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1 **Enhanced levels of atmospheric low-molecular weight monocarboxylic acids in gas and**
2 **particulates over Mt. Tai, North China, during field burning of agricultural wastes**

3

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18 Highlights:

19 We determined gaseous and particulate C₁-C₁₀ monocarboxylic acids over Mt. Tai.

20 Formic (C₁) and acetic (C₂) acids were dominant followed by lactic acid.

21 High concentrations of C₁ and C₂ were observed during field burning of wheat straw.

22 High particle-phase fractions of C₁ (0.50) and C₂ (0.31) were observed.

23 Field burning of agricultural wastes is an important source of monocarboxylic acids.

24

25 Keywords: Acetic acid, Formic acid, Lactic acid, Gas/particle partitioning, Wheat straw
26 burning, North East China

27 **Abstract**

28 To understand the source and atmospheric behaviour of low molecular weight
29 monocarboxylic acids (monoacids), gaseous (G) and particulate (P) organic acids were
30 collected at the summit of Mt. Tai in the North China Plain (NCP) during field burning of
31 agricultural waste (wheat straw). Particulate organic acids were collected with neutral quartz
32 filter whereas gaseous organic acids were collected with KOH-impregnated quartz filter.
33 Normal (C₁-C₁₀), branched (iC₄-iC₆), hydroxy (lactic and glycolic), and aromatic (benzoic)
34 monoacids were determined with a capillary gas chromatography employing
35 *p*-bromophenacyl esters. We found acetic acid as the most abundant gas-phase species
36 whereas formic acid is the dominant particle-phase species. Concentrations of formic (G/P
37 1570/1410 ng m⁻³) and acetic (3960/1120 ng m⁻³) acids significantly increased during the
38 enhanced field burning of agricultural wastes. Concentrations of formic and acetic acids in
39 daytime were found to increase in both G and P phases with those of K⁺, a field-burning
40 tracer ($r = 0.32-0.64$). Primary emission and secondary formation of acetic acid is linked with
41 field burning of agricultural wastes. In addition, we found that particle-phase fractions (F_p
42 $=P/(G+P)$) of formic (0.50) and acetic (0.31) acids are significantly high, indicating that
43 semi-volatile organic acids largely exist as particles. Field burning of agricultural wastes may
44 play an important role in the formation of particulate monoacids in the NCP. High levels (917
45 ng m⁻³) of particle-phase lactic acid, which is characteristic of microorganisms, suggest that
46 microbial activity associated with terrestrial ecosystem significantly contributes to the
47 formation of organic aerosols.

48 **1. Introduction**

49 Low molecular weight (LMW) monocarboxylic acids (monoacids) are known to
50 exist in the atmosphere as gases, aerosols, cloud, fog, and snow (Kawamura et al., 1985;
51 Chebbi and Carlier, 1996; Kawamura et al., 2012). Formic and acetic acids are dominant
52 organic species in the atmosphere. Concentrations of formic and acetic acids in gas phase are
53 higher than those in particle phase (e.g., Andreae et al., 1988, Liu et al., 2012) due to their
54 high vapour pressures. Global gaseous formic acid emission is estimated to be 100-120 Tg
55 year⁻¹ (Stavrakou et al., 2011), where formic acid comprises about 10% of the total emissions
56 of non-methane volatile organic compounds. However, there is a significant missing source of
57 organic acids. In particular, there is a limitation on our knowledge on how LMW monoacids
58 exist in particles and what parameters control the gas/particle partitioning of those organic
59 acids.

60 Particulate formic and acetic acids are water-soluble and thus play an important role
61 in the balance of radiative forcing acting as cloud condensation nuclei (CCN) (Yu, 2000).
62 Gaseous formic and acetic acids are adsorbed on the existing alkaline particles. Möhler et al.
63 (2008) reported that ice nuclei (IN) efficiency of mineral dust particles is decreased by the
64 adsorptive uptake of organics on the particle surface. However, effects of organics on CCN
65 and IN largely depend on their chemical composition or polarity. The organic aerosols that
66 are enriched with organic acids contribute to the formation of CCN and IN and thus
67 precipitation process, affecting the global water cycle and climate changes. On the other hand,
68 high abundances of LMW organic acids in the atmosphere can adversely affect air quality and
69 human health and also increase the acidity of rainwater (Keene et al., 1983, 1989; Kawamura
70 et al., 1996).

71 Previous studies have shown that LMW monoacids are derived from both primary
72 and secondary sources. LMW monoacids are directly emitted from fossil fuel combustion
73 (Kawamura et al., 2000), vegetation (Kesselmeier and Staudt, 1999), soils (Asensio et al.,

74 2008), and ocean microorganisms (Miyazaki et al., 2014). The secondary sources include
75 photooxidation of biogenic and anthropogenic organic compounds (Kawamura et al., 2000;
76 Lim et al., 2005; Paulot et al., 2011; Brégonzio-Rozier et al., 2015). In addition, biomass
77 burning is an important source of formic and acetic acids in gas and aerosol phases (Andreae
78 et al., 1988; Kesselmeier and Staudt, 1999). However, gas/particle partitioning of formic and
79 acetic acids emitted from biomass burning has rarely been studied. Further, there is no study
80 on gas/particle partitioning of normal (C_3 – C_{10}), branched chain (iC_4 – iC_6), and hydroxy (lactic
81 acid) monoacids in the atmosphere.

82 East China is one of the most polluted regions in the world, where many mega cities
83 such as Beijing and Shanghai are located. Our previous studies reported that concentrations
84 of levoglucosan, a biomass burning tracer (Fu et al., 2008), dicarboxylic acids (Kawamura et
85 al., 2013a), and α -dicarbonyls (Kawamura et al., 2013b) significantly increased over the
86 summit of Mt. Tai in the North China Plain (NCP) during the field burning of agricultural
87 wastes. The inflow of those polluted air masses may adversely affect the air quality even in
88 urban areas such as Beijing and outflow regions in East Asia and the western North Pacific.

89 In this study, we collected gas and particle samples at the summit of Mt. Tai during
90 June 2006 to better understand the distributions of LMW monoacids during the field burning
91 of agricultural residue. We utilized two-stage filter pack sampling technique (see the
92 experimental section) followed by derivatization to *p*-bromophenacyl esters to determine gas
93 and particulate monoacids simultaneously (Kawamura et al., 1985). Further, utilization of a
94 capillary gas chromatograph (GC) provides very high resolution and sensitivity and hence
95 relatively minor species such as propionic acid (C_3) and C_4 – C_{10} monoacids, which are not
96 detectable by ion chromatography, can be detected by GC method (Kawamura et al., 1985).
97 By modifying the method of Kawamura et al. (1985), we determined hydroxy (lactic and
98 glycolic) acids together with normal (C_1 – C_{10}), branched (iC_4 – iC_6), and aromatic (benzoic)
99 monoacids in both gas and aerosol phases. Although this technique has been used for air and

100 rainwater samples from Los Angeles (Kawamura et al., 2000, 2001), its application has not
101 been conducted to air samples from Asian regions. Here we report, for the first time, gaseous
102 and particulate LMW monoacids that are most likely derived from the field burning of
103 agricultural wastes. We also discuss gas/particle partitioning of monoacids during the field
104 burning of wheat straws in the NCP and potential contribution of microbial activity to organic
105 aerosols.

106

107 **2. Experimental**

108 Sampling of gaseous and particulate LMW monoacids was conducted at the top of Mt.
109 Tai (36.26° N, 117.11° E, elevation: 1534 m above sea level) located in Shandong Province,
110 North China (Figure 1) as part of the Mount Tai Experiment 2006 campaign (MXT2006)
111 (Kanaya et al., 2013). Sample collections were performed every 3 hours (00:00–03:00,
112 03:00–06:00, 06:00–09:00, 09:00–12:00, 12:00–15:00, 15:00–18:00, 18:00–21:00, 21:00–
113 24:00 LT) on June 2–5 and 23–25, 2006. Totally 36 sets of samples (72 filters) were collected
114 for both particulate and gaseous LMW monoacids employing neutral and KOH-impregnated
115 quartz fibre filters, respectively. Here, we compare the atmospheric abundances of gaseous
116 and particulate organic acids between more and less field-burning periods. The weather
117 conditions were sunny or cloudy during the campaign. Meteorological parameters such as
118 ambient temperature, relative humidity, wind direction, and wind speed were measured at the
119 Mt. Tai Observatory and reported previously (Kanaya et al., 2013; Kawamura et al., 2013a).
120 The height of planetary boundary layer (PBL) are provided by the Global Data Assimilation
121 System (GDAS1) (Air Resources Laboratory, NOAA). Average PBL in daytime and
122 nighttime were about 2200 m and about 600 m, respectively. Daytime samples were collected
123 within the PBL and nighttime samples were collected in the free troposphere during the
124 sampling period.

125 To collect ambient monoacids in gas and particle phases, a low-volume air sampler

126 equipped with the two-stage filter packs (URG-2000-30FG) were set in a series at the
127 balcony of two-story building of the observatory (height: 10 m) near the summit of Mt. Tai.
128 Total particulate monoacids were collected on precombusted (450 °C, 6 h) neutral
129 quartz-fiber filter (47 mm diameter) (first stage) whereas gaseous monoacids were collected
130 on the quartz-fiber filter impregnated with potassium hydroxide (KOH) (second stage) at an
131 airflow rate of 8.4 L min⁻¹. The KOH impregnated filters were prepared by rinsing the
132 precombusted quartz filter in a 0.2 M KOH solution and then dried in an oven at 80 °C. The
133 details of filter preparation and analytical method were described previously (Kawamura et
134 al., 1985). Both the neutral and alkaline filters were individually placed in clean glass vials
135 (50 ml) with a Teflon-lined screw cap to avoid a gas exchange with ambient air before and
136 after the sampling. After the sampling, filter samples were transported to the lab in Sapporo
137 and stored in a freezer room at -20 °C prior to analysis.

