



Inland Water Bodies in China: New Features Discovered in the Long-term Satellite Data

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1 **Inland Water Bodies in China: New Features Discovered in the**
2 **Long-term Satellite Data**

3

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24 Running title: Inland water bodies in China

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28 **Water bodies (WBs) - lakes, ponds, and impoundments, provide essential**
29 **ecosystem services for human society, yet their characteristics and changes over**
30 **large areas remain elusive. Here we used unprecedented data layers derived from**
31 **all Landsat images available between 1984 and 2015 to understand the overall**
32 **characteristics and changes of WBs between two epochs (i.e., 1984-1999 and**
33 **2000-2015) in China. Results show that the abundance estimate of WBs greater**
34 **than 1 km² and the total WB surface area were 0.3-1.5 times and 0.2-0.5 times**
35 **more than the previous estimates, respectively. The size-abundance and**
36 **shoreline-area relationships of WBs in China conformed to the classic power**
37 **scaling law, in contradiction to most previous studies. WB changes with various**
38 **occurrence probabilities show widespread co-existence of disappearance of**
39 **existent and emergence of new WBs across China driven primarily by human**
40 **activities and climate change. Our results highlight the importance of using**
41 **appropriate long-term satellite data to reveal the true properties and dynamics of**
42 **WBs over large areas, which is essential for developing scaling theories and**
43 **understanding the relative impacts of human activities and climate change on**
44 **water resources in the world.**

45 **Keywords: inland water bodies; size-abundance; fractal dimension; climate**
46 **change; land use change**

47

48 **Significance Statement**

49 Inland water bodies (WBs) provide essential ecosystem services for human society,
50 yet their characteristics and changes over large areas remain elusive. We used
51 unprecedented data layers derived from Landsat imagery to quantify and decipher the
52 spatial distributions and contemporary changes of WBs in China. The results shed new
53 light on the characteristics (e.g., abundance, and size-abundance and shoreline-area
54 relationships) and spatiotemporal dynamics of WBs. This study highlights the
55 importance of using appropriate long-term satellite data to reveal the true properties and
56 dynamics of WBs over large areas, which is essential for developing theories and
57 understanding the predominant impacts of human activities and climate change on
58 water resources not only in China but also in other regions of the world.

59

60 **Introduction**

61 The provisioning of freshwater is an essential ecosystem service for humankind,
62 and such services include water supply, biomass and food production regulation,
63 communication and transportation, wildlife habitats, biogeochemical cycles, and
64 climate regulation (1-3). Fresh water bodies (WBs) are under constant change because
65 of the impacts of anthropogenic activities and climate change (4). Monitoring the
66 dynamic changes of WBs and understanding their causes and consequences have
67 important implications to the resilience of social-ecological systems (5, 6).

68 The abundance and area of WBs over large regions are critical information for
69 water resource management but remain highly elusive. For example, the number of
70 lakes larger than 0.1 km² at the global scale was estimated to range from 246,146 to
71 4,123,551 (7, 8), more than one order of magnitude difference among them. The
72 estimated area of lakes larger than 0.1 km² varied from 2,428,100 km² to 4,080,000 km²
73 respectively, and the largest difference was 68%. These huge differences among various
74 studies call for innovative investigations to reduce the uncertainties in the number and
75 area of WBs.

76 Biogeochemical processes in WBs represent major components in the cycle of
77 carbon, nutrients, and other materials at the regional to global scales (9, 10). Estimating
78 biogeochemical processes such as greenhouse gas fluxes from WBs from regional to
79 global scales has been challenging because these processes are size dependent (11).
80 Finding the size-abundance relationship therefore has been a persistent effort for over a
81 century to facilitate the scaling of heterogeneous processes across various WBs (12).
82 Although a power law relationship has been reported in many studies, deviations exist,
83 particularly for small WBs (8, 13-15) and in mountainous regions (16) where
84 topographic relief, geology, and complexity of landscapes are thought to affect the
85 goodness of fit to the power scaling law (14). The existence of the power law and its
86 dependence on topography and geology needs to be evaluated more extensively.

