Forest-Tundra Neighbouring the North Pole: Plant and Insect Remains from the Plio-Pleistocene Kap København Formation, North Greenland

O. BENNIKE1 and J. BÖCHER2

(Received 19 June 1989; accepted in revised form 22 February 1990)

ABSTRACT. The Kap København Formation in northeast Peary Land, Greenland, is believed to be 2.0-2.5 million years old, i.e., from the Plio-Pleistocene transition. The dating is primarily based on biostratigraphical correlation of lower marine fauna and a few fragments of terrestrial mammals. Although deposited in marine and coastal environments, the sediments contain abundant remains of terrestrial and limnic organisms. This paper examines macroscopic plant and insect remains. About 60 taxa of vascular plants and 120 insect taxa have so far been identified. Nearly all of the named insect species are extant, extralimital forms, generally of a recent subarctic/boreal and more or less circumpolar distribution. The species composition shows that upland areas were covered with forest-tundra and heathland and that mesotrophic, well-vegetated lakes and a number of other wetland localities existed in the area. The presence of arctic plants in the formation puts some time constraints on their origin.

Key words: Plio-Pleistocene, North Greenland, palaeoecology, forest-tundra

RÉSUMÉ. La Formation de Kap Kobenhavn dans le nord-est de la terre de Peary, au Groenland, daterait de 2,0 à 2,5 Ma, c'est-à-dire de la transition du Plio-Pléistocène. La datation repose essentiellement sur une corrélation biostratigraphique de la faune marine inférieure et aur quelques fragments de mammifères terrestres. Même s'ils se sont déposés dans des milieux marins et côtiers, les sédiments contiennent d'abondants vestiges d'organismes terrestres et limniques. La présente communication examine des vestiges macroscopiques de plantes et d'insectes. Environ 60 taxons de plantes vasculaires et 120 taxons d'insectes ont jusqu'à ce jour été identifiés. Presque toutes les espèces d'insectes identifiés sont des formes exotiques existantes qui ont en général une distribution subarctique/boréale et plus ou moins circumpolaire récente. La composition des espèces indique que les terres hautes étaient recouvertes de forêt-toundra et de lande, et que la région était parsemée de lacs mésotrophes à végétation abondante et de plusieurs autres terres humides. La présence de plantes arctiques dans la formation en précise l'origine dans le temps.

Mots clés: Plio-Pléistocène, Groenland septentrional, paléoécologie, forêt-toundra

РЕФЕРАТ. Формация Кап Кебенхавн в северо-восточной части Земли Пири, Гренландия, насчитывает, предположительно, 2-2,5 миллиона лет, т. е. относится ко времени плио-плейстоценового перехода. Эта датировка основана главным образом на биостратиграфической корреляции низшей морской фауны и нескольких остатков наземных животных. Составляющие формацию осадки изобилуют остатками наземных и озерных организмов, хотя их накопление происходило в обстановке морских и прибрежных районов. В настоящей статье рассматриваются макроскопические остатки растений и насекомых. На настоящее время идентифицировано около 60 таксонов сосудистых растений и 120 таксонов насекомых. Почти все из идентифицированных насекомых относятся к ныне существующим, экстразональным формам, встречающимся в современную эпоху, как правило, в субарктиче-ских/северных и более или менее циркумполярных районах. Видовой состав указывает на то, что плоскогорья были покрыты в то время лесотундрой и вересковыми пустошами и что в этом районе существовали мезотрофные, заросшие растительностью озера и заболоченные земли. Присутствие в формации арктических растений накладывет на них некоторые временные ограничения.

Ключевые слова: плиоплейстоцен, Северная Гренландия, палеоэкология, лесотундра

INTRODUCTION

While Neogene and Early Pleistocene sediments (primarily the Beaufort Formation) containing evidence of the former terrestrial and limnic environments have been known for several decades in northern Canada (see review in Matthews, 1987; Matthews and Ovenden, 1990-this issue), similar sediments (the Kap København Formation) have only recently been located in Greenland (Funder and Hjort, 1980; Fig. 1). The Kap København Formation is of great significance with regard to Late Cenozoic high latitude biological and climatic evolution

This paper primarily discusses macroscopic remains of vascular plants and insects in the Kap København Formation, latitude 82°30'N (see also Böcher, 1989, and Bennike, in press). Other fossil terrestrial and limnic organisms present in the sediments include plants such as various thallophytes (Bennike, in press) and mosses (Mogensen, 1984) and animals such as cladocerans, freshwater bryozoans, notostracas, oribatid mites and vertebrates (Fredskild and Røen, 1982; Repenning et al., 1987; this paper).

The Kap København Formation (Fig. 2), which covers about 300 km² and is exposed at elevations up to 230 m above sea level, is divided into two members (Funder *et al.*, 1984).

The lower (A) is up to 50 m thick and rather homogeneous, consisting of laminated marine mud with scattered dropstones and lenses of non-sorted sediment (diamicton) interpreted as representing evidence for nearby glaciation. This member contains a sparse mollusc and Foraminifera fauna, indicating arctic/high arctic conditions. No macroscopic plant or insect remains are present; the pollen flora is depauperate and dominated by *Betula*.

The upper member (B) is up to at least 70 m thick, rather heterogeneous but dominated by coastal and nearshore sand (units B1 and B3). It also contains an offshore silt unit (B2) at about 100 m above sea level with abundant and diverse in situ marine faunas (molluscs, foraminifers [Feyling-Hanssen, in press] and ostracodes [Brouwers et al., unpubl. ms]). The bivalve mollusc Cyrtodaria kurriana may point to a nearby output of fresh water.

