



## Hydrocarbon Formation in Immature Sediments

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**Abstract:** Immature sediments ( $R_o < 0.6$ ) and hydrates commonly contain low concentrations of  $C_{2-8+}$  alkanes/alkenes, higher alkanes, cycloalkanes and aromatics (temperature  $< 373$  K; Pressure  $< 100$  MPa). Their origin is enigmatic. Traditionally they are interpreted as migrated thermogenic oil. Water treatment experiments have established that they could be formed through the interaction of water and organic carbon by Fe catalysis at 298 K. This study investigates the Eh and pH associated with low temperature (263-298 K) hydrocarbon formation in saline pore-waters containing Ca-montmorillonite and  $Fe^0$  (ZVI) over a 300 day period in order to identify the principal reaction mechanisms. The interaction of flowing gaseous carbon dioxide-hydrocarbon mixtures with halite promoted with  $Fe_xO_y$ ,  $Fe_x[OH]_y$  at 288 – 308 K is examined experimentally. The study established that halite and mixtures of halite with organic material, Fe-montmorillonite,  $CaCO_3$ ,  $Ca(OH)_2$ ,  $MgSO_4$ ,  $(NH_4)_2SO_4$ ,  $K_2SO_4$ , pyroclastics, ash, phosphate enriched organic material, and coal can facilitate the removal of  $CO_2$ , the formation of  $H_2O$  on the catalyst surface, and the formation of hydrocarbons incorporating the  $CO_2$ .

**Key words:** ZVI; Oil formation;  $CO_2$ ; Eh; pH; NaCl; Halite; Montmorillonite

### 1. INTRODUCTION

Some light hydrocarbons ( $C_{2+}$ ) found in immature sediments ( $R_o < 0.6$ ) and with methane hydrates (temperature =  $< 373$  K; Pressure =  $< 100$  MPa; 0-500 mbsb) from the Lower Congo Basin (East Atlantic), West Pacific Continental Slope, Pacific Abyssal Plain, and Indian Ocean Abyssal Plain may have been formed in situ by catalysis within the pore-waters (Antia, 2008a; 2009a; 2009b).

The principal catalytic processes for hydrocarbon formation are:

- Sabatier Process** (also known as the Fischer Tropsch (FT) Process (e.g. Elworthy & Williamson, 1902; Sabatier, 1908; 1910; Storch *et al.* 1951; Steynberg & Dry, 2004)) where  $aCO + bH_2 = C_xH_y + cH_2O + dCO_2$  (alkanes and alkenes (gas + oil)). Variants include  $aCO + bH_2 + C_xH_y = C_{x+n}H_{y+m} + cH_2O + dCO_2$ . By-products include  $C_xH_yO_z$ . This process requires a gaseous reaction environment (423-673 K, 0.1 – 5 MPa).
- Kolbel-Engelhardt (KE) Process** (e.g. Kolbel & Englehardt, 1959; O'Rear, 2005) where  $3nCO + nH_2O = C_nH_{2n} + 2nCO_2$  and  $nCO_2 + 3nH_2 = C_nH_{2n} + 2nH_2O$ . i.e.  $mCO_2 + 3nCO + nH_2O + 3mH_2 = C_{(n+m)}H_{2(n+m)} + 2mH_2O + 2nCO_2$ . This process requires a gaseous reaction environment (423-673 K, 0.1 – 5 MPa).
- Polymerisation Process** (e.g. Baxendale *et al.* 1946a; 1946b; 1946c; Hori *et al.* 1989; Hardy & Gillham, 1995; Campbell *et al.* 1997; Deng *et al.* 1997; 1999; Schrick *et al.* 2002; Lim *et al.* 2007; Antia, 2008a; 2009a; 2009b; 2010; 2011; Bokare & Choi, 2009) where hydrocarbons form catalytically in an aqueous environment from the reductive/oxidative interaction and polymerisation of  $H_xC_yO_z$ ,  $C_xH_y$ ,  $C_x$ ,  $CO_x$ ,  $H_xO_y$ , organo-halides and metal carbonyls/carbides. For example 74.7 API Lower Congo Basin oil found in association with methane hydrates in methane vents (pockmarks), may have been formed (Antia, 2008a) by the generic equation  $4.09 m^3 CH_4 + 3.09 m^3 CO_2 + 6.18 m^3 H_2 = 4.96 kg H_2O + 4.58 kg Oil$  [where  $3.09 CO_2 + 6.18 H_2 = 3.09 HCO + 3.09 H_3O^+ = 3.09$

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H<sub>2</sub>CO [formaldehyde] + 3.09 H<sub>2</sub>O]. This process requires an aqueous reaction environment (<260->373 K, <0.1 ->80 MPa).

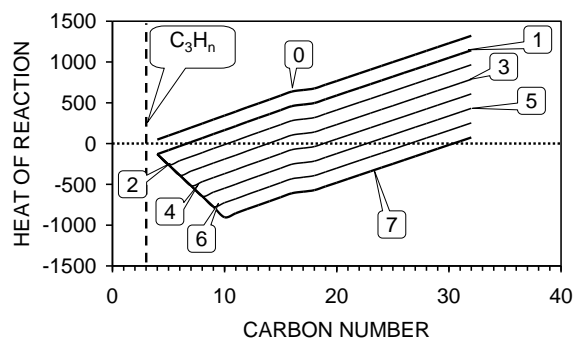
The FT and KE processes are unlikely to be responsible for the formation of low temperature hydrocarbons (Storch *et al.* 1951; Steynberg & Dry, 2004). Experiments (Deng *et al.* 1997) have established that C<sub>2+</sub> hydrocarbons can be formed over short time interval at 298 K using a carbon source (C, or metal carbide, or CH<sub>4</sub>) + an oxidant source (H<sub>2</sub>O or CO<sub>2</sub>).

This study uses two experiments to establish the mechanism for aqueous phase hydrocarbon chain growth utilising Fe catalyst + H<sub>2</sub>O [the oxidant] + C [the carbon source] (Experiment I) and for gaseous phase hydrocarbon chain growth utilising NaCl catalysts + CO<sub>2</sub> (the oxidant) + CH<sub>4</sub> [the carbon source] (Experiment II) at temperatures of 260 – 308 K. Both reaction routes may be present in the immature sediments of the continental slope and abyssal plain.

## 2. CHEMICAL BACKGROUND

In the geological environment, products and reactants are present as solids, liquids, gases and supercritical fluids. These phases affect the heat of formation (and reaction) [ $\Delta H$ ], the Gibbs free energy [ $\Delta G$ ], the equilibrium constants, the reaction rates and the availability of intermediate reaction routes.