87 Water-resource management and planning as well as climate change adaptation
88 rely heavily on the knowledge of long-term trends and spatial dynamics of surface
89 WBs (1) . However, progress on characterizing the dynamics of WBs over large

90 regions has been hindered by data availability, dynamic nature of WBs, complexity of
91 driving forces, and geospatial heterogeneity of environments (8). Most previous studies
92 do not have enough data points in time to capture the temporal changes of WBs, and
93 consequently the overall dynamics and size dimensions of the WBs, particularly the
94 intermittent WBs, are not well represented (17, 18). Recently, an unprecedented effort
95 by Pekel *et al.* (2016) has generated a global surface water dataset that describes the
96 changes in WBs from 1984 to 2015 at 30 m resolution (19). Although the overall global
97 patterns of WB changes have been analyzed, more detailed and elaborate analyses are
98 sorely needed at regional and continental scales to better understand the spatial and
99 temporal patterns of WB changes and to test the validity of existing theories such as the
100 controversial power scaling law for the size-abundance relationship (12, 20, 21). China
101 represents an ideal place for such as study because of its vast heterogeneous territories,
102 intensified human activities in recent decades, and large uncertainties in the
103 characterization of its dynamic water resources (22).

104 **Abundance, area, and scaling of WBs**

105 Our estimate of WB abundance in China exceeds all previous estimates (*SI*
106 *Appendix*, Table S1 and Table S2). The abundance of WBs ($> 1 \text{ km}^2$) is estimated to be
107 6,821 in our study, 1.3 times the previous highest estimate (5,535) and 2.5 times the
108 lowest estimate (2,693) (7, 17, 23, 24). The number of WBs larger than 0.0036 km^2 in
109 China is 688,617 in this study, much higher than 275,029 estimated previously (24).
110 Our estimates of WB area in China are higher than those from previous studies. For
111 example, our estimate of total area of WBs larger than 0.0036 km^2 in China is $134,158$
112 km^2 , 23% higher than $109,102 \text{ km}^2$ estimated by Yang *et al.* (2014) (24). The total area
113 of WBs larger than 1 km^2 is estimated to be $123,342 \text{ km}^2$ in our study, 22-51% larger
114 than the previous estimates (7, 17, 23, 24). Similar differences can be found across the
115 regions in China as well. For example, the number of medium and large WBs in the
116 YTR (Yangtze River basin, acronyms for the river basins are listed in *SI Appendix*,
117 Table S3) is estimated to be 1952 in this study, much higher than 1395 estimated by
118 Yang *et al.* (2013) (25). The medium and large WBs in the TP amount to 1289 in our
119 study, which is very close to Mao *et al.* (2018) (1291) (26) but higher than Wan *et al.*

120 (2014) (1055) (27) and Lu *et al.* (2011) (1044) (28). We found that the total area of WBs
121 larger than 0.0036 km² in the YTR is 36,710.83 km², which is 48% higher than Yang *et al.*
122 *al.* (2013) (24,842 km²) (25). Our estimate of the WBs larger than 1 km² in the TP is
123 56,247.41 km², much higher than Mao *et al.* (2018) (46,264.5 km²) (26) or Wan *et al.*
124 (2014) (41,831 km²) (27).

125 The overall exceedance of our WB estimates can largely be attributed to the
126 differences in methodology. Conventionally, WBs over large areas are characterized
127 using one or a few snapshots of remotely sensed images, which could not consistently
128 capture intermittent WBs and historical maximum water extents (29). Previous results
129 are therefore contingent to the snapshot-specific climate and hydrological conditions at
130 the time that images were taken.

131 The size-abundance relationships in China conform to the power law with a Pareto
132 coefficient of -0.835 (*SI Appendix*, Fig. S1F and Table S4), similar to -0.85 reported for
133 the contiguous United States (CONUS) (30), but higher than that for the world in 2004
134 (-0.992) (7) and in 2016 (-1.054) (17). At a regional scale, the Pareto coefficient in the
135 TP is -0.662, the highest among all regions, signifying a higher fraction of large WBs in
136 the TP compared to other regions (*SI Appendix*, Table S5). In contrast, the scaling
137 exponent in the PR is -0.941, the lowest among all regions, representing a relatively
138 higher share of small lakes. Comparatively, the log-transformed WB size-abundance
139 distribution in China presented a shallower slope or Pareto coefficient than those in the
140 world and CONUS, suggesting that small WBs are relatively more abundant in China
141 than in the CONUS and the world (Fig. 1C). Our results showing the power scaling law
142 between abundance and size of WBs in China, despite its complex geography and
143 topographic characteristics, are different from a number of recent studies that have
144 found deviations from the power scaling law, particularly for small WBs (8, 13-15).
145 Seekell *et al.* (2013) show that the size-abundance relationship in a mountainous region
146 (Adirondack, USA) did not follow the power scaling law while it fit well in a flat region
147 (Gotland, Sweden), and argue that the topographic relief, geology, and complexity of
148 the landscapes could impact how well data fits the simple power scaling law (13). It is
149 quite surprising to see that our results fit the power scaling law well in China even
150 though some regions such as the TP, and northwestern and southwestern China are