138 An aliquot of filter (4.3 cm²) was extracted with organic free ultrapure water (5 ml ×
139 3, > 18.2 MΩ cm) under ultrasonication. To remove particles and filter debris, the extracts
140 were passed through a Pasteur pipette packed with quartz wool. To avoid the evaporative loss
141 of volatile monoacids during analytical procedures, the pH of water extracts was adjusted to
142 8.5-9.0 with 0.05 M KOH solution to form organic acid salts (e.g., CH₃COO⁻K⁺). The organic
143 acid salts were concentrated using a rotary evaporator under vacuum at 50 °C. To make all
144 the organic acid salts to RCOO⁻K⁺ form, the concentrates were passed through a cation
145 exchange resin (DOWEX 50W-X4, 100-200 mesh, K⁺ form) packed in a Pasteur pipette. The
146 eluents were finally checked for pH=8.5–9.0 and then concentrated using a rotary evaporator
147 under vacuum. The dried RCOO⁻K⁺ salts were derivatized in acetonitrile (4 ml) with
148 *α,p*-dibromoacetophenone (0.1 M, 50 μl) as a derivatization reagent and
149 dicyclohexyl-18-crown-6 (0.01 M, 50 μl) as a catalyst at 80 °C for 2 hours (Kawamura and
150 Kaplan, 1984).

151 The derivatives (*p*-bromophenacyl esters) of monoacids including normal (C₁–C₁₀),

152 branched (iC₄–iC₆), and aromatic (benzoic acid) structures were measured using a capillary
153 gas chromatograph (GC) (HP GC6890, Hewlett-Packard, USA) equipped with a flame
154 ionization detector. The GC peaks were identified using a GC-mass spectrometer (GC/MS)
155 (Agilent GC7890A and 5975C MSD, Agilent, USA). The mass spectra of *p*-bromophenacyl
156 esters of monoacids are presented in Kawamura and Kaplan (1984).

157 The OH functional groups in *p*-bromophenacyl esters of organic acids were further
158 reacted with N,O-bis-(trimethylsilyl)trifluoroacetamide (BSTFA) with 1% trimethylsilyl
159 chloride and 10 µl of pyridine at 70 °C for 3 hours to derive trimethylsilyl (TMS) ethers of
160 *p*-bromophenacyl esters. The derived TMS ethers were determined with a capillary GC and
161 GC/MS. We determined hydroxymonocarboxylic acids including lactic and glycolic acids.
162 Details of the methods have been described in Kawamura et al. (2012).

163 A mixture of organic acids was prepared in our laboratory from the authentic
164 standards: C₁–C₇ and iC₄–iC₆ (10 mM in water, assay 98%, Sigma-Aldrich), lactic acid (assay
165 >85%, Wako), glycolic acid (assay 97%, Wako), and benzoic acid (assay 99%, Tokyo
166 Chemical Industry Co.). To check the recoveries of organic acids during the analytical
167 procedures, authentic monoacid standards (C₁–C₁₀, benzoic, lactic, and glycolic acids) were
168 spiked to KOH-impregnated quartz filters and the filters were analysed as a real sample. The
169 results showed that the recoveries were more than 80%. Analytical errors using authentic
170 standards were within 12%. Detection limits of organic acids in our analytical method were
171 ca. 0.02 ng m⁻³. Filed blank filters were prepared by the same sampling procedures without
172 pumping air. We found that blank levels of monoacids were less than ~5% of those of
173 ambient samples in both gaseous and particulate phases. In addition, we analyzed laboratory
174 blanks for monoacids, in which the same analytical protocols were used. The blank levels
175 were less than ~5% of those of ambient samples. The concentrations reported here are
176 corrected against the field and laboratory blanks, but not corrected for the recoveries.

177 Particulate organic acids collected on the neutral filter may be in part evaporated

178 during sampling and be collected on the second filter (KOH-impregnated). Thus, particulate
179 organic acids could be underestimated and gaseous organic acids could be overestimated.
180 However, we believe that evaporative loss should be minimal based on the experimental
181 results from Los Angeles (Kawamura et al., 1985, 2000). Samplings of total suspended
182 particles over the Mt. Tai and Los Angeles were conducted by the same sampling system.
183 Particle phase (TSP) concentrations of organic acids from Mt. Tai (C_1 : 1410 ng m⁻³, C_2 : 1120
184 ng m⁻³) are higher than those (C_1 : 163 ng m⁻³, C_2 : 120 ng m⁻³) from Los Angeles (Kawamura
185 et al., 2000). Particulate organic acids collected on a neutral quartz filter (first stage) have
186 been reported to be less than ~17% of total (gas + particulate) organic acids in Los Angeles
187 (Kawamura et al., 1985). We found that particulate monoacids are sometimes more abundant
188 than gaseous monoacids in this study, which may be in part associated with lower ambient air
189 temperature on the top of Mt. Tai (12-20 °C, average 16°C) than in Los Angeles (summer)
190 and other parameters as discussed later.

191 Keene et al. (1989) showed that alkaline (NaOH) impregnated filter with glycerol
192 (C₃H₈O₃) results in positive interferences of formic and acetic acids. Concentrations of formic
193 and acetic acids in gas phase may be overestimated in the previous studies where glycerol
194 was used because glycerol may be reacted with oxidants to result in formic and acetic acids.
195 However, we did not use glycerol in this study. Gaseous monoacids may be subjected to
196 atmospheric titration by alkaline dust particles (Ca, Na, K and Mg) in the atmosphere.
197 Organic acids in aerosols may exist in the form of salts such as RCOOK, RCOONa, and
198 others. Vapor pressures of those salts are significantly lower than those of free monoacids. In
199 this study, aerosol samples were collected to differentiate free organic acids and organic acid
200 salts using KOH-impregnated and neutral quartz filters, respectively. We calculated ion
201 balance in the aerosols over Mt. Tai (Please see discussion section). Total cations are slightly
202 higher than total anions. We think that absorption of gas phase organic acids in aerosol phase
203 on the first filter is limited.

204 For the measurements of inorganic ions (cations: Na^+ , NH_4^+ , K^+ , Mg^{2+} , and Ca^{2+} and
205 anions: F^- , MSA^- , Cl^- , NO_2^- , NO_3^- , Br^- , PO_4^{3-} , and SO_4^{2-}), ambient aerosols were collected
206 onto quartz-fiber filter (20×25 cm) (Tissuquartz 2500QAT-UP, Pallflex, USA) using a
207 high-volume air sampler during the same sampling period. A portion of filter was extracted
208 with ultrapure water under ultrasonication. The water extracts were filtered through a
209 membrane disk filter ($0.22 \mu\text{m}$, Millipore Millex-GV, Merck, USA) and then analyzed using
210 an ion chromatograph (Model 761 compact IC, Metrohm, Switzerland) (Miyazaki et al.,
211 2009). Anion analysis was conducted using a Shodex SI-90 4E column and a 1.8 mM
212 $\text{Na}_2\text{CO}_3 + 1.7 \text{ mM NaHCO}_3$ solution as an eluent. Cation analysis was performed using a
213 C2-150 column and a 4.0 mM tartaric acid plus 1.0 mM dipicolinic acid solution as an eluent.
214 The analytical errors in duplicate analyses of samples were within 4%.

215 Carbon monoxide (CO) concentration was continuously monitored with a non
216 dispersive infrared gas analyzer (Model 48C, Thermo Scientific, USA). The CO data were
217 obtained from Kanaya et al. (2009).

218 Seven-day backward and forward air mass trajectories were calculated over the
219 summit of Mt. Tai at a level of 1500 m and 750 m a.s.l. using the Meteorological Data
220 Explorer (METEX) from the National Institute for Environmental Studies, Tsukuba, Japan
221 (<http://db.cger.nies.go.jp/metex/index.html>). Meteorological data were obtained from the
222 National Centers for Environmental Prediction (NCEP) Reanalysis data. In addition, we have
223 checked weather conditions in Mt. Tai, Beijing, and Shanghai. Weather information were
224 obtained from the National Oceanic and Atmospheric Administration. Extreme weather
225 events did not occur during the period 28 May to 30 June 2006. Fire spot data sets were
226 downloaded from MODIS fire spot measurement data
227 (<http://earthdata.nasa.gov/data/near-real-time-data/firms>).

228

229 3. Results

230 **3.1. Characteristics of air mass and fire spots**

231 Figure 2 shows typical ten-day air mass back trajectories (1500 m a.s.l.) with fire
232 spots. In the North China Plain (NCP), harvest season of wheat starts in late May and then
233 agricultural wastes (wheat straw) are burned in the open field. The field burning was heavily
234 observed from the satellite over the NCP from 2 to 4 June, especially, in the south of Mt. Tai.
235 The field burning declined during 23–25 June with a shift of the hot spot areas to the north
236 and northwest of the sampling site. The air masses at a levels of 1500 m and 750 m (Please
237 see Figure A1) during 3–4 June have passed over the heavy fire spot regions. These days are
238 strongly influenced by the agricultural waste burning. On the other hand, the air mass (1500
239 m and 750 m) of 2 June was mainly originated from the Sea of Japan. The air masses at a
240 level of 1500 m on 23 and 25 June were transported from the NCP and South China,
241 respectively. Although the air masses at a level of 1500 m on 23 and 25 June were not
242 consistent with those of 750 m, the air masses of 2, 23 and 25 June were less affected by the
243 field burning of agricultural wastes. In addition, concentrations of K^+ , a good tracer of field
244 burning (Andreae, 1983), increased on 3–5 June and decreased on 23 and 25 June (Figure 3).
245 These results are consistent with the spatial distributions of fire spots and air mass trajectories
246 arriving over the sampling site. Hence, the air masses of 3–4 June were primarily delivered
247 from the areas of the field burning of agricultural residue. In this study, the data were
248 categorized into two periods: more field-burning (MFB) (3–5 June) and less field-burning
249 (LFB) (2 and 23–25 June) influenced periods.

