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PETROLOGY OF THE
MOUNT SHIELDS FORMATION
(BELT SUPERGROUP) WESTERN MONTANA
NORTHERN IDAHO

By

David A. Barlow

B.S., West Virginia University, 1981

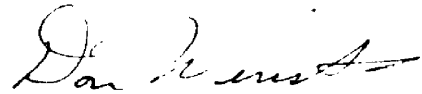
Presented in partial fulfillment of the requirements for the degree of

Master of Science

UNIVERSITY OF MONTANA

1983

Approved By:



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Barlow, David A., M.S., August 1983

Geology

Petrology of the Mount Shields Formation
(Belt Supergroup)
Western Montana and Northern Idaho (140 pp.)

Director: Dr. Donald Winston



The detrital mineralogy of the Mount Shields Formation is characterized by a moderately sorted, subangular to subrounded fine-grained population (62 to 250 microns) and a poorly sorted, subrounded to well rounded coarse-grained population (25 to 2000 microns). The aerial distribution of grain types reflect source terranes located south, southwest, northeast and north of the Belt basin. The southern terrane contained predominantly granitic rocks and associated igneous injected terranes. Very fine to medium-and coarse-grained chert-bearing quartz sandstone also existed south of the basin. The southwestern terrane contained coarse, chert-bearing quartz sandstone and metaquartzite. Granitic rocks were subordinate. Granitic rocks and associated igneous injected terranes, and coarse, chert-bearing quartz sandstone were exposed northeast of the basin. A very fine to fine-grained chert-bearing quartz sandstone, perhaps the upper Snowlip and lower Shepard, provided grains and rock fragments from the north. Volcanic rock fragments were reworked from the lower Shepard and/or eroded directly from the Purcell lava exposed north of the basin. The composition of the four source terranes remained relatively constant throughout deposition of the sediments which comprise the Mount Shields Formation. Abrupt increase in sandstone content from the Mount Shields I to the Mount Shields II reflect an episode of rapid basin subsidence relative to the surrounding source terranes.

Inferred weathering products include smectite and perhaps hematite and chlorite. Carbonate cements in the lower and upper Mount Shields Formation were precipitated in a subaqueous environment. Diagenetic hematite formed from oxidation of biotite, magnetite and ilmenite. Smectite was converted to illite and chlorite and eventually to 2M illite (200°C to 350°C) providing silica for quartz cement growth. Degredation of potassium feldspar and perhaps biotite supplied the constituents necessary for potassium feldspar precipitation.

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INTRODUCTION

Sedimentologic Setting

The Mount Shields Formation is comprised of sediments which were deposited on alluvial aprons in the Middle Proterozoic Belt basin of western Montana, northern Idaho, eastern Washington, and British Columbia (Figs. 1 and 2). These sediments are part of an enormously thick package of conglomerate and coarse, crossbedded quartzite, red, green and black argillite, and dolomite which were deposited in braided stream channels, on sheet wash fans, subaerial and subaqueous mudflats, and shallow water carbonate banks respectively. The Mount Shields Formation consists of red argillite and fine, horizontally laminated quartzite with lesser thicknesses of crossbedded quartzite.

Near Alberton, Montana, the Belt Supergroup reaches a thickness over 20 kilometers (Harrison, 1972). The great thickness of the Belt Supergroup sediments has been attributed to their deposition on slowly sinking tectonic blocks, bounded by high-angle faults (Winston et al., 1982; Harrison et al., 1974). The thickness of these sediments, and more specifically of the Mount Shields sediments, varies regionally and is a reflection of differential subsidence of these tectonic blocks.

Slover (1982) analyzed in detail the fining upward sequences in the lower part of the Mount Shields Formation and proposed that they result from cyclic alternation of abrupt rainy periods followed by gradual increased in aridity. Her thesis requires that the Belt basin was internally drained. Slover's findings support a similar conclusion proposed by Grotzinger (1981).

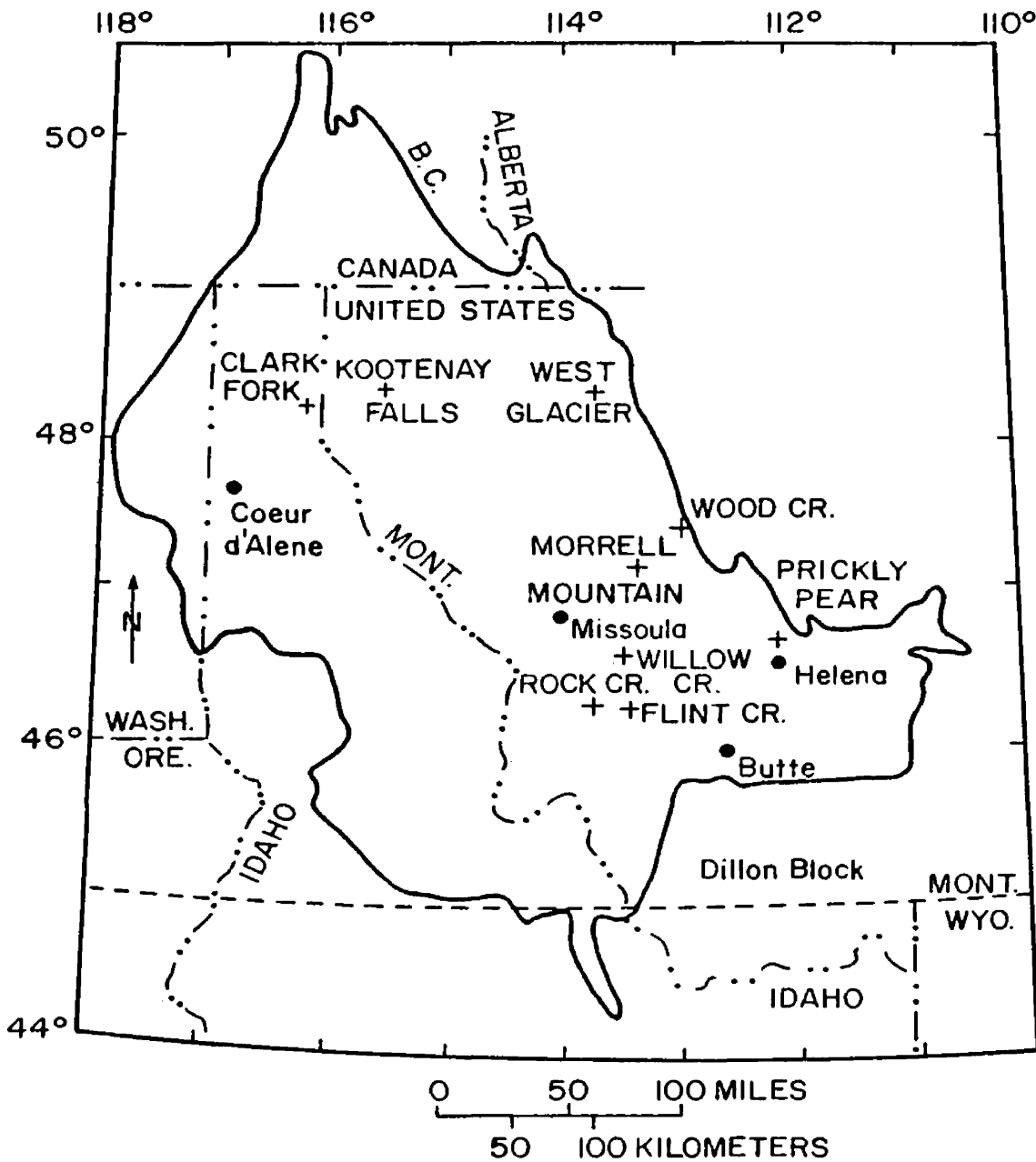
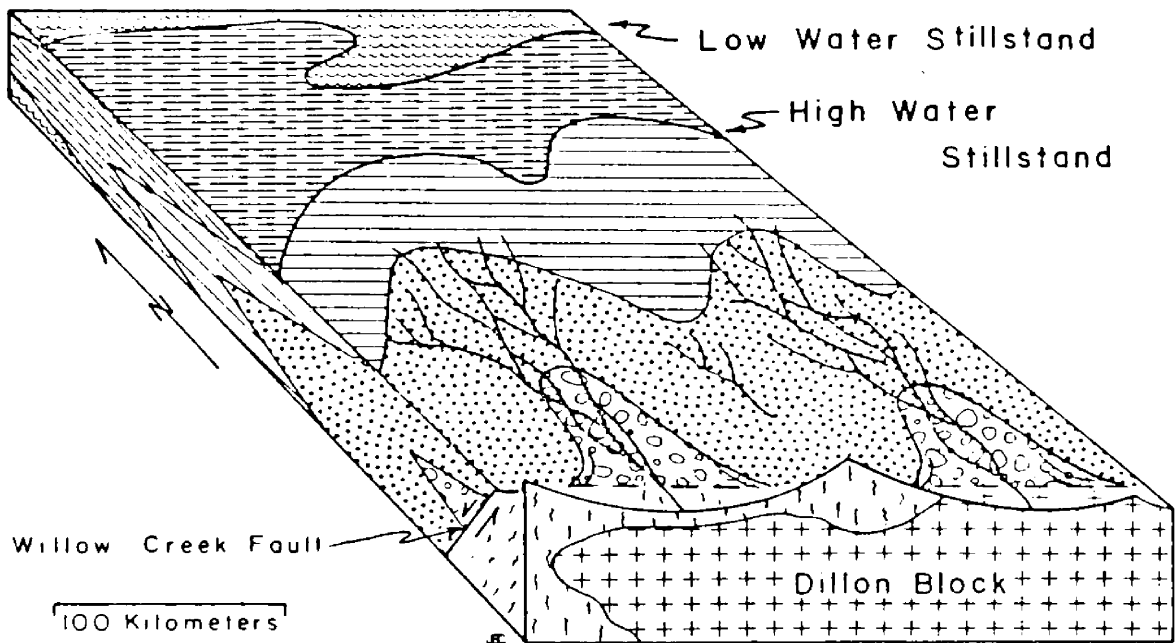


Figure 1. Index map outlining the Belt Basin and locating measured sections (modified after Harrison, 1972).








-  Conglomeratic Rock Type
-  Coarse, Crossbedded Rock Type
-  Fine, Horizontally Laminated Rock Type
-  Red Argillite Rock Type
-  Green Argillite Rock Type

Figure 2. Block diagram showing the facies and corresponding rock types of Belt alluvial fans (Winston, 1978).

Stratigraphic Setting

Winston et al. (1982) have constructed stratigraphic columns for the different tectonic blocks they have proposed for the Belt basin (Fig. 3). These structural blocks are discussed more fully in the following section. The importance of these stratigraphic relationships and their tectonic implications will become evident to the reader when the results of this study are discussed and interpreted.

The Mount Shields Formation was subdivided into four units in the Jocko Mountains, Montana. Winston and Jacob (1977) have informally named them: 1) Mount Shields I (interbedded fine-grained feldspathic quartzite and red argillite), 2) Mount Shields II (fine- and medium-grained feldspathic quartzite), 3) Mount Shields III (red argillite and abundant salt casts), and 4) Mount Shields IV (green argillite). The Mount Shields Formation is correlative with the Striped Peak Formation of the Coeur d'Alene district in northern Idaho (Fig. 3). At Clark Fork, Idaho (Harrison and Jobin, 1963) the Mount Shields I and II are correlative with the Striped Peak I; the Mount Shields III correlates with part of the Striped Peak II and III (Winston, 1977). In the northern, eastern and western sections of the Mount Shields Formation a stromatolite bed or calcareous beds mark the base of the Mount Shields III, and form an important stratigraphic marker. Extending below the stromatolite bed and the calcareous beds is a coarse-grained interval of the Mount Shields II ranging from a few centimeters thick at the northwestern sections to several meters thick in the south. These coarse grained beds occur in all sections of this

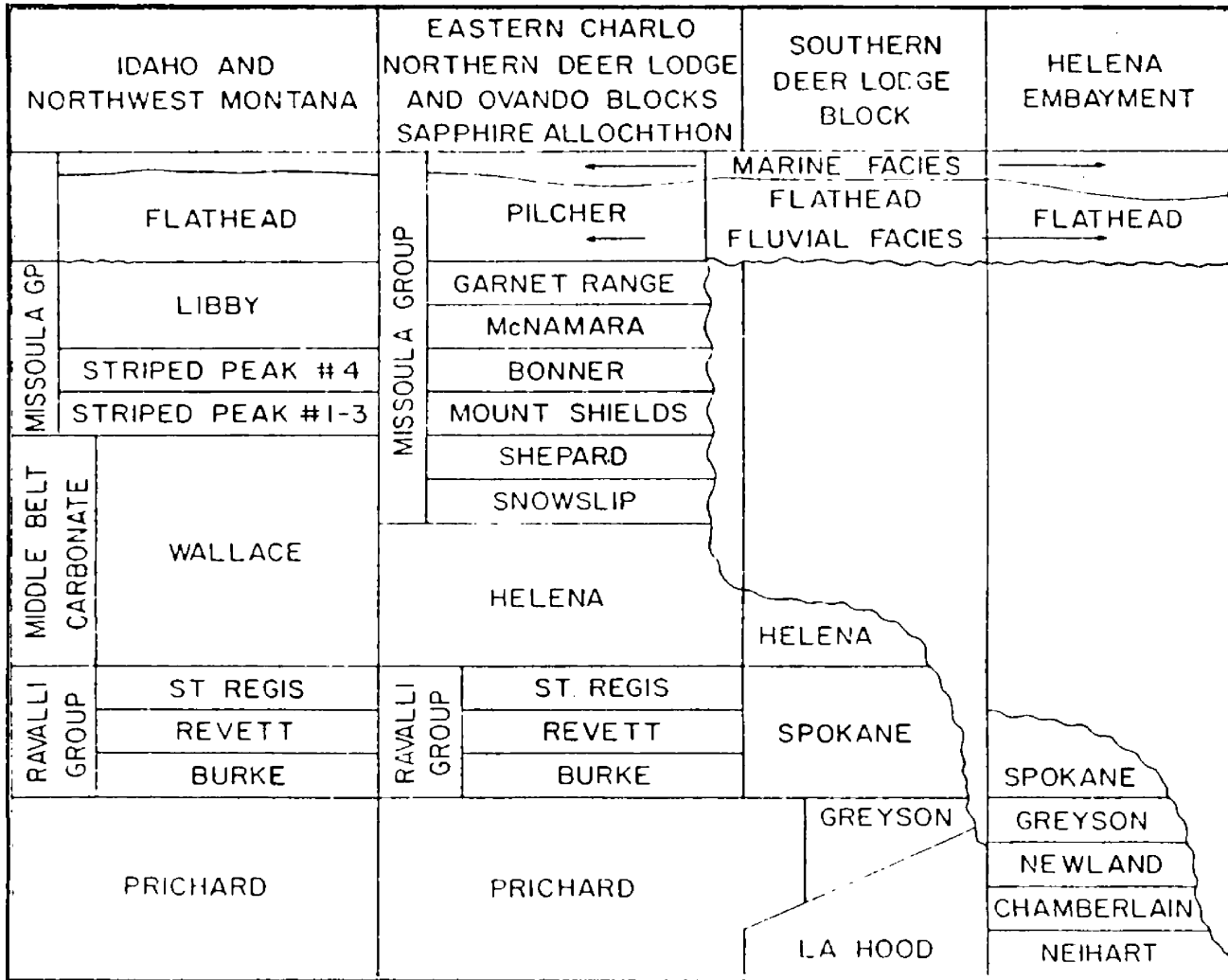


Figure 3. Stratigraphic sections of the Belt Supergroup for each of the proposed tectonic blocks of the Belt basin (Winston et al., 1982).

study except at Clark Fork, Idaho where the grains are medium. The significance of this course interval will be discussed in a following section.

Structural Setting

As mentioned above, the Belt Supergroup sediments were deposited in a block faulted basin. Winston et al. (1982) have proposed three zones of high-angle block faults which they call lines. These lines are in addition to the previously described east-west trending Perry line which bounds the south edge of the Belt basin and the north edge of the Dillon block (Fig. 4). The Perry line and the nearly east-west trending Greenhorn and Jocko lines are intersected on the east by the northwest-southeast trending Townsend line which marks the eastern boundary of the Deer Lodge, Ovando, and Charlo blocks, and the western end of the Helena Embayment (Fig. 4). Winston et al. (1982) have proposed that the western boundary of these blocks be marked by a western thrust belt. Harrison et al. (1972) have proposed that the Hope fault (Clark Fork, Idaho) coincides with a Proterozoic fault zone, and McMannis (1963) studied the Willow Creek Precambrian fault which lies along the Perry line.

The tectonic blocks proposed by Winston et al. (1982) not only controlled the structure of the Belt basin but may have acted as buttresses and ramps for Late Cretaceous thrusting. Winston et al. (1982), as well as other geologists (Harrison et al., 1974) speculate that the emplacement of other structures, such as the Idaho and Boulder

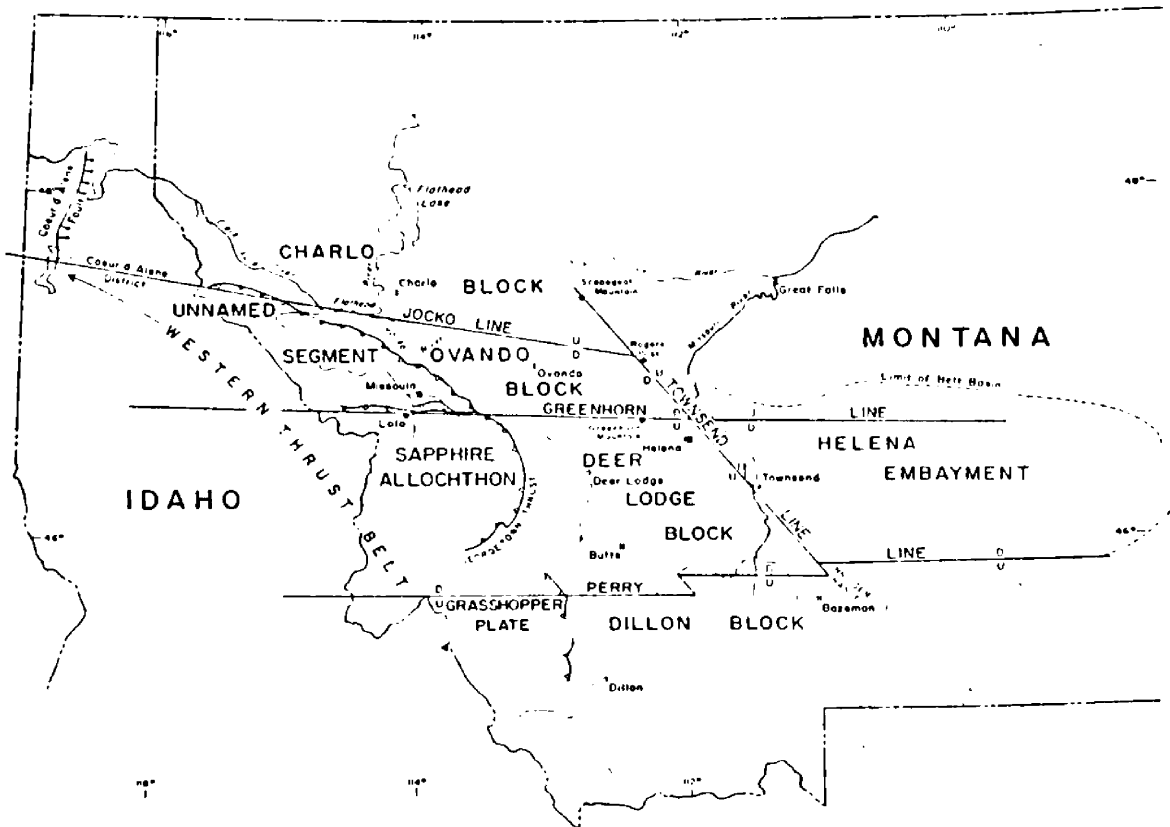


Figure 4. Proposed Precambrian tectonic blocks of western Montana and northern Idaho (Winston et al., 1982).

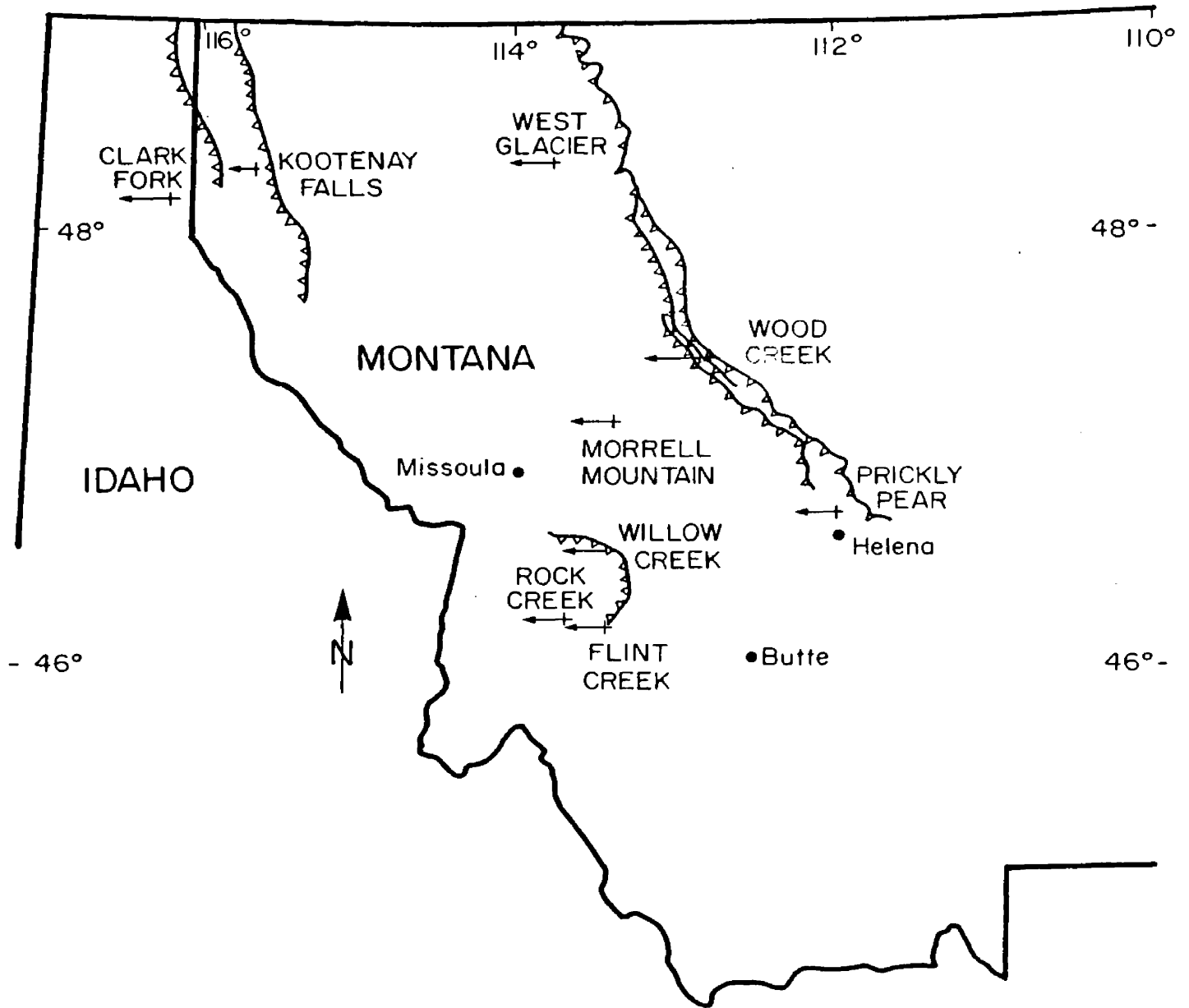


Figure 5. The spatial relationships of the nine measured sections with the major thrust faults of western Montana. The arrows indicate the direction of the original locations of these sections (modified after Winston et al. 1982).

batholiths and Late Cretaceous extensional faults, was controlled by inhomogeneities and faults in the Precambrian basement.

Figure 5 shows the locations of the nine measured sections and their relationships with some of the major thrust faults in the region. The arrows extending west from each section indicate the direction but not necessarily the distance to their original location. The present east-west relationships of the measured sections are the same as the original relationships that is, the Clark Fork section is the furthest west today and was probably the western-most section prior to thrusting. The original north-south relationships are also the same as today. The mineralogical and textural trends presented in this study are therefore valid for interpretation of source terrane locations in their relation to the measured sections.

Purpose of Study

Previous workers (Fenton and Fenton, 1937; Ross, 1963; Harrison, 1972; and others) have proposed several source terranes for the sediments of the Belt Supergroup (including the Mount Shields Formation) on the basis of regional facies relationships. One source area is northeast of the basin, another is south or southwest of the basin. These workers, however, have not systematically analyzed the mineralogy and textures in the Belt rocks for further evidence of source rock locations and compositions. Harrison and Grimes (1970) examined mineralogy and geochemistry of the Belt Supergroup but only sampled the rocks in the Mission Mountains, Montana and Pend

Oreille, Idaho. This thesis is a systematic study of the mineralogy and textures of the lower part of the Mount Shields Formation throughout most of the Belt basin and will therefore propose solutions to several interesting questions not necessarily answered by previous workers. Some of these questions are:

1. Where were the source terranes for the Mount Shields Formation sediments?
2. Did these source rocks have similar or different compositions and textures?
3. Was there a significant change in source terranes during deposition of the Mount Shields sediments?
4. What do the mineralogical and textural trends in the Mount Shields Formation imply about the tectonic history during deposition of these sediments?

