

1 **Impact of climate change and land use on groundwater salinization in**
2 **southern Bangladesh- implications for other Asian deltas.**

3 *M. A. Islam¹, M. A. Hoque², K. M. Ahmed³, and A. P. Butler⁴

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5 ¹Department of Oceanography, University of Dhaka, Bangladesh

6 Email: atik.ocn@du.ac.bd

7 ²School of Environment, Geography & Geosciences Sciences, University of Portsmouth, UK

8 Email: mo.hoque@port.ac.uk

9 ³Department of Geology, University of Dhaka, Bangladesh

10 Email: kmahmed@du.ac.bd

11 ⁴Department of Civil and Environmental Engineering, Imperial College London, UK

12 Email: a.butler@imperial.ac.uk

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16 **Abstract:**

17 **Pervasive salinity in soil and water is affecting agricultural yield and the health of millions**
18 **of delta dwellers in Asia. This is also being exacerbated by climate change through**
19 **increases in sea-level and tropical storm surges. One consequence of this has been a**
20 **widespread introduction of salt water shrimp farming. Here, we show, using field data**
21 **and modeling, how changes in climate and land use are likely to result in increased**
22 **salinization of shallow groundwater in SE Asian mega-deltas. We also explore possible**
23 **adaptation options. We find that possible future increase of episodic inundation events,**
24 **combined with salt water shrimp farming, will cause rapid salinization of groundwater**
25 **in the region making it less suitable for drinking water and irrigation. However, modified**
26 **land use and water management practices can mitigate the impacts on groundwater, as**
27 **well as the overlying soil, from future salinization. The study therefore provides guidance**
28 **for adaptation planning to reduce future salinization in Asian deltas.**

29 **Keyword:** Salinization; delta; land use; climate change; groundwater; shrimp farm

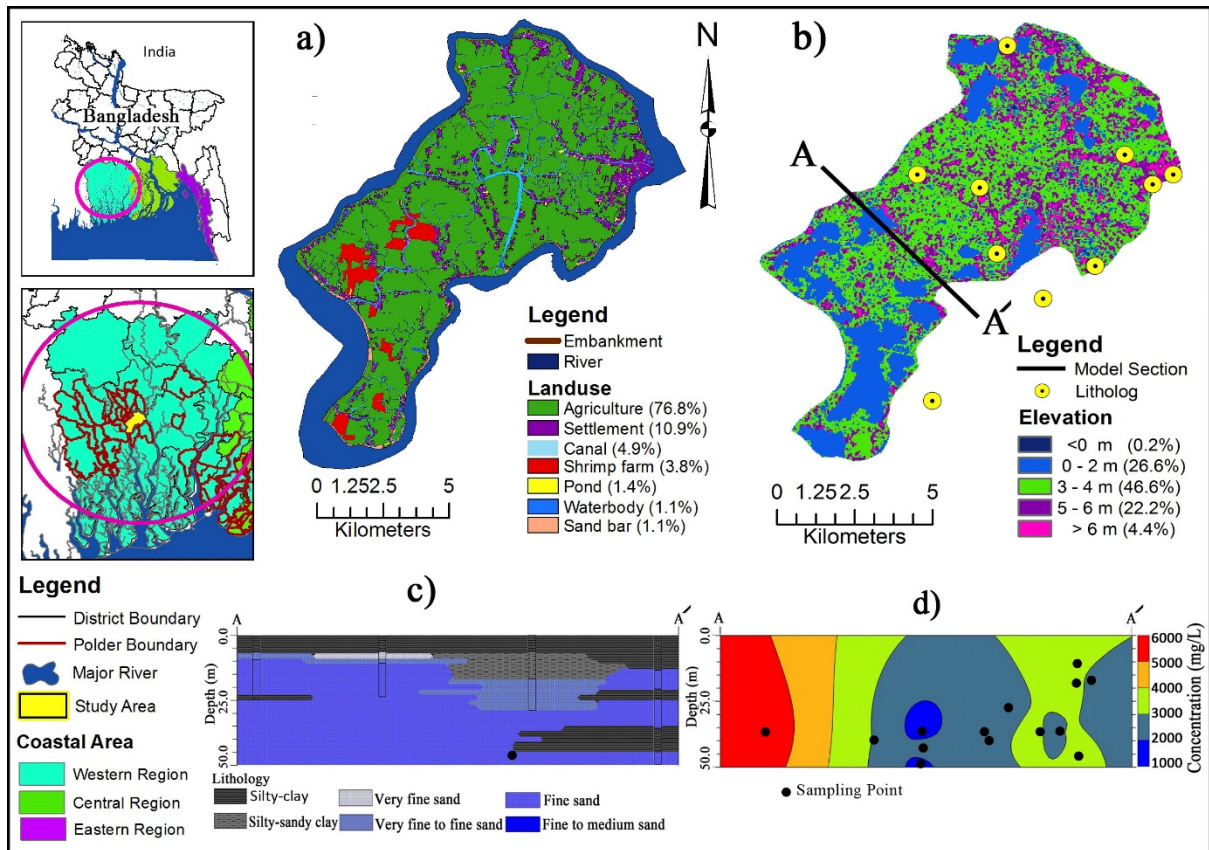
30 **1. Introduction**

31 Deltaic coastal plains all over the world are highly vulnerable to impacts from climate change,
32 such as: sea level rise, tidal flooding, storm surges. They are also subject to land subsidence,
33 coastal erosion and land loss, groundwater salinization, and environmentally damaging land
34 use practices (Ericson, et al., 2006, IPCC., 2007). There are eleven mega-deltas along the Asian
35 coastline (Milliman and Meade, 1983) which share similar geo-physical characteristics and are
36 home to millions of people (Woodroffe, et al., 2006). These are particularly vulnerable to
37 salinization, and future climate change is expected to exacerbate this (Hoque, et al., 2016).
38 Various Southeast Asian deltas show that salinity, particularly in surface water, has increased
39 and intruded inland over the last few decades (Mondal, et al., 2013, Noh, et al., 2013). High

