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# A multi-lake comparative analysis of the General Lake Model (GLM): Stress-testing across a global observatory network

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## Highlights

- The General Lake Model (GLM) is stress tested against 32 globally distributed lakes.
- There was low correlation between input data uncertainty and model performance.
- Model performance related to lake-morphometry, light extinction and flow regime; deep, clear lakes with high residence times had the lowest model error.
- Predictions of temperature were less sensitive to model parameters than thermocline depth and Schmidt stability.

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183 **Abstract**  
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186 The modelling community has identified challenges for the integration and assessment of  
187 lake models due to the diversity of modelling approaches and lakes. In this study, we  
188 develop and assess a one-dimensional lake model and apply it to 32 lakes from a global  
189 observatory network. The data set included lakes over broad ranges in latitude, climatic  
190 zones, size, residence time, mixing regime and trophic level. Model performance was  
191 evaluated using several error assessment metrics, and a sensitivity analysis was  
192 conducted for nine parameters that governed the surface heat exchange and mixing  
193 efficiency. There was low correlation between input data uncertainty and model  
194 performance and predictions of temperature were less sensitive to model parameters  
195 than prediction of thermocline depth and Schmidt stability. The study provides guidance  
196 to where the general model approach and associated assumptions work, and cases where  
197 adjustments to model parameterisations and/or structure are required.  
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206 Key Words: lake model, stratification, GLM, model assessment, global observatory data,  
207 network science.  
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## 1 Introduction

Vörösmarty et al. (2000) urged the international “water sciences community” to work together in the collation and dissemination of hydrological data and modelling techniques to improve our understanding of freshwater ecosystems and “secure a more complete picture of future water vulnerabilities”. Lakes, in particular, are highly valued ecosystems as they provide important water and food resources, and numerous other ecosystem services (Wilson and Carpenter 1999). Human activities such as fresh water diversion and increased nutrient loading, in addition to indirect pressures from climate change, have led to an increased vulnerability of lakes on a global scale (Folke et al. 2004). These challenges have given rise to international networks of scientists such as the Global Lake Ecological Observatory Network (GLEON: [gleon.org](http://gleon.org)). Collaborative networks can take advantage of shared data, techniques, and expertise to enable scientists to address the ecological challenges facing lakes globally (Eigenbrode et al. 2007; Adams 2012; Goring et al. 2014). GLEON was initiated in 2005 as a grassroots science community with a vision to observe, understand and predict freshwater systems at a global scale (Weathers et al. 2013).

Collaboration between scientists and synthesis of data collected through international networks has led to advances in our understanding of how lake ecosystems respond to external changes and contribute to effective lake management on a local (Gal et al. 2009), regional (Read et al. 2014; Trolle et al. 2015) and global scale (O’Reilly et al. 2015). Analyses based on data from a broad spectrum of lakes across the globe have provided insight into metabolism and carbon cycling in lakes (Hanson et al. 2011; Solomon et al. 2013), the role of wind and heat exchange in lake physics (Read et al. 2012), the impact of climate change (Adrian et al. 2009), response and recovery of lakes to extreme events (Jennings et al. 2012; Klug et al. 2012), incorporation of high frequency data for model validation (Hamilton et al. 2015) and assisted in development of models (Staehr et al. 2010; Read et al. 2011; Kara et al. 2012; Hipsey et al. 2017). Further interrogation of the emerging multi-lake datasets offers the potential to advance our understanding of how lakes respond to pressures such as climate or land use change from the individual to global scales.

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303 The collaborative network also creates opportunities for developing and testing  
304 modelling tools. Aquatic ecosystem models are recognised as essential instruments to  
305 improve understanding of processes, analyse relationships, test hypotheses and predict  
306 the state of a system (Trolle et al. 2012). These models have evolved since the first  
307 attempts in the early 1920s, with a recent review of aquatic ecosystem models revealing  
308 the diversity of existing models from simple 0-D to complex 3-D coupled hydrodynamic-  
309 biogeochemical models (Janssen et al. 2015). This diversity creates challenges for  
310 integration and synthesis of model approaches (Mooij et al. 2010). The Aquatic Ecosystem  
311 Modelling Network (AEMON: <https://sites.google.com/site/aquaticmodelling/home>)  
312 originated to foster collaboration and improve model development, predictability,  
313 transparency and reliability. One of the major challenges facing modellers is how to  
314 develop generic models that can capture the diversity of ecosystems while allowing  
315 prediction with confidence of the processes of each system. In order to undertake  
316 analytical synthesis across multiple sites, there is a need to assess the transferability of  
317 the underlying model and standardise its structure, parameterisation, development and  
318 examination. While the need to develop a set of standards for model assessment and  
319 reporting is widely recognized (Bennett et al. 2013; Grimm et al. 2014), the ability to test  
320 these standards across multiple systems and highlight both strengths and limitations of a  
321 particular model remains a challenge.  
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336 For lakes and reservoirs in particular, one-dimensional (1-D) models that resolve vertical  
337 profiles of temperature and density have found widespread use due to their  
338 computational efficiency and minimal calibration requirements. The reduced complexity  
339 of 1-D models is advantageous whenever greater computational efficiency is needed, e.g.,  
340 in ensemble modelling (Trolle et al. 2014), model inter-comparison projects such as  
341 LakeMIP (<http://www.unige.ch/climate/lakemip>) (Stepanenko et al. 2010; Thiery et al.  
342 2014), probabilistic studies (Schlabing et al. 2014), long-term scenario analysis (Gilboa et  
343 al. 2014) or when linking lake models to global climate models (Balsamo et al. 2012) or  
344 catchment models (Hipsey et al. 2015). Moreover, lake managers and reservoir operators  
345 prefer models having a simpler application and often rely on 1-D models for this reason  
346 (Kerimoglu and Rinke 2013; Weber et al. 2017).  
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363 Here we introduce the Multi-Lake Comparison Project (MLCP) undertaken within  
364 AEMON. The MLCP is a community driven project, where teams of modellers simulate  
365 lakes using common approaches for model setup, assessment and analysis. The  
366 underlying purpose of the project was to bring together an international network of  
367 scientists and modellers with diverse experience in order to improve our ability to predict  
368 how lake ecosystems respond to external drivers. In the first stage, the MLCP took  
369 advantage of GLEON and AEMON member data from numerous, diverse lakes to stress  
370 test the recently developed General Lake Model (GLM) (Hipsey et al. 2017). GLM is a 1-D  
371 hydrodynamic model for use in a broad spectrum of enclosed aquatic ecosystems such as  
372 lakes, reservoirs and wetlands. The model is simple in nature and is based on assumptions  
373 that are common to previous model applications (Imberger and Patterson 1989; Hamilton  
374 and Schladow 1997; Coats et al. 2006). The model conducts a lake mass and energy  
375 balance to compute vertical profiles of temperature, salinity and density while accounting  
376 for the effect of inflows and outflows, surface heating and cooling, mixing and ice cover on  
377 the lake. GLM can be coupled with biogeochemical models to explore the impact of  
378 temperature, stratification, and vertical mixing on the dynamics of lake ecology (e.g.  
379 Snortheim et al. 2017).

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392 This paper summarises the first phase of the MLCP to develop and stress-test GLM. The  
393 stress-test involved applying a single standardised procedure for model set-up,  
394 simulation, performance testing and analysis to 32 lakes from across the global network.  
395 The main objective of this study was to undertake comparative analysis of model  
396 performance using an unprecedented diversity of lake types in order to advance our  
397 understanding of limnology and contemporary modelling practices. The specific aims of  
398 the study were to:

- 403 1. ascertain levels of model performance and relate it to model input uncertainty;
- 404 2. identify lake attributes (e.g. depth, inflows, and climate) that correspond with high  
405 (or low) prediction accuracy;
- 406 3. relate sensitivity of model output variables to changes in surface exchange, heating  
407 and mixing parameters that characterise 1-D lake models;
- 408 4. document the transferability of the model without recalibration of individual  
409 parameters among lakes, even where these lakes may strongly differ in their  
410 properties; and  
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423 5. provide guidance to lake modellers as to how to focus data collation and model  
424 application efforts to improve predictions for lake ecosystems.  
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426 To ease readability, this main section of the paper includes all text as well as tables and  
427 figures relevant to the major methodology and results from the study. Additional data  
428 have been provided in the following four appendices as supplementary material to the  
429 main study:  
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432 A, describing uncertainty error associated with the model set up;

433 B, extended results describing model performance;

434 C, extended results of the sensitivity analysis; and

435 D, a summary of acknowledgements for each lake.  
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## 441 **2 Methods**

### 442 **2.1 Study site selection**

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445 Lakes were not chosen a priori based on their attributes, but rather AEMON and GLEON  
446 members were invited to participate in the MLCP by volunteering details of their  
447 candidate lake to the group (shared via open access spreadsheet). The requirement for  
448 inclusion of a lake was based on the following three conditions:  
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- 452 1. sufficient temperature data were available for validation (at least 2 years of  
453 monthly/regular thermistor chain and/or profile data);  
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- 455 2. high-resolution meteorological forcing data from an on-lake buoy or local  
456 terrestrial based station were available; and  
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- 458 3. gauged or well-estimated inflows and outflows were available over the simulation  
459 period to form a reliable lake water balance.  
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464 Participants were also required to have a basic knowledge of lake modelling. Instructions  
465 as to how to set-up the GLM test cases, and a common binary executable (GLM v2.2.0)  
466 were made available for download from the Aquatic EcoDynamics (AED) website  
467 (<https://github.com/AquaticEcoDynamics/GLM>). Pre- and post-processing MATLAB  
468 scripts were provided to all participants to ensure a common model setup and assessment  
469 approach (<https://github.com/AquaticEcoDynamics/GLMm>), and all GLM lake setups  
470 were available to other members via a cloud-based, shared folder.  
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483 A total of 32 lakes was chosen for the analysis, with an alphabetic listing of the lakes and  
484 their physical characteristics in Table 1. Each lake is associated with a two letter  
485 abbreviated code, and for brevity when presenting model results, the lakes are frequently  
486 referred to by this code. To illustrate the range of sizes in the lakes included in this study,  
487 lake outlines have been drawn to scale in Figure 1. With the exception of lakes Geneva and  
488 Kinneret, all lake simulations were run for two years, with the start year and date  
489 indicated in Table A3. For Lake Geneva and Lake Kinneret, analyses were performed  
490 separately for two alternative 2-year time periods with significant differences in climate  
491 and inflows. For Lake Geneva, 2003 to 2004 had higher than average summer air  
492 temperatures, precipitation and inflows as well as an uncharacteristically high winter  
493 inflow in early 2004. In contrast, 2001 to 2002 experienced closer to the “normal”  
494 seasonal cycles of climate and inflows (Anneville et al. 2010). These simulations are  
495 referred to as Geneva03 and Geneva01 respectively. For Lake Kinneret, 1997 to 1998 had  
496 generally average climatic conditions (Bruce et al. 2006). In contrast, 2003 to 2004 had a  
497 rainy winter (Feb-Mar 2003, Jan-Feb 2004), large changes to lake level and lower than  
498 normal water temperatures (Berger and Telzch 2005). These simulations are referred to  
499 as Kinneret97 and Kinneret03, respectively.  
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512 Lake depths ranged from 2.4 to 440 m, and lake surface areas from 104,000 m<sup>2</sup> to  
513 579,000,000 m<sup>2</sup> (Table 1). A comparative plot of the hypsographic curves for each of the  
514 32 lakes shows diversity in lake size and bed slope (Figure A1). Annual average inflows  
515 ranged from 0 to  $3.3 \times 10^7$  m<sup>3</sup> d<sup>-1</sup> and residence times from 1 month to 67 years (Table A3).  
516 Lake elevation ranged from 209 m below to 4718 m above sea level (Table 1). Annual  
517 average air temperature ranged from below freezing (-9.1°C) to 22.4°C (Table A3). While  
518 the majority of the lakes in the MLCP are mid-latitude (both northern and southern  
519 hemisphere), two lakes are located in the Arctic (Emaiksoun and Toolik).  
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## 526 2.2 GLM set-up

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528 GLM has several configuration options for simulating surface heating, mixing and inflow  
529 and outflow (Hipsey et al. 2017). For this assessment, model set-ups were configured  
530 based on the site-specific conditions (e.g., hypsographic curve and number of inflows and  
531 outflows), but all simulations adopted the same model algorithms and parameters for  
532 mixing, surface heat fluxes, and ice cover. Default parameters adopted are summarised in  
533 Table 2.  
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545 All simulations were run for 2 years or 730 days starting with initial conditions in the  
546 winter or when the lake was most nearly well mixed. For the northern hemisphere lakes  
547 the start date was the 1<sup>st</sup> of January and for lakes located in the southern hemisphere the  
548 start date was set at 1<sup>st</sup> July. The initial conditions were taken from the closest field profile  
549 measurements to the start date. The standardised start date was chosen to simplify cross  
550 lake comparisons. For the majority of the lakes in the MLCP, mid-winter is also associated  
551 with complete mixing thus reducing error associated with uncertainty in initial profiles.  
552 A spin up period of 28 days was eliminated from model analysis to further reduce error  
553 associated with uncertainty in initial conditions.  
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561 Box plots are used to present monthly means and range of input data across all 34  
562 simulations (Figure 2). For input data for each lake, refer to references listed in Table 1  
563 and/or the institutions listed in Table D1. Inflows and outflows are also plotted as  
564 monthly averages based on time from the beginning of the simulation (Figure 3a&b).  
565 There are no seasonal patterns apparent in the monthly inflows and outflows averaged  
566 over the MLCP lakes due to the large variation in peak flow months.  
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572 While an effort was made to use lakes with high quality input data, lakes where input data  
573 had to be estimated were still selected for the MLCP in order to ensure a sufficient  
574 variation in lake characteristics. For seven lakes either inflow, outflow or both were  
575 estimated (Bourget, Emaiksoun, Feeagh, Mendota, NamCo, Stechlin and Woods) and the  
576 parameter of light attenuation ( $K_w$ ) was estimated for three lakes (Alexandrina,  
577 Muggelsee and Woods). Meteorological data for short wave radiation, air temperature,  
578 relative humidity, wind speed and precipitation were supplied either from an on lake  
579 station or the closest meteorological station to the lake. Long wave radiation was either  
580 measured directly (net or incident) or calculated by GLM using cloud cover data.  
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589 In an attempt to assess the errors associated with input data limitations, a qualitative  
590 weighting system was used to assess each input variable or constant, where a minimum  
591 score is associated with the best available input or observation data (Table A1). Table A2a  
592 lists the method of determining the hypsographic curve, distance from lake and frequency  
593 of meteorological data and observed data and method of determining inflow, outflow and  
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extinction coefficient for each lake in the MLCP. This information is used to determine the relative error scale associated with boundary forcing and observed data for each lake (Table A2b), where low refers to low uncertainty in forcing data and high indicates a higher level of error associated with model input. Input error associated with the determination of long wave radiation was not included in the error scaling method.

### 2.3 Model assessment approach

Measures of model fit used to evaluate model performance included five alternatives listed below. This set of measures of model fit enabled us to standardise comparisons among lakes, track trends in deviations from observed data (Bennett et al. 2013) and to compare with similar lake modelling studies previously published (e.g. Rigosi et al. 2010).

Measures of model fit were calculated as:

- 1) Root mean square error (*RMSE*):

$$RMSE = \sqrt{\frac{\sum_{i=1}^N (P_i - O_i)^2}{N}} \quad (2-1)$$

- 2) Model Efficiency (*MEFF*; Murphy, 1988; Nash and Sutcliffe, 1970):

$$MEFF = 1 - \frac{\sum_{i=1}^N (P_i - O_i)^2}{\sum_{i=1}^N (O_i - \bar{O})^2} \quad (2-2)$$

- 3) Correlation coefficient (*r*):

$$r = \frac{\sum_{i=1}^N (P_i - \bar{P})(O_i - \bar{O})}{[\sum_{i=1}^N (P_i - \bar{P})^2 \sum_{i=1}^N (O_i - \bar{O})^2]^{1/2}} \quad (2-3)$$

- 4) Percent relative error (*PRE*) :

$$PRE = \frac{\sum_{i=1}^N (P_i - O_i)/O_i}{N} * 100 \quad (2-4)$$

- 5) Normalised mean absolute error (*NMAE*) :

$$NMAE = \frac{\sum_{i=1}^N |(P_i - O_i)/O_i|}{N} \quad (2-5)$$

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663 where  $N$  is the number of observations,  $O_i$  and  $P_i$ , the “i<sup>th</sup>” observed and model predicted  
664 data and  $\bar{O}$  and  $\bar{P}$  the mean observed and model predicted data, respectively.  
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667 A further advantage of calculating alternative measures of model fit is that different  
668 methods of model evaluation highlight different aspects of model performance (Bennett  
669 et al. 2013). *RMSE* is a standard measure of the average deviation of simulated values from  
670 observations with values near zero indicating a close match and units that correspond to  
671 those of the variable. *MEFF* is the square of the deviation of simulated values from  
672 observations, normalized to the standard deviation of the observed data, such that one  
673 indicates perfect fit and zero indicates that the model provides equal predictive skill as  
674 the mean of the observed data. The correlation coefficient  $r$  gives an indication of the  
675 linear relationship between observed and predicted data and is the most common  
676 measure for assessing aquatic models (Arhonditsis and Brett 2004). *PRE* is a measure of  
677 the relative deviation of simulated from observed values and can be used to determine  
678 the bias in predictions (Bennett et al. 2013). Finally, *NMAE* is both normalised to the mean,  
679 enabling like comparisons between variables and is absolute so that under and over  
680 estimations do not cancel each other out.  
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691 Initial manual calibration focused on refining input data by adjusting the wind scaling  
692 factor and river inflow slope parameters for each lake (the river slope is indicated as  $\phi_{inf}$   
693 in Hipsey et al. (2017), and they are denoted as `wind_factor` and `strmbd_slope` in the  
694 configuration file, respectively). Wind factor adjustment was required where wind  
695 stations were located some distance from the lake and/or to account for wind sheltering  
696 effects (Markfort et al. 2010). River inflow slope was adjusted to correct the magnitude of  
697 momentum and entrainment associated with plunging inflows. For lakes where few or no  
698 light attenuation or Secchi depth readings were available,  $K_w$  was also adjusted until  
699 simulated thermocline depth matched that of observed data. Initial calibration was  
700 carried out until an *RMSE* (calculated for all observed temperature data over the  
701 simulation period) of less than 2°C was achieved.  
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711 We chose a range of thermal metrics to assess model performance at each site: observed  
712 full profile temperature data; epilimnion temperature; hypolimnion temperature;  
713 thermocline depth and Schmidt Stability (Idso 1973). Schmidt Stability ( $S_T$ ) and  
714 thermocline depth (*thermD*) were calculated for both model output and observed  
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thermistor data using Lake Analyzer (<http://lakeanalyzer.gleon.org/>), an open source software tool that computes indices of mixing and stratification for lakes and reservoirs (Read et al. 2011). The comparison of *thermD* calculations was included in the analysis as it is a simple, widely-used metric of mixed layer depth, while acknowledging the calculation of *thermD* can be challenging for weakly stratified and polymictic lakes. Also, the approach used in Lake Analyzer identifies the strongest thermal gradient, and may miss important thermal structure.  $S_T$  represents resistance to mechanical mixing due to the potential energy inherent in the stratification of the water column, calculated as:

$$S_T = \frac{g}{A_s} \int_0^{z_D} (z - z_v) \rho_z A_z dz \quad (2-6)$$

where  $g$  is the acceleration due to gravity,  $A_s$  is the surface area of the lake,  $A_z$  is the area of the lake at depth  $z$ ,  $z_D$  is the maximum depth of the lake, and  $z_v$  is the depth to the centre of volume of the lake, and  $\rho_z$  is the water density at depth  $z$ . While not used as a direct gauge of model performance, the daily Lake Number ( $L_N$ ) output as a GLM diagnostic parameter was also used in the cross lake comparison analysis as a measure of the validity of the one-dimensional assumption of the model.  $L_N$  balances the strength of stratification to wind induced mixing across the thermocline and is a measure of the potential for mixing across the thermocline (Imberger and Patterson 1989).

$$L_N = \frac{S_T(z_e + z_h)}{2\rho_h u_*^2 A_s^{1/2} z_v} \quad (2-7)$$

where  $z_e$  and  $z_h$  are the depths to the top and bottom of the metalimnion, respectively,  $\rho_h$  is the average density of the hypolimnion and  $u_*$  is the surface friction velocity.

## 2.4 Sensitivity analysis

Sensitivity of model output to nine parameters of mixing and heat exchange was evaluated for each lake. Three of the parameters influence surface heat and momentum exchange: bulk aerodynamic coefficient for sensible heat transfer ( $C_H$ ), bulk aerodynamic coefficient for latent heat transfer ( $C_E$ ) and coefficient of wind drag ( $C_D$ ). The remaining six parameters control surface and hypolimnetic mixing: mixing efficiency for convective

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783 overturn ( $C_c$ ), mixing efficiency of wind stirring ( $C_w$ ), mixing efficiency of shear  
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785 production ( $C_s$ ), mixing efficiency of unsteady turbulence ( $C_T$ ), mixing efficiency of Kelvin-  
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787 Helmholtz turbulent billows ( $C_{KH}$ ), and mixing efficiency of hypolimnetic turbulence  
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789 ( $C_{HYP}$ ), (Table 2). To gauge a response to parameter change, the one-at-a-time (OAT)  
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791 method (Bruce et al. 2008) was adopted for the first stage of the MLCP, where the model  
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793 was first run with the model default value to each parameter and then run again  
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795 increasing and decreasing parameter values by 20%.