The diagenetic and low grade metamorphic features of the Mount Shields Formation are also examined in this study to understand better the alteration processes which have affected these rocks.

The Mount Shields Formation was chosen for this study because it is well exposed throughout the basin, it is only slightly metamorphosed, and the sand-sized grains lend themselves to examination under the petrographic microscope.

Methods of Study

Stratigraphic sections of the Mount Shields Formation, three measured by Winston (personal communication), one measured by McGill and Sommers (1967), two measured by Slover (1982), and three measured by myself, were representatively sampled every 20 to 50 feet; approximately 130 samples were collected. The nine measured sections are: Rock Creek (includes Flint Creek), Willow Creek, Morrell Mountain, Prickly Pear, Wood Creek, West Glacier, Kootenai Falls, and Clark Fork (Fig. 1).

Thin sections of the samples collected were stained with sodium cobaltinitrite for potassium feldspar identification, and examined under a petrographic microscope to characterize rock compositions and textures. Sample grain sizes and percentages were visually estimated in thin section, and the histogram in Appendix III is based on grain size range frequencies. Diagenetic features, such as grain overgrowths, were also studied under the petrographic microscope.

X-ray diffraction methods were used to identify the general clay mineralogy, including illite polytypes. The proportion of 2M illite polytype to total illite was estimated using Velde and Hower's (1963) method. The relative intensities of the 3.74A and the 2.58A illite peaks were determined from randomly oriented samples using the method described by Schultz (1964).

Previous and Related Work

There are no previous petrographic studies on the Mount Shields Formation. Winston (personal communication) has examined thin sections of the Mount Shields sandstone but has not systematically analyzed them. Hernden (personal communication) is presently studying the petrology of the Revett Formation which is similar depositionally to the Mount Shields Formation but has undergone greater diagenetic and metamorphic alteration.

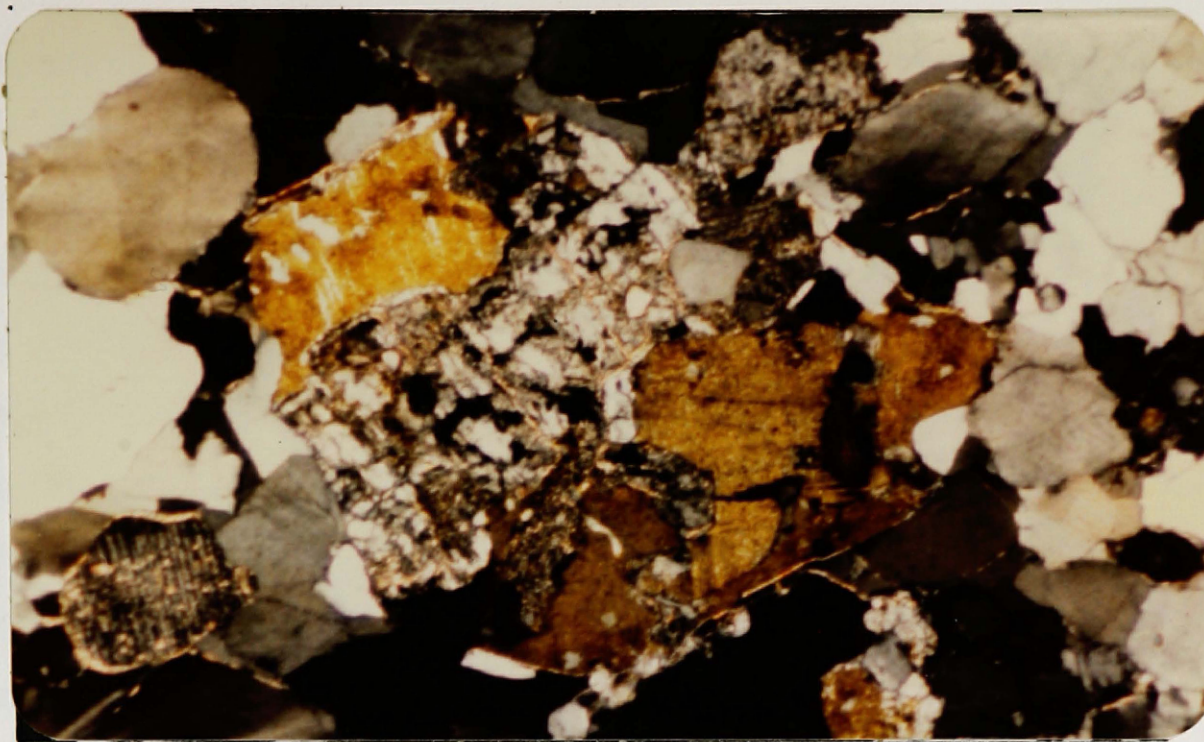
Burial metamorphism in the Belt Supergroup has been studied by Maxwell and Hower (1967) using illite polytypes. They collected samples ranging from the lower Belt to the upper Missoula Group from sections in the Little Belt Mountains and Glacier Park, Montana and at Clark Fork, Idaho. Eslinger and Savin (1973) also studied burial metamorphism of Belt rocks at Glacier Park and concluded from their oxygen isotope studies that metamorphic grade increases down section, and confirmed that illite polytypes can be used as a measure of geothermal grade in Belt rocks. Eslinger and Sellars (1981) examined evidence for conversion of illite from smectite during burial metamorphism of the Belt Supergroup at Clark Fork, Idaho. These studies provide important mineralogic and chemical data, and significant diagenetic interpretations which support the results and interpretations given in this paper for the diagenetic and low grade metamorphic processes in the Mount Shields Formation.

MINERALOGY OF DETRITAL GRAINS

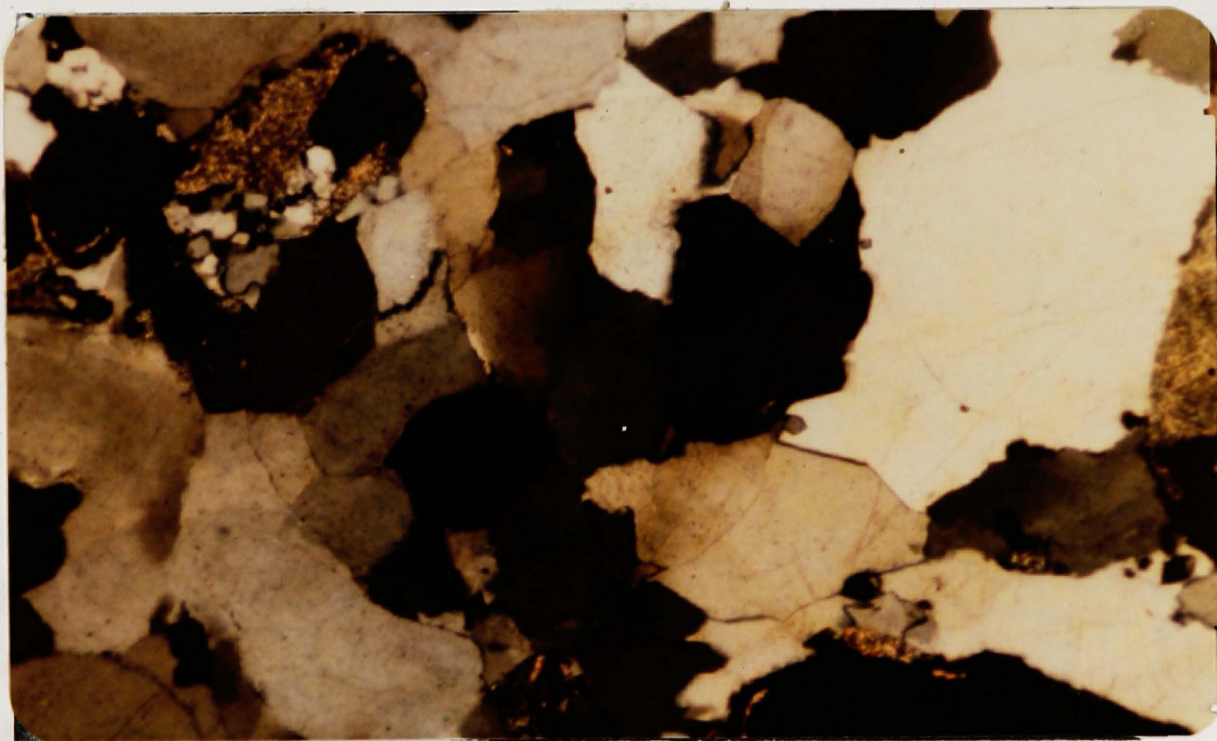
The following grain descriptions apply to the entire Mount Shields Formation except where specified. Grain percentages are tabulated in Appendix I, and Appendix II gives statistical support for some of the grain distribution discussed.

Granite Fragments

Description. Granite fragments in the Mount Shields Formation occur in poorly sorted beds of the coarse-grained population (greater than 250 microns; Appendix III) (Fig. 6a). These fragments range from 350 microns (medium sand) to 2000 microns (very coarse sand), and consist of grains 125 to 1000 microns in diameter (fine- to medium-grained granite). A few pebbles as large as 5000 and 9000 microns with 2000 to 4000 micron (medium granite) sized grains occur at Rock Creek. The maximum size and the angularity of the granite clasts at approximately the same stratigraphic level (top of the Mount Shields II) decrease in a general northward direction from Rock Creek (9000 microns; subrounded), Willow Creek (2000 microns; subrounded) and Prickly Pear (2000 microns; rounded) to Morrell Mountain (710 microns; round to well rounded) and Wood Creek (1410 microns; round to well rounded). Although the fine to medium granite fragments at West Glacier are also smaller (710 microns) than those in the south they are equally angular (subrounded to rounded) as the southern granite fragments. Granite fragments are most abundant (more than one percent) at Prickly Pear, Wood Creek and Rock Creek, and less common (less than



(a)



(b)

Figure 6. Partly sericitized granite fragment from Rock Creek (35X) (K-feldspar stained yellow) (a). Quartz sandstone fragment from the upper Mount Shields II at Prickly Pear (42X) (b).

one percent) at all other sections except Kootenai Falls and Clark Fork where these grains are absent.

Clear quartz occurs in each granite fragment. These quartz grains have straight extinction. Approximately 75 percent of the granite clasts contain orthoclase. At the southern localities (Rock Creek, Willow Creek, Prickly Pear) granite grains containing sericitized orthoclase are commonly mixed with those which have fresh orthoclase; the orthoclase in granite fragments at other sections is mostly sericitized. Five percent or less of the granite clasts contain fresh microcline. About 10 percent of the granite grains in the Mount Shields Formation contain microperthite; they occur primarily at Rock Creek and Willow Creek. Fresh and sericitized antiperthite occur only at Willow Creek. Twinned plagioclase (albite-oligoclase) occurs in some granite clasts at Rock Creek (approximately one percent), and is sericitized. Biotite flakes in the granite casts are fresh and occur in about 3 to 5 percent of the granite fragments (only at Rock Creek and Prickly Pear). One granite grain at Willow Creek contains a clear, brown and green zoned tourmaline crystal. These zones parallel the c-axis of the crystal. Green tourmaline also occurs in other granite fragments at Willow Creek. Two granite clasts at Rock Creek have muscovite flakes and one grain has zircon.

Interpretation. Fine- and medium-grained granite supplied the granite fragments to the Mount Shields Formation. The decrease in size and angularity of the grains to the north from Rock Creek, Willow Creek and Prickly Pear suggests that they were transported from a source

terrane in the south. Another granitic source in the northeast may be implied by the subrounded to rounded, coarse granite fragments at West Glacier.

Antiperthite is common at Willow Creek but absent at West Glacier, thereby implying that a granitic source nearby Willow Creek may have been more sodic than the northeastern granite which supplied detritus to the West Glacier area. Sericitic alteration probably resulted from burial metamorphism.

Sandstone Rock Fragments

Description. Fine- to very coarse grained (125 to 2000 microns) sandstone rock fragments in the Mount Shields Formation are subrounded to rounded, and contain moderately sorted round grains of very fine to fine-grained (62 to 250 microns) or fine to medium (177 to 350 microns) (only at Rock Creek) quartz sand cemented in clear quartz with straight extinction (Fig. 6b). One fragment contains a recycled quartz grain (doubly overgrown). Chert and polygonized quartz occur in approximately 20 percent of the sandstone clasts and one grain at Rock Creek contains orthoclase.

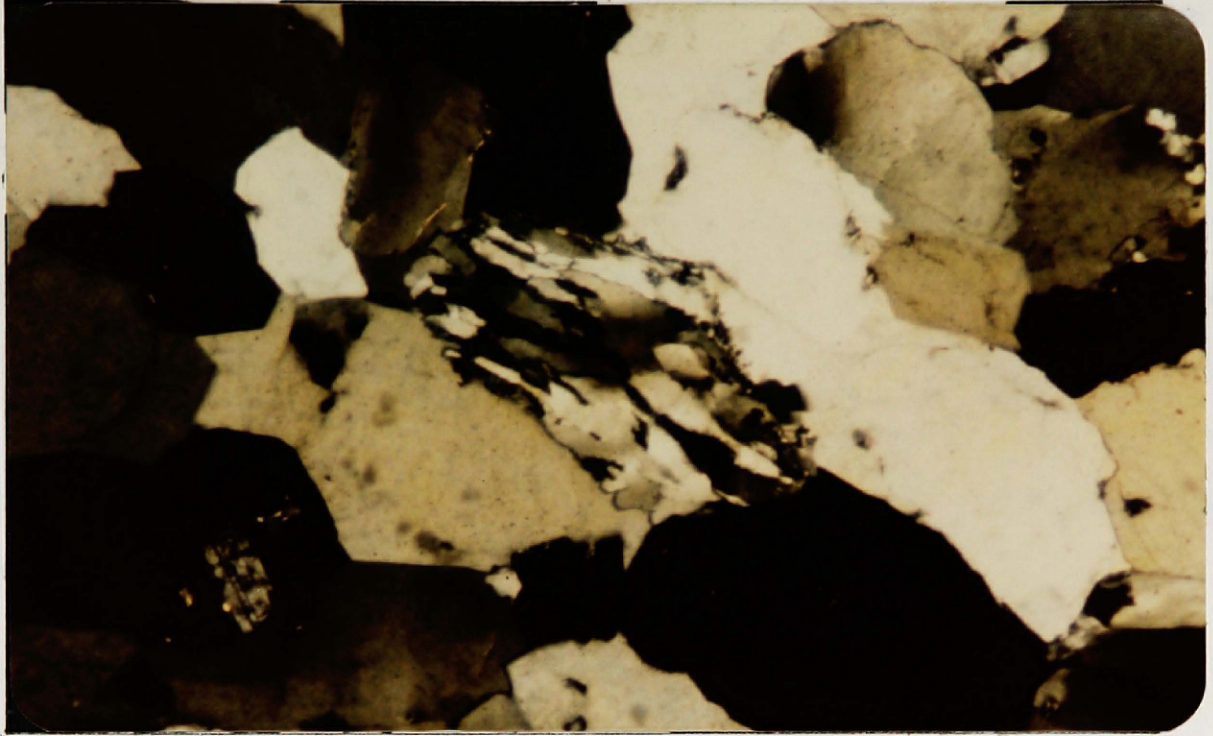
Sandstone fragments occur in moderately sorted beds of the fine-grained population (less than 250 microns; Appendix III) at Morrell Mountain and Kootenai Falls, and in coarse-grained beds at Prickly Pear, Morrell Mountain, West Glacier and Kootenai Falls. Sandstone clasts are absent at Clark Fork and Wood Creek. Sandstone clasts are most abundant in the coarse-grained population at Kootenai Falls and West Glacier (0.4 percent), and less common at Rock Creek

(0.2 percent), Willow Creek (less than 0.1 percent), and Prickly Pear and Morrell Mountain (0.1 percent) (Appendix I). Sandstone fragments make up 0.1 percent and less than 0.1 percent of the fine-grained population at Kootenai Falls and Morrell Mountain respectively.

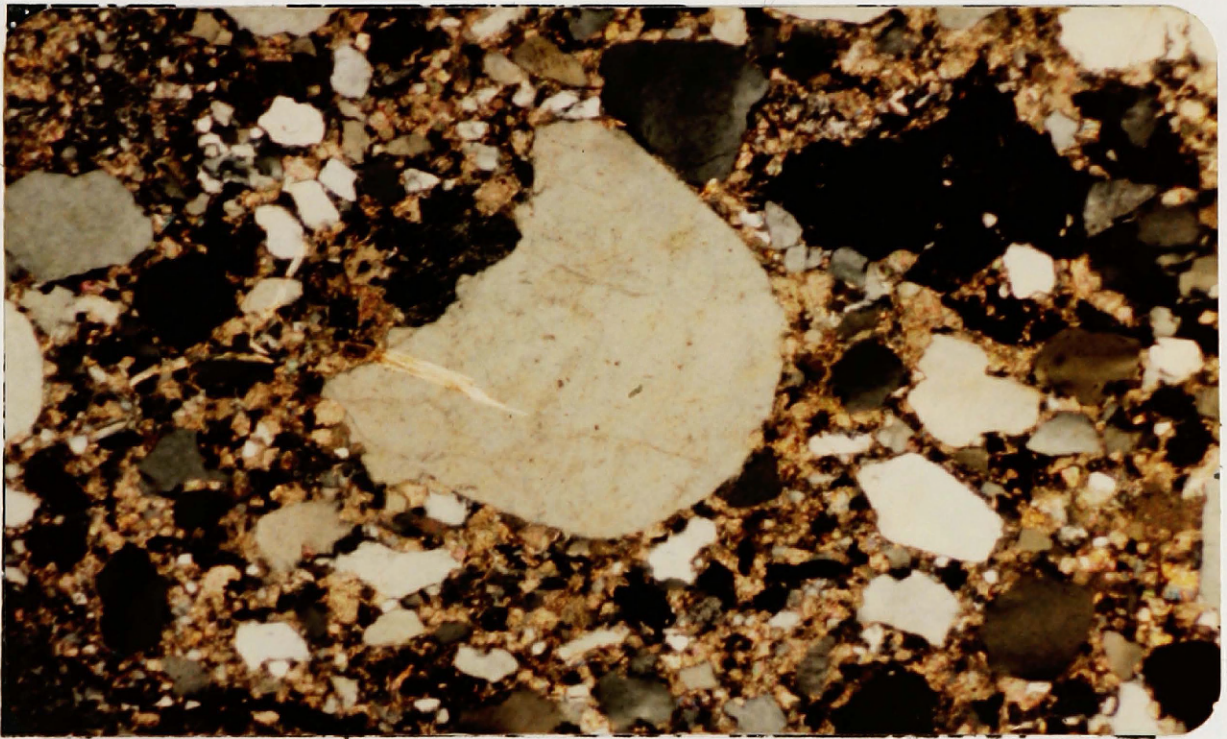
Interpretation. The sandstone fragments in the Mount Shields Formation were derived from a very fine to fine-grained and fine- to medium-grained sandstone, probably a unit lower in the Belt Supergroup such as the lower part of the Shepard Formation. In addition to quartz, this sandstone contained small amounts of chert, polygonized quartz, and orthoclase. This sandstone may have been a more important source of sediment to the Kootenai Falls and West Glacier localities than to the other localities. Sediment from sandstone exposed in the northern end of the basin could have been transported southeast to West Glacier and southwest to Kootenai Falls. Sandstone in the southwest probably shed fragments and grains northeastward.

Quartzite Fragments

Description. Quartzite fragments are distinguished from sandstone fragments by the undulatory extinction and sutured grain boundaries of the quartz grains which comprise these fragments; these quartz grains are very fine to coarse sand (88 to 1410 microns) (Fig. 7a). Quartzite fragments in the Mount Shields Formation are subrounded to rounded, approximately 500 to 2000 microns in diameter, and are only present at Prickly Pear, Willow Creek and Rock Creek. Quartzite grains are most common at Willow Creek (0.8 percent) and less abundant at Prickly Pear



(a)



(b)

Figure 7. Quartzite fragment from Willow Creek (55X) (a). Broken round quartz grain from Flint Creek (35X) (b).

(0.6 percent) to the northeast, and Rock Creek (0.4 percent) in the south. Appendix II shows that the abundance of quartzite fragments is statistically higher at Willow Creek.

Interpretation. A metamorphic source terrane in the southwest provided the quartzite fragments to the sediments comprising the Mount Shields Formation. The greater abundance of these fragments at Willow Creek might imply that their source rock was closest to this location. Sediment from this source could have therefore been transported southward to Rock Creek and northeast to Prickly Pear.

Quartz

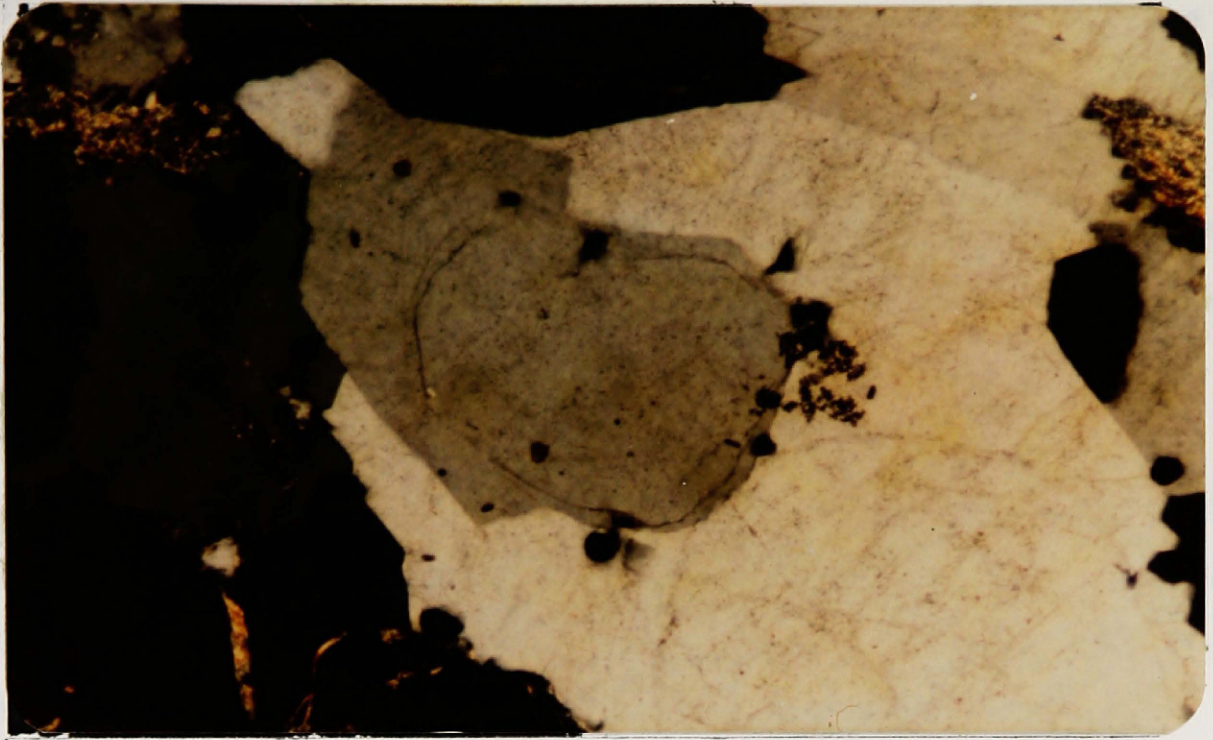
Description. Thin section examination of the Mount Shields Formation reveals two distinct populations of quartz, a moderately sorted fine-grained population (62 to 250 microns) and a poorly sorted coarse-grained population (250 to 2000 microns). The characteristics of the two modes are discussed below.

Quartz is evenly distributed throughout the fine-grained population. Approximately 53 percent of the detrital grains in the fine-grained population at Kootenai Falls, Prickly Pear, West Glacier, and Wood Creek is quartz; Clark Fork contains a little less quartz (48.6 percent) and Morrell Mountain slightly more (55.6 percent). All quartz grains in this population are subangular to angular, clear, and have straight extinction. Many grains have small overgrowths; some grains appear doubly overgrown. Polygonized quartz occurs at all sections and is included in the percentages given for quartz in Appendix I.