40 levels of soil salinity affect 20-30% of net cultivable land, resulting in declines in agricultural
41 productivity (Baten, et al., 2015, Dasgupta, et al., 2015). Whilst, elevated salinity levels (i.e. >
42 1000 mg/L) in drinking water (Rahman, 2014, Worland, et al., 2015) have been shown to be
43 associated with prevalent hypertension (Scheelbeek, et al., 2017) - a risk factor for pre-
44 eclampsia in pregnant women (Khan, et al., 2011, Khan, et al., 2014) and cardio-metabolic
45 diseases in general, as well as increased rate of infant mortality (Dasgupta, et al., 2015). In
46 addition, high groundwater salinity combined with raised water tables in coastal areas is
47 causing increased maintenance costs for roads (Dasgupta, et al., 2014). Furthermore, as salinity
48 in coastal surface waters increases, this is putting many freshwater ecosystems under threat
49 (Herbert, et al., 2015).

50 The risk of groundwater salinization is likely to increase further due to rising sea levels, more
51 frequent and/or intense tropical cyclones and land use modification (e.g., Hoque, et al., 2016,
52 Rahman, et al., 2018). According to Singh (2002), the rate of sea level rise is much higher
53 along the Bay of Bengal (4 to 7.8 mm/year) than the global average of 1.8 ± 0.3 mm/year. Karim
54 and Mimura (2008) show that this will significantly widen the extent of future storm surge
55 flooding for similar magnitude cyclones. Understanding the impact of future changes in climate
56 and land use on the coastal hydrological system is important when prioritizing adaptation and
57 long-term planning. In this paper, we are taking the southern coastal zones of Bangladesh as a
58 study area (**Figure 1**) in order to address the above issues using field-based data and modeling.

59



60

61 **Figure 1:** Map of southern coastal zones of Bangladesh. Three broad coastal regions and their
 62 associated polders are shown. The study area is marked in yellow. a) Land-use map and b) elevation
 63 map of the study area, showing respectively land-use types and elevation ranges and their relative
 64 percentages. c) Lithology and d) distribution of groundwater salinity along cross section A-A'.

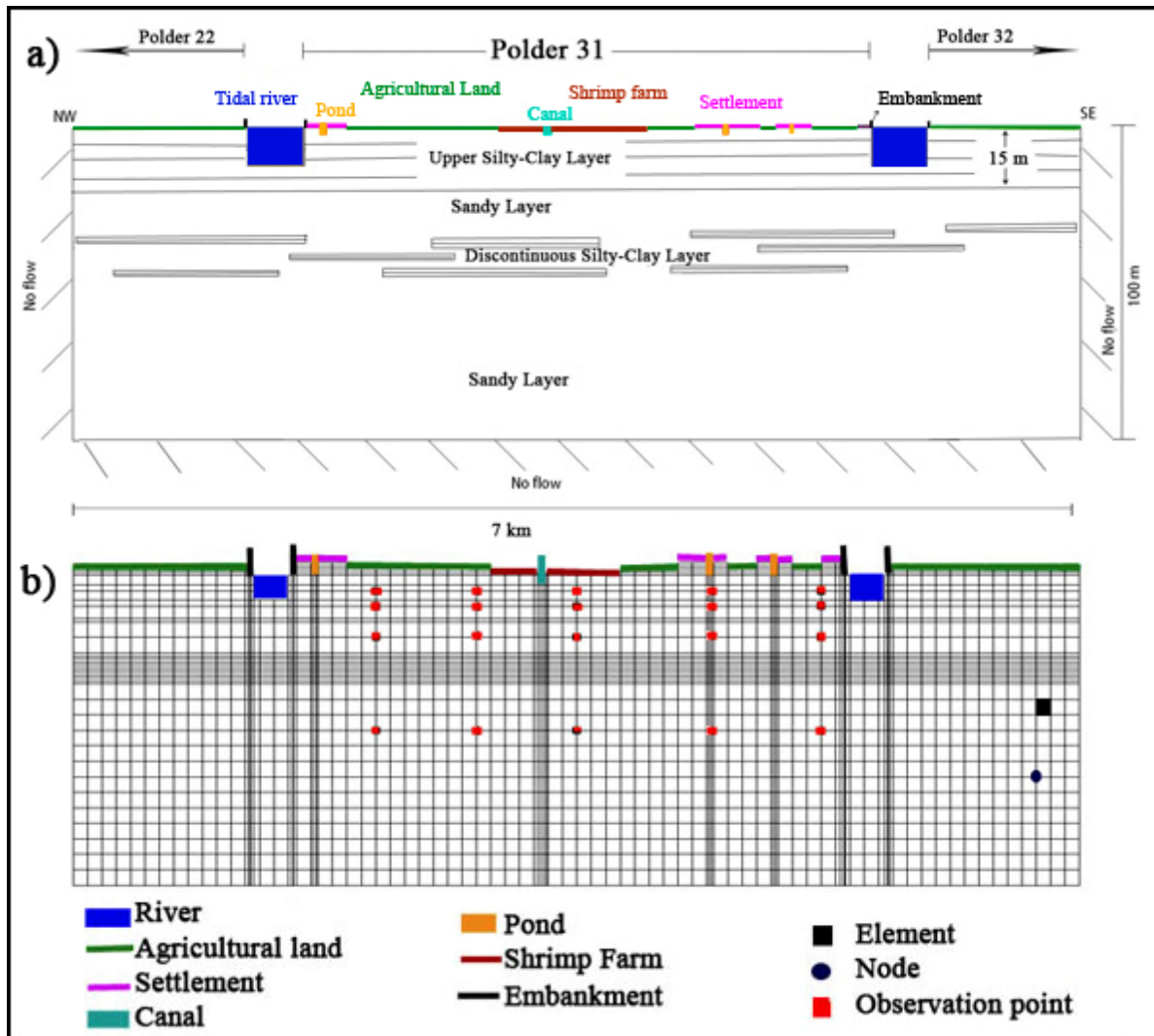
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66 The coastal region of Bangladesh is part of the Ganges-Brahmaputra Delta (GBD), the world's
 67 largest delta (Ericson, et al., 2006, IPCC., 2007), and most vulnerable to salinization among
 68 the Asian deltas (Hoque, et al., 2016). Like all Asian deltas, the GBD is regularly impacted by
 69 tropical cyclone induced storm surges and coastal flooding. Since the 1970s, these natural
 70 hazards, coupled with reductions in upstream river flows, have been causing widespread
 71 increases in soil and water salinity in southwest Bangladesh (Salehin, et al., 2018). The actual
 72 location and geometry of the saltwater freshwater interface in coastal Bangladesh is unknown
 73 but likely to be highly variable because of the complex geological structure, varied submarine