- Heats of Formation** change with product or reactant phase. The Heat of Reaction [ $\Delta H_R$ ] =  $\Sigma\Delta H(\text{Products}) - \Sigma\Delta H(\text{Reactants})$  (Kotz & Treichel, 1996). Reactions with a negative  $\Delta H_R$  (exothermic) are likely to progress at low temperatures.
- Gibbs Free Energy**  $\Delta G = \Sigma\Delta G(\text{Products}) - \Sigma\Delta G(\text{Reactants})$  (Kotz & Treichel, 1996). The equilibrium constant ( $K_r$ ) defines the extent to which a reaction occurs at a specific temperature ( $T$ , K). i.e.  $\Delta G(T) = -RT \ln[K_r(T)]$ ;  $R$  = gas constant. Endothermic reactions become more thermodynamically favourable at higher temperatures as  $\ln[(K_r(T_2))/(K_r(T_1))] = \Delta H_R/R[(1/T_2)-(1/T_1)]$ .  $\Delta G$  changes with product or reactant phase. Reactions with a negative  $\Delta G$  are spontaneous.



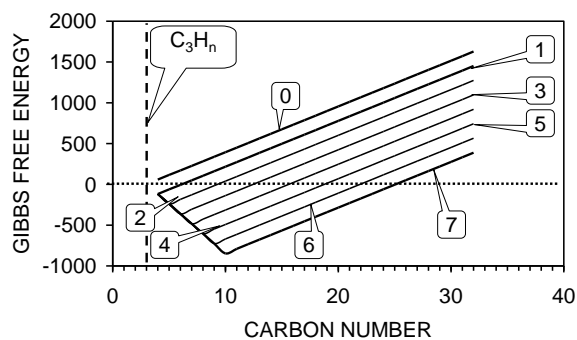
**Fig. 1: Heat of Reaction (kJ mol<sup>-1</sup> product alkane)**

Chain growth is from an alkane/alkene nucleus and occurs by adding  $a\text{CH}_4 + b\text{CO}_2$ , where  $b$  varies between 0 and 7, and  $a = n-m-b$ ;  $n$  = number of carbon atoms in the finished n-alkane;  $m$  = number of carbon atoms in the starting alkane.  $m$  for C<sub>3</sub>H<sub>8</sub> = 3.

In this study, [ $\Delta H$ ,  $\Delta H_R$ ,  $\Delta G$ ] are referenced to  $T = 298.15\text{K}$ , pressure,  $P = 0.1 \text{ MPa}$ .  $\Delta H$ ,  $\Delta H_R$ ,  $\Delta G$  vary with temperature.  $K_r$  varies with  $T$  and  $P$  and can be approximated (Antia, 2009b) as:  $K_p/K_c = (R P_{at} T)^c$ ;  $P_{at}$  = pressure (atmospheres);  $c$  = moles gaseous reactants – moles gaseous products;  $K_c$  = equilibrium constant at STP,  $K_p$  = equilibrium constant at  $P$  and  $T$ .

### 2.1 Low Temperature Thermodynamics

Reactions involving CO<sub>2</sub> and H<sub>2</sub> (at 423- 573 K, 0 – 5 MPa) are interpreted (Kuster, 1936a; 1936b; Puskas, 1997) as a two-stage Fischer Tropsch (FT) reaction where Stage 1 hydrogenates CO<sub>2</sub>. i.e. CO<sub>2</sub> + H<sub>2</sub> = CO + H<sub>2</sub>O (g)  $H_R^{298.15K} = 41.2 \text{ kJ mol}^{-1}$ ; CO<sub>2</sub> + H<sub>2</sub> = CO + H<sub>2</sub>O (l)  $H_R^{298.15K} = -2.8 \text{ kJ mol}^{-1}$ . The second stage follows a standard FT hydrogenation of CO to form oil. The first stage is exothermic when the H<sub>2</sub>O product is maintained as a liquid or supercritical fluid. The thermodynamics for chain growth on a C<sub>3</sub>H<sub>8</sub> nucleus by the addition of CO<sub>2</sub> and CH<sub>4</sub> (C<sub>3</sub>H<sub>8</sub> + bCO<sub>2</sub> + aCH<sub>4</sub> + bH<sub>2</sub> = C<sub>n</sub>H<sub>2n+2</sub> + cH<sub>2</sub>O + dH<sub>2</sub>; n=3+a+b) is both exothermic and spontaneous at  $T = 298.15\text{K}$  (Fig. 1, 2). The thermodynamics indicate that a mixture of catalyst (e.g. zero valent iron (Fe<sup>0</sup>), or Fe<sub>x</sub>O<sub>y</sub>), water, methane and



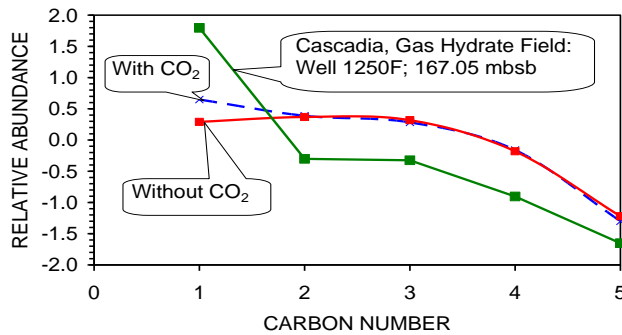
**Fig. 2: Gibbs Free Energy (kJ mol<sup>-1</sup> product alkane)**

For n-alkanes where chain growth by accretion is from an alkane/alkene nucleus and accretion occurs by adding  $a\text{CH}_4 + b\text{CO}_2$ . See Fig.1 caption for key.

CO<sub>x</sub> (or H<sub>y</sub>CO<sub>x</sub>) will result in the formation of hydrocarbons at ambient temperatures and pressures. H<sub>2</sub>=H<sub>2</sub>(aq, g), 2H<sub>3</sub>O<sup>+</sup> or 2H<sup>+</sup>.

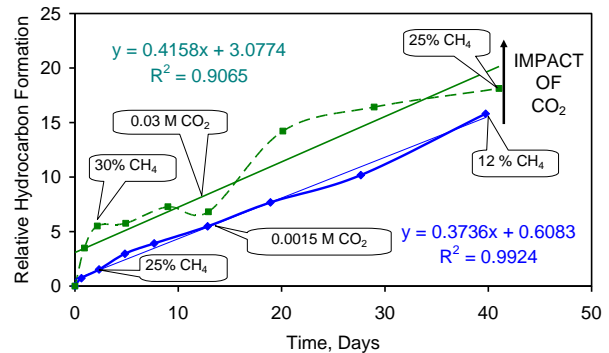
## 2.2 Formation of Hydrocarbons from Water, Carbon, and Fe<sup>0</sup>

Placement of 5gm iron filings (3.09 wt% C; surface area = 1.18 m<sup>2</sup>/g) + 15 ml pure H<sub>2</sub>O (0.0015 M CO<sub>2</sub> and 0.03M CO<sub>2</sub>) in sealed containers at 293 K demonstrate (Fig. 3, 4) a progressive increase in evolved hydrocarbons with time (Deng *et al.* 1997). After 40 days similar levels of ethane to pentane hydrocarbons were present. The methane product concentration was higher when high initial concentrations of dissolved CO<sub>2</sub> (Fig. 3) were present. The overall number of carbon atoms incorporated in hydrocarbon formation increases with increased CO<sub>2</sub> concentration (Fig. 4) and availability of carbon. Carbon incorporation into new alkanes/alkenes follows a linear relationship.