250

251 **3.2. Gaseous and particulate LMW monoacids**

252 We identified LMW monoacids; such as normal (formic: C_1 , acetic: C_2 , propionic:
253 C_3 , butyric: C_4 , pentanoic: C_5 , hexanoic: C_6 , heptanoic: C_7 , octanoic: C_8 , nonanoic: C_9 , and
254 decanoic: C_{10}), branched chain (isobutyric: iC_4 , isopentanoic: iC_5 , and isohexanoic: iC_6),
255 hydroxy acids (lactic: Lac, and glycolic: Glyco), and aromatic (benzoic: Benz) acids in both

256 gas and particle phases. Figure 4 shows average molecular compositions of detected LMW
257 monoacids in the MFB and LFB periods. Acetic acid was the dominant species in gas phase
258 in the MFB (61%) and LFB (41%) periods, followed by formic acid (MFB: 24%, LFB: 29%)
259 and lactic acid (MFB: 5%, LFB: 14%). In particle phase, formic acid was the dominant
260 species in the MFB (34%) and LFB (31%) periods, followed by acetic acid (MFB: 27%,
261 LFB: 27%) and lactic acid (MFB: 22%, LFB: 23%). In addition, isopentanoic acid in particle
262 phase was the fourth most abundant monoacid in the MFB (8%) and LFB (7%) periods.
263 Nonanoic acid (C₉) is the fifth most abundant monoacid in both gas- and particle-phases,
264 showing a peak in the range of C₅-C₁₀ monoacids during the MFB and LFB periods (Table 1).
265 C₉ is an oxidation product of oleic acid (C_{18:1}) containing a double bond at C-9 position
266 (Kawamura and Gagosian, 1987).

267 Figure 3 shows temporal variations in the concentrations of major LMW
268 monoacids in gas and particle phases. The gaseous concentrations of acetic acid during the
269 MFB period (3–5 June) are significantly higher than those during the LFB period (2 and 23–
270 25 June). Similarly, the gas-phase concentrations of formic acid are higher during the MFB
271 period than the LFB period. The particle-phase concentrations of formic and acetic acids
272 during the MFB period are slightly higher than those during the LFB period. We found that
273 particle-phase formic and acetic acids were less abundant in the morning and nighttime hours
274 but maximized around noontime or in the afternoon. Two peaks of isopentanoic (iC₅) and
275 lactic (Lac) acids in particle-phase were observed at 6:00–9:00 on 3 June and 9:00–12:00 on
276 4 June when the field burning of agricultural wastes increased (Figure 2a). However, these
277 acids did not show any diurnal trend in both gas and particle phases.

278 Concentrations of total monoacids in gas-phase ranged from 1530 to 12,100 ng m⁻³
279 whereas those in particle-phase ranged from 960 to 9680 ng m⁻³. Mean concentrations of total
280 monocarboxylic acids in gas-phase were 6540 ng m⁻³ during the MFB and 3070 ng m⁻³
281 during the LFB period, whereas those in particle-phase were 4120 ng m⁻³ during the MFB

282 and 2850 ng m⁻³ during the LFB period. Overall, gas-phase concentrations of total monoacids
283 are higher than those of particle phase in both MFB and LFB periods. Table 1 summarises
284 average concentrations of individual monoacids over Mt. Tai. In gas-phase, acetic acid was
285 found as the dominant monoacid species (MFB: 3960 ng m⁻³, LFB: 1270 ng m⁻³) followed by
286 formic acid (MFB: 1570 ng m⁻³, LFB: 890 ng m⁻³) and lactic acid (MFB: 319 ng m⁻³, LFB:
287 433 ng m⁻³). In particle-phase, formic acid is the dominant species (MFB: 1410 ng m⁻³, LFB:
288 883 ng m⁻³) followed by acetic acid (MFB: 1120 ng m⁻³, LFB: 763 ng m⁻³), and lactic acid
289 (MFB: 917 ng m⁻³, LFB: 661 ng m⁻³). Mean concentrations of major monoacids in both gas
290 and particle phases are greater during the MFB period than the LFB period, except for
291 gaseous lactic acid and isopentanoic acid (iC₅).

292 The particle-phase fractions (F_p) of each organic acid were calculated as $F_p =$
293 $P/(G+P)$, where P is the particle-phase concentration and G is the gas phase concentration.
294 Figure 5 shows temporal variations in the F_p of major monoacids. F_p of acetic (C₂) and
295 propionic (C₃) acids are slightly lower during the MFB period (3–5 June) than other period.
296 The F_p for other LMW monocarboxylic acids over Mt. Tai did not show any clear temporal
297 or diurnal variations. Table 1 also summarizes mean F_p for individual monoacids during the
298 MFB and LFB periods. F_p of individual monoacids during the MFB and LFB periods ranged
299 from 0.19 (C₃) to 0.78 (C₆) and 0.27 (C₄) to 0.93 (iC₆), respectively. F_p of acetic acid was
300 0.24 during the MFB and 0.38 during the LFB, whereas F_p of formic acid was 0.47 during the
301 MFB and 0.52 during the LFB. F_p of acetic and formic acids during the MFB period are
302 significantly lower than those during the LFB period. On the other hand, F_p of
303 hydroxymonoacids, e.g., lactic acid, and branched chain monoacids, e.g., isopentanoic acid,
304 are significantly high (0.62–0.92), however, there is no systematic trend of the F_p values
305 between the MFB and LFB periods (Table 1).

306

307 4. Discussion

308 **4.1. Impact of field burning on the concentrations of formic and acetic acids**

309 To evaluate the effect of biomass burning of agricultural wastes on gaseous and
310 particulate formic and acetic acids, we plotted formic and acetic acids against levoglucosan (a
311 good tracer of biomass burning) (Simoneit, 2002) (Figure 6 and Figure A3). Particulate
312 formic and acetic acids in daytime show a weak correlation with levoglucosan during the
313 MFB (C_1 : $r = 0.29$, C_2 : $r = 0.28$) and LFB (C_1 : $r = 0.28$, C_2 : $r = 0.24$) periods. Gaseous formic
314 and acetic acids in daytime did not show a positive correlation with levoglucosan ($r = -0.62 -$
315 0.35). In addition, K^+ is emitted from field burning and agricultural activities, which is an
316 open field burning tracer. We plotted formic and acetic acids against K^+ (Figure 7 and Figure
317 A4). Particulate formic and acetic acids in daytime showed a positive correlation with K^+
318 during the MFB (C_1 : $r = 0.59$, C_2 : $r = 0.64$) and LFB (C_1 : $r = 0.44$, C_2 : $r = 0.63$) periods,
319 whereas particulate formic and acetic acids in nighttime did not show any correlations with
320 K^+ ($r = -0.29 - 0.04$).

321 We found that concentrations of gaseous acetic acid in daytime increased with K^+
322 during the MFB ($r = 0.63$) and LFB ($r = 0.63$) periods. Gaseous acetic acid in nighttime
323 showed a positive correlation with K^+ (MFB: $r = 0.54$, LFB: $r = 0.39$). However, formic acid
324 in both gas and particle phases in nighttime showed no correlation with K^+ during the MFB
325 and LFB periods ($r = -0.29 - 0.38$). We found that particulate formic and acetic acids and
326 gaseous acetic acid showed good correlations with K^+ and no correlations with levoglucosan.
327 Our results suggest that particulate formic and acetic acids and gaseous acetic acid are
328 influenced by the open field burning in the NCP. Although the sampling was conducted in the
329 MFB and LFB periods at this study, the obtained data must be interpreted with caution
330 because of a limited number of samples.

331 During the sampling periods, concentrations of formic and acetic acids and K^+ in
332 daytime were higher than those in nighttime. Because the top of Mt. Tai in daytime is below
333 the PBL, organic acids as well as other gaseous and aerosols are uplifted to Mt. Tai from the

334 lowland agricultural areas in the NCP where field burning of wheat straws is widely operated.

