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THE COMMUNITY LEVERAGED UNIFIED ENSEMBLE (CLUE) IN THE 2016 NOAA/HAZARDOUS WEATHER TESTBED SPRING FORECASTING EXPERIMENT

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THE COMMUNITY LEVERAGED UNIFIED ENSEMBLE (CLUE) IN THE 2016 NOAA/HAZARDOUS WEATHER TESTBED SPRING FORECASTING EXPERIMENT

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The CLUE system represents an unprecedented effort to leverage several academic and government research institutions to help guide NOAA's operational environmental modeling efforts at the convection-allowing scale.

The National Severe Storms Laboratory (NSSL) and Storm Prediction Center (SPC) coorganize annual Spring Forecasting Experiments (SFEs), which are conducted in NOAA's Hazardous Weather Testbed (HWT) at the National Weather Center in Norman, Oklahoma, for five weeks during the climatological peak of the severe weather season. The SFEs are designed to test emerging concepts and technologies for improving the prediction of hazardous convective weather with the primary goals of accelerating the transfer of promising new tools and concepts from research to operations, inspiring new initiatives for operationally relevant research, and identifying and documenting sensitivities and performance characteristics of state-of-the-art experimental convection-allowing

modeling (CAM) systems. Over the last decade, the SFEs have emerged as an international resource for developing and evaluating the performance of new CAM systems, and major advances have been made in creating, importing, processing, verifying, and extracting unique hazardous weather fields while providing analysis and visualization tools including probabilistic information, for these large and complex datasets. For example, during the 2010 experiment (Clark et al. 2012), in addition to providing a 26-member, 4-km grid-spacing CAM-based ensemble, the Center for Analysis and Prediction of Storms (CAPS) at the University of Oklahoma provided a 1-km contiguous U.S. (CONUS) domain forecast that required over 10,000 computing cores. In the 2015 SFE (Gallo et al. 2017), six unique and

independently designed CAM-based ensembles were contributed by CAPS, the National Center for Atmospheric Research (NCAR), NSSL, SPC, and the Air Force Weather Agency (AFWA; now called the 557th Weather Wing). Figure 1 provides a summary of CAMs examined since 2007, along with a timeline of CAM guidance milestones.

Through the SFEs, much has been learned about how to utilize and configure CAMs and CAM ensembles, and since 2007 the number of CAM systems (including ensembles) examined in the HWT has increased dramatically. Meanwhile, new technologies and physical understanding have been migrated to the SPC, enhancing the timeliness and accuracy of their severe weather forecasts. Despite these advances, progress toward identifying optimal CAM ensemble configurations has been inhibited because HWT collaborators have independently designed contributed CAM systems, which makes it difficult to attribute differences in performance characteristics. For example, during the 2015 SFE, CAPS and NSSL contributed mixed- and single-physics ensembles, respectively, but because of other differences in the configurations (e.g., initial condition perturbations, data assimilation, grid spacing, domain size, and model version), the impacts of single versus mixed-physics configurations could not be isolated. Thus, after the 2015 SFE it was clear to SFE leaders that more controlled experiments were needed. Furthermore, around the same time period, the international University Corporation for Atmospheric Research Community Advisory Committee for the National Centers for Environmental Prediction (UCACN) Model Advisory Committee, which is charged with developing recommendations for a unified NOAA modeling strategy to advance the United States to world leadership in numerical modeling, released a

comprehensive set of recommendations¹ that included the following: 1) the NOAA environmental modeling community requires a rational, evidence-driven approach toward decision-making and modeling system development; 2) a unified collaborative strategy for model development across NOAA is needed; and 3) NOAA needs to better leverage the capabilities of the external community. Thus, in the spirit of these recommendations, organizers of the HWT SFEs made a major push to coordinate efforts among its large group of collaborators in 2016. Specifically, instead of each group providing a separate, independently designed CAM-based ensemble, all groups agreed on a set of model specifications so that the simulations contributed by each group could be viewed as one large, carefully designed “superensemble.” This design facilitated a number of controlled experiments geared toward finding optimal configuration strategies for CAM ensembles and has been termed the Community Leveraged Unified Ensemble (CLUE, hereafter). The superensemble concept has been used in previous works for tropical cyclone, weather, climate, and seasonal prediction systems (e.g., Krishnamurti et al. 1999; Palmer et al. 2004; Krishnamurti et al. 2016 and references therein), but has yet to be applied within a CAM ensemble framework. However, the philosophy behind the CLUE design is different from these previous works on superensembles. Specifically, the CLUE has a more coordinated design with goals that are more focused on identifying impacts of different ensemble design strategies, rather than generating a single “best” forecast from independent ensemble datasets.

¹ The full report is available at www.ncep.noaa.gov/director/ucar_reports/ucacn_20151207/UMAC_Final_Report_20151207-v14.pdf.

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The CLUE system represents an unprecedented effort to leverage several academic and government research institutions to help guide NOAA's operational environmental modeling efforts. In future SFEs, the CLUE will be reconfigured based on results from previous years, advances in technology, and feedback from the operational and research communities. Furthermore, the CLUE framework

will help test initial convection-allowing versions of the Finite Volume Cubed Sphere Model (FV3; Putman and Lin 2007) developed at NOAA's Geophysical Fluid Dynamics Laboratory. The FV3 has been selected as the dynamic core to replace the Global Forecast System (GFS) model as part of the Next Generation Global Prediction System (NGGPS; www.weather.gov/sti/stimodeling_nggps) program

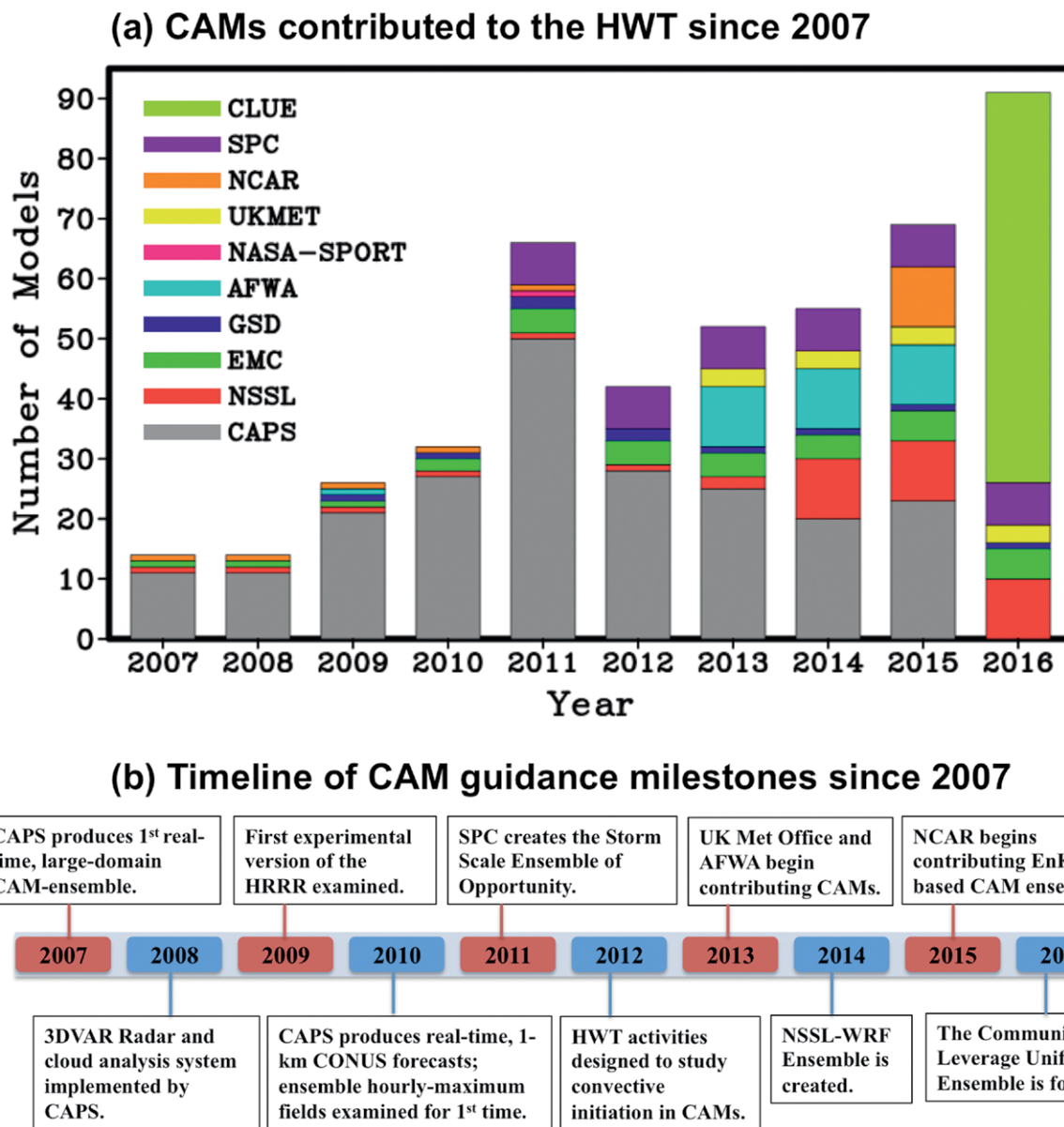


FIG. 1. (a) Stacked bar graph indicating the number of unique CAMs used each year since 2007 in the HWT SFEs. The different colors denote the number of models contributed by the different agencies. A legend is provided at the top left. Abbreviations are defined as follows: CLUE, Community Leveraged Unified Ensemble; SPC, Storm Prediction Center; NCAR, National Center for Atmospheric Research; UKMET, Met Office; NASA SPORT, National Aeronautics and Space Administration Short-term Prediction Research and Transition Center; AFWA, Air Force Weather Agency; GSD, Global Systems Division of NOAA's Earth System Research Laboratory; EMC, NOAA's Environmental Modeling Center; NSSL, National Severe Storms Laboratory; and CAPS, Center for Analysis and Prediction of Storms. (b) Timeline of CAM guidance milestones at the HWT since 2007.

and FV3 is envisioned as the eventual foundation for NOAA's regional models and ensemble systems. This will require much research, development, and testing to ensure that FV3 performs equal to or better than existing regional short-term forecasting systems.