**Fig. 3: Iron Catalysed Hydrocarbons**

Formed in a sealed reactor (5 g Fe<sup>0</sup> + 15 cm<sup>3</sup> H<sub>2</sub>O) at 283 K (in CO<sub>2</sub> (0.0015 M) poor and CO<sub>2</sub> (0.03 M) enriched water) after 40 days. A typical gas sample from the Cascadia hydrate field (263 – 290 K, >9 MPa) shows a similar pattern of relative ethane to pentane molar abundance. The methane yield without CO<sub>2</sub> (0.0015 M) = 8.7 cm<sup>3</sup> kg Fe<sup>0</sup> 40 d<sup>-1</sup> = 282.8 cm<sup>3</sup> kg C 40 d<sup>-1</sup>; The methane yield with CO<sub>2</sub> (0.03 M) = 20.8 cm<sup>3</sup> kg Fe<sup>0</sup> 40 d<sup>-1</sup> = 674.5 cm<sup>3</sup> kg C 40 d<sup>-1</sup>; Data Sources: Deng *et al.* 1997; Antia, 2009b.

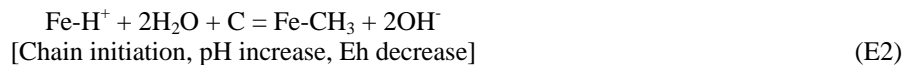
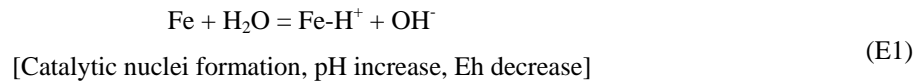


**Fig. 4: Iron Catalysed Hydrocarbons**

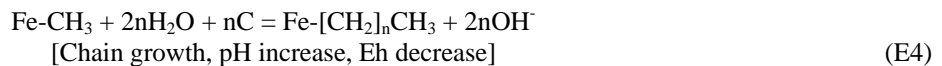
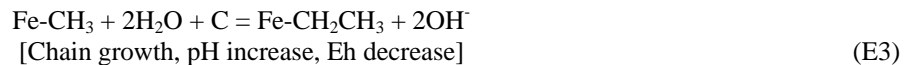
Formed at 283 K (in CO<sub>2</sub> poor and CO<sub>2</sub> enriched water). Data Source: Deng *et al.* 1997. Relative Hydrocarbon Formation = A<sub>1</sub> + 2A<sub>2</sub> + 3A<sub>3</sub> + 4A<sub>4</sub> + 5A<sub>5</sub>. A<sub>1</sub> to A<sub>5</sub> = molar abundance of methane to pentane.

## 2.3 Interpreted Hydrocarbon Formation Reactions

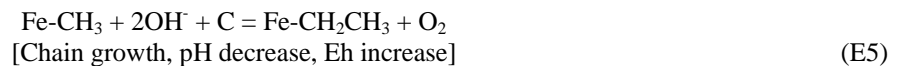
The interpreted primary oxidative and hydrogenation chain growth reactions (Baxendale *et al.* 1946a; 1946b; Mwebi, 2005; Antia, 2009a; 2009b; 2010; 2011; Braun, 2009; Kang & Choi, 2009) are:

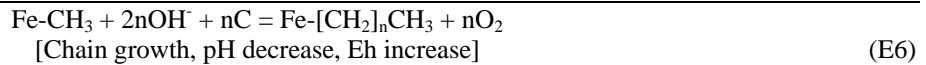


Chain growth on a specific site may be associated with an increase in OH<sup>-</sup> and a corresponding increase in pH.

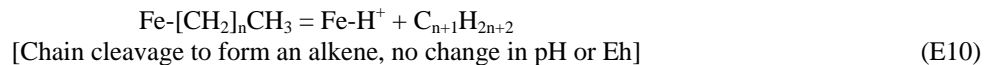
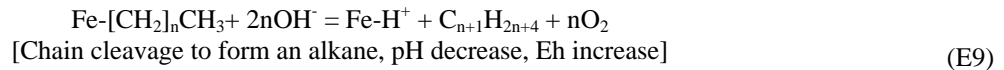
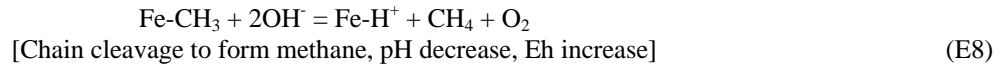
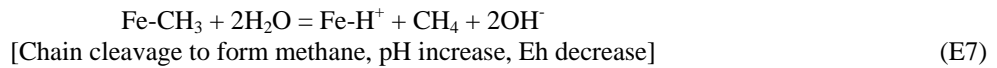


Alternatively, chain growth is associated with a decrease in OH<sup>-</sup> and a corresponding decrease in pH.

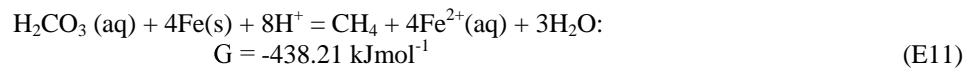




The O<sub>2</sub> is present as O<sub>2</sub>(g), O<sub>2</sub>(aq), O<sub>2</sub>H<sup>-</sup> (aq), Fe[OH]<sub>x</sub>, Fe<sub>x</sub>O<sub>y</sub> (Antia, 2010; 2011). Chain cleavage to release the hydrocarbon may take the form:



The effect of CO<sub>2</sub> is to reduce the availability of H<sup>+</sup> ions, thereby increasing pH (Deng *et al.* 1997),



Consequently sediments containing biogenic activity (or carbonates) can be expected to contain high concentrations of Fe catalysed methane while the relative molar abundance pattern for catalysed ethane to pentane remains unchanged. Similar ethane – pentane molar abundance patterns are present in the Cascadia gas hydrate field (East Pacific continental slope) (Fig. 3). The oxidation-reduction reactions associated with hydrocarbon formation are oscillatory Fenton Reactions (Barbusinsky, 2009; Antia, 2010; 2011). They may also involve an electron shuttle mechanism (Antia, 2010; 2011; Bokare & Choi, 2009; Kang & Choi, 2009).

