 743 1127 654 41 +	1 2 ARTI + + + + + 1 5 + 15 +	CLES + + + + + + + + + + + + + +	1 3 + + + + + + + + + + + + + + + + + + +	1 4 328 + + + + + + + + + + + + + +			17	
ACCU CONC 16 15 14 13 12 11 10 9 7 6		B TIED TATJE + + + + + + + + + + + + +	TIMI DN 111 CDORI + + + + + + + + + + + + + + + + +	57 E = 1 D I N/ + + + + + + + + + + + + + +	30 IN FI ATES + 127 295 + + + + + + + + +	1 O DAY PM ARE + + 75 473 785 625 + + + + + + +	(S IN F + + 126 563 1029 1335 4631 31 + +	1 P EET + + + 2 42 + + 247 743 12? 654 41 +	1 2 ARTI + + + + + 1 5 + 15 +	CLES + + + + + + + + + + + + + +	1 3 + + + + + + + + + + + + +	1 4 328 + + + + + + + + + + + + + + +			1 7	
ACCU CONC 16 15 14 13 12 11 10 7 6		B TIED RATIO + + + + + + + + + + + + 8	TIMI DN HI CDORI + + + + + + + + + + + + + 7	57 E = 1 DIN + + + + + + + + + + + 1	30 IN FI ATES + + 127 295 + + + + + + + + 1	1 O DAY M ARE + + 75 473 785 625 + + + + + + + +	(5 IN F + + 126 563 1029 1335 4631 31 + + 1	1 FEET + + 42 + 247 743 112? 654 41 +	1 2 ARTIC + + + + + 15 + 15 + 1	CLES + + + + + + + + + + + + + + 1	1 3 = : + + + + + + + + + + + + + + + + + + +	1 4 328 + + + + + + + + + + + 1			1 7	

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DO A SCREEN FRINT NON OR PRESS (RETURN) TO GO BACK TO THE HEHU. NAY

ACCL Conc	INUL Ent	ATED Ratii	717 1 HQ	te = Iap	9t) 14 Pi	DA1 PH	15	I	PARTI	CLES	:	552
		(C00F	tD I N/	ATES	ARE	[H	FEET				
24	1	+	+	+	+	+	+	*	•	+	+	•
23	1	٠	٠	+	+	٠	+		+	+	٠	•
22	1	٠	+	+	•	+	+	+	+	+	+	+
21	:	+	+	+	•	•		•	+	+	+	+
20	1	+	+	+	+	+	+	+	+	•	+	+
- 19	1	÷	٠	63	•	32	•	•	+	+	٠	+
18	1	+	+	63	42	+	+	•	•	+	•	+
17	1	+	٠	+	42	157	+		+	+	+	+
15	I.	+	+	168	209	94	31	٠	+	+	+	+
15	1	+	+	187	749	280	31	21	+	+	+	•
14	t	+	+	62	705	590	155	83	+	+	+	+
13	Ł	+	+	•	331	836	278	82	+	+	+	٠
12	:	+	+	+	+	2781	1139	184	+	+	+	+
11	:	+	٠	+	- 41	154	1779	388	+	+	+	٠
10	;	+	+	•	+	31	1722	1220	+	+	+	+
9	:	+	+	+	+	+	304	2045	+	+	٠	٠
9	1	+	+	+	+	+	30	948	15	+	+	+
7	1	+	+	+	+	+	+	69	15	+	٠	+
6	:	+	+	+	+	+	+	•	15	+	+	۰
			•					,				
		4	1	0	1	2	3		5	6	7	8

