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PALEOMAGNETIC INVESTIGATION OF THE LOWER CRETACEOUS KOOTENAI FORMATION, WESTERN MONTANA

By

Kelly M. Brunt

B.S., Syracuse University, 1993

Presented in partial fulfillment of the requirements for the degree of

Master of Science

University of Montana

1997

Approved by Chairman, Board of Examin

Dean Graduate School

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PALEOMAGNETIC INVESTIGATION OF THE LOWER CRETACEOUS KOOTENAI FORMATION, WESTERN MONTANA

Director: Steven D. Sheriff

The Lower Cretaceous (Aptian) Kootenai Formation was sampled in two localities of western Montana for a paleomagnetic investigation of a possible rotation of the Purcell Anticlinorium during Laramide deformation in Montana. Eight sites from the Ford Creek locality and three sites from the Marias Pass locality were sampled and subjected to standard paleomagnetic techniques in an attempt to test the rotational model. One site from the Marias Pass locality was rejected as it had a site-mean direction that is coincident with the secondary, present-day, direction of the other two. The two remaining sites of the Marias Pass locality yielded a structurally corrected, characteristic, direction that is seemingly rotated counterclockwise 45±16°, with no latitudinal translation. However, based on only two reliable sites, the statistics from this locality are not very robust. Thus results from the Marias Pass locality were not used to test the extent of the rotational model. Six of the eight sites sampled in the Ford Creek locality passed statistically based arbitrary rejection criteria and were compared with Aptian (108-115 Ma) reference poles in the literature. The sites from this locality have been rotated clockwise 23±9°, with little latitudinal translation. While the results of this locality support previously devised rotational models and are statistically coincident with paleomagnetic results from the Purcell Anticlinorium, it can not be conclusively determined that the Purcell Anticlinorium rotated as a coherent unit clockwise through $23\pm9^{\circ}$. The extent of the rotation can not be determined based on one reliable locality. While a regional rotation is one possible explanation, another is that the locality has experienced a small scale vertical axis rotation.

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Introduction

Paleomagnetism is the method used to determine rotations and translations of tectonic plates and terrains (Grubbs and Van der Voo, 1976; Beck, 1980; Schwartz and Van der Voo, 1984; Irving et al., 1986; Eldredge and Van der Voo, 1988; Beck, 1991; Irving and Wynne, 1991; Jolly and Sheriff, 1992; Symons and Timmons, 1992; Wynne et al., 1992; Link et al., 1993; Irving et al., 1996). While techniques prove easier when applying paleomagnetism to igneous rocks, rotational and translational models can also be tested using sedimentary units. Paleomagnetic field and laboratory techniques determine the declination and inclination of the geomagnetic field at the rock's formation. From the direction of the field at that point in time one can determine an instantaneous estimate of a dipole axis known as a virtual geomagnetic pole (VGP). Calculating inclinations and declinations from the measured virtual geomagnetic pole and then comparing that direction from rocks in an allocthonous area with those expected from the average paleopole for a continent yields a measurement of relative rotation and translation. Anomalous inclinations may indicate latitudinal translation. Similarly, differences in declinations may indicate that the terrain has rotated around a nearby vertical axis. This study uses paleomagnetism of the sandstones of the Lower Cretaceous (Aptian) Kootenai Formation (Suttner, 1969; DeCelles, 1986), of western Montana to examine a possible rotation of the Purcell Anticlinorium (Symons and Timmons, 1992; Sears, 1994; Cook and Van der Velden, 1995; Van der Velden and Cook, 1996) within the Rocky Mountain fold and thrust belt.

Many workers have been interested in, and proposed models for, the orogenic development of the Rocky Mountain fold and thrust belt (Smith, 1965; Hoffman et al., 1976; Monger and Price, 1979; Woodward, 1981; Price and Carmichael, 1986; Symons and Timmons, 1992; Link et al., 1993; Sears, 1994; Colpron and Price, 1995; Cook and Van der Velden, 1995; Van der Velden and Cook, 1996). Symons and Timmons (1992), following proprietary work of the Amoco Production Company, proposed a rotational model for the Purcell Anticlinorium (Figure 1). They completed a paleomagnetic study of layered argillites and the Moyie sills of the Middle Proterozoic Prichard Formation within the anticlinorium and compared their declinations to those calculated from the Apparent Polar Wander Path for the Middle Proterozoic (Irving, 1979). Symons and Timmons (1992) concluded that the tectonic history of the Purcell Anticlinorium included, among other structural complexities, a clockwise rotation of 37±12° with little or no latitudinal translation. They proposed that this deviation was caused as the anticlinorium was thrust in a northeasterly direction onto the craton during Laramide deformation at about 72 to 56 Ma (Hoffman et al., 1976). Theirs being the first published paleomagnetic study of a possible rotation of the Purcell Anticlinorium, Symons and Timmons (1992) cited many structural reasons for this apparent rotation. Included in their reasoning was the recognition of a regional north northwest structural trend which, in closer detail, changes to a northern trend within the anticlinorium. The detailed structural mapping of Mudge et al. (1982) also indicates this in measured strikes and dips in the vicinity of the anticlinorium. Thus changes in the trend of regional structures



Figure 1. Regional relationship between the two localities of this study area (Ford Creek, FC, and Marias Pass, MP) and four others. Irving et al. (1986) measured a clockwise rotation of $24\pm10^\circ$. Eldredge and Van der Voo (1988) measured clockwise rotations (EV) of 35° in the Helena Salient, 23° in the Southwest Montana Traverse Zone, 54° in the McCarthy Mountain Salient, and a counterclockwise rotation (-EV) of 30° in the Helena Salient. Jolly and Sheriff (1992) measured a counterclockwise rotation of $25.2\pm14^\circ$. Symons and Timmons (1992) measured a clockwise rotation of $37\pm12^\circ$.

seems to indicate the possibility of a rotation similar to the observations of Symons and Timmons (1992).

Link *et al.* (1993), in a review of Proterozoic sedimentary strata of the western Cordillera, examined the Middle Proterozoic Belt Supergroup (Höy, 1989; Ross *et al.*, 1991; Ross *et al.*, 1992; Chamberlain and Doughty, 1993) and compared its paleomagnetic data to those of other synchronological Proterozoic sedimentary units. Link *et al.* (1983) provide a contrary view to that of Symons and Timmons (1992). Because the paleomagnetic pole positions of the Belt units and other Proterozoic units were coincident, Link *et al.* (1993) concluded that the Belt units had not been significantly rotated. The paleopoles of the Belt units were obtained from many different thrust sheets and yield a regional result. While the data indicated no regional rotation of the Belt units, Elston and Bressler (1980) measured declination deviations near thrust sheet margins and concluded that these were only localized vertical axis rotations. Thus the older data reviewed by Link *et al.* (1983) are in conflict with the more recent determination of Symons and Timmons (1992).

More recently Sears (1994) elaborated on rotation models for the development of the Purcell Anticlinorium and the tectonic history of the structure by compiling the results of various structural and geophysical studies. He proposed that the Lewis and Clark line (Smith, 1965; Woodward, 1981), southwest of the anticlinorium, sits on top of a southwest-facing thrust ramp that helped form the anticlinorium as a fault-bend fold. Sears (1994) further implied that the ramp formed at the base of the marginal facies of the system. These marginal facies are now found on the northeast edge of the structure. Thus the Purcell Anticlinorium was thrust over the ramp. During thrusting, it rotated clockwise $37\pm12^{\circ}$ based on the measurements of Symons and Timmons (1992).

Although much of the strata of the Purcell Anticlinorium is Precambrian, the structure formed during Laramide deformation and units as young as Paleocene are carried on the thrust (Mudge et al., 1982). Several scientists have completed paleomagnetic investigations of Cretaceous aged sedimentary units in the Disturbed Belt of Montana and Alberta, Canada. Irving et al. (1986) studied rotations of the Early Cretaceous Crowsnest Formation in Alberta, Canada (Figure 1). Along with a northward translation of 17±6°, they found a clockwise rotation of 24±10° from the expected mid-Cretaceous paleomagnetic direction. Eldredge and Van der Voo (1988) sampled the Lower Cretaceous Kootenai Formation and found clockwise and counterclockwise rotations within salients of western Montana and eastern Idaho (Figure 1). They measured maximum clockwise rotations of 54° in the McCarthy Mountain Salient, 23° in the southwest Montana Traverse Zone, and 35° in the Helena Salient. Eldredge and Van der Voo's (1988) rotations were calculated from a reference pole averaged from two localities they sampled on the stable foreland. Thus Eldredge and Van der Voo (1988) and Irving et al. (1986) established the paleomagnetic reliability of Cretaceous-aged sedimentary units in the northern Rocky Mountains and found some rotations coincident with those of Symons and Timmons (1992).

Despite the reliability of sedimentary units and the generally clockwise rotations determined by Eldredge and Van der Voo (1988) and Irving *et al.* (1986), a significant number of counterclockwise rotations within the nose of the Helena salient are also

presented by Eldredge and Van der Voo (1988). These include a maximum of 30° of counterclockwise rotation in the Helena Salient. The authors interpret the varying degree and direction of the rotations as noncoherency of the thrust plate due to buttressing from the underlying craton. Northwest of the Helena Salient, and thus northwest of the work of Eldredge and Van der Voo (1988), Jolly and Sheriff (1992) also found counterclockwise rotations of 25.2±14° within the Late Cretaceous Two Medicine Formation near Wolf Creek, Montana (Figure 1). Their results are relative to the Late Cretaceous paleomagnetic reference pole of Gunderson and Sheriff (1991) from the Adel Mountain Volcanics of west central Montana.

Further afield, other studies have also focused on rotations in thrusted units and investigated the interaction between thrust sheets and their underlying rocks. Grubbs and Van der Voo (1976) completed a paleomagnetic study of Triassic units within the Prospect Thrust Sheet of the Idaho-Wyoming Overthrust Belt. Schwartz and Van der Voo (1984) followed up this investigation and added paleomagnetic data from the Lower Cretaceous. They found that the rotations of the Prospect thrust sheet were neither translated nor transmitted to the older, western, Darby and Absaroka thrust sheets. Irving and Wynne (1991) investigated rotations within the Lower Jurassic and Triassic units of the Wrangellia, Stikinia, and Quesnellia terrains of the Canadian Cordillera in British Columbia. Each of these studies reported considerable variance of the paleomagnetically determined rotations.

Given mixed clockwise and counterclockwise rotations (Grubbs and Van der Voo, 1976; Schwartz and Van der Voo, 1984; Eldredge and Van der Voo, 1988; Irving and Wynne, 1991; Jolly and Sheriff, 1991; Symons and Timmons, 1992) it becomes clear that near the margin of the thrust sheet, regional rotations may be masked by small scale vertical axis rotations. To test models of regional rotation requires a regional distribution of paleomagnetic data. This is particularly true when one considers the rotational models of Symons and Timmons (1992) and Sears (1994). To further test and substantiate the observations of Symons and Timmons (1992), with respect to local versus regional rotations, I conducted a paleomagnetic investigation of Cretaceous aged units, thrust with the Purcell Anticlinorium, at two localities approximately 100 km apart. The two localities, named Ford Creek (Figures 2a) and Marias Pass (Figure 2b), are within the Aptian (108-115 Ma) Lower Cretaceous Kootenai Formation (Figure 3) (Suttner, 1969; DeCelles, 1986). Ford Creek is 100 km north of the Helena Salient and the Lewis and Clark Line. The Marias Pass locality, just south of Glacier National Park, was chosen 100 km further north to investigate the possible coherency of the rotation.



(Mudge, et al., 1982)

Figure 2a. Geologic map of the Ford Creek locality.









Figure 3. Stratigraphy of the Lower Cretaceous Kootenai Formation in western Montana (after DeCelles, 1986 and Eldredge and Van der Voo, 1988).

Field Methods

I chose two localities at which to sample the Kootenai Formation, in order to test the regional model of rotation, in the Montana Fold and Thrust Belt. The equigranular, well-sorted, nonmarine, oxidized sandstones of the lower siliceous member of the Kootenai outcrop with good exposure in both localities. The Ford Creek locality (Figure 2a) was chosen from the extensive mapping of Mudge *et al.* (1982) and Dolberg (1986) because of its exposure, a fairly tight fold, and close proximity to the road and water. The second locality, near Marias Pass on the southern tip of Glacier National Park (Figure 2b), was chosen because of the relatively undisturbed nature (Whipple, 1992) of the Middle Proterozoic through Cretaceous section.

At the Ford Creek locality I used standard paleomagnetic techniques (Butler, 1992) in collecting seven to ten cores from each of eight sites. In the northern, Marias Pass, locality I collected six to eight cores from each of three sites. At each of these sites the samples were distributed in less than 2.5 m of stratigraphic section. Ideally three samples were collected at the same stratigraphic level to test the paleomagnetic consistency of what is presumed to be contemporaneous material. In actuality, due to fracturing within the outcrop and availability, only 28 of the 80 samples were averaged into 16 specimens because they were believed to be penecontemporaneous. These samples were averaged, giving unit weight to the resultant specimens. The minimum stratigraphic separation of sites at a locality was approximately a half a meter while the maximum was approximately 18 m.

Coring was accomplished using a chainsaw, altered by Pomeroy Industries Unlimited, to spin a 2.5 cm diameter diamond bit. Samples were oriented in the field using a magnetic compass and, when possible, a sun compass. I was unable to obtain suncompass orientation for 25 of the 80 samples. However, for those samples with both orientations, the average angular difference between the two measurements was close to zero indicating that the samples were not altered by lightning strikes, the magnetic intensities of the samples were not high enough to affect the accuracy of the magnetic compass, and that there are not any localized anomalies associated with the magnetic field. Thus the samples lacking suncompass orientations are as reliably oriented as those with suncompass orientations.