### 3. EXPERIMENT I

Experiment I establishes if the hydrocarbon chain growth and chain termination (E1 to E10) resulting from the interaction of Fe<sup>0</sup> and H<sub>2</sub>O is associated with the formation of OH<sup>-</sup>, or O<sub>2</sub>H, or O<sub>2</sub> or Fe[OH]<sub>x</sub> or Fe<sub>x</sub>O<sub>y</sub>.

#### 3.1 Reactors

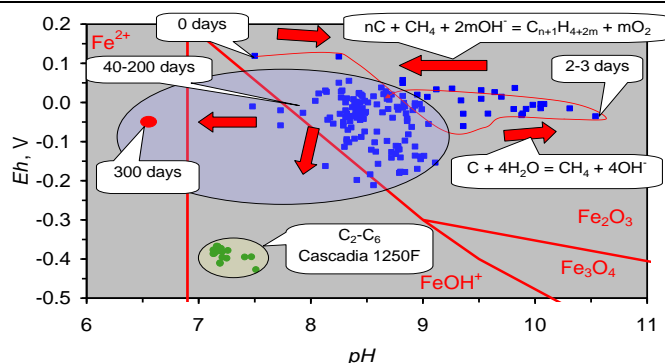
- a. *Base Case Sealed Static Diffusion Reactor*: 60 g/L Ca-montmorillonite + 90 g/L n-Fe<sup>0</sup> (60-120 K nm) were placed in 0.2 L of 0.96 PSU water. Temperature was maintained between 263 and 298 K for a period of 300 days.
- b. *Control Case Sealed Static Diffusion Reactor*: 60 g/L Ca-montmorillonite was placed in 0.2 L of 0.96 PSU water. Temperature was maintained between 263 and 298 K for a period of 300 days.

#### 3.2 Measurements

Measurements were made of Eh, pH, and temperature using Hanna HI-98120 (ORP/Temperature, Hanna Calibration Number = 24230), Hanna HI-98129 (pH, EC, TDS, Temperature; Hanna Calibration Number = 33999). HI-98120 is factory calibrated. HI-98129 was calibrated using Hanna test reagents (pH buffer solutions 4.01, 7.01, 10.01 (HI-7004, HI-7007, HI-7010) and conductivity solution 1.413 mS cm<sup>-1</sup> (HI-7031)).  $Eh_{298.15K}$  calculated from ORP as  $Eh = \text{ORP} (298.15/T_o) [1]$ . T<sub>o</sub> = measured temperature, K. The evolution of gases (O<sub>2</sub>, C<sub>x</sub>H<sub>y</sub>, H<sub>2</sub>) was monitored.

#### 3.3 Eh-pH Changes –Base Case

The initial formation of CH<sub>4</sub>, Fe-H<sup>+</sup> and Fe-CH<sub>3</sub> (Fig.3,4) is associated (Fig. 5) with a very significant increase in pH (i.e. ratio of OH<sup>-</sup>:H<sup>+</sup>). This is associated with a decrease in Eh (i.e. a decrease in the O<sub>2</sub>H:OH ratio). These experimental observations are consistent with equations E1, E2 and E7.

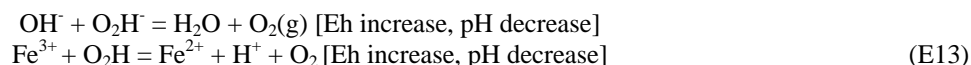
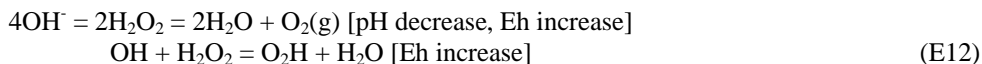


**Fig. 5: Iron Catalysed Changes**

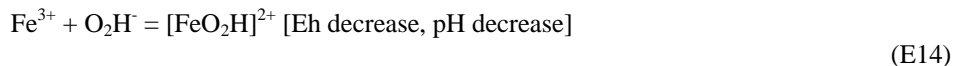
In water Eh & pH associated with hydrocarbon formation at 255 – 298 K at 0.1 MPa in a sealed static diffusion reactor containing 0.2 L saline H<sub>2</sub>O. The Eh & pH of sediments which have been interpreted (Antia, 2009b) as containing in situ ethane to hexane (Cascadia, gas hydrate field (263-290 K), ODP Leg 204, well 1250 F, 155 – 179 mbsb; water depth = 795.4 m) are shown for comparison.

Chain growth and the progressive formation of higher hydrocarbons (Fig. 3, 4) is associated with a decrease in pH and a stable or decreasing Eh (Fig. 5). This indicates consumption of OH<sup>-</sup> to reduce pH and a faster consumption of O<sub>2</sub>H than OH<sup>-</sup> to reduce Eh. These experimental observations suggest that equations E3, E4 only apply to a minor extent, and that E5, E6, E9, and E10 dominate during the formation of higher hydrocarbons. During this phase CH<sub>4</sub> may be formed in accordance with either E7 or E8.

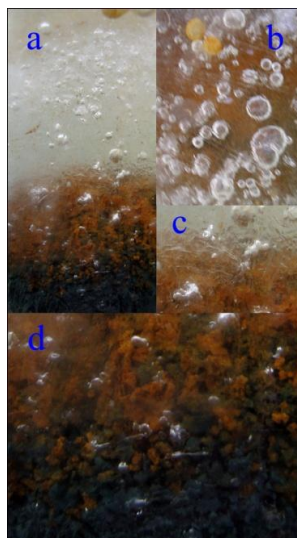
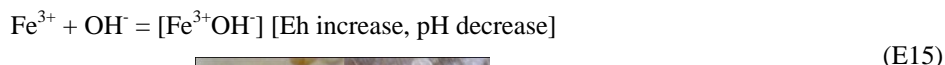
O<sub>2</sub>H and OH are relatively unstable ions. The relevant reactions (Barbusinski, 2009; Bokare & Choi, 2009; Kang & Choi, 2009; Antia, 2010) include:-



The characteristic deep brown-red iron complexes (Fig. 6a, d) are an ion pair complex (E14) formed (Evans *et al.* 1949; George, 1952; Behar & Stein, 1966) as:-



The colour due to the hydroxyl reaction (E15) is negligible (Evans *et al.*, 1949):