74 groundwater discharge rate (e.g., Werner et al 2013), and the impacts from storm inundation
75 events and monsoonal flushing. The numerical modelling indicates the saline wedge to be
76 extensive at shallower depths in the central coastal region because of infiltration from repetitive
77 storm surge induced inundation (Yu, 2010), and less from lateral seawater intrusion, which is
78 more dominant at deeper depths. The landward extend of this deeper interface is restricted by
79 the substantial submarine groundwater discharge (Faneca et al 2015), but could be affected by
80 rising sea level (Yu, 2010). With soil salinities becoming too great for rice production, local
81 farmers turned to farming salt-water shrimps in, an increasing number of, saline ponds. These
82 ponds are often extensive (*ca.* 0.3 to 1 km²) and normally maintain a persistent water column
83 of around 1 m deep throughout the year. The practice dramatically expanded during the 1990s
84 and currently the total area of shrimp farming in southern Bangladesh is around 2,000 km²
85 (Islam, et al., 2018).

86 The south-western and south-central coastal islands of Bangladesh are low elevation regions
87 that were enclosed by high (*ca.* 4 m) embankments to create a network of polders in the 1960s.
88 These act as a protection from regular spring high tides and consequently help increase food
89 security (Auerbach, et al., 2015, Rahman and Salehin, 2013). Sluice gates and culverts, located
90 at various places along the embankments, control water flow into and out of the polder areas.
91 Polder-31, with an area of 80 km², has been chosen as the specific study area (**Figure 1**) in the
92 southwest coastal region of Bangladesh. We analyzed satellite images for land use, lithological
93 data for hydrogeological characterization, and carried out field work to validate these analyses.
94 In addition, a reconnaissance field survey of groundwater salinity within the polder was also
95 conducted. We simulated the groundwater flow and salt mass transport along a 2D vertical
96 section across the polder, in order to investigate the effect of different land use types and
97 hydrological attributes (**Figure 2**) on the temporal behavior of groundwater salinity. This

98 provides a basis for assessing the long-term impacts of soil, water and groundwater salinity
 99 under various potential adaptation scenarios.



100

101 **Figure 2:** Groundwater model development. a) Conceptual model of the study area showing
 102 different land-use features, lithology and polder boundaries. b) SUTRA grid mesh showing different
 103 land-use types and observation points as they are simulated in the modelling.

104

105 2. Methods

106 2.1. Land-use

107 A land use map of the study area was produced using satellite images, dated 12 January 2015,
108 on Google Earth. Various land use classes were identified and geo-referenced and then merged
109 together in ArcGIS for further analysis and field validation. An elevation map was prepared
110 using SRTM 30 m DEM data (EROS, 2002).

111

112 **2.2. Hydrogeology and salinity**

113 Lithological data for 7 boreholes from Polder-31 were collected and 2 boreholes (Worland, et
114 al., 2015) drilled in Polder-32 were also included as these were in the neighboring region of
115 the studied polder. During the field survey the Electrical Conductivity (EC), as a proxy for
116 salinity, of various water sources were measured in the study area. The EC was measured from
117 14 tube-wells (Supplementary Information, hereafter SI, **Table S1**) to derive groundwater
118 salinity at various locations and depths. At each location a GPS position was recorded using a
119 Garmin handheld GPS, referenced to WGS 84 datum. Depth of the tube well screens were
120 noted by interviewing the well owner. The EC data were converted to total dissolved solids
121 (TDS) using a standard empirical relationship (i.e., multiplying the EC ($\mu\text{S}/\text{cm}$) by 0.5 to yield
122 a TDS value in mg/L (Freeze and Cherry, 1979)). Linear kriging interpolation was used to
123 generate the vertical salinity profile (Figure 1d).

124

125 **2.3. Model development**

126 The USGS open access computer code SUTRA 2.2 (which simulates variable density
127 groundwater flow and transport under saturated-unsaturated conditions (Voss and Provost,
128 2010) has been used to model a set of scenarios related to land use and climate change. The
129 graphical user interface ModelMuse 3.6.3 was used to facilitate input and output generation
130 (Winston, 2014) and Model Viewer 1.6 and GW Chart 1.29.0 were used to visualize the results

131 (Hsieh and Winston, 2002, Winston, 2000). Assumptions we made in the model development
132 are no unsaturated-flow, static boundary conditions, constant land use, and no pumping. The
133 dimensions of a vertical slice conceptual model were set as 7 km (width) by 100 m (depth).
134 Although the length of the study area is 5 km along the cross-section (**Figure 1b**), the model
135 was extended to 1 km outwards on each side to reduce the effect of the boundary conditions.
136 Similarly, although the main area of interest is the top 50 meters, the model was extended to
137 twice this depth in order to reduce any impact from the impermeable lower boundary. As the
138 upper part of the study area is covered by a thick silty-clay layer, a 15 m silty-clay layer was
139 implemented in the model. Beneath was an 85 m thick layer of fine-sand, within which was a
140 laterally discontinuous zone of silty-clay between the depths of 25 and 35 m (**Figure 2**). We
141 took this simplified geometry because lithological record from the area shows that geological
142 structure is less variable in the area (Figures 1c and S1 in SI). This supports our assumption
143 that the sediments have a 'layer-cake' stratigraphy and therefore it was reasonable to represent
144 this using a 2D model structure. In addition, a recent study from the delta demonstrated that
145 these laterally discontinuous silt-clay layers have an inter-leaved pattern, giving rise of similar
146 influence of 'layer-cake' stratigraphy on groundwater flow systems (Hoque et al 2017).