This article describes the design of the 2016 CLUE system and the eight specific experiments that were conducted within the CLUE framework. Additionally, as an example of the research enabled by the CLUE framework, results are presented from one of the experiments that examined the impact of using single versus multicore CAM ensemble configurations.

CLUE CONFIGURATION. The idea for the CLUE system was formulated in fall of 2015. At this time, plans were already in place for several groups of collaborators to contribute model data to the 2016 SFE through NOAA-funded research-to-operations projects. For example, NCAR and CAPS had projects funded by NOAA's Oceanic and Atmospheric Research (OAR) Office of Water and Air Quality (OWAQ), and the University of North Dakota (UND) had a project funded by the National Weather Service Research to Operations Initiative. While the model runs NSSL contributed to the CLUE were not supported by a specific grant, the Texas Advanced Computing Center (TACC) provided generous computing resources for their contribution. Since participation in the CLUE would require work beyond that outlined in their already-existing projects, leaders from each group of collaborators were approached individually to gauge whether they had the resources and willingness to participate. Fortunately, because of the mutually beneficial research that the CLUE system would enable, along with the potential to provide evidence to help optimize NOAA's first operational CAM-ensemble configuration, all collaborators were eager and willing to participate.

The CLUE configuration was formulated by considering some basic research questions, such as how to optimize CAM-ensemble configurations and how to build around each collaborator's already existing plans for model data contributions. Ultimately, the CLUE was designed to have 66 members: 35 contributed by CAPS, 15 by NSSL, 10 by NCAR, 5 by the UND, and 1 from the Earth System Research Laboratory/Global Systems Division (ESRL/GSD). The runs were conducted on several different high-performance computing systems. CAPS used TACC's Stampede system and the University of Tennessee's National Institute for Computation Science's (NICS) Darter system, NSSL used TACC's Lonestar5 system, NCAR used the Yellowstone

supercomputer, UND used TACC's Stampede system, and ESRL/GSD used NOAA's Jet system.

All members were initialized at 0000 UTC on weekdays with forecasts to 36 h using 3-km grid spacing over a CONUS domain. Members included the Advanced Research version of the Weather Research and Forecasting (WRF-ARW) Model (Skamarock et al. 2008), as well as the Nonhydrostatic Multiscale Model on the B grid (NMMB; Janjić and Gall 2012). The CAPS, UND, and NSSL members all shared a set of common model versions, domain specifications (including vertical levels), physics parameterizations, and postprocessing methods. The ESRL/GSD member was a developmental version of the High Resolution Rapid Refresh (HRRR) model (Benjamin et al. 2016) run to 36 h, which had a slightly different domain than the other members. The NCAR members also had a slightly different domain and used a 1-yr-older version of WRF, which was necessary because their members were from an already established ensemble system whose configuration was based on extensive testing and verification (Schwartz et al. 2015a). The NCAR group did not want to risk introducing changes to their system by adhering exactly to the CLUE specifications, since it could introduce unwanted systematic biases. Despite some minor differences in the NCAR and ESRL/GSD members, postprocessing was standardized across all ensemble subsets (described later).

The basic strategy in designing the CLUE was to formulate several subsets of up to 20 members that could be used to test specific configuration strategies in controlled experiments. Ten unique subsets were formulated, with CAPS contributing five subsets, NSSL two, and ESRL/GSD, NCAR, and UND each contributing one. Some experiments utilized combinations of these subsets. These subsets are described as follows:

- 1) core (CAPS)—Nine WRF-ARW members were designed to account for as many error sources as possible. The control member used initial conditions (ICs) and lateral boundary conditions (LBCs; 3-h updates) from 12-km grid-spacing North American Mesoscale Forecast System (NAM) analyses and forecasts, respectively. Radar reflectivity and velocity data and other traditional data, including surface observations and rawinsondes, were assimilated into the ICs using the Advanced Regional Prediction System (ARPS) three-dimensional variational data assimilation (3DVAR; Xue et al. 2003; Gao et al. 2004) and cloud analysis (Xue et al. 2003; Hu et al. 2006) system. The other core subset members also used ARPS-3DVAR, but IC perturbations

were derived from evolved (through 3 h) perturbations of 2100 UTC initialized members of the National Centers for Environmental Prediction (NCEP) Short-Range Ensemble Forecast (SREF) system (Du et al. 2006) and added to the control member ICs, with corresponding SREF forecasts used for LBCs. Mixed physics were implemented in the core subset using various combinations of microphysics and planetary boundary layer (PBL)/turbulence schemes.

- 2) s-phys-rad (CAPS)—Ten WRF-ARW members (including the control member of core) were configured the same as core but used a single set of physics.
- 3) caps-enkf (CAPS)—Ten WRF-ARW members used the same set of physics and LBCs as core, but with ICs that were derived from an ensemble Kalman filter (EnKF) system.
- 4) caps-nmmb-rad (CAPS)—A single NMMB run used the same ICs/LBCs as the core control member.
- 5) caps-nmmb (CAPS)—Five NMMB members had the same ICs/LBCs as five of the s-phys-rad members but did not use ARPS-3DVAR (i.e., a “cold start” was used).
- 6) s-phys-norad (NSSL)—Ten WRF members were the same as s-phys-rad, but without ARPS-3DVAR (i.e., cold start).
- 7) nssl-nmmb (NSSL)—Five NMMB members were configured the same as the caps-nmmb members,

except they shared a different set of the s-phys-rad ICs/LBCs.

- 8) HRRR36 (ESRL/GSD)—A development version of the HRRR was configured to provide 36-h forecasts. The HRRR is a 3-km grid-spacing, ARW-based model that is initialized hourly and provides 18-h forecasts.
- 9) ncar-enkf (NCAR)—Ten WRF members used single physics and ICs/LBCs derived from NCAR’s Data Assimilation Research Testbed (DART; Anderson et al. 2009) software (Schwartz et al. 2015a).
- 10) mp (UND)—Five WRF members had the same ICs/LBCs as the core control member, but with different microphysics parameterizations in each member.

Table 1 provides a summary of the specifications for each CLUE subset, and further details including specifications for every member can be found in the online supplement (<https://doi.org/10.1175/BAMS-D-16-0309.2>).

CLUE EXPERIMENTS. The design of CLUE allowed for eight unique experiments, which are described as follows:

- 1) ARW versus NMMB—A direct comparison of the subjective and objective skill of ARW and NMMB dynamic cores was conducted. These

TABLE 1. Summary of CLUE subsets. IC/LBC perturbations labeled SREF indicate that IC perturbations were extracted from members of NCEP’s SREF system and added to 0000 UTC NAM analyses. In subsets with “yes” indicated for mixed physics, the microphysics and turbulence parameterizations were varied, except for subset mp, which only varied the microphysics. Note that the control member of the core ensemble was also used as the control member in the mp and s-phys-rad ensembles. Thus, although the total number of members adds to 67, there were 66 unique members. Further, one member planned for the core subset was not ready for real-time implementation; thus, only nine core members were actually run. The HPC column provides the names of the high-performance computers used for each set of simulations. The agencies that maintain each system are given in the text.

CLUE subset	No. of members	IC/LBC perturbations	Mixed physics?	Data assimilation	Model core	Agency	HPC
core	10 (9)	SREF	Yes	ARPS-3DVAR	ARW	CAPS	Stampede
s-phys-rad	10	SREF	No	ARPS-3DVAR	ARW	CAPS	Stampede
caps-enkf	10	EnKF (CAPS)	Yes	EnKF (CAPS)	ARW	CAPS	Darter
caps-nmmb-rad	1	None	No	ARPS-3DVAR	NMMB	CAPS	Stampede
caps-nmmb	5	SREF	No	Cold start	NMMB	CAPS	Stampede
s-phys-norad	10	SREF	No	Cold start	ARW	NSSL	Lonestar5
nssl-nmmb	5	SREF	No	Cold start	NMMB	NSSL	Lonestar5
HRRR36	1	None	No	RAP-GSI/DFI	ARW	ESRL/GSD	Jet
ncar-enkf	10	EnKF (DART)	No	EnKF (DART)	ARW	NCAR	Yellowstone
mp	5	None	Yes	ARPS-3DVAR	ARW	UND	Stampede

direct comparisons were possible because 10 pairs of NMMB and ARW members within the caps-nmmb and nssl-nmmb, and s-phys-norad, subsets had different model cores but shared the same ICs/LBCs. The optimal dynamic core for CAM applications is still an open question. NMMB is known to be more computationally efficient than ARW, but ARW has been preferred by severe weather forecasters and the severe weather research community because of its more realistic depiction of storm structure and evolution.

- 2) Multicore versus single-core ensemble design—Three ensembles were compared to test the effectiveness of single-core versus multicore configurations. The first ensemble used five ARW and five NMMB members from the s-phys-norad and nssl-nmmb subsets, respectively; the second used the 10 ARW members from the s-phys-norad subset; and the third used 10 NMMB members from the caps-nmmb and nssl-nmmb subsets. The effectiveness of multicore (or multimodel) ensemble configuration strategies has been demonstrated for seasonal [e.g., the North American Multimodel Ensemble (NMME); Kirtman et al. (2014)], medium-range [e.g., the North American Ensemble Forecast System (NAEFS); Candille (2009)], and short-range (e.g., the SREF; Du et al. 2006) forecasting applications. The use of multiple models with different but equally valid methods for initialization and integration helps to better sample the range of future states than a single modeling system. Although the Storm Scale Ensemble of Opportunity (SSEO; Jirak et al. 2012) has been shown to be a skillful multimodel CAM ensemble, the multimodel strategy has not been tested for CAM applications in controlled experiments. Furthermore, given the push toward model core unification that will better focus model development efforts (e.g., UCAR 2015), it is preferred that a future operational CAM ensemble will be single core. Thus, it is important to quantify how much skill (if any) is sacrificed from a single-core configuration within the context of a controlled experiment.
- 3) Single physics versus multiphysics—Two ensembles with the same set of perturbed ICs/LBCs were compared to test the impact of single versus multiphysics. One ensemble, core, used varied turbulence and microphysics schemes, while another, s-phys-rad, used a common set of physics. Although past SFEs have quantified the error growth from varied physics within a perfect analysis framework (i.e., nonperturbed ICs/LBCs; e.g., Clark et al. 2010b),

there has not been an experiment designed in the SFE to examine the impact of varied physics with perturbed ICs/LBCs in a CAM ensemble. Furthermore, while multiple physics schemes have been shown to increase spread, leading to improved forecast skill (e.g., Stensrud et al. 2000; Hacker et al. 2011; Berner et al. 2011, 2015), there are theoretical and practical disadvantages to multiphysics approaches, including the resource-intensive need to develop and maintain multiple parameterizations, as well as the introduction of systematic biases (e.g., Jankov et al. 2017). Thus, it is important to quantify the gain in skill (if any) from using multiphysics. Future SFEs will explore whether stochastic physics perturbations (e.g., Jankov et al. 2017 and references therein) in a single-physics ensemble can match or exceed the spread and skill from the multiphysics approach.