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797 Sensitivity to changes in parameter values for each of the five lake variables used in the  
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799 model assessment described above (temperature of the full water column, epilimnion,  
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801 hypolimnion,  $thermD$  and  $S_T$ ) was analysed. Normalised sensitivity coefficients ( $S_{ij}$ ) to  
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803 assess the relative sensitivity of variable  $i$  to parameter  $j$  were calculated according to:

$$804 \quad S_{ij} = \frac{\Delta C_i / C_{is}}{\Delta \beta_j / \beta_{js}} \quad (2-8)$$

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808 where  $\Delta C_i$  is the change in output variable  $i$ , averaged over the simulation period, from  
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810 the standard or reference value  $C_{is}$  (Table 2) and  $\Delta \beta_{js}$  is the change in parameter  $j$  from the  
811  
812 reference value  $\beta_j$  (Fasham et al. 1990).

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815 Sensitivity coefficients were then compared relative to ten characteristics describing the  
816  
817 morphometry, climatic conditions and trophic state of the lakes. These properties were  
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819 the maximum depth, lake volume, ratio of area to maximum depth, ratio of length to  
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821 width, annual average inflow, residence time, mean air temperature, mean short wave  
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823 radiation, mean wind speed and extinction coefficient (Table A3).

## 824 825 **3 Results**

### 826 827 **3.1 Model Performance**

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829 Using the simulated results from running GLM with the standard set of parameters, five  
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831 model fit metrics ( $RMSE$ ,  $MEFF$ ,  $r$ ,  $PRE$  and  $NMAE$ ) were calculated for five data sets (full  
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833 profile, epilimnion, hypolimnion temperature,  $thermD$  and  $S_T$ ) for each lake. The full set of  
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835 results is provided in Appendix B (Table B1) with  $NMAE$  results given in Table 3. A

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843 comprehensive description of model performance for each lake can be found in the plots  
844 of modelled versus observed temperature data included in Appendix B.  
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848 An analysis of model performance in the prediction of temperature profiles (full profile)  
849 demonstrated a robust fit for GLM across the selected metrics, with an average *RMSE* of  
850 1.34°C, *MEFF* of 0.88, *r* of 0.96, *PRE* of -0.16% and *NMAE* of 0.11 (Table B1). The lakes with  
851 the lowest *RMSE* included Feagh, Tarawera and Emaiksoun. The highest *RMSE* values  
852 were calculated for Ravn, Ammersee and Woods. Ammersee also recorded the lowest  
853 values for *MEFF* along with NamCo and Toolik. All values of *r* were > 0.9, with the  
854 exception of Toolik. The *PRE* values ranged from +18% for NamCo to -15% for  
855 Rassnitzersee. Because lakes had both positive and negative *PRE* (representing a  
856 temperature bias, warm and cold respectively) the mean *PRE* was -0.16%. The lowest  
857 absolute *PRE* was for GrosseDhuenn (0.33%) which also performed well on all five  
858 measures of model fit.  
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867 In general, the model performance predicting the epilimnion temperatures was of similar  
868 magnitude to the full-profile temperatures (*RMSE* mean = 1.62°C). By analysing the *PRE*,  
869 it is clear that the GLM tended to produce both warm and cold temperature biases in the  
870 epilimnion, slightly favouring a cold bias (mean *PRE* = -0.84%). For most lakes, model  
871 performance metrics were similar for the epilimnion as the full profile with the exception  
872 of Windermere and Zurich which performed worse and Oneida which performed better  
873 in the computation of epilimnion temperatures.  
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879 For the hypolimnetic temperature simulations, average *RMSE* and *NMAE* values were  
880 relatively low, 1.31°C and 0.14 respectively. Typically small seasonal variation across all  
881 lakes led to greater percentage error between model and simulated data with both warm  
882 and cold temperature biases and a tendency to a warm bias (mean *PRE* = 1.97%). The  
883 mean *r* value of 0.73 was the lowest of the three temperature-associated properties. Lakes  
884 with the highest model performance for hypolimnion temperature included Geneva01,  
885 Geneva03 and Como with the lowest being Rassnitzersee, Esthwaite and Blelham. Model  
886 efficiency values for the calculation of hypolimnion temperatures were poor with less  
887 than a third greater than 0.5 and 44% of lakes recording a value of less than zero.  
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903 Thermocline depth (*thermD*) was a difficult parameter to model with the poorest *PRE* and  
904 *NMAE* values (Tables 3 & B1). Measures of model performance comparing calculations of  
905 observed and simulated *thermD* ranged in value across the lakes with *PRE* values from -  
906 16% to +52% and *NMAE* ranging from 0.10 to 0.76 (Tables 3 & B1). The *PRE* values  
907 indicate a bias towards over prediction of *thermD* by the model compared to the observed  
908 data. This was most apparent in Lake Geneva over the winter months when GLM  
909 predicted full mixing (i.e. *thermD* = lake depth) and the field data recorded a shallow  
910 *thermD* (<5m). As the lake depth was >300m this resulted in large relative error of greater  
911 than 6000%, leading to unfavourable mean measures of fit.  
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916 The *NMAE* values for calculation of  $S_T$  were generally low. The higher values of *NMAE*  
917 were associated with lakes such as Ammersee, Oneida and Pusiano which all had  
918 relatively low  $S_T$  during the simulated period. The mean *MEFF* and *r* were both quite high  
919 (0.83 and 0.96, respectively) indicating that the general seasonal patterns for  $S_T$   
920 prediction across the majority of lakes were well simulated by the model.  
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925 Analysis of the relationship between indices of model fit and input quality showed some  
926 correlation for the prediction of full profile, epilimnion and hypolimnion temperatures  
927 and *thermD* (Table B2). Analysis of measures of *PRE* indicated a cold bias in prediction of  
928 both full profile and hypolimnion temperatures when input uncertainty is greatest  
929 (Figure 4b). In addition, for lakes where the meteorological measurement station was  
930 near or at the lake edge, there was a warm bias and for lakes where meteorological input  
931 was sourced from further away, there was a cold bias (Figure 4a). Similarly, there was a  
932 warm bias for the prediction of hypolimnetic temperatures for lakes with high frequency  
933 meteorological data and a cold bias for lakes with daily meteorological data (Figure 4c).  
934 Lakes with lowest input uncertainty associated with the estimation of  $K_w$  corresponded  
935 with lowest values of *r* with respect to the prediction of full-profile temperatures (Figure  
936 4d) and similarly lakes that had close to ideal ranking of overall input uncertainty scored  
937 the lowest values of *r* for epilimnion temperatures (Figure 4e). This would be attributed  
938 to the use of  $K_w$  as a calibration parameter for lakes where there were no measurements  
939 for light attenuation. High frequency observed data also correlated with high *NMAE*  
940 scores for the prediction of hypolimnion temperatures (Figure 4f).  
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963 Analysis of model performance revealed a number of significant correlations linking  
964 model performance to lake characteristics (Table B3). For comparison of absolute model  
965 performance, the *RMSE* metric was used for temperatures and *MEFF* for *thermD* and *S<sub>T</sub>*.  
966 Whilst measurements of *PRE* can be a deceptive measure of model performance for lake  
967 variables where under and over-prediction occurs in equal measure, they are useful to  
968 observe patterns of bias in model prediction. A number of significant correlations  
969 between lake characteristics and model error are illustrated in Figure 5 and Figure 6 and  
970 described below.  
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978 The *RMSE* error associated with the prediction of both full profile and hypolimnion  
979 temperatures was generally higher for lakes with high light extinction ( $K_w > 0.8 \text{ m}^{-1}$ ) and  
980 lower for clear lakes ( $K_w < 0.3 \text{ m}^{-1}$ ) (Figure 5a&b). A correlation was observed between  
981 the *RMSE* associated with the prediction of hypolimnion temperatures and lake depth  
982 (Figure 5c), with deep lakes (>100 m) having the lowest values of *RMSE* (<1°C). In terms  
983 of relative measures of model performance, for lakes with both low inflows ( $< 10^5 \text{ m}^3\text{s}^{-1}$ )  
984 and low levels of incident short wave radiation averaged over the entire simulation period  
985 ( $< 120 \text{ Wm}^{-2}$ ) there was a cold bias in prediction of full profile and epilimnion  
986 temperatures, respectively (Figure 5c&d). Whilst correlation was relatively low, there  
987 was some indication that for lakes with low residence time there was a cold bias in the  
988 GLM-predicted hypolimnetic temperatures (Figure 5f).  
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997 For prediction of *S<sub>T</sub>*, the lake depth, residence time and extinction coefficient all had a  
998 significant impact on model performance (Figure 6a, b & c). Generally, clear deep lakes  
999 (>100 m), with residence times > 2 years recorded the lowest values of *NMAE*. A reverse  
1000 pattern of correlation was observed for the prediction of *thermD*, with deep lakes having  
1001 the highest values of *NMAE* and shallow lakes (<40m) showing highest levels of *thermD*  
1002 predictive accuracy (Figure 6d). There was a small but significant trend where GLM over  
1003 estimated *S<sub>T</sub>* in lakes with high incident short wave radiation ( $> 200 \text{ Wm}^{-2}$ ) (Figure 6e).  
1004 For prediction of *thermD*, GLM tended towards over-prediction which was more  
1005 pronounced in colder lakes (air temperature < 10°C) (Figure 6f).  
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1014 Model performance for the prediction of *thermD* and *S<sub>T</sub>* was better for lakes when mean  
1015  $L_N > 10$ , while these lakes tended to record reduced measures of model fit for the  
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1021 prediction of epilimnion and hypolimnion temperatures (Figure 7a,c,e,g). Conversely, for  
1022 the small number of lakes with a significant proportion of the stratification period under  
1023 a regime of  $L_N < 1$ , prediction of epilimnion and hypolimnion temperatures improved but  
1024  $thermD$  and  $S_T$  decreased (Figure 7b,d,f,h).

### 1030 3.2 Sensitivity Analysis

1031 The sensitivity analysis (SA) on each of the nine surface exchange and mixing parameters  
1032 highlighted differences both between lakes and thermal properties (Figure 8a-e). For all  
1033 three temperature metrics (full profile, epilimnion or hypolimnion) there was little  
1034 sensitivity to perturbations in physical parameters, when the SA was averaged over the 2  
1035 year simulation period. There was some degree of sensitivity to changes in  $C_d$  in the  
1036 calculation of hypolimnion temperatures and to  $C_e$  in the calculation of epilimnion  
1037 temperatures. Sensitivity index ( $SI$ ) for prediction of both  $thermD$  and  $S_T$ , were significant  
1038 ( $>1$ ) across a broader range of lakes (Figure 8d-e). While there was some variability  
1039 across the lakes and parameters, model output for both  $thermD$  and  $S_T$  had greatest  
1040 sensitivity to perturbations of  $C_d$ . Additionally, for  $S_T$  there was a consistent level of  
1041 sensitivity to perturbations of  $C_e$ .

1050 The sensitivity of each parameter was compared to a gradient of physical and climate lake  
1051 properties (Table C1-5) and a number of significant correlations were observed. For each  
1052 thermal metric, the three most significant correlations to lake characteristics were  
1053 compared (Figure 9). A common significant ( $p < 0.05$ ) trend was recorded for maximum  
1054 lake depth (Figure 9e, Figure 9m). For the prediction of full profile and epilimnion  
1055 temperatures, deeper and larger lakes were more sensitive to changes in  $C_{KH}$  than small,  
1056 shallow lakes (Figure 9e). Similarly, for the prediction of  $thermD$ , deeper lakes were more  
1057 sensitive to changes in  $C_c$ ,  $C_w$  and  $C_{KH}$  than shallow lakes (Figure 9).

1065 A significant correlation with air temperature indicated that lakes with low air  
1066 temperatures were more sensitive to changes in  $C_h$ ,  $C_s$  and  $C_{KH}$  than lakes in warm climates  
1067 (mean air temperature  $> 10^\circ\text{C}$ ) for the prediction of full-profile temperature (Figure 9c),  
1068 epilimnion (Figure 9d) and hypolimnion temperatures (Figure 9g). Lakes with low  
1069 inflows were more sensitive to changes in  $C_h$  for the prediction of hypolimnion  
1070 temperatures than those with larger inflows (Figure 9i). Finally, lakes with highest wind  
1071 speed recorded greatest  $SI$  to  $C_e$  in the prediction of  $S_T$  (Figure 9m).

## 4 Discussion

Historically, lake modellers have adopted simple methods to justify model performance and suitability, rarely reporting statistical measures of model fit (Arhonditsis and Brett 2004; Arhonditsis et al. 2006). For individual lake applications, these have been adequate to undertake scenario simulations and further our understanding of site specific dynamics. However, a common approach to model assessment, both in terms of metrics that should be applied and identification of a commonly agreed level of model performance, is necessary to further enhance model development (Bennett et al. 2013). Undertaking a standardized method of assessment of the community lake model, GLM, over a diversity of lakes has led to an improved level of understanding of the strengths and weaknesses in the predictive capacity of simple 1-D lake models. By first ascertaining an acceptable model error, we were able to elucidate the relation between model performance and data input uncertainty or lake characteristics (Figure 4; Figure 5).

The quality of input data was not as significantly related to model performance as expected. Lakes modelled using daily meteorological input, rather than hourly, did have the largest values of *NMAE* in the prediction of full profile temperature and *thermD* (Figure 4), which is not surprising given the importance of diurnal forcing in 1-D model predictive capability. The greater the meteorological observation distance to the lake tended to result in both cold-biased temperatures and under prediction of *S<sub>T</sub>* (Figure 4). The cause of warm-biased temperatures and over-prediction of lake stability when meteorological observations were obtained near or on-lake requires further investigation (Figure 4). The strong correlation between accuracy of *K<sub>w</sub>* measurements and model performance in the prediction of both full profile temperature and *thermD* (Figure 4) emphasises both the importance of light extinction in the determination of thermocline depth and the need to include measurements of *K<sub>w</sub>* in routine lake monitoring. The GLM can be coupled to water quality models such as the Aquatic EcoDynamics Model (AED; Hipsey et al. 2013) such that seasonal changes in *K<sub>w</sub>* would feedback in the model to potentially improve model prediction particularly in relation to *thermD*; this link is expected to further improve model accuracy in most circumstances.

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1143 The 1-D nature of the model implicitly assumes that the mixing within the lake can be  
1144 constrained by processes acting in the vertical and that processes which vary in the  
1145 horizontal, such as the degree of upwelling of the thermocline, have minimal impact on  
1146 vertical transport. This assumption is quantified by computation of the Lake Number  
1147 ( $L_N$ ) (Imberger and Patterson, 1989; eq. 2.7). As the  $L_N$  is a relative measure of the strength of  
1148 stratification to surface wind energy, the 1-D model assumption is said to hold true for  $L_N$   
1149  $\gg 1$  (Imberger and Patterson 1989; Yeates and Imberger 2003). Over the past three  
1150 decades, the 1-D model approach has been applied to a wide diversity of sites due to its  
1151 simplicity and tractability relative to 3-D models. However, given that  $L_N$  can be highly  
1152 variable, it has remained unclear what significance the 1-D assumption has on model  
1153 prediction error for various lake attributes and under what conditions this assumption  
1154 would no longer hold. The strong correlation ( $r^2=[0.70,0.82]$ ) between the percent of time  
1155  $L_N < 1$  during the stratified period and the model performance of both *thermD* and  $S_T$   
1156 endorses the use of  $L_N$  as an indicator of the validity of the 1-D model assumption, and  
1157 should be considered when modellers are deciding on model suitability.

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1169 A comparison of *PRE* against  $L_N$  for the calculation of simulated versus observed  $S_T$   
1170 indicated that lakes with mean  $L_N < 1$  tended to underestimate  $S_T$ . For these lakes, the 1-  
1171 D assumption as defined by  $L_N$  does not hold. One would expect mixing to be  
1172 underestimated and  $S_T$  to be higher, unless the resulting warmer near surface  
1173 temperatures led to greater heat losses by evaporation. Yeates and Imberger (2003)  
1174 demonstrated that for lakes where deep mixing is important, a 1-D lake model mixing  
1175 scheme similar to that used in GLM tended to overmix the water column and thus  
1176 underestimate lake stability and therefore  $S_T$ . A solution put forward by Yeates and  
1177 Imberger (2003) included a pseudo two-dimensional algorithm in the 1-D model DYRESM  
1178 to parameterise internal and boundary fluxes. Similarly Gaudard et al. (2016) proposed a  
1179 method of adding a seasonal component in the parameterisation of internal seiches that  
1180 led to improved accuracy in the prediction of deep mixing in the 1-D model SIMSTRAT.  
1181 Whilst compromising computational efficiency, lake modellers could consider a similar  
1182 approach when conditions for improved deep mixing accuracy are necessary. For  
1183 example, this approach could be valuable where upwelling or internal nutrient loading is  
1184 deemed important or when specific distribution phenomena such as deep chlorophyll  
1185 maxima are the focus of the modelling study.

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1205 Further exploration of how individual lake properties relate to measures of model  
1206 performance indicated the strongest correlations against  $K_w$  and lake depth (Figure 5;  
1207 Figure 6). Lakes with high  $K_w$  ( $> 0.5$ ), recorded greatest error in the prediction of lake  
1208 temperatures particularly in the hypolimnion. While there was no significant correlation  
1209 between the accuracy in prediction of epilimnion temperatures and lake depth, there was  
1210 a strong positive correlation for measures of model performance in prediction of  
1211 hypolimnion temperatures and depth (Figure 5). That is, for deeper lakes ( $>40$  m) where  
1212 surface mixing dynamics have less influence on hypolimnion temperatures, GLM predicts  
1213 hypolimnion temperatures with greater accuracy. This suggests that while the surface  
1214 thermodynamics are better represented by the model, prediction of rates of mixing across  
1215 the metalimnion requires attention and further development to enable more confident  
1216 prediction across the diversity of lake types. Relatively shallow, well-mixed lakes, such as  
1217 Feeagh and Emaiksuon, had the highest overall model performance. These lakes are  
1218 dominated by surface exchange with no thermocline and associated deepening.

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1229 The prediction of the lake thermocline depth proved harder to achieve than the lake  
1230 temperatures. Particularly in moderately deep lakes, small relative deviations in  
1231 predictions can result in large changes to error magnitude. As the GLM-predicted *thermD*  
1232 was both deeper and shallower than the observed *thermD* in different lakes, there does  
1233 not appear to be a consistent bias in the mixing algorithms, and rather, it may be driven  
1234 by high sensitivity to input parameter uncertainty and require site specific calibration.  
1235 The positive correlation between *NMAE* of thermocline prediction and lake depth was  
1236 significant with best fit occurring for lakes less than 50-80 m deep (Figure 6). A tendency  
1237 to over-predict thermocline depth in the majority of lakes could be attributed to an over-  
1238 prediction of penetrative heat and may be related to both the application of a standard  
1239 minimum layer thickness for all lakes and the use of a single average  $K_w$  value over 2  
1240 annual seasonal cycles. The positive correlation with  $K_w$  indicates that a single  $K_w$  for all  
1241 seasonal conditions is not appropriate, particularly for lakes with high mean or seasonally  
1242 variable  $K_w$  values. A consideration for using a  $K_w$  weighted towards the summer stratified  
1243 period could be a solution or coupling to a water quality model with explicit light  
1244 extinction feedback properties could improve thermocline prediction particularly in lakes  
1245 with high light extinction ( $K_w > 0.5$ ) (Shatwell et al. 2016).