Laboratory Techniques

Once back in the lab I cut the 80 samples collected in the field into 91, 2.5 cm specimens. Then I began preliminary, pilot demagnetization experimentation on representative specimens from the Ford Creek locality. These preliminary specimens were demagnetized using both alternating field (AF) and thermal (Th) techniques and measured using a Schonstedt SSM-2A Spinner Magnetometer at the University of Montana. For five of the specimens, a progressive alternating field demagnetization sequence, using a Molspin two-axis tumbling demagnetizer with a maximum induction of 100 mT, did not remove more than 50% of the magnetization. Therefore I employed thermal demagnetization on the remaining 13 specimens, using a homemade oven and a cooling chamber with a maximum induction of less than 6 nT, to finish the cleaning steps. However, due to the low intensities of the red beds, the Schonstedt SSM-2A Spinner Magnetometer was not sensitive enough to accurately measure the magnetization of the specimens in a reasonable amount of time. To circumvent this problem, I made arrangements to complete lab work at the University of New Mexico in Dr. John Geissman's lab. His magnetically shielded laboratory included a 2-G Enterprises 706R three axis, superconducting, cryogenic magnetometer, with the capability of accurate direction measurement below 1.0×10^{-5} A/m. I readily and accurately measured magnetizations as low as 5.0×10^{-5} A/m in Geissman's lab.

At the University of New Mexico, alternating field and thermal demagnetization were applied to a pilot specimen from each site from both localities. Alternating field

cleansing was accomplished using a 2-G Enterprises demagnetizer. Thermal demagnetization was accomplished in two thermally separate chambers of a Schonstedt TSD-1 furnace. Six previously untreated specimens from the Ford Creek locality were subjected to alternating field demagnetization and again an induction of up to 100 mT was unable to remove more than about 50% of the magnetization. However, three pilot specimens from the Marias Pass locality reacted well to this form of demagnetization. The Natural Remanent Magnetization (NRM) intensities of the three Marias Pass sites were on the order of 1.0×10^{-2} A/m. Each specimen from the Marias Pass locality underwent between 12 and 25 cleaning steps with an induction of 2 to 120 mT. I chose demagnetization steps which eliminated 5 to 10% of the NRM at each step. The brick red specimens from the Ford Creek locality, with NRM intensities between 1.0×10^{-3} and 1.0×10^{-4} A/m and an unwillingness to react well to the alternating field technique, required thermal demagnetization. Cleaning temperatures ranging from 100° C to 685° C were applied in 8 to 15 steps. The heating time, or cleaning, of the specimens was increased with each sequential step. Specimens were heated for one minute per every ten degrees.

Interpretation of the demagnetization data was simplified using orthogonal projections as discussed by Zijderveld (1967). Statistics were calculated according to the methods of Fisher (1953). As with most paleomagnetic studies, I applied some arbitrary rejection criteria to specimens and sites. At the site level, a maximum of 20° was chosen for the radius (α_{95}) of the 95% confidence cone. This eliminated sites FC1 and FC9 at the Ford Creek locality. At the sample level, specimens giving a direction of more than two

standard deviations away from the site mean were eliminated. This removed a total of seven specimens from the remaining nine sites. After application of these arbitrary rejection criteria, nine sites, with a total of 71 specimens, remained within the data set. Twenty-eight specimens were averaged because they were either obtained from the same core, or believed to be penecontemporaneous. Forty-one specimens from the Ford Creek locality and 14 from the Marias Pass locality remained for a total of 55 specimens for statistical analysis.

Paleomagnetic Results

The Ford Creek specimens demagnetized via thermal techniques to less than 15% of their NRM while the Marias Pass specimens decayed, via alternating field techniques, to less than 10% of their NRM. At these levels the signal to noise ratio had become too high to recover additional accurate directional information. Figures 4a and 4b present representative demagnetizations of specimens from the Ford Creek and Marias Pass localities.

To determine the age of acquisition of magnetization, the specimens from sites 3 and 4 of the Ford Creek locality were subjected to a fold test (Figure 5) as discussed by Butler (1992). The in-situ directions have an α_{95} of 12.7° with a precision parameter, k, of 11.6. After structural correction, α_{95} decreases to 6.5° and the precision parameter increases to 41.7. The expected F distribution value for N=13 and F_{24,24} is 2.0. The calculated k_a/k_b is 3.6. This is well above the expected F distribution and the two sites pass the statistical test. Thus the sites successfully passed the fold test and the magnetization predates deformation. By inference, and similar demagnetization behavior, the same is true for the remainder of the sites sampled at Ford Creek.

It is necessary to adequately average paleosecular variation in any structural application of paleomagnetic directions (Cox, 1970; McFadden *et al.*, 1988; Merrill and McFadden, 1988). I used Cretaceous aged sedimentary units and sampled the units with the intent of sampling a substantial amount of geologic time. Within a site, it is optimum to have low dispersion of directions; between sights, tight clustering is not expected if



Figure 4a. Structurally corrected orthogonal projection (Zijderveld, 1967) of a representative specimen from the Ford Creek locality. Ford Creek demagnetization accomplished via thermal techniques.



Figure 4b. Structurally corrected orthogonal projection (Zijderveld, 1967) of a representative specimen from the Marias Pass locality. Marias Pass demagnetization accomplished via alternating field techniques.



Figure 5. Equal area projection of in-situ (left, α 95=12.7°, k=11.6) Ford Creek 3 directions (\bullet) and Ford Creek 4 directions (*) with respect to their structurally corrected (right, $\alpha 95=6.5^{\circ}$, k=41.7) directions.

90°

0°

sites are spread to sufficiently sample the longest period of paleosecular variation. The amount of expected dispersion varies as a function of paleolatitude (Cox, 1970; McFadden, *et al.*, 1988). Although McFadden *et al.* (1988) revise latitudinal estimates of Cox (1970), the values of Cox (1970) may be more appropriate for this study. This is in part due to the Cretaceous Normal (85-119 Ma), a geologically long amount of time during which the paleomagnetic field remained almost continuously in a normal state. McFadden *et al.* (1988) concentrate on the last 5 Ma while the rocks of this study are Lower Cretaceous. The expected δ_{63} for latitudes of 47° N and 48° N is approximately 17° (Cox, 1970). The δ_{63} values for the Ford Creek and Marias Pass localities are both within 5° of this value (Table 1). This increases one's confidence that the study has adequately averaged paleosecular variation.

One expects δ_{63} values to be low within a given site and increase when calculated between sites. My site-level δ_{63} values (Table 2) are below the expected δ_{63} for the paleofield (Cox, 1970). Yet they are somewhat higher than are typically found in a paleomagnetic study of volcanic rocks. Thus it is probable that paleosecular variation has been partially averaged at the site level. This further increases one's confidence that the study has adequately averaged paleosecular variation.

Another way to determine whether or not a study has adequately averaged paleosecular variation is to document the existence of polarity reversals within the sampled rocks. If reversals occur within a locality or between localities, then these samples must be separated in time by at least the maximum period of paleosecular variation. Unfortunately, there were no reversals recorded at the Ford Creek and Marias

Table 1.	Structurally Corrected locality mean directions for the
	Ford Creek and Marias Pass localities.

Locality	Dec	Inc	VGP lat	VGP long	N/N _c	α95	k	δ ₆₃
Ford Creek	-9.2	62.5	78.3	32.1	6/8	3.7	37.2	13.3
Marias Pass	-39.5	60.2	45.6	245.4	3/3	11.3	13.3	21.6

Dec indicates locality mean declinations, clockwise is positive. Inc indicates locality mean inclinations. VGP lat, VGP long indicate latitude and longitudes of the VGP, reported north and east. N/N_c is the number of sites used in statistical analysis versus number sampled. α_{95} =semiangle of 95% confidence, k=Fisher (1953) precision parameter, δ_{63} =angular standard deviation.

•

site	latitude	longitude	dec	inc	N/Nc	VGP lat	VGP long	α95	k	δ ₆₃
FC3	47.381 N	247.318 E	-24.4	69.6	6/9	73.4	187.4	10.2	44.6	11.1
FC4	47.381 N	247.318 E	-7.1	55.1	7/12	77.1	93.9	5.2	136.7	6.4
FC5	47.381 N	247.318 E	0.1	55.9	8 /15	79 .0	66.9	8.9	39.4	12.1
FC6	47.381 N	247.318 E	6.0	70.0	7/ 9	82.4	275.2	5.1	141.7	6.3
FC7	47.381 N	247.318 E	-6.7	64.3	8/8	85.2	143.8	8.9	39.9	12.0
FC8	47.381 N	247.318 E	-23.0	60.4	5/7	72.6	145.7	16.1	23.5	15.0
MP1	48.411 N	246.767 E	-77.2	67.0	5/6	41.7	189.1	12.5	38.4	11.7
MP2	48.411 N	246.767 E	-78.9	66.7	4/6	40.5	189.3	8.9	107.4	6.8
MP3	48.411 N	246.767 E	37.7	46.9	5/5	54.7	357.9	16.0	23.9	14.9

Table 2. Structurally Corrected site mean directions for 9 sites.

FC sites are from the Ford Creek locality. MP represents Marias Pass sites. Dec indicates site mean declinations, clockwise is positive. Inc indicates site mean inclinations. N/N_c is the number of specimens used in statistical analysis versus number sampled. VGP lat, VGP long indicate VGP latitude and VGP longitudes, reported north and east. α_{95} =semiangle of 95% confidence, k=Fisher (1953) precision parameter, δ_{63} =angular standard deviation. Pass localities. Because of the age of the Kootenai (Aptian, 108-115 Ma), the lack of reversals is not surprising as it is within the Cretaceous Normal (Butler, 1992).

In using sedimentary units for a paleomagnetic rotational analysis, it is important to determine whether the magnetization is chemical or detrital in nature to best estimate its age. Suttner (1969) and DeCelles (1986) limit the deposition and age of the Kootenai to Aptian (108 to 115 Ma). Eldredge and Van der Voo (1984) point out that there is a short amount of time between the deposition in the Early Cretaceous of the Kootenai Formation and the thrusting events of the Late Cretaceous Laramide Orogeny in Montana. Montana thrusting has been limited by Hoffman *et al.* (1976) to be within 56-72 Ma. This leaves approximately 40 Ma and narrowly limits the age of magnetization. Given that the magnetization predates deformation, the question of deposition rates and whether magnetization is of chemical or detrital origin is not particularly important in this case.

Table 2 presents the site mean directions, derived from the Characteristic Remanent Magnetization (ChRM) of the specimens for the sites used in statistical analysis. As specified in the rejection criteria, Table 2 indicates that for each site that was not removed, the α_{95} values are low and well below the maximum 20°. Since α_{95} is a statistical confidence level, the results suggest that the virtual geomagnetic poles (Table 2) are reliable.

The mineralogy of specimens in paleomagnetic studies can be estimated using thermal demagnetization techniques and determination of the Curie temperature of the specimens. The typical Curie temperatures of magnetite and hematite are 580° and 680° C, respectively. In the Ford Creek locality, approximately 35% of the magnetization

remains after 580° C, implying that the dominant carrier of the characteristic component of magnetization in these red rocks is hematite. This can be seen in Figure 6a which presents the ratio of magnetization intensity and original magnetization intensity (J/J_o) versus demagnetization temperature for a representative specimen from the Ford Creek locality. Similarly, Figure 6b displays this same plot for a representative specimen from the Marias Pass locality. Less than 10% of the magnetization remains after 580° C, implying that the dominant carrier of the characteristic component of magnetization in the Marias Pass locality is magnetite. The observation that the Marias Pass specimens have a high percentage of magnetite than the Ford Creek specimens explains why each locality had a different reaction to demagnetization techniques. Because of the high coercive force of hematite, far greater than that of magnetite, it does not react well to alternating field techniques. Therefore, Ford Creek specimens had to be demagnetized via thermal demagnetization techniques while the Marias Pass specimens were demagnetized using alternating field demagnetization techniques.

The three sites of the Marias Pass locality, all of which passed the rejection criteria at the specimen level (Table 3), have an average structurally corrected declination and inclination of 320.5°, 60.2° (α_{95} =11.3°, k=13.3). The declination of each of the three sites is not close to this locality value. This is because the site mean declination of Marias Pass site 3 (MP3) differs from that of the other two sites by 116°.

Two of the Marias Pass sites shows an obvious and easily determined secondary component of magnetization (Figure 4b). Marias Pass sites 1 and 2 (MP1 and MP2) have






Table 3. Structurally corrected primary and secondary directions for the Marias Pass locality.								
site	dec	inc	N/N _c	α95	k	δ ₆₃		
MP1	-77.2	67.0	5/6	12.5	38.4	11.7		
MP2	-78.9	66.7	4/6	8.9	107.4	6.8		
MP3	37.7	46.9	5/5	1 6.0	23.9	14.9		
MP1s	52.7	54.1	5/6	7.9	94.8	7.4		
MP2s	46.0	53.8	3/6	14.7	70.9	7.9		

MP represents site mean directions from Marias Pass sites. MP1s and MP2s represent secondary directions from Marias Pass sites. Dec indicates site mean declinations, clockwise is positive. Inc indicates site mean inclinations. N/N_c is the number of specimens used in statistical analysis versus number sampled. α_{95} =semiangle of 95% confidence, k=Fisher (1953) precision parameter, δ_{63} =angular standard deviation.

prominent secondary directions (MP1s and MP2s) which can be separated from their characteristic components and were successfully removed in a 50 mT field. Although MP3 is statistically different from the directions of MP1 and MP2, it is statistically coincident with their secondary direction (Figure 7). Removing MP3 from the site mean statistics decreases the α_{95} from 11.3° to 6.7° and increases k from 13.3 to 60.5. At the same time, combining the direction of MP3 with the secondary components of MP1s and MP2s produces a mean declination and inclination of 45.5°, 51.6° (α_{95} =6.6°, k=40.9). Site 3 seems to have recorded the stable, overprinted, secondary direction that was isolated in analysis of the other two sites of this locality and is probably the result of Viscous Remanent Magnetization (VRM). Although MP3 did not fail the statistical criteria listed prior, based on its recording of an overprint direction, it will not be considered in further statistical analysis as it appears to be VRM from the present-day field.