- 4) Comparison of ensembles with and without radar data assimilation—Two single-physics ensembles with perturbed ICs/LBCs were identically configured, except one, s-phys-rad, used ARPS-3DVAR to assimilate radar data and other observations in all members, while another, s-phys-norad, used a cold start in all members. Previous studies have documented the impact of radar data assimilation by comparing deterministic models with and without radar data assimilation (e.g., Kain et al. 2010; Stratman et al. 2013), finding that the positive impact of the assimilation is strongest within the first 3–6 h of the forecast but can last up to 12 h. However, these comparisons have not been conducted within an ensemble framework to determine the time length and magnitude of the positive impact of radar assimilation.
- 5) 3DVAR versus EnKF data assimilation strategies—The core, caps-enkf, and ncar-enkf subsets were compared. Although it is much more computationally expensive than 3DVAR, the EnKF data assimilation method is advantageous because it provides flow-dependent background error covariances that result in higher correlations between the model state and observed variables (e.g., Johnson et al. 2015). Despite the theoretical advantages to EnKF, subjective and objective comparisons of CAM ensembles from past SFEs did not find that those using EnKF performed any better than other data assimilation methods (Jirak et al. 2015). Thus, more work is needed to optimize EnKF for CAM ensemble applications. However, this experiment was not as controlled as the others, because aspects of the subset of configurations other than the data assimilation methods also differed.

- 6) GSD radar versus CAPS radar assimilation—Two methods for assimilating radar data were compared. One used ARPS-3DVAR and the other used the Digital Diabatic Filter Initialization (DDFI; Benjamin et al. 2016) system used in the HRRR. It was planned to include a core member configured the same as HRRR36, but using the ARPS-3DVAR system to generate the ICs. Because of time constraints, the core member planned for this experiment was not ready for implementation in real time. Thus, this experiment was not conducted.
- 7) Microphysics sensitivities—Using the five members of the mp subset, the impact of different microphysics parameterizations on forecast storm structure and evolution was examined. This experiment has been conducted in SFEs since 2010 (e.g., Clark et al. 2012, 2014), and through the participation of microphysics scheme developers each year, parameterizations have been improved and valuable interactions have occurred with forecasters and modelers.
- 8) Ensemble size experiment—A comparison of ensembles with equal contributions of NMMB and ARW members using 2, 4, 6, 10, and 20 members was conducted to examine the impact of ensemble size. The ensembles used combinations of members from the caps-nmmb, nssl-nmmb, and s-phys-norad subsets. While very large ensembles (e.g., hundreds of members) would be ideal if computational expense were not an issue, the “optimal” ensemble size is generally considered one in which only relatively small gains are achieved by adding additional members. Using a 50-member CAM-ensemble, Schwartz et al. (2014) found that only small gains in precipitation forecast skill are attained after about 20 members and argue that in an operational setting with limited computational resources, sizes greater than 10 would be difficult to justify. Clark et al. (2011) also found that a CAM ensemble of around 10 members has similar quantitative precipitation forecast (QPF) skill to larger ensembles and point out that the optimal number of members varies as a function of forecast length and spatial scale.

CLUE POSTPROCESSING. In past SFEs, the members from each of the unique CAM ensembles contributed to the HWT were postprocessed² by each

² Postprocessing refers to the procedure used to convert raw model output to standard grids and pressure levels, as well as to compute diagnostic quantities (e.g., convective available potential energy and storm relative helicity).

collaborator using their own software. Furthermore, some collaborators, such as CAPS, provided separate sets of postprocessed files containing ensemble-derived fields (e.g., probabilities, ensemble maximum, and ensemble mean). Thus, ingesting the datasets into the HWT workstations required different procedures to account for different file formats, fields, and grids. Furthermore, combining ensemble members from different contributors was cumbersome and rarely done, since it required an extra regridding step before computing any ensemble-derived field. Thus, standardizing the postprocessing procedure was one of the most important aspects of the CLUE since it streamlined the workflow and allowed for consistent postprocessed fields, visualization, and verification.

To standardize the postprocessing, NSSL worked closely with scientists at the Developmental Testbed Center (DTC) and NCEP’s Environmental Modeling Center (EMC) to modify the most recent version of the Unified Post-Processor (UPP) software, which is maintained by the DTC (information on the most recent version is available at www.dtcenter.org/upp/users/index.php). The UPP was modified to output a set of 107 fields from each CLUE member in gridded binary (grib2) format over a 3-km grid-spacing CONUS domain. The fields match the two-dimensional fields output by the operational HRRR and were chosen because of their relevance to a broad range of forecasting needs, including aviation, severe weather, and precipitation. Additional output fields, which were requested by NCEP’s Weather Prediction Center (WPC), SPC, and Aviation Weather Center (AWC), were also included. This special version of the UPP was distributed by NSSL to collaborators in February 2016 to allow time for testing and implementation. All contributors were asked to supply all 107 fields but were also allowed to add additional diagnostics based on their own research interests. The online supplement contains a table listing all postprocessed fields.

CLUE RESULTS. Given the sheer volume of data composing the CLUE, it is impossible to present results in this article from each of the experiments. Additionally, active research is still being conducted to examine several of the CLUE datasets. For example, DTC Visitor Program projects (www.dtcenter.org/visitors/) are currently under way, examining the value of radar data assimilation using object-based verification methods, as well as the impact of mixed physics in the CLUE. Preliminary findings and results from the 2016 SFE, including some preliminary CLUE results, can be found in Clark et al. (2016). It is important to recognize that annual HWT assessment activities typically include

a combination of subjective and objective evaluation methods (Kain et al. 2003), which together provide a more complete picture of the potential utility of new forecast techniques in an operational environment. However, given the space limitations, this section will focus on results from the single versus multicore CLUE experiments as an example of the research enabled within the CLUE framework.

Single versus multicore experiment: Severe weather verification. Objective verification of four ensemble subsets was conducted for severe weather occurrence, which included 1) NMMB, a 10-member, single-physics NMMB ensemble with perturbed ICs/LBCs; 2) ARW, a 10-member, single-physics ARW ensemble with the same perturbed ICs/LBC as NMMB; 3) MIX_{10-mem}, a combination of five of the NMMB and ARW members; and 4) MIX_{20-mem}, a combination of all 10 NMMB and ARW members. Complete datasets from these ensembles were analyzed for each day the SFE operated (24 days; 2 May–3 June, excluding weekends/holidays).

To verify severe weather (defined as a tornado, damaging winds, or large hail), ensemble-derived severe

weather probabilities were computed by considering extreme values of hourly maximum updraft helicity (UH; e.g., Kain et al. 2010) as severe storm proxies following the “surrogate severe” approach outlined by Sobash et al. (2011, 2016b). This approach has been increasingly utilized for verifying CAM-based forecasts of severe weather (e.g., Schwartz et al. 2015a,b; Sobash et al. 2016a; Gallo et al. 2016; Loken et al. 2017; Dawson et al. 2017). The basic idea behind the surrogate severe approach is that “extremes” in simulated storm diagnostics are strongly correlated with observed severe weather. However, given the inherent uncertainty associated with convection forecasts at 12–36-h lead times, coarsened grids and spatial smoothing must be applied to account for timing and displacement errors. Furthermore, the skill and reliability of surrogate severe forecasts are heavily dependent on the threshold or percentile chosen to represent extremes, as well as the amount of smoothing applied. Thus, in the methods described below, a range of UH thresholds and smoothing levels are chosen, which are known to produce reliable forecasts based on previous work.

For application of the surrogate severe approach, the maximum UH at each grid point was computed