DO A SCREEN PRINT NOW OR PRESS (RETURN) TO GO BACK TO THE HEND. JUL

DHC		NTED	TH	1E 3	50	DA	75	P	ARTI	CLES	* :	3:0	
	ENTI	RATIC) HK 2001	18P 2012	IN PI ATES	PH AFE	 IN 1	FEET					
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2	Ť		+	•	+	+	+	+		+	+	٠	
1	1	+	+	+	+	+	+	+	٠	ŧ	٠	+	
0	:	+	+	+	÷	÷	+	+	٠	+	ŧ	+	
9	1	+	•	+	+	+	÷	÷	+	٠	+	+	
B	ł	+	+	+	+	+	+	+	+	+	٠	f	
7	1	+	٠	+	43	+	32	+	F	۲	+	+	
٨	:		+	129	128	+	+	+	+	+	٠		
s	÷	+		255	299	127			+	۲	٠		
4	1	+	+	+	847	412	32	21	,	٠	+	٠	
13	1	+	+	127	1434	726	94	21		F.	٠	۰	
12	1	+	+	63	992	1382	345	164	+	٠	+	٠	
1	:	÷	+	+	167	532	718	18?	٠	۲	٠	٠	
9	;	+	+	+	+	93	559	251	٠	٠	٠	٠	
	:	+	+	+	•	٠	31	411	31	۴	۲	٠	
8	ł.	+	+	+	+	+	62	161	15	÷	+	+	
7	4	+	+	+	+	+	+	41	٠	٠	۲	+	
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DO A SCREEN PRINT NOW OR PRESS (RETURN) 10 60 PACK TO THE MENU. JUN

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ACC'U CONCI	HUL En1	ULATED TIME = 120 DAYS NIRGTION MAP IN PPN							PART	ICLE	5 2	701
			:001	RDIN	ATES	AKE	[N	EET				
24	:	+	+	- +	٠	+	+	٠	+	٠	+	+
23	:	+	٠	•	•	+	+	+	+	٠	٠	+
22	:	+	+	+	31	+	•	+	+	+	+	+
21	;	+	ŧ	+	•	+	+	+	٠	+	+	+
20	\$	+	+	+	+	+	+	+	+	+	+	+
17	1	٠	+	68	45	- 34	+	+	+	+	٠	+
18	1	+	ŧ	•	91	102	- 34	+	+	+	٠	+
17	:	+	+	136	588	169	67	•	•	+	+	•
16	1	+	+	203	811	405	34	+	+	٠	+	+
15	:	+	+	63	1034	673	134	•	+	+	÷	+
14	:	+	٠	•	12111	141	368	89	ŧ	+	+	+
13	:	+	+	•	4021	605	701	22	+	+	+	+
12	1	+	٠	+	45	9001	297	265	•	+	+	+
11	:	+	٠	•	- 44	5321	459	375	+	+	+	+
10	1	+	+	•	+	+1	322	957	+	÷	٠	+
9	:	٠	٠	•	•	+	5602	2561	ló	+	+	+
9	:	+	+	+	•	•	•	1134	•	+	+	+
7	:	÷	ŧ	+	•	+	+	65	32	+	٠	٠
6	:	• 	•	•	•	•	•	• 	•	+	•	+ •
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		-		e	l	2	2	4	5	5	7	9

SO A SCREEN PRINT NOW OR PRESS "RETURN: TO GO BACK TO THE NERU. AUG

ACEU Conc	NUL) En 11	4/20 RATIO	FIN M M	ie = Iap I	I N E IN PP	n N	15	ļ	PART	LES	Ξ	309
		(:úOF	DINA	125	AFE	IN F	EE1				
24	:	+	+	+	+	. +	+	٠	+	+	+	+
23	:	٠	+	۰.	٠	٠	٠	+	÷	+	٠	+
22	:	÷	+	+	+	+	•	+	+	+	+	+
21	:	٠	۴	+	٠	+	34	+	+	٠	ŧ	+
20	:	+	٠	+	91	34	+	+	+	+	٠	+
19	:	÷	٠	68	91	+	+	+	+	ŧ	+	۲
10	ł	+	٠	204	91	34	34	+	+	+	٠	+
17	1	+	+	68	136	169	•	٠	+	+	٠	٠
16	;	+		- 53	541	236	34	+	+	٠	٠	+
15	:	+	٠	135	1034	303	34	+	ŧ	+	+	÷
14	:	+	+	494	942	738	201	÷.	+	+	٠	٠
13	1	+	+	336	1341	970	167	22	٠	÷	÷	*
12	:	+	۰	134	624	1667	865	68	٠	+	٠	+
- 11	4	+	+	+	133	532	497	132	33	+	+	+
10	ł	+	+	+	•	33	363	132	٠	+	۲	٠
9	;	+	+	+	+	- 33	99	306	16	+	٠	+
8	:	+	÷	+	٠	+	+	153	٠	٠	٠	٠
7	:		÷	+	•	+	+	•	36	+	+	+
6	:	•	+	+	+	+	*	+	•	, 	•	+
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		8	9	ı	1	1	1	I.	1	1	1	1
				ŋ	1	2	3	4	5	6	1	9