335 Table 2 compares the concentrations of formic and acetic acids in gas- and
336 particle-phases from multiple sites in the world. Particle-phase concentrations of formic
337 (1410 ng m^{-3}) and acetic (1120 ng m^{-3}) acids over Mt. Tai during the MFB period are
338 significantly higher than those reported in Los Angeles (Kawamura et al., 2000), megacities
339 of China such as Beijing (Wang et al., 2005) and Shanghai (Wang et al., 2006), Amazon
340 forest (Andreae et al., 1988) and Taiwan forest (Tsai and Kuo, 2013), and Alaska (Li and
341 Winchester, 1989). On the other hand, the gaseous levels of formic (1570 ng m^{-3}) and acetic
342 (3960 ng m^{-3}) acids over Mt. Tai during the MFB period are lower than those reported in
343 tropical forests from Amazon (Andreae et al., 1988), biomass burning plumes from the
344 Yangtze River Delta region, China (Kudo et al., 2014), urban air from Pasadena (Yuan et al.,
345 2015), and oil and gas fields from Utah (Yuan et al., 2015). The concentrations of formic and
346 acetic acids at Mt. Tai are 2–5 times higher than those reported from Los Angeles
347 (Kawamura et al., 2000), Greenland (Dibb and Arsenault, 2002), and the French Alps
348 (Preunkert et al., 2007), and 1–2 orders of magnitude higher than those reported from the
349 Pacific Ocean (Miyazaki et al., 2014) and Antarctica (Legrand et al., 2012). These
350 comparisons demonstrate that concentrations of formic and acetic acids in gas and particle
351 phases during field burning of wheat straw are relatively high compared with those from
352 other sites.

353 Carbon monoxide (CO) is originated from biomass burning (Wang et al., 2002) and
354 anthropogenic sources such as fossil fuel combustion. Average concentrations of CO in the
355 MFB and LFB periods were 553 ppb and 467 ppb, respectively (Kanaya et al., 2009), which
356 levels were very high and the air over Mt. Tai during the sampling periods was highly
357 polluted. Mean concentrations of total monoacids in both gas and particle phases are greater
358 during the MFB period (P: 4130 ng m^{-3} , G: 6540 ng m^{-3}) than the LFB period (P: 3070 ng m^{-3} ,
359 G: 2860 ng m^{-3}). Field burning of agricultural wastes is an important source of monoacids in

360 the atmosphere over Mt. Tai. In addition, Figure A2 shows seven-day air mass forward
361 trajectories during the MFB period. The dominant air mass flow at a levels of 1500 m and
362 750 m was northeasterly and easterly, indicating outflow of air masses to East Asia and the
363 western North Pacific. We suggest that high abundances of monoacids in the particle phases
364 not only degrade the air quality in East Asia but also largely contribute to enhance the
365 hygroscopic properties of ambient aerosols because they are highly water-soluble. A possible
366 formation of gaseous and particulate acetic acid will be discussed in section 4.3.

367

368 **4.2. Gas/particle partitioning of formic, acetic, and lactic acids**

369 Generally, F_p increases with an increase of carbon numbers of monoacids (Yatavelli
370 et al., 2014). However, the F_p values of formic (0.47–0.52) and acetic (0.24–0.38) acids at the
371 top of Mt. Tai are larger than the values expected from the high vapour pressures of formic
372 ($V_p = 5.6 \times 10^{-2}$ atm) and acetic ($V_p = 2.1 \times 10^{-2}$ atm) acids. In addition, F_p values of C₁ and C₂
373 over Mt. Tai are significantly larger than those reported from the Pacific Ocean (C₁: $F_p = 0.04$,
374 C₂: $F_p = 0.06$) (Miyazaki et al., 2014) and urban Los Angeles (C₁: $F_p = 0.16$, C₂: $F_p = 0.06$)
375 (Kawamura et al., 2000), in which the same sampling and analytical protocols were used. On
376 the other hand, the F_p of lactic acid (0.61 – 0.69) over Mt. Tai is slightly lower than that
377 reported from the Pacific Ocean (0.76 – 0.82) (Miyazaki et al., 2014).

378 Khan et al. (1995) reported that ambient temperature and relative humidity (RH) are
379 important factors to determine the gas-particle partitioning of organic acids. Gas-particle
380 partitioning of organic acids depend on temperature (related to Henry's law) and RH (related
381 to aerosol liquid water content indicator). In this study, diurnal variations of temperature and
382 RH were observed. However, there is no consistent relationship between F_p of organic acids
383 and temperature or RH. The ambient temperature and RH did not affect aerosol surface
384 characteristics over Mt. Tai.

385 As a potential reason of higher F_p values for organic acids, chemical states of those

386 acids in ambient aerosols should be considered. For example, ammonia reacts with gaseous
387 acidic species (e.g., oxalic acid) (Paciga et al., 2014). Gas phase acetic acid can be adsorbed
388 on calcite (CaCO_3) that is a major component of mineral dust (Alexander et al., 2015).
389 Because the vapour pressures of organic salts are lower than those of free organic acids, gas
390 to particle conversion of organic acids may be an important factor to control their gas/particle
391 partitioning. However, the F_p of formic and acetic acids did not show correlation with
392 alkaline species such as Na^+ , K^+ , Ca^{2+} , Mg^{2+} , and NH_4^+ (MFB: $r < -0.12$, LFB: $r < 0.24$). We
393 calculated total cation equivalents (Na^+ , NH_4^+ , K^+ , Mg^{2+} , and Ca^{2+}) and total anion
394 equivalents (F^- , MSA^- , Cl^- , NO_2^- , NO_3^- , Br^- , PO_4^{3-} , and SO_4^{2-}) including normal (C_1 - C_{10}),
395 branched chain (i C_4 -i C_6), aromatic (benzoic), and hydroxyl (lactic and glycolic) monoacids,
396 although CO_3^- and HCO_3^- , as well as unidentified organic anions, were not considered. The
397 slope of more than unity (1.21) indicates that excess cations exist in the aerosols over Mt. Tai.
398 We calculated total cation equivalents minus total anion equivalents (cations - anions) in the
399 particles from Mt. Tai. F_p of formic (C_1) and acetic (C_2) acids did not show a positive
400 correlation with cation - anions (C_1 : $r = -0.36$, C_2 : $r = -0.28$). Gas to particle conversion of
401 formic and acetic acids over Mt. Tai was not controlled by the above-mentioned excess
402 alkaline species.

403 *Barsanti et al.* [2009] reported that amines can more easily form organic acid salts
404 compared to ammonia. Unfortunately, we did not measure amines in this study. Amines that
405 are derived from vehicle exhaust and animal husbandry (e.g., Cape et al., 2011) and from
406 biomass burning (Ge et al., 2011) could act as a counterpart for free organic acids and those
407 amine-organic acid reactions may contribute to a new particle formation and condensation
408 growth of particles, enhancing the F_p values in the NCP. We suggest that gas/particle
409 partitioning of formic and acetic acids may be influenced by unidentified factors during
410 atmospheric transport over Mt. Tai, which need further study to be clarified in the future.

411

412 **4.3. Primary and secondary sources of formic and acetic acid**

413 Formic acid (C_1) is largely derived from secondary sources (Khwaja et al., 1995,
414 Pommier et al., 2016), whereas acetic acid (C_2) is largely emitted from primary sources
415 (Kawamura et al., 1985, 2000). The ratio of C_1/C_2 in particle phase is an indicator for primary
416 or secondary source of organic acid aerosol. Talbot et al. (1988) reported that high ratio of
417 C_1/C_2 (> 1) means the dominance of secondary sources, whereas low ratio (< 1) means the
418 dominance of primary sources. In this study, average ratio of C_1/C_2 in daytime in the MFB
419 and LFB periods was 1.3 and 1.2, respectively. C_1/C_2 ratios of Mt. Tai are higher than those
420 reported from urban in China (Shanghai: 0.5) (Wang et al., 2006), tropical rain forest (1.0)
421 (Andreae et al., 1988), and Ocean (0.3) (Miyazaki et al., 2014) and are comparable to those
422 reported from urban Los Angeles (1.4) (Kawamura et al., 2000) and China (Beijing: 1.1)
423 (Wang et al., 2005). High ratio of C_1/C_2 in the aerosols of Mt. Tai indicates that secondary
424 sources are important for the formation of organic acids over the Mt. Tai. Daytime maxima of
425 formic and acetic acids in particle phase (Figure 3) also indicate a secondary source for
426 formic and acetic acids.

427 Kawamura et al. (2013) reported that oxalic acid is the most abundant dicarboxylic
428 acid and largely emitted by the field burning of agricultural wastes in the NCP during the
429 same campaign, but its major sources are photochemical processes (secondary sources).
430 Concentrations of formic and acetic acids in particle phase show positive correlations with
431 concentrations of oxalic acid (C_1 : $r = 0.63$, C_2 : $r = 0.54$).

432 CO is an excellent tracer of primary combustion product. To evaluate the primary
433 and secondary sources of formic and acetic acids during the MFB and LFB periods, we
434 examined the relationship between formic acid or acetic acid and CO as shown in Figure 8.
435 Concentrations of particulate acetic acid in daytime were found to increase linearly with CO
436 during the MFB ($r = 0.70$) and LFB ($r = 0.44$) periods. Concentrations of gaseous acetic acid
437 also showed a positive correlation with CO (MFB: $r = 0.62$, LFB: $r = 0.49$). Concentrations of

438 particulate formic acid were found to increase linearly with CO during the MFB ($r = 0.64$).
439 However, concentrations of gaseous formic acid during the MFB and LFB periods and
440 particulate formic acid during the LFB period did not correlate with CO ($r < 0.39$). These
441 correlation analyses suggest that emissions of acetic acid in both particle and gas phases over
442 Mt. Tai are enhanced with an increase in the intensity of biomass burning during the MFB
443 periods.