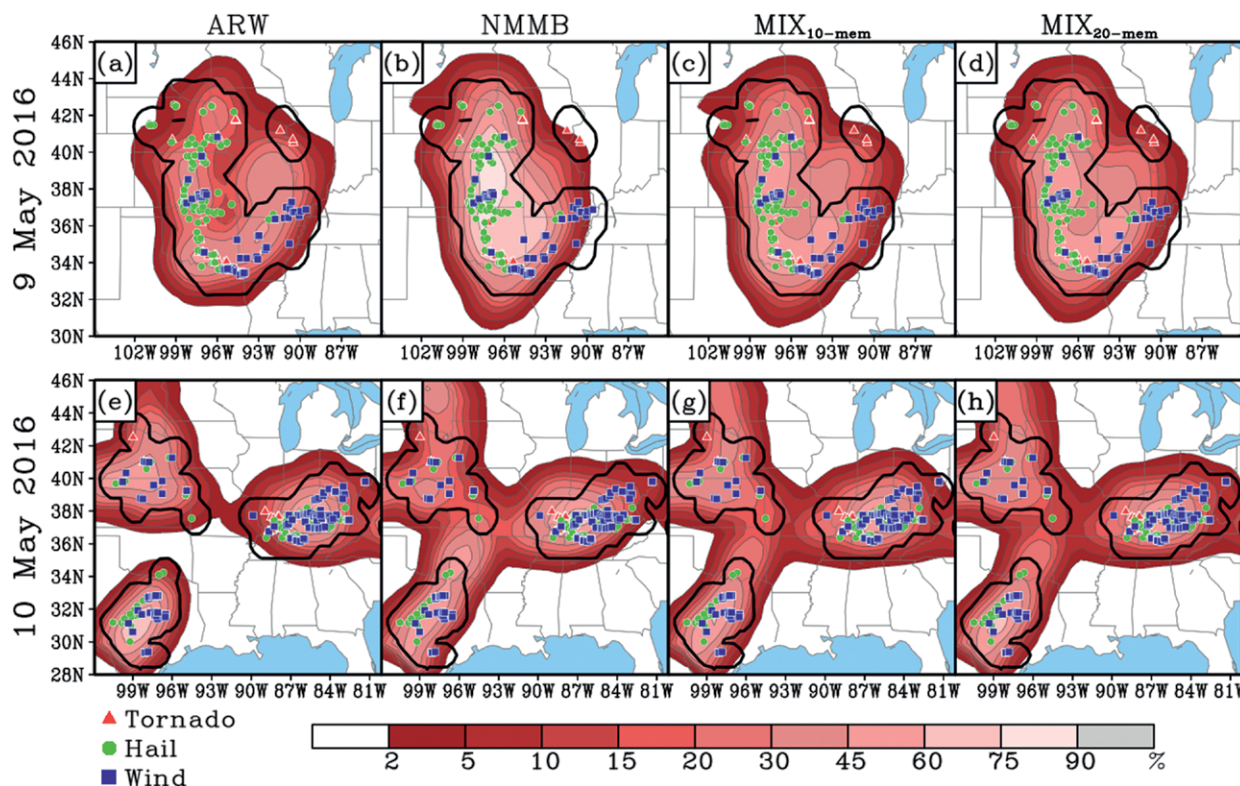


FIG. 2. SPPFs (shaded) using $\sigma = 80$ km and $p = 0.99$ for ensemble forecasts initialized at 0000 UTC 9 May 2016 and valid over forecast hours 13–36 for the ensemble subsets (a) ARW, (b) NMMB, (c) MIX_{10-mem}, and (d) MIX_{20-mem}. (e)–(h) As in (a)–(d), respectively, but for 10 May 2016. Locations of storm reports are overlaid with a legend indicating the type of report at the bottom left. The thick black contour indicates the area within 40 km of any storm report.

over the 24-h period 1200–1200 UTC (forecast hours 13–36) for each ensemble member. Then, for each member, these maximum UH values were remapped onto the 81-km NCEP 211 grid by assigning each 81-km grid box the maximum value of UH out of all 3-km grid points within the 81-km boxes. This methodology is consistent with the SPC operational day 1 convective outlook, which provides categorical and probabilistic forecasts for the 1200–1200 UTC time period and represents severe weather threats within 25 mi (~40 km) of a point. Next, severe weather probabilities [hereafter, surrogate severe probabilistic forecasts (SSPFs)] were computed by finding the ratio of members with UH greater than or equal to a specified percentile p and then applying a two-dimensional Gaussian filter to these ratios. The UH percentiles were computed separately for the set of members in each ensemble subset with the same model core using the distribution of UH values from the 81-km grids over all 24 cases. The percentiles, rather than thresholds, were used to avoid giving more weight to ensemble members with climatologically higher values of UH in the computation of SSPFs. In this dataset, the NMMB tended to have slightly higher UH than ARW (e.g., at $p = 0.99$, the UH values in NMMB and ARW were 152 and 141 $\text{m}^2 \text{s}^{-2}$, respectively).

The percentiles from 0.80 to 0.998 in increments of 0.02 (100 unique percentiles) were examined, and for each percentile, a range of standard deviations σ in the Gaussian filter from 40 to 300 km in increments of 5 km were tested (i.e., 53 unique σ values). Physically, 1σ can be thought of as the radius containing 68% of the Gaussian kernel weights. Thus, for each case and ensemble subset, there were $100 \times 53 = 5,300$ sets of SSPFs. Examples of these SSPFs using $\sigma = 80$ km and $p = 0.99$ for 9 and 10 May 2016 along with the verifying storm reports are shown in Fig. 2. To verify the SSPFs, preliminary observed storm reports from SPC (accessible at www.spc.noaa.gov/climo/reports/) were mapped onto the same 81-km grid as the SSPFs. Any grid box with one or more reports over the 1200–1200 UTC time period was assigned 1 while boxes with zero reports were assigned 0. Verification metrics were computed over the masked area displayed in Fig. 3, which was chosen to limit verification to land and near-coastal areas, as well as to eliminate the Intermountain West, where storm reports and precipitation estimates are not as reliable.

Three metrics are used for objective verification. 1) Area under the relative operating characteristic curve (AUC; Mason 1982) is computed by plotting the probability of detection (POD) versus the

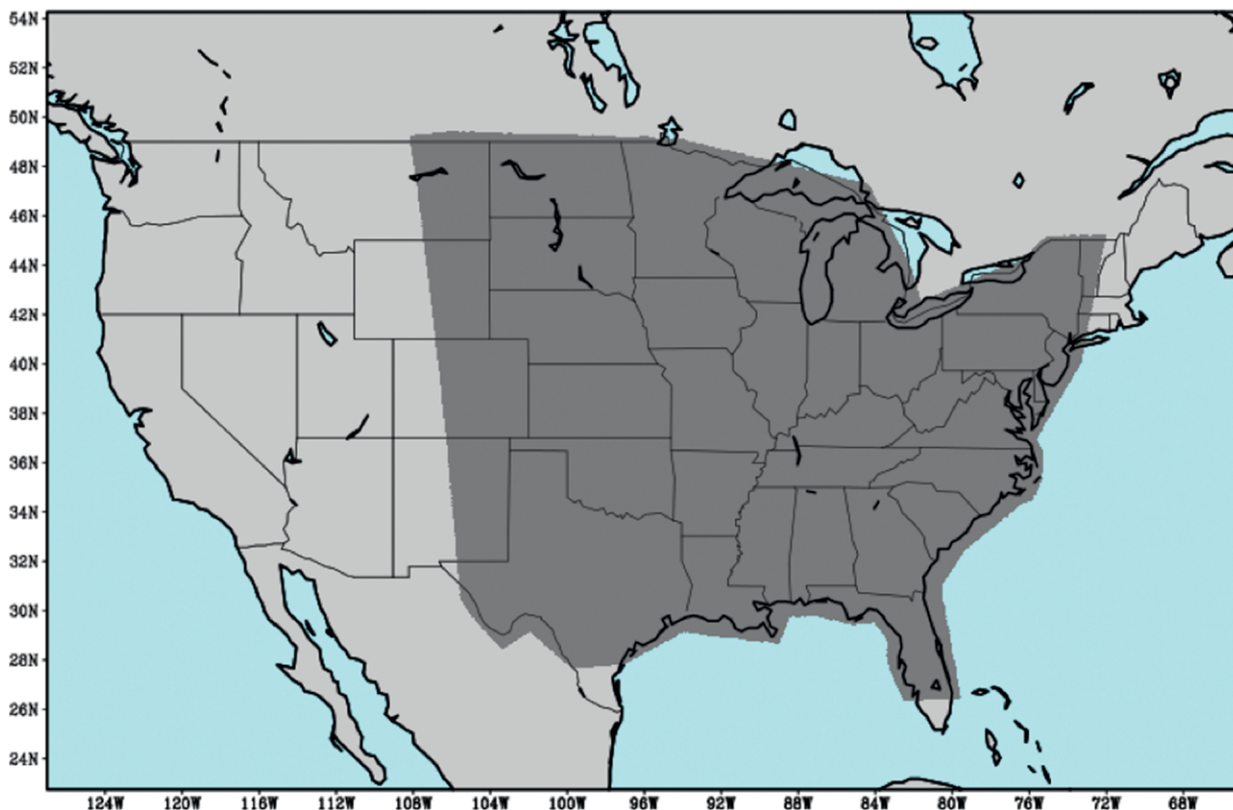


FIG. 3. Area over which verification metrics were computed.

probability of false detection (POFD) for a range of probabilistic thresholds (herein, 2% and 5%–95% in increments of 5% are used). The area under the curve connecting each POD–POFD pair is computed using a trapezoidal approximation (e.g., Wandishin et al. 2001). The AUC measures the ability of the forecast system to discriminate between events and nonevents. A value of 1.0 is considered a perfect AUC, while 0.5 and below is considered to have no skill. 2) Fractions skill score (FSS; Roberts and Lean 2008) is calculated by computing the mean-square error (MSE) of the SSPFs relative to “practically perfect” observations (e.g., Hitchens et al. 2013), which are constructed by applying a Gaussian filter with $\sigma = 120$ km to the 81-km grid of storm reports. The MSEs of the SSPFs are normalized by a worst-case reference forecast and subtracted from 1.0 to get the FSS [see Eqs. (3)–(5) in Sobash et al. (2011)]. The FSS ranges from 0 (no skill) to 1 (perfect forecast). 3) The reliability component of the Brier score (BS_{rely} ; Brier 1950; Murphy 1973) is computed by taking the squared difference of the probabilities within specified bins and their corresponding observed frequencies [see Eq. (2) in Atger (2003)]. The BS_{rely} essentially measures how closely the points within a reliability diagram follow the perfect reliability line, where the squared error for each point is weighted according to the number of forecasts within each probability bin. Lower BS_{rely} indicates increasing reliability, with $BS_{\text{rely}} = 0$ indicating perfect reliability. These three metrics were chosen because they are very well known and provide complementary information on discriminating ability (AUC), forecast accuracy (FSS), and reliability (BS_{rely}).