151 famous for their high topographic relief and steep slopes that are suspected to cause
152 deviations from the power scaling law (Fig. 1B, *SI Appendix*, Fig. S2).

153 The shoreline fractal dimension (SFD) derived from dimensional analysis (log
154 perimeter-log area analysis) is often used to indicate the influences of human
155 modifications or geology on the shape of WBs (31). The SFD ranges from 1.178 in TP
156 to 1.321 in PR (Fig. 1B, *SI Appendix*, Fig. S3). Theoretically, the fractal dimension of
157 size-abundance (D) ($D = 2 * \text{abs}(c)$, c is the scaling exponent or Pareto coefficient
158 mentioned above) is supposed to be similar to the SFD (13, 32). We compared the D
159 and SFD derived from China and found a strong linear relationship between these two
160 fractal dimensions ($D = 3.13 \times \text{SFD} - 2.27$, $R^2 = 0.73$, $P < 0.05$). However, the slope is
161 significantly different from 1, suggesting that our results do not support the theoretical
162 equivalency of these two fractal dimensions (Fig. 1D). The deviation of D and SFD
163 from the theoretical equivalency might be caused by increased frequency of smaller
164 WBs and/or effort in shaping WBs more regularly and reducing the sinuosity of WB
165 boundaries. Economic growth and aquaculture intensification have propelled the
166 increase of smaller WBs and geometric regularization of the WBs, particularly in the
167 southeastern and coastal parts of China (33).

168 <Fig. 1 roughly here>

169 **Various changes between two epochs**

170 Overall changes of WBs in China between the two epochs (i.e. 1985-1999 and
171 2000-2015) are shown in Fig. 2, which reveals that 53.9% of the WBs remain
172 unchanged nationwide, and further shows the various degrees of change between the
173 epochs. A total of 5197.09 km² or 3.92% of the WB area in China disappeared
174 (absolutely decreased); whereas the very likely (with 75-99% probability), likely (with
175 25-74% probability), and unlikely (with 1-24% probability) decreases accounted for
176 1.6%, 6.23%, and 7% of the total WB area, respectively (Fig. 3, *SI Appendix*, Table S6).
177 A total of 9040.26 km² were converted from non-water surfaces into WBs (absolutely
178 increased), corresponding to 6.81% of the total WB area in China. At the same time,
179 very likely, likely, and unlikely increase occurred on 3.76%, 8.18%, and 5.71% of the

180 total WB area, respectively. The large WBs ($>10 \text{ km}^2$) explain most of the very likely,
181 likely, and unlikely increase or decrease, followed by the small WBs ($< 1 \text{ km}^2$) (*SI*
182 *Appendix*, Fig. S4 and Table S7). The WB changes vary widely across regions (Fig. 2).
183 For example, the highest absolute increases in area in percentage (13.72%, 16.41%, and
184 9.69%) are found in the TP, NWR, and SER, respectively. The region changed the most
185 was the HR since only 11.74% of WBs remained unchanged, and the most stable region
186 is the SWR with 75.99% persistently covered by water. The largest absolute decreases
187 in WB occurrence in percentage are found in the HR, SHR, and YTR account for
188 14.26%, 7.83% and 5.47% of their total WB areas, respectively.

189 <Fig. 2 roughly here>

190 Our analysis shows a net gain (absolute change) of $3,843 \text{ km}^2$ in WB between the
191 two epochs. Although some previous studies have quantified an increase (34, 35),
192 decrease (22, 24) or net change of the WB area in China, they are not comparable to our
193 estimate because of the differences in time span and definitions of change. Our change,
194 following Pekel *et al.* (2016), is defined as the difference between two epochs or
195 periods while it is defined as the difference between two time points or years in other
196 studies (19). Consequently, our estimate of change likely represents permanent changes
197 whereas the estimate using other definitions would be more likely influenced by the
198 interannual variability in the climate and hydrologic conditions. Starting with a dry year
199 and ending with a wet year would contain false signals of WB increase, for example,
200 when using the second definition.