444 In addition, the slopes of regression lines drawn for acetic acid in gas- and
445 particle-phases and CO are different between the MFB and LFB periods. In particle-phase,
446 higher ratio of acetic acid/CO (0.003) was obtained during the MFB period, whereas lower
447 ratio (0.001) was obtained during the LFB period. In gas-phase, higher ratio of acetic acid/CO
448 (0.006) was obtained during the MFB period, whereas lower ratio (0.001) was obtained
449 during the LFB period. K^+ as a biomass tracer shows a significant correlation with CO,
450 whereas the slopes of the regression lines for K^+ and CO are different between the MFB and
451 LFB periods (Figure A5). Ratio of K^+ /CO (0.007) obtained from the slope of K^+ and CO
452 during the MFB period was lower than that during the LFB period (0.001). SO_4^{2-} as a fossil
453 fuel combustion and industrial emission tracer shows a significant correlation with CO,
454 whereas the slopes of the regression lines for SO_4^{2-} and CO are different between the MFB
455 and LFB periods (Figure A6). Ratio of SO_4^{2-} /CO (0.112) obtained from the slope of SO_4^{2-}
456 and CO during the MFB period was lower than that during the LFB period (0.056). The
457 higher ratios of K^+ /CO during the MFB period indicate a strong influence of field burning of
458 agricultural wastes in the NCP. The lower ratios of K^+ /CO during the LFB period suggest an
459 influence of contribution of CO from other sources such as fossil fuel combustion in the
460 surroundings of North China. These results again suggest that abundant gas- and
461 particle-phase acetic acid is directly emitted from the field burning of agricultural wastes.

462 During the campaign, ethane, ethene, propane, and propene (highly volatile
463 hydrocarbons) are emitted from open burning of crop residue in the NCP (Kanaya et al.,

464 2009), which are important precursors of acetic acid (Warneck, 2005). Concentrations of
465 these hydrocarbons were reported to increase under the influence of biomass burning plumes
466 in the NCP (Kanaya et al., 2009). The daytime peaks of particulate and gaseous acetic acid
467 can be explained by secondary processing via photochemical oxidation of volatile
468 hydrocarbons over Mt. Tai. We plotted the ratios of acetic acid in total monocarboxylic acids
469 (C_2 /total MCA) against K^+ (Figure 9). Ratios of C_2 to total MCA in both gas- and
470 particle-phases increased with an increase in CO concentrations during the MFB and LBF
471 periods. In particular, contributions of C_2 to total MCA during the MFB period showed a
472 significant correlation with CO (gas: $r = 0.93$, particle: $r = 0.80$). These positive correlations
473 suggest that acetic acid is selectively produced in the atmosphere by the photochemical
474 oxidation of volatile hydrocarbons when organic precursors are transported over the summit
475 of Mt. Tai from the field burning areas in lowland. We consider that primary emission and
476 secondary formation of gas- and particle-phase acetic acid are associated with field burning
477 of agricultural wastes in the NCP.

478

479 **4.4. High abundances of biogenic monocarboxylic acids**

480 Lactic acid is produced by lactobacillus (Cabredo et al., 2009), which is known to
481 exist in soil (Huysman and Verstraete, 1993) and emitted from plant tissues in agricultural
482 farm (primary origin) (Raja et al., 2008). This organic acid can also be produced by the
483 oxidation of isoprene with ozone (Nguyen et al., 2010) (secondary origin). Gaseous and
484 particulate lactic acid did not correlate ($r < -0.15$) with 2-methyltetrols, good tracers of
485 isoprene-derived secondary organic aerosols (Fu et al., 2012). Isopentanoic acid can be
486 produced by bacteria such as *Bacteroides ruminicola* and *Megasphaera elsdenii* (Allison,
487 1978). We found a strong positive correlation between lactic and isopentanoic acids in both
488 gas ($r = 0.91$) and particle ($r = 0.93$) phases (Figure 10). This result is similar to that obtained
489 for snow pit samples from high mountain site in central Japan, although the snow samples are

490 influenced by the outflow of Asian dust (Mochizuki et al., 2016). We suggest that major
491 portion of lactic and isopentanoic acids were directly derived from the terrestrial ecosystem.
492 Microbial productivity involved with lactic and isopentanoic acids should widely exist in the
493 terrestrial ecosystems in China. Hence, we conclude that these LMW monoacids are mainly
494 derived from microbial sources.

495 Average particle phase concentration of lactic acid (789 ng m^{-3}) is one order of
496 magnitude higher than those in marine aerosol samples from the Pacific (33.1 ng m^{-3})
497 (Miyazaki et al., 2014), pasture samples (22.1 ng m^{-3}) and forest samples (9.2 ng m^{-3}) in
498 Brazil (Graham et al., 2002). In addition, average gas phase concentration of lactic acid (376
499 ng m^{-3}) is significantly higher than that reported in the marine atmosphere from the Pacific
500 Ocean (7.1 ng m^{-3}) (Miyazaki et al., 2014). These results suggest that the NCP is an important
501 source of particulate and gaseous lactic acid in the atmosphere. Furthermore, high
502 abundances of isopentanoic acid (iC_5) in particle phase (average, 250 ng m^{-3}) in the ambient
503 aerosols from Mt. Tai were obtained, which is the same order of magnitude as particulate
504 lactic acid. To our knowledge, this is the first observation on the abundances of branched
505 chain monoacids in atmospheric aerosols. In previous studies, hydroxy monoacids such as
506 lactic acid and branched chain monoacids such as isopentanoic acid are scarcely reported
507 (Kawamura et al., 2012; Miyazaki et al., 2014; Mochizuki et al., 2016) We propose that
508 measurements of biogenic hydroxy and branched chain monoacids are important to evaluate
509 the source of organic aerosols and their atmospheric chemical processing.

510

511 **5. Summary and conclusions**

512 Gaseous (G) and particulate (P) low molecular weight (LMW) monocarboxylic acids,
513 including normal ($\text{C}_1\text{-C}_{10}$), branched ($\text{iC}_4\text{-iC}_6$), hydroxyl (lactic and glycolic), and aromatic
514 (benzoic) structures, were detected in the ambient air samples collected at the summit of Mt.
515 Tai in the North China Plain (NCP). We found that acetic acid was the most abundant

516 gaseous species whereas formic acid was the most abundant particulate species. Significantly
517 high concentrations of formic (average, G: 1570 ng m⁻³, P: 1410 ng m⁻³) and acetic (G: 3960
518 ng m⁻³, P: 1120 ng m⁻³) acids were detected during the period of inflow of air plumes from
519 the field burning of agricultural waste (wheat straw) in the NCP. Because gas- and
520 particle-phase concentrations of formic and acetic acids in daytime are correlated ($r = 0.32$ –
521 0.64) with K⁺ (a field burning tracer), field burning of agricultural waste should be an
522 important primary source of LMW monoacids over Mt. Tai. In addition, we found a large
523 fraction (F_p) of particulate formic (0.49) and acetic (0.32) acids in the total (G+P)
524 concentrations. We suggest that primary emission and secondary formation of gaseous and
525 particulate acetic acid are associated with field burning of agricultural waste in the NCP. We
526 also found high concentrations of isopentanoic acid (331 ng m⁻³) and lactic acid (917 ng m⁻³)
527 in particle phase, suggesting that biological sources are also important for LMW monoacids
528 in organic aerosols. LMW monoacids emitted from the field burning of agricultural wastes
529 importantly contribute to the formation of organic aerosols in Central East China.

530

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539

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- 731

732 **Figure captions**

733 Figure 1. A map of East Asia with the sampling site at Mt. Tai in the North China Plain.

734 Figure 2. Seven-day air mass back trajectories arriving at the summit of Mt. Tai (1534 m
735 above sea level) at a level of 1500 m a. s. l. with fir spots in Central East China for
736 selected periods in June 2006.

737 Figure 3. Temporal variations in the concentrations of major monocarboxylic acids in gas and
738 particle phases and K^+ . Open circles indicate gas phase samples and solid circles indicate
739 particle phase samples. Shaded areas indicate nighttime.

740 Figure 4. Compositions of monocarboxylic acids in gas and particle phases during less field
741 burning (LFB) and more field burning (LFB) influenced periods.

742 Figure 5. Temporal variations in particle phase fraction (F_p) for major organic acids. Shaded
743 areas indicate nighttime.

744 Figure 6. Correlations of formic and acetic acids in gas and particle phases with levoglucosan
745 in daytime during less field burning (LFB) and more field burning (LFB) influenced
746 periods.

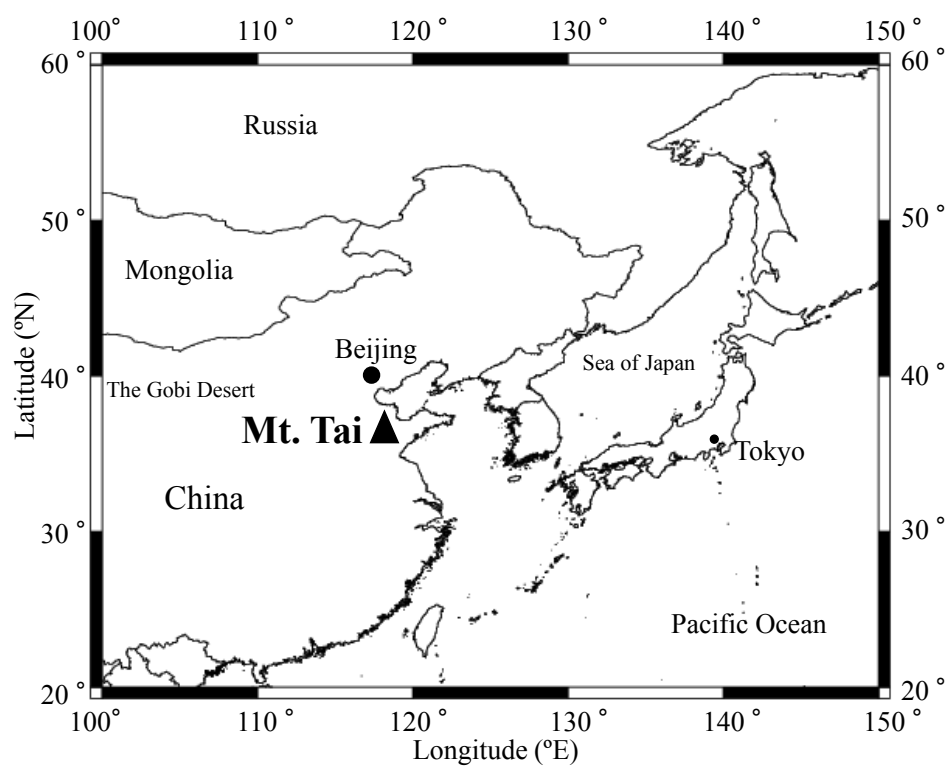
747 Figure 7. Correlations of formic and acetic acids in gas and particle phases with K^+ in
748 daytime during less field burning (LFB) and more field burning (LFB) influenced
749 periods.