Each skill metric for each ensemble is presented as a function of σ and the UH percentile in Fig. 4. The metrics behave quite differently in terms of where the best scores fall within the σ –UH percentile phase space. AUC has the highest scores at relatively low σ values (60–100 km) and UH percentiles (0.82–0.86), FSS maximizes at higher σ (150–180 km) and UH percentiles (0.92–0.94), and BS_{rely} is best at the highest σ (240–300 km) and UH percentiles (0.95–0.96). For reference, in each panel in Fig. 4, the UH percentile at which the number of surrogate severe storm reports is approximately equal to the number of observed severe reports over all cases is shown by the turquoise dashed line ($p = 0.974$; i.e., bias = 1.0). Thus, AUCs maximize at UH percentiles associated with biases well above 1.0. In fact, the biases at these lower ranges of UH percentiles range from 6.0 to 7.8 (not shown). The high biases associated with the maximum AUCs

are not surprising because AUC does not account for bias or reliability. Furthermore, for rare-event forecasts, increasing the number of forecast events almost always acts to increase the POD more than the POFD, thereby increasing the AUC, because correct negatives so heavily weight the POFD. For FSS and BS_{rely} , the scores maximize at UH percentiles closer to bias = 1.0 than AUC. For BS_{rely} , it may seem intuitive that the best reliability would occur when bias = 1.0; however, underdispersion causes probabilities to be too high, and the additional spatial uncertainty provided by a bias slightly higher than 1.0 along with very strong smoothing apparently achieves the best reliability.

For AUC and FSS, the MIXED_{10-mem} and MIXED_{20-mem} have slightly higher maximum scores than ARW and NMMB, which are very similar to each other. The best BS_{rely} values are nearly identical among the four ensembles. To evaluate whether any of the differences in maximum scores were significant, the resampling approach of Hamill (1999) was utilized and it was found that none of the differences between ensembles were significant at $\alpha = 0.05$. Thus, although the multicore approach has slightly higher scores than the single-model approach for severe weather forecasting, a larger sample is necessary to determine whether these differences can be attributed to more than just randomness. Continuation of CLUE-related experiments in subsequent years will contribute to larger data samples and lead to more robust statistical results.

Single versus multicore experiment: QPF verification.

Similar to severe weather, accurate precipitation forecasting is notoriously difficult for numerical weather prediction (NWP) models (e.g., Carbone et al. 2002; Roebber et al. 2004), but CAM-based systems have led to major improvements in QPFs over convection-parameterizing models (e.g., Clark et al. 2009, 2010a, 2012; Weisman et al. 2008; Iyer et al. 2016). In many ways, QPF verification is simpler than severe weather, because more reliable and higher-resolution observational precipitation datasets exist (e.g., NCEP’s Stage IV dataset), and QPFs are directly output from models and thus do not require surrogates like severe weather. Although these differences allow QPF verification to be reliably performed at higher spatial and temporal resolution than severe weather, to compare QPF and severe weather performance, verification is performed at the same scale as severe weather. Follow-up work will perform QPF verification at higher spatial and temporal resolution.

To perform QPF verification analogously to severe weather, 24-h accumulated precipitation from

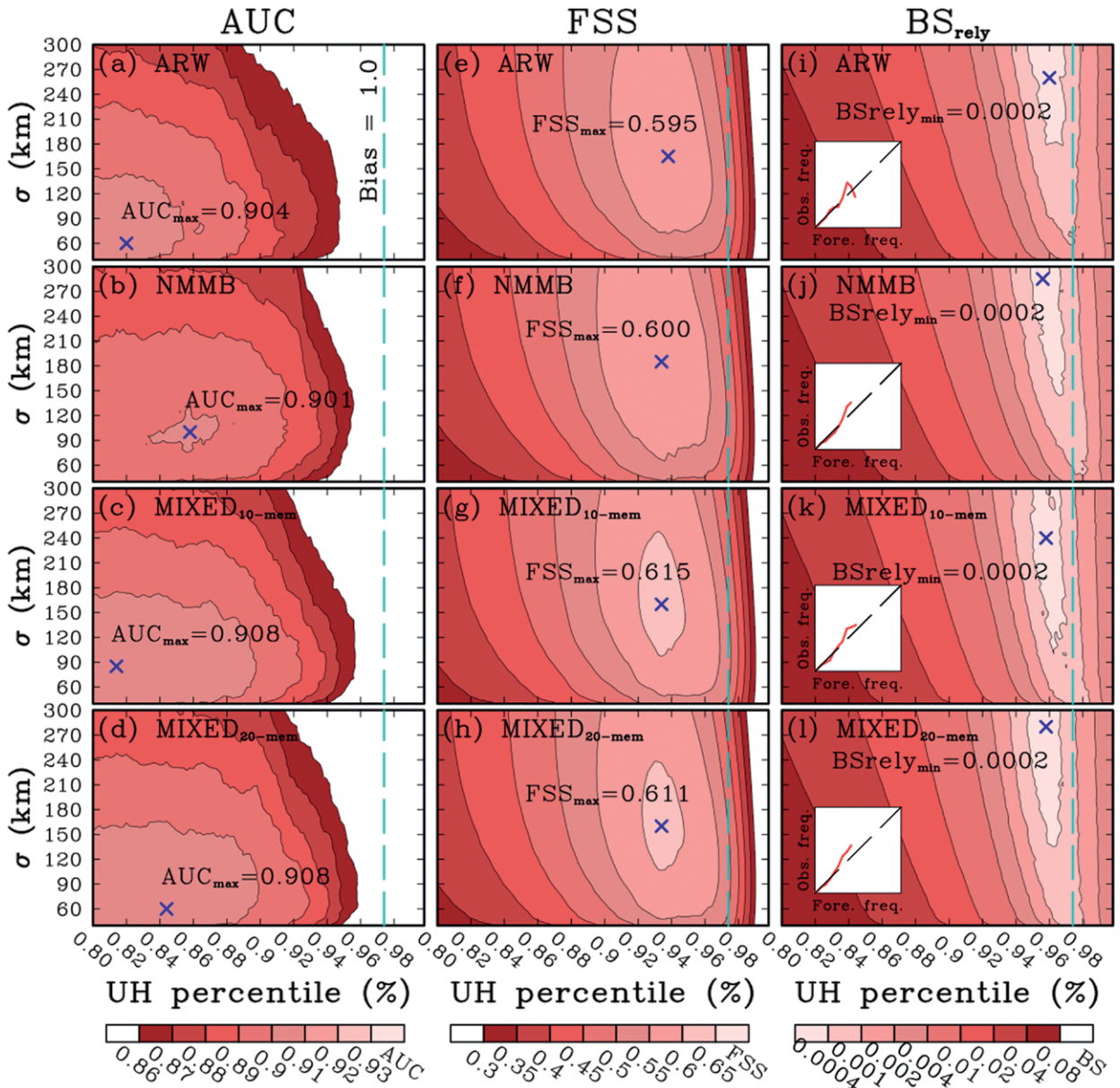


FIG. 4. AUC as a function of σ and UH percentile for the ensembles (a) ARW, (b) NMMB, (c) MIXED_{10-mem}, and (d) MIXED_{20-mem}. (e)–(h) As in (a)–(d), respectively, but for FSS. (i)–(l) As in (a)–(d), respectively, but for BS_{rely}. In each panel, a blue x marks the best score, which is indicated in the text; the vertical dashed turquoise line marks the UH percentile at which bias = 1 (i.e., the number of surrogate severe reports approximately matches the number of observed reports). In (i)–(l), the reliability diagrams are shown corresponding to σ and UH percentile at which BS_{rely} is minimized.

NCEP’s 4-km grid-spacing Stage IV dataset (Lin and Mitchell 2005; Nelson et al. 2016) was remapped onto the NCEP 211 grid using the maximum 24-h precipitation amount from all 4-km Stage IV grid points within each 81-km NCEP 211 grid box. Then, the precipitation amount that resulted in the same number of observed severe weather events was found, which was 2.69 in. (1 in. = 2.54 cm). In other words, the total number of 81-km grid boxes with an observed severe weather report was equal to the total number

of 81-km grid boxes in which the maximum observed precipitation was 2.69 in. or greater. Then, maximum 24-h accumulated QPFs from each ensemble member were remapped onto the 81-km grid in the same manner as UH, and heavy rainfall probabilities were also computed similarly to UH.

Figure 5 shows that all the precipitation skill metrics computed using the 2.69-in. threshold were noticeably higher than those for UH; thus, these CAM ensembles provide more skillful forecasts of extreme