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1265 The absence of strong sensitivity to parameterisation of surface exchange and mixing  
1266 algorithms in the prediction of temperature profiles (Figure 8) is indicative of the  
1267 dominance of surface boundary conditions in the thermal budget of individual lakes and  
1268 negative feedbacks in the surface heating sub-model. In contrast, the prediction of  
1269 thermocline depth and Schmidt Stability were more sensitive to changes in  
1270 parameterisation. In particular, the model was sensitive to the shear mixing efficiency and  
1271 wind drag coefficient parameters. Both parameters are directly related to the transfer of  
1272 wind energy to mixing. The errors in computing these terms again points to the need for  
1273 more effort in parameterizing the processes operative when  $L_N$  is low and shear increases  
1274 across the thermocline. Additionally, wind increases in magnitude as it flows across a lake.  
1275 This effect is important for small and large lakes and is not included when wind is  
1276 modelled with bulk drag coefficients. Care should be taken in both the accuracy of wind  
1277 speed measurements as well as the parameterization and classification of these  
1278 parameters in relation to lake characteristics to improve model performance across a  
1279 wide variety of lake properties.  
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1291 In general, simulations of deep lakes with large volumes and residence times were most  
1292 sensitive to changes in mixing efficiency parameters (as measured by changes in *thermD*  
1293 and *S<sub>T</sub>*) (Figure 9), which was expected since larger lakes require greater efficiency in  
1294 transfer of surface momentum input to thermocline deepening and subsequent mixing.  
1295 Lakes with low  $K_w$  were most sensitive to changes in surface exchange parameters. This  
1296 sensitivity is logical given that in lakes with low  $K_w$ , light will penetrate deeper causing a  
1297 deeper thermocline. Processes which moderate depth of mixing in the epilimnion, such as  
1298 convection, become important. Being able to model changing dynamics of lakes as  $K_w$   
1299 changes with modified hydrology and altered loading of chromophoric dissolved organic  
1300 matter is critical for quantifying the changes associated with climate variability (Snucins  
1301 and Gunn 2000).  
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1310 An appealing alternative to the minimal calibration presented here (i.e., input data  
1311 refinement, wind factor and river inflow slope adjustment) will be the relaxation of the  
1312 assumption of globally common parameter values for the core hydrodynamic parameters  
1313 and the adoption of a Bayesian hierarchical calibration framework that reflects the more  
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1323 realistic notion that each lake (or group of lakes) is peculiar but shares some commonality  
1324 of behavior with other lakes (Zhang and Arhonditsis 2009; Cheng et al. 2010; Shimoda  
1325 and Arhonditsis 2015). The proposed approach represents a pragmatic compromise  
1326 between system- or group-specific and globally common parameter estimates and may  
1327 be a conceptually sound strategy to accommodate within- and among-lake variability in  
1328 the context of model application within the global observatory network (Figure 10).  
1329 Recent work has shown that the delineation of more homogeneous subsets of lakes with  
1330 respect to their morphological characteristics/hydraulic regimes and their subsequent  
1331 integration with hierarchical frameworks may give models with better predictive  
1332 capacity (Cheng et al. 2010; Shimoda and Arhonditsis 2015). In particular, sensitivity  
1333 analysis patterns identified in this study could be used to identify groups with similarities  
1334 in behavior (e.g., deep versus shallow lakes, high versus low water transparency) as well  
1335 as to identify the candidate parameters for the calibration exercise. The prior  
1336 distributions of the hyper-parameters (or global priors) can be easily formulated on the  
1337 basis of existing knowledge (e.g., field observations, laboratory studies, and information  
1338 from the modeling literature) of the relative plausibility of their values. Moreover, the  
1339 proposed incorporation of mathematical models into Bayesian hierarchical frameworks  
1340 can also assist the effective modeling of systems with limited knowledge by enabling the  
1341 transfer of information across systems. With the hierarchical model configuration, we can  
1342 potentially overcome problems of insufficient local data by “borrowing strength” from  
1343 well-studied lakes on the basis of distributions that connect systems in space (Zhang and  
1344 Arhonditsis 2009). Another advantage of a Bayesian calibration configuration will be the  
1345 ability to express the input uncertainty in the form of probability density functions which  
1346 can then be propagated through the model structure and may ultimately shape the  
1347 moments of the posterior predictive distributions.  
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1365 Through international collaboration, this work allowed us to test and to improve the  
1366 process and performance of a 1-D open source model by simulating thermal structure in  
1367 lakes with varying physical and climatic characteristics. Initial efforts in setting up a  
1368 collaborative network of lake modellers were rewarded with improved user support and  
1369 feedback, refinements and testing to the development team. From its initiation as v1.0 in  
1370 the MLCP, using feedback and re-coding by network members, the GLM evolved through  
1371 numerous improvements to the current v2.2 described in this study. The study also  
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1383 identified the most sensitive parameters related to surface exchange and mixing that  
1384 affect model prediction and therefore performance for each individual lake. These  
1385 sensitivities could then be correlated to lake characteristics such as residence time,  
1386 meteorological conditions and trophic status. Additionally, this work opens a new  
1387 challenge for the community of limnologists involved in ecosystem modelling. Indeed the  
1388 next step would be cross lake comparison projects including biogeochemical processes  
1389 simulation using a similar open source community biogeochemical model such as the  
1390 Framework for Aquatic Biogeochemical Models (FABM: Bruggeman and Bolding, 2014)  
1391 and/or AED (Hipsey et al. 2013). The establishment of well-defined standards for  
1392 modelling techniques (set up, output analysis), and a diversity of lakes and scientists  
1393 provides enormous opportunity for further advances by aquatic ecosystem modellers.  
1394 The significance of the MLCP resides in a common and collaborative approach to  
1395 answering globally relevant lake science questions, and providing a benchmark for model  
1396 performance and an associated parameter set that future applications can refer to.  
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1413 networks as well as discussions and working groups held during AEMON workshops and  
1414 GLEON meetings.  
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Table 1 - Lakes included in the Multi-Lake Comparison Project Stage 1, abbreviation, maximum depth, surface area at maximum depth, crest elevation latitude (°N) and longitude (°E).

Lake Name	Abv.	Maximum Depth (m)	Surface Area at Crest (m <sup>2</sup> )	Crest Elevation (m)	Latitude	Longitude	Reference
Lake Alexandrina	AL	9.4	655,755,315	3.4	-35.4	139.1	(Hipsey et al. 2014b)
Ammersee	AM	83.7	47,250,000	533.5	48.0	11.1	(Weinberger and Vetter 2014; Bueche et al. 2017)
Blelham	BL	14.5	104,000	14.0	54.4	-3.0	(Woolway et al. 2015)
Lake Bourget	BO	146.0	42,575,000	230.5	45.4	5.9	(Vinçon-Leite et al. 1989, 2014; Kerimoglu et al. 2016)
Cannonsville Reservoir	CA	52.0	19,000,000	351.0	42.1	-75.3	(Samal et al. 2012)
Lake Como	CO	440.0	147,012,649	410.0	46.0	9.3	(Laborde et al. 2010; Copetti et al. 2013; Guyennon et al. 2014)
Lake Constance	CN	253.3	472,650,000	395.0	47.6	9.4	(Wessels 1998; Frassl et al. 2014)
El Gergal	EG	55.0	4,732,669	50.0	37.0	-2.5	(Rigosi et al. 2011)
Emaiksoun	EM	2.4	1,860,000	2.4	71.2	-156.8	(Potter 2011)
Esthwaite	ES	15.5	1,000,000	15.5	54.4	-3.0	(Woolway et al. 2015)
Feeagh	FE	43.0	3,942,266	9.0	53.4	-9.6	(Dalton et al. 2014)
Lake Geneva 2001-2	G1	309.0	578,560,865	371.4	46.4	6.1	(Anneville et al. 2010)
Lake Geneva 2003-4	G3	309.0	578,560,865	371.4	46.4	6.1	(Anneville et al. 2015)
Grosse Dhuenn	GD	48.5	3,750,100	177.5	51.1	7.2	(Weber et al. 2017)
Harp Lake	HA	37.5	713,800	327.0	45.4	-79.1	(Yao et al. 2014)
Lake Iseo	IS	256.0	60,880,350	185.2	45.7	10.1	(Pilotti et al. 2013, 2014; Valerio et al. 2015)
Lake Kinneret 2003-4	K3	44.0	173,000,000	-208.9	32.0	35.6	(Gal et al. 2009)
Lake Kinneret 1997-8	K7	44.0	173,000,000	-208.9	32.0	35.6	(Bruce et al. 2006)
Lake Mendota	ME	25.0	39,581,170	259.0	43.0	-89.4	(Magnuson et al. 2006)
Mount Bold Reservoir	MB	45.4	3,080,000	246.9	-35.1	138.7	(van der Linden and Burch 2016) Rigosi et al. 2015
Muggelsee	MG	8.0	7,318,000	32.4	52.0	13.6	(Huber et al. 2008)
Lake Nam Co	NM	98.9	2,018,230,000	4718.0	30.7	90.6	(Wang et al. 2009)
Oneida	ON	17.0	207,100,000	112.0	43.0	-75.9	(Hetherington et al. 2015)
Lake Pusiano	PU	30.9	8,123,699	27.0	45.8	9.3	(Copetti et al. 2006, 2013; Carraro et al. 2012)

2201	Rappbode	RP	85.6	4,344,724	423.6	51.7	10.9	(Bocaniov et al. 2014)
2202	Rassnitzersee	RS	40.0	3,033,057	85.0	51.3	12.0	(Böhrer et al. 1998; Boehrer et al. 2014)
2203								
2204	Ravn	RV	33.0	1,820,000	33.0	56.0	4.8	(Trolle et al. 2008a; b)
2205								
2206	Rotorua	RO	22.0	79,722,140	280.0	-38.0	176.3	(Burger et al. 2008)
2207	Stechlin	ST	69.5	4,231,549	60.0	53.2	13.0	(Kirillin et al. 2013)
2208								
2209	Tarawera	TA	88.0	40,996,000	297.8	-38.2	176.4	(Hamilton et al. 2006, 2010)
2210	Toolik	TO	24.0	940,119	740.0	68.6	-149.6	(MacIntyre et al. 2009)
2211	Windermere	WI	66.8	14,779,600	66.8	54.4	-3.0	(Woolway et al. 2015)
2212								
2213	Woods Lake	WO	10.4	15,000,000	738.2	-42.0	147.0	(Hydro Tasmania 2003)
2214	Lower Lake Zurich	ZU	136.0	66,600,000	406.0	47.3	8.8	(Peeters et al. 2002; Schmid and Köster 2016)

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Table 2 - Description, symbols and initial values of the parameters used in the sensitivity analysis.

Symbol	Description	Reference	Initial value
<b>Surface Heat Exchange</b>			
$C_h$	Bulk aerodynamic coefficient for sensible heat transfer	(Fischer et al. 1979)	0.0013
$C_e$	Bulk aerodynamic coefficient for latent heat transfer	(Fischer et al. 1979)	0.0013
$C_d$	Bulk aerodynamic momentum transfer coefficient	(Fischer et al. 1979)	0.0013
<b>Mixing</b>			
$C_c$	Mixing efficiency - convective overturn	(Yeates and Imberger 2003)	0.2
$C_w$	Mixing efficiency - wind stirring	(Spigel et al. 1986)	0.23
$C_t$	Mixing efficiency - unsteady turbulence (acceleration)	(Sherman et al. 1978)	0.3
$C_s$	Mixing efficiency - shear production	(Sherman et al. 1978)	0.51
$C_{KH}$	Mixing efficiency - Kelvin-Helmholtz turbulent billows	(Sherman et al. 1978)	0.3
$C_{hyp}$	Mixing efficiency of hypolimnetic turbulence	(Weinstock 1981)	0.5



Table 3 - *NMAE* for base simulations using standard parameter set against full profile temperature (Full Prof. Temp.) [°C], epilimnion temperature (Epi. Temp.) [°C], Hypolimnion temperature (Hyp. Temp.) [°C], thermocline depth (*thermD*) [m] and Schmidt Stability (*S<sub>t</sub>*). Note that for fully mixed lakes or for lakes where temperature profiles were shallower than the thermocline depth, *NMAE* values are listed as not applicable (N/A). N refers to the number of profiles used in the calculation of model performance.

Lake	Full Prof. Temp. (°C)	Epi. Temp. (°C)	Hyp. Temp. (°C)	thermD (m)	ST	N
Alexandrina	0.07	0.07	N/A	N/A	N/A	
Ammersee	0.19	0.20	0.13	0.40	0.17	
Blelham	0.12	0.13	0.31	0.18	0.45	
Bourget	0.08	0.11	0.07	0.32	0.09	
Cannonsville	0.10	0.05	0.15	0.39	0.12	
Como	0.10	0.17	0.06	0.64	0.19	
Constance	0.08	0.09	0.07	0.11	0.16	
ElGergal	0.08	0.06	0.07	0.30	0.27	
Emaiksoun	0.08	0.08	N/A	N/A	N/A	
Esthwaite	0.13	0.11	0.35	0.15	0.24	
Feeagh	0.06	0.04	0.09	0.14	0.30	
Geneva01	0.09	0.11	0.04	0.41	0.22	
Geneva03	0.08	0.05	0.04	0.52	0.20	
GrosseDhunn	0.07	0.05	0.09	0.37	0.09	
Harp	0.18	0.12	0.27	0.68	0.19	
Iseo	0.08	0.10	0.07	0.76	0.16	
Kinneret03	0.07	0.07	0.07	0.28	0.20	
Kinneret97	0.05	0.06	0.05	0.15	0.21	
Mendota	0.11	0.10	0.11	0.30	0.23	
MtBold	0.08	0.08	0.06	0.25	0.43	
Muggelsee	0.07	0.06	N/A	N/A	N/A	
NamCo	0.23	0.17	0.22	0.28	0.35	
Oneida	0.04	0.03	0.06	0.19	0.86	
Pusiano	0.14	0.11	0.26	0.24	0.19	
Rappbode	0.14	0.08	0.12	0.23	0.16	
Rassnitzersee	0.17	0.15	0.23	0.15	0.17	
Ravn	0.19	0.14	0.21	0.27	0.34	
Rotorua	0.07	0.08	0.08	0.09	0.43	
Stechlin	0.13	0.11	0.11	0.33	0.14	
Tarawera	0.04	0.04	0.03	0.27	0.10	
Toolik	0.25	0.26	0.25	0.61	0.43	
Windermere	0.14	0.23	0.26	0.22	0.21	
Woods	0.17	0.17	N/A	N/A	N/A	
Zurich	0.12	0.09	0.16	0.42	0.17	
<b>Mean</b>	<b>0.11</b>	<b>0.10</b>	<b>0.14</b>	<b>0.32</b>	<b>0.25</b>	

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<b>Median</b>	<b>0.09</b>	<b>0.09</b>	<b>0.10</b>	<b>0.28</b>	<b>0.20</b>	
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2470	observatory network. Future applications can improve parameter accuracy
2471	within a Bayesian hierarchical framework based on suitable groupings of lakes
2472	into distinct archetypes. Other lakes with limited data for robust calibration, can
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adopt standard model parameters depending on the lake archetype, to which it  
best relates. .... 53



Figure 1 - Lake outlines to scale for all lakes in the current MLCP GLM assessment.

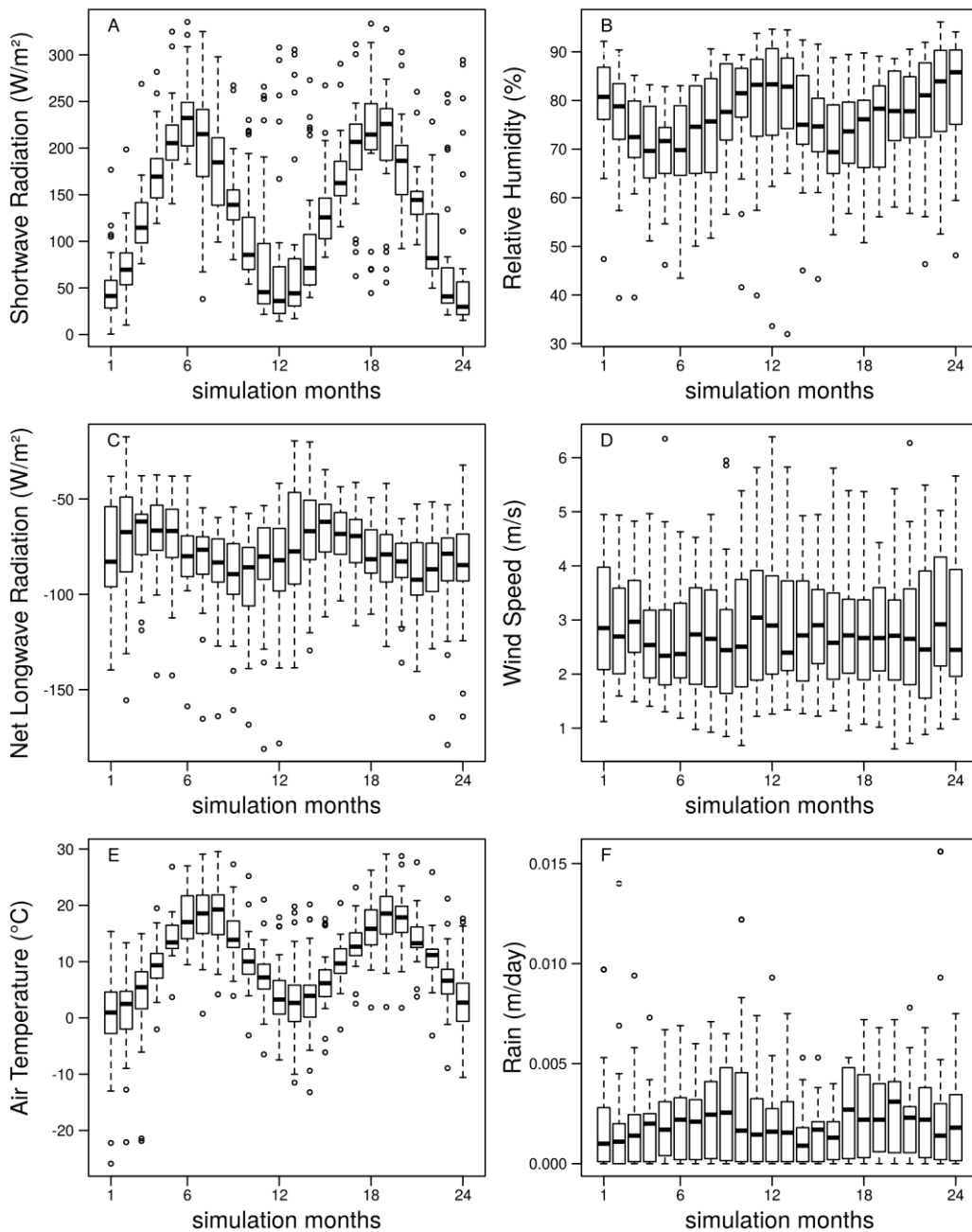


Figure 2 - Time series of monthly mean values across all lakes for (a) short wave radiation, (b) relative humidity, (c) net longwave radiation, (d) wind speed, (e) air temperature and (f) precipitation. For each box, horizontal lines represent median, 25<sup>th</sup> and 75<sup>th</sup> percentile, whiskers < 1.5 times the interquartile, and outliers ( $\circ$ ) values > 1.5 times the interquartile range. Note that lakes from the Southern Hemisphere start with a shift of 6 months relative to the Northern Hemisphere lakes.

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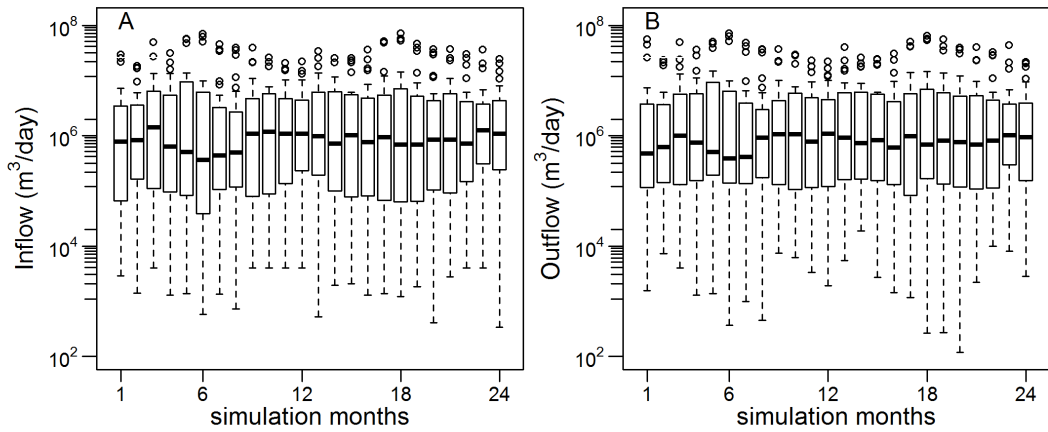


Figure 3 – Time series of monthly mean values across all lakes for (a) inflows and (b) outflows. For each box, horizontal lines represent median, 25<sup>th</sup> and 75<sup>th</sup> percentile, whiskers < 1.5 times the interquartile, and outliers (o) values > 1.5 times the interquartile range.