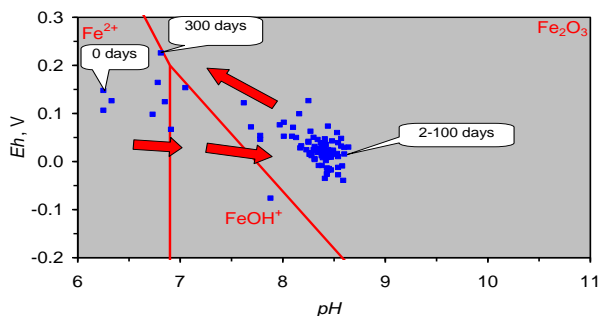


**Fig. 6: Iron Catalysed Hydrocarbon and Oxygen Formation**

At 273 K at Day 300 (a) Vertical section through the sediment-water interface.  $O_2 + H_2 + CH_4/C_{1+}$  degassing from n- $Fe^0$  (black) through  $Fe_2O_3/Fe(OH)_x + Ca$ -montmorillonite to  $H_2O$  (s). (b) View from upper surface of water showing  $Fe_2O_3 / Fe(OH)_x$  (principally  $[FeO_2H]^{2+}$  (goethite/limonite group of minerals)) +  $Ca$ -montmorillonite in some of the expelled ascending gas bubbles. (c) Detail of the sediment-water interface showing turbulent effect of micro-vent degassing. (d) Detail of the sediment ( $Fe^0$  overlain by  $Ca$ -montmorillonite +  $Fe_2O_3/Fe(OH)_x$ ) showing the close relationship between gas bubble formation and  $Fe_2O_3/Fe(OH)_x$ ; width about 2 cm. The largest bubbles have diameter of 2-4 mm. Some ice/ice slush is present in the water column due to desalination associated with the presence of  $Fe^0$  (Antia, 2010). Date 11/1/11.

During periods of pH decline and/or Eh decline, significant  $O_2(g)$  degassing (Antia, 2010) may accompany hydrocarbon formation (Fig. 6). Part of the  $O_2$  discharge is associated with the formation of  $Fe(OH)_x$  and  $Fe_2O_3$  (Fig. 6). These observations suggest that methane formation is associated with increasing alkalinity while higher alkane formation is associated with decreasing Eh and/or decreasing pH.

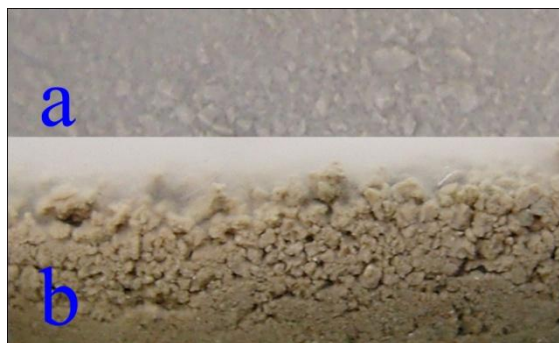
### 3.4 Eh-pH Changes – Control Case



**Fig. 7: Control Changes**

In water Eh & pH associated with addition of 60 g/L  $Ca$ -montmorillonite at 263 – 298 K at 0.1 MPa in a sealed static diffusion reactor containing 0.2 L saline  $H_2O$ .

The control sample showed an initial minor increase in pH as the water adjusted to the presence of Al and Ca ions associated with the  $Ca$ -montmorillonite (Fig. 7). Over a period of 300 days the pH increase was reversed and the Eh increased (relative to the initial Eh). The final Eh was about 0.25 mV higher (Fig. 7) than in the base case containing  $Fe^0$  (Fig. 5). Significantly no  $O_2(g)$ ,  $H_2(g)$  or  $CH_4/C_{1+}(g)$  bubbles (Fig. 6) were observed in the control reactor (Fig. 8).



**Fig. 8: Control Sample**

60 g/L  $Ca$ -montmorillonite, 272 K at Day 300. (a) plan view through water to clay surface showing an absence of bubbles; width c.3 cm. (b) vertical section through the clay/water interface showing an absence of bubbles and a well developed network of clods and macropores at the clay/ $H_2O$  interface; width about 2 cm: Date 11/1/11.

## 4. EXPERIMENT II

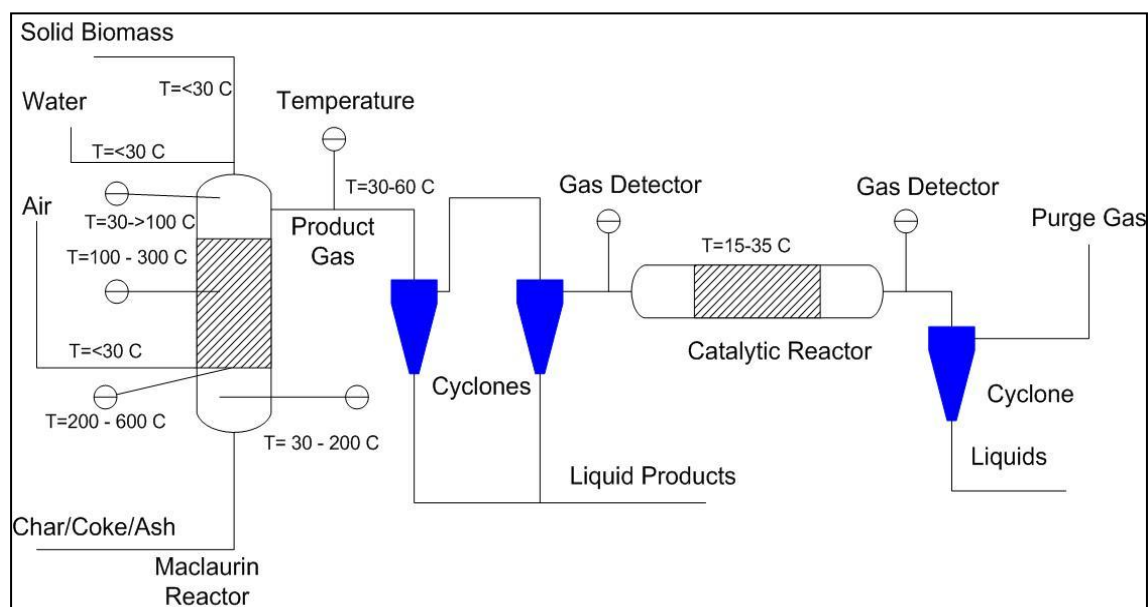
Experiment I established that  $Fe^0$  can catalyse the formation of light hydrocarbons at 0 – 30 C in an aqueous environment. However, most sedimentary sequences only contain minor concentrations of Fe, typically as  $Fe^{2+}$  ions,  $FeOH^+$ ,  $Fe_2O_3$  and  $Fe[OH]_x$ . These sequences may contain high concentrations of NaCl, either in solution or as halite. Halite, when combined with a zero valent metal ( $Fe^0$  or  $Ni^0$ ) will catalyse hydrocarbon formation from  $CO_2$ , CO and  $H_2$  (Gagnon, 2003; 2004; Antia, 2008b). Thermal and biogenic decomposition of organic matter releases  $CO_2 + CO + C_xH_yO_z$  (formaldehydes, acetates) +  $HCO_3^- + CH_4 + H_2$  (e.g. Verhaart *et al.* 2010). Experiment II considers

whether NaCl is able to facilitate CO<sub>2</sub> removal from a flowing gas containing CH<sub>4</sub>, and minor quantities of CO, H<sub>2</sub>, H<sub>2</sub>O, C<sub>x</sub>H<sub>y</sub>, C<sub>x</sub>H<sub>y</sub>Cl<sub>z</sub> and C<sub>x</sub>H<sub>y</sub>O<sub>z</sub> produced by the thermal decomposition of organic matter at low temperatures.