That rivers did indeed enter the area is clearly shown by accumulations of detrital organic matter in parts of member B. Only a few smaller such accumulations have been located below unit B2 (the silt unit), while above the silt unit, in unit B3, organic-rich lenses and layers are widespread. They include logs of small, stunted trees.

At present the Kap København area supports a flora of about 70 species of vascular plants (Bennike, unpubl.). The

¹Geological Museum, University of Copenhagen, Øster Voldgade 5-7, DK-1350 Copenhagen K, Denmark

²Zoological Museum, University of Copenhagen, Universitetsparken 15, DK-2100 Copenhagen Ø, Denmark

[©]The Arctic Institute of North America

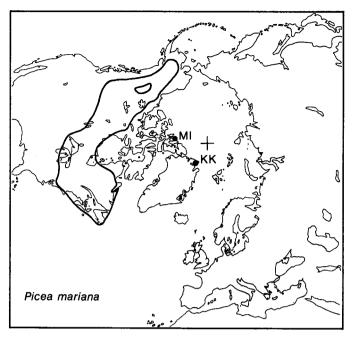


FIG. 1. The location of the Kap København Formation (KK) in North Greenland and the present range of *Picea mariana* (black spruce) (modified after Little, 1971). Also shown is the location of Meighen Island (MI).

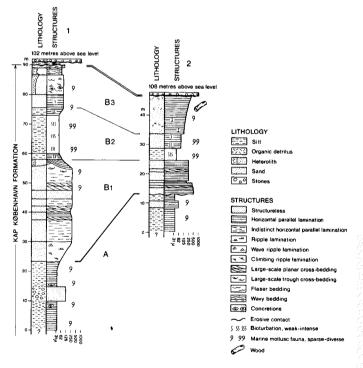


FIG. 2. Stratigraphy of the Kap København Formation to the southeast (left) and north (right) (adapted from Funder et al., 1984, with permission).

only woody plants are Salix arctica, Dryas and Cassiope tetragona. The latter forms small patches of heath in a sheltered area. Along the coast no woody plants can thrive, and the vegetation is dominated by Papaver radicatum and Saxifraga oppositifolia. There are no climatic data from the area, but as judged from the vascular plant flora the mean temperature for the warmest month is 3.5-4°C in the sheltered area with Cassiope heaths (Edlund, 1983). The present insect fauna is dominated by chironomids and flies. In spite of

intensive hunting, no beetles were caught in the area. The lakes of the area are ultra-oligotrophic.

Reconnaissance field work was carried out at Kap København in 1979 and 1980. More detailed work was conducted in 1983 (Funder *et al.*, 1984) and in 1986 by the present authors (Bennike, 1987).

METHODS

Most of the plant and animal remains were isolated by washing the samples through two sieves (plants: meshes of 0.42 and 0.21 mm; insects: meshes of 0.6 and 0.35 mm). Most samples dispersed well in water, but a few were treated with cold 5% KOH. After sieving, the insect remains were concentrated by flotation with kerosene and subsequent washing with detergent (Coope, 1986). Plant and insect remains were sorted out under a dissecting microscope. The plant material and part of the insect material is kept in 70% alcohol, but a large fraction of the identifiable insect fragments are mounted dry on standard cardboard rectangles on insect pins.

Both plant and insect remains are represented by fragments, which are generally well preserved. Small and thin-walled seeds and fruits are poorly preserved and presumably underrepresented. The same is the case with fragile insects (e.g., Diptera and Lepidoptera).

RESULTS

A preliminary list of plant and insect remains in the Kap København Formation is presented in Appendices 1 and 2. Some plant and insect remains are shown in Figures 3, 4 and 5.

All but the following three of the plant remains are classified as extant species. The larch Larix groenlandii is closely related to Larix occidentalis (western larch) of the Rocky Mountains. Myrica arctogale is closely related to the extant species M. gale (sweet gale) and the extinct species M. eogale described from the Neogene of Siberia (V.P. Nikitin, 1976). Two fossils are classified as Aracites globosa. These extinct plants are described in Bennike (in press).

As usual regarding subfossil assemblages, beetles (Coleoptera) dominate the insect record due to the highly

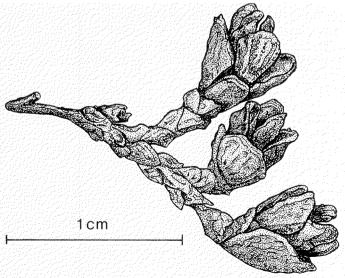


FIG. 3. Twig of *Thuja occidentalis* (white cedar) with scaly leaves and three small cones.

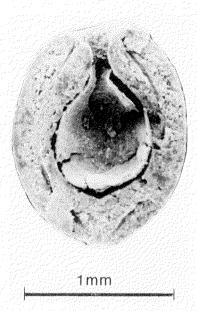


FIG. 4. SEM photo showing internal view of sectioned specimen of Aracites globosa.

resistant exoskeleton of most species. Among the Coleoptera the ground beetles (Carabidae) are prevalent. It is not apparent from the list that the rove beetles (Staphylinidae) are at least as well represented, but until now only a minor part of the numerous fragments have been identified. This is also true to some degree with the predaceous diving beetles (Dytiscidae).