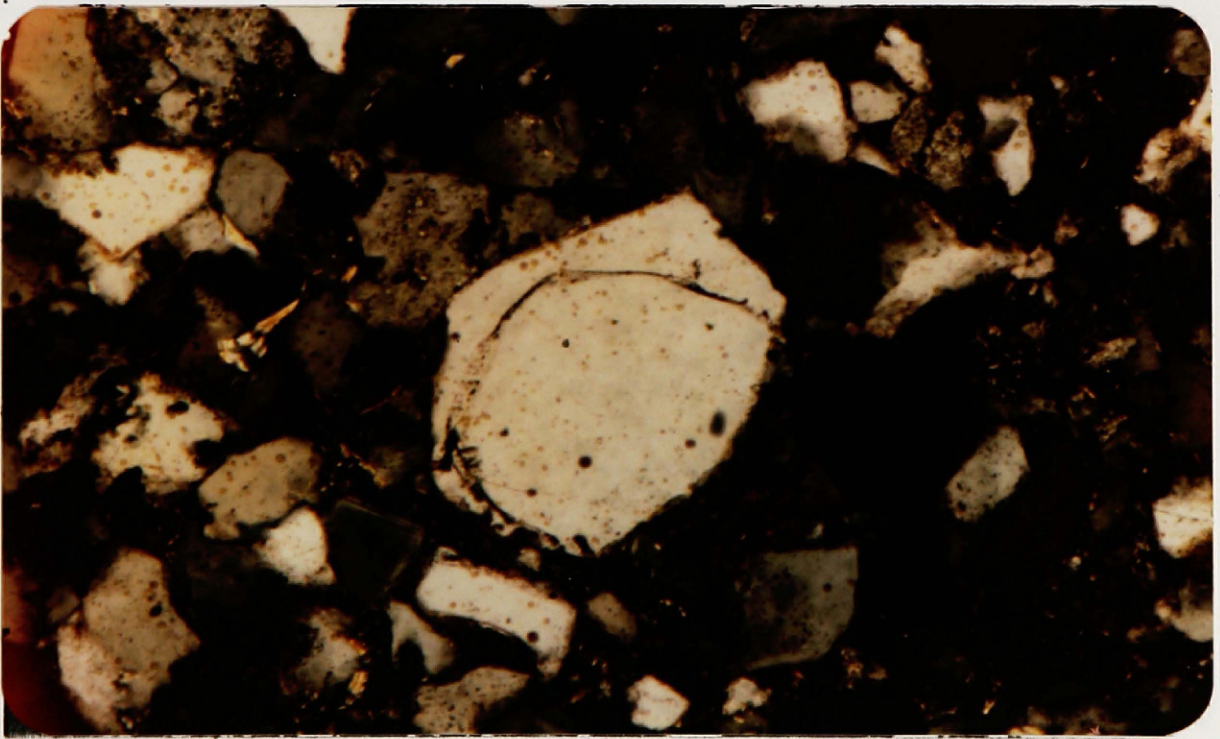
Quartz is not evenly distributed in the coarse-grained population. It is most abundant at Wood Creek (62.2 percent), Willow Creek (57.5 percent), and Rock Creek (58.7 percent). Quartz at other locations is less common and averages 51.4 percent. Appendix II shows that the greater abundance of quartz at Wood Creek, Willow Creek and Rock Creek is statistically significant. Quartz in the coarse-grained population is subrounded to well rounded, clear, and has straight or slightly undulatory extinction. Broken round grains, and large syntaxial overgrowths, some double or abraded, are also common (Figs. 7b, 8a, b). One quartz grain at Clark Fork has numerous bubble trains, and a fragment of composite, vein quartz occurs at Wood Creek.

Interpretation. Because the quartz in the granite fragments is clear and has straight extinction, the quartz in the fine-grained population which has the same characteristics was probably derived from erosion of the same granitic source rocks. The even distribution, and the uniform size and shape of these grains indicate that they were deposited evenly throughout the basin. The uniform abundance of fine-grained quartz also implies that the southern granite and the northeastern granite had similar quartz contents. The fine quartz grains with double overgrowths were clearly eroded from a sandstone, most likely the same sandstone that provided the sandstone fragments. Some polygonized grains may have also been eroded from this sandstone; others were polygonized after deposition, probably during compaction.

The high degree of rounding, and the occurrence of abraded and double overgrowths and broken round grains is convincing evidence that the coarse, common quartz grains were eroded from a sandstone coarser



(a)



(b)

Figure 8. Double quartz overgrowth from Prickly Pear (55X) (a).
Reworked quartz overgrowth from West Glacier (143X) (b).

than the sandstone which provided the fine-grained population. The relatively high percentages of coarse quartz at Wood Creek, and Willow Creek and Rock Creek suggests that the coarse-grained sandstone was exposed nearby these sections (i.e. in the east and the south). Coarse sandstone was probably a particularly important source of coarse quartz at Wood Creek. Less rounded coarse quartz resembles that in the coarse-grained granite fragments and was probably derived from the fine- to medium-grained granitic source rocks south and northeast of the Belt basin. The composite, vein quartz fragment and possibly the quartz grain with bubble trains were eroded from a hydrothermal vein (Krynine, 1946a) most likely in the northeastern granitic terrane.

Orthoclase

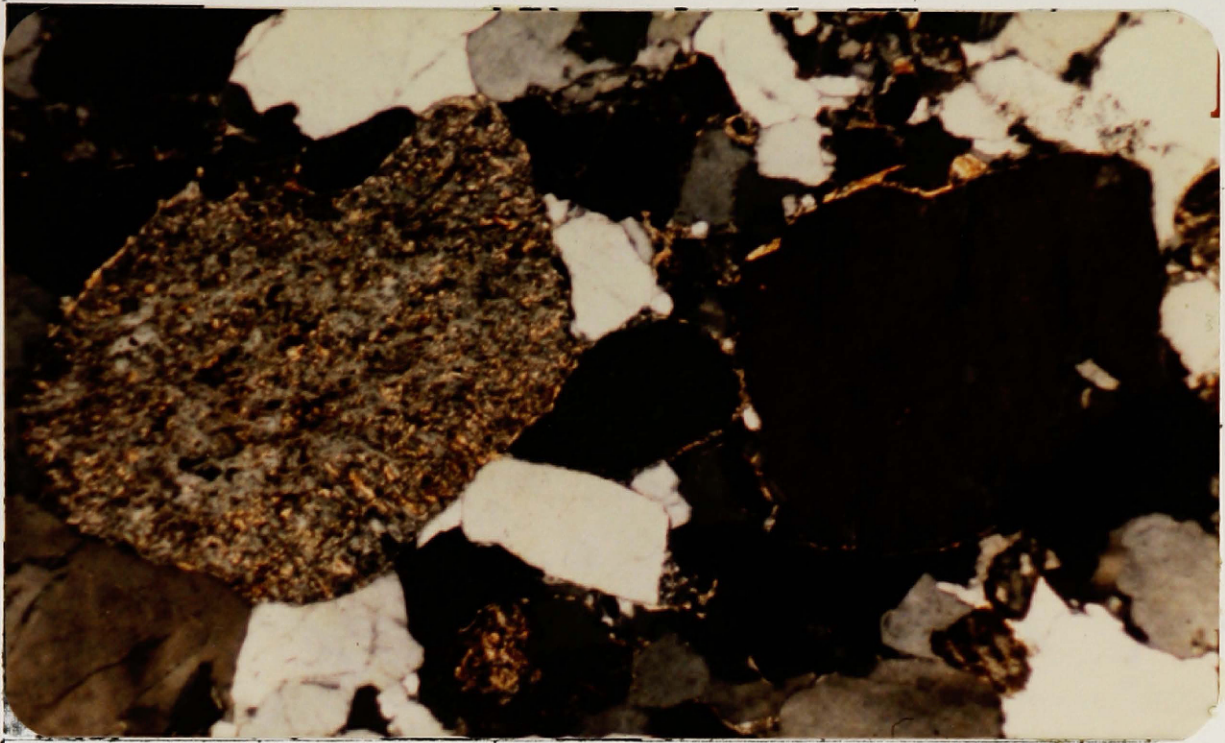
Description. There are two populations of orthoclase, a moderately sorted fine-grained population (62 to 250 microns) with a mode at 88 to 125 microns and a poorly sorted coarse-grained population (250 to 2000 microns) with a mode at 710 to 1000 microns (Appendix IV).

In the fine-grained population orthoclase is most abundant at Clark Fork (10.4 percent) and least common at Kootenai Falls (6.8 percent). The abundance of fine-grained orthoclase at the other sections is within one percent of the overall average of 8.6 percent. Specific percentages are given in Appendix I. Very fine and fine orthoclase is subangular to angular and is fresh or altered to sericite. Fresh syntaxial overgrowths surrounding fresh and altered detrital orthoclase occur in many samples. A few orthoclase grains at

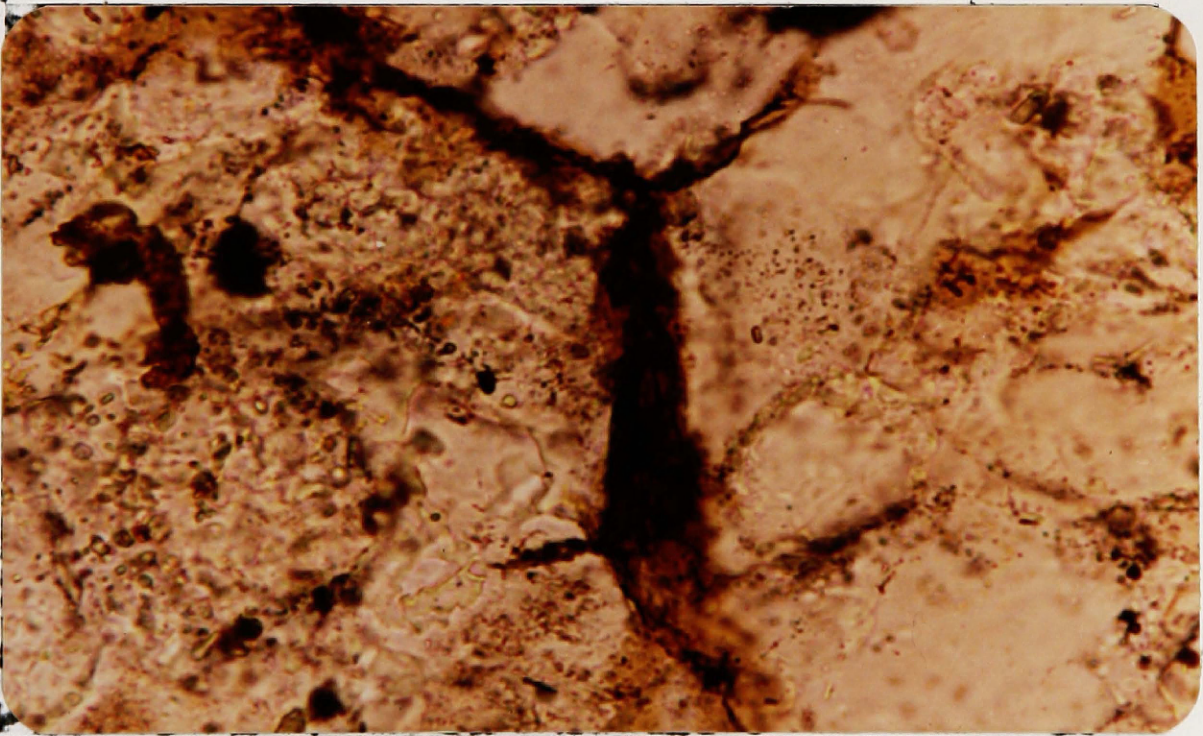
Kootenai Falls appear to have double overgrowths. Although the stratigraphic and aerial distribution of the very fine and fine-grained orthoclase is relatively uniform, the percentages of altered (sericitized) and fresh orthoclase vary inversely. Appendix II shows orthoclase is most altered (4.0 percent) at Kootenai Falls, Prickly Pear, and Morrell Mountain and freshest (8.8 percent) at Clark Fork, West Glacier and Wood Creek.

In the coarse-grained population orthoclase is subrounded to well rounded and averages 6.5 percent (Appendix II). The percentage of coarse orthoclase at Wood Creek (4.2 percent) varies farthest from this average. Orthoclase is coarsest at Rock Creek (2000+ microns), Willow Creek (2000 microns), Prickly Pear (2000 microns), Wood Creek (1410 microns), and West Glacier (1000 microns) and finer (710 microns) at Morrell Mountain, Kootenai Falls, and Clark Fork. Coarse orthoclase is fresh at Rock Creek, Wood Creek, West Glacier, and Clark Fork but is more altered at Willow Creek, Prickly Pear, Morrell Mountain and Kootenai Falls (Fig. 9a).

Interpretation. The percentage distribution and texture of the very fine and fine-grained orthoclase does not indicate source terrane locations, however, the granitic source rocks in the south and northeast are likely sources for most of this orthoclase. Doubly overgrown orthoclase at Kootenai Falls is evidence that some orthoclase was eroded from the sandstone source possibly in the north. The distribution of sericitized orthoclase probably is not a reflection of weathering or depositional processes but was probably produced by



(a)



(b)

Figure 9. Sericitized (left) and fresh, stained orthoclase (right) from Rock Creek (55X) (a). Fresh biotite flake sampled at Clark Fork (570X) (b).

differing diagenetic processes discussed in the chapter entitled Weathering and Diagenesis.

Orthoclase in the coarse-grained population bears evidence of source rocks in the south and northeast. The coarser orthoclase at Rock Creek, Willow Creek and Prickly Pear imply a southern source terrane. Orthoclase is less common in the coarse population than in the fine population because quartz from the coarse-grained sandstone source rock has "diluted" the population of orthoclase. Coarse orthoclase is least abundant at Wood Creek because this is where coarse quartz derived from the sandstone source is most common. Also, some orthoclase in the coarse population is in granite fragments and included in the percentages given for these fragments. The distribution of sericitized coarse orthoclase is similar to that of fine orthoclase. Coarse orthoclase has therefore undergone similar diagenetic alteration.

Microcline

Description. Microcline occurs in beds of all grain sizes (62 to 2000 microns). These grains are subangular to angular in the very fine and fine-grained rocks, and subrounded to rounded in the coarse and very coarse rocks. Appendix II shows that microcline is statistically more abundant (1.4 percent) at Prickly Pear and Rock Creek than at Clark Fork, Kootenai Falls, West Glacier, Wood Creek, Morrell Mountain and Willow Creek where only 0.6 percent of the detrital grains are microcline; specific percentages are tabulated in Appendix I.

Altered (sericitized) microcline occurs at all localities but is less common than altered orthoclase. Overgrowths are not common; they are small, not twinned and some at Prickly Pear are more altered than the grains.

Interpretation. A southeastern source of microcline would account for the relatively high abundance of this mineral at Rock Creek and Prickly Pear, and the smaller amount of microcline at Willow Creek. Granite fragments containing microcline are good evidence that the microcline grains in the Mount Shields Formation were eroded from southern (southeastern?), and probably northeastern granitic source rocks. Most sericitic alteration of the microcline probably resulted from burial metamorphism. Microcline is less altered than orthoclase because microcline is more stable under weathering and diagenetic conditions than orthoclase. Microcline overgrowths are more altered than detrital microcline probably because the overgrowths have monoclinic crystal symmetry and are therefore less stable (Baskin, 1956; Goldsmith, 1953).

Microperthite

Description. The abundance of microperthite is uniform within individual measured sections. These grains are very fine to very coarse sand (62 to 2000 microns), and subangular to rounded. The basin-wide distribution of microperthite is not uniform. Rock Creek, Willow Creek, Prickly Pear and Morrell Mountain contain an average amount of microperthite (1.2 percent) statistically greater than the

average percentage of microperthite (0.8 percent) at Wood Creek, West Glacier, Kootenai Falls and Clark Fork (Appendices I and II). The sodic phases of feldspar occur as "blebs" in orthoclase and less commonly in microcline. Three samples at Willow Creek contain approximately one percent antiperthite. A few grains with symplektitic texture also occur in these rocks. Overgrowths are absent and most grains are sericitized, especially the albite intergrowths.

Interpretation. Microperthite and antiperthite in granite fragments at Rock Creek and Willow Creek, and the greater abundance of these minerals at the southern-most sections suggest that the southern granite may have been more sodic than the other granitic sources. Microperthite and antiperthite were probably replaced by sericite during high grade diagenesis or low grade metamorphism.

Plagioclase

Description. The table in Appendix I shows that plagioclase (albite-oligoclase) is most common at Rock Creek (1.0 percent), comprises less than one percent of the grains at Willow Creek, Prickly Pear, Morrell Mountain, Wood Creek and West Glacier, and less than one tenth of one percent at Clark Fork, and is absent at Kootenai Falls. Harrison and Campbell (1963), however, report 10 percent or more albite-oligoclase in the Striped Peak Formation at Clark Fork, Idaho. The reason for this discrepancy is not certain. Plagioclase grains are equally abundant in the fine-grained and the coarse-grained beds, range from very fine to coarse-grained (approximately 62 to 1000 microns),

and are subangular to subrounded. All grains are twinned, not overgrown, and most are slightly to highly sericitized.

Interpretation. Plagioclase-bearing granite fragments at Rock Creek certainly indicate that the southern granitic source provided plagioclase to the Belt basin. Plagioclase at West Glacier and perhaps Wood Creek was eroded from the northeastern granite. The paucity of plagioclase grains at Clark Fork and their absence at Kootenai Falls is the result of prolonged transport and removal of these grains by weathering. Sericitic alteration of plagioclase is the result of burial metamorphism.

Muscovite

Description. Muscovite is ubiquitous throughout the Mount Shields Formation. The average length of these muscovite flakes ranges from approximately 125 microns in the very fine and fine-grained beds to 1000 microns in the coarse and very coarse beds. Notice from Appendix I that the distribution of this mineral is relatively uniform from section to section. Muscovite comprises 0.9 percent of the detrital grains in these rocks. The muscovite grains are not visibly altered.

Interpretation. The even distribution of muscovite does not reflect source terrane directions, however, the occurrence of muscovite in granite fragments at Rock Creek indicates that at least some of the muscovite grains were eroded from the southern granite. The northeastern granite most likely contained muscovite and is another probable source for these grains. The lack of alteration of the musco-

vite is not surprising considering its chemical stability under weathering and diagenetic conditions.

Biotite

Description. Less than 0.5 percent of the detrital grains in the Mount Shields Formation are biotite. These grains are commonly aligned subparallel to bedding and are approximately 250 to 500 microns long. Many biotite grains are crushed between other detrital grains. Biotite is equally abundant in beds of all grain sizes. Appendix I shows that all sections except Willow Creek contain this mineral. Most biotite flakes are oxidized to hematite but the Rock Creek, Kootenai Falls and Clark Fork sections contain some fresh biotite (Fig. 9b).

Interpretation. Biotite flakes oriented subparallel to bedding, and crushed between other grains are definitive evidence that the biotite in the samples collected is detrital. Considering that granite fragments at Rock Creek and Prickly Pear contain biotite, the southern granite is a likely source for the biotite in the southern part of the basin. The northeastern granite probably supplied biotite to the Wood Creek and perhaps the Kootenai Falls and Clark Fork sections. Absence of biotite at Willow Creek may imply that the grains at this section were derived from a source separate from that which shed grains into the Rock Creek area. Biotite oxidation is discussed in the chapter entitled Weathering and Diagenesis.

Chlorite

Description. Detrital chlorite in the Mount Shields Formation is commonly associated with biotite and has a similar distribution (Appendix I). Chlorite comprises less than 0.5 percent of the detrital grains in these rocks and is absent at Willow Creek. Chlorite flakes are approximately the same length as the biotite flakes (250 to 500 microns). They are clean and have light to medium green pleochroism.

Interpretation. The association of detrital chlorite with biotite, and their similar distribution and grain size may imply that they came from the same source terranes. Chlorite is a common alteration product of ferromagnesium minerals such as biotite, therefore, some of the biotite in the granitic source rocks may have been altered to chlorite before erosion and transport. Low grade metamorphic rocks are other possible sources of chlorite.

Chert

Description. Chert is most common at Willow Creek (1.1 percent) and less abundant (0.5 percent) at the other sections (Appendices I and II). These grains are commonly subangular to subrounded and range from approximately 62 to 1000 microns in diameter. A few chert grains at Willow Creek contain chalcedony.

Interpretation. Recall that several sandstone fragments in these rocks contain chert. The sandstone which provided these rock fragments probably also supplied most or all of the chert grains. The apparent concentration of chert at Willow Creek suggests that sandstone was a

more important source rock near Willow Creek and/or this sandstone contained more chert than the eastern and northern sandstones.

Magnetite and Leucoxene

Description. Magnetite and leucoxene (an alteration product of ilmenite) grains in the Mount Shields Formation are very fine to fine-grained (62 to 250 microns), subangular to angular, and occur in heavy mineral bands with zircon and tourmaline. Appendices I and II show that magnetite and leucoxene are most abundant (1.5 percent) at Clark Fork and Kootenai Falls and less common (0.3 to 1.0 percent) at the other sections. Magnetite is commonly oxidized to hematite. At Rock Creek, and Kootenai Falls and Clark Fork where the beds are mostly drab white or green, magnetite is not as oxidized as at the other sections where red beds predominate.

Interpretation. Magnetite and ilmenite are common accessory minerals in granite and were most likely eroded from the southern and northeastern granites. The abundance of magnetite and leucoxene at Clark Fork and Kootenai Falls probably reflects a favorable diagenetic environment for the preservation of these minerals. The section entitled Weathering and Diagenesis discusses the conditions under which the magnetite was oxidized and the association of fresh magnetite with drab and green beds, and oxidized magnetite with red beds.

Zircon

Description. Zircon is evenly distributed and comprises approximately 0.4 percent of the detrital grains in the Mount Shields

Formation. These grains are very fine to fine-grained (62 to 250 microns), clear and occur in heavy mineral bands. There are two populations of zircon, a subhedral and euhedral population, and a rounded population which also includes broken round grains. Grains from both populations occur at all sections.

Interpretation. Zircon in a granite fragment at Rock Creek is conclusive evidence that euhedral and subhedral zircon was eroded from the southern granitic source. The northeastern granite is a likely source of zircon. Zircon is a highly durable mineral, therefore, the round and broken round grains could not have attained their degree of rounding in a single sedimentary cycle. These grains were eroded from another sedimentary rock, probably sandstone in the south, east or north.

Tourmaline and Tourmaline Rock Fragments

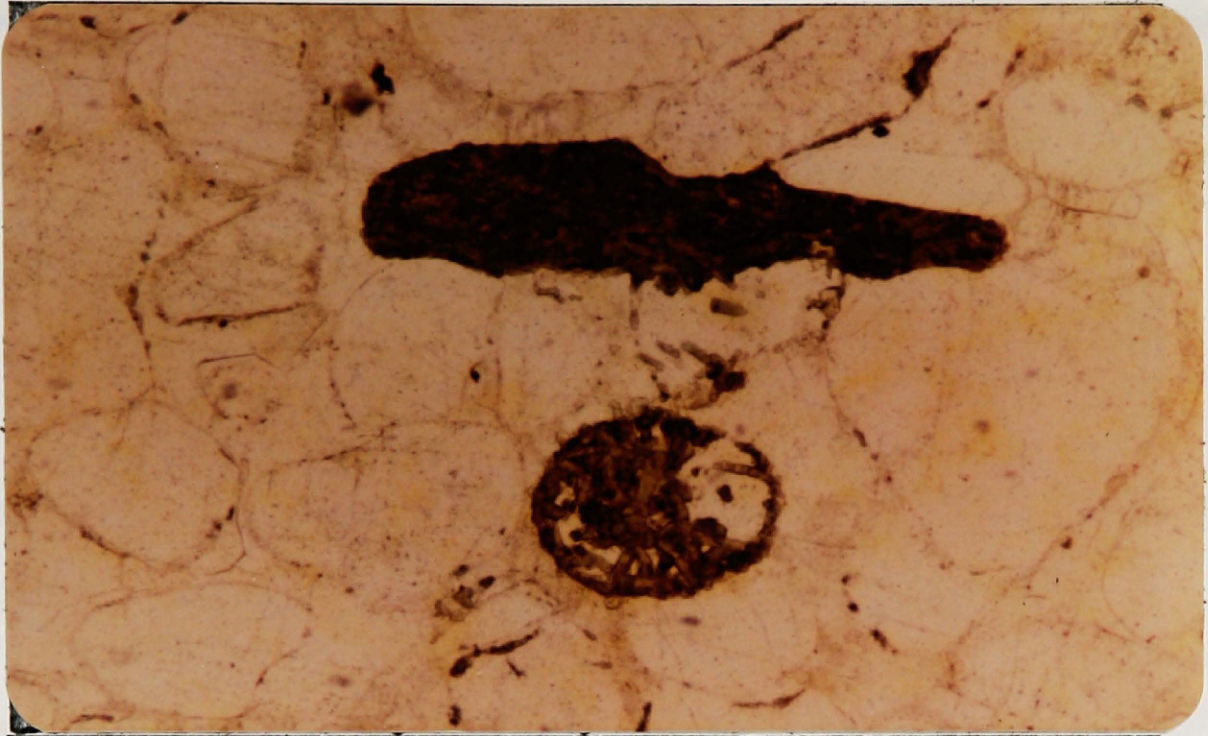
Description. The table in Appendix I shows that tourmaline is evenly distributed and comprises about 0.5 percent of the detrital grains in the Mount Shields Formation. These grains are very fine to fine sand (62 to 250 microns), subangular or subrounded to rounded, and range from mostly green, blue, brown or pink to pale yellow or black; some grains are colorless, others are zoned brown and green parallel to the c-axis. Bubbles and black inclusions are common in all varieties. Tourmaline overgrowths are ubiquitous; some are tens of microns long.