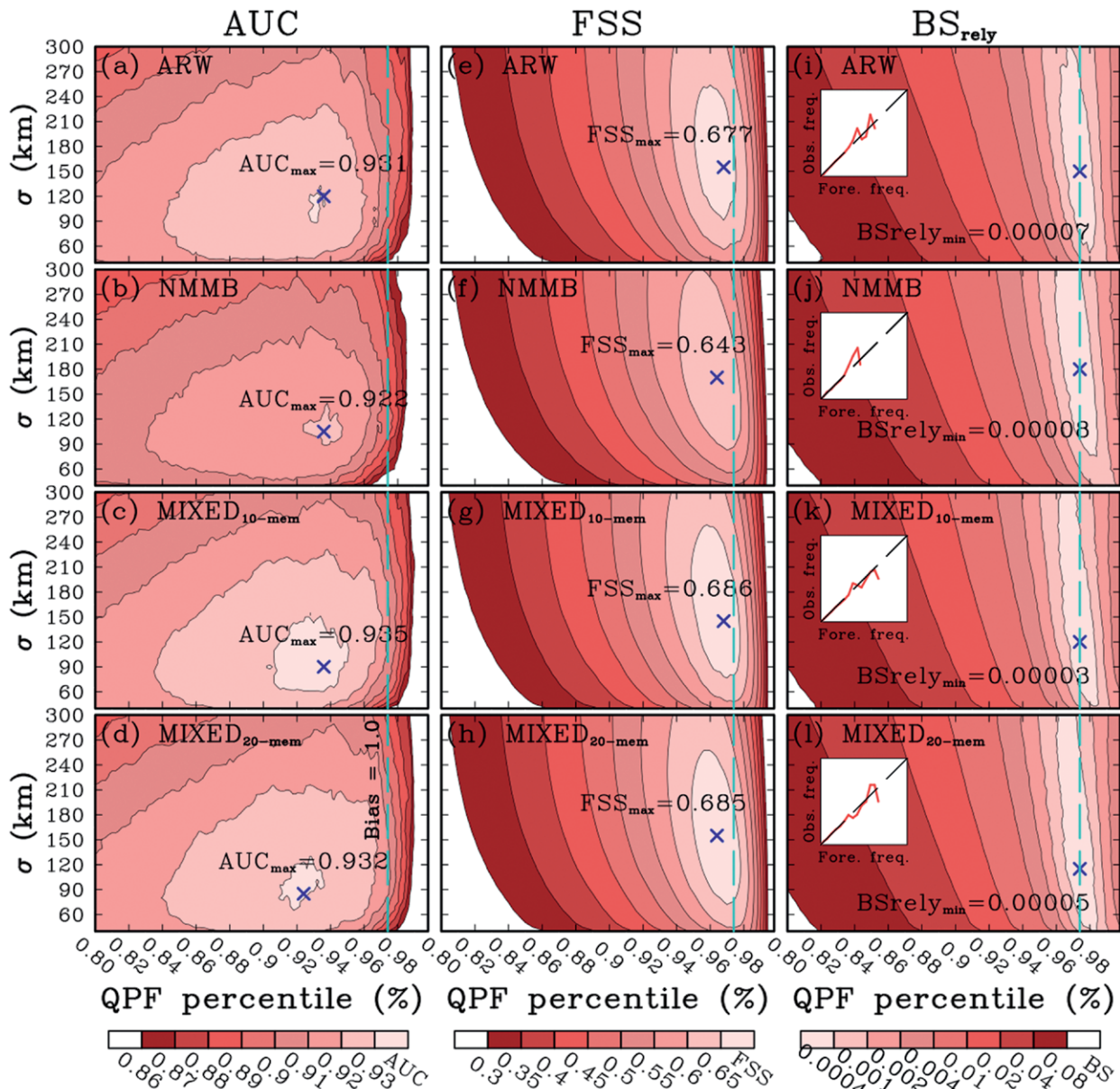


FIG. 5. As in Fig. 4, but for 24-h precipitation forecasts.

rainfall than severe weather. Also, the multicore systems (MIXED_{10-mem} and MIXED_{20-mem}) had better AUC, FSS, and BS_{rely} results than the single-core systems, but as with UH, none of the differences between the single- and multicore subsets were statistically significant. There are also noticeable differences in σ and the rainfall percentiles at which the metrics are maximized. Namely, relative to UH, there is a shift toward higher percentiles (biases closer to 1.0) and smaller σ (i.e., less smoothing) for QPF. The reason for this shift is not clear, but it is speculated that it is because of differences in the spatial characteristics of heavy rainfall and severe weather. However, further work is needed for substantiation.

CONCLUDING REMARKS. The 2016 SFE marks year 17 of annual SFEs organized by the SPC and NSSL, which aim to accelerate the advancement of new technologies and concepts from research to operations for improving hazardous convective weather prediction. Since 2004, a main focus of SFEs has been on evaluating performance characteristics of CAMs, as well as making advances in creating, importing, processing, verifying, and extracting unique hazardous weather fields, as well as providing analysis and visualization tools for CAMs. With increasing numbers of CAMs contributed to SFEs every year, and the strong community call for evidence-driven decision-making as EMC and the modeling

community configure the first generation of operational CAM-based ensemble prediction systems, a major initiative was started during the 2016 SFE to coordinate and standardize the CAM contributions from each of our external collaborators, so that each group of CAMs could be considered part of one large ensemble termed the Community Leveraged Unified Ensemble (CLUE).

The CLUE was designed to enable up to eight different controlled experiments focused on optimizing CAM ensemble configurations. Results from one of these experiments—single versus multicore ensemble design—were reported upon herein, while research is in progress for several other CLUE experiments. For the single versus multicore results, objective metrics for severe weather forecast skill indicated small differences in forecast skill, with multicore systems having slightly higher scores than those for a single core. Additionally, a 20-member mixed-core ensemble performed almost identically to a 10-member mixed-core ensemble. None of the differences were statistically significant, but with only 24 cases, significance would likely require a larger sample size.

For precipitation verification, probabilistic QPFs were found to be more skillful than those for severe weather when the verification was performed similarly. Additionally, the mixed-core ensembles had slightly better objective metrics for QPF than the single-core ensembles. Future work is planned to perform the precipitation verification from the multicore versus single-core ensemble design experiment at higher spatial and temporal resolution.

HWT has a long and productive history of bringing together different parts of the research, operational, and academic meteorological communities to work collaboratively in a real-time simulated forecasting environment, focusing on severe weather forecasting problems. We envision continuing the CLUE system in subsequent experiments, and there is ample reason to believe that it can further enhance effective engagement between the modeling and operational communities, as well as provide important scientific evidence necessary for informed decision-making, so that future U.S. hazardous weather prediction capabilities are the best possible.

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CLUE MEMBER SPECIFICATIONS AND OUTPUT FIELDS

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This document is a supplement to “The Community Leveraged Unified Ensemble (CLUE) in the 2016 NOAA/Hazardous Weather Testbed Spring Forecasting Experiment,” by Adam J. Clark, Israel L. Jirak, Scott R. Dembek, Gerry J. Creager, Fanyou Kong, Kevin W. Thomas, Kent H. Knopfmeier, Burkely T. Gallo, Christopher J. Melick, Ming Xue, Keith A. Brewster, Youngsun Jung, Aaron Kennedy, Xiquan Dong, Joshua Markel, Matthew Gilmore, Glen S. Romine, Kathryn R. Fossell, Ryan A. Sobash, Jacob R. Carley, Brad S. Ferrier, Matthew Pyle, Curtis R. Alexander, Steven J. Weiss, John S. Kain, Louis J. Wicker, Gregory Thompson, Rebecca D. Adams-Selin, and David A. Imy (*Bull. Amer. Meteor. Soc.*, **99**, 1433–1448) • ©2018 American Meteorological Society • *Corresponding author*: Adam J. Clark, adam.clark@noaa.gov • DOI:10.1175/BAMS-D-16-0309.2

All Community Leveraged Unified Ensemble (CLUE) members were initialized weekdays at 0000 UTC with 3-km grid spacing covering a contiguous U.S. (CONUS) domain. Members of the Advanced Research version of the Weather Research and Forecasting (WRF) Model (ARW) had 1,680 grid points in the east–west direction, 1,152 grid points in the north–south direction, and 51 vertical levels^{ES1} with a model top of 50 hPa and used the Rapid Radiative Transfer Model (RRTM; Mlawer

et al. 1997) for general circulation models (RRTMG; Iacono et al. 2008) for short- and longwave radiation. All Nonhydrostatic Multiscale Model on the B grid (NMMB) members have 1,568 grid points in the east–west direction, 1,120 grid points in the north–south direction, and 50 vertical levels with a model top of 50 hPa and use the RRTM short- and longwave radiation. Depending on the CLUE subset, forecast lengths range from 36 to 60 h. Specifications for the members within each subset are detailed in Tables ES1–ES10.

Developmental Testbed Center (DTC) Unified Post-Processor (UPP) software was used to output a set of 107 fields from each CLUE member, which are listed in Table ES11. Additionally, the Center for Analysis and Prediction of Storms (CAPS) and National Severe Storms Laboratory (NSSL) output six additional experimental diagnostics that are listed in Table ES12.

^{ES1} WRF-ARW sigma levels are set to the following values: 1.0, 0.998, 0.994, 0.987, 0.975, 0.959, 0.939, 0.916, 0.892, 0.865, 0.835, 0.802, 0.766, 0.727, 0.685, 0.64, 0.592, 0.542, 0.497, 0.4565, 0.4205, 0.3877, 0.3582, 0.3317, 0.3078, 0.2863, 0.267, 0.2496, 0.2329, 0.2188, 0.2047, 0.1906, 0.1765, 0.1624, 0.1483, 0.1342, 0.1201, 0.106, 0.0919, 0.0778, 0.0657, 0.0568, 0.0486, 0.0409, 0.0337, 0.0271, 0.0209, 0.0151, 0.0097, 0.0047, and 0.0.

TABLE ES1. Specifications for the core members of the CLUE that use the ARW, mixed-physics parameterizations, perturbed initial conditions (ICs) and lateral boundary conditions (LBCs), and radar data assimilation using Advanced Regional Prediction System (ARPS) three-dimensional variational data assimilation (3DVAR). The 12-km North American Mesoscale Forecast System (NAM) analysis and forecast are referred to as NAMa and NAMf, respectively. Here, 3DVAR refers to ARPS 3DVAR and cloud analysis. The 13-km Rapid Refresh (RAP) model analysis is RAPa. In the IC column, the model names appended with “pert” refer to perturbations extracted from a 16-km grid-spacing SREF member. For members core03–10, entries in the BC column refer to SREF member forecasts. The core02 member was not ready for real-time implementation; thus, it was not run in real time. Microphysics parameterizations include Thompson et al. (2004), the Predicted Particle Properties scheme (P3; Morrison and Milbrandt 2015), Milbrandt and Yau (2005a,b), and Morrison et al. (2009); planetary boundary layer parameterizations include Mellor–Yamada–Janjić (MYJ; Mellor and Yamada 1982; Janjić 2002), Mellor–Yamada–Nakanishi–Niino (MYNN; Nakanishi 2000, 2001; Nakanishi and Niino 2004, 2006), and Yonsei University (YSU; Noh et al. 2003); and land surface models (LSMs) included the Noah model (Chen and Dudhia 2001) and the Rapid Update Cycle (RUC; Smirnova et al. 1997, 2000).

Member	IC	BC	Microphysics	LSM	PBL	Model
core01	NAMa + 3DVAR	NAMf	Thompson	Noah	MYJ	ARW
core02	RAPa + 3DVAR	GFSf	Thompson	RUC	MYNN	ARW
core03	core01 + arw-pl_pert	arw-pl	P3	Noah	YSU	ARW
core04	core01 + arw-nl_pert	arw-nl	Milbrandt–Yau	Noah	MYNN	ARW
core05	core01 + arw-p2_pert	arw-p2	Morrison	Noah	MYJ	ARW
core06	core01 + arw-n2_pert	arw-n2	P3	Noah	YSU	ARW
core07	core01 + nmmb-pl_pert	nmmb-pl	Milbrandt–Yau	Noah	MYNN	ARW
core08	core01 + nmmb-nl_pert	nmmb-nl	Morrison	Noah	YSU	ARW
core09	core01 + nmmb-p2_pert	nmmb-p2	P3	Noah	MYJ	ARW
core10	core01 + nmmb-n2_pert	nmmb-n2	Thompson	Noah	MYNN	ARW

TABLE ES2. Specifications for the s-phys-rad members of the CLUE that use the ARW dynamic core, single physics, perturbed ICs and LBCs, and radar data assimilation using ARPS 3DVAR. The core01 member is repeated from Table ES1 because it is also the control member of the s-phys-rad ensemble subset.