#### 4.1 Low Temperature Thermogenic Gas Composition

Low temperature (373-673 K) thermal decomposition of organic matter (in the presence of CO, CH<sub>4</sub>, CO<sub>2</sub>, O<sub>2</sub> or H<sub>2</sub>O) produces a gas with a high CO<sub>2</sub>:CO ratio and a high H<sub>2</sub>O:H<sub>2</sub> ratio (Gaur & Reed, 1998). Prolonged heating periods result in almost complete gasification of the organic matter, leaving a residual char/ash. The rate of organic decomposition is accelerated by increasing temperature.

#### 4.2 Reactor and Process Flow

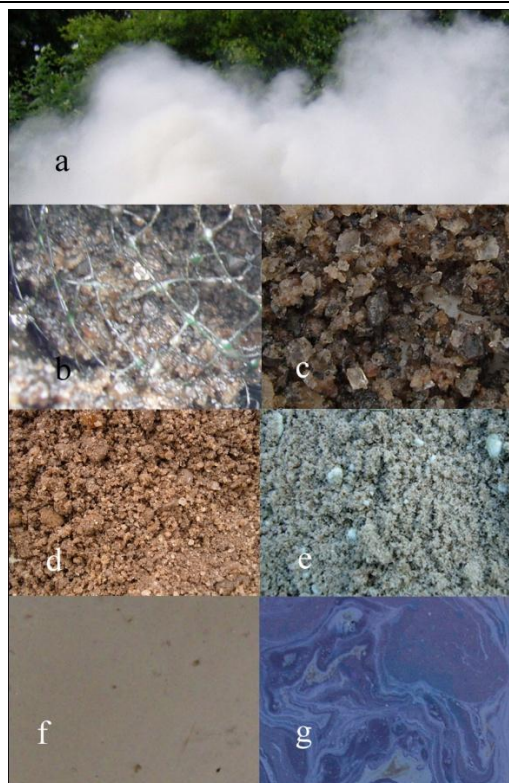


**Fig. 9: Process Flow for Experiment II**

The thermogenic feed gas was constructed using a non-catalytic, internally heated, Maclaurin carbonisation reactor (Maclaurin, 1915) fed by angiosperm wood and foliage. The dominant gaseous products are H<sub>2</sub>O, CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>, with minor quantities of CO, H<sub>2</sub>, C<sub>x</sub>H<sub>y</sub>, C<sub>x</sub>H<sub>y</sub>Cl<sub>z</sub> and C<sub>x</sub>H<sub>y</sub>O<sub>z</sub>. Solid waste (char + ash) formed 10-30% by weight of the product. Entrained liquids and gaseous products (estimated effective reactor raw gas discharge rate = <0.10 - >60 Nm<sup>3</sup> hr<sup>-1</sup>) were passed through four cyclones/condensers, placed in series, to remove entrained and condensable liquids (Fig. 9). The residual gas (estimated 1 - 40% of raw gas volume) was passed through a packed bed reactor (0.06m internal diameter) containing 0.25 kg NaCl catalyst. The residual gas was passed through a cyclone/condenser, prior to purge. The CO<sub>2</sub> and CO content of the gas entering and leaving the catalyst bed was measured [Anton Telegan Sprint VI. Measurement limits: O<sub>2</sub> = 0-25%; CO<sub>2</sub> = 0-25%; CO - direct measurement = 0-0.1%, CO:CO<sub>2</sub> ratio measurement = 0 - 0.99 (for CO concentrations of 0.1 - 7%); T = -50-1100 C; Poole Instrument Calibration Ltd. Calibration Certificate No. 109406]. Measured O<sub>2</sub> readings were typically <0.2%. The process flow and associated flowline temperatures are summarised in Fig. 9. The solid feedstock is maintained at a temperature of <300 C for the bulk of its residence in the reactor.

#### 4.3 NaCl Catalyst

Halite granules (Triassic, Cheshire) [0.5-5 mm diameter] are used as the catalyst, or were mixed (50 wt%) with Fe-montmorillonite or FeSO<sub>4</sub>, or Mg<sub>2</sub>SO<sub>4</sub>, or K<sub>2</sub>SO<sub>4</sub>, or CaCO<sub>3</sub>, or organic matter, or pyroclastics or ash (e.g. Fig. 10b,c,d,e). The halite contains, and is coated by, Fe<sub>x</sub>O<sub>y</sub> and Fe<sub>x</sub>[OH]<sub>y</sub> (Figure 10c). Details of catalyst manufacture are provided elsewhere (Antia, 2008b).

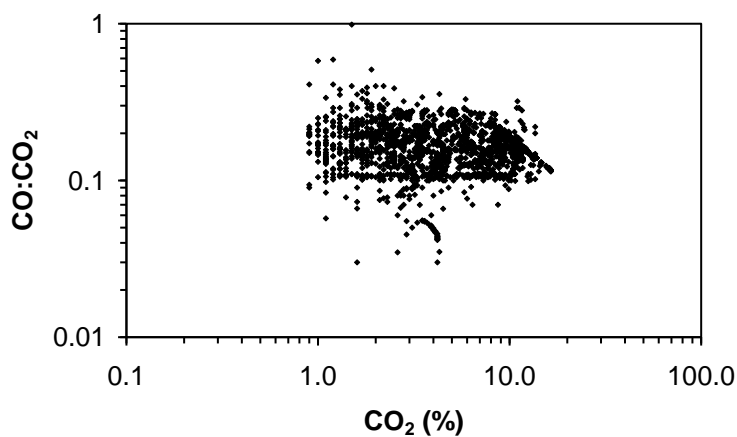


**Fig. 10: Low Temperature Oil Formation**

(a) Raw gas vented from the Maclaurin Reactor. (b) Halite catalyst in operation within the reactor showing H<sub>2</sub>O formation on the catalyst surface. (c) Halite catalyst removed from the reactor contained bitumens; (d) Halite + Fe-montmorillonite catalyst; (e) Halite + Fe<sub>2</sub>SO<sub>4</sub> catalyst. (f) solid alkanes produced by the halite catalyst. (g) liquid alkanes/hydrocarbons produced by the halite catalyst.