Brown and green zoned tourmaline predominantly occur as crystal groupings and aggregates or small euhedra in two kinds of rock fragments. In addition to the granite fragment at Willow Creek, this

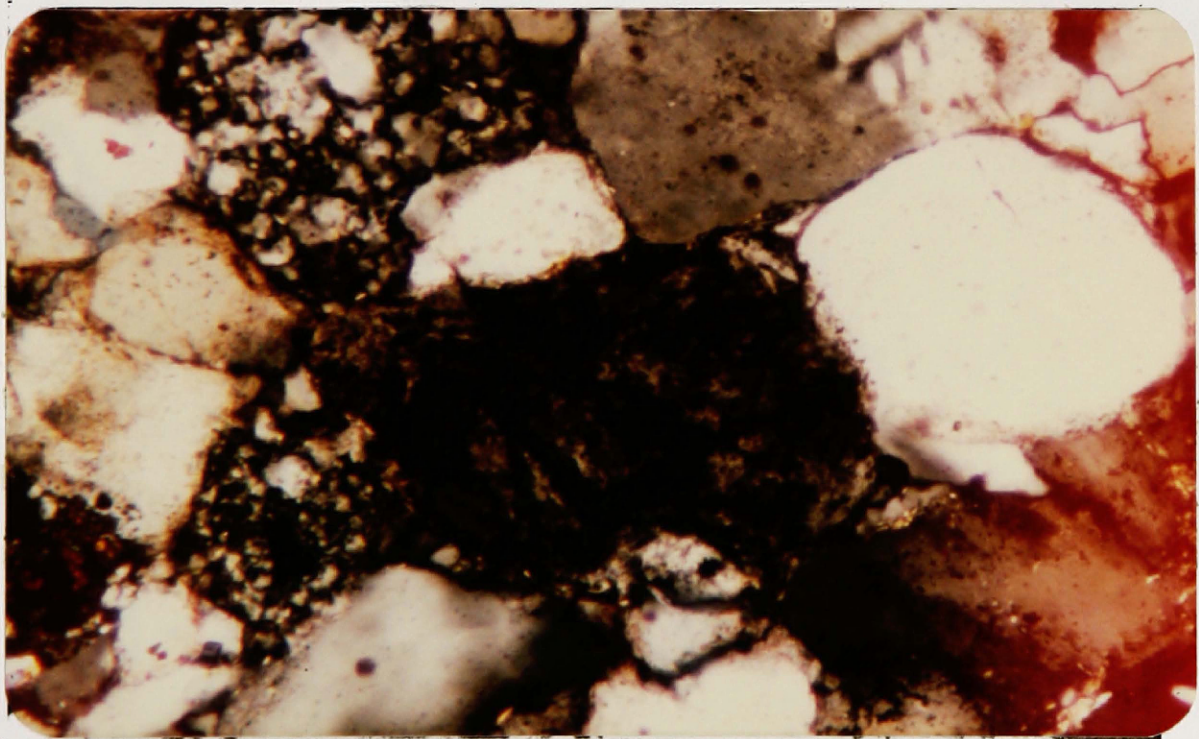
distinctive variety of tourmaline is commonly contained in very fine grained chert or amorphous-looking material, and mudchips (Fig. 10a). These rock fragments contain as much as 50 percent tourmaline and are medium- to coarse-grained; fine-grained fragments occur at Clark Fork. In Appendix I, tourmaline rock fragments are tabulated separately from tourmaline. They are found only in the middle to upper part of the Mount Shields II and are most abundant at Clark Fork (0.2 percent fine grains; 0.7 percent coarse grains) and less common at West Glacier (0.2 percent), Willow Creek (0.4 percent) and Rock Creek (less than 0.1 percent); these grains are absent at other sections.

Interpretation. Green, brown and pink, and blue tourmaline are typical granite and pegmatite varieties. The other types of tourmaline are diagnostic of pegmatite injected terranes (Krynine, 1946b). The subangular fragments were probably eroded directly from granite and pegmatites. Tourmaline-bearing granite fragments at Willow Creek support this interpretation. The occurrence of zoned tourmaline in a granite fragment at Willow Creek and in rock fragments certainly suggests that they were derived from a granite and its country rock. The southern and northeastern granitic terranes are likely sources. Schofield (1915) reports very fine grained chert bearing tourmaline crystals near Cranbrook, British Columbia (Sullivan mining district). Perhaps some of the tourmaline-bearing chert in the Mount Shields Formation was eroded from a Sullivan type terrane, although this is highly speculative.

Because of its high durability the subround and round tourmaline probably could not have attained its degree of rounding in a single



(a)



(b)

Figure 10. Tourmaline rock fragments from Willow Creek (plane light) (55X) (a). Volcanic rock fragment from West Glacier (225X) (b).

sedimentary cycle, therefore, these grains were eroded from a sedimentary rock. The southern, eastern and northern sandstones are likely sources.

Volcanic Rock Fragments

Description. Subangular volcanic rock fragments, one at Kootenai Falls (177 microns) and another at West Glacier (710 microns), occur in the upper Mount Shields II. These grains are glass containing gray or white feldspar laths (Fig. 10b).

Interpretation. Basalt, probably the Purcell lava, provided the volcanic rock fragments to the Mount Shields Formation. The Purcell lava extends from the central part of Glacier Park and northwest into British Columbia, therefore, the basalt fragments at West Glacier and Kootenai Falls were transported from the north.

EXPANDED ANALYSIS AND INTERPRETATION

This chapter analyzes the detrital mineralogy of the fine-grained population and the coarse-grained population, and fully develops and relates the interpretations discussed here to previous studies.

Fine-Grained Population

Grains in the fine-grained population range from 62 to 250 microns in diameter (Appendix III), are subangular to subrounded, and moderately sorted. Listed below are the grain types in the fine-grained population.

quartz	(52.9 percent)	chlorite	(0.2 percent)
orthoclase	(8.6 percent)	chert	(0.5 percent)
microcline	(0.8 percent)	magnetite/leucoxene	(0.9 percent)
microperthite	(1.0 percent)	zircon	(0.4 percent)
plagioclase	(0.4 percent)	tourmaline	(0.5 percent)
muscovite	(0.9 percent)	tourmaline rock fragments	
biotite	(0.3 percent)		(0.2 percent)
sandstone rock fragments	(trace)	volcanic rock fragments	
			(trace)

The size and shape of the grains in the fine-grained population is uniform throughout the Mount Shields Formation, however, the detrital mineralogy is not. The uneven distribution of fresh and sericitized orthoclase is the most striking mineralogic inhomogeneity of the fine-grained population. The abundance of microcline, microperthite, plagioclase, biotite, chlorite, chert, and magnetite and leucoxene, as well as tourmaline rock fragments, sandstone rock fragments and volcanic rock fragments, also vary within this population. The fine-grained population occurs alone or mixed with the coarse-grained population.

Coarse-Grained Population

The coarse-grained population contains grains which range from 250 to 2000 microns in diameter (Appendix III). A few grains at Rock Creek are as large as 5000 and 9000 microns. Grains in this population are subrounded to well rounded and poorly sorted. Grain types in the coarse-grained population include:

quartz	(54.3 percent)	chlorite	(0.2 percent)
orthoclase	(6.5 percent)	chert	(0.5 percent)
microcline	(0.8 percent)	granite fragments	(0.7 percent)
microperthite	(1.0 percent)	tourmaline rock	(0.2 percent)
plagioclase	(0.4 percent)	fragments	
muscovite	(0.9 percent)	quartzite fragments	(0.2 percent)
biotite	(0.3 percent)	volcanic rock fragments	
sandstone rock	(0.1 percent)		(trace)
fragments			

The texture, as well as the detrital mineralogy, of the coarse-grained population varies. The grains in this population are largest (up to 9000 microns) and most angular (subrounded) in the south and southwest, but smaller (up to 2000 microns) and more rounded (rounded to well rounded) in the other parts of the basin. Quartz in the coarse-grained population is most abundant at Rock Creek, Willow Creek and Wood Creek, whereas the percentage of orthoclase is not only low at these sections but is less common in the coarse-grained population than in the fine population.

Provenance

The mineralogy, texture and distribution of grains in the fine-grained population and the coarse-grained population reflect the com-

position and locations of four source terranes for the Mount Shields Formation. These terranes are located in Figure 11 and discussed below.

Southern Source Terrane. A southern source terrane is indicated by both textural and mineralogical evidence in the Mount Shields I and II. The maximum grain size of these rocks decreases north and north-westward from Rock Creek (9000 microns), Willow Creek (2000 microns), and Prickly Pear (2000 microns) to Morrell Mountain (710 microns), Kootenai Falls (710 microns) and Clark Fork (710 microns) certainly suggesting source rocks at the southern end of the basin (Fig. 11). The concentration of granite fragments and microcline at Rock Creek and Prickly Pear, and microperthite at Rock Creek, Willow Creek, Prickly Pear and Morrell Mountain is evidence of granitic source rocks in the south and perhaps southeast. Very fine to medium-grained quartz sandstone and coarse-grained sandstone also shed grains and rock fragments from the south.

A southern source terrane for the Belt Supergroup and therefore the Mount Shields Formation has been proposed by McMannis (1963) who demonstrates that the sediments of the LaHood Formation (lower Belt) were eroded from an uplifted crystalline terrane (Dillon Block) south of the Willow Creek fault (along the Perry line). Ruppel et al. (1981) suggests a southwestern source for the Mount Shields Formation and other Missoula Group Formations. Stratigraphic and sedimentologic evidence in the Bonner Formation cited by Quattlebaum (1980) and his paleocurrent analysis of planar cross-beds in these rocks show that the

sediments comprising the Bonner Formation were derived from the south or southwest.

Southwestern Source Terrane. Unlike the southern source rocks, coarse, chert-bearing quartz sandstone and quartzite provided the abundant quartz, chert and quartzite fragments at Willow Creek. The relatively low amount of granite fragments and the absence of biotite and chlorite (possibly a weathering product of biotite) in these rocks suggest that granite was a subordinate source of sediment. Detritus from the southwestern terrane was most likely transported north and northeastward and deposited with sediment derived from the south (Fig. 11).

Northeastern Source Terrane. Coarse-grained granite fragments and orthoclase at West Glacier and Wood Creek and abundant coarse round quartz at Wood Creek are evidence of granitic and coarse sandstone source rocks east of these sections (Fig. 11). Eastward thinning and coarsening of the Mount Shields Formation (Harrison, 1972) also implies a source terrane east of the basin. The Canadian Shield is a likely source of granite fragments at West Glacier and possibly Wood Creek. Harrison and Grimes (1970) found that the average composition of Belt rocks closely approximates the average composition of the Canadian Shield (granite-granodiorite) (Shaw et al. 1967). An eastern sandstone source rock is contrary to Harrison's (1972) assertion that sedimentary source rocks for the Belt Supergroup were limited to the southwestern part of the basin. At West Glacier, sediment from the northeastern terrane was deposited with grains eroded from the northern source area.

Northern Source Terrane. Thinning of the Snowslip Formation and the Purcell lava northwest from the Whitefish Range into southeastern British Columbia (Smith, 1963) shows that a low-relief "positive" area lay at the northern end of the basin. Price (1962) found that the Purcell basalt and the Snowslip Formation had been eroded prior to deposition of the upper part of the Shepard Formation. Pebbles of Purcell basalt in the lower part of the Gateway Formation (Shepard equivalent) near Cranbrook, British Columbia (Schofield, 1915) confirm Price's findings. The very fine to fine-grained sandstone rock fragments at West Glacier and Kootenai Falls could have been eroded from the upper most Snowslip or other fine-grained sandstone and probably a little coarse sandstone in the Shepard Formation exposed in the "positive" area at the north end of the Belt basin (Fig. 11). Basalt fragments in these sections may have been reworked from the Shepard and eroded directly from the Purcell lava.

Tectonics

The overall mineralogy of the Mount Shields Formation at a single stratigraphic section is relatively uniform and therefore does not provide evidence of tectonism or changes in source terranes during deposition of these sediments. Slover (1982), however, attributes the abrupt increase in grain size from the Mount Shields I to the Mount Shields II to a major episode of basin subsidence relative to the surrounding source terranes. The uniform texture within the Mount Shields I and within the Mount Shields II reflect gradual and constant subsidence of

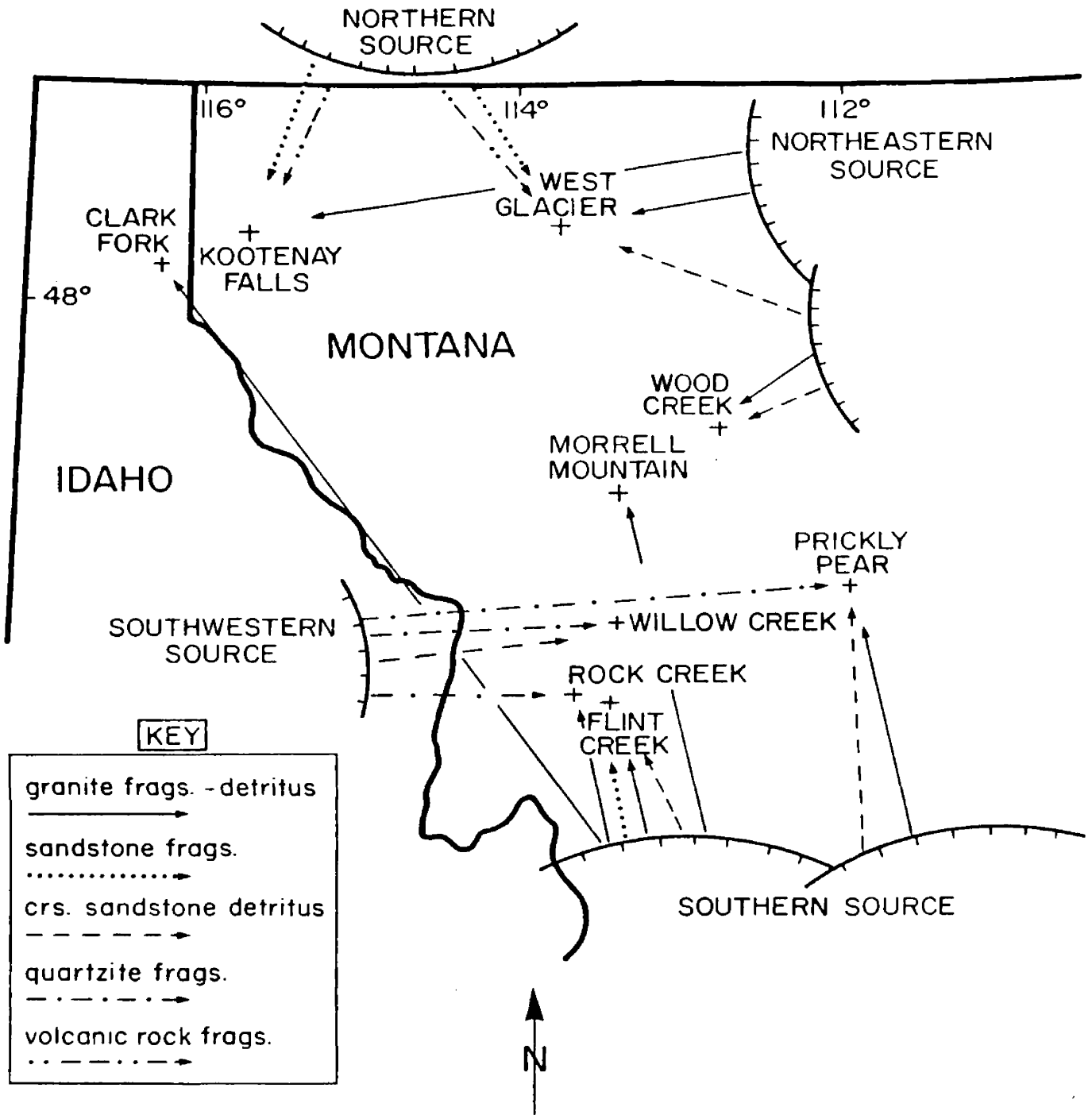


Figure 11. Schematic diagram of approximate source terrane locations showing transport directions of selected grain types.

the basin and/or uplift of the source areas. Textural variations within these units were produced by climatic changes (Slover, 1982).

WEATHERING AND DIAGENESIS

Given a granitic source and long transport, one would deduce that the sediment of the Mount Shields Formation was weathered during transport. Although most weathering products have been diagenetically altered, the climatic conditions during deposition of these sediments (Slover, 1982) and diagenetic characteristics in these rocks imply the pre-diagenetic mineralogy that includes smectite, magnetite and ilmenite, and biotite.

Weathering Products

Smectite. Although smectite was not detected in the samples collected, Eslinger and Sellars (1981) propose that smectite was originally present in the Belt Supergroup. Smectite is usually the primary weathering product of feldspars in arid environments (James et al. 1981), and salt casts in the Mount Shields III certainly suggest high evaporative conditions during deposition of these sediments. Wilson et al. (1971) report that the formation of smectite is favored by an alkaline environment with minimal leaching. The sediments of the Mount Shields Formation may have been deposited in such an environment.

Hematite. Hematite produced from the oxidation of iron in magnetite, ilmenite and biotite commonly coats detrital grains and grain overgrowths and comprises 20 percent or more of the red pigmented beds but is less common in green and white strata. Evidence that the hematite in the Mount Shields Formation was formed prior to deposition (Van Houten, 1973) is inconclusive. Although some of the hematite in

these rocks may be pre-diagenetic, evidence indicates that most hematite formed by diagenetic oxidation of iron-bearing minerals. Diagenetic hematite is discussed in a subsequent section.

Chlorite. The amount of chlorite weathered from detrital biotite and the process of this transformation are uncertain. Some chlorite probably formed by aluminum, iron, or magnesium hydroxyls replacing the interlayer potassium ions in the biotite. Vermiculite or montmorillonite may have been intermediate phases (Birkeland, 1974).

Diagenetic Products

Carbonates. Calcite and dolomite commonly occur together in the Mount Shields Formation; siderite is not common. Calcite crystals are anhedral and range from 62 to 88 microns in the fine-grained rocks and up to 500 microns in the coarse beds. Dolomite crystals are rhombic, 62 to 88 microns in diameter in fine-grained rocks and 125 to 177 microns in coarse rocks. Calcite and dolomite occur primarily as interstitial cement, and commonly have replaced detrital grains, particularly feldspar. These carbonates show no conclusive evidence of having replaced each other.

Dolomite and calcite occur low in the Mount Shields I, at the top of the Mount Shields II and in the Mount Shields III. The gradational decrease in the amount of carbonate cement from the base of the Mount Shields Formation up into the lower Mount Shields I reflects the transition from the subaqueous depositional environment of the Shepard Formation to the fluvial environment of the Mount Shields I and II. Dolomite and calcite at the top of the Mount Shields II and in the

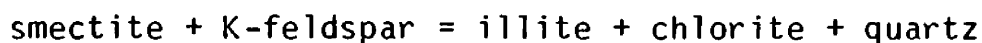
Mount Shields III also indicate a subaqueous depositional environment. Although there is no evidence that calcite precipitation preceded dolomite, the high evaporative conditions during deposition of these sediments was probably favorable for calcite followed by precipitation of dolomite.

Hematite. Hematite is most abundant in the red beds of the Mount Shields Formation and less common in the green and white beds (Rock Creek, Kootenai Falls, and Clark Fork). This mineral coats detrital grains and overgrowths and rarely forms dust rims around quartz grains. These coatings are thin or absent at grain contacts. Hematite coatings are particularly thick around biotite, magnetite, and leucoxene grains as if iron were "bleeding" from these minerals. Walker (1967; 1974) and Walker et al. (1978) also found hematite-lined interstices and coated iron-bearing minerals in desert alluvium and red beds in tropical climates. These workers propose a diagenetic origin for the hematite. Conditions during deposition of the sediments in the Mount Shields Formation were probably ideal for the production of hematite. The cyclic alternation of rainy and dry periods proposed by Slover (1982) would have repeatedly raised and lowered the water table allowing oxidation of the iron-bearing minerals. Conditions at Clark Fork, Kootenai Falls and Rock Creek were apparently unfavorable for the precipitation of ferric oxide. Here there is no evidence of hematite oxidation followed by reduction (McBride, 1974). Instead, most magnetite and biotite grains at these localities are fresh and do not appear etched or eaten away. The sediments comprising the rocks at Clark Fork and Kootenai Falls were deposited on the distal reaches of

alluvial aprons. The water table at these locations may have therefore remained high prohibiting oxygenation of the sediments. The rocks at Rock Creek were probably originally quite permeable and formed relatively high on the depositional slope. Hematite that might have been deposited in these sediments may have been flushed out and deposited farther down the depositional slope.

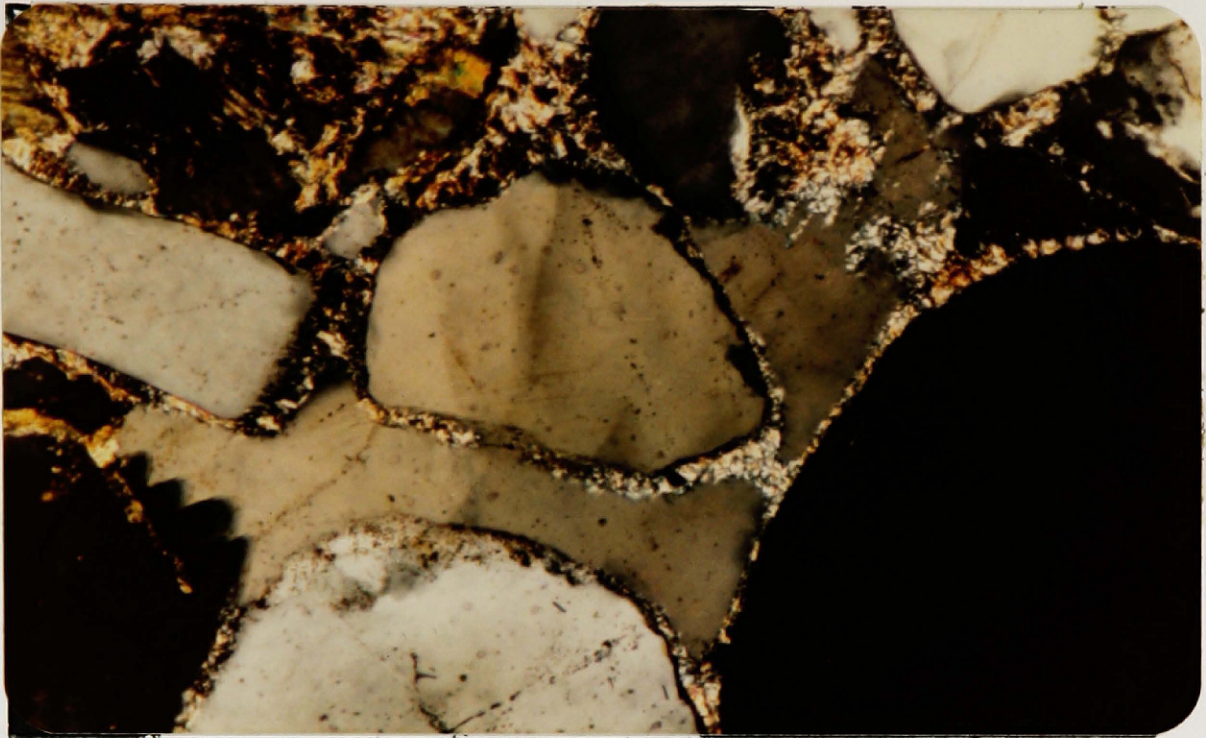
Illite and Chlorite. Illite and chlorite occur at all localities. The green beds owe their color directly to illite and chlorite (Keller, 1953) and indirectly to the absence of other coloring agents such as hematite (McBride, 1974). In addition to illite (sericite) and chlorite replacing feldspars, illite and chlorite occur intergrown and form "furry" coatings on detrital grains, grain overgrowths and less commonly between quartz overgrowths and their detrital nuclei (Fig. 12a). Smectite formed from weathered feldspars was the probable precursor of the illite (Eslinger and Sellars, 1981). Through progressive burial, temperature and pressure increased and brought about the conversion of smectite to mixed layered illite/smectite (50°C to 200°C) and finally to 1Md and 2M illite (200°C to 350°C). Eslinger and Savin (1973) determined from their oxygen isotope studies that the Belt rocks at Glacier Park reached temperatures between 225°C and 310°C. These mineralogic changes were accomplished by substitution of Al^{3+} for Si^{4+} in the tetrahedral sites (Hower et al. 1976) and fixation of K^+ in the interlayer sites of smectite (Hower et al. 1976; Hoffman and Hower, 1979). These workers propose that the aluminum and potassium were derived from potassium feldspar. Silicon released from

decomposed potassium feldspar, and iron and magnesium expelled from the octahedral layer of smectite contributed to the formation of quartz and chlorite respectively. These mineralogic transformations are represented in the equation below.

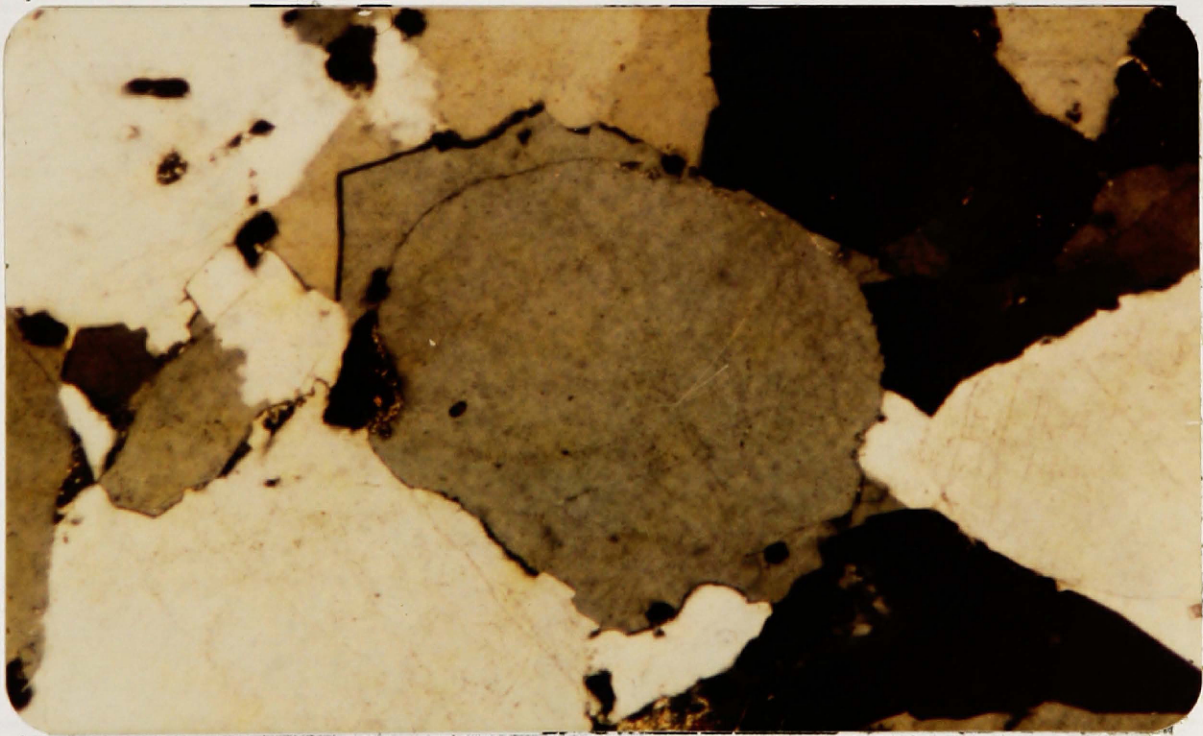


The penecontemporaneous relationship of illite, chlorite and quartz represented in the equation above agrees with the textural relationships of these minerals in thin section. High proportions of 2M illite do not necessarily accompany large amounts of sericitized feldspar therefore, the higher percentages of altered orthoclase at Willow Creek, Prickly Pear, Morrell Mountain and Kootenai Falls probably did not result from the same diagenetic processes which produced 2M illite. Apparently, local unexplained geochemical conditions controlled feldspar sericitization. Further quantitative clay mineral and geochemical analyses is needed to solve this problem.