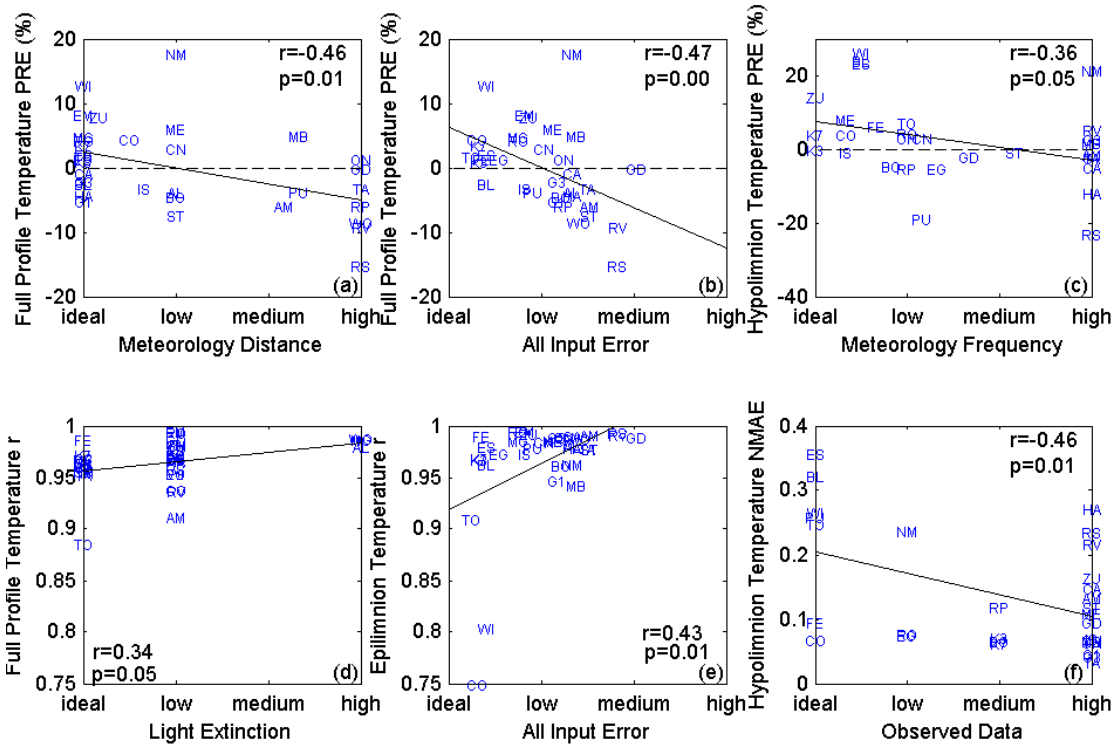


Figure 4 – Correlation between GLM model performance metrics PRE (a-c),  $r$  (d-e) and NMAE (f) for prediction of full profile temperatures (a, b & d), epilimnion temperatures (e) and hypolimnion temperatures (c & f) against rankings of input data uncertainty where 0-ideal, 1-low, 2-medium and 3-high level of uncertainty. Refer to Table 1 for lake acronyms and Table A1 for details of input uncertainty ranking system.



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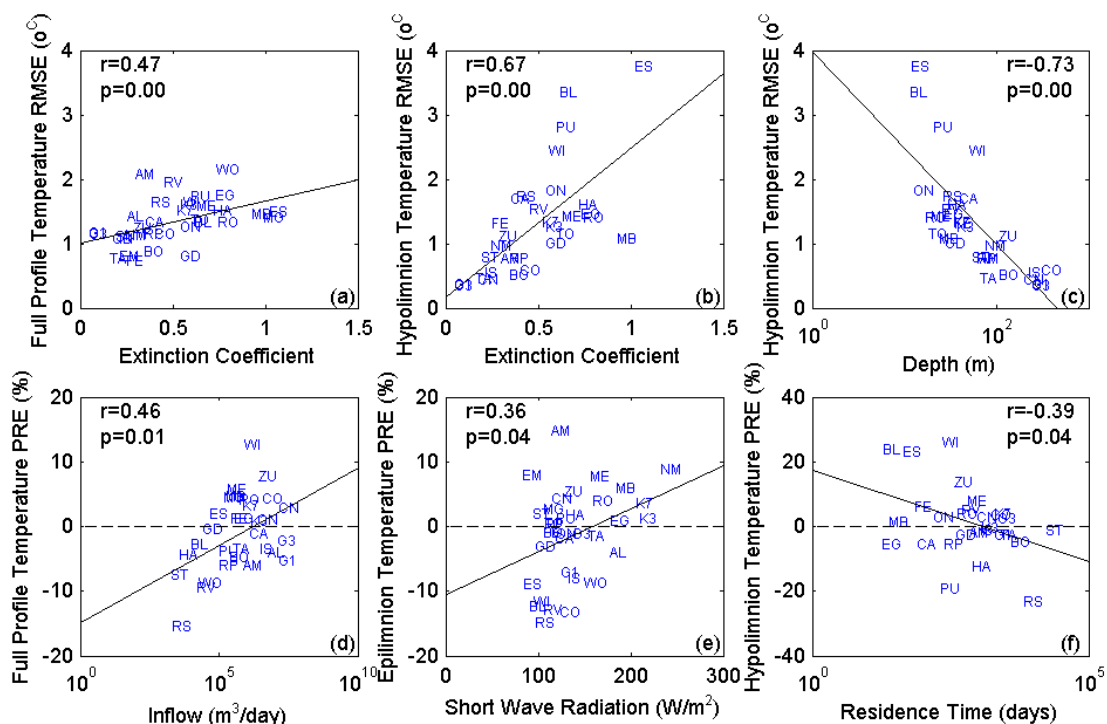


Figure 5 - GLM model performance metrics for prediction of full profile temperature (a&d), epilimnion temperature (e) and hypolimnion temperature (b,c&f) against lake characteristics. Refer to Table 1 for lake acronyms.

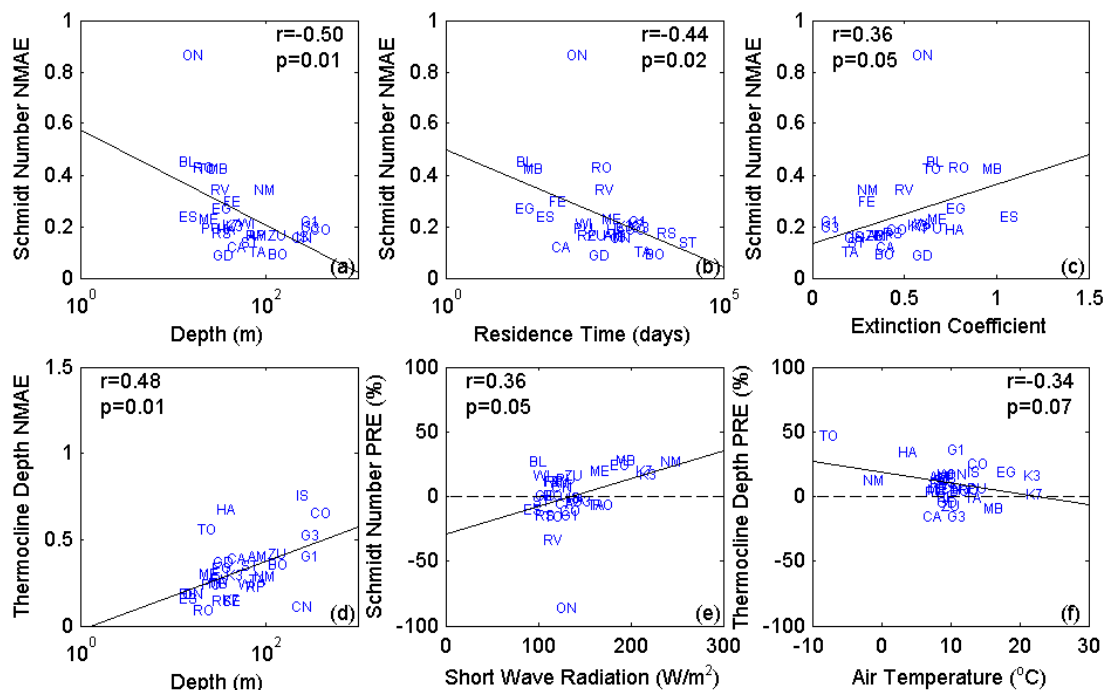


Figure 6 - GLM model performance metrics for prediction of thermocline depth (d,f) and Schmidt stability (a,b,c&e) against lake characteristics. Refer to Table 1 for lake acronyms.

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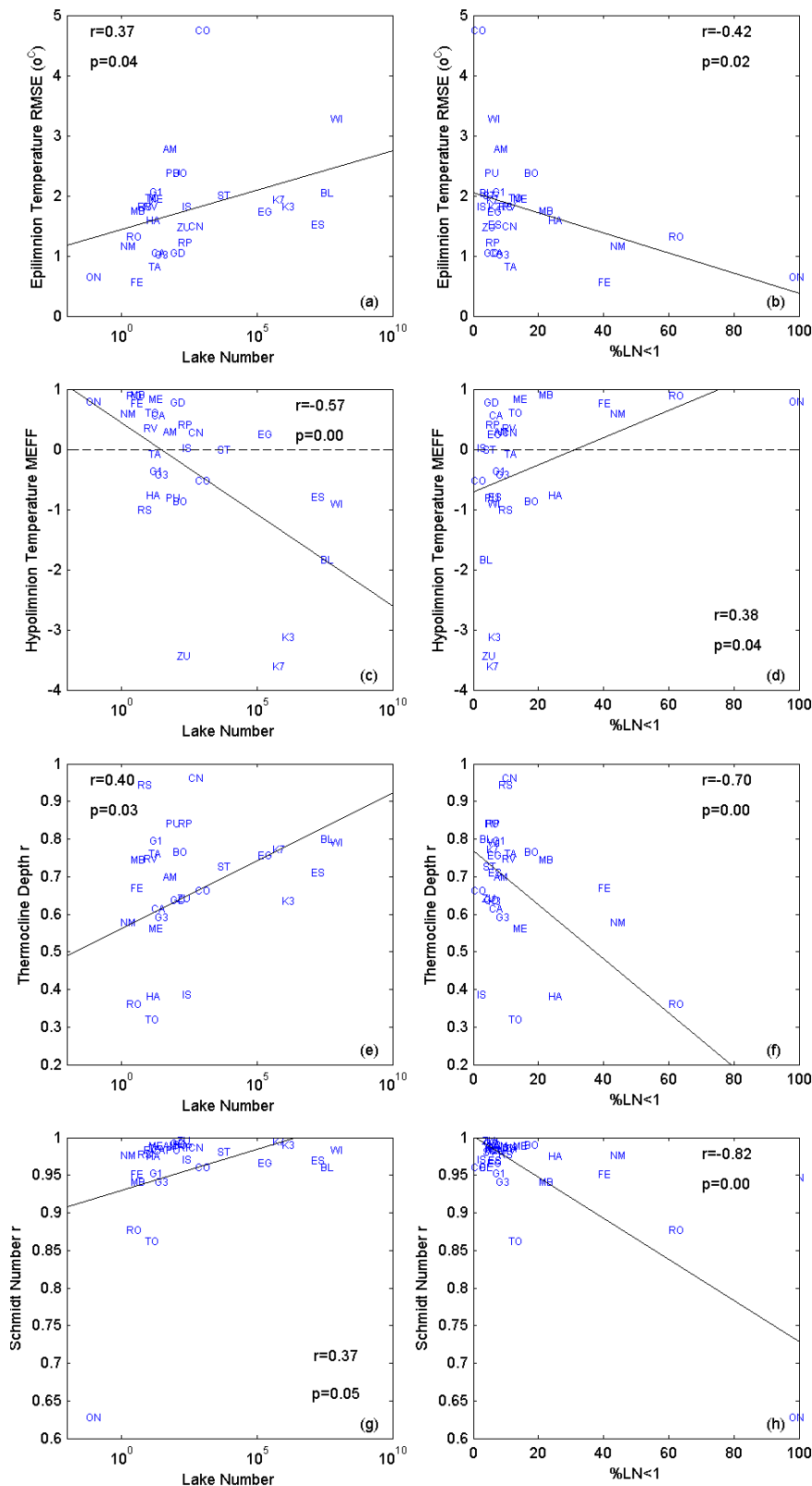


Figure 7 - GLM model performance metrics for prediction of epilimnion temperature (a,b), hypolimnion temperature (c,d), thermocline depth (e,f) and Schmidt stability (gh) against Lake Number and %LN<1. Refer to Table 1 for lake acronyms

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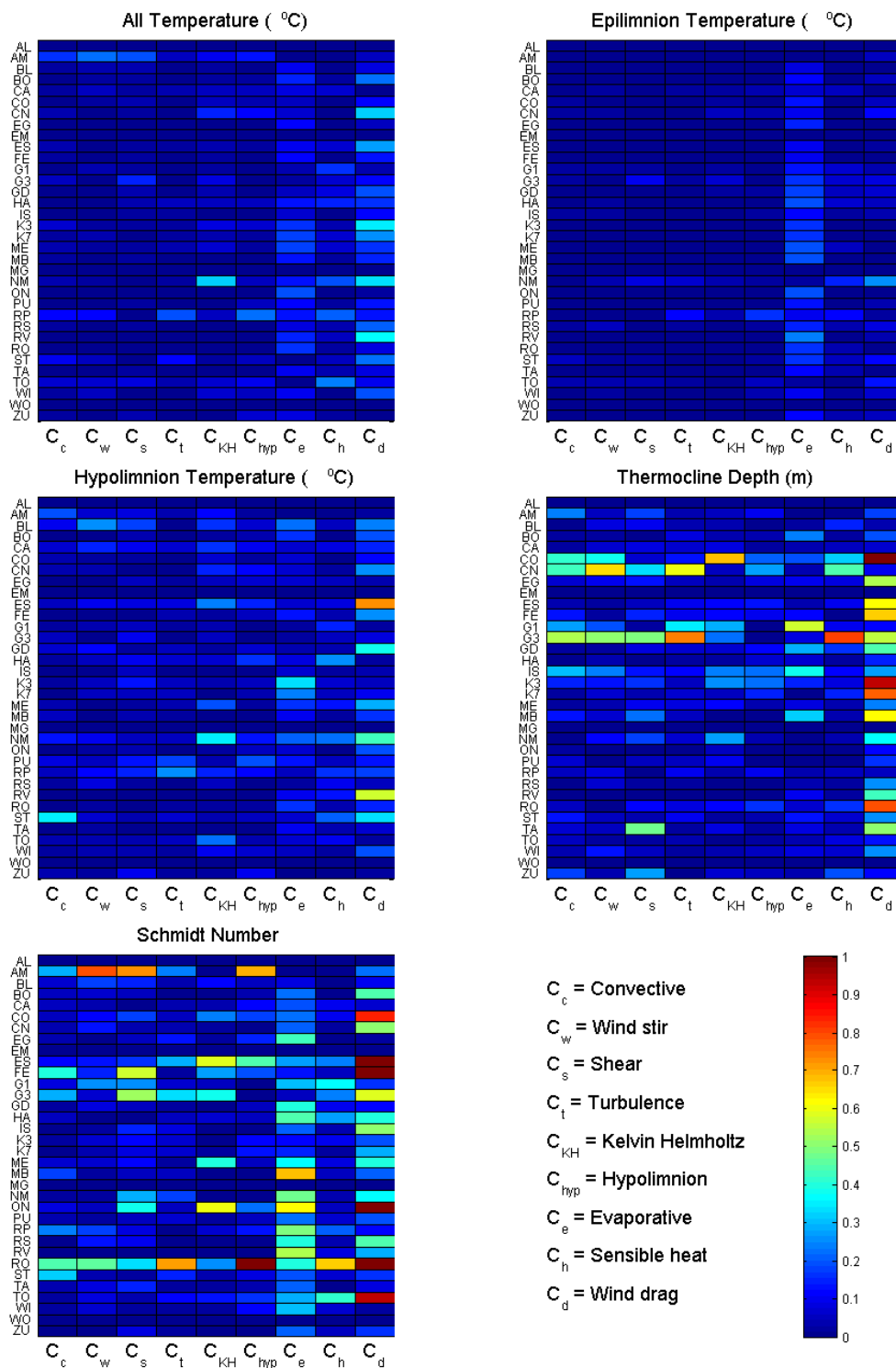
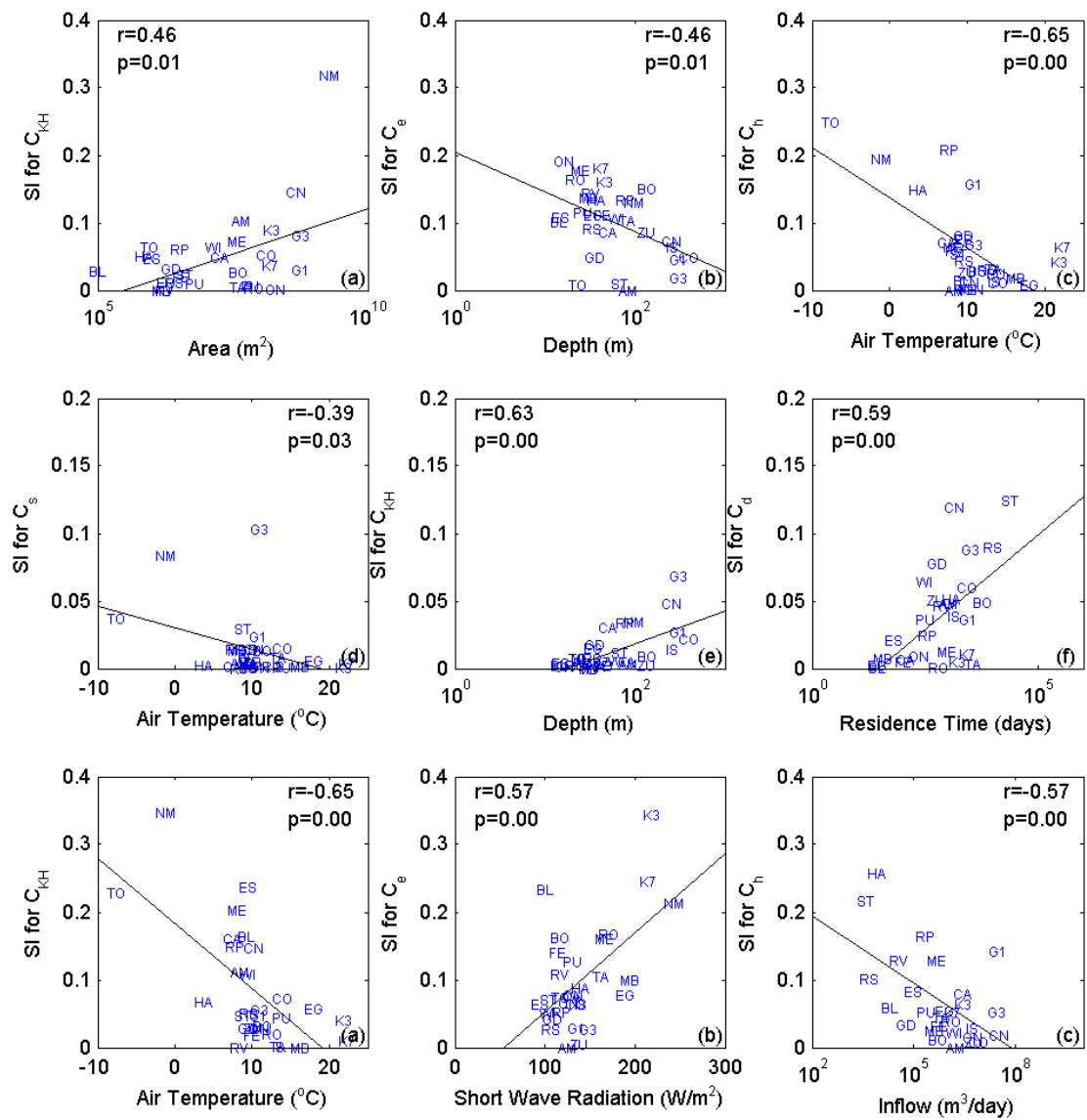


Figure 8 - Sensitivity indices for a) full profile temperature, b) epilimnion temperature, c) hypolimnion temperature, d) thermocline depth and e) Schmidt stability. The colour bar has been limited to a value of 1 so that any sensitivity index (SI) greater than one (indicating the percent response in thermodynamic metric is greater than the change in physical parameter) has been highlighted.