147

148 The various land use types were incorporated into the top 1 m grid layer of the model according
149 to their position on the cross-section, and their elevation was based on the SRTM derived
150 elevation data. The sizes of the rivers vary in the NW and SE boundary of the model and
151 therefore, the depth of the rivers was set at 8 m (NW) and 10 m (SE) below sea level,
152 respectively. Recharge on the land surface is generally controlled by net rainfall and was
153 generally implemented in the model as a time-dependent flux boundary. Those areas, however,
154 where there is standing water (e.g., Shrimp farm, drinking water pond, river) were represented
155 using time-dependent pressure heads and associated concentrations. No flow boundaries were

156 applied at the bottom and sides of the model. For the specified pressure boundary the following
157 equation was used.

158

159 $p(z) = p_0 - [\rho_b + \partial\rho/\partial C.C(z)]gz$ ----- (1)

160 where,

161 $p(z)$ = Hydrostatic fluid pressure [N/m²],

162 p_0 = Atmospheric pressure [N/m²]

163 ρ_b = Density of fresh water [kg/m³],

164 $\frac{\partial\rho}{\partial C}$ = Coefficient of fluid density change with concentration,

165 $C(z)$ = Concentration of fluid [kg/m³],

166 g = Gravitational acceleration [m/s²] and

167 z = elevation in reference to sea level [m], i.e., $z = 0$ is sea level

168 The entire model domain (7 km x 100 m) was divided into 3,268 elements and 3,444 nodes
169 using a fishnet mesh (Winston, 2015)(**Figure 2b**). The size of each general rectangular element
170 of the mesh has been kept to 100 m (horizontal) x 5 m (vertical). Elements situated above sea
171 level have a size of 100 m x 1 m. Moreover, elements representing different features have been
172 modified based on their size (**Table S2**). Each of the features has the same elevation.

173 The initial groundwater concentration in the model was set at $C_0=2,500$ mg/L, as groundwater
174 concentrations over most of the study area range from 2,000 mg/L to 3,000 mg/L (**Figure 1d**).

175 Time dependent river salinity was set using average seasonal values calculated from BWDB
176 (Bangladesh Water Development Board) data at the Chalna monitoring station (**Table S3**).

177 Starting from hydrostatic conditions, a steady state model was used to derive an initial pressure

178 field for the transient model, which was then run for 200 years, using 1 month time steps, to
179 assess long-term impacts.

180

181 **2.4. Parameterisation, validation and sensitivity analysis.**

182 In order to make the model implementation as simple as possible, parameter values were
183 derived from available data for the study area, supplemented by literature values (**Table S4**).
184 Some of the parameters, like recharge, water level, and water salinity are different for
185 different land use types and vary with time (these are tabulated in **Table S5**). Rainfall data for
186 the area, provided by the Bangladesh Meteorological Department (BMD), were obtained by
187 averaging monthly rainfalls for 3 sites in the vicinity of Polder-31 over the period 2000-2014.
188 River water level and salt concentration data were set from monthly average data for 2014
189 provided by the BWDB. Pond water levels and concentrations have been collected from data
190 given in Rahman et al. (2018). The model was not calibrated, as our objective was to use it in
191 an explanatory manner following the approach encouraged by Voss (2011a, b), however, an
192 extensive sensitivity analysis was undertaken in order to provide robust conclusions. Model
193 performance was assessed through a comparison of simulated hydraulic heads with those
194 monitored by the BWDB for Polder-31 (**Figure S1**).

195 The sensitivity of the base case model to different aquifer and fluid properties, i.e. porosity,
196 intrinsic permeability, dispersivity, and different side boundary conditions, was undertaken by
197 comparing head and salinity values at different locations in the model (**Table S6**). Among the
198 various parameters, groundwater head is slightly sensitive to changes in the side boundary
199 conditions, while salinity is most sensitive to the reduction of intrinsic permeability. The model
200 output, both in terms of head and salinity, was largely insensitive to the other parameters.

201

202 **3. Results**

203 **3.1. Land use and land form.**

204 The 2015 satellite data indicate that most (77%) of the land area for Polder-31 is agriculture,
205 whilst around 11% is habitable settlements/villages (**Figure 1a**). Each village has several
206 purpose-built ponds, which cover around 1% of the total land and are normally used for
207 drinking and other household purposes. Shrimp farming, around 4% of the total area, is limited
208 to the south-western part of the polder, where ground elevation is relatively lower. In other
209 polders, however, shrimp farming is more extensive (Johnson, et al., 2016) and is often the
210 dominant land use class. The variation in elevation over the area is shown in **Figure 1b**. The
211 land surface is generally below 6 m with only around 4% of the area having an elevation above
212 this. About a quarter (27%) of the land area has an elevation of less than 2 m, while (47%) lies
213 between 3 and 4 m and the remaining 22% is between 5 and 6 m. Areas having an elevation <
214 2 m are typically flooded during the monsoon and generally used for agriculture during the dry
215 season (Oct-Apr). Ponds and shrimp farms also are generally located in these low-lying areas
216 (**Table S7**). The areas with elevations between 3 and 4 m are mainly used for agriculture. Those
217 with elevations > 4 m are typically artificial constructions, such as the polder embankment,
218 communal settlements and local roads.