Maxwell and Hower (1967) studied burial diagenesis in the Belt Supergroup and documented an increase in the porportion of 2M illite from the east side of the basin at the Little Belt Mountains westward to Glacier Park and Clark Fork, Idaho. They attribute this increase to westward thickening of the supergroup and therefore progressively higher grades of burial diagenesis from the east side of the basin to the west side of the basin. Proportions of 2M illite in the Mount Shields Formation follow a similar trend. Where the Missoula Group is thinnest the percent 2M illite in the Mount Shields Formation is



(a)



(b)

Figure 12. Sericite coatings on quartz grains and overgrowths from Flint Creek (143X)(a). Euhedral quartz overgrowth sampled at Prickly Pear (55X) (b).

lowest; 2M illite in these rocks is most common at sections where the Missoula Group is thickest. Listed below are the localities sampled with their corresponding proportions of 2M illite and approximate thicknesses of the lower Missoula Group at these locations. The proportions of 2M illite at each section are also listed in Appendix I.

SECTION	MISSOULA GROUP THICKNESS	PERCENT 2M ILLITE 2M/2M + 1Md x 100
WOOD CREEK	1500 FEET (MCGILL AND SOMERS, 1967)	.18 PERCENT
MORRELL MOUNTAIN	UNKNOWN	19 PERCENT
PRICKLY PEAR	4800+ FEET (KNOPF, 1963)	22 PERCENT
WEST GLACIER	5400+ FEET (ROSS, 1959)	24 PERCENT
ROCK CREEK	5000+ FEET (CALKINS AND EMMONS, 1915)	27 PERCENT
WILLOW CREEK	5000+ FEET (CALKINS AND EMMONS, 1915)	27 PERCENT
KOOTENAI FALLS	7000+ FEET (WINSTON, PERS. COMM.)	32 PERCENT
CLARK FORK	7000+ FEET (WINSTON, PERS. COMM.)	34 PERCENT

Table 1. Table showing correlation of highest proportions of 2M illite in the Mount Shields Formation with the thickest Missoula Group sections.

Maxwell and Hower (1967), as well as this thesis, do not consider additional burial of the Belt Supergroup by stacked thrust sheets. A future study examining the affects of Cretaceous-Tertiary thrusting on the low grade metamorphism of the Belt Supergroup may provide insight to this problem.

Quartz. Quartz overgrowths occur in most of the samples collected. These overgrowths range from tens of microns thick on fine grains to hundreds of microns thick on coarse grains (Fig. 12b). Some coarse-grained samples contain grains which are loosely packed and cemented in quartz. Large volumes of pore water circulating through the sediments are required for precipitation of silica cement (Blatt, 1979). The abundance of quartz cement in the coarse-grained beds is therefore not surprising considering the original permeability of these beds was probably much higher than that of the finer grained beds. Quartz cementation may have occurred over a wide range of temperatures. Detrital grains loosely packed in quartz cement certainly suggest early cementation before deep burial and therefore at near surface temperatures (Blatt, 1979). Quartz overgrowths on illite and chlorite coated grains implies that silica may have also precipitated at temperatures as high as 200°C to 350°C. Silicon released by the conversion of smectite to illite is one likely source of silica for the later quartz cement in the Mount Shields Formation. Evidence of pressure solution was not observed in these rocks and can therefore be discounted as a source of silica.

K-Feldspar. Thin (a few tens of microns) potassium feldspar overgrowths occur on detrital orthoclase and less commonly on detrital

microcline. The source of potassium, aluminum and silicon ions needed for the growth of diagenetic feldspar in the Mount Shields Formation is not certain. Ali and Turner(1982), Stablein and Dapples (1977) and other workers consider interstitially degraded detrital potassium feldspars important sources of potassium, aluminum and silicon ions. Degraded biotite may have also contributed these constituents. Baskin (1956) proposes that diagenetic feldspar forms at low temperatures (less than 100°C). Oxygen isotope studies by Savin and Epstein (1970) confirm a low crystallization temperature for diagenetic feldspar. Unlike some quartz overgrowths in the Mount Shields Formation, illite and chlorite occur only on the outside surfaces of feldspar overgrowths not between overgrowths and their detrital nuclei. Diagenetic potassium feldspar in these rocks therefore, probably precipitated prior to illite and chlorite growth at temperatures less than 200°C.

CONCLUSIONS

The detrital mineralogy of the Mount Shields Formation is characterized by a fine-grained population and a coarse-grained population. Grains in the fine-grained population are subangular to subrounded, moderately sorted and range from 62 to 250 microns in diameter. Grains in the coarse-grained population are characteristically subrounded to well rounded, poorly sorted, and range from 250 to 2000 microns in diameter. A few pebbles (5000 to 9000 microns) occur at Rock Creek. Four source terranes supplied the detritus in the Mount Shields Formation. These terranes were located south, southwest, northeast and north of the Belt basin. Granite fragments, abundant subrounded quartz and feldspar and muscovite reflect granitic source rocks in the south, southwest and northeast. Tourmaline rock fragments and tourmaline were eroded from igneous injected rocks within these granitic terranes. Coarse, rounded and well rounded quartz grains, some of which are broken, doubly overgrown or have abraded overgrowths, are common at Rock Creek, Willow Creek and Wood Creek, and were eroded from coarse, chert-bearing sandstone exposed nearby these sections. Very fine to medium sandstone also outcropped in the south. Very fine to fine-grained sandstone, perhaps the upper Snowlip and lower Shepard, shed grains and rock fragments south from the northern source terrane. Volcanic rock fragments were reworked from the lower Shepard and/or eroded directly from the Purcell basalt exposed north of the basin.

The detrital mineralogy of the Mount Shields I and II at a single locality is relatively uniform, thus reflecting constant source terrane compositions. The abrupt increase in sandstone content from the Mount Shields I to the Mount Shields II was produced by an episode of rapid basin subsidence. Textural variations within these units were climatically controlled (Slover, 1982).

Some feldspar in the Mount Shields Formation was probably originally weathered to smectite, although this mineral was converted to illite and therefore not detected in any samples. Some hematite in these rocks may have formed from oxidized magnetite, ilmenite and biotite before deposition of these minerals. Some detrital biotite might have altered to chlorite prior to deposition.

Abundant diagenetic hematite is responsible for the red coloration of these rocks. Repeated raising and lowering of the water table brought about by cyclic alternation of rainy and arid periods (Slover, 1982) oxygenated the sediments and allowed formation of iron oxide from biotite, magnetite and ilmenite. Hematite is absent from the distal green beds because a high water table produced reducing conditions. Hematite was flushed through the permeable white coarse beds at Rock Creek. Quartz overgrowths formed at low, near surface temperatures and penecontemporaneously with illite and chlorite at temperatures perhaps as high as 200°C to 350°C. The conversion of smectite to illite is one likely source of silica (quartz). Potassium feldspar overgrowths formed early at relatively low temperatures. Potassium, aluminum and silicon in these overgrowths were probably derived from degraded potassium feldspar and possibly biotite.

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APPENDIX I
TABLE OF DETRITAL MINERALOGY

GRAIN TYPES	GRAIN POPULATION	CLARK FORK	KOOTENAI FALLS	PRICKLY PEAR	MORRELL MOUNTAIN	WEST GLACIER	WOOD CREEK	WILLOW CREEK	ROCK CREEK
quartz	fine	48.6	53.6	52.5	55.6	53.6	53.4	-	-
	coarse	54.6	50.0	49.5	48.7	53.5	62.2	57.5	58.7
orthoclase (fresh)	fine	9.6	3.5	3.6	4.3	8.0	8.8	-	-
	coarse	6.5	2.5	3.2	2.0	5.8	3.7	2.6	5.4
orthoclase (altered)	fine	0.8	3.3	4.3	4.5	0.4	0.8	-	-
	coarse	1.3	3.0	3.8	3.0	2.7	0.5	3.6	2.1
microcline	general	0.5	0.6	1.4	0.8	0.5	0.7	0.3	1.3
microperthite	general	0.9	0.8	1.3	1.3	0.9	0.5	1.1	1.3
plagioclase	general	trace	-	0.2	0.7	0.3	0.6	0.4	1.0
muscovite	general	1.2	1.0	0.9	1.0	1.1	0.9	0.6	0.9
biotite	general	0.1	trace	0.5	0.6	0.1	0.6	-	0.3
chlorite	general	0.2	trace	0.2	0.4	0.5	0.6	-	0.1
chert	general	0.4	0.2	0.8	0.4	0.1	0.6	1.1	0.6
magnetite and leucoxene	general	1.5	1.5	0.8	0.9	1.0	0.3	0.3	0.8

GRAIN TYPES	GRAIN POPULATION	CLARK FORK	KOOTENAI FALLS	PRICKLY PEAR	MORRELL MOUNTAIN	WEST GLACIER	WOOD CREEK	WILLOW CREEK	ROCK CREEK
zircon	general	0.5	0.5	0.4	0.4	0.4	0.2	0.4	0.4
tourmaline	general	0.5	0.5	0.5	0.5	0.5	0.4	0.6	0.4
sandstone rock frags.	fine	-	0.1	-	trace	-	-	-	-
	coarse	-	0.4	0.1	0.1	0.4	-	trace	0.2
granite rock frags.	coarse	-	-	1.8	0.4	0.2	1.3	0.3	1.5
tourmaline rock frags.	fine	0.2	-	-	-	-	-	-	-
	coarse	0.7	-	-	-	0.2	-	0.4	trace
volcanic rock frags.	fine	-	trace	-	-	-	-	-	-
	coarse	-	-	-	-	trace	-	-	-
quartzite rock frags.	coarse	-	-	0.6	-	-	-	0.8	0.4
illite $\frac{2M}{2M+1Md} \times 100$	general	34	32	22	19	24	18	27	27

values given are percent detrital grains
trace--less than 0.1 percent

APPENDIX II
STATISTICAL COMPARISONS (t TEST)
OF DETRITAL GRAIN MEANS

FINE POPULATION

Orthoclase (Fresh)

Clark Fork, West Glacier, vs. Kootenai Falls, Prickly Pear,
Wood Creek Morrell Mountain

mean
8.8 percent

mean
3.8 percent

degrees of freedom
72

t value
8.35

probability of random occurrence
 1×10^{-15}

Orthoclase (Altered)

Clark Fork, West Glacier, vs. Kootenai Falls, Prickly Pear,
Wood Creek Morrell Mountain

mean
0.7 percent

mean
4.0 percent

degrees of freedom
69

t value
7.00

probability of random occurrence
 1×10^{-12}

COARSE POPULATION

Quartz

COARSE POPULATION (cont.)

Quartz (cont.)

Rock Creek, Willow Creek, Wood Creek, vs Clark Fork, West Glacier, Kootenai Falls, Prickly Pear, Morrell Mountain

mean
58.3 percent

mean
51.4 percent

degrees of freedom
46

t value
3.39

probability of random occurrence
 1×10^{-3}

Quartzite

Willow Creek vs. Rock Creek, Prickly Pear

mean
0.8 percent

mean
0.5 percent

degrees of freedom
29

t value
1.85

probability of random occurrence
 1×10^{-2}

GENERAL POPULATION

Orthoclase

Fine Population vs. Coarse Population

GENERAL POPULATION (cont.)

Orthoclase (cont.)

mean
8.6 percent

mean
6.5 percent

degrees of freedom
124

t value
5.20

probability of random occurrence

1×10^{-7}

Microcline

Prickly Pear, Rock Creek, vs. Clark Fork, Kootenai Falls,
West Glacier, Wood Creek,
Morrell Mountain, Willow Creek

mean
1.4 percent

mean
0.6 percent

degrees of freedom
127

t value
6.22

probability of random occurrence

1×10^{-9}

Microperthite

Rock Creek, Willow Creek, vs. Wood Creek, West Glacier
Prickly Pear, Kootenai Falls, Clark Fork
Morrell Mountain

GENERAL POPULATION (cont.)

mean
1.2 percent

mean
0.8 percent

degrees of freedom
127

t value
3.70

probability of random occurrence

1×10^{-4}

Plagioclase

Clark Fork, Kootenai Falls vs. Rock Creek, Willow Creek,
Morrell Mountain,
Prickly Pear,
Wood Creek, West Glacier

mean
0.03 percent

mean
0.5 percent

degrees of freedom
128

t value
6.03

probability of random occurrence

1×10^{-9}

Chert

Willow Creek vs. Rock Creek, Morrell Mountain
Prickly Pear, Wood Creek
West Glacier, Kootenai Falls
Clark Fork

GENERAL POPULATION (cont.)

Chert (cont.)

mean
1.1 percent

mean
0.5 percent

degrees of freedom
127

t value
3.53

probability of random occurrence

1×10^{-4}

Magnetite and Leucoxene

Clark Fork, Kootenai Falls vs. Rock Creek, Willow Creek,
Morrell Mountain,
Prickly Pear,
Wood Creek, West Glacier

mean
1.5 percent

mean
0.7 percent

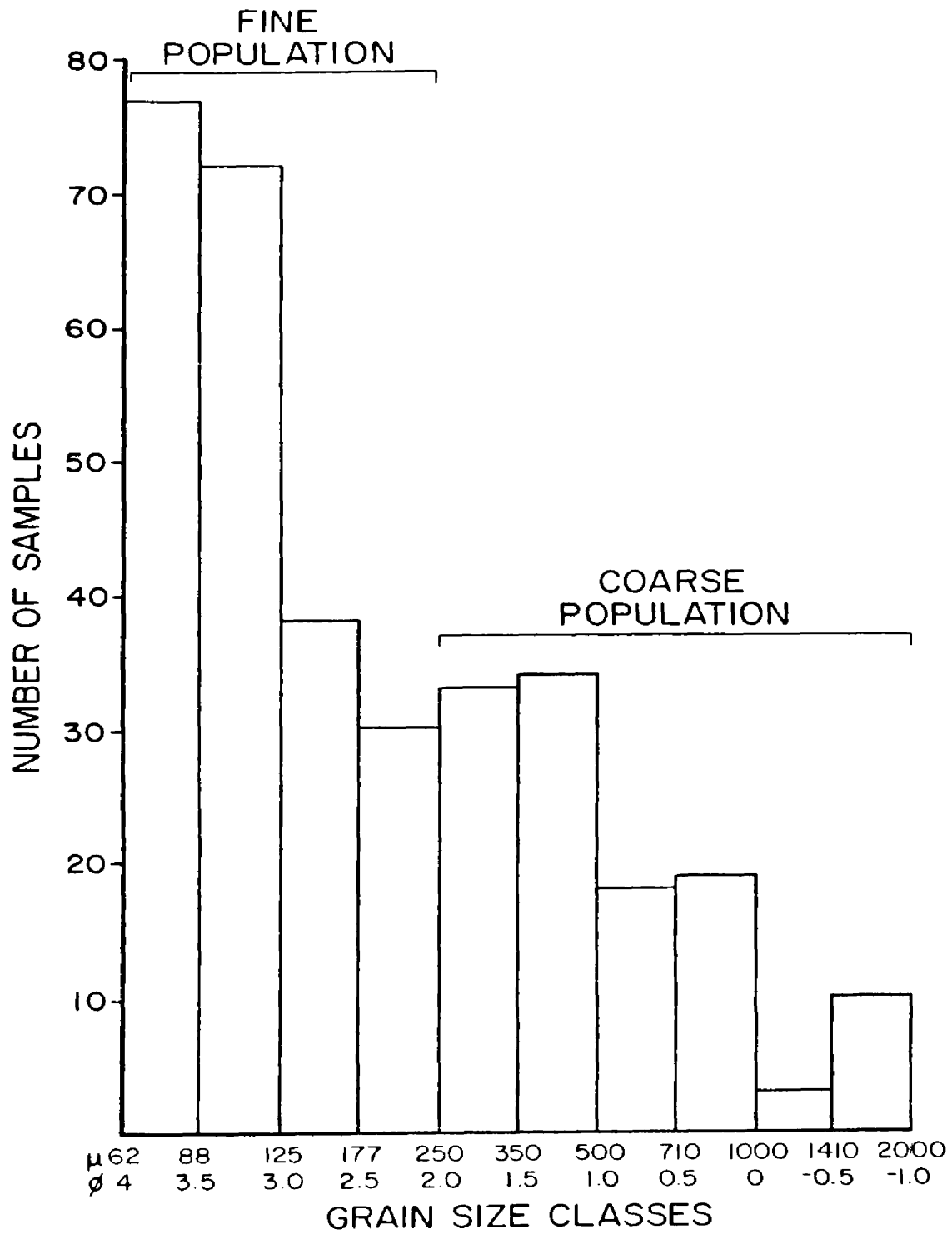
degrees of freedom
128

t value
5.00

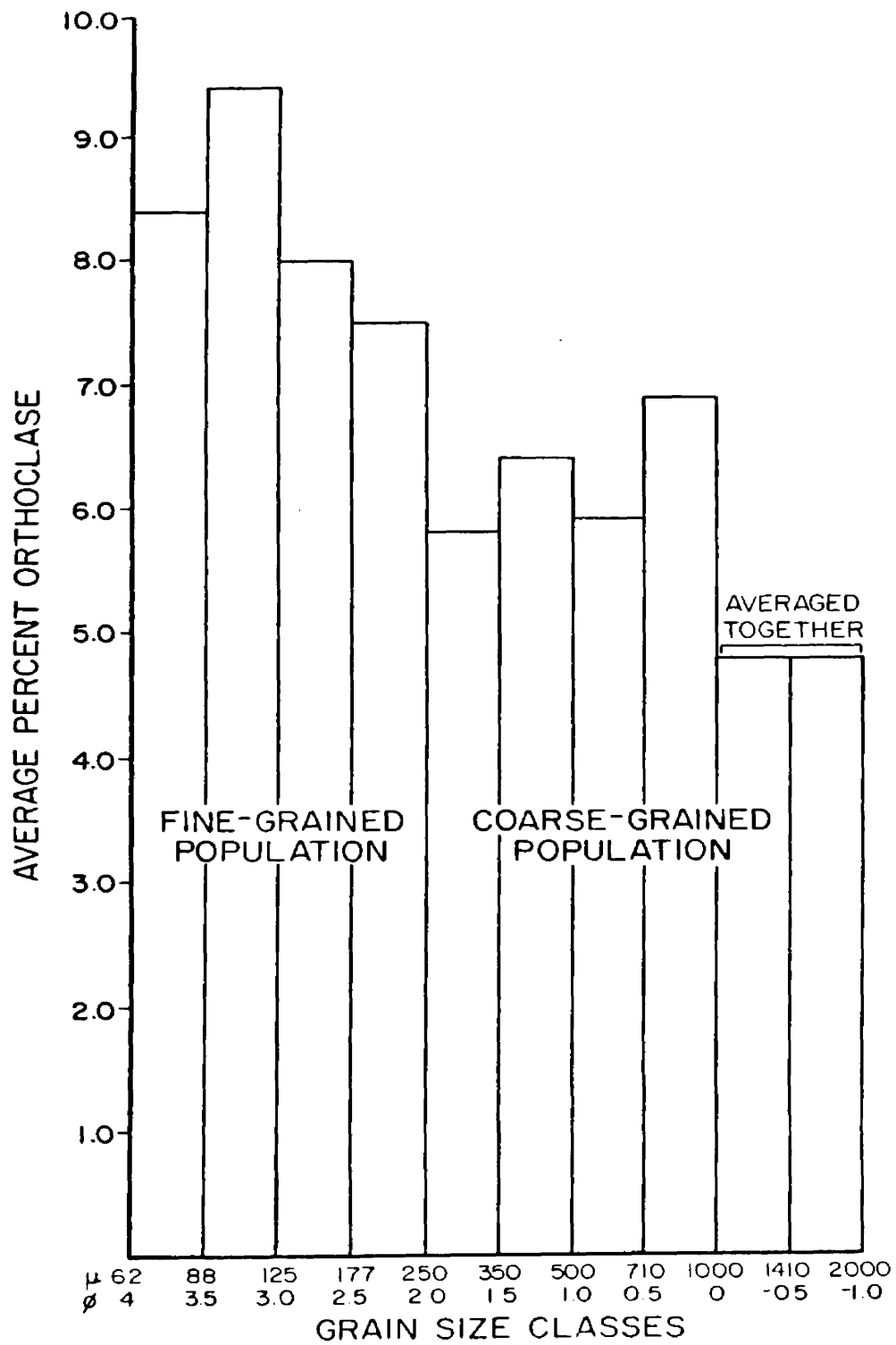
probability of random occurrence

1×10^{-6}

APPENDIX III
HISTOGRAM OF APPROXIMATE
GRAIN SIZE DISTRIBUTION



APPENDIX IV
HISTOGRAM OF APPROXIMATE
GRAIN SIZE DISTRIBUTION
OF ORTHOCLASE



APPENDIX V
THIN SECTION DESCRIPTIONS

CF0

CLARK FORK

Grains: 50%

size: $\leq 62-83\mu$ (silt to very fine sand)

shape: subangular to subrounded

sorting: moderate

composition qty -- 30%

maybe higher
(some possibly
caliche)→ K-feld -- $\approx 10\%$ (mostly clean, some slightly dirty)

perthite -- none

tour -- trace (one green grain)

musc -- 7%

leucosene -- 2%

chlorite -- 2%

Matrix: 50%

caliche (probably replaces original matrix + some grains), a little dolomite.

CF 45

fair photo of

double arrangements

Grains: 75%

size: $62-250\mu$ (very fine - fine sand) w/ $\leq 1\%$ $350-500\mu$ (med. sand)

shape: subrounded to subangular (larger grains more rounded)

sorting: moderate to poor (some grading)

composition: qty -- 62% (some w/ double arrangements) some large grains
w/ lot of bubble
trains

maybe 1-2% replace K-feld -- 7%

(pretty clean)

by caliche

microcline -- 1% clean

perthite -- 1% clean

musc trace -- 1%

tour -- trace

Matrix: 25% magnetite - 1% (scattered); zircon trace subangular
(over)

CF 49

Grains: ~50%⁺ in coarse layer; 70% in fine-grained portion

size: 62-710 μ (very fine to coarse); 62-125 μ (very fine sand)

shape: rounded \rightarrow well rounded; subangular to subrounded

sorting: poor ; moderate

composition: qtz - 39% (overgrown) ; 55% qtz

K-feld \approx 8% (some very dirty); K-feld - 12%

misc - 1%

micro - trace (diam)

perthite - 1% (diam)

zircon - trace (subrounded)

tour - trace

hematite - trace (scattered)

Matrix: 50% ; 30%

CF 100

Grains: 75%⁺

size: 62-125 μ (very fine to fine sand)

shape: subangular

sorting: moderate

composition: qtz - 55%⁺

K-feld - 10%

perthite - 2%

micro clino - trace

misc 1% biotite - trace

tour - trace

muschops - 6%

zircon - trace

mag / hematite 1% (scattered)

C.F 200

Grains: 65%

size: 88 177 μ (very fine to fine sand)

Shape: subangular to angular

Sorting: moderate

composition: qtz - 48% (a few anisogranes)

may some replace by K-feld - 12% (most is pretty clean / some ser.)

cal. mica - 1% (clean)

pyrite - 1% (clean)

foss - trace

musc - $\leq 1\%$ biotite? (trace)

mud chips - 3%

opx grains may, lens (19%) (scattered) (one subhedral grain)

chert - trace zircon - trace (subhedral, subangular)

plagioclase - trace (dirty)

C.F 231Too fine-grained (a few grains are very fine sand)

Calcic feldspathic siltstone.