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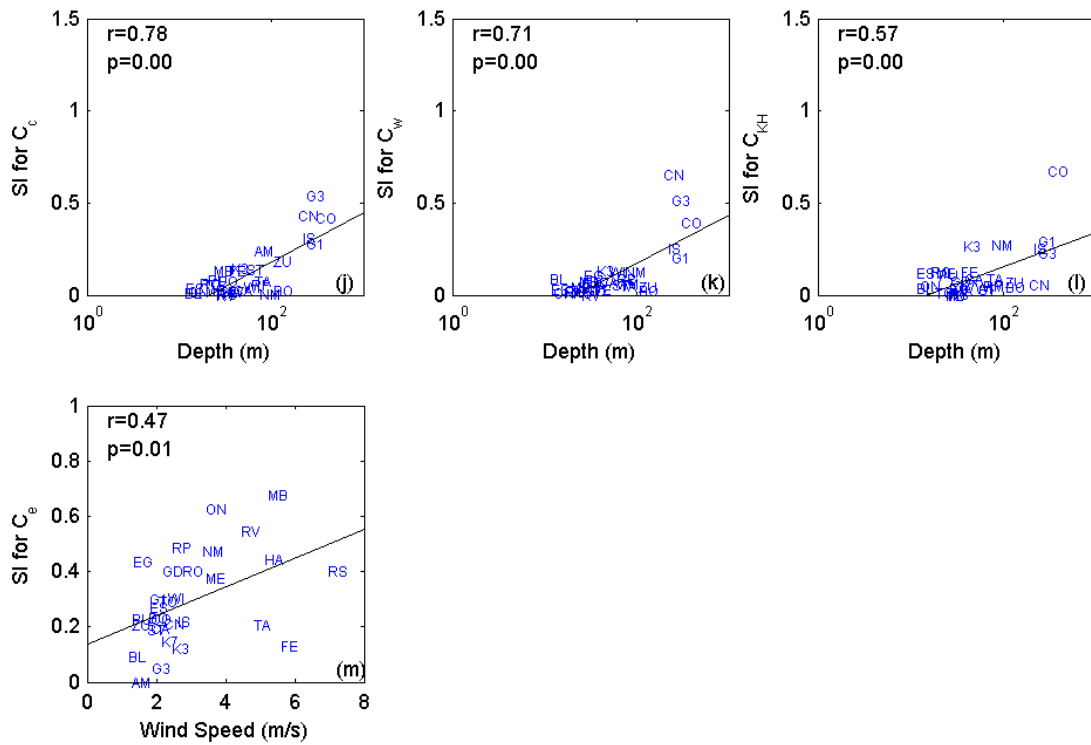
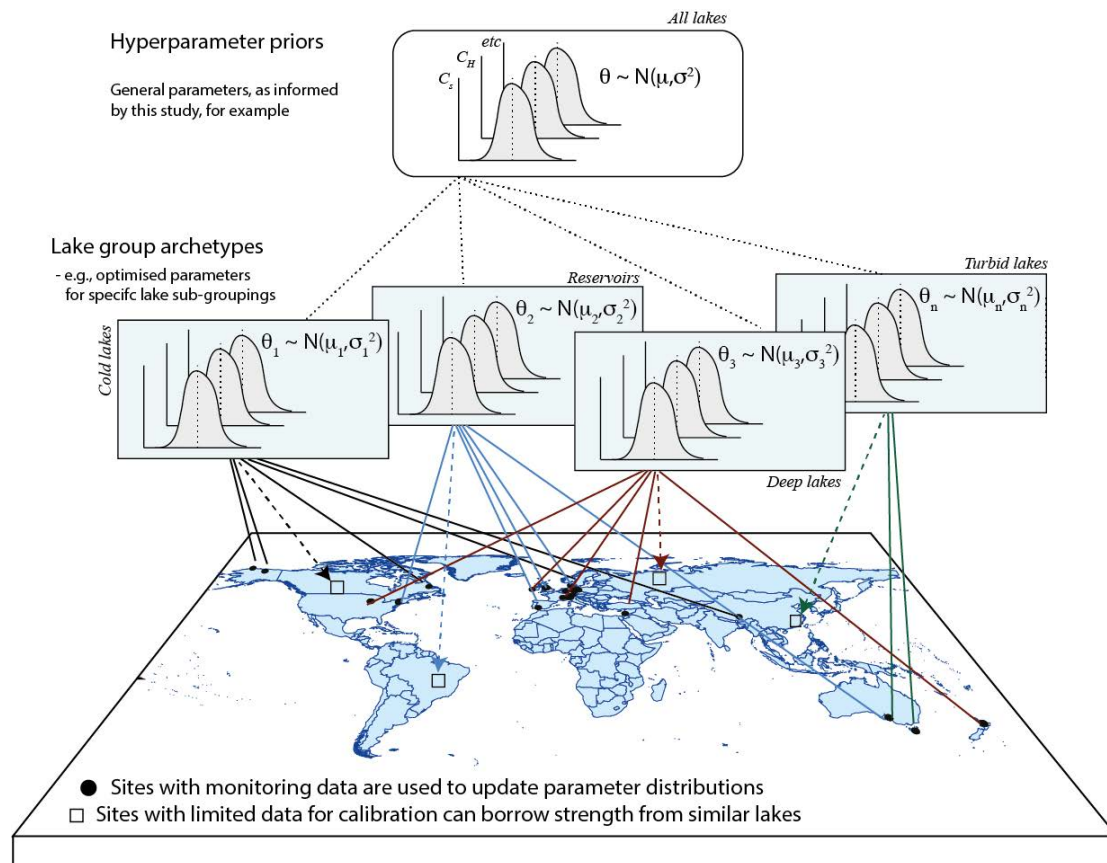


Figure 9 - Significant correlation between sensitivity indices of GLM physical parameters for the prediction of: full profile temperatures and (a) surface area, (b) lake depth and (c) air temperature; epilimnion temperature and (d) air temperature, (e) lake depth and (f) residence time; hypolimnion temperatures and (g) air temperature, (h) short wave radiation and (i) inflow; thermocline depth as a function of lake depth (j-l); and Schmidt stability as a function of wind speed (m).



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Figure 10 - A conceptual overview of future lake modelling applications to best integrate model applications with the increasing volumes of sensor data. In this study no parameter fitting was undertaken for GLM and parameters presented herein could be used as the hyperparameter prior for all lakes within the observatory network. Future applications can improve parameter accuracy within a Bayesian hierarchical framework based on suitable groupings of lakes into distinct archetypes. Other lakes with limited data for robust calibration, can adopt standard model parameters depending on the lake archetype, to which it best relates.

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## 5 Appendix A – Model Input

Table A1 – Input uncertainty ranking system

Rank	0 - ideal	1 - high	2 - medium	3 - low
Morphometry	digitised		estimated from topographic drawing	estimated
Meteorology - Distance	on lake (< 1km)	<5km from lake	<10km from lake	>10km from lake/ estimated
Meteorology - Frequency	sub-hourly	hourly	sub-daily	daily
Flow	gauged	modelled		estimated
Kw	mean from light measurements	Secchi depth mean from > 12 measurements/year	Secchi depth mean from < 12 measurements/year	estimated
Frequency of Observed Data	>=daily	>= weekly	>=monthly	< monthly

Table A2 – Input data quality for each lake, (a) measurement, (b) rank. A value of 9999 indicates that input data has been estimated.

Lake Name	Morphometry	Distance (km)						Sampling interval (hours)						Kw	Number of Obs Data		
		Short Wave Rad.	Long Wave Rad.	Air Temp.	Rel. Hum.	Wind speed	Precipitation	Short Wave Rad.	Long Wave Rad.	Air Temp.	Rel. Hum.	Wind speed	Precipitation			Inflow	Outflow
Alexandrina	digital	1.2	1.2	1.2	1.2	1.2	1.2	0.25	0.25	0.25	0.25	0.25	0.25	model	model	estimated	14
Ammersee	digital	10	12	10	10	10	10	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	24
Blelham	contour	0	0	0	0	0	0	0.03	24.00	0.03	0.03	0.03	0.03	gauge	gauge	secchi	2920
Bourget	contour	2	2	2	2	2	2	1.00	1.00	1.00	1.00	1.00	0.10	gauge	estimated	secchi	74
Cannonsville	contour	0	0	0	0	0	0	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	31
Como	digital	0	50	0	0	0	1	0.02	1.00	0.02	0.02	0.02	1.00	gauge	gauge	secchi	2916
Constance	digital	1.2	1.2	1.2	1.2	1.2	1.2	1.00	1.00	1.00	1.00	1.00	6.00	gauge	gauge	secchi	31
ElGergal	digital	0.3	0.3	0.3	0.3	0.3	0.3	1.00	1.00	1.00	1.00	1.00	24.00	gauge	gauge	secchi	124
Emaiksoun	digital	0	0	0	0	0	0	0.00	0.00	0.00	0.00	0.00	0.00	estimated	estimated	secchi	467
Esthwaite	contour	0	0	0	0	0	0	0.03	24.00	0.03	0.03	0.03	0.03	gauge	gauge	secchi	2799
Feeagh	digital	0.7	0.7	0.35	0.35	0.35	0.35	0.03	24.00	0.00	0.00	0.00	1.00	gauge	estimated	light	2913
Geneva03	contour	1	1	1	1	1	1	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	39
Geneva05	contour	1	1	1	1	1	1	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	30
GrosseDhuenn	estimated	37	9999	22.5	22.5	22.5	9999	1.00	9999.00	1.00	1.00	1.00	9999.00	model	gauge	secchi	17
Harp	contour	0.5	0.5	0.5	0.5	0.5	0.5	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	20
Iseo	digital	0	20	0	0	0	3	0.03	1.00	0.03	0.03	0.03	1.00	gauge	gauge	secchi	24
Kinneret03	digital	0	0	0	0	0	0	0.02	0.02	0.02	0.02	0.02	0.02	gauge	gauge	light	93
Kinneret97	digital	0	0	0	0	0	0	0.02	0.02	0.02	0.02	0.02	0.02	gauge	gauge	light	78
Mendota	contour	2.5	2.5	2.5	2.5	2.5	2.5	<0.01	1.00	<0.01	<0.01	<0.01	1.00	gauge	estimated	light	30
MtBold	digital	12	22	10	10	12	5	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	46
Muggelsee	digital	0	0	0	0	0	0	0.08	24.00	0.08	0.08	0.08	0.08	model	model	estimated	1747
NamCo	digital	1.6	1.6	1.6	1.6	1.6	1.6	24.00	24.00	24.00	24.00	24.00	24.00	estimated	estimated	light	731
Oneida	digital	21	21	21	21	21	21	1.00	1.00	1.00	1.00	1.00	1.00	model	gauge	light	40
Pusiano	digital	11	9999	4	17	11	4	1.00	3.00	1.00	1.00	1.00	1.00	model	model	secchi	2320
Rappbode	digital	16	16	16	16	16	16	1.00	1.00	1.00	1.00	1.00	1.00	model	gauge	secchi	58



3221	Rassnitzersee	digital	12	12	12	12	12	12	24.00	24.00	24.00	24.00	24.00	24.00	model	model	secchi	16
3222	Ravn	contour	50	50	50	50	50	50	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	44
3223	Rotorua	digital	0	0	0	0	0	0	1.00	1.00	1.00	1.00	1.00	1.00	gauge	estimated	secchi	677
3224	Stechlin	digital	5	5	5	5	5	5	3.00	3.00	3.00	3.00	3.00	24.00	estimated	estimated	light	41
3225	Tarawera	digital	15	15	15	15	15	15	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	21
3227	Toolik	digital	0.05	0.05	0.05	0.05	0.05	0.05	1.00	1.00	1.00	1.00	1.00	1.00	gauge	model	light	2920
3228	Windermere	contour	0	0	0	0	0	0	0.03	24.00	0.03	0.03	0.03	0.03	gauge	gauge	secchi	2920
3229	Woods	digital	9999	9999	33.6	33.6	33.6	33.6	1.00	1.00	1.00	1.00	1.00	1.00	estimated	gauge	estimated	731
3230	Zurich	contour	0.5	2.0	0.5	0.5	0.5	0.5	0.17	0.17	0.17	0.17	0.17	0.17	gauge	gauge	secchi	24

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		Distance		Sampling interval				Number of											
Lake Name	Morphometry	Short Wave Rad.	Long Wave Rad.	Air Temp.	Rel. Hum.	Wind speed	Precipitation	Short Wave Rad.	Long Wave Rad.	Air Temp.	Rel. Hum.	Wind speed	Precipitation	Inflow	Outflow	Kw	Obs Data	Mean	
3238	Alexandrina	0	1	1	1	1	1	0	0	0	0	0	0	1	1	3	3	1.33	
3239	Ammersee	0	2	3	2	2	2	2	3	3	3	3	3	3	0	0	1	3	1.53
3240	Blelham	1	0	0	0	0	0	0	3	0	0	0	0	0	0	1	0	0.42	
3241	Bourget	1	1	1	1	1	1	1	1	1	1	1	1	0	0	3	1	2	1.22
3242	Cannonsville	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	1	3	1.33
3243	Como	0	0	3	0	0	0	0	0	1	0	0	0	1	0	0	1	0	0.31
3244	Constance	0	1	1	1	1	1	1	1	1	1	1	1	2	0	0	1	3	1.03
3245	ElGergal	0	0	0	0	0	0	0	1	1	1	1	1	3	0	0	1	1	0.56
3246	Emaiksoun	0	0	0	0	0	0	0	0	0	0	0	0	0	3	3	1	1	0.83
3247	Esthwaite	1	0	0	0	0	0	0	0	3	0	0	0	0	0	0	1	0	0.42
3248	Feeagh	0	0	0	0	0	0	0	0	3	0	0	0	1	0	3	0	0	0.36
3249	Geneva03	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	0	3	1.17
3250	Geneva05	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	0	3	1.17
3251	GrosseDhuenn	3	3	3	3	3	3	3	1	3	1	1	1	3	1	0	1	3	2.03
3252	Harp	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	1	3	1.33
3253	Iseo	0	0	3	0	0	0	1	0	1	0	0	0	1	0	0	1	3	0.83

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3261	Kinneret03	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	2	0.33
3262	Kinneret97	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	2	0.33
3263	Mendota	1	1	1	1	1	1	1	0	1	0	0	0	1	0	3	0	3	1.14
3264	MtBold	0	3	3	2	2	3	1	3	3	3	3	3	3	0	0	0	3	1.39
3265	Muggelsee	0	0	0	0	0	0	0	0	3	0	0	0	0	1	1	3	0	0.75
3267	NamCo	0	1	1	1	1	1	1	3	3	3	3	3	3	3	3	0	1	1.33
3268	Oneida	0	3	3	3	3	3	3	1	1	1	1	1	1	1	0	0	3	1.25
3269	Pusiano	0	3	3	1	3	3	1	1	2	1	1	1	1	1	1	1	0	0.92
3270	Rappbode	0	3	3	3	3	3	3	1	1	1	1	1	1	1	0	1	2	1.25
3272	Rassnitzersee	0	3	3	3	3	3	3	3	3	3	3	3	3	1	1	1	3	1.83
3273	Ravn	1	3	3	3	3	3	3	3	3	3	3	3	3	0	0	1	3	1.83
3274	Rotorua	0	0	0	0	0	0	0	1	1	1	1	1	1	0	3	1	1	0.75
3276	Stechlin	0	1	1	1	1	1	1	2	2	2	2	2	3	3	3	0	3	1.53
3277	Tarawera	0	3	3	3	3	3	3	3	3	3	3	3	3	0	0	0	3	1.50
3278	Tarawera	0	3	3	3	3	3	3	3	3	3	3	3	3	0	0	0	3	1.50
3279	Toolik	0	0	0	0	0	0	0	1	1	1	1	1	1	0	1	0	0	0.25
3280	Windermere	1	0	0	0	0	0	0	0	3	0	0	0	0	0	0	1	0	0.42
3281	Woods	0	3	3	3	3	3	3	1	1	1	1	1	1	3	0	3	0	1.42
3282	Woods	0	3	3	3	3	3	3	1	1	1	1	1	1	3	0	3	0	1.42
3283	Zurich	1	0	1	0	0	0	0	0	0	0	0	0	0	0	0	1	3	0.83

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Table A3 – Input lake characteristics used for comparative analysis. Values for lake depth, surface area, volume and meteorological data are averaged over the simulation period. Area: Depth Ratio, is the ratio between average lake surface area and lake depth.

Abbrev.	Lake Name	Depth (m)	Surface Area (m <sup>2</sup> )	Area: Depth Ratio	Volume (m <sup>3</sup> )	Length: Width Ratio	Inflow (m <sup>3</sup> /d)	Res. Time (days)	Short Wave Radiation (W/m <sup>2</sup> )	Air Temp. (°C)	Wind Speed (m/s)	K <sub>w</sub> (m <sup>-1</sup> )	Start Date
AL	Alexandrina	6.1	587,158,666	9.66E+07	1.11E+09	1.14	1.12E+07	98.96	186.98	14.88	3.81	0.30	1-Jul-10
AM	Ammersee	83.8	46,427,854	5.64E+05	1.81E+09	3.00	1.70E+06	1061.06	125.18	8.40	1.58	0.35	1-Jan-09
BL	Blelham	14.7	103,802	7.07E+03	5.78E+05	3.75	2.05E+04	28.14	100.78	9.41	1.46	0.67	1-Jan-08
BO	Bourget	140.5	40,379,890	3.03E+05	3.31E+09	5.83	5.48E+05	6035.65	117.27	11.48	2.08	0.40	1-Jan-09
CA	Cannonsville	49.7	18,014,685	3.83E+05	3.28E+08	27.40	2.80E+06	116.90	128.97	7.36	2.11	0.40	1-Jan-03
CO	Como	400.2	133,890,923	3.67E+05	2.21E+10	15.00	8.42E+06	2618.03	134.28	13.96	2.14	0.46	1-Jan-05
CN	Constance	253.4	465,703,202	1.87E+06	4.70E+10	4.00	3.26E+07	1443.46	126.05	10.26	2.51	0.23	1-Jan-94
EG	El Gergal	34.9	1,951,384	1.35E+05	2.47E+07	19.38	7.40E+05	33.43	189.57	18.14	1.64	0.79	1-Jan-01
EM	Emaiksoun	2.2	1,871,484	8.77E+05	2.89E+06	2.30	0.00E+00	N/A	93.80	-9.09	4.10	0.27	1-Jan-12
ES	Esthwaite	15.0	1,003,558	6.69E+04	6.38E+06	7.67	9.79E+04	65.16	94.79	9.50	2.09	1.07	1-Jan-08
FE	Feeagh	44.9	3,620,712	8.78E+04	6.49E+07	3.86	6.02E+05	107.72	115.20	10.11	5.87	0.30	1-Jan-11
G1	Geneva01	308.9	577,950,600	1.87E+06	9.12E+10	5.21	3.13E+07	2914.51	135.43	10.91	2.10	0.10	1-Jan-03
G3	Geneva03	308.8	577,404,068	1.87E+06	9.12E+10	5.21	2.77E+07	3292.98	148.04	10.95	2.14	0.10	1-Jan-01
GD	GrosseDhuenn	33.8	2,214,418	1.11E+05	3.26E+07	8.36	1.05E+05	309.21	108.20	9.51	2.50	0.60	1-Jan-96
HA	Harp	37.2	706,244	1.92E+04	9.31E+06	1.50	7.96E+03	1170.14	139.26	3.73	5.39	0.77	1-Jan-92
IS	Iseo	260.5	60,829,158	2.34E+05	7.91E+09	8.33	5.48E+06	1443.16	140.28	13.55	2.83	0.25	1-Jan-10
K3	Kinneret03	47.5	171,786,293	3.64E+06	4.96E+09	1.62	2.93E+06	1690.75	219.52	22.02	2.74	0.59	1-Jan-03
K7	Kinneret97	43.6	166,425,534	3.97E+06	4.29E+09	1.62	1.62E+06	2643.24	215.43	22.37	2.41	0.57	1-Jan-97
ME	Mendota	25.0	39,229,728	1.58E+06	4.96E+08	2.00	4.96E+05	1000.35	167.57	8.23	3.74	0.69	1-Jan-09
MB	MtBold	31.7	1,703,574	1.07E+05	1.99E+07	7.33	1.70E+05	116.81	195.67	16.36	5.55	0.98	1-Jul-03
MG	Muggelsee	7.8	7,173,756	9.43E+05	3.38E+07	1.69	3.65E+05	92.55	117.60	10.26	3.85	1.05	1-Jan-04
NM	NamCo	98.9	1,942,514,246	2.04E+07	1.00E+11	2.38	0.00E+00	N/A	244.03	-1.13	3.64	0.30	1-Jan-12
ON	Oneida	16.4	199,785,009	1.26E+07	1.35E+09	3.79	5.68E+06	236.69	131.55	10.96	3.73	0.60	1-Jan-11
PU	Pusiano	26.8	6,537,710	3.03E+05	7.72E+07	1.97	2.39E+05	322.45	131.21	13.91	1.62	0.66	1-Jan-02

3341	RP	Rappbode	79.4	3,453,019	5.47E+04	8.60E+07	16.00	2.31E+05	371.77	118.44	7.82	2.77	0.40	1-Jan-08
3342	RS	Rassnitzersee	34.8	2,714,136	8.71E+04	5.42E+07	1.24	5.14E+03	10551.99	107.95	9.66	4.26	0.44	1-Jan-01
3343	RV	Ravn	32.3	1,748,402	6.00E+04	2.61E+07	1.33	3.57E+04	729.93	116.19	8.37	4.72	0.50	1-Jan-03
3344	RO	Rotorua	21.8	78,779,626	3.66E+06	7.95E+08	1.17	1.23E+06	645.34	170.60	12.64	3.07	0.80	1-Jul-07
3346	ST	Stechlin	69.5	4,230,060	6.11E+04	9.75E+07	0.76	4.00E+03	24364.00	104.49	8.93	2.02	0.25	1-Jan-01
3347	TA	Tarawera	82.8	39,406,707	4.95E+05	2.18E+09	1.17	6.11E+05	3570.44	162.49	13.34	5.07	0.21	1-Jul-02
3348	TO	Toolik	23.5	920,947	3.99E+04	6.18E+06	1.50	1.29E+04	479.13	117.31	-7.54	2.36	0.65	1-Jan-06
3349	WI	Windermere	63.7	14,360,584	2.32E+05	5.15E+08	12.13	1.61E+06	319.72	104.08	9.44	2.60	0.60	1-Jan-08
3351	WO	Woods	9.2	13,451,648	1.63E+06	5.55E+07	1.31	4.93E+04	1125.18	162.37	6.62	4.97	0.80	1-Jul-11
3352	ZU	Zurich	136.0	66,593,085	4.90E+05	3.37E+09	11.20	6.20E+06	543.19	138.46	10.06	1.57	0.34	1-Jan-03