It has so far been possible to identify 38 species and 65 genera (27 families, 6 orders) of insects. All fossil species are classed with extant species, but unidentified material might include extinct taxa.

Only a few logs exceed 12 cm in cross section and a few metres in length (maximum is 18 cm in cross section and 460 cm in length). Growth rings are extremely narrow, about 0.3 mm in width on average, but highly variable. Logs of Larix (n=76), *Picea* (n=22) and *Betula* (n=1) have been identified. A number of insects (Hadrobregmus, Pissodes, Rhyncolus, Scolytus, Pityophthorus, Camponotus herculeanus) are intimately associated with trees. Camponotus herculeanus (carpenter ant) is confined to forests or at least to the presence of trees. It avoids the dark and cool interior of forests, preferring sunny edges and stony slopes. The nest is almost exclusively built in living or dead trees, both in conifers and hardwoods (Larsson, 1943; Collingwood, 1979). Hadrobregmus is confined to dead, dry wood.

Fly holes and galleries in wood reveal the presence of still more arboricolous forms (Cerambycidae, Buprestidae, Urocerus, Megastigmus). A few species prefer woodland but are not dependent on trees (Notiophilus biguttatus, Bembidion grapii) (Lindroth, 1961-69, 1985-86).

In addition to Larix groenlandii and Picea mariana, boreal upland plants include Thuja occidentalis (Fig. 3), Taxus sp., Cornus spp. and Viburnum cf. edule. Common dwarf shrubs are Betula nana, Dryas octopetala and Vaccinium uliginosum. Insect species such as Notiophilus aquaticus, Blethisa catenaria, Diacheila polita, Miscodera arctica, Pterostichus brevicornis, P. vermiculosus and Agonum exaratum prefer open tundra or fairly xeric heathland (Lindroth, 1961-69, 1985-86).

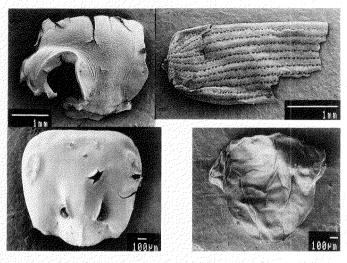


FIG. 5. SEM photos of some insect remains from the Kap København Formation: a) head of a Cicindela species (tiger beetle); b) left elytron of a Hedrobregmus species (death watch beetle) probably closely related to H. pertinax (L.), but with smaller and more dense strial punctures and more coarse, warty-spiny surface structure; c) head of Camponotus cf. herculeanus (L.) (carpenter ant); d) head of a leaf hopper (Cicadellidae gen. and sp. indet.).

The upland vegetation during the sedimentation of the formation was dominated by forest-tundra and heaths, and an increase of summer temperatures of c. 7°C as compared to the present day is implied from floral composition and growth ring analyses (Bennike, in press).

The trees probably grew in sheltered areas with rather heavy snow cover during the winter period. Where basal parts of branches are preserved they are always bent downwards as a result of loading by snow.

It is remarkable that none of the insect species from the Kap København Formation occurs in the High Arctic today and only one species (Pterostichus vermiculosus) is confined to the (Low) Arctic. However, a number of species (14) extend into the Low Arctic but are found as well in subarctic environments (Danks, 1981). By far the majority of the named species (90%) are mainly or partly found in the subarctic zone. Some species are temperate-boreal, extending into the subarctic zone (e.g., Notiophilus biguttatus and Grypus equiseti), and some are today boreo-alpine (Nebria rufescens, Tachyporus nimbicola, T. rulomus and Camponotus herculeanus).

Modern tree lines are made up of only one or two tree species, but the Kap København Formation contains remains of four or five species. The dominating tree was larch, which presently forms the tree line in Siberia, while in North America today spruce is the dominating conifer of the boreal

Although upland plants and insects are well represented in the Kap København Formation, wetland organisms are undoubtedly heavily overrepresented, because their remains are the most likely to be dispersed by streams. Seeds and fruits of the water plants Menyanthes trifoliata, Hippuris vulgaris and Potamogeton spp. are very common. Remains of freshwater invertebrates include statoblasts of freshwater bryozoans (Cristatella mucedo and Plumatella spp.), ephippia, head shields and shells of Cladocera, mandibles of Lepidurus sp. and egg cocoons of Turbellaria. Aquatic insects are represented by head capsules of chironomid larvae, fragments of larval Trichoptera and fragments of at least 23 species of beetles (many Dytiscidae and some species of Gyrinidae and Hydrophilidae). Overall these aquatic organisms confirm the picture of a subarctic climate and indicate the presence of an abundance of different freshwater biotopes in the ancient landscape, notably that well-vegetated, mesotrophic lakes were present. Especially interesting in this connection is the concise description by Larson (1975:367) of the typical habitat of Agabus bifarius as: "shallow temporary ponds situated in rough fescue prairie or in parkland areas. These ponds are typically shaded, at least in part, by willow, aspen and emergent vegetation as well as dense accumulations of Carex and grass stalks."

The presence of *Tolypella* cf. *nidifica* (Characeae) oospores may indicate that ion-rich waters were also present in the general area.