~~too~~ ~~fine~~

General composition - to calcite

Tones

1. brown-green basal

2. yellow-brown (x); med. yel-

brown (z), reddish brown biotite, z

3. yel. tint (x), med. green (z); biotite

some lens (heavy mineral lenses)

qtz

K-feld

mica

zircon

brown biotite, z

Most of the rock is calcite (carbonate) and muscovite w/ maybe 30% grains

CF 300

Grains: 75%

size: 8-125µ (very fine sand)

shape: subangular to subrounded

sorting: moderate to well

composition: qtz - 51%

 K-feld - 10% (2% dirty)

 microcline - trace

 perthite - ~~trace~~ ≤ 1%

 musc - 1% biotite - trace

 tan - trace

 (mudchips) - 7%

 zircon - trace

 tan rock frag^s - trace (quite small)

 mag - 3% (scattered)

CF 355

Grains: 75%

size: 62-125µ (very fine sand)

shape: subangular to subrounded

sorting: moderate 55%

composition: qtz - (awigean)

 K-feld - 12% (-highly dirty,) - no awigean

 microcline - trace (clean)

 perthite - 1%

 musc ≤ 1%

 zircon - trace


might approach → tan - trace
1% chert

tan rock frag^s (mudchips) - ^{small als} ≤ 1%

mag / musc / mica - 2% (scattered)

mudchips - 2%

CF 405

Photo of double overgrowths, tour. rock frags.
try to find 

Grains: 75%

Size: 250-350µ (med. sand)

Shape: subrounded to rounded, a few well rounded

Sorting: moderate

Composition: qtz - 65% (double overgrowths)

maybe 9% replaced by calcite? → K-feld - 3% (slightly dirty, perthite - trace)

microcline - trace

chert - ≤ 1%

1gr. tour. rock frags (s. 1st row w/ tour) - 3%*

tour. - trace zircon - trace (larger rounded - well rounded)

Matrix: 25%

clayite - 2%

qtz - 18%

calcite

- 5%

CF 500

2-40, 2-3, 3-5

25% med. sand

Photo of rounded grains

clayite

Grains: 70%*

Size: 88-125µ (very fine sand); w/ some layers 80-350µ (very fine to med)

Shape: subangular to subrounded (smaller grains more angular)

Sorting: moderate; (coarse layers less well sorted)

Composition: qtz - 48% ← 57% in coarse layers

K-feld - 15%* (coarse layers w/ 5% coarse qtz)

perthite - ≤ 1% (clean)

micro - trace (clean)

chert - 1-2%

zircon trace (subhedral, euhedral, subangular)

micro - ≤ 1%

≈ 27% (assumed)

epaques - 2% (very rare) (some in layers)

tour. - trace

CF 545

Hayter photo of ooids?

Grains: ~60%±

ooids, qtz frags. 1% 2/1%

size: 62-83µm (v. fine sand); w/ pebble-sized mud chips, 350-500µm (med. s.)

shape: subangular; subrounded to rounded qtz + qtzite

sorting: poor

composition: qtz ~39%

clot of calcite, prob. = K-feld. ~5-7% (coarsely fine grained) (replaced clot by calcite).

Hayter? rutile -- trace

musc. ~1%

zircon - (rounded) - trace (in heavier layers)

mud chips: 10% tan - trace (approaches 1%) - brown + green + white (some as in rock frags)

magnetite ~2% in layers, esp. coarser area

ooids ~1% - many are larger and broken

(over)

Matrix 40%

FC 1

FLINT CREEK

Grains: 70%

size: 88-141µm (very fine to very coarse sand)

shape: rounded

sorting: poor

composition: qtz ~55%

K-feld. 10% (dirty, 7% stained) - some sericitized

musc. ~1% (slightly dirty)

rutile ~1% (clean)

granite frags - 1% (med. grained)

qtzite frags - 2% (fine to med. grained)

tan... trace

musc. trace

zircon trace anhedral

magnetite trace

FC-5

Grains: 70%

size: $\phi\phi - 250\mu$ (very fine to fine sand)

shape: subangular to subrounded

sorting: moderate

composition: qtz: 56%

K-feld: 10% (slightly dirty, some sensitized, all stained)

perthite: 2%

microcline: $\leq 1\%$

plag: ?-trace

zircon: trace (anhedral + subrounded) (some brown)

tour: ~~trace~~ ^{1%} (subrounded)

mag + kaurerene: trace

musc: trace siltite: trace

FC6

gtyte frags., granite frags., rounded grains

Grains: 65%

size: $177 - 1410\mu$ (fine to very coarse sand)

shape: subrounded to rounded

sorting: poor

composition: qtz - 54% (some overgrained)

K-feld: 10%

musc: $\leq 1\%$

perthite: trace

gtyte: trace (mostly v. fine grained, some coarse grained)

granite: trace (coarse grained)

musc: trace

Matrix: 35%

dolomite and calcite - 35% (replaces most of feld.) gty: trace

FC 9

photo of tan in rock frags.

Grains: 70%

size: 88 - 350 μ (very fine to mediums)

shape: subangular to subrounded

sorting: moderate

composition: qtz - 52%

K-feld - 12% (some is pretty dirty, soiled)

perthite - 2% clean

plag. - trace (dirty)

mica - 1% (clean)

musc. - trace

* mud chips - ^{siltstone} w/ tan xls - trace (a few grams)

mag. & biotite - 1% (scattered)

tan. - 1% siltite - trace

Zircon - trace (brown)

FC 14

Grains: 75%

size: 62 - 83 μ (very fine sand)

shape: subangular to subrounded

sorting: moderate to well sorted

composition, qtz - 66%

K-feld - 7% (dirty along cleavage)

mica - trace (clean)

perthite - trace

musc. - \leq 1%

plag - trace (dirty)

chert - trace

tan. - trace

opaques (oxidized grains, mag, etc.) - 1% (some under layers)

Zircon - trace

FC 15

photo of qtzite, etc.

Grains: 60%

size: 62-1410µ (very fine to very coarse sand)

shape: subrounded and rounded

sorting: very poor

composition: qtz - 44%

K-feld - 4%

(3% stained / clear)

micro - trace

perthite - 1%

chert - trace (1 grn)

qtzite - 1% (fine & coarse grns) (some / felds - granite?)

mud chips - 20%

tour - trace in 2 grns. colorless & light green (clear)

musc - trace

KF 70

KOOTENAY FALLS

Grains: 60%

size: 62-88µ (very fine sand), a few (3 or 4) fine grains

shape: subangular

sorting: moderate 50%

composition: qtz - (some little overgrowths)

+4% seric. & 12-feld - 4% (orig grain maybe higher, alot of sericite)

perthite - 1%

musc - 1%

tour - trace

magnetite/hauc - % (some in layers)

mud chips - % ignored in total

Matrix: ~~55%~~ 40%

sericite - 30%

calcite - 10%

KF 120

Photo of s.s. frag

Grains: 65% \leftarrow 10% med \rightarrow coarse $\frac{1}{2}$ size: 62-~~125~~¹²⁵ μ (very fine sand) and 350-710 μ (med. \rightarrow coarse sand)

shape: subangular-subrounded; rounded

sorting: poor

composition: gty. 59%

+ 2% epidoid \rightarrow il. feld. 3%; (some of these are grains once or twice)

by ser. carbonate

microcline - trace

pyrite - trace

zircon - trace

magnetite - 3% (scattered in layers)

177-250 μ fine sand \rightarrow sandstone frag. (a fragment of feld) - note grain w/ double overgrowth!

musc. - trace

KF 205

Grains: 65%

size: 62-125 μ (very fine sand)

shape: subangular

sorting: moderate

composition: gty. 48%

+ 7% sericitized \rightarrow il. feld. 4% (alot sericitized?)

microcline - trace

pyrite - trace

musc. - 3%

magnetite - 2% (some in layers)

zircon - (subrounded) trace

tan. - \leq 1%

Matrix: 35%

sericite 30%

chlorite - 5%

LF 265

Grains: 60%

size: 62-350 μ (very fine to med sand)

shape: subangular to subrounded

sorting: moderate to poor

composition: g/b - 53% some overgrown / may be reworked ^{clastic} overgrowth

\approx +5% replaced K-feld - 2% (seritized)

by ser. calcite partite - \leq 1%

muscline - trace

muscovite - trace

tour - trace zircon - trace (subrounded)

ilmenite - trace

Matrix: 40% mudchips -- not included in %

KF 300

Great Jay sample

Grains: 50%

size: \leq 62 μ (silt to very fine sand); 177-350 μ (fine to med. sand)

shape: subangular to subrounded; some of these are rounded

sorting: moderate to poor

composition: g/b - 39% some overgrown

slightly lower in fines
grain prob. maybe from
new stratigraphic

K-feld - 2%

partite - 1% maybe det higher (5%?) because of growth

muscline - trace

sandstone frags (very fine - fine) - \leq 1%

muscovite - 2%

some detrital chl?

tour - trace

zircon - trace (subrounded to rounded)

mag/ilmenite - 1% (scattered; mostly w/ med. grains)

Matrix: 50%

KF 350

Grains: 65%

size: 62-250 μ (very fine to fine sand)

shape: subangular to subrounded (maybe a few rounded)

sorting: moderate to poor

composition: qtz: ~~55%~~^{55%} (some overgrown)5% sericitized \rightarrow 12-feld: 3% (dirty, some sericitized)

pyrite: 1% (slightly dirty)

muscovite: trace (slightly dirty w/ some sericitized)

muscovite: $\leq 1\%$

tour: trace

zircon: trace (subangular & rounded)

magnetite/hem: trace $\rightarrow 1\%$

Matrix: 35%

sericite: 20%

chlorite: 15%

KF 400(Take photo of 'funky' sericite) (good clay sample)

Grains: 55%

size: 88-350 μ (very fine to medium sand)

shape: subangular to angular

sorting: moderate to poor

composition: qtz: 45% (a few overgrown)

2-3% sericitized \rightarrow 12-feld: 5% (dirty) some recrystallized

pyrite: 1% (dirty)

chart: trace

sillite: 1% (mud chips)

zircon: trace (subangular)

tour: trace (subangular) (brown)

muscovite: $\leq 1\%$

Matrix: 45%

KF 465 (take photo of ^(good clay sample) sericite + maybe rounded grains)

Grains: 50% (original may have been higher)

size: 62-85µ (very fine sand)

shape: subangular to angular

sorting: moderate 42%

composition: qtz - (some rounded)

biotite - trace

mus - 1%

opques (2%) mag + trace - trace (subhedral - green)

leucosene zircon - trace (subhedral)

X → no apparent K-feld - looks as though all replaced by sericite ≈ 5% sericite

Matrix: 50%

↳ various grains

composition: sericite 48% qtz - trace siderite - 2%

KF 500

Maybe a photo of s.s. frag

Grains: 70%

size: 88-177µ (very fine to fine) & 250-350µ (med. sand) < 1%

shape: subangular to subrounded; → subrounded & rounded

sorting: moderate to poor

composition: qtz 63% (some overgrown)

+ 3% rounded → K-feld - 3% (slightly dirty),

(rounded, fine-med)

perthite - 1% (clean)

muscovite - 1% (clean)

muscovite - trace

trace - trace ~~apparently~~

zircon - trace (subangular)

† sandstone frag 2 fragment (very fine to fine) well-sorted

magnetite - 1% (clean)

Matrix: overgrown grains probably sericite

KF 550

(good clay sample)

Grains: 60%

size: 88 - 250 μ (very fine to fine sand)

shape: subangular to subrounded

sorting: moderate

composition: gty - 49% (some small overgrowths)

+ 2% sensitized K-feld - 7%

micro - trace

opaque (1%) - mag. & siltite - 17%+

leucosane

tan. - trace (green) (subrounded)

zircon - trace (subrounded)

musc. - trace

Matrix: 40% sensitive - 25%

calcite - 15%

KF 600

Grains: 70%

size: 88 - 350 μ (very fine to med. sand)

shape: subangular to subrounded w a few rounded (med. sand)

sorting: moderate to poor

composition: gty - 54% (some sensitized, some overgrown)

+ 2% sensitized K-feld - 6% (some sensitized,

microcline - 2% (pretty clean)

perthite - 2% (pretty clean)

microcline - trace

tan. - trace

zircon - trace (subrounded)

magnetite - 2% (some in crack layers)

chert - trace

1 - 1% (microclips)

KF 650 (good clay sample) (photo of remobilized grains?)

Grains: 60%

size: 80-110µ (very fine to fine sand) some 250-350µ (med. sand)

shape: sub-rounded to subangular

sorting: moderate

composition: qtz - 50%

~10% sericitized K-feld - 5% (clean)

mus - trace (clean)

opaque (3%) mag. tour - trace (green) (subangular)

(patchy concentration) zircon - trace (subrounded) (some euhedral?)

perthite - trace (clean)

sillite - ~1% (clean?)

mus - trace

Matrix: 40%

composition: sericite - 3% calcite - 3%

KF 700 (good clay sample) (some remobilized grains (subround grains))

Grains: 65%

size: 62-250µ (very fine to fine sand) w/a few medium grains.

shape: subrounded to subangular

sorting: moderate

composition: qtz - 56% (some small overgrowths)

1-2% sericitized K-feld. - 5% (slightly dirty)

mus - trace (clean)

opaque (trace) mag. perthite - 1% (slightly dirty)

leucosine tour - trace (green) (subrounded)

zircon - trace (subrounded)

mud chips - 1%

clay - 1% mus. - trace

Matrix: 35% replaces qtz

composition: sericite - 3-17% qtz - trace calcite - 1%

KF 745

Grains: 65%

size: 62 - 250 μ (very fine to fine sand)

shapes: subangular to subrounded

sorting: moderate

composition: qtz ... 55%

+ 5% sericite K-feld - 3% (clean or sericitized)

epaques (1%), mag, perthite 1% (clean)

leuc. (may alter to musc. - trace (clean)

leucocrone) musc. - 1%

zircon - trace subrounded, euhedral

tour. - trace (green) subrounded

clast - trace

Matrix: 35% \leftarrow replaces grain edges

composition: sericite - 32% calcite - 3% (partly replaces sericite)

KF 800

Grains: 70%

Size: 62 - 125 μ (very fine sand); a few 177 - 250 μ (fine sand)

shape: subangular to subrounded

sorting: moderate

composition: qtz ... 55% (some overgrain)

4% sericite K-feld - 10% (clean)

perthite - trace (clean; green)

epaques (3%), mag, tour. - trace (dark brown) subangular

leucocrone, some zircon - trace (subangular)

in layers musc. - 1%

musc. - trace (clean)

Matrix: 30%

composition: sericite - 27%

detrital material (?) - 3%

KF 850

Take photos of vol. fragment (if you can find it).

Grains: 70%

size: 88-250µ (very fine to fine sand)

shape: subangular to subrounded; some rounded

sorting: moderate

composition: gty = 60% (polygonized grns, some overgrown)

±1% sericitized K-feld - 5% (some sericitized, others clean)

microcline - ≤1% (clean)

perthite - ≤1% (clean)

magnetite - 2%

(some in coarse layers)

muscovite - ≤1% biotite - trace (detrital)

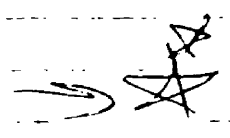
tour - trace

zircon - trace (subrounded to rounded)

chert - ≤1%

sandstone frags 20% (fine sand gty + unstained feld)

Volcanic rock frag - glass w/ feld laths 1 grn.



KF 900

Maybe take photo of detrital overgrowth.

Grains: 75%

size: 88-350µ (very fine to fine) ... few to just ~~med~~ about (250-350µ) med. ss.

shape: subangular to subrounded; some larger grains are rounded

sorting: moderate to poor

composition: gty = 62% (some ~~poly~~ polygonized; overgrown some)

+4% sericitized sand

1 sericitized

K-feld - 5% (sericitized or clean)

microcline - trace (clean)

perthite - trace (clean)

muscovite - 1%

chert - 1%

tour - trace zircon - trace (subrounded)

magnetite - ≤1% (scattered)

chlorite - trace (detrital)

sandstone frags (60° grns. cemented together) - ≤1%

KF 950

(original mineralogy obscured by calcite replacement)

Grains: 60% -

size: 62-125 (very fine sand)

shape: angular to subangular

sorting: moderate to good

composition: qtz - 52%

≈ 2% replaced by cal, K-feld - 3% ± (scattered or clean)

sericite - trace

perthite - trace (clean)

microcline - trace (clean)

musc - ≤ 1%

fau - trace

zircon - trace (subrounded)

ilmenite/magnetite - 2% (well layered w/ other leaves)

Matrix: 40% ±

calcite - 25%

sericite - 15% ± probably some chlorite too

HM 259a

MORRELL MOUNTAIN

Grains: 75%

size: ≤ 62.5 μm (very fine sand to silt)

shape: subangular

sorting: moderate

composition: qtz - 62%

K-feld - 5% (some loads dirty,)

perthite - trace (clean)

microcline - trace (clean)

plagioclase - trace (clean)

biotite - ≤ 1% (red stained)

musc - ≤ 1%

red mudclay 5% magnetite - trace - 1% (scattered)

chlorite - trace

zircon - trace (subrounded); fau - trace

MM 300

reworked ghrs.

Grains: 70%

size: 62-125 μ (very fine sand)

shape: subangular

sorting: ~~not~~ moderate

composition: qtz - 55%

original amt. prob. \rightarrow K-feld - 5%(highly altered ghrs. \approx 5-7%) (other than clean)

higher.

pyrite - 2% (slightly dirty)

plag. - 1% (slightly dirty) musc. trace (slightly dirty)

biotite trace (green) (brown)

sillite - trace (in clumps)

zircon trace (subangular)

musc. - 1%

tour. - trace

mag + kerogen 1% (scattered)

MM 390

Grains: 70%

size: 62-177 μ (very fine to fine sand)

shape: subangular to subrounded

sorting: moderate

composition: qtz - 57% (auger some)

original % \rightarrow K-feld - 5%5%+ - scattered dirt \leftarrow some stained

more dirty higher pyrite - 1% (slightly dirty)

plag - 1%

sillite - trace - maybe some dirt

tour. - trace

zircon - trace (subangular)

musc. - 1%

biotite - trace

inclusion - trace (dirty)

opaque - mag,

10% (trace)

 \uparrow
one under layer

MM 445

Grains 165%

size: 73-250 μ (very fine to fine sand)

shape: subangular, some rounded

sorting: moderate

composition: qtz - 48% (anagranum)

original high? \rightarrow K-feld - 8% (anagranum) $1^{+} 5\%$ ser.- plag - 10⁺ (clean, some dirty)- perthite - 1% microcline - ~~1%~~ $\leq 1\%$ (dirty)

- opaque - mag. zircon trace (subrounded)

(trace) tour. trace

sillite trace \leftrightarrow mud chips 1%

clast trace

biotite trace (brown) chl. trace

misc - 1%

MM 489

Grains 70%

size - 62-125 (very fine sand)

shape - subangular to angular

sorting - moderate 54%

composition - qtz - ~~54%~~ (anagranum) \rightarrow K-feld - 10%+ (pretty clean)

perthite - 2% (clean)

plag - $\leq 1\%$ (clean, slightly dirty)

microcline - trace (clean, some dirty)

clast - trace

tour. trace biotite trace (brown)

misc - $\leq 1\%$

opaque - mag, kaur, oxides (2%) in layers

zircon - subrounded (trace)

MM 533

Grains: 75%

size: 6.2-125 μ (very fine sand)

shape: subangular to subrounded

sorting: moderate 59%

composition: gtz

+6% sericitized \rightarrow K-feld - 5%

perthite - 2%

microcline - 1%

plagioclase - 2% (some sericite, some clean)

muscovite - 1%

chlorite - trace

biotite - 1% (red stained, weathered?)

chert - trace

tour - trace

zircon - trace subangular

magnetite: \leq 1% (scattered)MM 585

Grains: 70%

size: 88-172 μ (very fine to fine sand)

shape: subangular

sorting: moderate

composition: gtz 55% (some green)

K-feld - 10%

perthite 2% (clean)

microcline 1% (clean)

plagioclase \approx 1% (dirty)

tour - trace

siltite - trace

epher - 1% (mag, musc.) scattered

zircon - trace (subangular)

biotite - trace (green)

(sericite \leftarrow 3%)

MM 637

Grains: 70%
 size: 8 ϕ -17 μ (very fine to fine sand); about 4-6 grains of qtz (250-350 μ) med. sand
 shape: subangular to subrounded
 sorting: moderate
 composition: qtz 56%
 2% replaced by calc. & seric. K-feld - 7%
 microcline - trace (clean)
 perthite - \approx 2% (clean)
 plag - \approx 1% (dirty)
 musc. - 1%
 rut. - trace
 mica chips - 1%
 tour. - trace
 biotite - trace \approx 1% (brown, some green)
 opaques - mag, Fe-c. (\approx 1%) -- some under layers

MM 640

Grains: 75%
 size: 8 ϕ -17 μ (very fine to fine sand)
 shape: subangular to subrounded
 sorting: moderate
 composition: qtz 59% (lower grain)
 6% replaced by (O₂ & sericite) K-feld - 7% (sericitized - dot) (lower grain)
 perthite - 2% (clean)
 microcline - 1% (clean)
 musc. - 1%
 biotite - trace (red stained, under the dot?)
 tour. - trace
 mica chips - 3%
 plagioclase - trace (sericitized)
 chlorite - trace
 magnetite - 2% (lowered)

MM 740

Grains: 70%

size: 62-177 μ (very fine to fine sand)

shape: subangular

sorting: moderate

composition: qtz - 55% (overgrown)

K-feld - 7%

(3% altered to sericite)

microcline - trace (clean, overgrown)

plag - trace (dirty, clean)

orthite - 1% (slightly dirty)

horn - trace mg, leuc. - trace (scattered)

misc. - 2%

biotite - trace (brown)

chert - trace

mud chips - 3%

MM 790

Grains: 70%

size: 62-125 μ (very fine sand)

shape: subangular to subrounded

sorting: moderate - 57%

composition: qtz -

K-feld - 10%

(2% seritized)

microcline - $\leq 1\%$ (clean)orthite - $\leq 1\%$ (slightly dirty)

plag - trace (dirty, sericite)

misc. - $\leq 1\%$

horn - trace

biotite - trace

chert - trace

zircon - trace (angular-subangular)

opx, mg, leuc. (trace) scattered

MM 840

Grains: 70%

size: 88-172 μ (very fine to fine sand)

shape: subrounded to subangular

sorting: moderate

composition: qtz = 55% \leftarrow (some overgrown)

K-feld = 10% (1-20% sericitized grns.)

microcline = 1% (clean)

plagioclase = \leq 1% (clean)

pyrite = 1% (clean)

musc. = \leq 1%

tour. = trace zircon = trace (subangular)

opaque mag, ferr., oxides = 1% (scattered)

biotite (trace)

MM 890slide not stained \leftarrow feld all eaten upGrains: 70%⁺size: 88-35 μ (very fine to med.) (mostly very fine - fine)shape: subangular \Rightarrow angular (larger grains now rounded)

sorting: moderate

composition: qtz = 60% (some overgrown)

K-feld = 10% (locally sericitized)

zircon trace (subangular \Rightarrow subrounded)

biotite = trace

tour. = trace

chlorite = trace

difficult to see \rightarrow granite mag \gg (very fine grained, feld highly sericitized) mag (trace)

musc. = trace

Matrix: 30%

sercite = 30% ; 3 trace

MM 940

Grains: 70%

size: 88-172µ (v. fine to fine sand) w/a few ^{≤1%} 250-350 (med. sand)

shape: subangular

sorting moderate, poor in layer of coarse grains

- composition: qtz - 36% (some coarse grains)
- + 2% sericitized K-feld - 7% (sericitized)
- microcline - 1-2% (clean)
- perthite - 1% (dirty)
- musc - 2%
- biotite - trace
- clst - ≤1%
- mag, hematite - 1% (under layers)
- plag - trace (dirty)
- rock chips - 1%
- tour - trace
- Zircon - trace (subangular)

MM 990

same reworked grains. B

Grains: 75%+

size: 125-250µ (fine sand) w/o few med grains

shape: subangular to subrounded

sorting moderate

composition: qtz - 55%

this slide next ⇒ K-feld - 7% (approx) (sericitized clst) ≈ 4% sericitized
formed to K-feld % approx.

- microcline - 1% (slightly dirty)
- perthite - 1%
- musc - 2%
- biotite - 2% (red stained)
- tour - trace
- plagioclase - 1% (dirty)
- rock chips - 3%
- magnetite/hematite - 1% (some layers)

labital
chlorite - trace

MM 1035

Grains: 70%+

size: 125-250 μ (med. sand)

shape: subrounded

sorting: moderate

composition: qtz 57% (some overgrown) (some relict)

4%+ relict 12-feld - 2% (slightly dirty or sericitized)

micro - 1% (clean)

pyrite - 2% (clean)

may, trace (trace) plag - trace (dirty) + (clean)

tour - trace

mus. - 1% zircon - trace (anhedral)

zircon - trace (subangular)

biotite - trace

mud chips - 1%+ chert - trace

MM 1091

Grains: 70%

size: 62-125 μ (very fine sand)

shape: subangular to subrounded

sorting: moderate

composition: qtz 56%

3% sericitized 12-feld 8% (clean, relict)

micro - 1% (clean)

pyrite - 1% (clean, slightly dirty)

mus. - \leq 1%tour trace chert - \leq 1%

zircon - trace (subangular)

biotite - trace

may, trace - \leq 1% (conch layers)

plag - trace (dirty)

MM 1140

maybe s.s. and/or granite frags; remobilized grains.

Grains: 65%

mostly qtz w/ orth

size: 88-177 μ (very fine to fine sand); 1-2% (177-350 μ) (fine to med. sand)

shape: subangular to subrounded

sorting: moderate to poor

composition: qtz - 51%

K-feld - 6% ^{is stained} \rightarrow (5% ^{is stained} scattered)

perthite - 2% (clean)

microcline - 1% (clean)

musc - 1% biotite - trace

tour - trace

plag - trace (clean)

chert - 1%

zircon - trace (subangular)

crystals (mag, leuc., oxides) \leq 1% (scattered)

MM 1190

Take photo tour attached to qtz grains.