DO A SCREEN PRINT NOW OR PRESS (RETURN) TO GO BACK TO THE NEWC. SEP

194

accu Conc	MULI Enti	NTED RATIO	T() M)	1E = 1AP	18: IN PI) (H PH	NYS.		PART.	ICLES	:	290
		(200	201N	ATES	ARE	IN 1	FEET				
24	1	+	•	+	+	•	*	+	•	•	•	••••
23	:	+	÷		+	٠	+	+	+	+	٠	•
22	:	+	+	50	+	+	+	•	+	+	٠	•
21	1	٠	+	74	+	۴	+	+	+	+	+	٠
20	1	+	+	÷	148	37	+		+	٠	٠	+
19	1	+	+	148	147	٠	+	+	٠	+	+	+
18	:	+	+	73	391	145	37	+	٠	+	٠	+
17	1	٠	٠	146	925	146	+	24	+	+	+	٠
15	;	+	÷	291	1965	217	109	+	•	+	÷	٠
15	:	٠	٠	145	1301	541	144	24		٠	٠	٠
14	:	+	٠	216	757	574	323	72	18	•	٠	+
13	:		+	٠	382	\$00	428	- ?1	+	+	٠	٠
12	t	+	+	+	190	356	452	4?	12	٠	+	٠
11	t	÷	+	+	95	- 71	195	199	٠	٠	٠	
19	1	+	+	+	٠	+	141	211	+	+	٠	+
9	1	+	•	•	•		- F	233	+	+	٠	+
8	÷	+	+	+	+	+	•	46	+	+	٠	+
1	:	٠	+	•	+	+	+	+	٠	+	٠	+
6	:	+	•	•	*	+	*	•	+ 	+	*	+
		8	9	1	1	1	ł	1	t	1	1	1
				0	1	- 2	- 3	4	5	6	- 7	8

DO A SCREEN PRINT NOW OR PRESS (RETURN) TO BD BACK TO THE MENU. OCT

ACCUMULATED TIME * 210 DAYS PARTICLES = 187 EDMCENTRATION HAP IN PPH

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			C001	io Ini	ATES	ARE	1N	FEET				
24	;	•	•	•	•	•	•	+	••••	+	•	•••
23	1	٠	+	۴	+	٠	•			+	+	+
22	:	÷	٠	+	+		+	٠	•	+	+	•
21	:	٠	+	٠	50	٠		F	•	•	•	۰
20	:	٠	+	+	258	+	+	+	+	٠	٠	+
19	:	٠	+	149	247	111	+	25	+	٠	٠	+
18	1	+	+		345	184	74	+	+	+	+	٠
17	:	+	٠	-74	343	331	٠	24	٠	٠	٠	+
16	1	٠	+	+	537	549	193	9?	+	+	٠	+
15	1	+	+	+	399	291	169	24	٠	٠	+	٠
14	;	+	+		9?	598	191	48	+	٠	٠	+
13	1	+	+	+	96	108	253	48	•	+	٠	+
12	1	+	+	•	•	72	288	72	+	+	٠	+
11	1	+	+	+	+	36	251	72	+	•	+	+
10	ť	÷	+	٠	+	+	143	190	+	•	٠	•
9	:	+	+	+	+	+	•	142	٠	+	+	
8	ł.	٠	+	+	•	+	+	21	٠	+	+	
1	ł	+	•	+	+	+	+	+	+	+	+	+
6	1	+	+	•	•	+	٠	+	+	٠	٠	٠
			_	•								
		8	9	1	1	1	1	ŧ	ł	1	t	ł
				0	1	2	3	4	S	6	1	9