Member	IC	BC	Microphysics	LSM	PBL	Model
core01	NAMa + 3DVAR	NAMf	Thompson	Noah	MYJ	ARW
s-phys-rad02	core01 + arw-pl_pert	arw-pl	Thompson	Noah	MYJ	ARW
s-phys-rad03	core01 + arw-nl_pert	arw-nl	Thompson	Noah	MYJ	ARW
s-phys-rad04	core01 + arw-p2_pert	arw-p2	Thompson	Noah	MYJ	ARW
s-phys-rad05	core01 + arw-n2_pert	arw-n2	Thompson	Noah	MYJ	ARW
s-phys-rad06	core01 + arw-p3_pert	arw-p3	Thompson	Noah	MYJ	ARW
s-phys-rad07	core01 + nmmb-pl_pert	nmmb-pl	Thompson	Noah	MYJ	ARW
s-phys-rad08	core01 + nmmb-nl_pert	nmmb-nl	Thompson	Noah	MYJ	ARW
s-phys-rad09	core01 + nmmb-p2_pert	nmmb-p2	Thompson	Noah	MYJ	ARW
s-phys-rad10	core01 + nmmb-n2_pert	nmmb-n2	Thompson	Noah	MYJ	ARW

TABLE ES3. Specifications for the caps-enkf members of the CLUE. For these members, a 3-km Grid-point Statistical Interpolation analysis system–ensemble Kalman filter (GSI–EnKF) system was initialized at 1800 UTC each day, which assimilated the RAP/High Resolution Rapid Refresh (HRRR) GSI data stream hourly (except satellite data) from 1800 to 0000 UTC over the CONUS domain. Radar data were assimilated every 15 min from 2300 to 0000 UTC using the CAPS EnKF system. The ensemble consists of 40 ARW members with initial perturbations and mixed-physics options to provide input for the EnKF ensemble analyses. Each member uses Thompson microphysics, although with varied parameter settings. A nine-member ensemble forecast (run for 60 h) follows using the final EnKF analyses at 0000 UTC employing the same multiphysics configurations as are used for the core members. In addition, one deterministic forecast from the ensemble mean analysis at 0000 UTC is also produced.

Member	IC	BC	Microphysics	LSM	PBL	Model
caps-enkf01	enkf_m01a	NAMf	Thompson	Noah	MYJ	ARW
caps-enkf02	enkf_mna	GFSf	Thompson	Noah	MYNN	ARW
caps-enkf03	enkf_m08a	arw-p1	P3	Noah	YSU	ARW
caps-enkf04	enkf_m10a	arw-n1	MY	Noah	MYNN	ARW
caps-enkf05	enkf_m17a	arw-p2	Morrison	Noah	MYJ	ARW
caps-enkf06	enkf_m23a	arw-n2	P3	Noah	YSU	ARW
caps-enkf07	enkf_m24a	nmmb-p1	MY	Noah	MYNN	ARW
caps-enkf08	enkf_m12a	nmmb-n1	Morrison	Noah	YSU	ARW
caps-enkf09	enkf_m09a	nmmb-p2	P3	Noah	MYJ	ARW
caps-enkf10	enkf_m6a	nmmb-n2	Thompson	Noah	MYNN	ARW

TABLE ES4. Specifications for the caps-nmmb-rad member of CLUE that uses the Nonhydrostatic Multiscale Model on the B grid (NMMB) dynamic core with radar data assimilation and cloud analysis using ARPS 3DVAR with Ferrier–Aligo microphysics (Aligo et al. 2014).

Member	IC	BC	Microphysics	LSM	PBL	Model
caps-nmmb-rad	NAMa + 3DVAR	NAMf	Ferrier–Aligo	Noah	MYJ	NMMB

TABLE ES5. Specifications for the caps-nmmb members of the CLUE that use the NMMB dynamic core, single physics, IC and LBC perturbations extracted from SREF members, and no-radar data assimilation.

Member	IC	BC	Microphysics	LSM	PBL	Model
caps-nmmb01	NAMa + arw-p3_pert	arw-p3	Ferrier–Aligo	Noah	MYJ	NMMB
caps-nmmb02	NAMa + nmmb-p1_pert	nmmb-p1	Ferrier–Aligo	Noah	MYJ	NMMB
caps-nmmb03	NAMa + nmmb-n1_pert	nmmb-n1	Ferrier–Aligo	Noah	MYJ	NMMB
caps-nmmb04	NAMa + nmmb-p2_pert	nmmb-p2	Ferrier–Aligo	Noah	MYJ	NMMB
caps-nmmb05	NAMa + nmmb-n2_pert	nmmb-n2	Ferrier–Aligo	Noah	MYJ	NMMB

TABLE ES6. Specifications for the s-phys-norad members of the CLUE that use the ARW dynamic core, single physics, perturbed ICs and LBCs from SREF members, and no-radar data assimilation.

Member	IC	BC	Microphysics	LSM	PBL	Model
s-phys-norad01	NAMa	NAMf	Thompson	Noah	MYJ	ARW
s-phys-norad02	NAMa + arw-pl_pert	arw-pl	Thompson	Noah	MYJ	ARW
s-phys-norad03	NAMa + arw-nl_pert	arw-nl	Thompson	Noah	MYJ	ARW
s-phys-norad04	NAMa + arw-p2_pert	arw-p2	Thompson	Noah	MYJ	ARW
s-phys-norad05	NAMa + arw-n2_pert	arw-n2	Thompson	Noah	MYJ	ARW
s-phys-norad06	NAMa + arw-p3_pert	arw-p3	Thompson	Noah	MYJ	ARW
s-phys-norad07	NAMa + nmmb-pl_pert	nmmb-pl	Thompson	Noah	MYJ	ARW
s-phys-norad08	NAMa + nmmb-nl_pert	nmmb-nl	Thompson	Noah	MYJ	ARW
s-phys-norad09	NAMa + nmmb-p2_pert	nmmb-p2	Thompson	Noah	MYJ	ARW
s-phys-norad10	NAMa + nmmb-n2_pert	nmmb-n2	Thompson	Noah	MYJ	ARW

TABLE ES7. Specifications for the nssl_nmmb members of the CLUE that use the NMMB dynamic core, single physics, IC and LBC perturbations extracted from SREF members, and no-radar data assimilation.

Member	IC	BC	Microphysics	LSM	PBL	Model
nssl-nmmb01	NAMa	NAMf	Ferrier–Aligo	Noah	MYJ	NMMB
nssl-nmmb02	NAMa + arw-pl_pert	arw-pl	Ferrier–Aligo	Noah	MYJ	NMMB
nssl-nmmb03	NAMa + arw-nl_pert	arw-nl	Ferrier–Aligo	Noah	MYJ	NMMB
nssl-nmmb04	NAMa + arw-p2_pert	arw-p2	Ferrier–Aligo	Noah	MYJ	NMMB
nssl-nmmb05	NAMa + arw-n2_pert	arw-n2	Ferrier–Aligo	Noah	MYJ	NMMB

TABLE ES8. Specification for the HRRR36 member of CLUE that uses the ARW dynamic core with HRRR physics and data assimilation.

Member	IC	BC	Microphysics	LSM	PBL	Model
HRRR36	RAP	GFSf	Thompson	RUC	MYNN	ARW

TABLE ES9. Specifications for the ncar-enkf members of the CLUE. This ensemble provides forecasts to 48 h and uses the National Center for Atmospheric Research's (NCAR) Data Assimilation Research Testbed (DART) software with ARW, version 3.6.1, and a slightly different horizontal domain than other CLUE members. The analysis system is composed of 50 members that are continuously cycled using the ensemble adjustment Kalman filter (EAKF; Anderson 2001, 2003). New analyses are produced every 6 h with 15-km grid spacing. Other specifications include 51 vertical levels with a 50-hPa top (same vertical levels as other ARW CLUE members), a horizontal localization of 1,270 km and vertical localization of two scale heights, adaptive prior inflation, adaptive localization, sampling error correction, and freely evolving soil states. The following observational sources are utilized: Meteorological Assimilation Data Ingest System (MADIS) Aircraft Communications Addressing and Reporting System (ACARS), aviation routine weather reports (METARs), radiosondes, National Centers for Environmental Prediction (NCEP) real-time marine data (MARINE), Cooperative Institute for Mesoscale Meteorological Studies (CIMMS) cloud-track winds, Oklahoma Mesonet, and GPS radio occultation. All members have constant physics, which include Tiedtke (1989) cumulus parameterization, Thompson microphysics, MYJ PBL, Noah land surface model, and RRTMG shortwave and longwave radiation with aerosol and ozone climatologies. The 10-member forecasts are initialized at 0000 UTC daily with ICs provided by downscaled members of 0000 UTC WRF/DART EAKF analyses (described above). Perturbed LBCs from GFS forecasts are used. The physics are the same as those from the data assimilation system, but without cumulus parameterization (detailed below). Additional information on this ensemble can be found in Schwartz et al. (2015).

Member	IC	BC	Microphysics	LSM	PBL	Model
ncar-enkf01	anal01	GEFS01	Thompson	Noah	MYJ	ARW
ncar-enkf02	anal02	GEFS02	Thompson	Noah	MYJ	ARW
ncar-enkf03	anal03	GEFS03	Thompson	Noah	MYJ	ARW
ncar-enkf04	anal04	GEFS04	Thompson	Noah	MYJ	ARW
ncar-enkf05	anal05	GEFS05	Thompson	Noah	MYJ	ARW
ncar-enkf06	anal06	GEFS06	Thompson	Noah	MYJ	ARW
ncar-enkf07	anal07	GEFS07	Thompson	Noah	MYJ	ARW
ncar-enkf08	anal08	GEFS08	Thompson	Noah	MYJ	ARW
ncar-enkf09	anal09	GEFS09	Thompson	Noah	MYJ	ARW
ncar-enkf10	anal10	GEFS10	Thompson	Noah	MYJ	ARW

TABLE ES10. Specifications for the s-phys-norad members of the CLUE that use the ARW dynamic core, single physics, perturbed ICs and LBCs from SREF members, and no-radar data assimilation.