Specimens from the Ford Creek locality show a very minor secondary component of magnetization (Figure 4a). Many of the NRM directions are close to that of the present-day magnetic field. Their orthogonal projections, with curved trajectories during progressive demagnetization between 0° and 250° C, straighten out at temperatures greater than 250° C and head toward the origin. Because this component has a relatively low unblocking temperature and its in-situ direction is parallel to the present-day field, it appears to be a VRM.





Discussion

Marias Pass Locality

When comparing the in-situ directions of the Marias Pass locality to the presentday field, it becomes apparent that this locality has been subjected to some remagnetization (Table 4). The expected declination and inclination for this locality with respect to the present-day field are 16.9°, 72.3° respectively. The in-situ secondary declination and inclination of the Marias Pass locality are 6.7°, 71.1° (α_{95} =8.9°, k=34.3) and are statistically coincident with the present-day field (Figure 8). Thus the secondary direction is most likely VRM from the present-day magnetic field. However the characteristic direction is not statistically coincident with the direction of the present-day magnetic field. Therefore it is presumed that the structurally corrected, characteristic direction for the Marias Pass locality (Dec=281.9°, Inc=66.9°) represents insight into the tectonic history of the terrain. Unfortunately, since MP3 was interpreted as a complete VRM overprint from the present-day field and was rejected, only two reliable sites remain from the Marias Pass locality. Two sites, even with some within-site averaging of paleosecular variation, are insufficient to yield a robust average. A fold test could not be performed in this locality to assess the timing of acquisition of magnetization because no folds were present and the structural trends of both sites were very similar.

present day field.							
site	dec	inc	N/N _c	α95	k	δ ₆₃	
MPI	5.8	49.3	5/6	17.6	19.8	16.4	
MP2	2.3	51.8	4/6	8.8	109.6	6.7	
MPls	4.4	69.0	5/6	6.1	1 58.8	5.8	
MP2s	9.0	73.2	3/6	13.4	86.4	7.1	
Present day	16.9	72.3					

Table 4. In-situ site mean primary and secondary directions of the Marias Pass locality and the

MP represents Marias Pass sites. MP sites with an 's' indicated secondary directions. Dec indicates site mean declinations, clockwise is positive. Inc indicates site mean inclinations. N/N_c is the number of specimens used in statistical analysis versus number sampled. α_{95} =semiangle of 95% confidence, k=Fisher (1953) precision parameter, δ_{63} =angular standard deviation.



Figure 8. Equal area projection of the structurally corrected site-mean Marias Pass 1 & 2 directions (●) and the in-situ site-mean Marias Pass 1 & 2 directions (o) with respect to the present-day field (*). "c" represents characteristic directions (in-situ: α95=8.9°, k=34.3; structurally corrected: α95=6.7°, k=60.5).
"s" represents secondary directions (in-situ: α95=4.9°, k=129.6; structurally corrected α95=5.8°, k=93.5).

For the six Ford Creek sites that passed the arbitrary rejection criteria, paleomagnetic site mean directions were determined using ChRM directions (Table 2). The in-situ site-mean declination and inclination for the Ford Creek locality are 345.9° , 53.0° . The calculated declination and inclination of the present-day magnetic field for the Ford Creek locality are 16.2° , 71.6° . The direction of the Ford Creek locality differs from that of the present-day field by 23° ($\alpha_{95}=6.6^{\circ}$, k=12.3). Thus the Ford Creek locality is statistically different from the present-day magnetic field. This locality also passed a fold test, substantiating that the magnetization predates deformation and is not a record of a more recent, or present-day, field.

Cretaceous Reference Poles

Many paleomagnetic studies have been interested in a more accurate and continuous assessment of the Cretaceous Apparent Polar Wander Path (Runcorn, 1956; Larochelle and Black, 1963; Larochelle *et al.*, 1965; Van Alstine and de Boer, 1978; Irving, 1979; Irving and Irving, 1982; Gordon *et al.*, 1984; Globerman and Irving, 1988; Van der Voo, 1990; Gunderson and Sheriff, 1991). The Lower Cretaceous Kootenai Formation, as discussed above, has been limited in age to be Aptian (108-115 Ma) by Suttner (1969) and DeCelles (1986) by means of flora and fauna. Thus reference poles from the Lower and Middle Cretaceous are relevant to the paleomagnetic results from the Ford Creek locality.

Larochelle and Black (1963) and Larochelle *et al.* (1965), following the work of Runcorn (1956), determined a reference paleopole for the 109 ± 5 Ma Isachen diabasic rocks of the Northwest Territory, Canada. The pole is located at 69° N, 180° E (A₉₅=7.5°). Irving (1979) compiled the results of many studies to better determine the Apparent Polar Wander Path for the past 300 Ma. The path created contains a 110 Ma reference pole at 68° N, 185° E (A₉₅=5°). Later, Irving and Irving (1982) gathered paleomagnetic data for the same range of time in an attempt to better understand plate relations and Gondwana. To do so, they created Apparent Polar Wander Paths for North America, Northern Eurasia, Africa, South America, and Australia. The path for North America, which included the results of Larochelle and Black (1963), contains a 110 Ma reference pole at 69° N, 186° E (A₉₅=6°). The calculated Ford Creek declination and inclination, with respect to this paleopole, are 337.3°, 70.4°.

Gordon *et al.* (1984), following the work of Van Alstine and de Boer (1978), Irving (1979), Larochelle and Black (1963), and Larochelle *et al.* (1965) generated a synthetic continuous Cretaceous Apparent Polar Wander Path for North America. Their poles were calculated based on paleomagnetic Euler poles and rates of angular displacement of poles along their paths. Within the path created by Gordon *et al.* (1984) is a 110 Ma pole at 69.9° N, 189.6° E.

Globerman and Irving (1988) attempt to characterize the Middle Cretaceous paleomagnetic pole for North America by sampling Cretaceous-aged intrusive units in Arkansas. Beck (1991) cites this pole (λ , ϕ 74.1° N, 192.5° E, A₉₅=5.7°) as the product of "extensive mapping, painstaking laboratory demagnetization, and exacting selection criteria." It differs from the first Cretaceous reference pole of Runcorn (1956) by less than 5° and includes Larochelle and Black (1963), Larochelle *et al.* (1965), Irving (1979), and Gordon *et al.* (1984). Because of its encompassing nature, Beck (1991) cites this pole as one of high confidence.

Eldredge and Van der Voo (1988) recognized the lack of a tight constraint on the age of the Kootenai Formation. Although they cite Van Alstine and de Boer (1978), to solve the age problem they measured and averaged two virtual geomagnetic poles from relatively undeformed Kootenai sites near Great Falls, Montana. When comparing the virtual geomagnetic poles from their sites within the salients to that of the average of the two sites from the stable craton, they obtained results which removed the question of the age of their unit and the expected position of their pole on the Apparent Polar Wander Path. The paleopole obtained from the two sites that they averaged on the stable craton is located at 78.4° N, 170.6° E (A_{95} =5.3°).

The Cretaceous reference field is fairly well defined. The Middle Cretaceous reference pole of Globerman and Irving (1988) is well determined and is within 5° of the first reference pole of Runcorn (1956). Therefore, one would expect latitudinal differences in other Cretaceous reference poles to be minimal. However, the Globerman and Irving (1988) pole has a latitude nearly 5° north of the other three poles. This is a substantial amount and probably due to the fact that this is a Middle Cretaceous pole (88-

100 Ma) as opposed to an Aptian (108-115 Ma) pole. Although accurate, the Globerman and Irving pole (1988) is most likely of the wrong age for this study.

The reference pole position of Eldredge and Van der Voo (1988) differs from the mean of the other four poles (Larochelle and Black, 1963; Larochelle *et al.*, 1965; Irving, 1979; Irving and Irving, 1982; and Gordon *et al.*, 1984) by 10°. Based on this inconsistency, the pole of Eldredge and Van der Voo (1988) is probably inappropriate for this study.

Of the four remaining, similar, reference poles the earliest documented reference pole (Larochelle and Black, 1963, Larochelle *et al.*, 1965) is within 3° of the most recently published pole (Gordon *et al.*, 1984). Irving and Irving (1982) cite the work of Larochelle and Black (1963) and Irving (1979) in their Apparent Polar Wander Path study of the past 300 Ma. The Irving and Irving (1982) study can be viewed as an updated reference pole and is cited often in rotational studies. Their path is also based on real data as opposed to being synthetically generated. Thus errors based on their poles are easily determined and readily understandable with respect to other paleomagnetic studies. For this reason, I use the Irving and Irving (1982) pole as the reference pole for the Ford Creek directions in order to test the rotational hypothesis for the Purcell Anticlinorium.

Rotation of the Ford Creek locality

Symons and Timmons (1992) measured a clockwise rotation of 37±12° within the Proterozoic units of the Purcell Anticlinorium. A recent model of Sears (1994) predicts a

25-30° clockwise rotation of the anticlinorium primarily based on structural restrictions. The Ford Creek and Marias Pass localities were paleomagnetically sampled and measured to further test observations and hypotheses concerning the rotation of the Purcell Anticlinorium. Given the remagnetization at the Marias Pass locality and only two reliable sites, only the mean direction from the Ford Creek locality is compared to the reference pole determined by Irving and Irving (1982).

Based on the Lower Cretaceous reference pole of Irving and Irving (1982), the expected declination and inclination of the Ford Creek locality are 328.1°, 69.7°. The structurally corrected declination and inclination for the Ford Creek locality are 350.8°, 62.5° (Table 1). The declination and inclination differences and associated confidence intervals, calculated based on the methods of Demarest (1983) and Butler (1992), between the Irving and Irving (1982) pole and the Ford Creek locality are $23\pm9^{\circ}$ and $7\pm4^{\circ}$, respectively (Table 5) (Figure 9). This coincides with the clockwise rotation as measured by Symons and Timmons (1992) and provides a stronger basis for the hypothesis of Sears (1994). The apparent flattening of the observed inclination with respect to the expected inclination is statistically significant at the 95% confidence interval. Given my δ_{63} values are as expected, I suspect that the latitude discrepancy of these sedimentary rocks is due to a recording error known as inclination error (Butler, 1992). If true, this may explain why the reference pole of Eldredge and Van der Voo (1988) is also far sided with respect to Irving and Irving (1982).

The declination deviation is substantial and of similar magnitude as the rotational hypothesis of Sears (1994). It is also statistically coincident with the paleomagnetic

Table 5. Declination deviations between							
Ford Creek sites and a reference pole from							
Irving and Irving (1982).							
site	dec	inc	dec dif	inc dif			
FC3	-24.4	69.6	7+25	0+9			
FC4	-7.1	55.1	25+9	-15+5			
FC5	0.1	55.9	32+14	-14+8			
FC6	6.0	70.0	38+13	0+5			
FC7	-6 .7	64.3	25+18	-5+8			
FC8	-23.0	60.4	<u>9+28</u>	. -9 +13			
		Ford Creek:	23+9	-7+4			

FC represents Ford Creek sites. These are compared to the pole of Irving and Irving (1982) of latitude and longitude (69° N, 186° E (D/I 328.1°/69.7°). Dec dif indicates declination deviations, clockwise is positive. Inc dif indicates inclination deviations and are negative in the flattened direction.



Figure 9. Equal area projection of the structurally corrected Ford Creek locality direction (0, α 95=3.7°, k=37.3) with respect to the expected direction (\bullet , α 95=6°) based on Irving and Irving (1982).

evidence of Symons and Timmons (1992). Assuming the pole of Irving and Irving (1982) is representative of the Kootenai Formation, then the thrust sheet which houses the Ford Creek locality is likely to have undergone a rotation.

Paleomagnetic data from the Marias Pass locality are difficult to incorporate because one of three sites has been remagnetized. Yet the remaining two sites are likely to have some within-site averaging of paleosecular variation. Thus, based on the reference pole of Irving and Irving (1982), they suggest that this locality may have experienced a counterclockwise rotation of $45\pm16^{\circ}$ and $4\pm6^{\circ}$ of flattening. Symons and Timmons (1992) measured clockwise rotations of $37\pm12^{\circ}$ and this study measured $23\pm9^{\circ}$ of clockwise rotation in the Ford Creek locality. These two measurements are separated by approximately 300 km. The Marias Pass locality, while it is not a very robust indication of a counterclockwise rotation, is between these two measurements. Because the Ford Creek locality was the only reliable locality to compare with the reference pole of Irving and Irving (1982), it is difficult to determine the extent of the rotation. While the regional rotation model of Sears (1994) is one possible explanation for the results of this study, another explanation is that the thrust sheet has experienced a small scale vertical axis rotation based on impingement from the topography of the craton.