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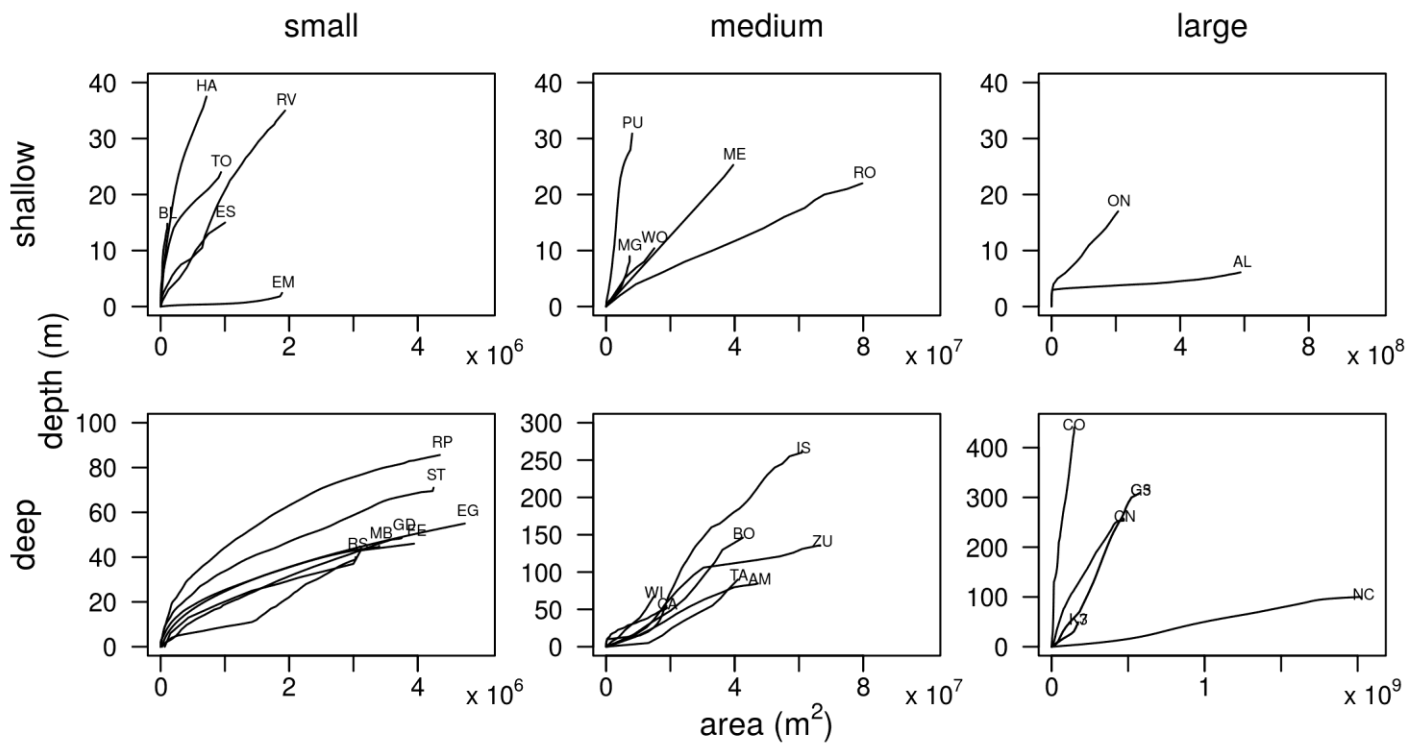


Figure A1 - Hypsographic curves for (a) small, shallow lakes, (b) medium, shallow lakes, (c) large, shallow lakes, (d) small, deep lakes, (e) medium, deep lakes, (f) large, deep lakes.

## 6 Appendix B – Analysis of Model Performance

Table B1 – Model performance metrics for base simulations using standard parameter set. Note that for fully mixed lakes or for lakes where temperature profiles were shallower than the thermocline depth, *NMAE* values are listed as not applicable (N/A).

	Full-profile Temperature					Epilimnion Temperature					Hypolimnion Temperature					Thermocline Depth					Schmidt Stability				
	RMS E	MEF F	r	PRE	NMA E	RMS E	MEF F	r	PRE	NMA E	RMS E	MEF F	r	PRE	NMA E	RMSE	MEF F	r	PRE	NMA E	RMS E	MEF F	r	PRE	NMA E
Alexandrina	1.43	0.86	0.98	-3.9	0.07	1.44	0.86	0.98	-4.0	0.07	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A
Ammersee	2.07	0.79	0.91	-6.0	0.19	2.78	0.84	0.99	14.8	0.20	0.78	0.30	0.59	-1.7	0.13	28.0	0.28	0.70	15.6	0.40	278	0.95	0.99	10.6	0.17
Blelham	1.32	0.89	0.97	-3.0	0.12	2.02	0.85	0.96	-12.0	0.13	3.24	-1.60	0.84	21.6	0.31	2.6	0.54	0.80	-0.2	0.18	24	0.65	0.96	27.0	0.45
Bourget	0.87	0.93	0.97	-4.5	0.08	2.01	0.91	0.97	0.1	0.11	0.53	-0.95	0.40	-4.9	0.07	36.4	0.56	0.79	3.9	0.32	823	0.98	0.99	0.9	0.09
Cannonsville	1.33	0.94	0.97	-1.0	0.10	1.06	0.97	0.99	-1.8	0.05	1.70	0.57	0.79	-5.3	0.15	11.6	0.32	0.62	-15.0	0.39	125	0.96	0.98	-5.7	0.12
Como	1.13	0.86	0.94	4.2	0.10	4.32	0.52	0.78	-10.0	0.17	0.57	-0.49	0.48	3.1	0.06	82.7	-0.43	0.67	26.6	0.64	5498	0.90	0.96	-10.1	0.19
Constance	1.08	0.95	0.98	2.8	0.08	1.49	0.95	0.98	4.4	0.09	0.44	0.25	0.74	2.8	0.07	31.0	0.92	0.96	6.7	0.11	1372	0.95	0.99	7.0	0.16
ElGergal	1.72	0.81	0.95	1.1	0.08	1.54	0.91	0.98	1.4	0.06	1.38	0.38	0.80	-5.6	0.07	8.5	0.55	0.79	16.8	0.30	328	0.74	0.97	24.1	0.27
Emaiksoun	0.80	0.95	0.99	8.0	0.08	0.80	0.95	0.99	8.0	0.08	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A
Esthwaite	1.49	0.89	0.98	1.4	0.13	1.50	0.93	0.98	-8.9	0.11	3.64	-0.66	0.92	21.4	0.35	2.7	0.40	0.71	10.3	0.15	12	0.93	0.97	-9.9	0.24
Feeagh	0.72	0.95	0.99	1.2	0.06	0.53	0.98	0.99	-0.9	0.04	1.30	0.80	0.97	5.8	0.09	9.7	0.41	0.67	6.3	0.14	34	0.90	0.95	11.6	0.30
Geneva01	1.18	0.92	0.96	-5.3	0.09	2.07	0.86	0.95	-7.0	0.11	0.36	-0.34	0.65	-2.6	0.04	94.4	0.46	0.80	36.3	0.41	3350	0.87	0.95	-14.0	0.22
Geneva03	1.16	0.93	0.97	-2.1	0.08	1.02	0.98	0.99	-1.0	0.05	0.34	-0.39	0.67	2.6	0.04	123.0	0.22	0.59	-15.6	0.52	3977	0.88	0.94	-3.6	0.20
GrosseDhuen n	0.81	0.97	0.99	-0.3	0.07	1.05	0.97	0.99	-3.1	0.05	1.02	0.78	0.90	-2.5	0.09	13.5	0.37	0.64	-4.1	0.37	69	0.98	0.99	0.7	0.09
Harp	1.54	0.92	0.96	-4.6	0.18	1.60	0.95	0.98	1.7	0.12	1.63	-0.79	0.70	-12.5	0.27	5.9	-0.48	0.38	34.3	0.68	60	0.94	0.98	-1.9	0.19
Iseo	1.07	0.96	0.98	-3.4	0.08	1.83	0.92	0.97	-8.0	0.10	0.55	0.03	0.56	-1.0	0.07	122.9	-0.33	0.39	19.3	0.76	2620	0.94	0.97	-0.5	0.16
Kinneret03	1.60	0.88	0.96	0.9	0.07	1.76	0.87	0.97	1.4	0.07	1.28	-3.04	0.28	-0.6	0.07	10.9	0.31	0.66	15.4	0.28	527	0.88	0.99	17.3	0.20

3461	Kinneret97	1.49	0.87	0.97	3.1	0.05	1.65	0.89	0.99	4.6	0.06	1.10	-2.16	0.46	3.0	0.05	7.8	0.56	0.79	2.6	0.15	571	0.87	0.99	19.1	0.21
3462	Mendota	1.60	0.92	0.97	5.9	0.11	1.94	0.94	0.98	7.9	0.10	1.42	0.84	0.95	7.8	0.11	7.8	0.15	0.56	5.3	0.30	96	0.88	0.99	20.0	0.23
3463	MtBold	1.47	0.87	0.96	4.8	0.08	1.74	0.80	0.94	6.0	0.08	1.08	0.90	0.96	1.7	0.06	11.4	0.50	0.75	-9.3	0.25	146	0.57	0.94	28.7	0.43
3464	Muggelsee	1.40	0.92	0.99	4.6	0.07	1.24	0.94	0.98	2.7	0.06	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A
3465	NamCo	1.13	0.85	0.95	17.6	0.23	1.04	0.93	0.98	11.1	0.17	0.85	0.70	0.93	19.1	0.22	29.6	0.16	0.58	13.4	0.28	304	0.81	0.98	25.7	0.35
3466	Oneida	1.26	0.91	0.96	1.1	0.04	0.65	0.97	0.99	-1.0	0.03	1.83	0.79	0.91	2.8	0.06	3.0	-1.49	0.10	17.1	0.19	26	-0.48	0.63	-86.3	0.86
3468	Pusiano	1.74	0.90	0.97	-3.8	0.14	2.38	0.90	0.98	1.2	0.11	2.81	-0.81	0.26	-19.3	0.26	5.3	0.69	0.84	6.4	0.24	146	0.88	0.98	13.3	0.19
3469	Rappbode	1.15	0.91	0.97	-6.0	0.14	1.22	0.96	0.99	0.4	0.08	0.77	0.41	0.74	-5.3	0.12	13.3	0.67	0.84	3.6	0.23	254	0.92	0.99	11.5	0.16
3470	Rassnitzersee	1.64	0.80	0.96	-15.3	0.17	1.82	0.90	0.99	-14.9	0.15	1.73	-1.00	0.77	-23.3	0.23	5.2	0.82	0.94	14.4	0.15	116	0.94	0.98	-14.9	0.17
3471	Ravn	1.94	0.85	0.94	-9.3	0.19	1.81	0.91	0.99	-12.9	0.14	1.53	0.36	0.88	5.2	0.21	8.3	0.34	0.75	10.4	0.27	149	0.80	0.98	-33.1	0.34
3472	Rotorua	1.33	0.91	0.99	3.8	0.07	1.33	0.91	0.99	3.9	0.08	1.38	0.89	0.99	3.9	0.08	3.9	0.04	0.36	4.9	0.09	11	0.74	0.88	-6.6	0.43
3473	Stechlin	1.11	0.91	0.96	-7.3	0.13	1.73	0.93	0.99	2.5	0.11	0.77	0.04	0.80	-1.0	0.11	21.2	0.42	0.74	11.9	0.33	159	0.96	0.98	-3.9	0.14
3474	Tarawera	0.77	0.86	0.95	-3.3	0.04	0.82	0.93	0.98	-1.4	0.04	0.47	-0.08	0.60	-2.4	0.03	21.4	0.46	0.76	-1.3	0.27	424	0.97	0.99	-6.6	0.10
3475	Toolik	1.36	0.77	0.88	2.1	0.25	1.94	0.82	0.91	1.4	0.26	1.15	0.61	0.81	7.4	0.25	11.3	-0.89	0.29	52.3	0.61	17	0.74	0.86	-15.7	0.43
3476	Windermere	1.61	0.82	0.95	12.4	0.14	3.21	0.54	0.81	-11.1	0.23	2.39	-0.82	0.85	25.2	0.26	10.2	0.20	0.79	15.8	0.22	271	0.90	0.98	16.2	0.21
3477	Woods	2.14	0.82	0.99	-9.0	0.17	2.13	0.82	0.99	-9.1	0.17	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A
3478	Zurich	1.24	0.94	0.98	7.6	0.12	1.48	0.96	0.99	5.3	0.09	1.11	-3.05	0.60	13.6	0.16	50.8	0.30	0.64	-7.6	0.42	816	0.94	1.00	16.2	0.17
3479	<b>Mean</b>	<b>1.34</b>	<b>0.89</b>	<b>0.96</b>	<b>-0.16</b>	<b>0.11</b>	<b>1.67</b>	<b>0.89</b>	<b>0.97</b>	<b>-0.84</b>	<b>0.10</b>	<b>1.31</b>	<b>-0.25</b>	<b>0.73</b>	<b>1.97</b>	<b>0.14</b>	<b>26.5</b>	<b>0.23</b>	<b>0.66</b>	<b>9.89</b>	<b>0.32</b>	<b>753</b>	<b>0.83</b>	<b>0.96</b>	<b>1.23</b>	<b>0.25</b>
3480	<b>Median</b>	<b>1.33</b>	<b>0.90</b>	<b>0.97</b>	<b>0.26</b>	<b>0.09</b>	<b>1.62</b>	<b>0.92</b>	<b>0.98</b>	<b>0.23</b>	<b>0.09</b>	<b>1.13</b>	<b>0.03</b>	<b>0.78</b>	<b>2.13</b>	<b>0.10</b>	<b>11.3</b>	<b>0.36</b>	<b>0.71</b>	<b>8.53</b>	<b>0.28</b>	<b>206</b>	<b>0.90</b>	<b>0.98</b>	<b>0.81</b>	<b>0.20</b>
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3501 Table B2 – Significance and correlation between model performance metrics and input uncertainty for morphometry (morph), distance to meteorological  
 3502 station (distmet), frequency of meteorological data (freqmet), determination of inflow and outflows (flow), light extinction coefficient (Kw), frequency of  
 3503 observed temperature profiles (obs) and average error ranking (mean). Significant correlations highlighted in red and corresponding r in yellow.

		Full Profile Temperature				Epilimnion Temperature				Hypolimnion Temperature				Thermocline Depth			Schmidt Stability									
		RMS	MEF			NMA	RMS	MEF		NMA	RMS	MEF		NMA	RMS		NSE	r	PRE	NMA	RMS	MEF			NMA	
		E	F	r	PRE	E	E	F	r	PRE	E	E	F	r	PRE	E	E				E	E	F	r	PRE	E
r	morph	0.05	0.26	0.03	-0.01	0.04	0.05	0.05	-0.10	-0.25	-0.02	0.09	0.12	0.22	0.22	0.14	-0.02	0.18	0.23	0.01	-0.01	-0.06	0.25	0.24	-0.09	-0.27
	distmet	0.27	-0.27	-0.16	-0.44	0.17	0.00	0.04	0.20	-0.18	0.00	0.14	-0.21	-0.22	-0.35	0.09	-0.18	0.09	0.07	-0.13	-0.14	-0.22	-0.15	-0.10	-0.38	0.04
	freqmet	0.11	-0.19	-0.27	-0.32	0.34	0.00	0.07	0.06	-0.01	0.11	-0.09	-0.11	-0.01	-0.35	0.11	0.24	-0.34	-0.18	0.32	0.50	-0.08	0.11	0.10	-0.21	-0.13
	flow	-0.21	0.06	0.23	0.17	0.16	-0.24	0.18	0.24	0.29	0.09	-0.05	-0.01	0.07	0.00	0.05	-0.10	-0.19	-0.17	0.14	-0.01	-0.25	-0.08	-0.08	0.16	0.14
	Kw	0.30	-0.03	0.38	-0.16	0.03	0.15	-0.10	0.08	-0.27	0.06	0.33	-0.17	-0.19	-0.16	0.41	-0.17	0.42	0.40	-0.27	-0.29	0.02	0.24	0.26	0.03	-0.29
	obs	-0.03	0.21	-0.02	-0.31	-0.16	-0.17	0.30	0.31	0.13	-0.32	-0.33	-0.13	-0.09	-0.29	-0.32	0.34	-0.18	-0.07	0.25	0.40	-0.01	0.07	0.10	-0.31	-0.25
	mean	0.17	-0.02	0.01	-0.44	0.20	-0.10	0.22	0.32	-0.06	-0.05	-0.06	-0.18	-0.09	-0.37	0.05	0.10	-0.11	0.00	0.17	0.24	-0.20	0.08	0.11	-0.33	-0.21
	p	morph	0.80	0.14	0.88	0.94	0.83	0.77	0.78	0.57	0.15	0.93	0.65	0.51	0.24	0.25	0.47	0.93	0.33	0.22	0.95	0.97	0.76	0.19	0.20	0.64
distmet		0.12	0.13	0.38	0.01	0.32	0.99	0.82	0.27	0.32	0.99	0.47	0.26	0.24	0.06	0.64	0.34	0.65	0.71	0.51	0.44	0.23	0.44	0.60	0.04	0.84
freqmet		0.54	0.29	0.12	0.07	0.05	0.98	0.69	0.72	0.94	0.54	0.64	0.55	0.96	0.06	0.57	0.19	0.07	0.33	0.09	0.01	0.69	0.56	0.58	0.26	0.48
flow		0.24	0.72	0.19	0.34	0.35	0.17	0.31	0.17	0.09	0.62	0.79	0.94	0.72	0.99	0.79	0.59	0.31	0.36	0.46	0.94	0.18	0.69	0.67	0.41	0.46
Kw		0.08	0.85	0.03	0.38	0.89	0.41	0.59	0.65	0.12	0.73	0.07	0.36	0.32	0.39	0.03	0.36	0.02	0.03	0.15	0.12	0.92	0.21	0.16	0.87	0.12
obs		0.85	0.23	0.92	0.07	0.37	0.33	0.09	0.07	0.46	0.07	0.08	0.51	0.65	0.12	0.09	0.06	0.33	0.72	0.18	0.03	0.96	0.71	0.61	0.09	0.18
mean		0.33	0.90	0.97	0.01	0.25	0.57	0.22	0.06	0.73	0.76	0.77	0.33	0.64	0.05	0.77	0.59	0.55	0.99	0.36	0.20	0.29	0.66	0.55	0.07	0.27



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3542 Table B3 – Significance (p) and correlation (r) between model performance metrics and mean values of lake volume (V), surface area (Area), depth (D),

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3544 surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature (T<sub>air</sub>),

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3546 wind speed (u<sub>wind</sub>), light extinction coefficient (K<sub>w</sub>), latitude (Lat), Lake Number (LN) and percent time when LN is less than one (%<1). Significant

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3548 correlations highlighted in red and corresponding r in yellow.