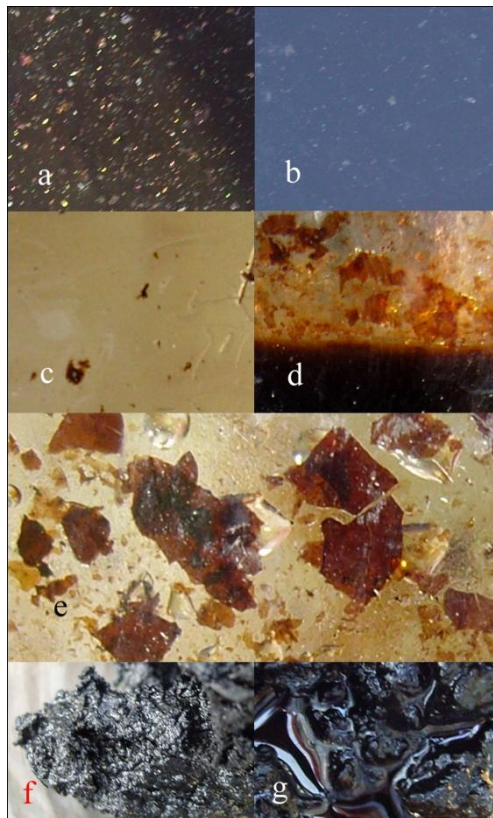
#### 4.4 Test Results: Maclaurin Reactor

The Maclaurin reactor (Fig.9) produced a product gas (Fig. 10a) containing entrained particulates (H<sub>2</sub>O, C<sub>x</sub>H<sub>y</sub>, C<sub>x</sub>H<sub>y</sub>O<sub>z</sub>). The gas contained 1-20% CO<sub>2</sub> (Fig. 11). The cyclones (Fig. 9) recovered a number of different coloured organic compound-water mixtures (Fig. 12a-e) [toluene, alkyl acids, phenols, furans]. Aromatic crystals exsolved in the liquids during storage (Fig. 12a-e). Minor quantities of bitumen (Fig. 12f) and paraffinic wax (Fig. 12g) were produced. The gases leaving the downstream cyclone were essentially free of entrained particulates.



**Fig. 11: CO<sub>2</sub> and CO Content of Product Gas from the Maclaurin Reactor**





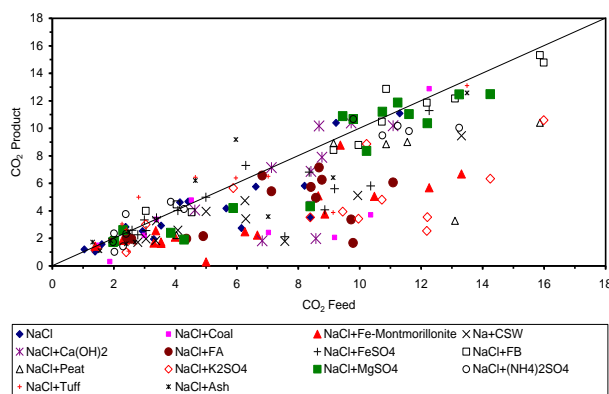
**Fig. 12: Primary Products**

From the Maclaurin Reactor. (a-c) Liquids (organic chemicals + water) containing crystallised aromatics. (d) Density separated liquid containing principally phenols + water overlain by water containing crystallised aromatics. (e) Examples of crystallised aromatics. Field of view = 3 cm. (f) Tar/bitumen, (g) Paraffinic waxes + aromatics.

#### 4.5 Test Results: NaCl Packed Bed Reactor

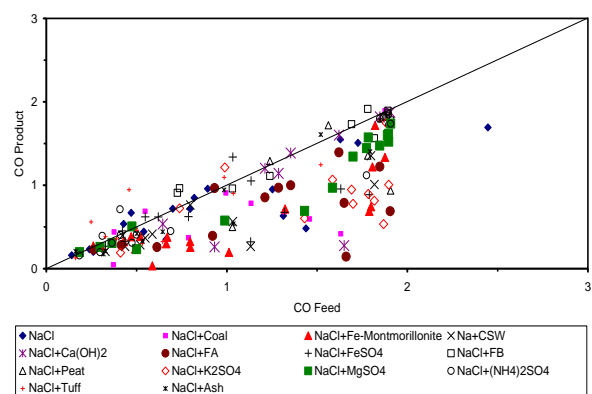
It was observed that the passage of CO<sub>2</sub>, CO through the catalyst bed (Fig.9; 10b,c,d) changed the molar abundance distributions [average (Fig. 13, 14, 15), standard deviation (Fig. 16), skewness (Fig. 17) and kurtosis (Fig.18) ].

#### 4.6 Interpretation of NaCl Test Results



**Fig. 13: Average Molar CO<sub>2</sub> Abundances (%) of Feed Gas and Product Gas**

Number of Measurements = 8740. Analysis is not corrected for the decrease in product gas volume relative to feed gas volume. Temperature = 16 – 35 C; Pressure = 0.1 MPa.



**Fig. 14: Molar CO Abundances (%) of Feed Gas and Product Gas**

CSW = nanoporous calcified algae skeletons; FA = organic rich nitrate fertiliser; FB = organic rich phosphate fertiliser; Ash = ash from the Maclaurin Reactor.

The molar abundances (Fig. 13, 14, 15) indicate that CO + CO<sub>2</sub> is removed from the feed gas within the catalyst bed. The feed gas discharge rate and feed gas composition from the Maclaurin reactor (Fig. 9) is in a state of constant flux (Fig. 11, 16-18). This variation creates a disequilibria within the catalyst bed. The surface of active NaCl catalyst within the catalytic reactor (Fig. 10b) is coated with product water. The solubility of CO<sub>2</sub> in this water is less than 0.03 M kg<sup>-1</sup> (Duan and Sun, 2003) and is unlikely to account for >0.002 cm<sup>3</sup> CO<sub>2</sub>. The CO<sub>2</sub> volumes removed are greater than the adsorption capacity of the NaCl (Saber, 1996).