219

220 **3.2. Hydrogeology and salinity.**

221 The main near-surface aquifer is semi-confined. Lithological data show a thick upper silty-clay
222 layer, 5 to 15 m thick, and the presence of some discontinuous silty-clay layers (with
223 thicknesses of about 5 to 10 meters) at a depth between 25 and 35 m. In between these, there
224 is a sandy layer containing very fine to medium sands. A lithological cross-section (**Figure 1c,**
225 **and Figure S2 in SI**) indicates the laterally discontinuous nature of the silt-clay layers

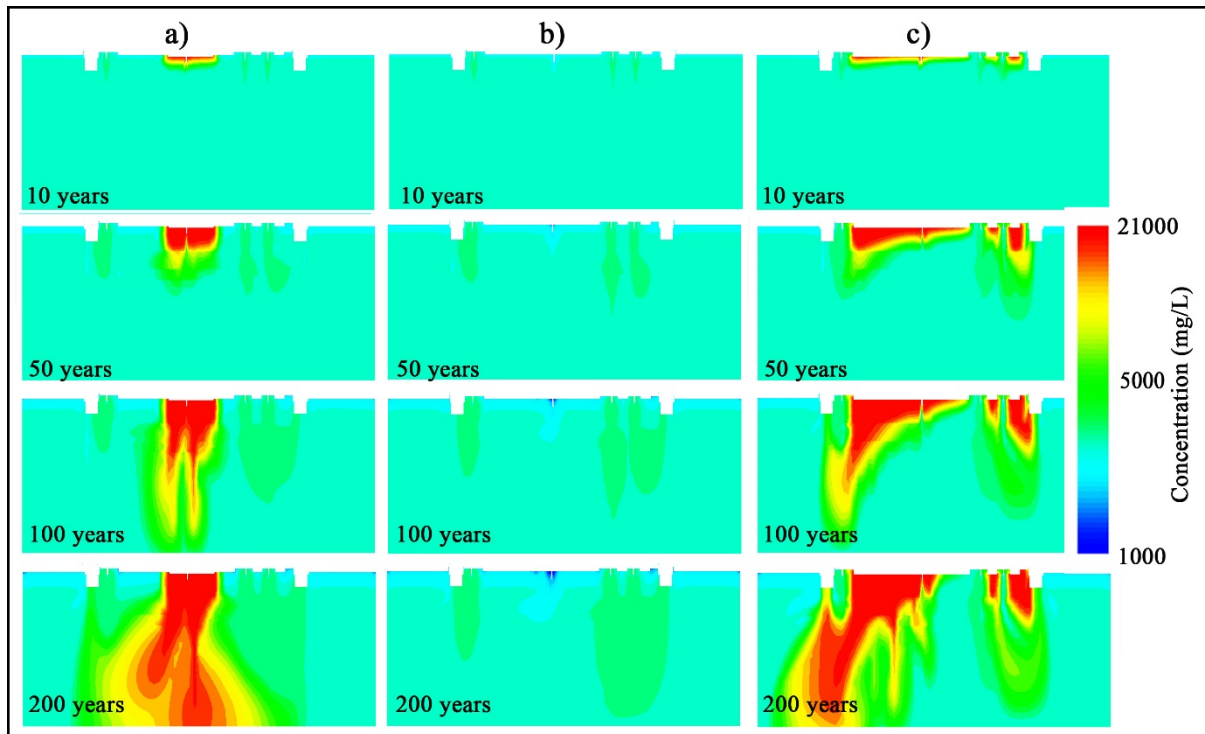
226 implying, at the local scale, a separation of this shallow aquifer from the relatively deeper
227 aquifer. The hydrogeological framework was corroborated with local drillers who confirm the
228 presence of numerous silt-clay layers embedded in the sandy aquifer materials. The drillers
229 also report that the silt-clay layers (between 30 and 35 m) are discontinuous, and are
230 intercalated with finer sands forming a regional hydrogeological feature. The salinity of the
231 groundwater typically ranges between 1,000 to 6,000 mg/L (Figure 1d). Wells near to
232 rivers/embankments have a relatively high concentration (ca. 5,000 mg/L). Regionally, shallow
233 groundwater is primarily brackish (ca. 3,000 mg/L) with isolated pockets of fresher (<1,500
234 mg/L) water, (Rahman, et al., 2012), which can be used for irrigation.

235

236

237 **3.3. Salinity and present land use condition.**

238 To explore the long-term relationship between salinity and land use, a set of modeling scenarios
239 were developed that consider different land use types found along the cross-section A-A'
240 (**Figure 1b**). Concentrations of salt (i.e., TDS) in groundwater after 10, 50, 100 and 200 years
241 are plotted (**Figure 3**). These show that, different land use patterns affect groundwater salinity
242 differently. Concentrations decrease slightly below agricultural land and habitable settlements
243 because of rain water recharge. As pond water has a relatively lower salinity compared to
244 groundwater, it also contributes to a decrease in salinity but the influence is limited to a small
245 area. The TDS of groundwater below and surrounding the shrimp farm shows a gradual
246 increase.



247

248 **Figure 3:** Effects of land-use on groundwater salinity for a) current land use, b) current land use, but
 249 without shrimp farms, c) all non-habitable land assumed to comprise shrimp farms. Distribution of
 250 TDS in groundwater after 10 years, 50 years, 100 years and 200 years.