Conclusion

Six reliable sites from the Ford Creek locality indicate that the thrust sheet which houses this locality has rotated 23±9° clockwise based on the reference pole of Irving and Irving (1982) (Figure 9). With the partial remagnetization of the Marias Pass locality, statistics from this locality proved to be insufficient for a confident interpretation of the apparent counterclockwise rotation. With reliable paleomagnetic results from one locality to compare to a reference pole, it is difficult to discern whether or not this is a localized, vertical axis rotation (e.g. Grubbs and Van der Voo, 1976; Schwartz and Van der Voo, 1984; Eldredge and Van der Voo, 1988; Jolly and Sheriff, 1992) or whether this is indicative of a regional rotation of the Purcell Anticlinorium about an Euler pole (Symons and Timmons, 1992; Sears, 1994).

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Appendix A

The Kootenai Formation

The Lower Cretaceous Kootenai Formation (Figure 3) was chosen for paleomagnetic sampling. It was selected based on the mapping and stratigraphy of Dolberg (1986) and Mudge *et al.* (1982) and because the unit yielded good paleomagnetic results in the study of Eldredge and Van der Voo (1989). Good results implies four things: the Kootenai contains a measurable amount of magnetization, it reacted well to laboratory demagnetization, the magnetization predated deformation, and the unit adequately averaged paleosecular variation.

The Kootenai is a predominantly siliceous unit of sandstones, siltstones, and mudstones. Its basal member is a chert conglomerate and the formation is capped by a calcareous marker bed containing numerous gastropods. This biomicritic member, deposited in a fresh water lake, is called the "Gastropod Limestone." Another calcareous member within the Kootenai Formation splits the sandstones, siltstones, and mudstones into upper and lower members. The beds of the siliceous members generally alternate between red and green in the upper siliceous unit and red and brown in the lower siliceous unit. The equigranular, well-sorted, nonmarine oxidized sandstones of the lower siliceous member of the Kootenai Formation were sampled. (Suttner, 1969; DeCelles, 1986)

When units alternate between red and green layers, this usually implies that both the oxidized and reduced form of iron are present. Hematite $([Fe^{3+}]_2[O^{2-}]_3)$ is more oxidized than its reduced counterpart, magnetite $([Fe^{2+}] [Fe^{3+}]_2[O^{2-}]_2)$. Since very red layers were sampled in the Ford Creek locality, a high hematitic content was expected. In the Marias Pass locality, the beds sampled were not as brick red as the beds in the Ford Creek locality.

Suttner (1969) dates the Kootenai Formation based on gastropods and pelecypods. The pelecypods abundant in most of the Kootenai are Aptian in age (108-115 Ma). Also, units of the Blairmore Group, of Alberta Canada, have been correlated with the Upper Kootenai. The top of the Blairmore Group has been dated through flora as Albian in age (97-108 Ma). DeCelles (1986) also dates the Kootenai as Aptian and brackets it by dating the underlying Morrison Formation and the overlying Blackleaf Formation. The Kootenai Formation rests unconformably on the Morrison Formation, which has been dated by fossils as Kimmeridgian age (143-146 Ma). The overlying, conformable, Blackleaf Formation has been dated as Albian age (97-108 Ma).

Appendix B

Future Studies

Based on the paleomagnetic data of this study, the hypothesis of Sears (1994) and model of Symons and Timmons (1992) could not be supported or refuted. One site from the Marias Pass locality was rejected based on an overprint form the present-day field. This left only two sites in the Marias Pass locality for statistics and analysis. The Ford Creek locality recorded a 23±9° clockwise rotation from the expected Early Cretaceous reference pole of Irving and Irving (1982). However, this measurement is not conclusive of a rotation of the Purcell Anticlinorium as a cohesive unit. Instead it may be associated with a small, localized, vertical axis rotation. To determine the extent of the rotation, either more data is needed from the Marias Pass locality, or paleomagnetic data from a new locality is needed.

The Cretaceous units of Marias Pass locality offered few outcrops without extensive fracturing. For these reasons, if Cretaceous units are to be sampled, Marias Pass may not an ideal choice in a locality. Another possibility is to choose non-Cretaceous units. In this case, it would be optimum to sample two new localities rather than try to correlate the results to those from Ford Creek. This may eliminate error.

Great care should be taken in choice of new localities. In light of the result of this study and of previous studies, the craton seems to have considerable effects on the thrust sheet at its margin. For this reason, localities should be chosen where the thrust sheet is substantially thicker and acting as a cohesive unit.

Appendix C

Data

Below is a compilation of Zijderveld (1967) diagrams for all of the specimens from the six non-rejected Ford Creek sites and all three Marias Pass sites. Each site has a cover page followed by the structurally corrected Zijderveld (1967) diagrams, for each specimen, with a best fit characteristic remanent magnetization direction.

A general explanation of the diagrams is as follows:

• Each diagram can be divided into two sections: data and figure.

• The data section is comprised of eight columns.

• Column 1 is the demagnetization step. Numbers in this column that are on the order of 10 imply AF demagnetization; Numbers in this column on the order of 100 imply Th demagnetization.

• Columns 2, 3, and 4 are the x, y, z vector component intensities of the specimen. The units for these values are 10E-6 A/m.

• Columns 5 and 6 are the directional components of declination and inclination, respectively.

• Column 7 is the total intensity of the specimen. The units for these values are 10E-6 A/m.

•Column 8 is the measurement error.

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• Between the data and the figure is the information gained from the trend of the best fit line, including declination and inclination.

• The figures are typical Zijderveld (1967) diagrams. The convention of this study is to plot declination (•) information with north on the vertical axis. This plots inclination (o) information on the up and east axes.

• The upper left corner of the figure contains a "th" for thermal demagnetization (which is inaccurate for a few of the Ford Creek specimens and nearly all of the Marias Pass specimens) and a "tc" for "tilt" corrected (this is accurate).

• The upper right corner contains the specimen number and, when appropriate, a "B" for the second specimen from that sample. If an "F" appears after the specimen, it means that the specimen was subjected to AF demagnetization. The "F" convention was only used in the Ford Creek locality as Th demagnetization was the norm.

Ford Creek 3

Specimens:

11, 12, 13, 14 & 18 (avg), 16, 17 & 17F (AF demag)

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In-situ:

Dec=-74.31°, Inc=55.17°

 $\alpha_{95} = 10.11^{\circ}$

k=44.88

δ₆₃=11.06°

Structurally corrected:

Dec=-24.44°, Inc=69.57°

 $\alpha_{95} = 10.15^{\circ}$

k=44.52

δ₆₃=11.10°





















Ford Creek 4

Specimens:

19, 20, 21 & 23 (avg), 22, 25 & 25B, 26 & 26B, 27 & 27B

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In-situ:

Dec=7.51°, Inc=57.08°

α₉₅=5.21°

k=135.05

δ₆₃=6.46°

Structurally corrected:

Dec=-7.06°, Inc=55.05°

 $\alpha_{95} = 5.18^{\circ}$

k=136.74

δ₆₃=6.42°






















Ford Creek 5

Specimens:

28 & 28B & 29 & 29B (avg), 30, 31, 32 & 32B, 33 & 33B, 34 & 35 & 35b, 36, 37 In-situ:

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Dec=3.42°, Inc=61.46°

α₉₅=8.96°

k=39.16

δ₆₃=12.14°

Structurally corrected:

Dec=0.10°, Inc=55.84°

α₉₅=8.94°

k=39.36

δ₆₃=12.10°























STEP	*	В	c	DECtc	INCtc	INTENS	ERR	DATE	TIME	
										80
100 0	-30	1289	663	310.4	44.9	1450	1.4			
100 -	-55	1329	548	311.1	46.3	1480	0.8			
200 -	-23	1235	546	314.7	46.4	1350	2.8			
300 •	-66	1121	424	317.3	43.8	1200	1 4			
350 *	40	997	464	315.4	43.2	1100	2 9			
400 *	51	913	315	323.3	45 0	967	2.6			
450 *	84	886	193	333 4	45 0	911	3 6			
500 +	35	568	191	124 1	44 9	600	3.3			
530 .	44	443	171	333 0	43.0	427	3.8			
545 -	67	437	60	340 7	42.0		3.1			
555 *	29	339		340.7	42.7	440	3.8			
675 •	-20	- 2	13	331.4	45.0	350	4.3			
••••	- 20	-2	14	208.0	26.5	24	13.9			
						••.	•			











Ford Creek 6

Specimens:

38, 39 & 41 (avg), 40, 42, 43, 45, 46

In-situ:

Dec=-16.98°, Inc=60.51°

 $\alpha_{95}=5.10^{\circ}$

k=141.07

δ₆₃=6.32°

Structurally corrected:

Dec=6.04°, Inc=70.01° α_{95} =5.09°

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k=141.68

δ₆₃=6.31°

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STEP	*	В	С	DECtc	INCTO	INTENS	ERR	DATE	TIME	
										91
0	-2277	985	-671	43.1	72.5	2570	2.9			
150	-2183	970	-801	31.7	72.5	2520	3.1			
300 *	-2275	1038	-1069	17.1	71.3	2720	2.3			
450 +	-2441	1094	-1119	18.9	71.2	2900	3 3			
S\$0 *	-1016	457	-472	18.1	71.1	1210	2 0			
570 +	-1111	483	-497	20.8	70.8	1310	2.1			
590 *	-1037	483	-479	17.8	71 9	1240				
610 -	-1130	381	-405	24.1	67.5	1160	2.0			
610 -	-967	382	-369	27.6	69.7	1110	0.3			

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STEP		*	B	с	DECtc	INCLC	INTENS	ERR	DATE	TIME	
•											95
160		-1074	1469	-643	16.4	79.9	1930	3.3			
150		-1034	1392	-727	5.4	77.4	1880	3.2			
300		-942	1271	-503	26.5	80.5	1660	2.9			
450	•	-738	935	-441	18.7	77.8	1270	3.5			
550	•	-375	425	-202	27.6	75.5	607	3 0			
570	*	-373	359	-147	44 R	72 1	519	2.0			
590	•	-280	290	-122	40.7	71.9	A11				
610	•	-240	191	-213	2 7	59.1	305	3.4			
610	•	-175	164	-146	6.7	50.1	383	3.1			
640	•	-153	142	-440	34.5	08.0	441	9.6			
650	•	-185	141	-97	44.0	69.3	227	1.8			
		-307	222	-138		59.7	282	3.9			
				~14	//.5	D4.8	385	1.6			
							-				
	**		*********								

INCLOR: dec = 16.4 inc = 78.4 int = 1929 mad = 3.9








step		*	в	c	DECtc	INCLC	INTENS	ERR	DATE	TIME	
											99
0	•	-2500	-798	284	348.9	72.5	2640	0.9			
150	*	-2447	-712	239	349.1	74.2	2560	1.4			
300	•	-2378	-647	172	346.4	75.7	2470	1.0			
450	•	-1898	-621	116	319 5	73.4	2000	1.0			
550	•	+1163	-109	164	350.7	70.7	2000	2.0			
570	•	+1155	-339	120	340.7	22.0	1240	0.16			
590	•	-1040	- 3 3 5	120	347.8	/3.9	1210	Q.8			
610	•	-1040	-382	156	349.7	69.3	1120	1.0			
010	-	~1007	-356	60	337.5	72.2	1070	0.8			
610	•	-973	-308	137	353.1	71.8	1030	3.5			
640	•	-1002	-371	215	357.2	67.4	1090	1 0			
650	•	-751	-211	76	350.9	74 6	794	1.7			
660	•	-704	-158	190	18.7	69.8	746	2.3			





Ford Creek 7

Specimens:

47, 48, 49, 50, 51, 52, 53, 55

In-situ:

Dec=20.46°, Inc=32.72°

α₉₅=8.88°

k=39.88

δ₆₃=12.02°

Structurally corrected:

Dec=-6.69°, Inc=64.28°

α₉₅=8.87°

k=39.92

δ₆₃=12.02°

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STEP		*	D	c	DECto	INCtc	INTERS	ĒRR	DATE	TIME	
											101
		-476	-669	86	38.8	41.3	826	1.9			
150	•	-492	-544	64	36.9	47.9	736	4.4			
300	•	-343	-522	79	40.5	39.2	629	2 4			
450	•	-383	-366	119	48 9	51 6	547	1.0			
550	•	-283	-252	45	38.6	54 1	343				
		-386	_10	- 710	30.0	24.4	301				
590	•	-336	- 59	-210	300.0	37.7	441	2.9			
		-336	* 3 9	-11	335.1	80.9	341	0.8			
610		-033	345	-342	261.1	44.3	798	2.1			
630	-	-316	-79	-21	346.9	77,1	326	1.0			
	-	~276	-123	3	20.4	70.8	302	7.7			
650	•	-312	-128	-32	1.7	69.8	339	2.4			
660	•	-287	-96	-69	340.3	67.1	311	3.6			

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INCLOR: dec = 36.6 inc = 50.9 int = 826 mad = 14.9





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STEP		A	B	с	DECto	INCto	INTENS	ERR	DATE	TIME	
_											103
0	•	-1056	-425	-345	353.0	60.4	1190	2.6			
150	•	-995	-486	-309	359.8	58.2	1150	1.4			
300	•	-804	-432	-338	344 4	53 6	973	1.0			
450	•	-603	-419	-157	12.7	52.5	751	1.1			
550	•	-499	-295	-162	4 1	54 5	603	1.9			
570	•	-498	+57	-187	321 6	65 3	602	V. 9			
590		-390	-142	-100	340 7	85.3	233	1.4			
610		-181	-761	-190	340.7	55.7	456	2.4			
630	•	-335	-431	-454	349.9	44.5	524	1.9			
640		-323	-/9	-105	338.9	65.0	351	1.9			
660		-244	-162	-125	356.9	48.0	319	4.6			
050	-	-223	-137	-102	357.5	50.6	281	4.0			
00V	•	-249	-99	-91	350.2	59.2	283	1.9			