A large fraction of the beetles are hygrophilous species living in damp, luxuriant places close to the water (e.g., Nebria rufescens, Patrobus stygicus, Bembidion transparens, Olophrum spp., Tachyporus spp. and Tachinus spp.) (Lindroth, 1961-69, 1985-86; Campbell, 1973, 1979). However, a similar fraction occurs exclusively on open, sandy or gravelly shores and banks of lakes and streams. These are generally thermophilous, diurnal and highly active predators hunting by means of vision and, accordingly, dependent on a high incidence of sunshine during the active season (species of Elaphrus, Asaphidion, Bembidion, Notiophilus, Cicindela and Stenus). Also Aegialia spp. are intimately connected with sandy areas, especially seashores.

The majority of the identified insect species today exhibit a circumpolar or holarctic distribution (56%), while 26% are nearctic and 18% palaearctic. Some are presently found far away from Greenland: Elaphrus sibiricus is distributed from central Siberia and Mongolia eastwards to northeast China, Japan and Kamchatka (Goulet, 1983); Asaphidion alaskanum is restricted to Alaska, the Northwest Territories and Yukon (Canada) (Lindroth, 1961-69); Bembidion planiusculum is restricted to the northwestern coastland of North America (Alaska and British Columbia) (Lindroth, 1961-69) and Helophorus niger to northwestern Siberia (Angus, 1970). Only two species, Nebria rufescens and Bembidion grapii, are found in Greenland today (Böcher, 1988).

NOTES ON PICEA

Picea (spruce) remains include three cones and numerous needles. The cones are small, broadly elliptic in outline and compare well with modern *P. mariana* (black spruce) cones. However, the needles show rather large morphological variability, from short, thick ones with rounded apexes to longer, more slender needles with acute apexes.

According to some students, needles of northern North America spruces can be identified by their anatomy: needles of *P. mariana* and *P. rubens* (red spruce) have two lateral resin channels, while needles of *P. glauca* (white spruce) only have resin sacs or no resin cavities (Durell, 1916; Duman, 1957; J.H. McAndrews, pers. comm. 1987; Birks, 1976:402 [unpubl. appendix II cited]). Others consider this character unreliable (R.G. Baker, pers. comm. 1987; J.V. Matthews, pers. comm. 1987; H.E. Wright, pers. comm. 1987; Marco, 1932).

Studies of modern reference material were in complete agreement with Duman's results: *P. mariana* and *P. rubens* needles have two resin channels, a feature never observed in *P. glauca* needles. The resin sacs of the latter, if paired, may

however easily be confused with the resin channels of *P. mariana* and *P. rubens* if the needles are not cleared or are only sectioned at a few levels. While *P. mariana* and *P. glauca* cannot be separated by their morphology, *P. rubens* needles are generally larger than *P. mariana* needles. Twenty spruce needles of different morphology from the Kap København Formation were sectioned transversely at about 20 different levels. They all had two resin channels and are referred to *P. mariana* on account of this character and their size.

NOTES ON ARACITES

Two fossils of the mysterious extinct plant Aracites globosa are present in the material (Fig. 4). They are identical to the fossils named Aracispermum sp. by Matthews (1987) and Aracites Johnstrupii by Matthews (1989).

There is a great deal of confusion about the names *Aracites*, Aracites johnstrupii (Hartz) Nikitin and Aracispermum. The names were originally proposed by the Russian palaeobotanist P.A. Nikitin in the late 1940s. He died in 1950, but his main works were not published until later (P.A. Nikitin, 1957, 1965). The fossils described as Carpolithes Johnstrupii by Hartz (1909) belong to the genus *Myrica* (Kircheimer, 1938, 1957; Friis, 1985). Dorofeev (1963) transferred Aracites johnstrupii to the genus Aracispermum, a genus name first proposed by P.A. Nikitin in 1947 but not formally established until 1965 (P.A. Nikitin, 1965). However, V.P. Nikitin (1979) proposed to keep the two genera separate, a view adopted here. P.A. Nikitin (1957) proposed that the fossils are seeds belonging in the family Araceae. Apparently, Reid and Reid (1915) were the first to describe this fossil type. They classified it as fruits of an extinct *Hippuris* species. Furthermore, Katz and Katz (1964) proposed that Aracites belong in the family Cucurbitaceae.

AGE ESTIMATE AND CORRELATION

The Kap København Formation is primarily dated by its remains of mammals, molluscs, foraminifers and ostracodes, and secondly by palaeomagnetic and amino acid studies.

At the 18th International Geological Congress in London, in 1948, it was recommended to place the Plio-Pleistocene boundary at the base of the Calabrian Stage in Italy (King and Oakley, 1949). The age of the boundary thus defined is presently dated to c. 1.6 Ma (Backman et al., 1983; Tauxe et al., 1983). However, in the Netherlands and in other parts of northwest Europe the boundary is usually placed at 2.3 or 2.5 Ma (e.g., Zagwijn and Hager, 1987; Mangerud, 1986; Watts, 1988). Because of this discrepancy it is important to try to put an age on deposits in this time range.

The simultaneous occurrence in the uppermost part of the formation of the extinct rabbit *Hypolagus* and the extant hare *Lepus* suggests an age of c. 2.0 Ma (Repenning et al., 1987). The Foraminifera faunas (Funder et al., 1985; Feyling-Hanssen, 1986, 1987, in press) include some extinct species, among which is *Cibicides grossa*, from the lower part of the formation. In the North Sea area this species became extinct slightly before the Reuverian/Praetiglian boundary (King, 1983), dated to about 2.3 Ma by palaeomagnetic studies (Montfrans, 1971).