Member	IC	BC	Microphysics	LSM	PBL	Model
core01	NAMa + 3DVAR	NAMf	Thompson	Noah	MYJ	ARW
mp2	NAMa + 3DVAR	NAMf	Morrison	Noah	MYJ	ARW
mp3	NAMa + 3DVAR	NAMf	MY	Noah	MYJ	ARW
mp4	NAMa + 3DVAR	NAMf	P3	Noah	MYJ	ARW
mp5	NAMa + 3DVAR	NAMf	WSM6	Noah	MYJ	ARW

TABLE ES11. The set of 107 required output diagnostics for the CLUE members, which are output at hourly intervals.

No.	Level/layer	Parameter	Description
1	Entire atmosphere	REFC	Composite reflectivity (dB)
2	Cloud top	RETOP	Echo top (m)
3	Entire atmosphere	VIL	Vertically integrated liquid (kg m^{-2})
4	Surface	VIS	Visibility (m)
5	1,000 m above ground	REFD	Reflectivity (dB)
6	4,000 m above ground	REFD	Reflectivity (dB)
7	Surface	GUST	Wind speed (gust) (m s^{-1})
8	500 hPa	HGT	Geopotential height (gpm)
9	500 hPa	TMP	Temp (K)
10	500 hPa	DPT	Dewpoint temp (K)
11	500 hPa	UGRD	<i>U</i> component of wind (m s^{-1})
12	500 hPa	VGRD	<i>V</i> component of wind (m s^{-1})
13	700 hPa	HGT	Geopotential height (gpm)
14	700 hPa	TMP	Temp (K)
15	700 hPa	DPT	Dewpoint temp (K)
16	700 hPa	UGRD	<i>U</i> component of wind (m s^{-1})
17	700 hPa	VGRD	<i>V</i> component of wind (m s^{-1})
18	850 hPa	HGT	Geopotential height (gpm)
19	850 hPa	TMP	Temp (K)
20	850 hPa	DPT	Dewpoint temp (K)
21	850 hPa	UGRD	<i>U</i> component of wind (m s^{-1})
22	850 hPa	VGRD	<i>V</i> component of wind (m s^{-1})
23	925 hPa	TMP	Temp (K)
24	925 hPa	DPT	Dewpoint temp (K)
25	925 hPa	UGRD	<i>U</i> component of wind (m s^{-1})
26	925 hPa	VGRD	<i>V</i> component of wind (m s^{-1})
27	1,000 hPa	TMP	Temp (K)
28	1,000 hPa	DPT	Dewpoint temp (K)
29	1,000 hPa	UGRD	<i>U</i> component of wind (m s^{-1})
30	1,000 hPa	VGRD	<i>V</i> component of wind (m s^{-1})
31	400–1,000 hPa above ground	MAXUVV	Hourly max upward vertical velocity—lowest 400 hPa (m s^{-1})
32	400–1,000 hPa above ground	MAXDVV	Hourly max downward vertical velocity—lowest 400 hPa (m s^{-1})
33	0.5–0.8 sigma layer	DZDT	Vertical velocity (geometric; m s^{-1})
34	Mean sea level	PRMSL	Pressure reduced to MSL (Pa)
35	1,000 hPa	HGT	Geopotential height (gpm)
36	1,000 m above ground	MAXREF	Hourly max of simulated reflectivity at 1 km AGL (dB)
37	5,000–2,000 m above ground	MXUPHL	Hourly max updraft helicity: 2–5 km AGL ($\text{m}^2 \text{s}^{-2}$)
38	Entire column	TCOLG	Total column-integrated graupel (kg m^{-2})
39	Surface	LTNG	Lightning (—)
40	80 m above ground	UGRD	<i>U</i> component of wind (m s^{-1})
41	80 m above ground	VGRD	<i>V</i> component of wind (m s^{-1})
42	Surface	PRES	Pressure (Pa)
43	Surface	HGT	Geopotential height (gpm)

TABLE ES11. Continued.			
No.	Level/layer	Parameter	Description
44	Surface	TMP	Temp (K)
45	0 m underground	MSTAV	Moisture availability (%)
46	Surface	WEASD	Water equivalent of accumulated snow depth (kg m ⁻²)
47	Surface	SNOWC	Snow cover (%)
48	Surface	SNOD	Snow depth (m)
49	2 m above ground	TMP	Temp (K)
50	2 m above ground	SPFH	Specific humidity (kg kg ⁻¹)
51	2 m above ground	DPT	Dewpoint temp (K)
52	10 m above ground	UGRD	<i>U</i> component of wind (m s ⁻¹)
53	10 m above ground	VGRD	<i>V</i> component of wind (m s ⁻¹)
54	10 m above ground	WIND	Wind speed (m s ⁻¹)
55	Surface	CPOFP	Percent frozen precipitation (%)
56	Surface	PRATE	Precipitation rate (kg m ⁻² s ⁻¹)
57	Surface	APCP	Total precipitation (kg m ⁻²)
58	Surface	WEASD	Water equivalent of accumulated snow depth (kg m ⁻²)
59	Surface	APCP	Precipitation (kg m ⁻²): hourly total
60	Surface	WEASD	Water equivalent of accumulated snow depth (kg m ⁻²)
61	Surface	CSNOW	Categorical snow (—)
62	Surface	CICEP	Categorical ice pellets (—)
63	Surface	CFRZR	Categorical freezing rain (—)
64	Surface	CRAIN	Categorical rain (—)
65	Surface	VGTYP	Vegetation type [integer (0–13)]
66	500–1,000 hPa	LFTX	Surface lifted index (K)
67	Surface	CAPE	Convective available potential energy (J kg ⁻¹)
68	Surface	CIN	Convective inhibition (J kg ⁻¹)
69	Entire column	PWAT	Precipitable water (kg m ⁻²)
70	Low-cloud layer	LCDC	Low-cloud cover (%)
71	Middle-cloud layer	MCDC	Medium-cloud cover (%)
72	High-cloud layer	HCDC	High-cloud cover (%)
73	Entire atmosphere	TCDC	Total cloud cover (%)
74	Cloud base	PRES	Pressure (Pa)
75	Cloud base	HGT	Geopotential height (gpm)
76	Cloud ceiling	HGT	Geopotential height (gpm)
77	Cloud top	PRES	Pressure (Pa)
78	Cloud top	HGT	Geopotential height (gpm)
79	Top of atmosphere	ULWRF	Upward longwave radiation flux (W m ⁻²)
80	Surface	DSWRF	Downward shortwave radiation flux (W m ⁻²)
81	3,000–0 m above ground	HLCY	Storm-relative helicity (m ² s ⁻²)
82	1,000–0 m above ground	HLCY	Storm-relative helicity (m ² s ⁻²)
83	0–6,000 m above ground	USTM	<i>U</i> component of storm motion (m s ⁻¹)
84	0–6,000 m above ground	VSTM	<i>V</i> component of storm motion (m s ⁻¹)
85	0–1,000 m above ground	VUCSH	Vertical <i>U</i> -component shear (s ⁻¹)
86	0–1,000 m above ground	VVCSH	Vertical <i>V</i> -component shear (s ⁻¹)
87	0–6,000 m above ground	VUCSH	Vertical <i>U</i> -component shear (s ⁻¹)

TABLE ES11. Continued.

No.	Level/layer	Parameter	Description
88	0–6,000 m above ground	VVCSH	Vertical V-component shear (s^{-1})
89	180–0 hPa above ground	4LFTX	Best (four layer) lifted index (K)
90	180–0 hPa above ground	CAPE	Convective available potential energy ($J kg^{-1}$)
91	180–0 hPa above ground	CIN	Convective inhibition ($J kg^{-1}$)
92	Surface	HPBL	Planetary boundary layer height (m)
93	Lifted condensation level	HGT	Geopotential height (gpm)
94	90–0 hPa above ground	CAPE	Convective available potential energy ($J kg^{-1}$)
95	90–0 hPa above ground	CIN	Convective inhibition ($J kg^{-1}$)
96	255–0 hPa above ground	CAPE	Convective available potential energy ($J kg^{-1}$)
97	255–0 hPa above ground	CIN	Convective inhibition ($J kg^{-1}$)
98	Equilibrium level	HGT	Geopotential height (gpm)
99	255–0 hPa above ground	PLPL	Pressure of level from which parcel was lifted (Pa)
100	Surface	LAND	Land cover (0 = sea, 1 = land; proportion)
101	Surface	ICEC	Ice cover (proportion)
102	250 hPa	UGRD	U component of wind ($m s^{-1}$)
103	250 hPa	VGRD	V component of wind ($m s^{-1}$)
104	250 hPa	HGT	Geopotential height (gpm)
105	250 hPa	TMP	Temp (K)
106	700 hPa	VVEL	Vertical velocity ($m s^{-1}$)
107	$-10^{\circ}C$	REFD	Reflectivity (dB)

TABLE ES12. The set of six experimental output diagnostics for the CLUE members contributed by CAPS and NSSL, which are output at hourly intervals.

No.	Level/layer	Parameter	Description
1	Surface	HAIL1	Max hail size from HAILCAST (mm)
2	Surface	HAIL2	Max hail size from Thompson method (mm)
3	3,000–0 m above ground	MXUPHL	Hourly max updraft helicity: 0–3 km AGL ($m^2 s^{-2}$)
4	5,000–2,000 m above ground	MNUPHL	Hourly min updraft helicity: 2–5 km AGL ($m^2 s^{-2}$)
5	3,000–0 m above ground	MNUPHL	Hourly min updraft helicity: 0–3 km AGL ($m^2 s^{-2}$)
6	1 km AGL	VVEL	Vertical velocity ($m s^{-1}$)

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