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		Full-profile Temperature					Epilimnion Temperature					Hypolimnion Temperature					Thermocline Depth					Schmidt Stability				
		RMS	MEF	r	PRE	NMA	RMS	MEF	r	PRE	NMA	RMS	MEF	r	PRE	NMA	RMS	MEF	r	PRE	NMA	RMS	MEF	r	PRE	NMA
		E	F			E	E	F			E	E	F			E	E	F			E	E	F			E
r	V	-0.20	0.14	0.00	0.19	-0.22	0.12	-0.09	-0.13	0.18	-0.10	-0.65	-0.11	-0.37	0.02	-0.60	0.65	-0.03	-0.02	-0.05	0.16	0.63	0.07	0.02	-0.05	-0.20
	Area	-0.13	0.12	0.08	0.26	-0.26	0.01	-0.04	-0.03	0.26	-0.16	-0.56	-0.07	-0.30	0.06	-0.58	0.52	-0.10	-0.14	-0.06	0.05	0.50	-0.07	-0.11	-0.10	-0.05
	D	-0.25	0.14	-0.21	-0.03	-0.02	0.31	-0.17	-0.29	-0.04	0.07	-0.74	-0.12	-0.42	-0.08	-0.50	0.84	0.11	0.23	0.02	0.47	0.80	0.38	0.31	0.10	-0.50
	A/D	0.00	0.03	0.21	0.32	-0.30	-0.17	0.05	0.13	0.32	-0.25	-0.35	0.00	-0.15	0.09	-0.51	0.22	-0.15	-0.26	-0.11	-0.18	0.22	-0.29	-0.29	-0.14	0.20
	L/W	-0.09	0.12	0.02	0.14	-0.13	0.12	-0.16	-0.28	-0.14	-0.12	0.01	0.10	0.04	0.01	-0.09	0.12	0.06	0.11	-0.24	0.18	0.19	0.10	0.14	0.12	-0.21
	Inf	-0.23	0.21	0.02	0.46	-0.53	0.03	-0.14	-0.21	0.23	-0.25	-0.43	-0.02	-0.26	0.19	-0.59	0.57	-0.19	-0.18	-0.05	0.17	0.58	-0.14	-0.21	-0.05	0.06
	RT	-0.16	0.01	-0.19	-0.38	0.13	0.20	-0.02	-0.02	0.00	0.19	-0.58	-0.21	-0.43	-0.38	-0.33	0.38	0.03	0.09	0.23	0.26	0.36	0.32	0.18	-0.21	-0.44
	sw	0.18	-0.21	0.00	0.32	-0.12	-0.10	0.01	0.10	0.38	-0.18	-0.29	-0.09	-0.14	0.03	-0.38	0.01	0.01	-0.09	-0.09	-0.06	0.02	-0.08	0.04	0.36	0.12
	T <sub>air</sub>	0.16	0.05	0.21	-0.18	-0.59	0.16	-0.12	-0.01	-0.16	-0.46	-0.04	-0.40	-0.43	-0.21	-0.52	0.08	0.34	0.29	-0.37	-0.28	0.19	-0.02	0.14	0.18	-0.13
	u <sub>wind</sub>	0.03	-0.12	0.12	-0.19	0.05	-0.29	0.13	0.18	-0.15	-0.09	-0.04	0.24	0.35	-0.27	0.02	-0.26	-0.03	-0.05	0.03	-0.18	-0.25	-0.11	-0.10	-0.18	0.16
	K <sub>w</sub>	0.47	-0.22	0.04	0.14	0.08	0.12	-0.14	-0.06	-0.05	0.08	0.67	0.02	0.27	0.15	0.43	-0.61	-0.14	-0.16	0.06	-0.23	-0.46	-0.26	-0.17	0.09	0.36
Lat	-0.17	0.05	-0.16	-0.11	0.32	0.04	-0.03	-0.13	-0.22	0.35	0.24	0.12	0.23	0.13	0.48	-0.03	-0.13	0.02	0.32	0.16	-0.05	0.11	-0.05	-0.27	-0.01	
LN	0.21	-0.18	0.09	0.14	-0.09	0.36	-0.39	-0.31	-0.30	0.17	0.46	-0.55	-0.19	0.43	0.31	-0.14	0.31	0.41	-0.02	-0.18	-0.03	0.25	0.37	0.42	-0.22	
%<1	-0.14	0.10	0.10	0.25	-0.16	-0.43	0.24	0.20	0.23	-0.24	0.04	0.39	0.42	0.08	-0.14	-0.25	-0.56	-0.70	0.07	-0.29	-0.25	-0.78	-0.82	-0.53	0.78	
p	V	0.25	0.41	0.99	0.27	0.21	0.50	0.60	0.46	0.31	0.57	0.00	0.57	0.04	0.92	0.00	0.00	0.89	0.90	0.79	0.38	0.00	0.70	0.92	0.81	0.30
	Area	0.45	0.52	0.65	0.14	0.14	0.96	0.82	0.86	0.15	0.36	0.00	0.71	0.11	0.76	0.00	0.00	0.60	0.47	0.74	0.78	0.00	0.72	0.55	0.60	0.81
	D	0.15	0.42	0.23	0.87	0.90	0.08	0.35	0.09	0.84	0.68	0.00	0.54	0.02	0.66	0.00	0.00	0.55	0.23	0.91	0.01	0.00	0.04	0.10	0.62	0.01
	A/D	0.98	0.85	0.22	0.06	0.08	0.34	0.79	0.45	0.06	0.15	0.06	0.99	0.43	0.65	0.00	0.24	0.41	0.16	0.56	0.34	0.23	0.12	0.12	0.45	0.30
	L/W	0.61	0.49	0.92	0.42	0.48	0.50	0.36	0.11	0.43	0.51	0.94	0.62	0.83	0.94	0.63	0.51	0.74	0.56	0.19	0.33	0.32	0.60	0.47	0.53	0.25
	Inf	0.22	0.25	0.91	0.01	0.00	0.86	0.45	0.25	0.21	0.18	0.02	0.91	0.18	0.33	0.00	0.00	0.33	0.37	0.79	0.38	0.00	0.49	0.29	0.79	0.78
	RT	0.40	0.95	0.32	0.04	0.47	0.28	0.92	0.94	1.00	0.31	0.00	0.29	0.02	0.05	0.09	0.05	0.90	0.64	0.24	0.18	0.06	0.10	0.35	0.29	0.02

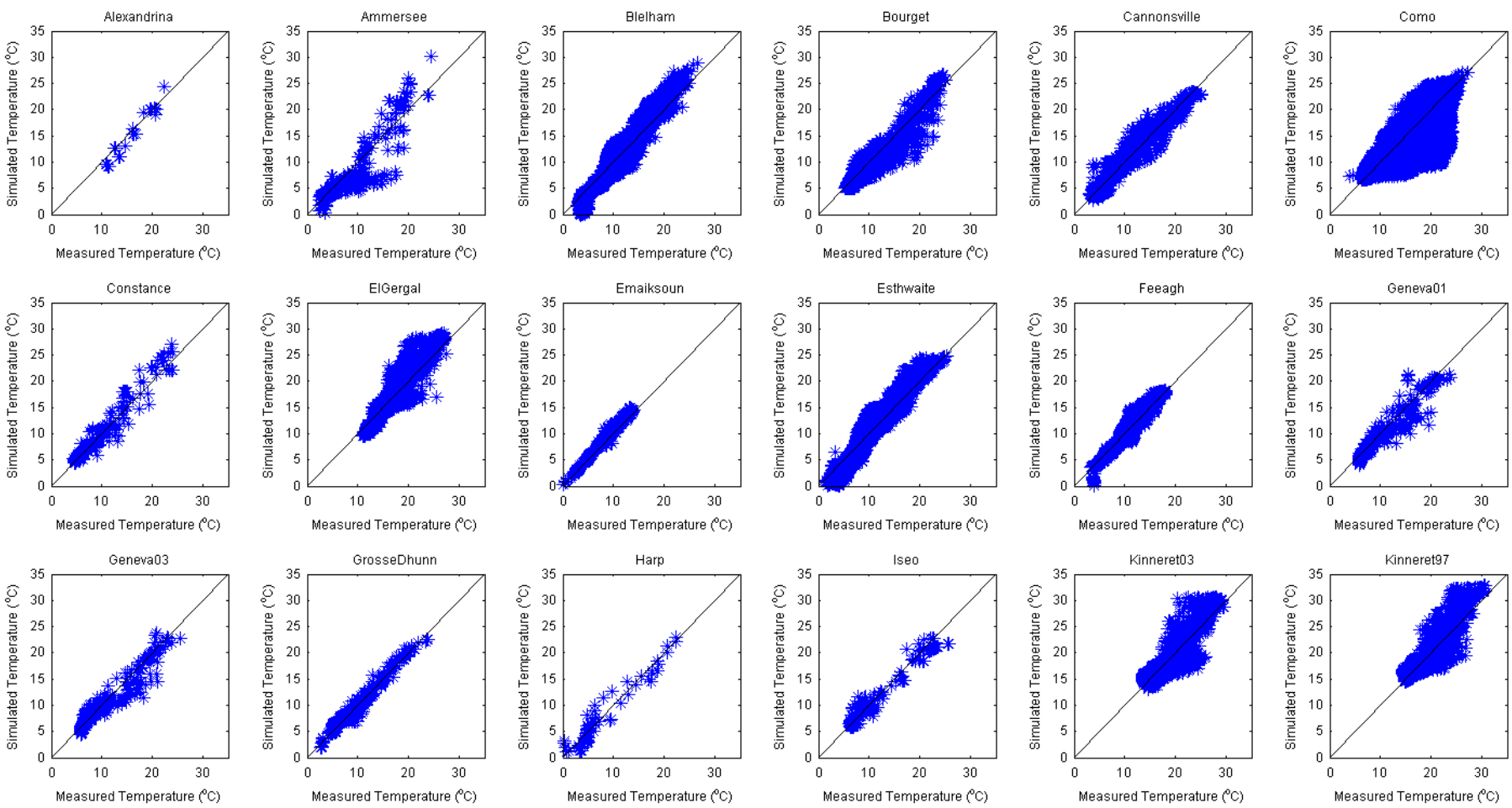
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3581	sw	0.30	0.24	1.00	0.06	0.49	0.59	0.98	0.57	0.02	0.30	0.12	0.63	0.45	0.88	0.04	0.96	0.94	0.62	0.63	0.75	0.93	0.66	0.82	0.05	0.54
3582	T <sub>air</sub>	0.37	0.78	0.23	0.32	0.00	0.37	0.52	0.98	0.38	0.01	0.85	0.03	0.02	0.26	0.00	0.69	0.07	0.12	0.04	0.14	0.31	0.91	0.47	0.35	0.49
3583	u <sub>wind</sub>	0.86	0.51	0.49	0.27	0.80	0.09	0.48	0.32	0.39	0.63	0.83	0.20	0.06	0.15	0.91	0.16	0.88	0.79	0.89	0.35	0.17	0.57	0.62	0.35	0.40
3584	Kw	0.00	0.20	0.84	0.42	0.66	0.52	0.44	0.72	0.77	0.67	0.00	0.90	0.14	0.42	0.02	0.00	0.47	0.39	0.74	0.22	0.01	0.17	0.38	0.64	0.05
3585	Lat	0.33	0.79	0.36	0.53	0.07	0.82	0.89	0.46	0.22	0.04	0.20	0.52	0.22	0.49	0.01	0.88	0.50	0.93	0.08	0.39	0.77	0.57	0.80	0.16	0.96
3586	LN	0.26	0.35	0.62	0.45	0.63	0.05	0.03	0.10	0.10	0.37	0.01	0.00	0.30	0.02	0.10	0.48	0.10	0.02	0.92	0.35	0.86	0.18	0.04	0.02	0.24
3587	%<1	0.45	0.60	0.61	0.19	0.41	0.02	0.20	0.29	0.23	0.19	0.82	0.03	0.02	0.66	0.45	0.19	0.00	0.00	0.72	0.12	0.18	0.00	0.00	0.00	0.00
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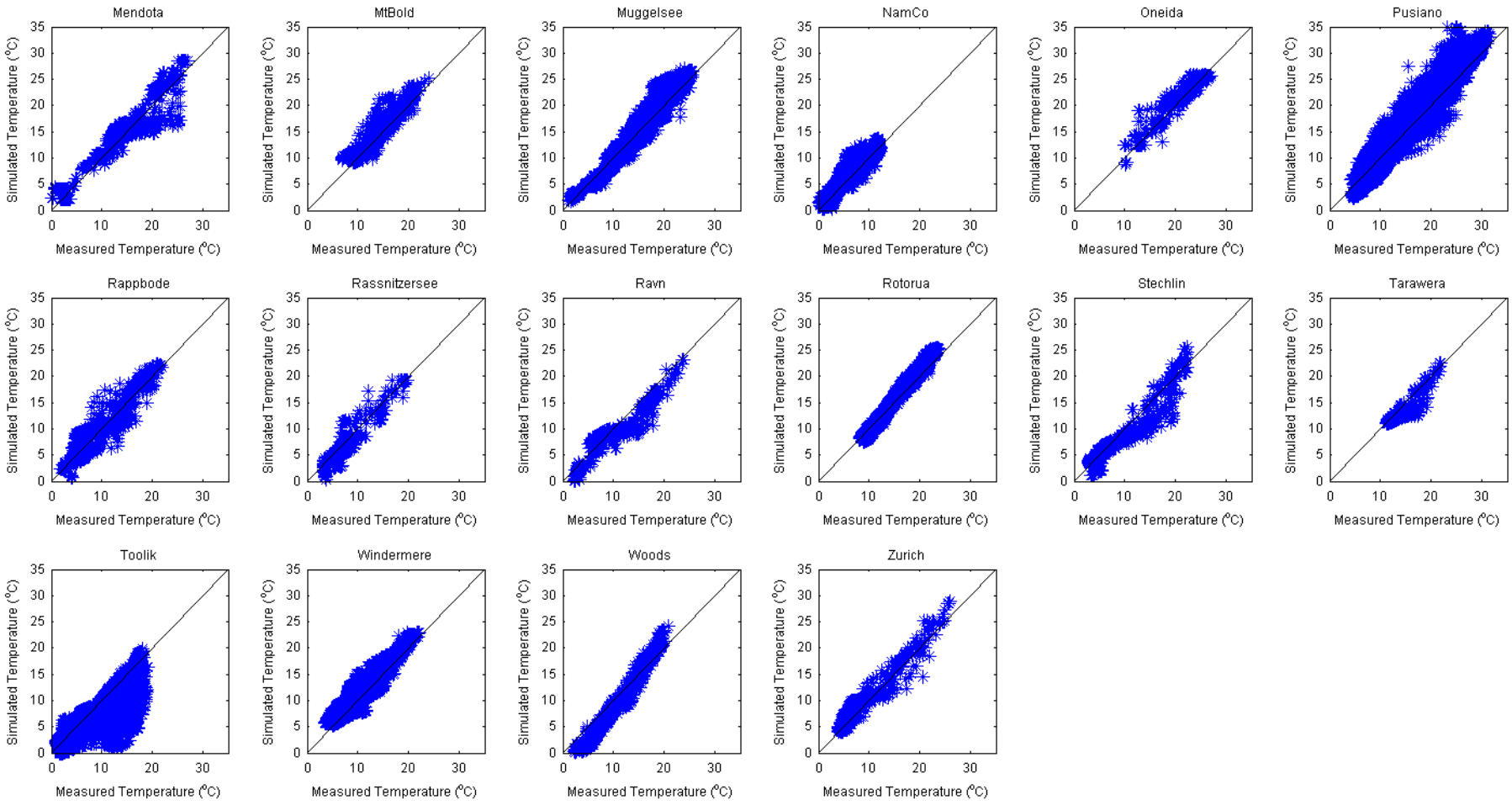


Figure B1 – Plot of modelled vs observed temperature data for each of the MLCP lakes.

## 7 Appendix C – Sensitivity Analysis

Table C1 - Significance (p) and correlation (r) between sensitivity indices for full profile temperature and mean values of lake volume (V), surface area (Area), depth (D), surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature (T<sub>air</sub>), wind speed (u<sub>wind</sub>), light extinction coefficient (K<sub>w</sub>), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red and corresponding *r* in yellow.

Attribute	C <sub>c</sub>	C <sub>w</sub>	C <sub>s</sub>	C <sub>t</sub>	C <sub>KH</sub>	C <sub>hyp</sub>	C <sub>e</sub>	C <sub>h</sub>	C <sub>d</sub>	
r	V	0.08	0.04	-0.09	-0.10	0.47	0.04	-0.07	-0.20	0.02
	Area	0.15	0.09	0.02	0.09	0.33	0.16	-0.43	-0.18	-0.06
	D	0.06	0.03	-0.12	-0.13	0.49	0.02	0.06	-0.16	0.03
	A/D	-0.02	-0.03	-0.16	-0.22	0.43	-0.08	0.30	-0.13	0.05
	L/W	-0.08	0.14	-0.10	0.15	-0.05	0.21	-0.15	-0.08	-0.40
	Inf	0.01	0.06	0.00	-0.24	0.43	0.07	-0.08	-0.41	-0.32
	RT	0.21	0.04	-0.01	0.18	0.18	0.06	-0.24	0.09	0.26
	sw	-0.10	-0.11	-0.18	-0.14	0.39	-0.12	0.41	0.12	0.18
	T <sub>air</sub>	-0.23	-0.24	-0.26	-0.22	-0.40	-0.30	0.34	-0.67	0.03
	u <sub>wind</sub>	-0.30	-0.32	-0.19	-0.15	-0.14	-0.30	0.33	0.09	0.22
	K <sub>w</sub>	-0.25	-0.16	-0.04	-0.16	-0.30	-0.14	0.42	0.12	0.11
Lat	0.22	0.16	0.24	0.18	-0.24	0.20	-0.49	0.19	-0.06	
p	V	0.68	0.85	0.64	0.60	0.01	0.84	0.73	0.30	0.91
	Area	0.42	0.65	0.90	0.65	0.08	0.40	0.02	0.34	0.75
	D	0.75	0.89	0.54	0.48	0.01	0.92	0.77	0.39	0.90
	A/D	0.90	0.87	0.39	0.24	0.02	0.67	0.11	0.49	0.79
	L/W	0.68	0.47	0.60	0.44	0.80	0.27	0.44	0.66	0.03
	Inf	0.95	0.74	0.98	0.23	0.02	0.72	0.68	0.03	0.10
	RT	0.29	0.83	0.97	0.36	0.37	0.76	0.21	0.67	0.18
	sw	0.61	0.56	0.35	0.47	0.04	0.52	0.02	0.54	0.35
	T <sub>air</sub>	0.22	0.21	0.17	0.25	0.03	0.11	0.07	0.00	0.88
	u <sub>wind</sub>	0.11	0.08	0.31	0.43	0.48	0.10	0.08	0.65	0.25
	K <sub>w</sub>	0.18	0.39	0.84	0.39	0.11	0.45	0.02	0.53	0.57

3741	Lat	0.25	0.41	0.20	0.35	0.20	0.28	0.01	0.31	0.76
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3781 Table C2 - - Significance (p) and correlation (r) between sensitivity indices for epilimnion temperature and mean values of lake volume (V), surface area  
 3782 (Area), depth (D), surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air  
 3783 temperature (T<sub>air</sub>), wind speed (u<sub>wind</sub>), light extinction coefficient (K<sub>w</sub>), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red  
 3784 and corresponding r in yellow.

3785		Attribute	C <sub>c</sub>	C <sub>w</sub>	C <sub>s</sub>	C <sub>t</sub>	C <sub>KH</sub>	C <sub>hyp</sub>	C <sub>e</sub>	C <sub>h</sub>	C <sub>d</sub>
3786	r	V	0.24	0.19	0.31	0.25	0.56	-0.10	-0.20	0.21	0.31
3787		Area	0.40	0.29	0.27	0.30	0.63	0.09	-0.25	0.21	0.33
3788		D	0.18	0.14	0.31	0.22	0.49	-0.13	-0.17	0.20	0.28
3789		A/D	0.00	0.01	0.24	0.12	0.30	-0.21	-0.05	0.13	0.17
3790		L/W	-0.07	0.10	-0.13	0.20	0.25	0.41	-0.18	0.14	-0.22
3791		Inf	0.13	0.09	0.07	0.14	0.51	-0.05	-0.31	-0.16	-0.09
3792		RT	0.44	0.29	0.40	0.09	0.22	-0.06	0.09	0.21	0.60
3793		sw	-0.13	-0.13	0.27	0.10	0.03	-0.20	0.09	0.19	0.08
3794		T <sub>air</sub>	-0.21	-0.50	-0.54	-0.40	-0.10	-0.34	0.43	-0.46	-0.61
3795		u <sub>wind</sub>	-0.18	0.08	-0.11	-0.05	-0.26	-0.10	0.35	0.08	-0.02
3796		K <sub>w</sub>	-0.43	-0.37	-0.33	-0.37	-0.58	-0.14	0.32	-0.31	-0.36
3797	Lat	0.26	0.23	-0.06	0.03	-0.05	0.26	-0.16	-0.15	0.12	
3798	p	V	0.19	0.31	0.10	0.19	0.00	0.62	0.30	0.27	0.10
3799		Area	0.03	0.12	0.15	0.10	0.00	0.63	0.18	0.26	0.08
3800		D	0.35	0.45	0.10	0.24	0.01	0.50	0.38	0.29	0.13
3801		A/D	0.99	0.94	0.20	0.51	0.11	0.27	0.79	0.50	0.37
3802		L/W	0.71	0.60	0.49	0.28	0.18	0.02	0.34	0.45	0.24
3803		Inf	0.50	0.64	0.73	0.48	0.01	0.80	0.11	0.40	0.66
3804		RT	0.02	0.14	0.03	0.67	0.25	0.75	0.66	0.28	0.00
3805		sw	0.50	0.49	0.14	0.62	0.86	0.30	0.65	0.31	0.67
3806		T <sub>air</sub>	0.27	0.00	0.00	0.03	0.58	0.06	0.02	0.01	0.00
3807		u <sub>wind</sub>	0.33	0.68	0.56	0.78	0.17	0.61	0.06	0.68	0.94
3808		K <sub>w</sub>	0.02	0.04	0.07	0.05	0.00	0.45	0.09	0.10	0.05
3809	Lat	0.17	0.22	0.77	0.88	0.80	0.16	0.39	0.43	0.51	

3821 Table C3 -- Significance (p) and correlation (r) between sensitivity indices for hypolimnion temperature and mean values of lake volume (V), surface area  
3822 (Area), depth (D), surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air  
3823 temperature (T<sub>air</sub>), wind speed (u<sub>wind</sub>), light extinction coefficient (K<sub>w</sub>), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red  
3824 and corresponding r in yellow.