The marine molluscan fauna includes a number of species that originated in the Pacific Ocean and first entered the Arctic Ocean and the North Atlantic by the first opening of the Bering Strait about 3 Ma ago (Albertsson, 1978; Gladenkov, 1979). The molluscan fauna correlates best with the lower part of the Gubik Formation in northern Alaska, having such species as *Macoma balthica*, *Cyrtodaria kurriana* and *Trichotropis bicarinata* in common (Bennike, 1989; Marincovich, pers. comm. 1987). The lower part of the Gubik Formation is dated to the Late Pliocene (Repenning, 1983; Nelson and Carter, 1985).

Only minor parts of the Kap København Formation have been studied palaeomagnetically. These parts show reverse polarity (Abrahamsen and Marcussen, 1986) and were deposited prior to the Brunhes/Matuyama boundary at c. 0.7 Ma. Unfortunately, no reversals have been detected in the sedimentary sequence.

We conclude that the Kap København Formation is probably between 2.5 and 2.0 Ma old.

The fossil flora and insect fauna show closest resemblance to those from the Beaufort Formation on Meighen Island (Matthews, 1974b, 1977, 1979, 1987). The flora from Meighen Island also contains a mixture of boreal and arctic species, many of which are the same as those from Kap København. A conspicuous common feature is the predominance of the two carabid genera *Bembidion* and *Pterostichus*. However, the presence on Meighen Island of the plants *Pinus*, *Picea banksii*, *Comptonia*, *Claytonia* and *Physocarpus* led Matthews (1987) to suggest that the Meighen Island flora and fauna is somewhat older than the Kap København Formation. Recent ⁸⁷Sr/⁸⁶Sr analyses on *Arctica* sp. shell fragments suggest an age between 2.5 and 5.1 Ma (Matthews and Ovenden, 1990-this issue).

It is suggested that the warm climate in North Greenland was primarily a North Atlantic phenomenon. Evidence for a similar large temperature increase is recorded from England (Jenkins and Houghton, 1987) and from the Vörring Plateau off Norway (Eldholm et al., 1987). On the other hand, a similar climatic warming in this late time interval is not seen in the Krestovka section of northeast Siberia (Sher et al., 1979), in the Late Cenozoic sequences of Yukon (C. Schweger, pers. comm. 1987) or in the Cape Deceit sequence of western Alaska (Matthews, 1974a; Repenning, 1980).

CONCLUSIONS

The flora and fauna found in northernmost Greenland at the Tertiary-Quaternary boundary were indeed very different from those at present. Most notable is the occurrence of tree species in the area. Only 33 species of modern native beetles are known from all of Greenland, just two of which occur in North Greenland, where they are restricted to the warmer, interior parts (Böcher, 1988). This is in striking contrast to the fossil record of at least 120 species of Coleoptera from the Kap København Formation. No ants exist in Greenland today, but at least three species are found as fossils. The great majority of the named fossil insect species are today found in the subarctic zone and thus confirm the picture of a forest-tundra environment derived from the floral remains.

Altogether the flora and fauna of the Kap København Formation account for a fairly detailed picture of the palaeoecological conditions: a varied environment, very much characterized by a diversity of freshwater localities, in some places bordered by rich vegetation, in other places with open, sunny and sandy shores. Trees and possibly small forests were present in sheltered places, separated by tundra vegetation and more xeric heath areas.

Presumably the rather similar fossil faunas and floras found at Kap København and in other parts of northern North America in the Late Tertiary constituted a coherent biota surrounding the Arctic Ocean. Extensive areas of lowland tundra did not yet exist, but arctic plants such as the widespread *Dryas* were well established. Other old records of *Dryas* come from Meighen Island (Matthews, 1987).

Some "neobotanists" have hypothesized that the arctic flora did not originate until the Quaternary (e.g., Tolmatchev, 1966), but the fossil records from northern North America and Greenland show that it must have existed already in the Late Tertiary.

ACKNOWLEDGEMENTS

Special thanks are due to John Matthews (Ottawa) for invaluable help and support regarding the identification of plant and insect remains. Thanks also to M. Nuorteva (Helsinki), who identified insect galleries in the wood, to B. Bejer-Petersen (Copenhagen), who identified borings in larch seeds, and to G.R. Coope (Birmingham), S. Elias (Boulder), H. Goulet (Ottawa), W. Hofmann (Plön), V. Mahler (Århus), S. Andersen, O. Lomholt, O. Martin, M. Hansen and M. Holmen (Copenhagen), who helped with the identification of insect remains. E.M. Friis (Stockholm), F.Ju. Velichkevich (Minsk), K. Wasylikowa (Krakow) and R.G. West (Cambridge) are thanked for help with Aracites. C.A. Repenning (Denver) kindly read and commented on an early version of the manuscript. O.B. Berthelsen, B.W. Rasmussen and G. Brovad did the photographic work. The Carlsberg Foundation and the Commission for Scientific Research in Greenland are thanked for generous financial support.

APPENDIX 1. VASCULAR PLANTS REPRESENTED BY MACROSCOPIC REMAINS IN THE KAP KØBENHAVN FORMATION

Equisetaceae
Equisetum sp.
Selaginellaceae
Selaginella selaginoides (L.) Link
Polypodiaceae
gen. and sp. indet.
Pinaceae
Picea mariana (Mill.) B.S.P.
Larix groenlandii Bennike
Cupressaceae
Thuja occidentalis L.
Taxaceae
Taxus sp.
Nymphaeaceae

Nuphar lutea type

Ranunculaceae

Ranunculus spp.
Anemone sp.
Papaveraceae
Papaver sect. Scapiflora sp.
Betulaceae
Betula nana L.
Betula alba L. s. l.
Alnus cf. crispa (Ait.) Pursh
Polygonaceae
Oxyria digyna (L.) Hill.
Polygonum sp.
Rumex acetosa L.
Caryophyllaceae
Cerastium cf. arcticum Lge./alpinum L.
Stellaria sp.