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751 monoxide (CO) during less field burning (LFB) and more field burning (LFB) influenced
752 periods.

753 Figure 9. Changes of contribution of particulate and gaseous acetic acid (C_2) to total
754 monocarboxylic acids (total MCA) in the Mt Tai during less field burning (LFB) and
755 more field burning (LFB) influenced periods.

756 Figure 10. Correlation plots of lactic acid and isopentanoic acid in gas and particle phases.

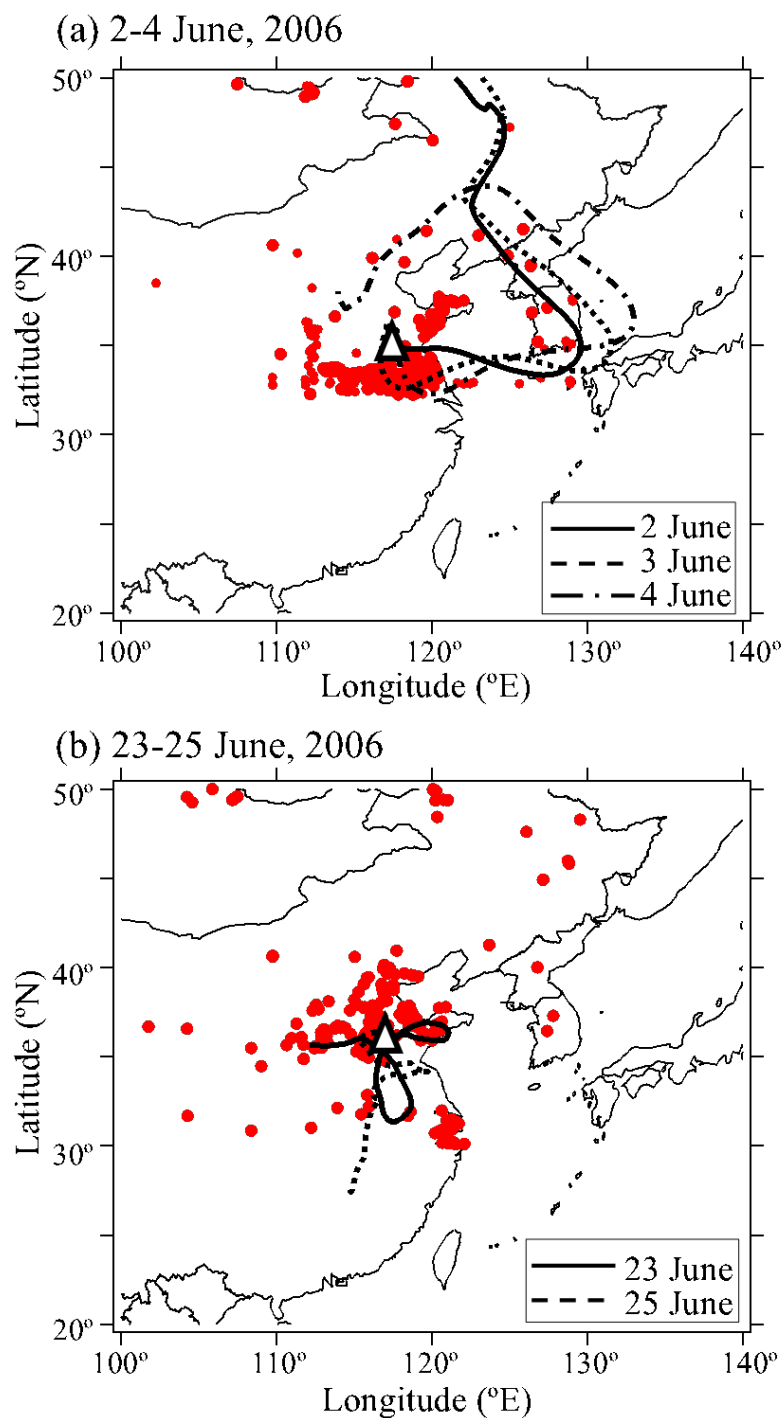
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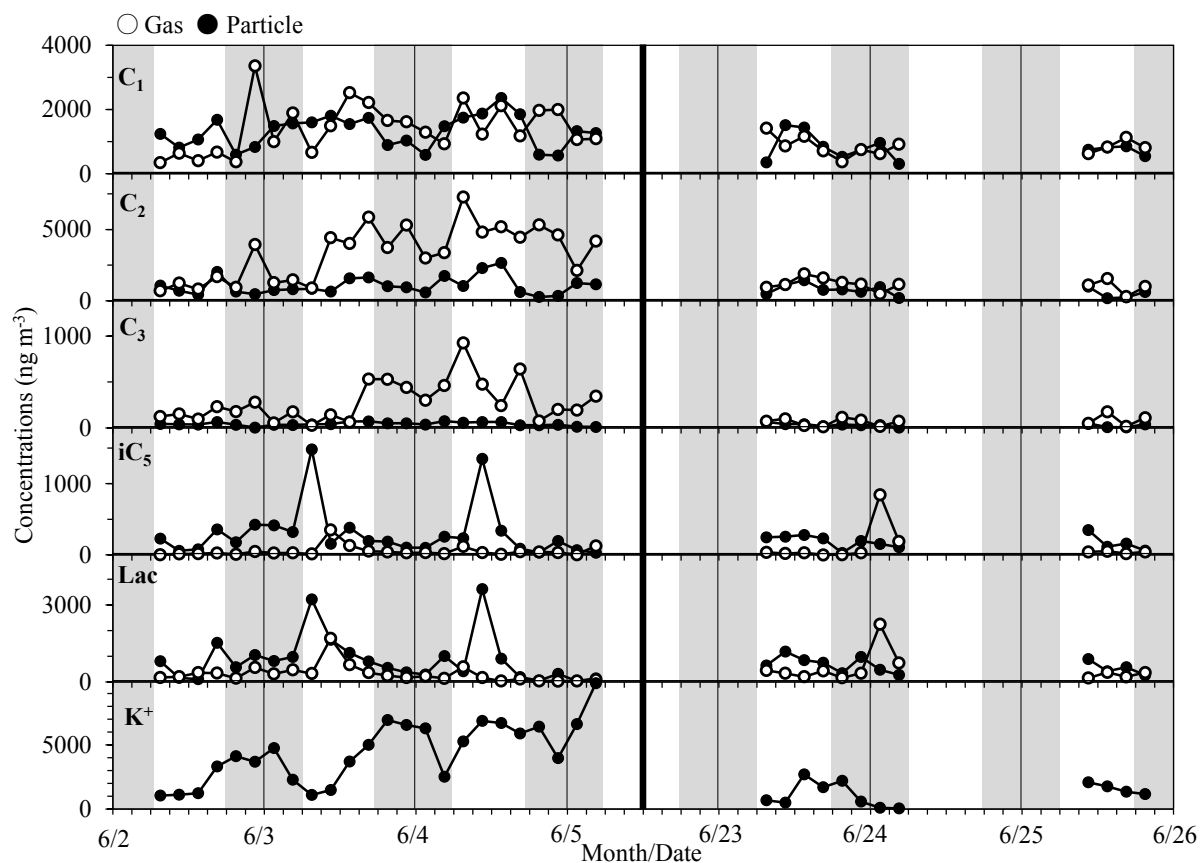
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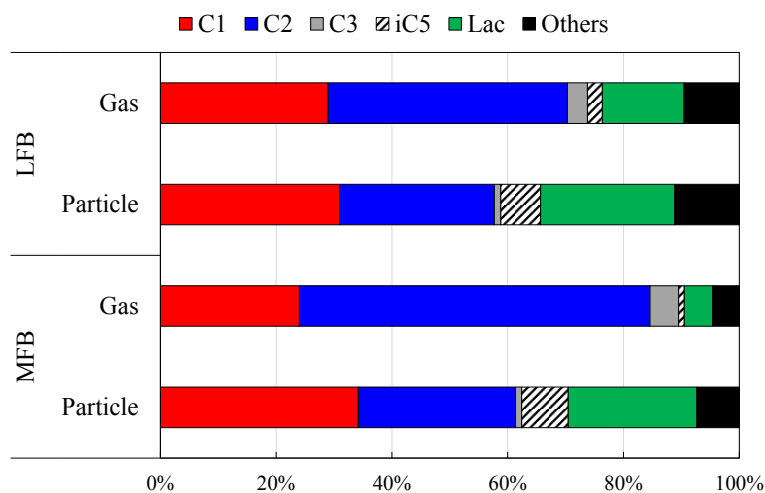
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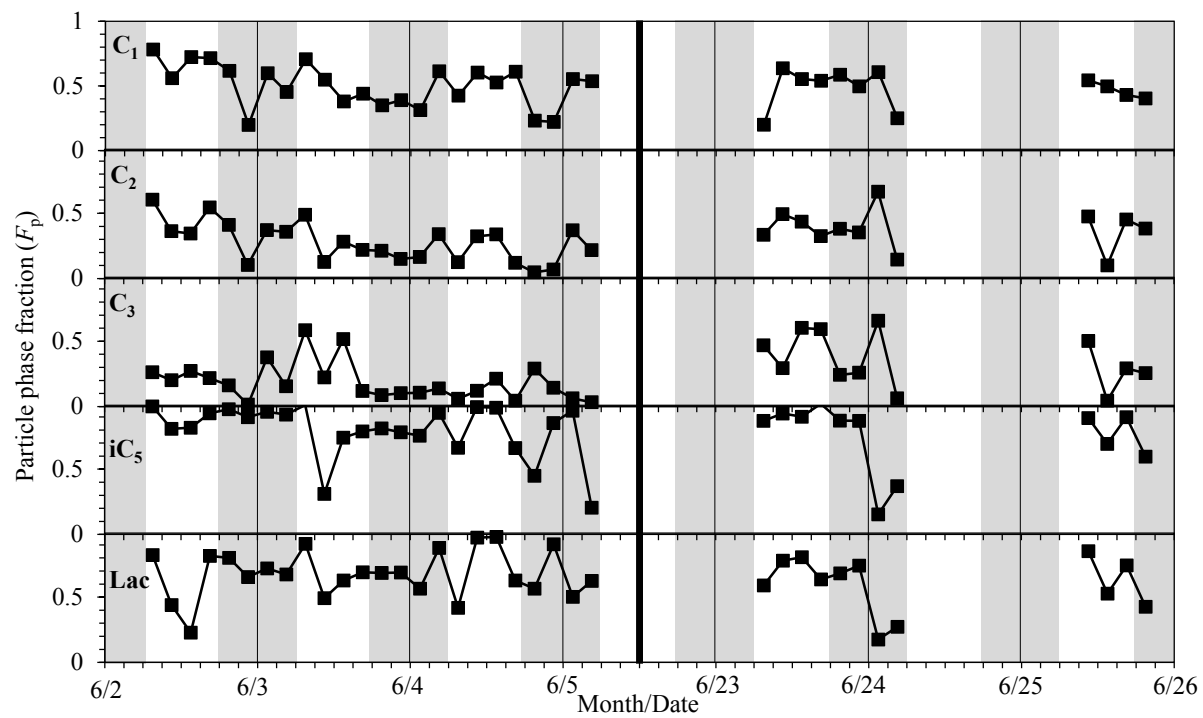
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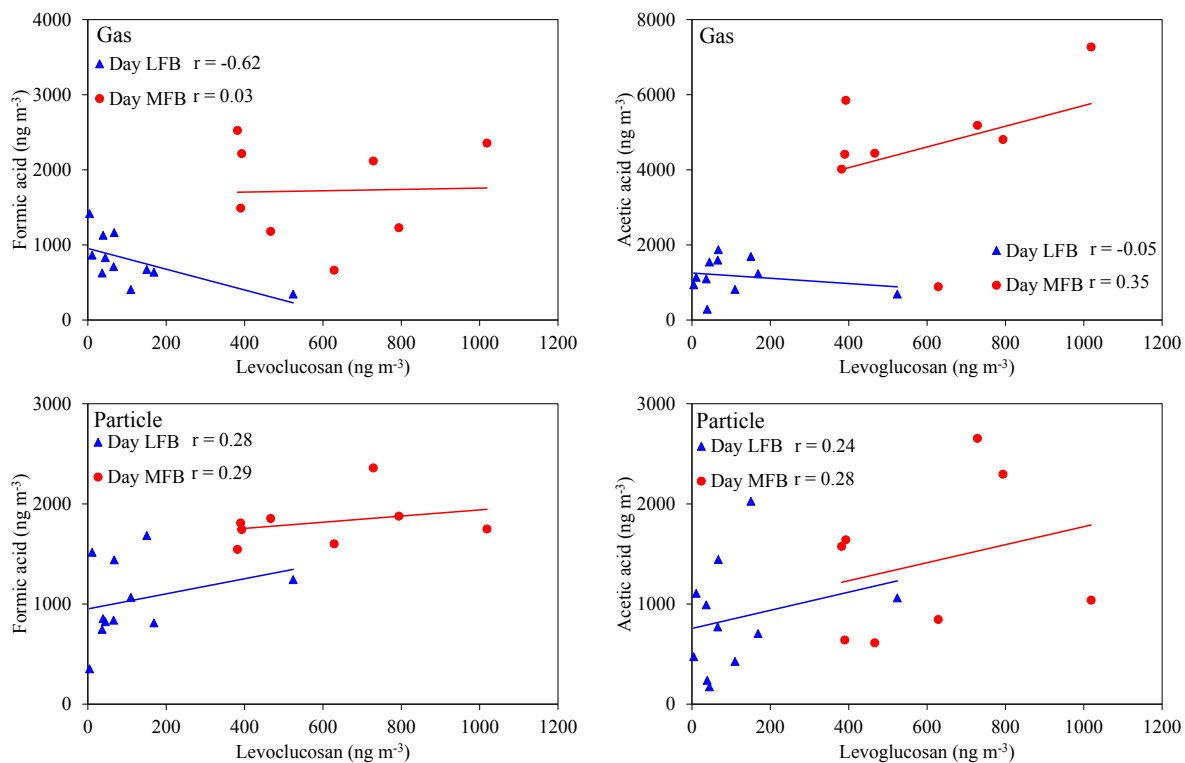
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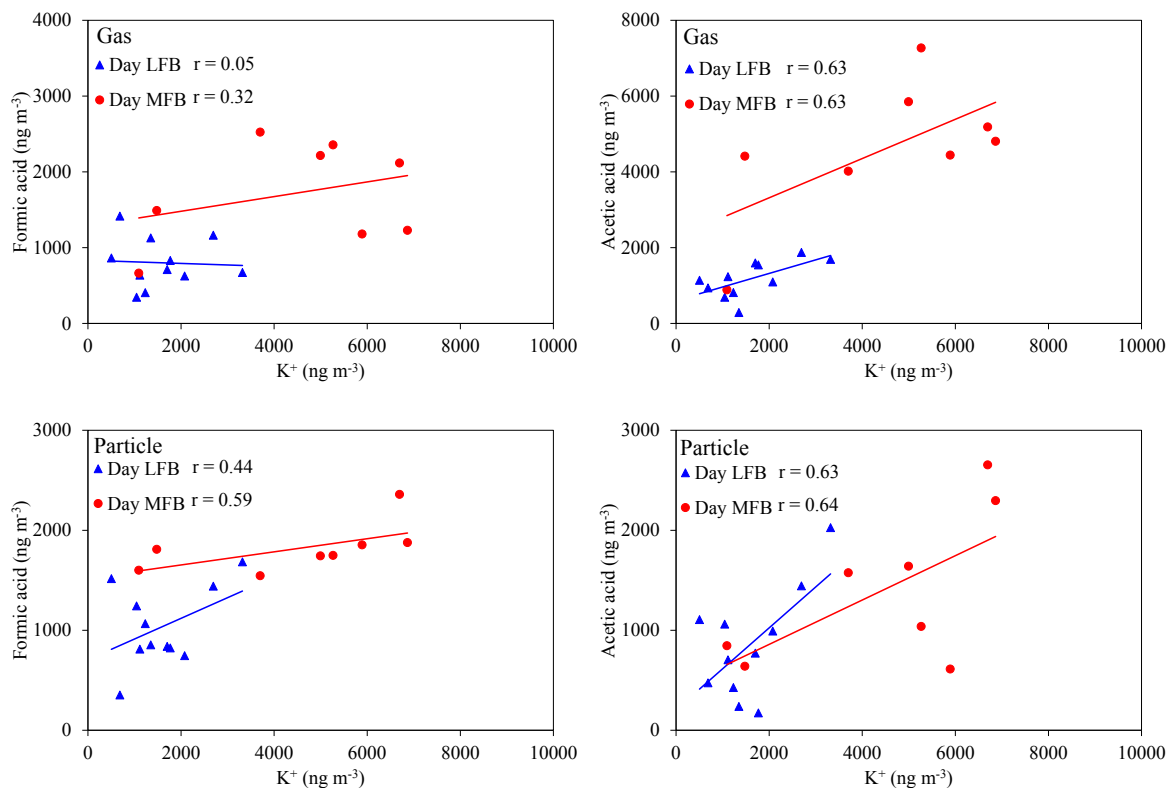