3825	Attribute	C <sub>c</sub>	C <sub>w</sub>	C <sub>s</sub>	C <sub>t</sub>	C <sub>KH</sub>	C <sub>hyp</sub>	C <sub>e</sub>	C <sub>h</sub>	C <sub>d</sub>	
3826	r	V	-0.11	-0.37	-0.33	-0.26	-0.07	-0.22	0.08	-0.27	-0.26
3827		Area	-0.03	-0.34	-0.29	-0.21	-0.14	-0.20	-0.23	-0.22	-0.32
3828		D	-0.11	-0.34	-0.31	-0.24	-0.02	-0.18	0.16	-0.25	-0.22
3829		A/D	-0.14	-0.29	-0.26	-0.19	0.02	-0.14	0.31	-0.22	-0.12
3830		L/W	-0.10	0.19	0.20	0.17	0.14	0.22	-0.39	-0.15	-0.09
3831		Inf	-0.39	-0.39	-0.20	-0.25	-0.08	-0.19	-0.02	-0.68	-0.38
3832		RT	0.18	-0.29	-0.39	-0.13	-0.34	-0.36	-0.03	0.17	-0.24
3833		sw	-0.14	-0.22	0.01	-0.05	0.03	-0.08	0.59	0.11	-0.22
3834		T <sub>air</sub>	-0.22	-0.23	0.01	-0.11	-0.64	-0.22	0.32	-0.50	-0.20
3835		u <sub>wind</sub>	-0.29	-0.21	-0.35	-0.17	-0.21	-0.22	0.01	0.27	0.04
3836		K <sub>w</sub>	-0.14	0.15	0.27	0.17	0.13	0.25	0.16	0.07	0.27
3837		Lat	0.11	0.14	-0.03	0.09	0.15	0.00	-0.48	0.03	0.21
3838	p	V	0.55	0.04	0.07	0.16	0.71	0.25	0.69	0.15	0.17
3839		Area	0.88	0.07	0.12	0.27	0.46	0.28	0.22	0.24	0.09
3840		D	0.55	0.06	0.10	0.20	0.92	0.34	0.40	0.19	0.25
3841		A/D	0.46	0.12	0.17	0.30	0.93	0.46	0.10	0.24	0.52
3842		L/W	0.61	0.32	0.29	0.36	0.46	0.25	0.04	0.42	0.65
3843		Inf	0.04	0.04	0.30	0.21	0.68	0.34	0.93	0.00	0.05
3844		RT	0.37	0.14	0.04	0.50	0.08	0.06	0.88	0.40	0.22
3845		sw	0.47	0.24	0.96	0.78	0.87	0.66	0.00	0.56	0.25
3846		T <sub>air</sub>	0.25	0.22	0.94	0.57	0.00	0.25	0.09	0.01	0.30
3847		u <sub>wind</sub>	0.12	0.26	0.06	0.36	0.27	0.25	0.96	0.15	0.85
3848		K <sub>w</sub>	0.46	0.42	0.15	0.38	0.48	0.18	0.38	0.71	0.16
3849		Lat	0.55	0.46	0.87	0.63	0.44	0.99	0.01	0.86	0.28



3861 Table C4 - - Significance (p) and correlation (r) between sensitivity indices for thermocline depth and mean values of lake volume (V), surface area (Area),  
 3862 depth (D), surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature  
 3863 ( $T_{air}$ ), wind speed ( $u_{wind}$ ), light extinction coefficient ( $K_w$ ), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red and  
 3864 corresponding  $r$  in yellow.

3865	Attribute	$C_c$	$C_w$	$C_s$	$C_t$	$C_{KH}$	$C_{hyp}$	$C_e$	$C_h$	$C_d$	
3866	r	V	0.48	0.43	0.30	0.35	0.58	0.29	0.23	0.31	0.25
3867		Area	0.57	0.54	0.29	0.42	0.61	0.27	0.40	0.36	0.17
3868		D	0.41	0.36	0.27	0.31	0.51	0.28	0.14	0.27	0.25
3870		A/D	0.22	0.19	0.22	0.18	0.33	0.20	0.02	0.15	0.25
3871		L/W	0.01	0.16	-0.01	0.04	0.04	0.06	0.06	0.15	-0.04
3872		Inf	0.50	0.43	0.34	0.38	0.53	0.33	0.23	0.40	0.23
3873		RT	0.25	0.13	-0.06	0.10	0.28	0.08	0.16	0.07	0.10
3874		sw	0.03	0.08	0.31	-0.06	0.10	0.17	0.05	0.03	0.38
3875		$T_{air}$	0.11	0.13	0.19	0.06	0.09	0.34	0.19	0.18	0.44
3876		$u_{wind}$	-0.19	-0.13	0.06	-0.18	-0.22	-0.19	0.00	-0.33	0.06
3877		$K_w$	-0.42	-0.24	-0.27	-0.26	-0.46	0.03	-0.23	-0.04	0.08
3878	Lat	-0.01	-0.11	-0.34	0.01	-0.09	-0.19	-0.06	-0.05	-0.32	
3879	p	V	0.01	0.02	0.11	0.06	0.00	0.12	0.21	0.10	0.17
3880		Area	0.00	0.00	0.12	0.02	0.00	0.15	0.03	0.05	0.38
3881		D	0.03	0.05	0.15	0.10	0.00	0.14	0.45	0.15	0.18
3882		A/D	0.24	0.31	0.24	0.35	0.08	0.29	0.93	0.42	0.18
3883		L/W	0.94	0.39	0.94	0.83	0.83	0.74	0.77	0.43	0.84
3884		Inf	0.01	0.02	0.07	0.04	0.00	0.09	0.25	0.04	0.24
3885		RT	0.21	0.50	0.77	0.61	0.15	0.70	0.40	0.71	0.60
3886		sw	0.86	0.68	0.10	0.75	0.60	0.37	0.79	0.88	0.04
3887		$T_{air}$	0.55	0.49	0.31	0.75	0.63	0.07	0.30	0.34	0.02
3888		$u_{wind}$	0.31	0.49	0.74	0.34	0.25	0.31	1.00	0.07	0.76
3889	$K_w$	0.02	0.20	0.14	0.16	0.01	0.89	0.22	0.83	0.68	
3890	Lat	0.95	0.58	0.07	0.96	0.64	0.30	0.77	0.81	0.08	

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3901 Table C5 - - Significance (p) and correlation (r) between sensitivity indices for Schmidt stability and mean values of lake volume (V), surface area (Area),  
 3902 depth (D), surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature  
 3903 ( $T_{air}$ ), wind speed ( $u_{wind}$ ), light extinction coefficient ( $K_w$ ), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red and  
 3904 corresponding  $r$  in yellow.

3905	Attribute	$C_c$	$C_w$	$C_s$	$C_t$	$C_{KH}$	$C_{hyp}$	$C_e$	$C_h$	$C_d$	
3906	r	V	0.02	0.13	0.27	0.15	-0.08	-0.06	-0.16	-0.07	0.00
3907		Area	0.00	0.07	0.10	-0.06	-0.30	-0.28	-0.25	-0.15	-0.31
3908		D	0.04	0.14	0.31	0.22	0.02	0.04	-0.09	-0.01	0.13
3909		A/D	0.05	0.12	0.33	0.30	0.15	0.17	0.07	0.03	0.30
3910		L/W	-0.15	-0.16	-0.21	-0.17	-0.06	-0.08	0.03	-0.12	-0.21
3911		Inf	0.02	0.17	0.31	0.16	0.13	0.07	-0.16	0.07	0.13
3912		RT	0.01	0.07	-0.05	0.00	-0.25	-0.14	-0.20	0.02	-0.13
3913		sw	-0.10	-0.14	0.04	0.16	-0.22	0.00	0.17	-0.04	0.00
3914		$T_{air}$	0.00	-0.05	-0.01	0.00	-0.08	0.03	-0.15	-0.29	0.00
3915		$u_{wind}$	0.04	-0.18	0.03	-0.21	0.00	-0.15	0.43	-0.05	0.14
3916		$K_w$	-0.11	-0.11	-0.25	0.15	0.31	0.32	0.32	0.24	0.26
3917	Lat	0.06	0.08	-0.07	-0.16	0.16	-0.05	-0.13	0.17	-0.05	
3918	p	V	0.93	0.48	0.14	0.43	0.67	0.77	0.41	0.73	0.98
3919		Area	0.99	0.70	0.62	0.74	0.10	0.13	0.19	0.44	0.10
3920		D	0.83	0.45	0.09	0.24	0.91	0.85	0.62	0.95	0.49
3921		A/D	0.80	0.53	0.08	0.11	0.44	0.37	0.70	0.87	0.11
3922		L/W	0.43	0.40	0.26	0.36	0.77	0.67	0.86	0.54	0.27
3923		Inf	0.92	0.40	0.11	0.43	0.52	0.74	0.42	0.72	0.52
3924		RT	0.95	0.72	0.78	0.98	0.20	0.48	0.31	0.94	0.50
3925		sw	0.60	0.45	0.84	0.40	0.24	1.00	0.38	0.85	0.98
3926		$T_{air}$	0.98	0.80	0.95	0.98	0.66	0.88	0.44	0.11	0.99
3927		$u_{wind}$	0.85	0.33	0.87	0.26	0.99	0.42	0.02	0.77	0.46
3928		$K_w$	0.57	0.55	0.19	0.41	0.09	0.08	0.08	0.20	0.17
3929	Lat	0.77	0.68	0.70	0.39	0.39	0.80	0.50	0.36	0.79	

## 8 Appendix D – Acknowledgements

Table D1 – Acknowledgements by individual lake. Note people includes additional staff and students who helped set up the GLM not included in the list of authors.

Lake Name	Morphometry	Meteorology	Flow	Field Data	Institutions	Funding	People
Alexandrina	Department of Environment, Water and Natural Resources	Natural Resources SA Murray-Darling Basin	Murray-Darling Basin Authority (MDBA)	Natural Resources SA Murray-Darling Basin	The University of Western Australia	ARC Discovery Grant DP130104078	Alex Perry
Ammersee	Bavarian Environment Agency	German Weather Service,Wielenbach; Bavarian Agency of Agriculture, Rothenfeld + Westerschondorf	Water Agency Weilheim	Water Agency Weilheim		Bavarian Environment Agency; Bavarian State Ministry of the Environment and Consumer Protection	Otfried Baume
Blelham	Ramsbottom, A.E. 1976. Depth charts of the Cumbrian Lakes. Freshwater Biological Association	Centre for Ecology and Hydrology	Environment Agency	Centre for Ecology and Hydrology	University of Reading; Centre for Ecology and Hydrology	UKLEON (NE/I007407/1)	Bernard Tebay
Bourget	Delebecque, 1898 and IFREMER 1992	Météo France	DREAL Rhône-Alpes	© SOERE OLA-IS, INRA Thonon-les-Bains, CISALB, [2012], developed by INRA's ORE Eco-Information system	INRA Thonon les Bains	CISALB, Agence de l'eau Rhône-Méditerranée-Corse, SOERE OLA, Ecole des Ponts ParisTech, INRA	Orlane Anneville
Cannonsville	New York City Department of Environmental Protection	Cannonsville Dam Station	Trout Creek and West Branch Delaware River	New York City Department of Environmental Protection, Kingston, NY	New York City Department of Environmental Protection	New York City Department of Environmental Protection	Karen E.B. Moore
Como	Lombardy Region	Water Research Institute - National Research Council of Italy (IRSA-CNR); Bergamo-Orio al Serio Aiport; Regional Authority for Environmental Protection (ARPA Lombardia)	Consorzio dell'Adda	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Centro Volta Como ; Istituto Nazionale della Montagna	Simulake Project; WAAESs Project	Gianni Tartari (Water Research Institute - National Research Council of Italy (IRSA-CNR)

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Constance	IGKB (Internationale Gewässerschutzkommission für den Bodensee	German meteorological service (DWD)	IGKB (Internationale Gewässerschutzkommission für den Bodensee	IGKB (Internationale Gewässerschutzkommission für den Bodensee	LUBW Landesanstalt für Umwelt, Messungen und Naturschutz Baden-Württemberg; Limnological Institute, University of Konstanz	DFG, grant Ri 2040/1-1	Thomas Wolf
El Gergal	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	University of Granada	CGL2005-04070/HID	Carmelo Escot
Emaiksoun	Kenneth M. Hinkel, University of Cincinnati	Brittany L. Potter, University of Nebraska-Lincoln		Brittany L. Potter, University of Nebraska-Lincoln; Kenneth M. Hinkel, University of Cincinnati	LimnoTech; University of Nebraska-Lincoln; University of Cincinnati	NSF ARC-1107792	Brittany L. Potter, Kenneth M. Hinkel
Esthwaite	Ramsbottom, A.E. 1976. Depth charts of the Cumbrian Lakes. Freshwater Biological Association	Centre for Ecology and Hydrology	Environment Agency	Centre for Ecology and Hydrology	University of Reading; Centre for Ecology and Hydrology	UKLEON (NE/I007407/1)	Bernard Tebay
Feeagh	Marine Institute	Met Eireann; Marine Institute	Marine Institute	Marine Institute	Marine Institute	Marine Institute	Eleanor Jennings; Elizabeth Ryder; Mary Dillane; Russell Poole; Burrishoole field staff
Geneva03	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL			Orlane Anneville
Geneva05	As Above	As Above	As Above	As Above	As Above	As Above	As Above
GrosseDhuen	Wupperverband (reservoir manager)	German Weather Service stations Luedenscheid, Cologne-Bonn	Wupperverband (reservoir manager)	Wupperverband (reservoir manager), Helmholtz Centre for Environmental Research UFZ	Helmholtz Centre for Environmental Research UFZ	District Council Cologne and Ministry of Environment North Rhine-Westphalia	Karsten Rahn, Martin Wieprecht, Wilfried Scharf
Harp		Dorset Environmental Science Centre; Environment Canada	Dorset Environmental Science Centre				

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Iseo	Regione Lombardia	Università degli Studi di Brescia; Bergamo-Orio al Serio Airport; Regional Authority for Environmental Protection (ARPA Lombardia)	Consorzio dell'Oglio	Università degli Studi di Brescia; Prof. Letizia Garibaldi (Università degli Studi di Milano-Bicocca)			
Kinneret03	Kinneret Limnological Laboratory	Isr. Meteorol. Service; Kinneret Limnological Laboratory	Isr. Hydrological Service	Kinneret Limnological Laboratory, IOLR	Israel Water Authority	Israel Water Authority	
Kinneret97	As Above	As Above	As Above	As Above	As Above	As Above	As Above
Mendota	Wisconsin Department of Natural Resources	SSEC-GAMIS-RIG UW-Madison; US National Climatic Data Center	U.S. Geological Survey	University of Wisconsin - Madison LTER program	University of Wisconsin, Madison	NSF grant DEB-0822700 (North Temperate Lakes Long-Term Ecological Research)	
MtBold	South Australian Water Corporation	Bureau of Meteorology; South Australian Water Corporation, Happy Valley Reservoir	Department of Water and Natural Resources; SA Water Major Systems	South Australian Water Corporation	University of Adelaide, Adelaide, South Australia	Water Research Foundation, Boulder, CO, USA; South Australian Water Corporation	Mike Burch; Rob Daly;
Muggelsee	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Senatsverwaltung für Stadtentwicklung und Umwelt Berlin	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Thomas Hintze (IGB) for operating the Müggelsee Lake Station.
NamCo	Institute of Tibetan Plateau Research, Chinese Academy of Sciences	Institute of Tibetan Plateau Research, Chinese Academy of Sciences			Institute of Tibetan Plateau Research, Chinese Academy of Sciences; Institut für Geographie, Friedrich-Schiller-Universität	the National Basic Research Program of China (2012CB956100), National Natural Science Foundation of China (41071123) and CADY project (TP2:03G0813F) from BMBF, Germany	
Oneida	National Oceanic and Atmospheric Administration; Cornell Biological Field Station	Northeast Regional Climate Center	United States Geological Survey	Cornell Biological Field Station	Cornell University	Cornell University Brown Endowment; New York State Department of Environmental Conservation; United States Department of Agriculture 0226747	Lars Rudstam

4061	Pusiano	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Regional Authority for Environmental Protection (ARPA Lombardia); National Oceanic and Atmospheric Administration	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Parco Valle Lambro; Fondazione CARIPOLO	Progetto PIROGA	Gianni Tartari and Franco Salerno (Water Research Institute - National Research Council of Italy (IRSA-CNR))
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4066	Rappbode	Reservoir authority of the state of Saxony-Anhalt (Talsperrenbetrieb Sachsen-Anhalt)	German Meteorological Service (DWD)	Rappbode Reservoir Authority (Talsperrenbetrieb Sachsen-Anhalt)	Fernwasserversorgung Elbaue Ostharz; Helmholtz Centre for Environmental Research - UFZ; Talsperrenbetrieb Sachsen-Anhalt	Helmholtz Centre for Environmental Research - UFZ		Karsten Rahn (UFZ, SEEFO); Martin Wieprecht (UFZ, SEEFO); Maren Dietze (Talsperrenbetrieb Sachsen-Anhalt); Dieter Noga (DWD); Marco Matthes (Fernwasserversorgung Elbaue Ostharz).
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4072	Rassnitzersee	Helmholtz Centre for Environmental Research - UFZ	German Meteorological Service (DWD)	Lausitzer and Mitteldeutsche Braunkohle Verwaltungsgesellschaft - LMBV	Uwe Kiwel(UFZ) and Karsten Rahn (UFZ)	Helmholtz Centre for Environmental Research - UFZ	Helmholtz Centre for Environmental Research - UFZ	Uwe Kiwel(UFZ) and Karsten Rahn (UFZ)
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4077	Ravn	National Monitoring Program for Water and Nature. Data hosted by Aarhus University.	Danish Meteorological Institute (DMI). Data made available for analyses relating to the National Monitoring Program for Water and Nature.	National Monitoring Program for Water and Nature. Data hosted by Aarhus University.	National Monitoring Program for Water and Nature. Data hosted by Aarhus University.	Aarhus University	CLEAR centre of excellence (Villum-Kann Rasmussen Foundation)	
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4082	Rotorua	Digitised bathymetry held by University of Waikato, Bay of Plenty Regional Council	Meteorological Service of New Zealand Limited. Data obtained via 'cliflo' database (National Institute of Water and Atmospheric Research (NIWA), New Zealand).	National Institute of Water and Atmospheric Research.	Bay of Plenty Regional Council	University of Waikato-Environmental Research Institute	Bay of Plenty Regional Council Chair in Lake Restoration at University of Waikato	
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4087	Stechlin		German Meteorological Service (DWD), Umwelt Bundesamt (UBA), Energiewerke Nord GmbH (Betriebsteil Kernkraftwerk Rheinsberg)		Leibniz-Institute of Freshwater Ecology and Inland Fisheries	Leibniz-Institute of Freshwater Ecology and Inland Fisheries	German Federal Ministry of Research and Education (BMBF Project KLIMZUG-INKABB TP22)	Peter Kasprzak
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Tarawera	Digitised bathymetry held by University of Waikato, Bay of Plenty Regional Council	National Institute of Water and Atmospheric Research	Bay of Plenty Regional Council	Bay of Plenty Regional Council	University of Waikato-Environmental Research Institute	Bay of Plenty Regional Council Chair in Lake Restoration at University of Waikato	Andy Bruere
Toolik	Toolik Field Station - Environmental Data Center (TFS EDC). Jason J. Stuckey (TFS GIS Manager) performed a bathymetrical study of the lake in 2008 using a Garmin GPSMAP 188 Sounder	Meteorological datasets provided by the Toolik Field Station Environmental Data Center are based upon work supported by the U. S. National Science Foundation (NSF) under grants #455541 and #1048361	Arctic Long Term Ecological Research (ARC LTER), funded by NSF, Division of Environmental Biology (DEB) 0423385	(a) Arctic Long Term Ecological Research (ARC LTER); (b) Toolik Field Station - Environmental Data Center (TFS EDC), University of Alaska Fairbanks (UAF)	(a) Department of Ecology, Evolution, and Marine Biology, University of California, Santa Barbara, California; (b) Marine Science Institute, University of California, Santa Barbara, California	NFS Support, from two different grants of Arctic Natural Sciences (ANS) to Sally MacIntyre: #0714085 and #1204267	
Windermere	Ramsbottom, A.E. 1976. Depth charts of the Cumbrian Lakes. Freshwater Biological Association	Centre for Ecology and Hydrology	Environment Agency	Centre for Ecology and Hydrology	University of Reading; Centre for Ecology and Hydrology	UKLEON (NE/I007407/1)	Bernard Tebay
Woods	Hydro Tasmania	Australian Bureau of Meteorology	Hydro Tasmania	Hydro Tasmania	University of Tasmania	ARC Linkage Grant LP130100756	Carolyn Maxwell, Leon Barmuta, Abhijeet Kulkarni, Aditya Singh
Zurich	David M. Livingstone	MeteoSwiss	Federal Office for the Environment (FOEN)	Wasserversorgung der Stadt Zürich (Zurich Water Supply)	Eawag: Swiss Federal Institute of Aquatic Science and Technology	Amt für Abfall, Wasser, Energie und Luft (AWEL) of the Canton of Zurich (Lake Monitoring)	