?Arenaria sp.

Melandrium affine J. Vahl/angustiflorum (Rupr.) Walp. Violaceae Viola sp. Cruciferae Arabis cf. alpina L. Salicaceae Salix reticulata L. Salix spp. Myricaceae Myrica arctogale Bennike Ericaceae Arctostaphylos uva-ursi (L.) Spreng. Cassiope tetragona (L.) D. Don. Ledum palustre L. Andromeda polifolia L.

APPENDIX 1. (continued)

Oxycoccus palustris Pers. Vaccinium uliginosum L. ssp.

microphyllum Lge.

Empetraceae

Empetrum nigrum L. s.l.

Crassulaceae

Sedum annuum L.

Fabaceae

Hedysarum sp.

Rosaceae

Potentilla palustris (L.) Scop.

Potentilla spp. Drvas octopetala L.

Rubus chamaemorus L. R. arcticus L./saxatilis L.

Hippuridaceae

Hippuris vulgaris L.

Geraniaceae Erodium sp.

Cornaceae

Cornus stolonifera Michx.

C. canadensis L. Cornus sp. A Menyanthaceae

Menyanthes trifoliata L.

Caprifoliaceae

Viburnum cf. edule (Michx.) Raf.

Scrophulariaceae ?Linaria sp.

Potamogetonaceae

Potamogeton natans L.

P. cf. perfoliatus L/richardsonii (Benn.)

Rvdh.

P. cf. gramineus L. P. alpinus Balb. P. cf. vaginatus Turcz.

Juncaceae Juncus sp. ?Luzula sp. Сурегасеае Carex spp.

Scirpus microcarpus Presl.

Gramineae

Gen. and sp. indet.

Sparganiaceae

Sparganium angustifolium type

Insertae sedis

Aracites globosa (Reid & Reid) Bennike

Carpolithes sp. A

APPENDIX 2. FOSSIL INSECTS FROM THE KAP KØBENHAVN FORMATION

COLEOPTERA

Carabidae

Cicindela sp.

Carabus spp.

Nebria cf. rufescens (Ström, 1768)

Nebria spp.

Notiophilus cf. aquaticus (Linnaeus,

Notiophilus cf. biguttatus (Fabricius,

1770) Blethisa cf. catenaria Brown, 1944 Diacheila polita (Faldermann, 1835) Elaphrus lapponicus Gyllenhal, 1810 Elaphrus sibiricus Motschulsky, 1846

Elaphrus tuberculatus Mäklin, 1877

Miscodera arctica Paykull, 1798 Patrobus cf. stygicus Chaudoir, 1871

Patrobus spp.

Asaphidion alaskanum Wickham, 1919

Bembidion (Chrysobracteon) cf. lapponicum Zetterstedt, 1828 Bembidion (Chrysobracteon) sp. Bembidion (Plataphus) cf. planatum

Leconte, 1848

Bembidion (Plataphus) cf. planiusculum

Mannerheim, 1843

Bembidion (Plataphus) spp. Bembidion (Hirmoplataphus) sp.

Bembidion (Peryphus) cf. mckinlevi Fall. 1926

Bembidion (Peryphus) cf. grapii Gyllenhal, 1827

Bembidion (Peryphus) cf. sordidum

Kirby, 1837

Bembidion (Diplocampa) cf. transparens Gebler, 1829

Pterostichus cf. adstrictus Eschscholtz,

1823

Pterostichus (Cryobius) cf. brevicornis Kirby, 1837

Pterostichus (Cryobius) spp.

Pterostichus cf. vermiculosus Menetries,

Pterostichus haematopus Dejean, 1831 Agonum cf. exaratum Mannerheim, 1853 Amara cf. glacialis Mannerheim, 1853

Amara spp.

Dromius sp. (palaearctic)

Dytiscidae

Hydroporus spp. Oreodytes spp.

Hydroporinae spp.

Agabus bifarius (Kirby, 1837)

Agabus spp.

Ilybius spp.

Colymbetinae spp.

Dytiscidae spp.

Gyrinidae

Gyrinus spp.

Hydrophilidae

Helophorus (Meghelophorus) cf. niger

J. Sahlberg, 1880 Helophorus (Rhopalelophorus) cf. frater

Smetana, 1985 Helophorus spp.

Hydrophilidae gen. and sp. indet.

Catopidae

cf. Catops sp.

Liodidae

cf. Anisotoma sp.

Silphidae

Heterosilpha ramosa (Say, 1823)

Staphylinidae

Philonthus sp. Quedius sp.

Pycnoglypta sp.

Olophrum cf. rotundicolle (C. R.

Sahlberg, 1834)

Olophrum cf. consimile (Gyllenhal, 1810)

Eucnecosum cf. brachypterum

(Gravenhorst, 1802) Eucnecosum spp. Omaliinae spp.

Bledius spp.