251

252 **3.4. Modification of land use and salinity.**

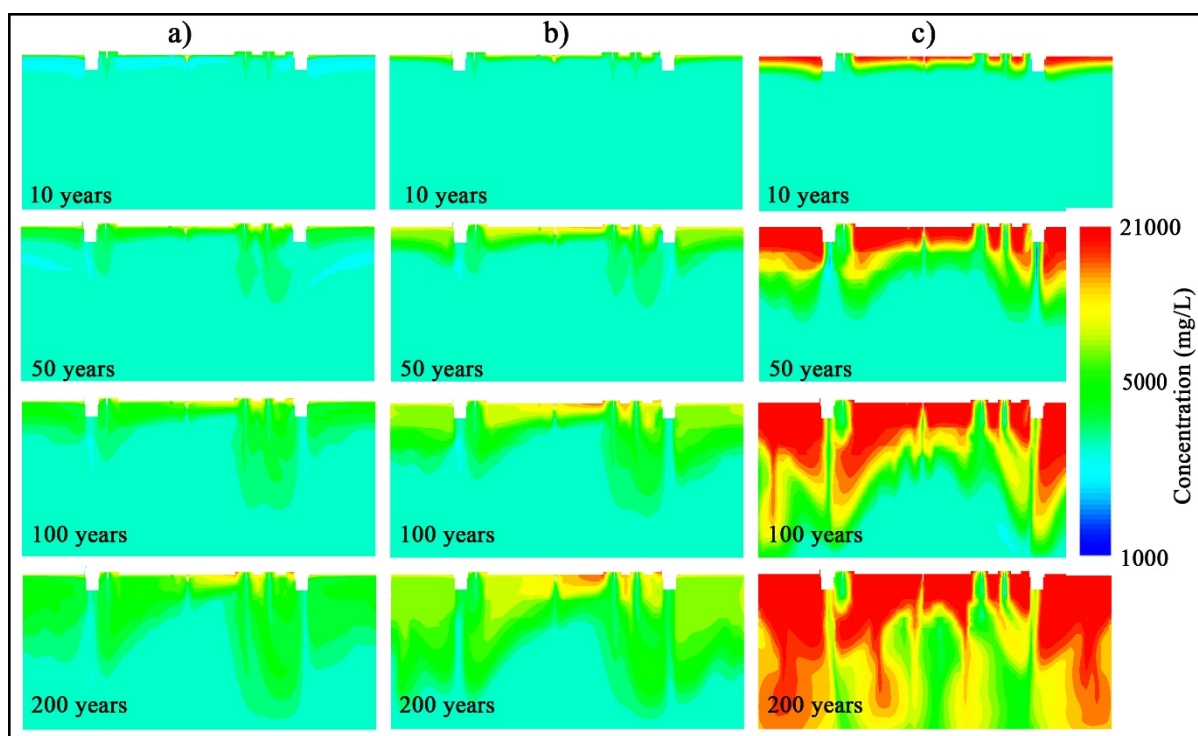
253 Shrimp farming in recent years has become less profitable due to disease outbreaks in salt water
 254 shrimps (Kabir, et al., 2015). As a result, traditional shrimp farming is decreasing in some
 255 coastal areas (Iqbal, et al., 2015), however, there are some indications of a reverse taking place
 256 in Polder-31 (see SI, **Figure S3**). In order to assess the range of potential impact on
 257 groundwater salinity from shrimp farming, scenarios were developed ranging from the case
 258 where there was no shrimp farming to that where all agricultural land has been converted to
 259 shrimp farms. In the case where there is no shrimp farming, it is seen that groundwater salt
 260 content is decreasing in most of the places (**Figure 3b**) apart from where there are drinking
 261 water ponds, which are assumed, based on field observations (Rahman, et al., 2018,

262 Scheelbeek, et al., 2017) to have variable TDS concentration ranging between 825 and 7,575
263 mg/L (**Table S5**). In addition, there is a slight increase in concentration below and surrounding
264 the canals/ivers. In the extreme case, where all agricultural land has been converted to shrimp
265 farming, it can be seen that this results in a progressive salinization of the underlying aquifers
266 over time (**Figure 3c**). The results indicate that prolonged salt water shrimp farming could lead
267 to a deterioration of groundwater quality by salinity over a time period of multiple decades.

268

269 **3.5. Salinity and frequency of cyclones.**

270 The study area is also affected by tropical cyclones on a regular basis; typically once every 3-
271 4 years (MoEF., 2009). However, although currently inconclusive, some studies have
272 suggested that the frequency and intensity of tropical cyclones could increase in the future due
273 to climate change (IPCC, 2014, Islam and Peter, 2009). On average, an intense cyclone
274 (categories 4 and 5) landfalls once every 10 years (Hoque, et al., 2016). We use the base case
275 model, but without any shrimp farms, to investigate the effect of tropical cyclone storm surge
276 frequency on groundwater salinity by considering recurrent intervals of storm surge inundation
277 of the Polder every 10 and 5 years and, as an extreme case, every year. We assume that, over
278 the relevant land area, the polder is flooded by saline river water (concentration = 28,000 mg/L)
279 to a depth of 1 meter for a period of one year. This is to represent the effect of regular saline
280 water flooding from a breached (and unrepaired) embankment (Mallick, et al., 2011). To
281 account for the effect of dilution from the seasonal monsoon, the model treats the salinity of
282 the infiltrating water as 80% sea water in the first year and decreasing at a rate 20% over the
283 next three years and by 10% in the fifth, according to the recurrence interval. The simulations
284 show groundwater salinity progressively increasing from the surface and extending downwards
285 over time (**Figure 4**). As it is assumed that a storm surge will flood the entire polder, more
286 frequent inundation has a major impact on groundwater salinity.



288

289 **Figure 4:** Effects of frequency of cyclone induced storm surges on groundwater for a) every 10 years
 290 b) every 5 years, c) every year for current land use without shrimp farming. TDS concentration in
 291 groundwater after 10 years, 50 years, 100 years and 200 years.