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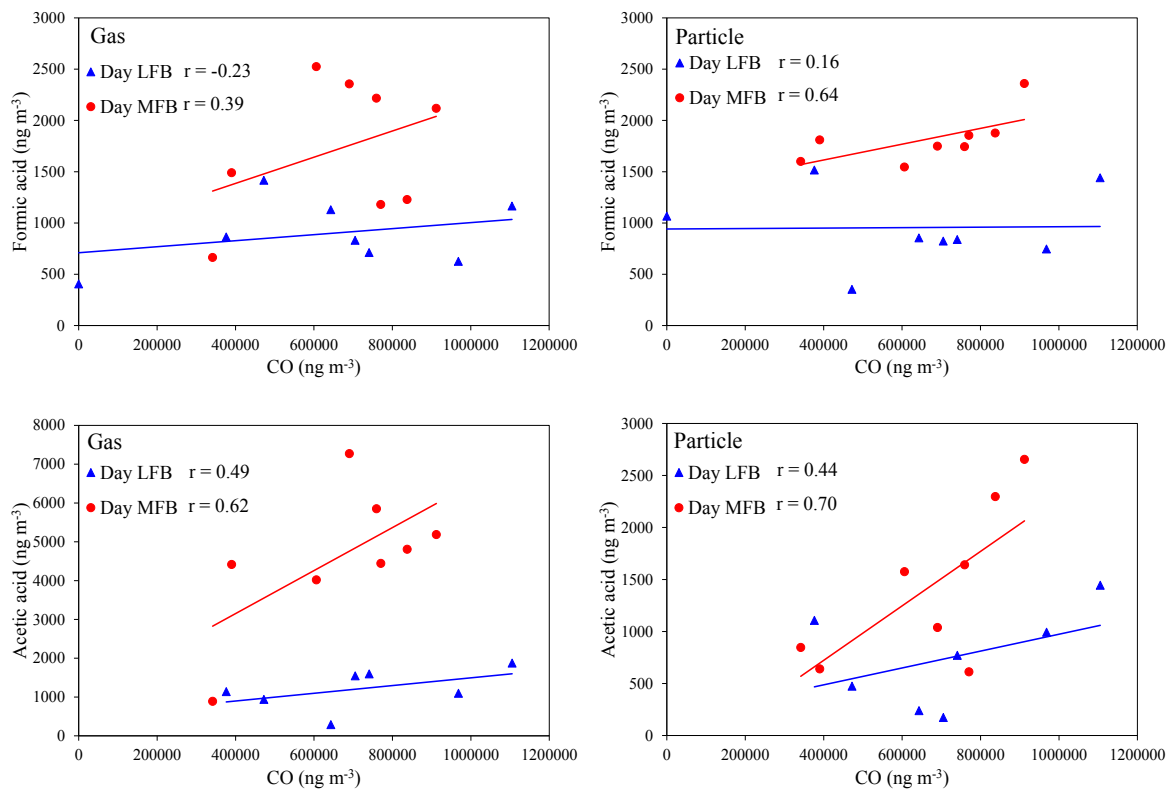


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784 Figure 7. Correlations of formic and acetic acids in gas and particle phases with K^+ in

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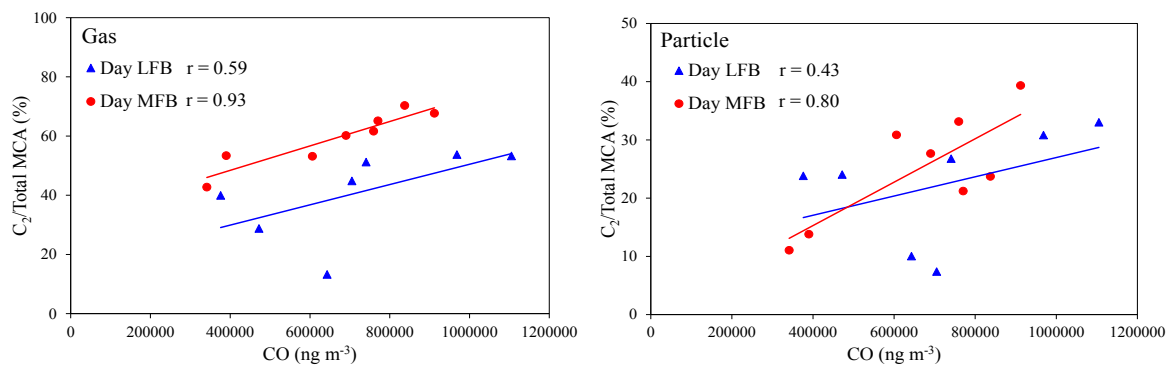


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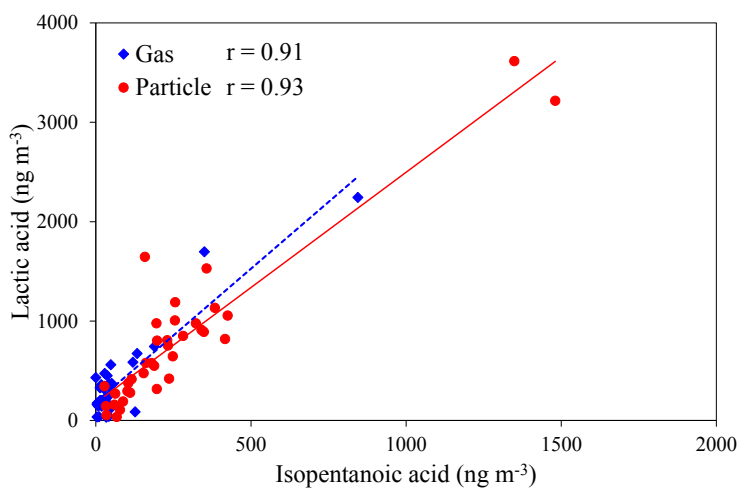


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793 Figure 9. Changes of contribution of particulate and gaseous acetic acid (C₂) to total
 794 monocarboxylic acids (total MCA) in the Mt Tai during less field burning (LFB) and more
 795 field burning (LFB) influenced periods.

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797



798

799 Figure 10. Correlation plots of lactic acid and isopentanoic acid in gas and particle phases.

800 Table 1. Average concentrations (ng m^{-3}) of gaseous and particulate low molecular weight
 801 monocarboxylic acids and particle phase fraction (F_p) during less field burning (LFB) and
 802 more field burning (MFB) influenced periods.

Acid species	MFB						LFB						MFB	LFB
	Gas			Particle			Gas			Particle			F_p	F_p
	min.	max.	average	min.	max.	average	min.	max.	average	min.	max.	average	average	average
Aliphatic acids														
Formic, C ₁	664	2520	1570	570	2360	1410	346	3360	890	307	1690	883	0.47	0.52
Acetic, C ₂	888	7270	3960	263	2650	1120	289	3940	1270	174	2030	763	0.24	0.38
Propionic, C ₃	28	927	324	12	74	45	12	280	106	4	65	33	0.19	0.30
Isobutyric, iC ₄	2	90	37	0	42	16	7	45	23	1	33	12	0.35	0.32
Butyric, C ₄	4	192	77	3	42	17	5	83	36	1	26	11	0.30	0.27
Isopentanoic, iC ₅	4	350	64	33	1480	331	0	844	79	28	425	196	0.76	0.80
Pentanoic, C ₅	0	12	4	0	8	2	0	9	2	0	11	3	0.46	0.58
Isohexanoic, iC ₆	0	8	2	0	4	2	0	4	0	0	5	2	0.62	0.92
Hexanoic, C ₆	0	31	4	2	17	9	0	30	5	0	17	5	0.78	0.61
Heptanoic, C ₇	0	5	2	0	10	2	0	5	2	1	7	2	0.48	0.50
Octanoic, C ₈	0	10	3	1	41	13	0	15	8	1	26	11	0.77	0.57
Nonanoic, C ₉	4	370	59	1	117	40	0	400	123	4	693	112	0.42	0.41
Decanoic, C ₁₀	0	9	2	0	18	5	0	9	3	0	20	4	0.59	0.36
Sub total			6110			3010			2550			2040		
Aromatic acid														
Benzoic, Benz	2	119	40	5	98	27	6	48	25	2	34	17	0.51	0.39
Hydroxyacids														
Lactic, Lac	33	1700	319	36	3615	917	144	2240	433	107	1530	661	0.69	0.61
Glycolic, Glyco	12	343	72.1	14	348	168	16	125	65.1	25	431	140	0.71	0.65
Sub total			391			1090			498			801		

803

804

805 Table 2. Comparisons of atmospheric concentrations of formic and acetic acids in the
 806 atmosphere over Mt. Tai (this work) with those in previous studies from multiple sites in the
 807 world.

Location		Formic acid (ng m^{-3})	Acetic acid (ng m^{-3})	References
Mt. Tai, China (wheat biomass burning)	G	1570	3960	This study
Los Angeles (urban)	G	860	1800	Kawamura et al. (2000)
California, Pasadena (urban)	G	4000	—	Yuan et al. (2015)
Utah (oil and gas producing region)	G	4700	—	Yuan et al. (2015)
Amazon (tropical forest)	G	3400	5900	Andreae et al. (1988)
Yangtze River Delta region, China (biomass burning)	G	—	5000	Kudo et al. (2014)
Greenland (mountain)	G	1070	1070	Dibb and Arsenault (2002)
France (Alps)	G	640	750	Preunkert et al. (2007)
Pacific Ocean	G	55	122	Miyazaki et al. (2014)
Antarctica	G	92	75	Legrand et al. (2012)
Mt. Tai, China (wheat biomass burning)	P	1410	1120	This study
Los Angeles (urban)	P	163	120	Kawamura et al. (2000)
Amazon (tropical forest)	P	46	48	Andreae et al. (1988)
Pacific Ocean	P	2	8	Miyazaki et al. (2014)
Taiwan (subtropical forest)	P	30	312	Tsai and Kuo (2013)
Alaska	P	244	744	Li and Winchester (1989)
Beijing, China (urban)	P	370	350	Wang et al. (2005)
Shanghai, China (urban)	P	80	150	Wang et al. (2006)

808

809 G: gas-phase, P: particle-phase