DO A SCREEN PRINT NOW OR PRESS (RETURN TO GO BACK TO THE MENU. 1989)

ACCUMULATED TIME - 240 DAYS PARTICLES = 145 EDREENTRATION MAE IN PEN LOOFDINATES ARE IN FEET -------14 : • • • • • • • • • • 23 : * * * * * * * * * * * 20 1 + + + 149 37 + 25 + + + + 19 : + + 149 595 297 37 25 - - + + 16 1 + + + 196 220 110 24 + + + + 15 1 + + + 90 256 293 24 + + + + 14 1 + + + 146 73 75 + + + 13 1 + + + + 109 291 48 + + + + 12 ; + + + 72 109 24 + + + + 11 1 + + + + + 108 144 + + . • 10 1 + + + + 36 167 • • •
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DO A SCREEN PRINT NOW OR PRESS (RETUPN) TO GO BACK TO THE HEND, DEC

CONCENTRATION MAP IN FEW												
•			2001	201N	ATES	AFE	111 1	FEET				
14	,	• • • •	•• ••	• • • •	• • • •	••••		•	****	••••	• • • •	
	÷		•		•		•		•			
	1	+			42		75			÷	÷	<u>.</u>
71			+	74	76.7	76		25	•		•	
20	÷	+	•		151	- 9	78	Ξ,	•	٠	•	•
17	÷	+	•	150	50	75	75	25	+	+	,	•
18	÷	+	٠	25	204	399	- 75	25	٠		+	
12	1	۴	+	75	147	299	112	•			+	
15	:	1	•	•	99	112	/1		•	*	٠	+
15		+	+	+	ŧċ.	111	162	74	•	•	•	•
19		•	*	+	+		111		+	+	٠	+
13	:	+	٠	•	+	110	37	147	+	+	+	+
12	:	٠	+	+	+	*	7:	73	+	•	٠	*
11	1	+	٠	+	٠	37	37	73	+	+	٠	+
10	ł	+	٠	+	+	+	٠	121	•	+	٠	+
9	1	+	+	+	٠	+	+	72	18	٠	+	٠
9	:	٠	٠	٠	+	•	F	٠	•	۰	٠	٠
7	1	٠	+	+	+	٠	+	+	٠	•	*	•
6	•	٠	٠	+	+	+	+	٠	+	ŧ	•	•
	-				• • • •				 .			
		ş	?	1	ι	1	1	1	1	1	1	1
		-		9	i	2	3	4	5	b	1	8

675989 ATCH LINE = 376 BARS - PARTICIES = 141

DO A SLHEEN PRINT NON OR FRESS (RETURN TO GO BACH TO THE MEND, JAN

032A 2010	MUL/ Enti	ATED Ratii	TII Dr n	1E = 1AP :	390 In Pf	e Di M	lrS		Phri	ICLES	j 3	97
			C901	RD1N	ATES	AFE	111 F	EET				
24	1	+	+	•	*	+	•	•	+	*	••••	•
23	1	٠	+	+	+	13			+			+
22	1	٠	+	51	136	51	25	17	+	•	٠	•
21	1	+	+	228	202	76	38	•	۴	٠	٠	
20	1	+	+	151	303	76	114	+	+	+	+	٠
19	t	+	ŧ	•	151	75	38	+	+	+	+	+
18	1	+	•	+	150	199	75	25	+	+	+	+
17	1	+	+	+	50	187	112	50	+	+	+	+
16	ł	+	•	+	50	18?	159	25	٠	•	+	•
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11	1	+	+	+	+	+	+	-74	٠	٠	+	•
10	1	+	+	+	+	+	+	24	+	+	+	+
9	:	+		+	+	+	٠	49	+	+	+	•
B	;	+	+	+	+	+	+	- 24	+	+	٠	+
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- 23	1	+	+	+	176	110	63	19	٠	+	•	٠
22	:	+	+	+	103	101	103	34	+	•	٠	+
21	:	+	٠	71	103	17	115	Ze	•	•	٠	+
20	:	+	+	+	+	+	192	26	19		+	٠
19	1	+	٠	+	51	76	76	51	+	+	٠	+
10	1	+	+	+	51	76	76	25	+	•	•	+
17	1	٠	+	٠	•	٠	76	25	٠	٠	٠	٠
16	:	+	+	+	+	76	76	50	+	+	+	٠
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DO A SCREEN PRINT NOW OR PRESS SPETURN. TO GO PAEK TO THE MEMU. FEB