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STEP	A	B	c	DECto	INCLC	INTENS	ERR	DATE	TIMZ	
										106
0 *	-1059	-721	-152	343.3	56.7	1290	0.8			
150 -	-1049	-562	-32	342.9	64.7	1190	1 3			
300 •	-949	-557	-13	347.7	63.6	1100	1 1			
450 +	-731	-164	141	329 5	84 2	767	1.4			
550 *	-470	7	-12	262.1	70.4	470				
570 •	-480	59	-25	251.0	65 6	484	7.0			
590 ·	-278	-65	-41	303 2	66 1	199				
610 -	-218	73	-119	261 0	39 5	260	1.4			
630 *	-195	9	-77	275 7	51 3	210				
640 *	-231	- 97	-52	370 6	59 7	366	0.4			
	-227	91	16	221.1	57.3	245	15.9			

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STEP	*	8	с	DECLC	INCLC	INTENS	ERR	DATE	TIME	
										107
0 •	-1779	-644	-382	346.0	67.6	1930	0.9			
150 •	-1673	-710	-287	356.1	67.2	1840	0.4			
300 •	-1488	-679	-335	353.7	64.4	1670	1 1			
450 *	-771	-307	-481	178 9	67 6	000				
550 -	-640	-690	-196	13.3	45 1	961	4.4			
570 -	-416	-501	- 246	13.3	42.1	201	0.9			
590 -	-687	-303	-440		30.7	638	3.0			
610 .		-353		30.8	00.0	795	1.0			
630 +		-321	60	37.9	62.4	587	2.3			
6.00	-101	-118	-10	21.3	58.5	201	2.9			
640 -	-227	-125	-68	354.6	59.1	268	2.1			
650 *	-335	-13	-157	300.7	58.9	370	3 0			



STEP	A	8	ç	DECtc	INCto	INTENS	ERR	DATE	TDE	•••••
										108
0 •	-2333	-608	-375	329.0	72.4	2440	1 4			
150 •	-2108	-618	-318	334.9	72.2	2220	1 0			
300 -	-1924	-553	~440	328.4	68.6	2050	1 1			
450 -	-1518	-364	-393	320.5	67.9	1610	2.5			
550 •	-711	-197	-284	320.2	60 7	791	1 7			
570 •	-557	116	-169	266 1	50 6	597	5 6			
590 *	-526	-54	-297	302 4	53 5	606	4.0			
	-693	605	418	188 6	40.5	1010	2.5			
630 *	-324	169	-152	256 6	44 6	196	10 6			
640 *	-617	172	-283	269 6	\$1.7	300	20.0			
650 -	-369	216	-4	225.0	52.8	445	11.0			



Ford Creek 8

Specimens:

58, 59, 60, 61, 64

In-situ:

Dec=-24.92°, Inc=50.90°

α₉₅=17.60°

k=19.85

δ₆₃=16.32°

Structurally corrected:

Dec=-23.03°, Inc=60.42°

α₉₅=16.11°

k=23.52

δ₆₃=14.99°

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FEP	A	в	с	DECte	INCEC	INTERS	ERR	DATE	TIME	
_										11
0	-474	1	-124	292.6	73.0	490	2.1			
2 *	-432	-7	-66	312.3	76.0	437	1.0			
4 *	-412	-17	-53	320.1	75.4	416	2.9			
6 •	-379	-22	-58	316.8	73.9	384	1 5			
	-357	-18	-74	306.5	72.8	365	3.1			
10 -	-337	-6	-62	307.1	75.0	343				
14 *	-278	-17	-56	308 7	72 4	284	6 6			
18 •	-264	-12	-47	310 7	74 0	268	1 1			
22 -	-226	-23	-12	321 7	72 0	220				
26 .	-208	-24	-65	300 4	66 5		2.1			
30 .	-186	-47	-78	302.0	57 0	202	3.2			
36 .	-172	-16	-63	204 6	65 0	107	3.0			
42 .	-116	- 24	-67	234.0	63.9	103				
		-44	-0/	407.1	24'5	130	7.1			



-	^	B	c	DECCC	INCLE	INTENS	ERR	DATE	TDE
0 2 •	-707 -734	-135 -54	-338 -213	332.7 343.4	45.9 54.4	795 766	2.6		
6 •	-642 -556	-95 -135	-144 -129	350.5 351.1	51.1 46.0	665 587	1.3 3.0		
2 •	-436	-194 -219	-98 -53	353.9 359.7	36.8	503 474	1.8 2.4		
5 •) •	-325	-198	-105	338.2	27.8	504 421	3.2 3.0		
•	-436 -359	-174	-170	342.9	37.3	499	6.3 3.1		
•	-458 -375	-133 -171	-52	0.3	43.8	480	3.6		
	-307 -273	-149 -70	-112	345.5	33.7	359	2.2		
•	-213 -153	-229 -205	-58 121	354.3 29.9	13.4 3.6	318 283	5.1 1.8		
•	-189 -211	-22 -200	-8 -73	7.0 350.8	53.1 17.0	190 - 300	3.4 3.1		
	-167	-58	-38	352.4	40.7	181	4.3 478		
CLOR:	de c = 3	350.9	inc =	39.4	int =	766	mad =	13.5	
 th / to	2				1			<u></u>	59.000
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step		*	В	c	DECto	INCto	INTENS	ERR	DATE	TIME	
											112
150	•	-2248	244	-743	333.0	80 4	2260	2.4			
300	•	-2058	251	-765	324 9	78.8	2300	4.4			
		-2131	504		147 7	80.3	2210	2.0			
550	•	-974	221	-962	202 6	60.4	2190	3.0			
570	•	-885	106	-604	303.0	30.0	1320	4.1			
590	•	-803	_44		319.2	03.0	1080	2.6			
610		-779	- 30		343.1	67.7	904	2.8			
630		-774	103	-491	313.5	67.2	921	3.5			
640	-	-3/6	366	-481	276.1	72.2	1150	2.7			
640		-657	535	-186	225.1	61.9	867	2.3			
650	•	-386	85	-273	303.2	64.9	481	1 7			
660	•	-16	- 94	111	85.6	-8.0	147	7.9			

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STEP		×	B	c	DECto	INCte	INTENS	ERR	DATE	TIME	
	_										114
		-886	-49	-482	320.9	61.5	1010	2 3			
150	•	-965	-10	-414	325.9	66.9	1050	2.0			
300	•	-543	137	-212	291.1	75.8	500	3.4			
450	•	-544	53	-194	321 2	73 5	500	3.4			
550	+	-363	49	-732	740 7	63.3	580	0.7			
570	•	-343	47	-207	277.3	23.2	434	1.0			
590	•	-374		-207	300.1		403	16.2			
610		-295	63	-83	312.6	82.6	393	6.0			
630		-203	-98	-136	337.0	56.3	323	4.3			
640	1	-214	7	-172	303.2	55.4	275	8.7			
560	-	-263	2	-222	303.7	53.7	344	4.6			
\$ 50	•	-81	50	-25	222.9	. 72.9	98	3.0			
•		-124	-75	26	32.2	39.8	147	4.8			

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Marias Pass 1

Specimens:

101 & 103 (avg), 102, 104, 105, 106, 107	
In-situ: (Characteristic)	(Secondary)
Dec=5.75°, Inc=49.34°	Dec=4.44°, Inc=68.95°
α ₉₅ =17.63°	α95=6.09°
k=19.78	k=158.84
δ ₆₃ =16.35°	δ ₆₃ =5.75°

Dec=-77.23°, Inc=67.03°	Dec=52.67°, Inc=54.13°
α ₉₅ =12.51°	α ₉₅ =7.9°
k=38.38	k=94.77
δ ₆₃ =11.72°	δ ₆₃ =7.45°

$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	STEP	X	3	c	DECto	INCto	INTENS	ERR	DATE	TIME	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	٥	14300									110
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		-16/33	25570	12552	261.6	55.1	32100	0.8			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		-12/82	28179	5107	239.2	63.5	32700	0.6			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	2	-14448	26693	4609	239.2	64.4	30700	1.4			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		-12741	24886	3685	238.0	66.0	28200	1 0			
		-11578	22674	3678	239.7	66.0	25900	à à			
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	10	-10051	21462	3789	243.7	67.2	24000	0.4			
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	14	-\$105	19161	3518	247.7	68.8	21100	10			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	18	-6483	16652	3452	253.4	69.4	18200	4.U			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	72	-5699	15440	2832	252.4	70.9	16700	0.0			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	28	-4783	13709	2867	257.7	70 7	14800	V.0			
	32	-4237	12390	2845	260.7	70.2	13400	4.4			
	40	-3062	11417	2069	264 7	74 6	12000	0.5			
$60 *$ -1090 8125 1578 272.3 75.0 8090 0.7 $70 *$ -2026 6294 1543 264.7 70.3 6790 0.8 $80 *$ -1555 5985 1761 277.0 69.8 6430 0.8 $90 *$ -1149 4879 1527 281.6 69.5 5240 1.5 $100 *$ -409 4181 1205 102.8 72.1 4370^{-1} 1.1 $120 *$ -470 2901 1078 294.2 67.6 3130 2.2	50 -	-2701	9729	2033	266 7	22.3	12000	4.1			
$70 + -2026$ 6294 1543 264.7 70.3 6690 0.7 $80 + -1555$ 5985 1761 277.0 69.8 6430 0.8 $90 + -1149$ 4679 1552 281.6 69.5 5240 1.5 $100 + -409$ 4181 1205 302.8 72.1 4370^{-1} 1.1 $110 + -1194$ 3978 240 228.6 67.6 110 2.2	60 -	-1890	8125	1578	272 3	75.0	10300	U.4			
80 * -1555 5985 1761 277.0 69.8 6430 0.8 90 * -1149 4879 1527 281.6 69.5 5240 1.5 100 * -409 4181 1205 302.8 72.1 4370- 1.1 110 * -1194 3978 240 238.6 77.2 4160 1.0 120 * -470 2901 1078 294.2 67.6 3130 2.2	70 *	-2026	6294	1543	264 7	20.3	6670	0.7			
90 • -1149 4879 1527 281.6 69.5 5240 1.5 100 • -409 4181 1205 102.8 72.1 4370 1.1 110 • -1194 1978 240 238.6 77.2 4160 1.0 120 • -470 2901 1078 294.2 67.6 1100 2.2	80 *	-1555	5985	1761	272 0	66.8	6790	0.0			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	90 -	-1149	4878	1677	301 6	63.0	6430	0.8			
110 * -1194 3978 240 238.6 77.2 4160 1.0 120 * -470 2901 1078 236.2 67.6 3130 2.2	100 +	-409	4181	1205	101.0	67.3	5240	1.5			
$120 \cdot -470 = 2901 = 1078 = 296.2 67.6 = 3130 = 2 2$	110 .	-1194	3976	240	304.6	14.1	4370-	1.1			
	120 •	-470	39/18	1070	238.6	77.2	4160	1.0			
			4301	10/8	296.2	67.6	3130	2.2			

INCLOR: dec = 272.2 inc = 72.8 int = 10300 mad = 4.2





TEP	*	8	c	DECtc	INCLE	INTENS	ERR	DATE	TIME	
0	-16734									
2	-14023	30200	440	259.9 253.1	60.8 71.0	31900 33300	1.7			
4	-12664	28637	533	256.4	71.6	31500	0.6			
8	-9960	23691	463	258.1 257.8	72.1	28900	0.9			
LO L4	-8575	21880	-56	260.5	73.9	23500	0.9			
8	-5175	17031	-26 114	270.9	75.4	20600	0.6			
2	-3946	15709	303	288.1	75.4	16200	0.5			
2 *	-2056	13405	445	290.3	74.7	13800	0.9			
0 *	-1505	11267	862	310.4	71.7	11400	1.1			
io +	-959	9162 7576	999 234	312.9	69.5 74 0	9270	0.9			
70 • 10 •	-956	6174	639	305.8	70.4	6280	0.5			
• •	-54	4379	262	310.3	72.9	5120	0.5			
	380	3818	-157	355.1	69.1	3840	0.5			
0 -	340	3560	180	325.5	71.3	3570	1.0			
							, 			
CLOR:	dec =	313.2	inc =	72.8	int :	= 12199	mad	= 4.1		
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EP	A	3	c	DECLC	INCLC	INTENS	ERR	DATE	TIME	
0	-14403									1
2	-14471	15394	1595	248.8	59.9	21200	1.6			
Ā	-13636	15976	589	245.2	62.1	21900	1.7			
6	-13030	15402	346	246.3	63.4	20700	1.6			
ě	-12199	14435	92	248.2	65.0	18900	1.3			
າດັ	-10/45	13174	102	250.2	65.6	17000	2.3			
14	-9425	12331	114	255.1	66.9	15400	1.6			
16	-7353	11322	77	263.7	68.9	13500	1.7			
22	-5924	10081	-422	267.8	72.2	11700	2.3			
20	-5053	9651	-386	276.7	73.0	10900	1.6			
40 20	-3722	9239	-425	296.6	73.9	9970	1.2			
32	-3419	8765	-196	298.4	72.6	9410	1.3			
	-3394	8357	12	295.4	71.4	9020	1.1			
50	-2634	7870	-119	307.1	71.7	8300	1.1			
60 ·	-1885	7272	-328	320.9	71.8	7520	1.3			
/0 -	~2160	6297	-554	309.5	75.6	6680	1.7			
80 -	-1329	6148	-0	323.4	68.5	6290	1.4			
30 -	-1766	5071	-176	305.0	72.9	5380	1.1			
00 *	-1557	4480	-258	307.0	74.1	4750.	1.2			