201 **Driving forces of WB changes**

202 WB changes across the country show great divergent trends (Fig. 3), which were
203 most likely driven by multiple compounding drivers such as climate change and direct
204 human land use change activities (22, 35) (Table 1). Climate change has been shown to
205 have a major impact on WB changes, particularly in the TP and NWR regions. Because
206 of the TP's high elevation and sparse population, its WBs are predominantly influenced
207 by climate change. Despite a decline only some areas, the widespread increase of WBs
208 in the TP is primarily entailed by multiple climate-related factors, including the

209 precipitation, temperature, glacier, and permafrost (36, 37). On the one hand, the
210 climate change in the TP is prevalently accelerating, as the mean annual precipitation
211 has increased 20 mm and mean annual air temperature has warmed 1.6°C from 2000 to
212 2014 (26), resulting in the expansion of existing WBs and formation of new WBs by
213 melting glacier and snow in high mountains. On the other hand, permafrost degradation
214 and WB shrinkage have also been observed in some localized areas with drier climate
215 and reduced precipitation, especially in the area between the Himalayas and Tanggula
216 Mountains and in north of the Hengduan Mountains (26). Similarly, increasing WBs are
217 observed in the less populated mountainous areas in the NWR region caused by glacier
218 melting and increase in precipitation (38). The dry climate and drought events in the
219 arid regions (e.g., Inner Mongolia Plateau and part of the Xinjiang Uygur Autonomous
220 Region) have also contributed to WB shrinkage (35, 39).

221 <Fig. 3 roughly here>

222 <Table 1 roughly here>

223 Land use change activities have contributed to the expansion and/or contraction,
224 disappearance, and emergence of WBs at different locations throughout the country.
225 For example, dam constructions can substantially change WB in the upstream and
226 downstream regions (34, 40). Using the YTR basin as an example, both the numbers
227 and surface areas of the reservoirs in the YTR have increased rapidly during past
228 decades, and the impoundment of the great Three Gorges Dam alone formed a new
229 large reservoir with an area of 1,084 km² in 2003 when it was put into full operation
230 (41). Intensification of agriculture also brings heavy pressure to WBs (40, 42). For
231 example, the northeast region has experienced a rapid increase (319.0%) in irrigated
232 cropland area in the past three decades (22, 43). Aquacultural change affects the
233 dynamics of WBs in China, particularly in the east, south central, and southeast regions
234 where water resources and climate are suitable for aquaculture (33). Land reclamation,
235 or WB impoldering (i.e., conversion of WBs into other land uses), has resulted in the
236 loss or shrinking of WBs (38, 41). Many natural wetlands and WBs have been drained
237 and transformed into croplands, grasslands, and woodlands (22, 43), resulting in the
238 loss of about 50,360 km² or 14.18% of the wetlands across the country from 1990 to
239 2000 alone. Many small WBs are infilled in the process of urban expansion (41).

240 Wuhan, the capital city of Hubei province and once known as the city with thousands of
241 lakes, has lost about 70% of its lakes to reclamation and development between 1994
242 and 2014 (44). Various policies are being put forward by Chinese governments to
243 promote wetland conservation and restoration in recent years (45). After joining the
244 Ramsar Convention in 1992, the Chinese government has implemented the National
245 Wetland Conservation Plan, trying to prevent natural wetlands from further loss and
246 degradation (43). The government had funded more than 200 pilot programs with 20.7
247 billion US\$ during 1990-2007 to restore degraded wetlands and create new ones to
248 make amends for lost wetlands (46).