OD & SCREEN PRINT NOW OR PRESS (RETURN) TO GO BACK TO THE RENU. APA

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ACCUMULATED TIME = 339 DAYS PARTICLES = 101 CONCENTRATION MAP IN PPH

COORDINATES ARE IN FEET

24	÷.	+	+	•	•	+	+	•	+	+	+	+
23	:	+	٠	+	69	39	•	9	٠	+	+	٠
22	1	+	+	+	204	127	51	+	+	+	+	٠
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20	1	+	+	+	59	114	- 39	25	+	+	٠	+
19	1	+	٠	+	151	115	20	25	+	٠	•	+
18	:	٠	+	•	50	263	301	25	+	٠	٠	٠
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DO A SCREEN PRINT NOW OR PRESS (RETURN. TO GO BACK TO THE MENU. MAP

	ACT. VS (EMPTY C)	5. SIM. CI DLUMN SPAC	CONC. IN ES REPRES	N ppm WITH SENT NO DA	RANDOM-W	JALK* (TION)
200.	ACT	MAY CIM		JUNE		JULY
мыт	HC1.	51M.	ACT.	SIM.	ACT.	SIM.
Мыд	1144	1124	391	411	2640	2078
MUS	028	560	660	344	1040	1160
MULL	14	41	64	127	100**	150
- MI47	•	0	<50	15	17	21
MUC	(TA	0	30	128	152	200
	34 200	0	<50	0	<36	0
	200	126	373	255		500
MW17	< 34	0	0	0		65
MWDO	U	0	27	21		32
riw20		94		411		602
LOC.		AUG.		SEP.		OCT.
	ACT.	SIM.	ACT.	SIM.	ACT.	SIM
MWJ	2764	2370		306	201	227
MW4	1254	1201		865	459	
MWS	225**	410		275	144	105
MWG	<36	83		16	100	201
MW7	123	188		200	204	23
MWR	16	100		200	204	263
MULLA	10	540		700	101	475
MU17		250		700	6V6	4/3
MILLE		124		68	99	144
		124		34	68	140
MW20		1036		/38	626	224
LOC.	A.C.T.	NOV.	60T	DEC.		JAN.
NAL 172	AUT.	514.	ACI.	SIM.	ACI.	SIM.
MWS	100**	140		/1		72
mw4	400**	284	200**	108	200**	72
MMD	150	95		36		36
MW6	< 34	24		24		73
MW7	150	250	132	195	41	98
MW8		<u> </u>				~
MW16		•		0	-	U
	779	96		0 97	683	109
MW1/	779 122	96 148	34	0 97 149	683	109 149
MW17 MW19	779 122 61	96 148 108	34 21	0 97 149 125	683	109 149 100
MW17 MW19 MW20	779 122 61 527	96 148 108 503	34 21 405	97 149 125 145	683 405	109 149 100 109
MW17 MW19 MW20 LDC.	779 122 61 527	96 148 108 503 FEB.	34 21 405	0 97 149 125 145 MAR.	683 405	109 149 100 109 APR.
MW17 MW19 MW20 LOC.	779 122 61 527 ACT.	96 148 108 503 FEB. SIM.	34 21 405 ACT.	0 97 149 125 145 MAR. SIM.	683 405 ACT.	109 149 100 109 APR. SIM.
MW17 MW19 MW20 LOC.	779 122 61 527 ACT. 14	96 148 108 503 FEB. SIM. 48	34 21 405 ACT. 40	0 97 149 125 145 MAR. SIM. 49	683 405 ACT. 4	109 149 100 109 APR. SIM. 24
MW17 MW19 MW20 LOC. MW3 MW4	779 122 61 527 ACT. 14 21	96 148 108 503 FEB. SIM. 48 73	34 21 405 ACT. 40 29	0 97 149 125 145 MAR. SIM. 49 70	683 405 ACT. 4 8.2	0 109 149 100 109 APR. SIM. 24 25