INCLOR: dec = 315.5 inc = 72.6 int = 7520 mad = 3.3



EP	*	8	c	DECtc	INCtc	INTENS	ERA	DATE	TIME	
0										1
š	-17037	19275	3003	231.1	59.4	25900	1.4			
2	-18153	19226	3705	232.5	57.2	26700	1.1			
2	+16909	18296	3793	234.1	\$7.6	25200	0.7			
	-15183	16747	3665	235.4	\$7.9	22900	1 6			
8	-13401	15153	3326	236.0	58.6	20500	1 3			
10	-11637	13530	2985	236.8	59.3	18100	1 9			
14	-9446	12552	2456	238.4	62.9	15900	1 2			
16	-7810	11152	2251	241.2	64.5	13800	1.0			
22	-6495	10173	1783	241.6	67 1	12200	1.0			-
28	-5476	9306	1492	242.8	69.7	10000	1.0			
14 · *	-4522	8678	1562	250.2	21 0	8910	0.8			
12 *	-4338	7305	1400	246 1	68 7	5510 8610	4.3			
50 *	-3712	6838	189	230.7	77 8	3300	1-2			
50 •	-3129	5824	719	241 3	74.9	7750	1.6			
70 +	-2242	4607	944	766 6	71.9	6630	1.0			
80 *	-1429	4961	222	239.3	71.4	5210	1.3			
• •	-1813	3617	443	203.4	83.6	5170	0.6			
)o +	-1388	3021		243.7	73.3	4070	2.5			
-	- 1966	3231	312	466.5	76.8	3530	1.6			
							~• •			
					-					

INCLOR: dec = 245.4 inc = 72.3 int = 9910 mad = 4.4



0 2	-14723								
2									12
•		19802	8193	257.5	58.1	26000	0.9		
	-15849	19435	10004	259.2	54.1	27000	1.0		
2	-14624	18775	9434	260.4	55.0	25600	1.4		
<u>,</u>	-13110	17833	8749	262.2	56.1	23800	0.B		
	-11854	16292	8324	263.6	55,7	21800	0.8		
10	-10256	15202	7729	266.6	56.6	19900	0.3		
14	-8809	13916	7262	269.8	56.9	18000	0.3		
18	-7855	12556	7043	272.0	55.7	16400	0.6		
22	-6498	11534	6838	277.3	55.4	14900	0.0		
28	-5616	10774	5884	278.7	57 5	13500			
34	-5037	9694	5652	280.0	56.2	12200	1.9		
42	-3905	8558	5099	285 1	56 7	10700	0.0		
50 *	-3360	7624	4649	786 6	55.0	10700	0.6		
60 •	-2643	6062	1601	286.5	33.8	3540	0.3		
70 •	-2129	4762	2044	200.0	30.4	7530	0.4		
80 -	-1945	4736	2346	450.3	\$5.4	5990	0.6		
90 *	-1438	4378	3174	388.5	51.2	5700	1.2		
100 .	-4440	4130	2050	294.5	54.8	5120	1.4		
120 -	-050	3/85	2006	307.9	58.8	4340	1.3		
120 .	.1159	3483	2103	295.0	56.3	4230	1.2		
130 -	-1419	2993	1674	282.8	57.4	3710	1.1		

INCLOR: dec = 289.1 inc = 55.7 int = 9540 mad = 3.6



STEP	×	B	c	DECCC	INCLC	INTENS	ERR	DATE	TIME	
٥	-74570									12
ž	-26462	8843	-7742	241.2	50.1	27200	1.0			
i i	-24763	7989	-7720	241.2	53.2	28700	1.3			
ŝ.	-29741	1516	-7349	241.5	53.0	26900	1.4			
Ä	-20720	6618	-6946	242.6	53.5	24700	1.4			
30	-20/20	5453	-6869	245.0	54.3	22500	1.8			
14	-10943	4340	-6791	249.4	55.2	20500	1.8			
18	-10/01	3459	-6192	251.3	56.0	18200	1.8			
22	-14333	2432	-5960	256.4	56.7	15900	1.4			
28	-13049	1957	-4970	255.2	58.2	14100	1.2			
32	-11662	1400	-\$333	262.6	57.2	12900	1.5			
40. *	-10665	1200	-5536	267.4	55.3	11900	1.5			
50 4	-9039	827	-4868	269.5	55.8	10300	1.6			
60 +	-7715	473	-4037	270.7	57.2	8720	1.7			
70 9	-6241	610	-3373	269.2	55.5	7120	1.4			
	-5348	~25	-2675	274.4	60.2	5980	1.5			
6V -	-4513	228	-3153	279.3	51.9	5510	1.3			
30 -	+4379	-75	-3293	285.4	\$1.8	5480	1.6			
100 -	-3361	1007	+3082	273.7	39.9	4670-	1.2			
110 *	-3104	-386	-2412	293.1	53.1	3950	1.7			
					•					
********	*********									
INCLOR :	dec =	273.7	inc =	55.0	int	 = 10300				-



Marias Pass 2

Specimens:

108 & 109 (avg), 110, 111, 112, 113, 11	4
In-situ: (Characteristic)	(Secondary)
Dec=2.26°, Inc=51.82°	Dec=9.01°, Inc=73.17°
α ₉₅ =8.81°	α ₉₅ =13.35°
k=109.62	k=86.35
δ ₆₃ =6.71°	δ ₆₃ =7.12°

Structurally corrected:

Dec=-78.87°, Inc=66.66°	Dec=45.96°, Inc=53.75°
α ₉₅ =8.91°	α ₉₅ =14.75°
k=107.36	k=70.93
δ ₆₃ =6.78°	δ ₆₃ =7.86°