Grains: 65%

size: 62-125 μ (very fine sand) w/ some 125-177 μ (fine sand)

shape: subangular to angular

sorting: moderate

composition: qtz - 52%

note low \rightarrow K-feld - 2% ^{is stained} (6% ^{is scattered})

%

perthite - trace

musc - 1%

tour - trace

opacities - 1% (mag, leuc., oxides, frags) (in ^{some} layers)biotite - \leq 1% (green, brown)

zircon - trace (subangular)

mud chips - 1%

chert - \leq 1%

s. frag trace (very fine s.)

MM 1200

also det gm. counted in ssf!
 photo of coarse grns, broken round grn., serrat

Grains: 65%

size: 88-125µ (very fine sand); 250-710µ (med. to coarse sand)

shape: subangular-subrounded; subrounded to rounded & some well-rounded

sorting: poor

composition: qtz - 57% (mostly serrat) ; feld - 37% (overgrown)

K-feld - 4% + 3% highly serratized; M-feld - 1%

trace - trace

pyrite - 1%

plagioclase - trace (rather clean)

biotite - trace

mud chips - 10%

chert (in ssf) - 1 gm. trace

det. s.s. frags:

quartz - 2 or 3 each

magnetite/illite < 1% granite frags (all qtz)

trace trace (in det. grns)

zircon - trace (subrounded?)

MM 1210

Maybe good clay sample

Grains: 65%

size: 88-175µ (very fine to fine sand); 350-710µ (med. - coarse sand)

shape: subangular; rounded & well rounded

sorting: poor

composition: qtz - 56%

feld - 43%

3%⁺ replaced by 12-feld - 5% also serratized; K-feld 1% stained,

pyrite - 1% (dark) granite frags - 1%

illite - 1% (dark) mud chips - 20%

muscovite 1%

plagioclase - trace (slightly dirty);

biotite - trace

Wolfe - trace

zircon - trace

trace - trace

PP0 Poor slide - PRICKLY PEAR

Grains. 60%

size -- 62-83µ (very fine sand)

sorting - moderate

shape -- subangular

composition -- qtz - 46%

K-feld - 10% (some sensitized)

misc. - 2%

misc. - 2%

* lithics (mud chips) - 1%

perthite - 1%

chlorite - 1% (detrital)

plag. - trace

tour. - trace

Matrix: 40% (over)

PP5C Slide not stained

Grains: 75%

size -- 62-83µ (very fine sand)

sorting -- moderate

shape - subangular to subrounded

composition -- qtz - 59% (overgrown)

replaced by $Si_2O_5(OH)_2$ - K-feld - 7% (dirty w/ clean overgrowth)

biotite - 1% (plaid nature to new +) (large grains) perthite - 1% (fresh)

plag. - 1% (fresh)

hematite grains (2%) misc. - 1% (Zircon trace rounded)

(some could be weathered) mica - 1% (fresh) tour. - trace

negritite chert - 1% (some dirty)

tour. - trace ... (mud chips 3%) (some are smashed into pseudomatrix)

Matrix 25% (over)

PP 145

Grains: 65%

size: 88-125 μ m (very fine sand)

shape: subangular to subrounded

sorting: moderate

composition: qtz - 51% (some double overgrowths)

6% sericitized qtz \Rightarrow K-feld - 5% (clean) or sericitized)

replaced by carbonate perthite - 1% (clean)

microcline - \leq 1% (clean)

tour - trace

musc - 2%

zircon - trace (subrounded)

chert - 1%

Matrix: 3.5%

qtz - trace mud (alut pseudomatrix) - 20% dolomite - 13% calcite - 2%

PP 210

Grains: 70%

size: 62-125 μ m (very fine sand)

shape: subangular to angular

sorting: moderate

composition: qtz - 58% (maybe some double overgrowths)

5% sericitized qtz replaced by carbonate K-feld - 5% (clean)

perthite - 1% (clean)

microcline - trace (clean)

musc - 1%

tour - trace

chert - trace

magnetite - \leq 1% (oxidized) (scattered)

zircon - trace (broken round grain)

Matrix: 30%

qtz 1% hematite (coating grains) 15% calcite 13% (also in vein)

PP 254

Too much replacement to tell correct original composition

Grains: 50%

Size -- 62-125 μ m (very fine sand)

sorting -- moderate to poor

shape -- subangular

composition -- qtz = 43% (some overgrowths)

musc = 2%

plag. -- trace

zircon -- trace subrounded

chlorite -- 1%

chert -- 1%

some kaolinite

biotite 2%+ (altered to
musc + k)

K-feldt -- 1%

tour. -- trace

Matrix: 50%

← some citric acid replaced
by carbonate

composition -- calcite -- 45%

dolomite -- 5%

quartz -- trace also some silicate trace

The percentage of grains is probably quite different from the original?

PP 280

This rock is \approx 90% carbonate (calcite)

Minerals identified include: (grains are silt to very fine sand)

qtz

biotite

tour. blue-green varieties

musc

moy

micaceous (looks clean)

Not practical or worthwhile to estimate grain %.

PP 350

~~is~~ broken round grns. & , Maybe epheto of sericite

Grains: 65%

bimodal

size: 88-500 μ m (very fine to med. sand) \leftarrow about 3% of coarse grns,
 shape -- subangular to rounded & some well rounded

sorting -- moderate to poor

composition -- qtz -- 49% (some poly, few microcrystals)

\approx 3% sericitized K-feld -- 10% (clear, or sericitized)

microcline -- 3% (clear)

perthite -- 3% (mostly clear)

plag -- trace

granite frags -- trace tour -- \leq 1%

(fine grained) mag / biot -- 2% (widetized; in coarse layers)

(all qtz) zircon -- trace (subrounded)

g. biot -- trace (stretched)

chert -- trace

PP 415

This slide has too much carbonate replacement.

(40-50% grains, remainder in carbonate (mostly calcite w/ dolomite))

grain composition includes:

qtz & orth. % not distinguished (slide not stained) -- 30%

perthite (clear) -- \approx 1% , microcline -- trace (clear)

3% \rightarrow musc

4% \rightarrow biotite } several % also some chlorite -- 2%

tour -- green -- trace

glauconite? \leftarrow could ^{be} ~~be~~ calcite

chert? -- trace

PP 450

broken round grains? (

Grains: 65%

size: 62-88µ (very fine sand); 1% 250-350µ (medium sand)

shape: subangular to angular

sorting: moderate-poor

composition qtz - 52% (pale; some overgrown)

4% scintillated rounded K-feld - 6% (a little dirty) = (med + of grains)

microcline - 1% (clean)

perthite - 1% (clean)

plag - ?

granite frags trace musc - 1%

(v. fine) (dirty) biotite - trace

mag (oxidized) - 1% (scattered)

tour - trace

chert - 1%

PP 495

May be good clay sample (from slide)

Grains: 65%

size - 62-88µ (very fine sand)

sorting - moderate

shape - subangular to angular

composition - qtz 54% (overgrown)

+ 7% scattered, altered K-feld - 3% (overgrown) (dirty centers)

biotite - trace musc 1%

Jent - trace perthite + musc - trace (clean)

Zircon - trace (overgrown) plag - trace

tour - trace

Matrix: 35%

composition - principally biotite (min, ser) 10%

quartz - 20%

feld - 5% (over)

PP550

Grains: 70%

size: 125-177 μ (fine sand)

shape: subrounded to subangular

sorting: moderate

composition: qtz - 59% (some poly. overgrown)

4% ⁺ cemented, K-feld - 7% (slightly dirty; overgrown)

microcline - 2% (clean)

perthite - 3% (clean)

musc - trace

zircon - trace (subangular)

tour - trace

mag - trace

biotite - trace

PP610

Good clay sample

Grains: 65%

size: 62-123 μ (very fine sand)

sorting: moderate

shape: subangular to subrounded

composition: qtz - 53% (overgrown) some poly. cemented

4% ⁺ cemented K-feld - 5% (dirty; some overgrownths)

musc - 2%

musc - 1%; biotite - trace

perthite 1% clay - trace (dirty) (small grains)

tour - trace

mag: (oxidized) 1%

zircon - trace (subrounded)

chert trace

3.5%

Mature

PP 650

May be good clay sample

Grains: 65%

size - 62-125µ (very fine sand) also some fine sand

sorting - moderate

shape - subangular to subrounded

composition - qtz - 50% (overgrown) (somopaly.)

7% sintered,

K-feld - 4%⁺

musc - 1%

biotite - 1%⁺ (altered to hematite)

trace trace

chert - 1%⁺

1% oxidized grains (mag.) perthite - 1% (slightly dirty)

zircon - trace (subang) musc line - 1% (some overgrown) (slightly dirty)

plagi - trace

Matrix: 35%

composition - sericite - 33%

quartz - 2%

feld - trace

hematite - trace

PP 700

Take photo of ^{overgrown musc.} (good clay sample)

Grains: 65%

size - 88 to 177µ (very fine to fine sand)

sorting - moderate

shape - subangular to subrounded

composition - qtz - 50% (overgrown) (somopaly.)

5% sintered, K-feld - 10%

(dirty) ^{some} overgrown

biotite trace

musc trace

zircon - trace (subang) perthite - 2%

trace trace

musc line - 1% (some overgrown)

chert - trace
(some with musc
smaller than grains)

Matrix: 35%

composition - sericite - 35% (some heavy)

quartz trace

feld - trace (over)

PP 795

Take photo of overgrown microcline.
(Good sample for clay)

Grains: 65%

size -- 88-177 μ (very fine to fine sand)

sorting -- moderate

shape -- subrounded to subangular

composition -- qtz -- 53% (poly, overgrown some)

4% sericitized a lot K-feld -- 7% (overgrown)

chert -- trace microcline -- 2% (overgrown -- more altered than grain)

zircon -- trace (ground) musc -- 1%

tan. trace (anhedral & euhedral in heavy larger) plaq. -- trace biotite -- trace

perthite -- 1% (clean) may -- (in heavy: es. layer)

Matrix: 35%

composition -- sericite -- 35% trace

quartz -- trace

feld -- trace

PP 810 Slide too thin, not stained

Grains: 65%

size -- 88-350 μ (very fine to med.)

shape -- subangular to subrounded

sorting -- moderate to poor

composition -- qtz -- 57% (poly, some stretched)

orth -- 5%+ dirty grains

microcline -- 2% (clean)

perthite -- 1% (clean)

musc -- trace

chert -- trace

tan. -- trace

Matrix: 35%

sericite & clays

PP 900

Grains: 60%

size: = 62 - 125 μ (very fine sand); 250 - 350 μ (med. sand) ^{a few grains.}

sorting: moderate

shape: subangular, some subrounded

composition: gtz = 47% (poly)

3% sericitized. K-feld - 6% (dirty)

* musc. line - 2% (clean), perthite = 1% (clean)

3% un I.D. heavy (oxidized) tour - trace

in layers musc. - 1% (smashed into pseudomorph)

biotite - trace zircon - trace (subround)

- chert - trace

Matrix: 40% ← sericite

composition: - - -

PP 950

Slide not stained

Grains: 70%

size: = 80 - 170 μ (very fine to fine sand) a few med. grains

sorting: moderate to poor

shape: subangular

composition: gtz = 50% some very small, poly

musc - 2% (angular) (clean)

orth - 6% (highly sericitized)

2% heavy (may oxidized)

in layers (poor) *

chert - 2%

perthite - 1% (clean)

inusc - \leq 1%

? plag - trace

biotite - trace (altered to hematite)

tour - trace

zircon trace (small) (rounded, subrounded)

Matrix 30%

composition: sericite - 30%

quartz

trace

K-feld - trace

hematite

- trace

PP 1000

65%

Grains:

size -- mostly 85-125µ (very fine), a half-full of 250-350µ (medium sand)
 sorting -- poor
 shape -- subangular to subrounded
 composition -- qtz = 52% (very few overgrowths)
 3% sericitized K-feld = 10% (clean)
 musc - trace
 detrital heavy minerals - 2% (clean); perthite - 1%+ (clean)

(trace)

Dioctite - trace
 tour - trace (approaches 1%)
 zircon - trace (rounded)
 chert - trace

Matrix 35%

composition -- sericite 35 (even)

PP 1150 Slide Not Stained!

Grains: 60% F/Q = 1.15

size -- 62-250µ (very fine to fine sand) and 350-1410µ (med. to v. coarse) separate sharp contact
 sorting -- poor
 shape -- large grains (rounded to well-rounded), small grains (subangular)
 composition -- qtz = 48% (poly, some stretched)
 ortho = 5%+ (dirty, sericitized) grains

2% heavy minerals (mostly rimed w/ large grains) oxidized, mag.
 granite = 2% (mostly quartz)
 perthite = 1% (clean)
 musc - trace
 chert - trace (med sized grains)
 mica = 2% (clean)
 tour - trace, zircon - trace (subangular)

Matrix 40% sericite = 37%, calcite = 1%, hematite = 2%

PP 1200

Maybe good clay sample, take photo of large grains w/ vtz + feld (granite)

Grains: 65%

size -- 88-1000µ (very fine to coarse sand) (bimodal)

sorting -- poor

shape -- large grains are rounded, small grains are subangular-subround

composition -- qtz -- 48% (a few amygdales/poly)

3% sericitized K-feld 10% (clean or sericitized) 1% (high grade)

some heavy mica -- 2% (clean) qtzite -- (stretched)

some mag, mostly perthite -- 2% (clean) granite frag -- 1% (f → c)

oxidized -- 1% zircon -- trace (subhedral?) qtz + altered K-feld.

tour -- trace chert -- ≤ 1%

Matrix: 35%

composition -- sericite 32%, hematite 3% (coarse grains), qtz -- trace

Sericite has replaced alot of the feldspar.

Few amygdales are present.

PP 1250

Maybe take picture of granite frag + also sandstone frag.

Grains: 70%

F/Q = 12

Double growth on only qtz !!!

size -- 88-350µ (very fine to med. sand), 500-2000µ (coarse to very coarse)

sorting -- poor

shape -- large grains (rounded to well rounded), small grains (subangular)

composition qtz -- 53% (a few amygdales) + double growth, some poly

3%+ sericitized K-feld -- 6% (highly sericitized + some)

mica -- trace muscovite -- 2% (slightly cloudy)

≤ 1% heavy, angular mag? have with qtzite 2% + fine sand sized qtz (some really stretched)

granite frag -- 3% (med-c, has biot, orth, NS with, trace)

chert -- 2% (some lath-like, nepheline)

perthite -- 1% (slightly cloudy + dirty)

tour -- trace zircon -- trace (rounded) (anhedral)

Matrix: 30%

(qtz + feld cement traces)

composition -- sericite 27% hematite 3% (some lath intergrown w/ sericite)

PP 1345

Grains: 70%

size - 62-83 μ (very fine sand) w/ some fine sand

sorting - moderate

shape - subangular to angular

composition - qtz - 50%

2% = sericitized

K-feld - 4%

perthite - 1% (clean)

1% opaques (mag, leucosens)

musc. - 2%

o. des / evenly scattered

microcline - 1% (clean)

zircon - trace (subround)

tan. - trace

biotite - $\leq 1\%$ (altered to hematite)

Matrix: 30%

composition - sericite - 15% (some is "ferry")

hematite - 15% (patchy) (some intergrown w/ sericite)

PP II 71 (1416)

Take photo of "granite" clast; Photo of chert

Grains: 70%

size - 177-710 μ (fine to coarse sand), 710 μ -2000 μ (coarse to very coarse)

sorting - poor

shape - subangular (small grains), rounded to well rounded (large grains)

composition - qtz - 55% (a few small overgrowths) (pale) \rightarrow (some a little stretched)

4% sericitized \rightarrow

K-feld - 6% (sericitized)

tan. - trace

musc. line 1% (slightly curly)

opaques (1%) - oriented

perthite 2% (slightly curly)

gms., mag., leucosens

scattered

granite flag 3% (musc., act., qtz, biot)

chert - 1%+

musc. - trace

zircon - trace (subround) large grains, also in chert

Matrix: 30%

(over)

PP II 82 (1427') Photo of s.s. frag, some ^{small} double overgrowths

Grains: 65%

size -- 88-500 μ (very fine to med. sand), 710-2000 μ (coarse to very coarse)

sorting -- poor

shape -- small (subangular), large grains (rounded to well rounded)

composition -- qtz - 54% (alot of overgrowths) (few double overgrowths)

5% highly sericitized \rightarrow K-feld - 3% (dirty + sericitized)

no opaques

chert -- $\leq 1\%$

s.s. frag (v.f. - kerolite)

biotite - trace \neq

perthite - 2% (dirty)

micro - 1% (some are overgrowths) (dirty)

(m-c; with qtz) \rightarrow glaucochane - 2%⁺ zircon - trace, tour - trace

Matrix: 35%

composition - quartz - 35% feld. - trace

trace \rightarrow sericite -

PP II 84 (1429')

Grains: 65%

size -- 88-250 (very fine to fine sand), 500-2000 μ (coarse to very coarse) ^{only 4 or 5 grains}

sorting - moderate

shape - small (subangular to subrounded), large (rounded)

composition - qtz - 53% (some overgrowths) (some pebbles)

3% sericitized \rightarrow K-feld - 6% (altered, sericitized)

perthite - 2% (clean)

no opaques

microcline - $\leq 1\%$ (clean)

micro - trace

chert - 3%⁺ zircon - trace

glaucochane - trace (mostly K-feld) (med)

Matrix: 35%

composition - sericite - 35% qtz - trace

PPJL 150 (149S)

biotite and grains ()

Grains: 50% (original was higher)

size: 88-172µ (very fine to fine), 350-200µ (med. to very coarse sand)

sorting: poor

shape: small (subangular), large (rounded to subrounded)

composition: qtz = 39% (some overgrown) poly (some etched)

pyrite = 2% (clean)

microcline = 1% (clean)

chert

opaque (1%) mag. oxides

1%+

trace

w/coarse grns.

K-feld = 4%

4%+ sericitized and/or calcitized

granite

1%+ (fine grained; w/ qtz and orth)

trace, zircon trace

Matrix: 50% (original prob. lower)

composition: calcite 50% (med size)

sericite: trace quantity - trace

(low)

RC 0

ROCK CREEK

Grains: 75%+

size: 88-250µ (very fine to fine sand)

shape: subangular to subrounded + some rounded

sorting: moderate

composition: qtz = 61% (some overgrown) / some sericitized

+ 24% sericitized K-feld = 7% (dirty, sericitized)

microcline = 2% (clean, slightly dirty)

trace, sericitized pyrite 2% (slightly dirty, clean)

plagioclase = 1%

microcline = 1% biotite trace

chert trace

granite trace (fine)

trace

mag. = 1% (mostly unetched)

chert trace

zircon trace (subangular, angular)

RC 26

Maybe try to find s.s. frag. for photo.

Grains: 70%

size: 172-100µm (fine to coarse sand) w/lon 2grns. of v. coarse
 shape: subrounded to rounded

sorting: poor

composition: qtz - 62% (polygonal grns., anisogranular)

+2% sericitized K-feld - 5% (some clean, some sericitized)

perthite - ≤1% (slightly dirty)

miccline - ≤1% (dirty or clean)

plagioclase - trace (dirty)

no foun.

granite frags - 1% (aid. highly sericitized, fine-med. grained)

muscovite - trace

zircon - trace

difficult to define chart - trace

grn. board. ~~st. s.~~ s.s. frag - 1 grn. (fine med sand) (has poly. qb)

RC 53

double eurygrains, rounded grains

Grains: 75%+

95%

size: 62-125µm (very fine sand); 250-100µm (med. to coarse sand) (5% of rock)

shape: subangular - subrounded; subrounded to rounded

sorting: poor

composition: qtz - 60% (some polygonal, double eurygrains maybe a triple one!)

+2% sericitized K-feld - 10% (clean or dirty) ^{sericitized}

miccline - ≤1% (clean to slightly dirty)

perthite - ≤1% (clean to slightly dirty)

plagioclase - ≤1% (sericitized or clean)

mag. juv. trace
 no larger

misc 1%

fau - trace

chart - trace

zircon trace (anhedral)

blende - trace

biotite - trace (looks altered, may be to hematite?)

granite frags trace - 1% (med.) w/ 1tz, 1th

RC 268

Grains: 75%+ slide is too thin
 size: 62-125µ (very fine sand) and a few grns. (250-350µ) med.
 shape: subangular to subrounded; subrounded
 sorting: moderate
 composition: qtz - 62% (some poly, some unaggr.)
 + 3% secondary K-feld - 2% (clean or sericitized),
 mica - < 1% (clean → slightly dirty)
 perthite - < 1% (clean → slightly dirty)
 plagioclase - trace (clean or dirty)
 tour - trace
 musc - 1% biotite? - trace (highly altered to hornblende)
 zircon - trace (euhedral)
 chert - trace - 1%
 mag/oxidized - trace

RC 316

Grains: 65%
 size: 88 - 352µ (very fine to med.) and 500 - 5800µ (coarser, top pebbles)
 shape: subangular to subrounded
 sorting: very poor 49%
 composition: qtz - (some grains (some plates), some w/ a lot of bubble trails)
 ≈ 39% secondary → K-feld - 10% (highly sericitized or clean),
 perthite - 2% (slightly sericitized)
 microcline - < 1% (sericitized) - albite
 plagioclase - < 1% (mostly in granite frags) (sericitized)
 tour - trace
 zircon - trace (subround)
 musc - trace
 biotite - trace (altered rock)
 mag/trace 2% (under layers)

large grns
 granite frags - 2 or 3
 (medium coarse grained)
 (w/ perth, plag, quartz)

RC 348

broken rounded grains §

Grains: 75%

size: 28-710µ (very fine to coarse sand)

shape: subrounded to rounded and well rounded

sorting: moderate to poor

composition: qtz - 58% (maybe some pale, overgrown)

±1% sericitized

K-feld - 10% (slightly dirty)

perthite ≤ 2% (clean)

microcline - 1% (clean)

plagioclase ≤ 1% (dirty & some sericitized)

zircon - trace (anhedral or rounded)

biot - 1 tiny flake

tour - trace

glauconite! - 1 grn. musc - 2%

mag - 1% (mostly oxidized)

clast - ≤ 1%

granite frags - 1% (fine-med. grnd) (has unsp. orth)

RC 375

maybe some reworked broken grains §

Grains: 75% +

size: 62-250µ (very fine to fine) and 350-710µ (med-coarse)

shape: subangular-subrounded; rounded to subrounded

sorting: poor

composition: qtz 55% (some pale, a few small overgrowths)

+ ≤ 2% sericitized →

K-feld - 10% (clean or slightly dirty)

perthite - 1% (clean)

microcline - 2% (pretty clean)

plagioclase ≤ 1% (a little sericitized)

musc - 2%

biotite - trace (altered to chlorite)

gylite - ≤ 1% (really stretched)

granite frags - 1% (med. grn, has qtz, perth, microcline)

clast - 1%

one broken grain

zircon - trace (subang)

tour - trace (round)

mag - 2% (scattered)

large grains mostly qtz - 1 perth, orth, clast

50-50
L - - - - -

RC 402

Grains: 75%

size: 8-125-1000 (fc) w/ a few 1000-1410 (very coarse) stained with.

shape: subangular to subrounded, larger grains are rounded.

sorting: moderate to poor

composition: qtz -- 63% (pretty ^{dirty} overgrown)

≈ 2% sensitized K-feld - 7%

microcline - 1%+ (pretty clean)

perthite - 1%+ (pretty clean)

plagioclase - trace (sensitized)

granite frags - 2% (fine grains, with ^{very sensitized} plag, qtz ^{close biot})

may: trace (scattered) chert - trace (not even grain size - fine and coarse)

musc - ≤ 1%

tour - trace

qtzite - trace (stretched)

RC 423

photo of granite, sandstone, qtzite

Grains: 70%+

size: 350-9000µm (med sand to pebbles)

shape: subrounded and rounded

sorting: very poor

composition: qtz -- 52% (overgrown a lot)

K-feld - 6%

micro - 2% (clean)

perthite - 1% (dirty, sens.)

plag - 1% (sensitized and dirty)

biotite - trace

musc - trace

qtzite - 2% (58-1410µm ^{ind.} grains) (highly stained)granite frags - 5% (500-700µm, one ^{ind.} about 4000µm grains)

sandstone frags - 1-2% (6-85µm grains)

; a few small overgrowths

chert ≤ 1%

(quartz)

RC 450 slide as too thin

Grains: 80%+

size: 250-7000µ (med sand to pebbles)

shape: subrounded and rounded

sorting: poor

composition: qtz - 67% (some overgrown)

≤ 1% sericitized K-feld - 8% (clean or very sericitized);

perthite - 1% (clean or slightly ser.)

microcline - 1% (clean)

plagioclase - ~~1%~~ 2% (sericitized)

trace - trace (one granite frags: 3%+ (fine to coarse grained, orth, micro, musc, ^{zircon} plagi)

chert - trace

g, ts, te - trace (stretched)

sericitized!
orth
matrix

Matrix: 20% (looks like matrix is blocked out) sericitized & qtz cement

RC 490

Grains: 75%

size: 250-710µ (med-coarse) ²⁻¹⁰ / 2000-4000µ (very coarse pebbles)

shape: rounded to subrounded

sorting: poor

composition: qtz - 62% (some large anhedral grains, overgrown clots)

3% sericitized K-feld - 7% (matrix highly sericitized)

perthite - 5% (clean)

microcline - 1% (clean)

plagioclase - 1% (sericitized)

granite frags - 2% (ser. orth, perthite, musc, micro)

chert - trace

zircon - trace (conchoidal & anhedral)

g, ts, te - trace (stretched)

musc - trace

W0 WILLOW CREEK

Grains: 60%

size: 88-2000µ (very fine to very coarse sand)

shape: subangular (small grains); subrounded → rounded (large grains)

sorting: very poor

composition: qtz 50%

K-feld 3%

one large euhedral grn. (smaller grains)

microcline: trace (clear) small grn.

plag: trace (pretty clear) small grn.

perthite: trace (clear) small grn.

qtzite: 2% (mostly 500-700µ grains, some w/ 85µ grains)

mudchips 5% (red w/v fine sand grains)

tan

- trace (green basal sect. clear)

chert: trace

granite frag. - trace (coarse-grn.)