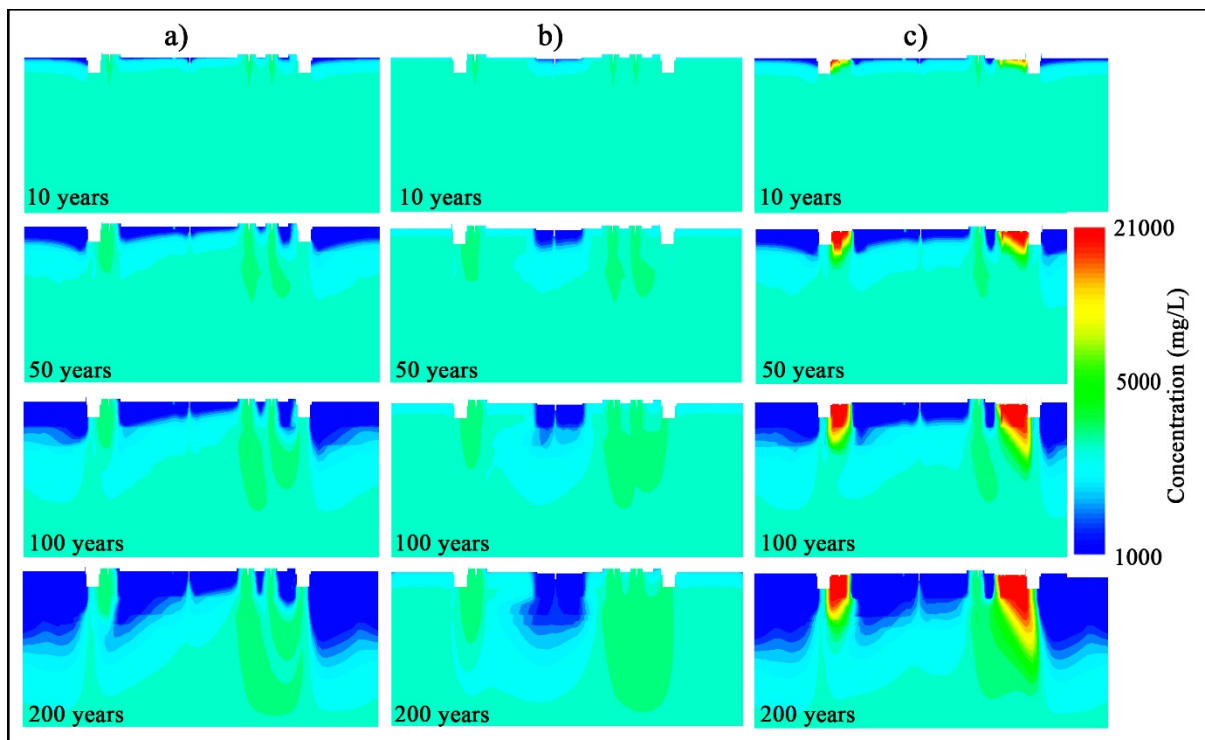
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293 **3.6 Adaptation options for salinity management.**

294 In view of the substantial impacts that both shrimp farming and storm surge inundation could
 295 have on groundwater and surface water salinity, and consequently on potential drinking water
 296 sources, the model was used to explore possible adaptation options (**Figure 5**). We considered
 297 three scenarios: i) current land use, without shrimp farms, inundated by fresh river water during
 298 the rainy season and protected during dry season through sluice gate controls; ii) replacing
 299 saline water shrimp farms with fresh water shrimp farms; and iii) saline shrimp farms located
 300 solely adjacent to the river, and therefore outside of the polder embankment, while the inner
 301 part of polder receives seasonal fresh river water inundation each rainy season. It is seen that,

302 if shrimp farms are removed and the polder area is seasonally flooded with fresher river water
 303 aquifers freshen up significantly. Replacing salt water shrimp farms with fresh water farms
 304 also improves the groundwater quality. In the third option, where salt water shrimp farms are
 305 located on the peripheral regions of the polder there is only limited salinization in these areas,
 306 whilst there is a general improvement in groundwater quality in the main area of the polder. It
 307 is important to note, however, that the modeling did not consider the effects of groundwater
 308 abstractions, which might facilitate transport of salt from peripheral zones to the central areas
 309 of the polder. On the other hand, this is not considered to be a major concern due to the limited
 310 use of groundwater pumping in Polder-31 (cf. Scheelbeek et al., 2017).

311



312

313 **Figure 5:** Simulation of various adaptation options. a) Inundated by freshwater during each rainy
 314 season and protected during dry season. b) Replacing saline water shrimp farming with fresh water
 315 shrimp farming, c) Shrimp farm only by the sides of the river while inner part of polder receives a
 316 seasonal freshwater inundation every rainy season.

317

318 **4. Discussion**

319 **4.1 Influence of land use and surface lithology on groundwater salinity**

320 The entire area of Polder-31 is covered by a thick silty-clay layer, with a thickness of 5 to 15
321 meters, below which is a very fine to fine sandy layer. An EM (Electromagnetic) survey in the
322 adjacent Polder-32 (Worland, et al., 2015) also shows the presence of an upper clay/silty-clay
323 layer of similar varying thickness. This topmost layer acts as an aquitard, making the
324 underlying shallow aquifer semi-confined, which is how it has been mapped in a previous
325 regional scale study (DPHE/DFID/JICA, 2006). Although it has been postulated that the
326 aquitard may protect the underlying aquifer from surface pollution (Yu, 2010), our results
327 indicate otherwise. This is also in agreement with small scale field observations (c. 0.05 km²)
328 made in Polders-31 and -32, which show salinity levels in the upper few meters of the silty-
329 clay unit that appear to be consistent with the types land use described above (Rahman et al.,
330 2018). One reason for this association could be the relatively permeable nature of the silty-clay
331 unit. Among the various parameters, intrinsic permeability is found to be the most sensitive for
332 salinity; therefore, a lower permeability may slow the process but not stop it. The vertical
333 hydraulic conductivity of which Rahman et al. (2018) found to be c. 0.03 m/day. This value is
334 consistent with the intrinsic permeability values used in the current model simulations (see
335 **Table S4**).