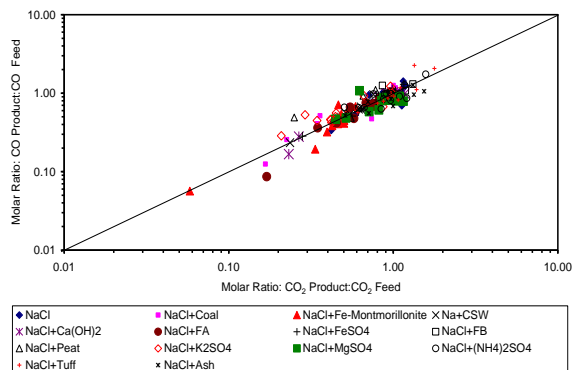


Fig. 15: Molar Product:Feed Ratios Assuming 1 Mole Feed Gas = 1 Mole Product Gas

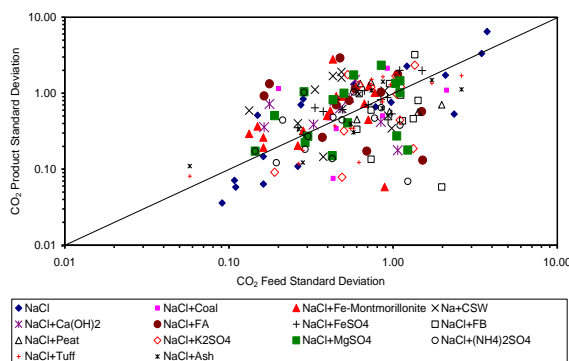


Fig. 16: Standard Deviation of Molar CO<sub>2</sub> Abundances of Feed Gas and Product Gas

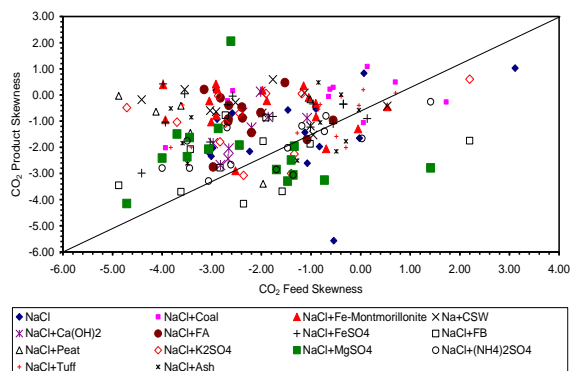


Fig. 17: Skewness Associated with Molar CO<sub>2</sub> Abundances of Feed Gas and Product Gas

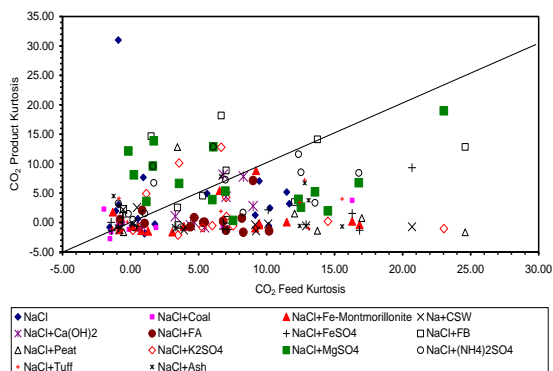
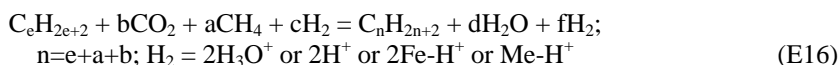


Fig. 18: Kurtosis Associated with Molar CO<sub>2</sub> Abundances of Feed Gas and Product Gas

The downstream collection cyclone (Fig. 9) contained a thin film of wax (Fig. 10f), minor quantities of oil (Fig. 10g) and no water. Bitumen was recorded within the spent catalyst (Fig. 10d). Fig. 1, 2 demonstrate that reaction E16 is both exothermic and spontaneous at atmospheric pressure and 298K. The observed water film on the NaCl (Fig. 10b) is interpreted as a reaction product (E16). The absence of water in the downstream cyclone (Fig. 9) is interpreted as indicating that any water produced via E16 is removed by reactions E1 to E10 (where Fe is substituted as appropriate by another cation).



## 5. CONCLUSIONS

The traditional oil formation model links hydrocarbon generation directly to the heating duration, heating rate and temperature of organic source rocks. It remains the most likely hydrocarbon generation model for most commercial oil and gas accumulations. However, most immature sediments ( $R_o < 0.6$ ) at temperatures of 0 – 100 C contain extractable C<sub>5+</sub> hydrocarbons and varying concentrations of C<sub>1</sub>-C<sub>5</sub> gaseous hydrocarbons. While some of these hydrocarbons can be explained by migration from a deeper more mature source rock or formation as a result of microbial activity, the origin of the majority of these dispersed, low concentration, hydrocarbons is unknown.

Zero Valent Iron (ZVI) water treatment studies have demonstrated that alkane formation can be catalysed at  $<30\text{ C}$  using  $\text{Fe}^0$  + water and a carbon source (Fig. 3, 4). 1 kg of C has the potential to produce a maximum of  $1.8\text{ m}^3\text{ CH}_4$ . The reaction rates measured in Fig. 3 indicate that this ultimate potential (if the reactions run to completion) may naturally be achieved within a short time period (weeks to years). Experiment I established the Eh and pH associated with this hydrocarbon generation and possible reaction routes. This experiment established that anoxic alkaline pore water is more likely to catalyse methane formation, while anoxic neutral pore waters are more likely to catalyse the formation of  $\text{C}_{2+}$  hydrocarbons.

Hydrocarbon accumulations are commonly associated with saline pore waters and in some locations (e.g. Middle East, North Sea, Gulf of Mexico) with halite deposits (and halite diapers). Halites conventionally form part of a carbonate-evaporite sequence (of limestones, gypsum, halite and poly-halite). The associated oils and gases may have high  $\text{H}_2\text{S}$  concentrations (derived from gypsum) but commonly have very low  $\text{CO}_2$  concentrations. Experiment II established that NaCl is able to facilitate the low temperature incorporation of  $\text{CO}_2$  (within a migrating hydrocarbon gas) into longer chain hydrocarbon molecules (through Equations E1-E16) at  $<308\text{ K}$ . Experiment I has established that  $\text{Fe}^0$ ,  $\text{Fe}_x\text{O}_y$ ,  $\text{Fe}(\text{OH})_x$  minerals can catalyse the incorporation of  $\text{CO}_2$  (and  $\text{H}_x\text{CO}_x$ ) and  $\text{H}_2\text{O}$  into hydrocarbons at these temperatures (through Equations E1–E12) under Eh and pH conditions which are common in many reservoirs (and hydrocarbon migration pathways). These observations may indicate that some of the dispersed  $\text{C}_{1-8+}$  hydrocarbons found in immature low temperature sediments (in the continental slope and abyssal plain) have an in situ catalytic origin, and may not indicate the presence of migrated hydrocarbons. The distribution of these low temperature hydrocarbons will be independent of sediment temperature, but may be directly linked to sediment facies/mineralogy and associated pore water chemistry.

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