249 Assessing the dynamics of WBs over large regions has been a major challenge
250 because of the constraints in data availability, heterogeneity of environments, and the
251 complexity of driving forces (8). Many studies have mapped WBs with multi-spectral
252 and hyperspectral images and built databases comprising locations, sizes, and
253 distributions of WBs over large regions (17, 47). However, these studies often suffer
254 from two weaknesses. First, they usually do not have enough data points in time to
255 capture the temporal changes of WBs. For instance, results are often generated from
256 subsamples of available remote sensing data at 5-year or longer time intervals (48),
257 representing only snapshots of WB conditions in the sampled years. As a result, the
258 overall dynamics and size dimensions of the WBs, particularly the intermittent WBs,
259 are not well represented (17, 18). Second, the minimal detectable WB size, defined by
260 the resolution of remotely sensed data (e.g., Terra /Aqua MODIS), is usually not fine
261 enough to detect the small WBs that are biologically more active than the large ones
262 (49). Our study shows that the use of dense long-term satellite data at appropriate
263 resolution has the potential to overcome some of the difficulties and reveal
264 unprecedented distributional details and dynamics of WBs over large areas. This
265 capability is essential for validating the relevant theories such as the power scaling
266 relationship between size and abundance and for adequately scaling many hydrological
267 and biogeochemical processes (e.g. the carbon cycle) from individual aquatic systems
268 to regional and global scales (50). It is also important to map the details of WB changes
269 and understand their causes and consequences to improve water resource management
270 and planning.

271 **Materials and Methods**

272 **Study area and region boundaries**

273 China (73°33'-135°2'E and 3°52'-53°33'E) is our study area (Fig. 1A). To
274 investigate the regional differences of WBs, we delineated China into 11 large regions,
275 primarily following the region boundaries from HydroSHEDS
276 (<http://www.hydrosheds.org>, last accessed in 2018) (51) and National Lake-Watershed
277 Science Data Center (<http://lake.geodata.cn>). The abbreviations of the regions are listed
278 in Table S3.

279 **Extraction of lakes, ponds, and impoundments**

280 The Global Surface Water Dataset (GSWD), published by the European
281 Commission's Joint Research Centre, was the basis for our study
282 (<https://global-surface-water.appspot.com>, last accessed on December 12 in 2018).
283 GSWD was generated using 3,066,102 (1823 terabytes of data) scenes from Landsat 5,
284 7, and 8 acquired between 16 March 1984 and 31 December 2015 (19). To address the
285 challenges in separating water from other surfaces on the global scale over multiple
286 decades, expert systems, visual analytics, and evidential reasoning techniques were
287 exploited. Details about the implementations of these techniques can be found in Pekel
288 *et al.* (2016). The resultant maps constitute the long-term water history that show the
289 “when and where” of the water presence during the observation period. At the same
290 time, the Water Occurrence, Occurrence Change Intensity, Seasonality, Recurrence,
291 Transitions, Maximum water extent are also shown in the dataset. The accuracy of the
292 water maps was assessed by the developers of GSWD in term of errors of omission and
293 commission at the pixel scale (i.e., 30 m) using a total of 40,124 control points
294 distributed both geographically (globally), temporally (across the 32 years). Overall,
295 errors of commission were less than 1% and omission less than 5%.

296 The data layer showing the maximum water extent (MWE) (all locations ever
297 mapped as water) during the 32 years was used to extract the locations and maximum
298 extents of WBs in our study. GSWD contains three types of waters: artificial paddy
299 fields, rivers, and others (i.e., lakes, ponds, wetlands, and impoundments). As our goal

300 was to investigate the characteristics and dynamics of lakes, ponds, wetlands, and
301 impoundments, rivers and rice paddy fields were removed from the MWE map using
302 the HydroSHEDS data through overlay operation in ArcGIS v10.2 (ESRI, Redlands,
303 CA, USA) and manually corrected by visual inspection against Google Earth
304 high-resolution images. Consequently, the temporal changes of rivers and streams were
305 also effectively excluded from this analysis (*SI Appendix*, Text S1). Paddy fields in the
306 MWE were removed by referring to the Global Land Cover dataset (GLCD)
307 (<http://www.globallandcover.com>, last accessed in 2018).