MW17 MW19 MW20 LOC. MW3 MW4 MW5	779 122 61 527 ACT. 14 21 21	96 148 108 503 FEB. SIM. 48 73 24	34 21 405 ACT. 40 29	0 97 149 125 145 MAR. SIM. 49 70 0	683 405 ACT. 4 8.2 110	0 109 149 100 109 APR. SIM. 24 25 0
MW17 MW19 MW20 LOC. MW3 MW4 MW5 MW5	779 122 61 527 ACT. 14 21 21	96 148 108 503 FEB. SIM. 48 73 24 49	34 21 405 ACT. 40 29	0 97 149 125 145 MAR. SIM. 49 70 0 25	683 405 ACT. 4 8.2 110 14	0 109 149 100 109 APR. SIM. 24 25 0 50
MW17 MW19 MW20 LOC. MW3 MW3 MW4 MW5 MW6 MW7	779 122 61 527 ACT. 14 21 21 81	96 148 108 503 FEB. SIM. 48 73 24 49 50	34 21 405 ACT. 40 29 77	0 97 149 125 145 MAR. SIM. 49 70 0 25 37	683 405 ACT. 4 8.2 110 14 51	0 109 149 100 109 APR. SIM. 24 25 0 50 50
MW17 MW19 MW20 LOC. MW3 MW3 MW4 MW5 MW5 MW6 MW7 MW8	779 122 61 527 ACT. 14 21 21 81 27 (?)	96 148 108 503 FEB. SIM. 48 73 24 49 50 0	34 21 405 ACT. 40 29 77 30	0 97 149 125 145 MAR. SIM. 49 70 0 25 37 20(?)	683 405 ACT. 4 8.2 110 14 51 11.4	0 109 149 100 109 APR. SIM. 24 25 0 50 50 30(?)
MW17 MW19 MW20 LOC. MW3 MW4 MW5 MW5 MW5 MW5 MW5 MW5 MW5 MW5 MW6 MW7 MW8 MW16	779 122 61 527 ACT. 14 21 21 81 27(?) 527	96 148 108 503 FEB. SIM. 48 73 24 49 50 0 49	34 21 405 ACT. 40 29 77 30 398	0 97 149 125 145 MAR. SIM. 49 70 0 25 37 20(?) 24	683 405 ACT. 4 8.2 110 14 51 11.4 248	0 109 149 100 109 APR. SIM. 24 25 0 50 50 30(?) 37
MW17 MW19 MW20 LOC. MW3 MW3 MW4 MW5 MW5 MW5 MW6 MW7 MW8 MW16 MW17	779 122 61 527 ACT. 14 21 21 81 27(?) 527 91	96 148 108 503 FEB. SIM. 48 73 24 49 50 0 49 150	34 21 405 ACT. 40 29 77 30 398 91	0 97 149 125 145 MAR. 5IM. 49 70 0 25 37 20(?) 24 50	683 405 ACT. 4 8.2 110 14 51 11.4 248 51	0 109 149 100 109 APR. SIM. 24 25 0 50 50 30(?) 37 51
MW17 MW19 MW20 LOC. MW3 MW4 MW5 MW4 MW5 MW6 MW7 MW8 MW16 MW17 MW19	779 122 61 527 ACT. 14 21 21 81 27(?) 527 91 34	96 148 108 503 FEB. SIM. 48 73 24 49 50 0 49 50 0 49 150 37	34 21 405 ACT. 40 29 77 30 398 91 29	0 97 149 125 145 MAR. SIM. 49 70 0 25 37 20(?) 24 50 75	683 405 ACT. 4 8.2 110 14 51 11.4 248 51 29	0 109 149 100 109 APR. SIM. 24 25 0 50 50 50 30(?) 37 51 38

*ALL VALUES USED IN ACT. COLUMNS ARE ESTIMATED BY SUBTRACTING 20 ppm FROM CONC. AS BACKGROUND AND ANY RESIDUAL C1 CONC. **ESTIMATES OF C1 CONC. IN SHALLOW NESTED WELLS FROM DEEP NESTED WELL CONC.