Mycetoporus sp. Tachyporus cf. nimbicola Campbell, 1979 Tachyporus cf. rulomus Blackwelder, 1936

Tachinus spp. Stenus spp.

cf. Atheta sp. Aleocharinae

gen. and sp. indet.

Scarabaeidae Aegialia cf. sabuleti (Panzer, 1797)

Aegialia spp.

Elateridae

Hypnoidus spp.

Buprestidae

Galleries and exit holes in Larix wood

Byrrhidae

Simplocaria spp.

Anobiidae

Hadrobregmus cf. pertinax (Linnaeus,

1758)

Lathridiidae

cf. Corticaria sp.

Lathridiidae

gen. and sp. indet.

Cerambycidae

gen. and sp. indet. (galleries in Larix

wood)

Chrysomelidae

Hydrothassa sp.

cf. Graphops sp.

cf. Chrysomela sp.

cf. Galeruca sp. Chrysomelidae

gen. and sp. indet.

Apionidae

Apion spp.

Curculionidae

cf. Otiorhynchus spp.

Pissodes sp.

Grypus equiseti (Fabricius, 1775)

Notaris sp. Rhynchaenus sp. Dorytomus sp.

Rhyncolus sp. Curculionidae spp.

Scolytidae

Scolytus sp.

Pityophthorus sp.

gen. and sp. indet. (galleries in Larix

wood)

HYMENOPTERA

Siricidae

Urocerus cf. gigas (Linnaeus, 1758) (fly

holes in Larix wood)

Tenthredinidae

gen. and sp. indet.

Ichneumonidae: cf. Pimplinae sp.

> Ichneumonidae gen. and sp. indet.

Chalcididae

gen. and sp. indet.

Diapriidae

gen. and sp. indet.

Megastigmus sp. (bore holes in Larix

seeds)

Formicidae

Camponotus cf. herculeanus (Linnaeus,

1758)

(continued)

APPENDIX 2. (continued)

Camponotus sp.
Formica sp.
Formicinae
gen. and sp. indet.
Hymenoptera
gen. and sp. indet.
DIPTERA
Chironomidae
Tanypodinae
gen. and sp. indet.
Orthocladius spp.

Psectrocladius sp.
Orthocladiinae
gen. and sp. indet.
cf. Diamesa sp.
Chironomus spp.
Endochironomus sp.
Microspectra sp.
Corynocera ambigua Zetterstedt 1840
Chironomini spp.

Brachycera

gen, and sp. indet.

Schizophora spp.

LEPIDOPTERA
gen. and sp. indet.

TRICHOPTERA
gen. and sp. indet.

HEMIPTERA
cf. Cicadellidae

Cyclorrhapha

ABRAHAMSEN, N., and MARCUSSEN, C. 1986. Magnetostratigraphy of the Plio-Pleistocene Kap København Formation, eastern North Greenland. Physics of the Earth and Planetary Interiors 44:53-61.

REFERENCES

ALBERTSSON, K.J. 1978. Um aldur jardlaga á Tjörnesi. Náttúrufrædingurinn 48:1-8.

ANGUS, R.B. 1970. A revision of the beetles of the genus *Helophorus* F. (Coleoptera: Hydrophilidae). Subgenera *Orphelophorus* d'Orchymont, *Gephelophorus* Sharp and *Meghelophorus* Kuwert. Acta Zoologica Fennica 129:1-62.

BACKMAN, J., SHACKLETON, N.J., and TAUXE, L. 1983. Quantitative nannofossil correlation to open ocean deep-sea sections from Plio-Pleistocene boundary at Vrica, Italy. Nature 304:156-158.

BENNIKE, O. 1987. News from the Kap København Formation, Plio-Pleistocene, North Greenland. Polar Research 5 n.s.:339-340.

. 1989. Trichotropis bicarinata (Gastropoda) from the Plio-Pleistocene Kap København Formation, new to the fossil fauna of Greenland. Mededelingen van de Werkgroep voor Tertiaire en Kwartaire Geologie 26:137-143.

_____. In press. The Kap København Formation: stratigraphy and palaeobotany of a Plio-Pleistocene sequence in Peary land, North Greenland. Meddelelser om Grønland, Geoscience 23.

BIRKS, H.J.B. 1976. Late-Wisconsinan vegetational history at Wolf Creek, central Minnesota. Ecological Monographs 46:395-429.

BÖCHER, J. 1988. The Coleoptera of Greenland. Meddelelser om Grønland, Bioscience 26. 100 p.

_____. 1989. Boreal insects in northernmost Greenland: palaeoentomological evidence from the Kap København Formation (Plio-Pleistocene), Peary Land. Fauna norvegica B36:37-43.

CAMPBELL, J.M. 1973. A revision of the genus *Tachinus* (Coleoptera: Staphylinidae) of North and Central America. Memoirs of the Entomological Society of Canada 90. 137 p.

. 1979. A revision of the genus Tachyporus Gravenhorst (Coleoptera: Staphylinidae) of North and Central America. Memoirs of the Entomological Society of Canada 109. 95 p.

COLLINGWOOD, C.A. 1979. The Formicidae (Hymenoptera) of Fennoscandia and Denmark. Fauna Entomologica Scandinavica 8. Klampenborg: Scandinavian Science Press. 174 p.

COOPE, G.R. 1986. Coleoptera Analysis. In: Berglund, B.E., ed. Handbook of Holocene Palaeoecology and Palaeohydrology. Chichester: John Wiley & Sons. 703-713.