308 **WB change between two epochs**

309 GSWD provides water occurrence change intensity (OCI) between two epochs
310 (i.e., 1984 to 1999, and 2000 to 2015), which was derived from water occurrence
311 difference between homologous pairs of months (19). Specifically, the water
312 occurrence difference for each homologous pair of month between epochs was
313 calculated, and then differences between all homologous pairs of months were
314 averaged to create the surface water OCI map between the epochs. The OCI map shows
315 where surface water occurrence increased, decreased, or remained invariant, providing
316 a summary of the location and persistence of water in space between the two epochs.
317 The OCI map represents the degree of change as a percentage with values ranging from
318 -100% to 100%; positive values indicate increase in occurrence while negative ones
319 show loss of occurrence, and 0% suggests no change in occurrence between the epochs.
320 To summarize the results, OCI values were grouped into several classes: 75-99% (high
321 OCI), 25-74% (medium OCI), 1-24% (low OCI), 0% (unchanged), and 100%
322 (absolutely changed). In essence, these OCI classes effectively represent the
323 probabilities or likelihoods of conversion from non-water to water (positive OCI) or
324 from water to other surfaces (negative OCI) between two epochs, which resulted in nine
325 generalized probability classes of conversion between water and other land covers:
326 absolutely decreased (OCI= -100%), very likely decreased (OCI= -99% to -75%),
327 likely decreased (OCI= -74% to -25%), unlikely decreased (OCI= -24% to -1%),
328 unchanged (OCI= 0%), unlikely increased (OCI= 1% to 24%), likely increased (OCI=
329 25% to 74%), very likely increased (OCI= 75% to 99%), and absolutely increased
330 (OCI= 100%).

331 The minimum water surface area detectable is the pixel size of Landsat (0.0009
332 km²). To facilitate comparison with other studies, whenever necessary, WBs were
333 classified into three categories of small (< 1 km²), medium (1-100 km²), and large (>
334 100 km²) following the classification scheme of most previous studies (35, 39). All data
335 processing and analysis were performed using ARCGIS 10.2 (ESRI, Redlands, CA,
336 USA) and Python 3.6.

337 **Data Availability**

338 All data used in support of this manuscript are available in Figshare
339 (<https://figshare.com/s/9e931d6db628e3f96689>; DOI: 10.6084/m9.figshare.9959516)
340

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Figure Legends

Fig. 1. Boundaries of regions, morphological and size-abundance scaling of water bodies in China. (A) locations and boundaries of the 11 regions in China; (B) shoreline fractal dimensions (SFD) of TP. (C) size-abundance distributions of water bodies in China, CONUS, and the world. (D) shoreline fractal dimensions (SFD) and fractal dimension of size-abundance (D) in China and the regions.

Fig. 2. Regional occurrence changes of WBs at various confidence levels. Subsets of four selected cases show WB changes. A, Dongting lake in the YTR, centered at 112.8° E, 29.1°N. B, Hulun lake in the SHR, centered at 117.5°E, 48.9°N. C, Seling Co lake in the TP, centered at 89.1°E, 31.8°N. D, Hongze lake in the HuR, centered at 118.6°E, 33.2°N. E, epochal changes in different regions.

Fig. 3. Epochal changes of water bodies in China and regions at various confidence levels. Changes are classified as Absolutely Decreased (AD), Very Likely Decreased (VLD), Likely Decreased (LD), Unlikely Decreased (ULD), Unchanged (UC), Unlikely Increased (ULI), Likely Increased (LI), Very Likely Increased (VLI), and Absolutely Increased (AI).

Table 1. Summary of main driving forces in different regions.

Region	Main driving forces	References
TP	variations of precipitation; warming-induced glacier-melting and permafrost degradation	Mao et al. 2018b, Zhang et al. 2017a
PR	land reclamation; waterbody impoldering; urbanization	Li et al., 2018; Guo et al., 2015
YTR	dam constructions; aquacultural change; land reclamation or impoldering; urban expansion; sand mining and illegal sand dredging	Cai et al., 2016; Xie et al., 2003
YR	intensification of agriculture; intensified groundwater exploitation	Shi et al., 2011; Zhang et al., 2003
HR	rapid urbanization; land reclamation or waterbody impoldering	Yang et al., 2011; Tian et al. 2016
HuR	dam and floodgate construction; land reclamation	Hu et al., 2008; Zhang et al., 2009
SHR	intensification of agriculture; irrigation; wetland draining or restoration	Xie et al., 2018; Ma et al., 2010
NWR	persistent drought; intensification of agriculture (irrigation); mining	Fang et al., 2018; Tao et al., 2015
SWR	variations of precipitation; warming-induced glacier-melting	Wu et al. 2007; Mao et al., 2018b
SER	aquacultural change; land reclamation or waterbody impoldering;	Lu et al. 2015; Wu et al., 2018
LR	intensification of agriculture (irrigation); wetland conversion	Xie et al. 2018; Ma et al., 2010

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