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APPENDIX W

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Cl LOAD METHODOLOGY AND CONCENTRATIONS

METHODOLOGY

To calculate Cl loading from the reserve pit, I separated each well site with nine discrete blocks of aquifer as shown in Exhibit 12. I used data from well nests MW 4 and MW 5 to stratigraphically separate layering of Cl concentrations. Well MW 6 and MW 19s were assumed to be located at the same place for deep and shallow Cl concentration data.

I calculated loading as follows:

- 1. (AREA OF EACH ZONE) x (.25 POROSITY) = AREA OF WATER (ft^2)
- 2. (AREA OF WATER) x (MONTHLY VELOCITY) = MONTHLY WATER VOLUME PER ZONE (ft³)
- 3. CONVERT ft³ INTO LITERS
- 4. (MONTHLY WATER x (MONTHLY Cl mg/l) = MONTHLY VOLUME PER ZONE) Cl LOAD (mg Cl)
- 5. (ADD EACH BLOCK OF C1 LOAD PER MONTH) = TOTAL MONTHLY C1 LOAD (mg C1)
- 6. (ADD EACH TOTAL MONTHLY C1 LOAD) = TOTAL YEARLY C1 LOAD (mg CL)

7. SUBTRACT BACKGROUND C1 LEVELS OF 20 mg/1 FROM TOTAL

8. CONVERT mg OF C1 LOAD PER YEAR INTO kg AND 1bs OF C1

200



CROSS SECTION DATA					
BLOCK SITI	<u>E WELL</u>	DATA CHE	MISTRY_	BLOCK AREA (ft^2)	
A1		MW	55	200	
A2		MW	4s	380	
A2		MW	19s	280	
B1		MW	5	200	
B2		MW	4 m	380	
B 3		MW	6	280	
Cl		MW	5/BACKGROUNI	D 550	
C2		MW	4d	1045	
C3		MW	6/BACKGROUN	D 770	
	GROUND WA	TER VELO	CITY DATA		
	MONTH	VEI	LOCITY (ft/da	av)	
	The second second second second second second second second second second second second second second second s				
	JUL		5.6		
	AUG		5.3		
	SEP		4.7		
	OCT		4.5		
	NOV		3.2		
	DEC		2.9		
	JAN		2.9		
	FEB		2.8		
	MAR		2.6		
	APR		2.4		
	MAY		5.6		
	JUN		5.9		
C1 CONCEN	TRATIONS (* = ESTI	MATED)		
			· .		
BLOCK		AUG	SEP	በርሞ	
DROOM	<u>~~~</u>	<u>404</u>		<u>M.M.A.</u>	
A1	225*	300*	250*	186	
A2	1060	1270	1000*	678	
A3	120*	100*	95*	88	

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A3	120*	100*	95*	88
B1	30	52	45*	40*
B2	650	895	400*	88
B3	37	30	28*	24*
C1	20	30	28*	25*
C2	106	36	36*	35*
C3	30	25	20*	20*

CONT. C1 CONCENTRATIONS

BLOCK	NOV	DEC	JAN	FEB
Al	170	150*	140*	120*
A2	400*	300*	200*	100*
A3	81	41	40*	54*
B1	34	30*	35*	41
B2 🛸	61	47	41	41
B3	22	20*	21*	21*
C1	24	22	20*	20*
C2	34	30*	30*	30*
С3	20	20*	20*	20*
<u>BLOCK</u>	MAR	APR	MAY	JUN
A1	100*	130	34	84
A2	100*	80*	658	890
A3	54	50	47	56
B1	50	60	50*	45
B2	49	28	30*	40*
B3	25*	34*	30*	40*
C1	22*	25	25*	30*
C2	30*	25	25*	35*
C3	21*	22*	25*	28*

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