$\frac{0}{2} - \frac{13247}{15992} = \frac{8955}{10663} - \frac{3966}{1981.5} = \frac{1381.5}{41.5} = \frac{43.2}{47.3} = \frac{17100}{1.0} = \frac{1.0}{1.2} + \frac{13902}{1.1} = \frac{13902}{15775} = \frac{13776}{3977} = \frac{233.9}{233.9} = \frac{61.4}{21200} = \frac{21400}{1.0} = \frac{1.0}{1.0} + \frac{13902}{1.1} = \frac{13642}{1577} = \frac{6305}{237.7} = \frac{263.7}{5.7} = \frac{21700}{21700} = \frac{1.0}{1.0} + \frac{13642}{100} = \frac{13642}{11204} = \frac{16784}{1212} = \frac{263.1}{263.1} = \frac{664.3}{64.3} = \frac{21100}{110} = \frac{0.5}{1.0} = \frac{13642}{110} = \frac{13642}{11205} = \frac{76464}{13836} = \frac{13642}{7514} = \frac{276.0}{276.0} = \frac{64.3}{64.3} = \frac{18400}{110} = \frac{0.3}{1.0} = \frac{13642}{11205} = \frac{1366.7}{225.0} = \frac{63.5}{63.0} = \frac{13500}{13} = \frac{0.5}{122} = \frac{13666}{6673} = \frac{297.5}{297.5} = \frac{63.0}{63.5} = \frac{13900}{110} = \frac{0.5}{122} = \frac{1366}{12986} = \frac{667.7}{2275} = \frac{263.1}{230.5} = \frac{136.4}{64.1} = \frac{14600}{1.1} = \frac{0.5}{1200} = \frac{0.5}{122} = \frac{1366.7}{1295} = \frac{136.7}{1200} = 1$
$\frac{1}{2} - \frac{1592}{1066} - \frac{1935}{1066} - \frac{193.5}{10} + \frac{17100}{10} - \frac{1.0}{12} - \frac{16846}{113698} - \frac{101}{10} + \frac{1122.3}{100} - \frac{56.7}{20200} - \frac{1.1}{12} - \frac{11460}{10} - \frac{1.0}{10} - \frac{11440}{10011} - \frac{16929}{1000} - \frac{6105}{2005} - \frac{243.2}{205.7} - \frac{63.7}{63.7} - \frac{21000}{21000} - \frac{1.0}{1.0} - \frac{100114}{10011} - \frac{16944}{10594} - \frac{7157}{70157} - \frac{257.7}{257.7} - \frac{63.7}{63.7} - \frac{21000}{21000} - \frac{1.0}{1.0} - \frac{100114}{10011} - \frac{16744}{10594} - \frac{7431}{7411} - \frac{265.6}{205.1} - \frac{14400}{100} - \frac{1.5}{100} - \frac{1.5}{100} - \frac{1.5}{1000} - \frac{1.5}{10000} - \frac{1.5}{1000} - \frac{1.5}{100} - \frac{1.5}{100} - \frac{1.5}{1000} - \frac{1.5}{100} - $
$\frac{2}{4} -\frac{14846}{13698} + \frac{136.5}{132.3} + \frac{47.3}{56.7} + \frac{18900}{20200} + \frac{1.2}{1.1} + \frac{13902}{15775} + \frac{13978}{3978} + \frac{233.9}{233.9} + \frac{61.4}{61.4} + \frac{21400}{1.0} + \frac{10}{1.6} + \frac{15934}{1440} + \frac{15934}{15934} + \frac{7157}{7157} + \frac{257.7}{257.7} + \frac{63.7}{63.7} + \frac{21700}{2100} + \frac{10}{0.6} + \frac{16414}{16784} + \frac{16784}{7421} + \frac{263.1}{263.1} + \frac{64.3}{64.3} + \frac{21100}{110} + \frac{0.5}{14} + \frac{16859}{15711} + \frac{7411}{7411} + \frac{269.6}{264.6} + \frac{64.5}{64.5} + \frac{13500}{15900} + \frac{0.5}{0.5} + \frac{128}{22} + \frac{-6668}{13039} + \frac{75314}{7571} + \frac{276.0}{277.6} + \frac{64.3}{64.8} + \frac{13600}{110} + \frac{0.3}{0.3} + \frac{128}{22} + \frac{-5663}{13039} + \frac{7356}{731} + \frac{276.0}{287.5} + \frac{65.0}{65.0} + \frac{13500}{285.0} + \frac{0.5}{0.5} + \frac{13500}{110} + \frac{0.5}{0.5} + \frac{13900}{110} + \frac{0.5}{0.5} + \frac{13900}{110} + \frac{1319}{295.2} + \frac{26.9}{295.0} + \frac{60.7}{60.5} + \frac{11200}{110} + \frac{0.2}{120} + \frac{297.3}{1120} + \frac{60.9}{110} + \frac{12699}{110} + \frac{12699}{110} + \frac{1080}{11} + \frac{1080}{110} + \frac{11080}{110} + \frac{11080}{110} + \frac{11080}{110} + \frac{110}{110} + \frac{110}{2525} + \frac{100.9}{297.3} + \frac{100}{56.3} + \frac{110200}{110} + \frac{110}{110} + \frac{110}{255.2} + \frac{1100}{297.3} + \frac{110}{56.0} + \frac{110}{295.0} + \frac{110}{110} + \frac{110}{200} + \frac{110}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{10200} + \frac{11}{110} + \frac{110}{200} + \frac{110}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{10200} + \frac{11}{11} + \frac{110}{200} + \frac{110}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{10200} + \frac{11}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{10200} + \frac{11}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{10200} + \frac{11}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{200} + \frac{110}{11} + \frac{110}{255.2} + \frac{110}{297.3} + \frac{110}{56.3} + \frac{110}{10200} + \frac{11}{11} + \frac{110}{10} + $
$\frac{1}{1200} + \frac{1}{1200} + 1$
$\frac{1}{10} - \frac{1}{1040} = \frac{1}{1094} - \frac{1}{105} - 1$
$\frac{1}{10} -\frac{11460}{16994} -\frac{1397}{157} -\frac{259}{257}, \frac{63.0}{63.7} -\frac{22100}{2100} -0.8$ $\frac{10}{10} -\frac{10414}{16784} -\frac{16784}{7421} -\frac{263.1}{263.1} -\frac{64.3}{64.3} -\frac{21100}{21100} -0.5$ $\frac{14}{14} -\frac{8685}{15711} -\frac{7411}{7411} -\frac{269.6}{246.3} -\frac{64.3}{18400} -0.5$ $\frac{18}{12} -\frac{-6468}{13836} -\frac{7691}{7691} -\frac{285.0}{285.0} -\frac{63.5}{15900} -0.6$ $\frac{10}{12} -\frac{-5206}{12966} -\frac{2673}{287.5} -\frac{285.0}{63.5} -\frac{15500}{15500} -0.6$ $\frac{10}{10} -\frac{-6223}{12295} -\frac{1295}{6563} -\frac{295.7}{63.9} -\frac{14100}{1400} -0.5$ $\frac{50}{10} -\frac{-3229}{10511} -\frac{6319}{6319} -\frac{296.7}{285.0} -\frac{63.7}{11200} -0.2$ $\frac{100}{10} -\frac{-3167}{2851} -\frac{9819}{5727} -\frac{297.1}{296.7} -\frac{63.4}{62.4} -\frac{11800}{1200} -0.2$ $\frac{50}{120} -\frac{-2678}{256} -\frac{8686}{5450} -\frac{299.0}{299.0} -\frac{60.6}{60.6} -\frac{10600}{1.0} -0.3$ $\frac{110}{120} -\frac{-2104}{2256} -\frac{8256}{5609} -\frac{304.6}{58.3} -\frac{10200}{1.1}$ $\frac{120}{120} -\frac{-2524}{7541} -\frac{5252}{2277.3} -\frac{297.3}{58.0} -\frac{9530}{0.3} -\frac{61.3}{0.3}$ $\frac{100}{1.1} -\frac{12699}{1.256} -\frac{12699}{0.3} -\frac{100}{0.3}$ $\frac{100}{1.1} -\frac{100}{1.1} -$
$\frac{10}{14} - \frac{10414}{16794} \frac{16374}{7421} \frac{253.7}{257.7} \frac{63.7}{63.7} \frac{21700}{21100} \frac{1.0}{0.5}$ $\frac{14}{14} - \frac{8859}{15711} \frac{15711}{7411} \frac{269.6}{269.6} \frac{64.3}{64.5} \frac{11500}{1500} \frac{0.5}{0.5}$ $\frac{18}{22} - \frac{-6468}{6681} \frac{13039}{7107} \frac{707}{285.0} \frac{282.4}{63.5} \frac{64.3}{15900} \frac{18400}{0.3}$ $\frac{132}{22} - \frac{5206}{12966} \frac{12966}{6673} \frac{6673}{287.5} \frac{287.5}{65.0} \frac{65.0}{15500} \frac{15500}{0.8}$ $\frac{10}{40} - \frac{4623}{4623} \frac{12195}{12965} \frac{6563}{290.5} \frac{290.5}{64.1} \frac{14600}{1400} \frac{1.1}{0.5}$ $\frac{60}{6} - \frac{-3298}{-3298} \frac{10511}{1098} \frac{6308}{296.7} \frac{296.7}{63.9} \frac{14100}{100} \frac{0.5}{0.5}$ $\frac{60}{50} - \frac{-32298}{-3298} \frac{10511}{6319} \frac{6319}{296.2} \frac{296.7}{61.7} \frac{112700}{11200} \frac{0.2}{0.2}$ $\frac{80}{50} - \frac{-3023}{-32298} \frac{9142}{5721} \frac{5721}{297.1} \frac{296.4}{62.4} \frac{11800}{10600} \frac{0.7}{0.2}$ $\frac{80}{50} - \frac{-2574}{-3124} \frac{6686}{5450} \frac{5450}{299.0} \frac{296.6}{61} \frac{10600}{0.2}$ $\frac{100}{-3124} \frac{2556}{5609} \frac{5609}{304.0} \frac{58.3}{10200} \frac{10.8}{1.1}$ $\frac{110}{210} - \frac{-2524}{7541} \frac{5252}{227.3} \frac{297.3}{58.0} \frac{9530}{9530} \frac{61.3}{0.3}$ $\frac{110}{530} - \frac{2574}{550} \frac{60.9}{100} \frac{10.8}{530} \frac{10200}{1.1}$ $\frac{110}{5252} - \frac{297.8}{56.0} \frac{10.9}{9530} \frac{10.2}{0.3}$
$\frac{14}{16} - \frac{16}{16} + \frac{16}{16} + \frac{1421}{12} + \frac{263.1}{16} + \frac{66.3}{6} + \frac{21100}{16} + \frac{0.5}{16} + \frac{1100}{16} + \frac{0.5}{16} + 0.5$
$\frac{18}{22} -\frac{7655}{16499} +\frac{7514}{7514} +\frac{269.6}{276} +\frac{64.5}{64.5} +\frac{19500}{18400} +\frac{0.5}{0.3} +\frac{1969}{12986} +\frac{7691}{282.4} +\frac{276}{62.8} +\frac{17100}{17100} +\frac{0.3}{0.3} +\frac{1936}{282.4} +\frac{62.8}{62.8} +\frac{17100}{17100} +\frac{0.3}{0.3} +\frac{1298}{298.2} +\frac{64.1}{285.0} +\frac{11800}{1.1} +\frac{11800}{0.5} +\frac{11800}{0.2} +\frac{1180}{0.2} +\frac{11800}{0.2} +\frac{11800}{0.2} +\frac{1180}{0.2} +\frac{11800}{0.2} +\frac{11800}{0.2} +\frac{11800}{0.2} +\frac{11800}{0.2} +\frac{1180}{0.2} +\frac{11800}{0.2} +\frac{1180}{0.2} +\frac{11800}{0.2} +\frac{1180}{0.2} +\frac{1180}{0.2} +\frac{1180}{0.2} +\frac{1180}{0.2} +\frac{1180}{0.2} +\frac{1180}{0.2}$
$\frac{1}{22} - \frac{6668}{668} \frac{138765}{138356} \frac{7514}{7691} \frac{276 \cdot 0}{285 \cdot 0} \frac{66.3}{61.5} \frac{18400}{13000} \frac{0.3}{0.3} \\ \frac{28}{25} - \frac{5563}{5506} \frac{13039}{12986} \frac{7107}{285 \cdot 0} \frac{285 \cdot 0}{55 \cdot 0} \frac{13590}{15900} \frac{0.6}{0.6} \\ \frac{40}{40} - \frac{4623}{4523} \frac{12195}{12986} \frac{6563}{290 \cdot 5} \frac{297 \cdot 5}{64 \cdot 1} \frac{14600}{1400} \frac{1.1}{0.5} \\ \frac{50}{50} - \frac{-3239}{31298} \frac{10511}{5511} \frac{6319}{5319} \frac{296 \cdot 2}{296 \cdot 7} \frac{63 \cdot 9}{63 \cdot 9} \frac{14100}{1400} \frac{0.2}{0.5} \\ \frac{60}{60} - \frac{-3239}{31298} \frac{10511}{5511} \frac{6319}{5327} \frac{297 \cdot 1}{296 \cdot 9} \frac{60 \cdot 7}{60 \cdot 7} \frac{11200}{1200} \frac{0.2}{0.2} \\ \frac{60}{50} - \frac{-32639}{3042} \frac{8686}{5450} \frac{5450}{299 \cdot 0} \frac{296 \cdot 7}{50 \cdot 6} \frac{106600}{10} \frac{0.2}{0.2} \\ \frac{80}{50} - \frac{-3124}{2039} \frac{8686}{5450} \frac{5450}{299 \cdot 0} \frac{296 \cdot 6}{51 \cdot 1} \frac{10200}{10200} \frac{1.1}{1.5} \\ \frac{100}{120} - \frac{-2124}{2556} \frac{8039}{5609} \frac{304 \cdot 0}{56 \cdot 3} \frac{10200}{153} \frac{1.1}{10200} \frac{1.1}{1.5} \\ \frac{120}{5252} \frac{-2524}{7541} \frac{5252}{297 \cdot 3} \frac{59 \cdot 0}{58 \cdot 0} \frac{9530}{9530} \frac{0.3}{0.3} \\ \frac{100}{13} \frac{1000}{13} \frac{1000}{13$
$\frac{28}{32} - \frac{5683}{5683} \frac{13039}{13039} \frac{7107}{282.4} = \frac{62.8}{62.8} \frac{7100}{1500} \frac{0.3}{0.3}$ $\frac{32}{32} - \frac{5206}{5206} \frac{12986}{12986} \frac{6673}{673} \frac{287.5}{287.5} \frac{65.0}{55.0} \frac{15500}{0.8}$ $\frac{60}{-4233} \frac{12195}{12195} \frac{6563}{2563} \frac{290.5}{290.5} \frac{64.1}{64.1} \frac{14600}{1.1}$ $\frac{11}{60} - \frac{-3238}{10511} \frac{15319}{5319} \frac{298.2}{296.2} \frac{61.7}{63.9} \frac{11200}{1400} \frac{0.5}{0.2}$ $\frac{70}{-3628} \frac{-3623}{9142} \frac{5727}{297.1} \frac{296.2}{62.4} \frac{61.7}{11200} \frac{11200}{0.2}$ $\frac{100}{-3228} \frac{-3628}{9142} \frac{8688}{5450} \frac{299.0}{299.0} \frac{60.6}{60.6} \frac{10660}{0.9}$ $\frac{100}{-3124} \frac{8039}{8039} \frac{4461}{290.0} \frac{299.0}{63.4} \frac{60.6}{1100} \frac{110}{0.8}$ $\frac{100}{-2104} \frac{8256}{5569} \frac{5609}{304.0} \frac{58.3}{10200} \frac{110}{1.1}$ $\frac{200}{-2524} \frac{7541}{5252} \frac{297.3}{297.3} \frac{58.0}{58.0} \frac{9530}{0.3} \frac{61.3}{0.3}$ $\frac{1082}{100} \frac{11}{100} \frac{110}{100} 110$
$\frac{32}{40} - \frac{5206}{40} = \frac{13039}{12966} = \frac{6773}{285.0} = \frac{285.0}{63.5} = \frac{15900}{1500} = 0.6$ $\frac{40}{40} - \frac{4623}{4023} = \frac{12986}{11908} = \frac{6563}{296.7} = \frac{295.0}{63.9} = \frac{64.1}{14600} = \frac{11}{14}$ $\frac{60}{40} - \frac{-3298}{4029} = \frac{10511}{1511} = \frac{6319}{296.7} = \frac{261.7}{12700} = 0.2$ $\frac{70}{40} - \frac{-3167}{203} = \frac{9142}{5721} = \frac{5727}{297.1} = \frac{297.1}{62.4} = \frac{61.7}{11200} = 0.2$ $\frac{80}{400} - \frac{-3124}{200} = \frac{8638}{5450} = \frac{299.0}{299.0} = \frac{60.6}{61.6} = \frac{10600}{10200} = 0.3$ $\frac{100}{4023} = \frac{-2578}{4668} = \frac{5669}{304.0} = \frac{299.0}{56.3} = \frac{10200}{1200} = \frac{11}{1.1}$ $\frac{120}{20} - \frac{-2524}{7541} = \frac{5252}{2522} = \frac{297.3}{297.3} = \frac{58.0}{58.0} = \frac{9530}{6.3} = \frac{6.3}{1200}$ $\frac{110}{200} = \frac{2.4}{1100} = \frac{12699}{100} = \frac{12.4}{100} = \frac{1000}{100} = $
$\frac{40}{50} -\frac{4623}{-4623} \frac{12796}{12195} \frac{6673}{5563} \frac{287.5}{290.5} \frac{68.0}{64.1} \frac{15500}{14000} \frac{0.8}{1.1} \\ \frac{50}{50} -\frac{-3829}{-3829} \frac{11908}{10511} \frac{6508}{2319} \frac{296.2}{296.7} \frac{63.9}{53.9} \frac{14100}{14100} \frac{0.5}{0.5} \\ \frac{-3167}{70} \frac{9819}{9139} \frac{5727}{277.1} \frac{297.1}{297.1} \frac{62.4}{62.4} \frac{11800}{11800} \frac{0.7}{0.2} \\ \frac{-3023}{90} \frac{9142}{-2678} \frac{6686}{5450} \frac{299.0}{299.0} \frac{60.6}{60.6} \frac{10600}{0.9} \frac{0.8}{0.9} \\ \frac{100}{10} -\frac{-2104}{-2526} \frac{8256}{5609} \frac{304.0}{304.0} \frac{58.3}{58.0} \frac{10200}{9530} \frac{1.1}{0.3} \\ \frac{20}{7} -\frac{2524}{7541} \frac{5252}{297.3} \frac{111}{58.0} \frac{9111}{9530} \frac{1.2}{0.3} \\ \frac{100}{11} \frac{110}{5252} \frac{100}{297.3} \frac{100}{58.0} \frac{111}{9530} \frac{1.2}{0.3} \\ \frac{100}{11} \frac{110}{5252} \frac{110}{297.3} \frac{110}{58.0} \frac{110}{9530} \frac{1.1}{0.3} \\ \frac{100}{11} \frac{110}{5252} \frac{100}{297.3} \frac{100}{58.0} \frac{1.1}{9530} \frac{1.1}{0.3} \\ \frac{100}{11} \frac{110}{5252} \frac{100}{297.3} \frac{100}{58.0} \frac{110}{9530} \frac{1.1}{0.3} \\ \frac{100}{11} \frac{100}{11} \frac{110}{5252} \frac{100}{297.3} \frac{100}{58.0} \frac{100}{9530} \frac{100}{1.3} \\ \frac{100}{11} \frac{100}{10} 10$
$\frac{50}{60} = -\frac{3629}{12906} = \frac{12439}{1906} = \frac{6563}{296.7} = \frac{290.5}{63.9} = \frac{14600}{1400} = \frac{1.1}{1.5}$ $\frac{60}{60} = -\frac{3298}{10511} = \frac{6519}{10511} = \frac{296.2}{61.7} = \frac{61.7}{12700} = \frac{12700}{0.2}$ $\frac{60}{10} = -\frac{3023}{122} = \frac{912}{5727} = \frac{297.1}{297.1} = \frac{62.4}{11200} = \frac{11200}{0.2}$ $\frac{60}{10} = -\frac{-2678}{1226} = \frac{8686}{5450} = \frac{296.9}{295.0} = \frac{60.6}{10600} = \frac{10600}{0.9}$ $\frac{10}{10} = -\frac{-2104}{1226} = \frac{8256}{5609} = \frac{300.0}{300.0} = \frac{51.4}{51.1} = \frac{9710}{0.8}$ $\frac{10}{20} = -\frac{-2524}{7541} = \frac{5252}{2522} = \frac{297.3}{297.3} = \frac{58.0}{9530} = \frac{9530}{0.3} = \frac{61.3}{0.3}$ $NCLOR: dec = 297.8 inc = 60.9 int = 12699 mad = 2.4$
$\frac{60}{7} = -\frac{1298}{1282} = \frac{11908}{10511} = \frac{6508}{5119} = \frac{296.7}{296.2} = \frac{61.7}{112700} = \frac{14100}{0.2} = \frac{0.5}{0.5}$ $\frac{70}{7} = -\frac{3167}{29819} = \frac{9619}{5727} = \frac{297.1}{297.1} = \frac{62.4}{11800} = \frac{11800}{0.7} = \frac{0.5}{0.7}$ $\frac{80}{7} = -\frac{2678}{200} = \frac{8688}{5450} = \frac{5450}{299.0} = \frac{296.7}{60.6} = \frac{106600}{10600} = \frac{0.9}{0.2}$ $\frac{100}{7} = -\frac{2104}{2556} = \frac{8568}{5252} = \frac{390.0}{297.3} = \frac{63.4}{58.0} = \frac{9710}{0.3} = \frac{0.8}{0.3}$ $\frac{100}{7} = -\frac{2104}{2524} = \frac{8256}{5252} = \frac{297.3}{297.3} = \frac{11200}{58.0} = \frac{0.8}{0.3}$ $\frac{100}{7} = -\frac{2524}{7541} = \frac{50.9}{5252} = \frac{11200}{9530} = \frac{12699}{0.3} = \frac{12699}{0.3} = \frac{12699}{0.3} = \frac{1000}{0.3} = \frac{1000}{0.3}$ $\frac{1000}{100} = \frac{1000}{100} = \frac{1000}{0.3} = 10$
$\frac{70}{80} = -\frac{3167}{-3023} + \frac{10511}{9142} + \frac{5727}{297.1} + \frac{297.1}{62.4} + \frac{11800}{11800} + \frac{0.7}{0.2} + \frac{3023}{292.2} + \frac{5727}{297.1} + \frac{297.1}{62.4} + \frac{11800}{11200} + \frac{0.7}{0.2} + \frac{1000}{200} + \frac{0.7}{200} + \frac{11200}{200} + \frac{0.7}{200} + \frac{1000}{200} + \frac{0.7}{200} + \frac{0.7}{200$
$\frac{10}{50} = -3023 - 9813 - 5727 - 297.1 - 62.4 - 11800 - 0.7 - 11200 - 2.7 - 2678 - 8688 - 5450 - 299.0 - 60.6 - 11200 - 0.3 - 1124 - 8039 - 4461 - 290.0 - 63.4 - 9710 - 0.8 - 1102104 - 8256 - 5609 - 304.0 - 58.3 - 10200 - 1.1 - 1202524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 11025242524 - 7541 - 5252 - 297.3 - 58.0 - 9530 - 0.3 - 1102524254 - $
$\frac{30}{100} = -\frac{100}{1100} = \frac{9142}{1100} = \frac{5721}{1100} = \frac{296}{9} = \frac{9}{60.7} = \frac{11200}{10600} = \frac{0.2}{0.9} = \frac{1000}{0.9} = \frac{0.2}{0.00} = \frac{1000}{0.9} = \frac{0.2}{0.00} = \frac{0.2}$
$\frac{300}{-3124} = \frac{3688}{8039} = \frac{5450}{4461} = \frac{299.0}{299.0} = \frac{60.6}{63.4} = \frac{10600}{9710} = \frac{0.9}{0.8}$ $\frac{110}{-2104} = \frac{2256}{8256} = \frac{5699}{304.0} = \frac{304.0}{58.3} = \frac{10200}{1.1} = \frac{11}{10200} = \frac{11}{1000}$ $\frac{120}{-2524} = \frac{7541}{7541} = \frac{5252}{297.3} = \frac{297.3}{58.0} = \frac{9530}{9530} = \frac{61.3}{61.3}$ $\frac{1000}{1000} = \frac{1000}{1000} = \frac{1000}{10$
$\frac{100}{-2104} = \frac{8039}{256} = \frac{4461}{5609} = \frac{290.0}{304.0} = \frac{61.4}{56.3} = \frac{9710}{10200} = \frac{0.8}{1.1} = \frac{10200}{1.1} = \frac{10200}{1.20} = \frac{1000}{1.20} = \frac{1000}{1.2$
$\frac{100}{220} - \frac{-2104}{-2524} = \frac{8256}{7541} = \frac{5609}{5252} = \frac{304.0}{297.3} = \frac{10200}{9530} = \frac{1.1}{0.3}$ $\frac{10200}{0.3} = \frac{1.1}{0.3}$ $\frac{10200}{0.3} = \frac{1.1}{0.3}$ $\frac{10200}{0.3} = \frac{1.1}{0.3}$ $\frac{10200}{0.3} = \frac{1.1}{0.3}$
NCLOR: dec = 297.8 inc = 60.9 int = 12699 mad = 2.4 th/tc 108
NCLOR: dec = 297.8 inc = 60.9 int = 12699 mad = 2.4 th/tc 10800
th / tc 108 000