W13 * (take photo of tan aggregates); double overgrowths!

Grains: 70%

size: 177-710µ (fine to coarse sand)

shape: subrounded to rounded

sorting: poor to moderate 2%

composition: qtz (overgrown) double overgrowths

maybe 1% ser. → K-feld 3% (overgrown) (1% is sericitized, others are very clear)

plag: trace (clear sericitized)

perthite: trace (clear)

qtzite: 1% (coarse grn; siltite) (stretched) some

rock frag w/ euhedral tan. (see back) 2%+

bas. - 1 discrete grain (large & rounded), abundant (1%)

tiny euhedral grains

granite frags (some w/ tan) trace musc. trace

Matrix: 30%

W 50

(Thin plate of coarse grs. - qtzite, broken sand grs.)

Grains: 65%

size: 62-82µ (very fine sand) to 710-1000µ (coarse sand)

shapes: subangular to angular | rounded & well rounded

sorting: very poor - 54%

composition: qtz - (some overgrown)

5% dirty → K-feld - 5% (mostly dirty & a large % is altered)

micro - trace

Notes: feld, & diff. grn sizes not evenly distributed; just "dumped" in places

perthite - 3%

opaques (trace) - magnetite

sillite - trace

qtzite - 1% (some well stretched); granite - trace

horn - trace (green) subangular, brownish green.

plag. - trace (dirty)

zircon - trace (subrounded) almost cubical

W 51

(Thin plate of qtzite, broken sand grs.)

Grains: 65%

size: 62-82µ (very fine sand) to 710-1000µ (coarse sand)

shapes: subangular | rounded

sorting: very poor

composition: qtz - 57% (some overgrown)

most is not stained because too altered → K-feld - 4% (very dirty).

micro - trace → lower in coarse round layers (near rounded pop)

qtzite - 1%

chrt trace (one grain has ~~in~~ inclusions)

zircon trace

horn - trace (green, subangular)

plag. - trace

W 100

Photo of diable augenite, unaltered perthite, tour. R.F.

Grains: 80%

size: 125-710 μ (fine to coarse sand)

shape: subrounded (smaller grns.); rounded - well rounded (larger grains)

sorting: poor

composition: qtz -- 70% (albit augenite, many relict) (polygonized grns.)

K-feld -- 2% (dirty, $\leq 1%$ not ^{dirty &} stained) = larger grns. (round)

most do not stain?!

\Rightarrow perthite -- 1% (round albit/grn. w/ jered tour.)

~~plagioclase~~ -- $\leq 1%$ (dirty ^{slightly} clean)

muscovite -- trace

* tour. rock frags -- $\leq 1%$ (some are ~~not~~ ^{some kind of} granitic) \leftarrow *
 \uparrow has albite-rich perthite

may trace

Zircon -- trace

tour. -- $\leq 1%$ microcline -- trace (albit)

chert -- 2% (some grains sensitized)

granite frags -- $\leq 1%$

W 153

photo of gylite, doleritic

Grains: 65%

size: 62-500 μ (some mixed, some on separate layers small grns, + large grns)

shape: subrounded

sorting: poor in some layers, moderate in others

composition: qtz -- 58% (some augenite, some diable)

2% albit ^{not stained} \Rightarrow K-feld 2%

plag trace (quite dirty) (also a clean grn.)

perthite trace (dirty & stained)

gylite frags 1% (v.f. & fine sand grns) (some highly stained)

tour. trace (round grn.)

magnetite (1%) scattered

musc. -- 1%

Zircon trace

chert trace (maybe cement)

W 200

Grains: 75%

size: 62-100 μ (very fine to coarse sand)

shape: subangular to subrounded (small grns.); rounded & well rounded (large grns.)

sorting: poor

composition: qtz -- 65% (auger grain (some doublets))

K-feld -- $\leq 1\%$ stained, $\geq 5\%$ ^{aligned to} (auger grain)

perthite -- 3% (dirty; does not stain?)

plagioclase -- trace (slightly dirty)

zircon -- trace (well rounded, brown)

muscovite -- trace

TRF (loads like tour -- trace mag -- trace

tour in w/qtz chert -- $\leq 1\%$

grains: i.e. pointed granite frags trace (fine-med. grained)

trace

sandstone frags?

W 250

(Fair photo of grain w/ tour !!, auger grain qtz, g/te)

Grains: 60%

size: 125 to 500 μ (fine to medium sand)

shape: subrounded & rounded

sorting: moderate to poor

(well cemented)

composition: qtz -- 53%

(lot of auger grains) (some doublets)

K-feld 3%

(dirty) + 3% dirty & not stained

plag -- trace

(slightly dirty)

perthite 1%

(slightly dirty)

chert -- 1%

qtzite trace

musc. line trace (clean)

lithics 0.5% (grains w/ lot of green tour) TRF's

tour -- trace (green)

W302

photo of gytite, numbered preserved grains.

Grains: 65%

size: 62 - 100µ (v. fine to coarse sand)

shape: subrounded & rounded

sorting: poor

composition: gyt - 49%

K-feld - 7% (clean)

misc trace

pyrite - 1% - alot doesn't stain + is clean

trace trace

clert - 1%

zircon (brown) trace
subrounded

gytite - 2% (v. f. to coarse grains) (tidgynitic)

misc. - 3%

mud chips (mucous, some w/ hair) - 2%

Matrix: 35%

hematite - 5%

(fine-grained) stuff (misc) - 25%

W365

Grains: 70%

size: 62 - 82µ to 350 - 500µ (very fine to medium sand)

shape: subrounded & rounded

sorting: moderate

composition: gyt - 58% (very coarse) (some double)

K-feld - 10% (alot are dirty, sensitized) 5% of this is too dirty to stain

pyrite - 1% (dirty)

opaque (1%) mag.
in under layers

mud chips - 1%

hair - trace (brown, green) subrounded

misc - trace (clean)

biot. ? - almost
to hematite
(trace)

phlog - trace (dirty)

gytite - trace (5% - 10% or more)

clert - trace

TRF - very small (trace)

misc. - trace zircon - trace

W 330

photo of gty, granite

Grains: 65%

size: 62-1000µ (very fine to coarse sand) → separate beds (more field in

shape: rounded to well rounded

fin rep.

sorting: poor

composition: gty -- 54% (overgrown) (double overgrowths)

K-feld -- 3% (4% ^{not} stained & dirty, other on large clean gms)

mica -- trace dirty

pyrite -- 1% dirty

chert -- 2% (some seem to have chert along too) (lengthwise)

gty frags -- 1% (stretched)

granite frags -- 1%

mica -- trace

zircon -- trace (angular)

tour. -- trace (rounded) muscov. -- trace

W 404

broken round grains Q

Grains: 65%

size: 62-1000µ (v. fine to coarse sand)

shape: rounded

sorting: poor

composition: gty -- 53% (some overgrown, some double

K-feld -- 3% ^{stained} (4% ^{stained} ^{rest} small gms. → altered i)

mica -- trace (clean)

pyrite -- trace (clean)

granite frags -- 1% (med. ground)

gty frags -- 1% (fine-med ground)

Zircon -- rounded (trace)

plag. -- trace (dirty) (clean)

tour. -- trace muscov. -- trace

Jert 3%

W 500

Grains: 70%

size: 62-500 μ (v fine to med. sand mostly in one layer)

shape: subrounded to subangular

sorting: moderate to poor

composition: qtz - 57%

K-feld - 5% stained + (5% altered)

micro - 1% (clean)

perthite - trace (dirty)

granite - trace

chert - 2%

gylite - trace

misc. - 1%

tour. - trace

zircon - trace

W 515

Grains: 75%

size: 677-710 μ (fine to coarse sand)

shape: rounded to well rounded

sorting: poor

composition: qtz - 67%

(\approx 10-2% saturated + polyoxidized; most are overgrown ^{rest of} detrital)

K-feld - 4% stained;

perthite - 2%

microcline - 1%

microcline - trace

tour. - trace

zircon - trace (subrounded - rounded)

chert - 1%

magnetite - trace

biotite - trace (altered to hornblende)

reworked grains of it with

2% ^{altered or} unstained (overgrown)

(clean) (some slightly dirty)

(clean)

plag. - trace (slightly dirty)

granite - 1%

gylite (streaked)
+ poly qtz - trace

some with chalcocyan

WC 290

WOOD CREEK

Grains: 65%

size -- 62-125 μ (very fine sand)

sorting -- moderate

shape -- subangular to angular

composition -- qtz - 45%

K-feld ~~10%~~⁺ - pretty fresh (some slightly dirty)

plag. - 1% - " " " " "

some very rounded \rightarrow micro - trace

musc - 1% biotite - trace and chlorite

opaque (trace) mgg, oxides

mud chips - 1% trace trace

chert (and/or siltite grains) - 1%

perthite - trace

(Maybe good clay sample)

WC 295

(take photo(s) of reworked (broken) grains)!

Grains: 65%⁺size -- 62-88 μ (very fine sand), w/ 177-500 μ layers of grains (fine-med)

sorting -- moderate

shape -- subrounded \rightarrow subangular, ^{poor} subangular \rightarrow subrounded, a few round.

composition -- qtz - 56%

K-feld - 9%⁺ - pretty clean (hard to tell from qtz)musc. - 5% overgrown (reworked?) (some larger, ^{some})

perthite -- 1% a little dirty

plag. - trace mostly very fine grained

siltite - trace; biotite - trace and chlorite

trace - trace

Matrix: 35%

composition: illite - (aligned) -- 35% -

WC 310

(take photos showing K-feld stain under normal light)

Grains: 70%

size - 62-125µ (very fine sand), about 12 or so grains 125-175µ (fine sand)

sorting - moderate to well

shape - subangular to angular (some K-feld slightly euhedral)

composition - qtz - 49% - large grains less angular

K-feld - 15%

perthite - 1%

microcline - 1%

plag. - 1%

musc. - trace

Zircon - trace (very small, round)

mud chips (coarse-grained) - 5%

feldspar all dirty, esp. K-feld. ^{seriate}
chlorite - trace

trace - trace

(over)

WC 330

Grains: 65%

size - 62-125µ (very fine sand)

sorting - moderate

shape - subangular to angular, some large grains subrounded

composition - qtz - 54%

* K-feld - 10% (seriated)

microcline - 1% (one grain)

plag - 1%

perthite - trace

musc. - trace

Zircon - trace

chlorite - (trace)

tan. - trace (light gr / med. gr)

(over)

w/ bivalves
rounded also light grey / very green

trace of scattered, undized opaquers

glauconite?

Matrix: 35%

WC 350

Grains: 60%

size: $\leq 62-83\mu$ (silt + very fine sand)

sorting: moderate

shape: subangular

composition: qtz -- 48%

K-feld -- 10%+ (quite altered, sensitized)

plag. -- trace

musc. -- 2%

muco. -- trace biot./chl -- trace

chert? -- trace (very fine-grained)

tan. -- trace

Matrix: 40%

composition: red mud 40%

WC 360

Grains: 65%

size: $\leq 62-83\mu$ (silt to very fine sand)

shape: subangular to angular

sorting: moderate

composition: qtz. 49%

maybe 3% replaced by calc.

K-feld. 10%+

(3/4 is clean, 1/4 fairly dirty)
 non-sensitized

plag. -- trace

(dirty)

opaque: 1%, in layers (mag, oxides)

musc. -- $\leq 1\%$

biotite -- $\leq 1\%$ and chlorite

zircon -- trace

tan. -- trace (one large green w/ brown & green together)

Matrix: 35%

(tan)

WC 370

(late photo of coarse grains)

Grains: 70%

size: ≤ 62 - 80μ (silt + w.f. sand), ± 350 - 1000μ (m. \rightarrow c.) w/ 62 - 125μ v.f.s.

shape: subangular to angular

sorting: moderate

composition: qtz - 58% (overgrown)

musc. - 1%

^{trace}
opaque oxides, mag.

chlorite - - 1%

zircon - trace

tour. - trace

~~iron trace~~

plag. - trace (very dirty)

(Claucaite?)K-feld 10% (dirty)
(overgrown)

Matrix: 30%

composition: qtz - 3%, 27% mud

70%

rounded, subrounded; subangular

very poor overgrown

qtz - 66% (some preserved)

K-feld - 2% (slightly dirty)

mud chips - 1%

plag. - 1% (fairly fresh)

musc. - trace

granite frag

Matrix: 30%

composition: qtz - 12%

siderite + calcite - 3%

mud - 15%

WC 390

Grains: 70%

size: 62-125 μ (very fine sand)

shape: subangular, subrounded

sorting: moderate

composition: qtz - 57% (overgrown)

K-feld - 13% (overgrown)

plag. - \leq 1%

musc. - 1%

musc. - 1%

sericite - 1%

mud chips - 1%

chert - trace

biotite - trace

no tour. found



WC 410

Grains: 70%

size: 62-88 μ (very fine sand) also see other side

shape: subangular, subrounded

sorting: moderate (layer along edge is very poor -- see comments)

composition: qtz -- 55% (overgrown)

+2% sericitized K-feld -- 10% " dirty

micro -- 1% (overgrown) " Fairly clean

opaque -- 1% magnetite, in
crude layers

pyrite -- 2% (overgrown) " dirty

chert -- 1%

zircon -- trace biotite -- trace chlorite

sillite? -- 1% " trace

plag -- trace dirtier

Matrix: 30% calcite -- 2%, hematite -- 2%, feld -- 3%, qtz -- 5%

WC 425

Grains: 65%

size: \leq 62 μ (silt > very fine sand)

shape: subangular to angular

sorting: moderate

composition: qtz -- 52% (overgrown)

K-feld -- 7% (seems quite dirty) (overgrown)

plag -- trace (dirty)

mud chips -- 1-2%

chert? -- 1% " trace

pyrite -- < 1%

biotite -- 1% " ; chl. 1%

micro -- 1% " trace, zircon -- trace

(over)

WC 480

Grains: 60-65%

size: $\leq 62-82\mu$ (silt \rightarrow very fine sand)

shape: subangular to angular

sorting: moderate

composition: qtz -- 40% (small ^{some} anorthoclase)

K-feld -- 10% (slightly dirty)

plag. -- trace

musc. -- 2%

 \rightarrow biotite -- 1%⁺

chert -- 1%

chl -- 2%⁺ // beddingglauconite ? trace

mudclays 2%

mudclays -- 2%

tour. -- trace

WC 491

Grains: 70%

size: $\leq 62-82\mu$ (silt to very fine sand)

shape: subangular to angular w/ mostly subangular & angular

sorting: moderate

composition: qtz -- 60% (anorthoclase)

maybe 1%⁺ repleat
by calcite

K-feld -- 7%

biotite -- trace (very small grains)

opaque -- trace

musc. -- trace

mag. oxides

chert -- 2%

(scattered)

plag. -- trace (altered)

musc. -- trace (highly altered)

tour. trace (green)

(any)

WC 515

Grains: 75%

size: $\leq 62-80 \mu$ (silt to very fine sand)

shape: subangular to subrounded

sorting: moderate

composition: qtz - 67% (omigraine)

K-feld - 4% (pretty clean)

musc - trace - 1%

±1% opaque (oxides,
pyroxene)

biotite - 1% and chlorite

tour. - trace

perthite - trace (clean)

plag. - trace (clean)

chert - trace - 1%

(over)

WC 520

Grains: 70%

size: $\leq 62-80 \mu$ (silt to very fine sand) & 350-500 μ (med. s. mud chips)shape: ~~sub~~ subangular to angular

sorting: moderate

composition: qtz - 57% (omigraine)

K-feld - 7% (omigraine) ^{all thin to med. (to thin) and good} (dirty, very coarse, plates)

musc. - trace (pretty clean)

trace
opaque - mag.

musc - 1%

biotite - 1%, chlorite trace - trace

tour. - trace (light-gray green) & (green)

plag. - trace

chert - 1% ; mud chips - 3%

(over)

WC 530(maybe take photo of thin x-section
near right end of slice)

Grains: 70%

size: $\leq 62-73 \mu$ (silt to very fine sand)

shape: subangular, subrounded (maybe some rounded)

sorting: moderate

composition: qtz - 58%

K-feld - 10%⁺ (prismatic clin) (also some altered)

plag - trace (dirty)

mus. - trace

hornbl. - trace (green to brown - euhedral x-section)

glaucophane - trace biotite - 1% and chlorite

chert - 1%

pyrite - trace zircon - trace

(avg)

WC 540

(take photo of granite)

Grains: 75%

size: 88-125 μ (very fine sand) + 1000-1400 μ (very coarse sand)

sorting: very poor

shape: subrounded to subangular

composition: qtz - 68% some polygonized

K-feld - 2%

granite fragments - 3% (med. sized) (qtz)

mus. 1% } in feld, smelted aggregates (incl chert?)

biotite 1%

zircon - trace (rounded) & subangular

chlorite - trace

No Tour. Found

Matrix: 25%

composition: qtz 20%, K-feld 5%

difficult to tell qtz because so well compacted + cemented

W60

WEST GLACIER

Grains: 65%

size: 62-125µ (very fine sand)

shape: subangular to ^{angular} subrounded, some are subrounded

sorting: moderate to poor

composition: qtz -- 53% (a lot of overgrowths)
~2% sericitized K-feld -- 7% (obvious thick overgrowths, quite dirty)
mica -- 1% (not too dirty)

opaque (3%) -- mag, oxides, plag. -- trace (pretty dirty)

leucosene (some in
crude layers) perthite -- 1% (dirty)

zircon -- trace (subangular), rounded

tour -- trace (brown to green grains), bluish green

biotite -- trace, chlorite

musc -- trace

W619

Grains: 75%

size: 88 - 250µ (fine to med. sand)

shape: subangular to subrounded (some circular)

sorting: moderate to poor

composition: qtz -- 66% (overgrown)

3% hematite coated K-feld -- 6% dirty (overgrown)

mica -- 19% + dirty (overgrown)

perthite -- 1% dirty (overgrown)

opaque (trace) mag
oxides, leucosene plag. -- trace dirty

musc -- trace

trace 1 tiny grain

musclips -- trace

zircon -- trace (rounded, subrounded)

W6 41

(take photo of broken round gm. Ⓟ)

Grains: 70%

Size: 88-710μ (very fine to coarse s.s.)

Shape: rounded to subrounded

Sorting: poor

composition: qtz -- 60% (overgrown)

K-feld -- 9% (maybe overgrown tho mostly clean, or (replaced by calcite))
(clean)

opaque (trace) oxides
scattered

plag -- trace (dirty)

pyrite -- 1% (dirty)

Volc. flag
(one grain)

tan -- trace (rounded qtz & subangular gm).
blue-green & brown, small overgrown

clast -- 1%

zircon -- trace (subrounded)

W6 60

(take photo of broken round gm. Ⓟ)

Grains: 75%

Size -- 88-177μ (v.f. to f.s.) layers & 250-500μ (med. s.s.) layers

Shape -- subrounded to rounded (finer grains a little more angular)

Sorting -- moderate to poor

composition -- qtz -- 64% (overgrown)

+ 5% altered

K-feld -- 5% (overgrown) (some dirty, clean)

micro -- 1% (pretty clean)

opaque (±1%) may
scattered

pyrite -- 1% (a little dirty)

minerals -- trace

tan -- trace (brown) subrounded (some anhedral like in TIR's)

zircon -- trace subrounded to subangular (one uniaxial)

Matrix: 25%

(dirty)

WG 92 (Maybe good clay sample) (take photo of broken round grain (f) granite frag -- fine s. ind. grains)

Grains: 55%

size -- 62 to 710 μ (v.f. to coarse ss - mixed layers in mud)

shape -- subrounded & rounded, smaller grains more angular

sorting -- very poor

composition: qtz -- 43% (more qtz in large grain size fraction)

1% hematite stained K-feld -- 8% (pretty clean, hematite coats)

plag -- 1% (dirty)

opaque (1%) magnetite in crude layers

pyrite -- 17% (a little dirty)

tour. -- trace (green) (subangular)

clit -- trace

musc. -- trace chlorite -- (trace) - green

zircon -- trace -- (subrounded)

1% granite (qtz grains) (fine)

WG 102

Grains: 70%

size -- $\leq 62-380 \mu$ (silt to very fine sand), a few 177-250 μ grains (f. sand)

shape -- subangular

sorting -- moderate

composition -- qtz -- 55%

K-feld -- 10% (clean)

plag -- 1% (slightly dirty)

opaque (trace) magnetite, pyrite -- trace (clean)

leucosomes

tour. -- trace (dark green) (subangular, subrounded?)

zircon -- trace (subrounded)

musc. -- 1%

sillite -- 1% chlorite -- 1%

Note: No muscovine found

WG 120

(Take photo of broken round grains) (3)
 (Take photo of T.R.F., granite.)

Grains: 70%

size - $\leq 62.83\mu$ - 1% 177-350 μ (very fine to coarse sand)
 shape - subangular
 sorting - moderate
 composition - qtz - 51% (some overgrowths) mostly in air of
 calcite may replace K-feld - 10% (some overgrowths)
 10-20%? plagioclase - trace
 perthite - trace
 * \Rightarrow tour. - 4% (euhedral, subhedral) TRF
 opaques (trace) musc. - trace
 leucosane, calcite mudchips - 5% (calcaneous) chlorite - trace
 granite - 2 grains seen (coarse)
 zircon - rounded - trace

WG 127

(Take photo of granite broken round grain)

* Photo of poly qtz + qtz cemented together TRF s.s. frag.

Grains: 50%

size, 62-710 μ (very fine to coarse s.s.)
 shape: subangular \rightarrow subrounded (some rounded)
 sorting - poor
 composition: qtz - 37% (some overgrowths) some poly
 maybe several % K-feld - 5% (some overgrowths)
 replaced by calc. perthite - 1% dirty, some clean
 muscovite - trace dirty
 S.S. frag - trace zircon - trace musc. trace (5% to 10% of slab - opaque)
 mudchips - 1% (qtz, plagioclase)
 granite - 2% chlorite (trace)
 chert - trace
 tour. trace (green) subhedral, subangular, euhedral

TRF trace

WG 143

Grains: 65%

size - 62-88 μ (very fine sand), 50 & 177-250 μ (fine sand) qtz grains.

shape - subangular, angular

sorting - moderate

composition - qtz - 50%

K-feld - 10% (pretty fresh)

pyrite - trace clean

opaque (1%) magnetite micro - trace pretty clean

in mud layers red mud chips - 3%

musc. - 1%

chlorite - trace

zircon - trace subangular

tour. - trace subangular

WG 160

Grains: 60%

size - 62-88 μ (very fine sand)

shape - subangular to angular

sorting - moderate

composition - qtz - 48% (amorphous)

K-feld - 5%? (pretty clean)

pyrite - 1% (clean)

opaque (1%) tour. - trace (green) subangular, some brown ~~and~~ ^{color, green}

magnetite, scattered musc. - 1%

red mud chips - 5%

zircon - trace (subangular) anhedral

biotite - trace & chlorite

micro-limo - trace (clean) (amorphous)

WG 180

Grains: 70%

size: $\leq 60-80\mu$ (silt to very fine sand)

shape: subangular to angular

sorting: moderate

composition: qtz - 55% (some small overgrowths)

K-feld - 10%+ (pretty clean)

perthite - trace

opaque (2%) magnetite zircon - trace (subrounded)

scattered in mud layers. mica - 2%

biot - trace and chlorite

tour - $\leq 1\%$ (green, bluish green, grey-green) subangularWG 200

Grains: 75%

size: $60-80\mu$ (very fine sand) w/ $350-500\mu$ (med. sand) mud chips

shape: subangular to angular

sorting: moderate

composition: qtz - 55% (overgrown)

K-feld - 15% (overgrown, pretty clean)

perthite - 2% (vacuolated)

there seems to be compositional layering of K-feld (can be very high % thin) mica - trace (clean) (some overgrowths)

zircon - trace (subrounded)

tour - trace (green, subangular) (subangular)

mud chips - 1%

opaque (1%) mica, red mud chips - 3%

leucosane, scattered

mica - 1% chlorite - trace

WG 220

Grains: ~~65~~ 70%

size: $\leq 62-88 \mu$ (ilt to very fine sand)

Shapes: subangular to angular

sorting: moderate

composition: qtz - 58% (overgrown) } slide not stained
K-feld - 5%? (dirty, overgrown) } approx. %

plithite - 1% (pale, clean, very dirty)

apatite (trace)

microcline - trace (clean or slightly dirty)

magnetite, scattered

plag. - trace (pretty clean)

tour. - trace (green to gray green) subangular, euhedral

zircon - trace (subrounded)

musc. - 1%

chlorite - trace

med
muscl. chips - 5%