336 In light of these results, we suggest that surface land use features, and their associated water
337 salinities, can play an important role in the salinization of groundwater in coastal polders.
338 Different land uses, such as agriculture, human settlement and drinking water ponds are sources
339 of fresh water recharge in such aquifers, which contribute to the freshening of the aquifer over
340 time. In contrast, salt water shrimp ponds, and any drainage canals affected by these (i.e. via

341 sluice gates), have relatively high salinities, whose higher relative density results in convective
342 plumes of downward migrating saline water (Xie, et al., 2011) causing increased salinization
343 of the underlying aquifer (**Figure 3**). Our results indicate that the effect of shrimp farms on
344 groundwater could be several times higher than any other kind of land use. This assessment
345 indicates that shrimp farming, if used extensively, could lead to enhanced salinization of
346 groundwater and a deterioration of near surface aquifers as a water resource.

347

348 **4.2 Influence of climate change on groundwater salinity**

349 Our results indicate that climate change impacts, such as sea-level rise, changes in precipitation
350 patterns, and more frequent and/or intense storm surges could lead to a deterioration in
351 groundwater quality. The study area is frequently affected by tropical cyclones (Ali, 1996,
352 Hoque, et al., 2016, Islam and Peterson, 2008), some of which can over-top and/or breach
353 polder embankments leading to inundation with saline water (Hoque, et al., 2016, IMD, 2009,
354 Mallick, et al., 2011). This inundated saline water can then infiltrate into the subsurface causing
355 groundwater salinization, which will continue until the embankment breaches are fixed and the
356 saline surface water is drained out of the polder. It is anticipated that with climate change there
357 might be higher frequencies of storm surge events capable of breaching or over-topping polder
358 embankments and this is likely to lead to additional salinization of the underlying aquifer
359 (Islam and Peter, 2009). In addition, sea level rise can modify the dynamics of tropical cyclones
360 in this region and further increase the risk of flooding and bank erosion (Karim and Mimura,
361 2008).

362 It has been projected that Bangladesh rainfall will increase by 12% in the monsoon and
363 decrease by 3% in winter by 2100 (Agrawala, et al., 2003). However, our results appear to
364 show that these changes in rainfall would have minimal effect on groundwater salinity (**Figure**

365 **S4**), as recharge is around 10-15% of the total rainfall (Shamsudduha, et al., 2011), and
366 therefore only represents a minor change.

367

368 **4.3 Salinity management options**

369 Our results imply that both land use and climate change can have an impact on groundwater
370 salinization. In some cases groundwater salinity has been found to decrease for a given land-
371 use (e.g., rice cultivation using rainy season flood irrigation, **Figure 5a**) or increase for other
372 (e.g., installing shrimp farm, **Figure 3c**). Land use planning and hydrological management
373 scheduling may significantly help in combatting soil and groundwater salinization in these
374 regions. Embankments have a great influence on the groundwater salinity, as these were made
375 to protect the enclosed areas from regular tidal flooding and tropical cyclone storm surges. Sea
376 water inundations of any sort are a potential source of saline water recharge and thus
377 responsible for increasing groundwater salinity. In our simulations, we have considered how
378 controls on the flow of water in and out of a polder via sluice gates placed within the polder's
379 embankments can be used as a salinity management tool. We see that, if the embankments
380 prevent tidal water to enter into the polder, salinization of groundwater decreases remarkably.
381 Furthermore, if saline river water intrusion into the polder is prevented during the dry season
382 but the ingress of fresher river water (i.e. when salinity is <300 mg/l), via sluice gates installed
383 within the embankment, is allowed during the rainy season (Auerbach, et al., 2015, Willcocks,
384 1930) then over time the aquifer will become fresher (**Figure 5**). In addition, this open tidal
385 hydrological regime during the rainy season will also allow the polder to gain fertile sediments
386 and help it to promote elevation equilibrium between the inside and outside of the polder, as
387 suggested by Auerbach et al. (2015). According to Kibria, et al. (2015) agricultural land in
388 these coastal polders, utilising rainy season flood irrigation for rice, has increased from 22% to
389 31% post Aila (2009). We suggest that this practice of the ingress of fresher river-water for

390 rice cultivation during the rainy season will be good for both the natural environment and the
391 local resource economy.

392

393 **5. Conclusion**

394 The study was done in Bangladesh in order to explore the effects of land use and climate change
395 on groundwater salinization. In this study, we consider that the understandings are applicable
396 to other Asian deltas too. The results will facilitate to identify knowledge gaps and design
397 appropriate study for management decisions. The study shows that land use management that
398 has resulted in ponding of saline water at the surface is mainly responsible for the salinization
399 of underlying groundwater. In contrast, we have shown that there are adaptation measures that
400 policy makers can adopt to reduce groundwater salinization (such as relocating shrimp farms,
401 encouraging land use for agricultural purposes, and particularly allowing the polder area to be
402 part of tidal plain during the rainy season). We also find that the embankments are an essential
403 landscape management solution in the face of climate change and sea level rise to protect the
404 polder area from the ingress of saline water, and by allowing the polder area in rainy season to
405 be part of tidal plain when river water is fresher.

406 **Acknowledgement**

407 We thank Managed Aquifer Recharge (MAR) project team at the University of Dhaka for
408 sharing the lithological data, and Bangladesh Water Development Board (BWDB) for
409 providing the water level and salinity data, and Bangladesh Meteorological Department (BMD)
410 for rainfall data. This work was funded by The Leverhulme Trust (grant no. RPG-314) and The
411 Wellcome Trust (Institutional Strategic Support Fund: Networks of Excellence Scheme 2014),
412 whose support is gratefully acknowledged.

413

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