STRP	A	B	c	DECte	INCtc	INTENS	ERR	DATE	TIME	
0	-17686	11700	1							127
2	-13864	12000	-1004	238.7	64.1	17400	2.1			
4	-13166	12070	-1580	239.6	63.6	19000	1.5			
6	-17197	14740	-1438	243.0	64.2	16500	1.2			
Ř	-11400	12748	~1467	246.2	65.3	17700	1.0			
10	-11600	12403	-1350	248.8	65.7	16900	1.0			
14	-10460	12032	-1348	252.1	66.6	16000	0.6			
18	-9404	11259	-947	255.8	66.0	14700	1.1			
22	-8470	10718	-1031	258.9	67.1	13700	1.0			
28 +	-7471	10333	-1117	264.7	68.4	12800	1.5			
17 +	-6801	9956	-1020	268.8	68.5	12100	1.0			
40 .	-6421	9138	-842	267.1	67.8	11200	0.9			
50 -	-5978	8354	-755	265.8	67.6	10300	0.9			
40 .	-5005	7680	-493	272.6	66.9	9180	0.6			
20 .	- 6293	6749	-578	274.0	68.0	8020	1.1			
70 -	-3844	5794	-314	271.6	66.4	6960	0.9			
00 -	-3297	5444	-912	277.1	72.1	6430	1.5			
30 -	-2426	4459	-491	284.8	69.4	5100	0.9			
100 •	-1938	4143	~486	294.0	69.3	4600	1.0			
110 -	-2685	4014	-1089	265.9	76.3	4950	0.8			
	-		_							





CP	A	В	c	DECtc	INCtc	INTENS	ERR	DATE	TIME	
0	-17788	1								
2	-12122	10030	2299	250.5	55.6	19400	2.0			
4	-13173	15314	+226	241.0	61.3	20200	0.7			
6	-11390	14977	-81	244.2	61.9	19300	0.9			
A	-10070	14459	-80	245.8	62.4	18400	0.7			
	-10108	13673	185	249.2	62.4	17100	1.1			
4	-3623	12906	204	249.7	62.4	16100	0.9			
	-6453	12391	102	254.0	63.8	15000	1.4			
	-7879	11814	14	254.9	64.3	14200	0.9			
	-7087	11254	-27	258.0	65.1	13300	0.6			
	-6318	10550	246	261.9	64.4	12300	0.6			
10	-5633	9907	272	264.9	64.5	11400	0.3			
	-\$377	8664	242	260.1	63.9	10200	0.7			
	-4084	7725	612	270.1	62.4	8760	0.8			
	-3798	7161	273	269.0	64.4	8110	1.3			
-	-3486	5947	479	2 64 .B	61.8	6910	Q. 8			
	-2797	5734	74	273.0	66.0	6380	0.7			
	-2375	4908	554	274.9	60.9	5480	0.7			
	-2291	4372	207	269.8	64.0	4940 -	0.5			
	-2240	4300	-126	268.2	67.8	4850	0.9			
20 •	-1900	2978	518	263.2	57.0	3570	1.1			
								•		





Marias Pass 3

Specimens:

115, 117, 118, 119, 120

In-situ:

Dec=12.80°, Inc=75.15°

α₉₅=15.91°

k=24.07

δ₆₃=14.81°

Structurally corrected:

Dec=37.70°, Inc=46.92°

α₉₅=15.96°

k=23.94

δ₆₃=14.85°

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Step	*	8	c	DECCC	INCte	INTENS	ERR	DATE	TIME	
0	-7584	420								1
2 •	-3346	447	-238	204.3	34.6	2630	0.5			
4 .	-3140	577	-423	204.6	38.4	3250	1.6			
6 •	- 25 40	203	-377	203.6	37.5	2900	1.6			
Å •	- 2265	639	-349	202.2	36.4	2610	1.8			
30 +	-6203	404	-300	200.0	33.9	2300	1.5			
14 *	-2039	336	-290	200.8	35.4	2100	2.0			
10 .	-1/11	103	-353	193.8	33.8	1750	2.0			
33 #	-136/	99	-165	198.8	31.4	1600	1.5			
20 0	-1328	108	+273	194.6	34.8	1360	1.2			
20 -	-1196	16	-184	194.7	30.0	1210	1.4			
34 -	-1107	~87	-6	198.6	20.7	1110	1.5			
	-925	25	45	205.4	23.4	926	2.7			
50 -	-772	-112	33	198.4	16.1	761	2.4			
30 .	-867	-121	0	195.1	15.9	678	2.1			
/0 -	-725	78	-47	203.2	31.1	731	2.1			
80 .	-561	111	30	212.9	30.3	573	15			
90 ·	-630	59	-114	196.4	34.5	643	0.5			
*	-456	-154	-85	181.6	15.1	489.	5.7			
									•	
					••••••	**************	*******		*	
INCLOR:	dec =	201.2	inc =	34.9	int :	= 3250	mad	= 5.4		



STEP	λ	8	c	DECto	INCLC	INTENS	ERR	DATE	TIME	
										133
o	-5182	16	-1445	229.2	55.5	5380	1 2			
2 •	-5295	662	-950	220 8	48 5	\$430				
6 •	-4386	639	+779	220 #	47 3	3420	4.3			
10 •	-3609	519	-736	222.0	47 6	1000	4.0			
16 •	-2786	386	-745	217 6		3720	2.7			
22 +	-2784	302		227.9	40.0	2910	1.9			
28 •	-2041	347	-375	447.4	48.2	2380	2.3			
36 *	-1940	347	-616	230.5	46.4	2160	1.7			
	.1.40	723	-475	228.3	47.5	1830	1.5			
	-1413	140	-484	234.1	50.1	1500	2.0			
52	-1147	-198	-396	237.9	64.6	1230	1.4			
60	-1009	131	-362	235.1	48.4	1080	1 5			
70	-956	14	+235	226.1	54.9	985	1 6			

INCLOR: dec = 223.4 inc = 47.9 int = 5419 mad = 2.5

STEP	*	В	с	DECLO	INCLC	INTENS	ERR	DATE	TINE	****
0 2 * 8 * 14 * 20 * 36 * 52 * 62 * 72 * 87 102	-5617 -5689 -4621 -3671 -2996 -1947 -1755 -1251 -1092 -872 -555 -486	1019 1244 667 622 603 533 535 436 355 205 78 72 226	-602 -737 -894 -900 -916 +821 -644 -570 -396 -308 -308 -352 -41	216.0 218.2 229.3 234.2 236.7 235.9 238.0 243.3 239.3 241.5 260.0 215.4	55.2 53.3 54.7 55.5 52.0 49.4 50.5 47.8 50.5 47.8 58.5 58.5 51.0 40.5	5740 5870 4790 3830 2590 2130 1920 1420 1420 180 928 661 538	1.5 1.5 1.2 1.6 2.5 1.2 2.5 1.2 2.4 1.4 3.1 0.3 0.5 1.5			134

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STEP	▲ .	8	c	DECto	INCte	INTENS	ERR	DATE	TIME	
0										13
2	- 7182	-1232	-6814	271.1	55.7	11500	0.8			
Ĩ.	-0/01	-269	-4428	251.7	56.8	9820	1.8			
ř.	-0381	25	-3362	242.1	56.8	9030	1.4			
Å	-//86	283	-2684	236.4	55.6	8240	1.4			
10	-7260	349	-2043	230.5	55.4	7550	1.6			
12	-6825	434	-1832	229.2	54.6	7080	1 5			
14 .	-6441	392	-1480	225.6	54.8	6620	1 6			
14 4	-6123	536	-1381	225.1	53.3	6300	2.9			
10 -	-5751	415	-1249	224.4	54.2	5900	2.0			
10 -	+5530	356	-1245	225.1	54.6	5690	2.0			
20 -	-5176	509	-1254	226.6	52.7	5350	4.0			
	-4993	418	-1146	225.5	53.5	5140	1.5			
28	-4584	494	-1057	225.6	52.2	4730	4.3			
32 *	-4268	378	-910	224.0	53.3	47.50	1.0			
36 *	-4019	368	-769	222.0	52 B	4350	2.0			
40 *	-3759	399	-728	222 3	52.2	3860	1.6			
46 •	-3435	396	-657	222 1	51 7	3650	2.5			
52 *	-3198	345	-737	776 6	83 5	3520	1.9			
58 *	~2696	433	-787	220.2	40 3	3300-	2.0			
64 *	-2619	389	-474	230.3	49.3	2840	2.5			
70 +	-2210	403	-654	220 6	40 1	2690	2.4			
78 •	-2167	442	-101	230.0	40.1	2340	1.9			
86 *	-1838	248	-131	210.1	60.2	2220	1.8			
94 *	-1270	102	-444	217.0	50.3	1870	1.2			
102 -	-1572	144	*400	719.8	53.6	1290	1.5			
112 *	-201	374	-65	208.3	51.8	1580	0.7			
-122-4			-03	216.6	33.1	775	1.7			
	-687	73	-249	236.8	47+7 51.8	751 734	1+0 0.8	****		••••••
INCLOR:	dec =	224.5	inc =	52.8	int	= 6300	mad	= 2.6		