DANKS, H.V. 1981. Arctic Arthropods. A review of systematics and ecology with particular reference to the North American fauna. Ottawa: Entomological Society of Canada. 608 p.

DOROFEEV, P.I. 1963. Tretichnye flory Zapadnoj Sibiri. Moskva: Nauka. 345 p.

DUMAN, M.G. 1957. Resin cavity pattern in the needles of Black, Red and White Spruce. Bulletin of the Torrey Botanical Club 84:388.

DURELL, L.W. 1916. Notes on some North American conifers based on leaf characters. Proceedings of the Iowa Academy of Science 23:519-582.

EDLUND, S.A. 1983. Bioclimatic zonation in a High Arctic region: central Queen Elizabeth Islands. Geological Survey of Canada, Paper 83-1A:381-390.

ELDHOLM, O., THIEDE, J., and TAYLOR, E. 1987. Summary and preliminary conclusions, ODP Leg 104. In: Eldholm, O., Thiede, J., and Taylor, E., eds. Norwegian Sea. Proceedings Initial Reports of the Ocean Drilling Project 104A:751-771.

FEYLING-HANSSEN, R.W. 1986. Grænsen mellem Tertiær og Kvartær i Nordsøen og i Arktis, fastlagt og korreleret ved hjælp af benthoniske foraminiferer. Dansk Geologisk Forening, Årsskrift for 1985:19-33. . 1987. The biostratigraphic position of the Kap København Formation based upon its foraminifera. Polar Research 5 n.s.:345-346.

gen. and sp. indet.

_____. In press. Fossil foraminiferal assemblages from Kap København, North Greenland, and their stratigraphic position. Meddelelser om Grønland, Geoscience.

FREDSKILD, B., and RØEN, U. 1982. Macrofossils in an interglacial peat deposit at Kap København, North Greenland. Boreas 11:181-185.

FRIIS, E.M. 1985. Angiosperm fruits and seeds from the Middle Miocene of Jutland (Denmark). Det Kongelige Danske Videnskabernes Selskab, Biologiske Skrifter 24(3). 102 p.

FUNDER, S., and HJORT, C. 1980. A reconnaissance of the Quaternary geology of eastern North Greenland. Grønlands Geologiske Undersøgelse, Rapport 99:99-105.

FUNDER, S., BENNIKE, O., MOGENSEN, G.S., NOE-NYGAARD, B., PEDERSEN, S.A.S., and PETERSEN, K.S. 1984. The Kap København Formation, a late Cainozoic sedimentary sequence in North Greenland. Grønlands Geologiske Undersøgelse, Rapport 120:9-18.

FUNDER, S., ABRAHAMSEN, N., BENNIKE, O., and FEYLING-HANSSEN, R.W. 1985. Forested Arctic: Evidence from North Greenland. Geology 13:542-546.

GLADENKOV, Y.B. 1979. Comparison of late Cenozoic molluscan assemblages in northern regions of the Atlantic and Pacific Oceans. International Geology Review 21:880-890.

GOULET, H. 1983. The genera of Holarctic Elaphrini and species of *Elaphrus* Fabricius (Coleoptera: Carabidae): classification, phylogeny and zoogeography. Quaestiones Entomologicae 19:219-482.

HARTZ, N. 1909. Bidrag til Danmarks tertiære og diluviale flora. Danmarks Geologiske Undersøgelse, II. række, No 20. 292 p.

JENKINS, D.G., and HOUGHTON, S.D. 1987. Age, correlation and palaeoecology of the St. Erth Beds and the Corraline Crag of England. Mededelingen van de Werkgroep voor Tertiaire en Kwartaire Geologie 24:147-156.

KATZ, N., and KATZ, S. 1964. Die Eigentümlichkeiten der Pleistozänen Pflanzengesellschaften. Report of the VIth International Congress on the Quaternary, Warsaw 1961. Lódz:439-445.

KING, C. 1983. Cainozoic micropalaeontological biostratigraphy of the North Sea. Institute of Geological Sciences, Report 82(7). 40 p.

KING, W.B.R., and OAKLEY, K.P. 1949. Definition of the Pliocene-Pleistocene boundary. Nature 163:186-187.

KIRCHEIMER, F. 1938. Über das Alter der Braunkohle Jütlands. Braunkohle 1938(24):405-409.

______. 1957. Die Laubgewächse der Braunkohlenzeit. Halle: Veb Wilhelm Knapp Verlag. 672 p.

LARSON, D.J. 1975. The predacious water beetles (Coleoptera: Dytiscidae) of Alberta: Systematics, natural history and distribution. Quaestiones Entomologicae 11:245-498.

LARSSON, S.G. 1943. Myrer. Danmarks Fauna 49. København: G.E.C. Gads Forlag. 190 p.

LINDROTH, C.H. 1961-69. The Ground-Beetles (Carabidae, excl. Cicindelinae) of Canada and Alaska. 1-6. Opuscula Entomologica Supplementum 20, 24, 29, 33, 34, 35:1-1192. I-XLVIII.

_____. 1985-86. The Carabidae (Coleoptera) of Fennoscandia and Denmark. Fauna Entomologica Scandinavica 15, 1-2. 497 p.

LITTLE, E.L., Jr. 1971. Atlas of United States trees. Vol. 1. Conifer and important hardwoods. U.S. Department of Agriculture, Forest Service, Miscellaneous publication 1146. 200 maps.

MANGERUD, J. 1986. Spor efter istider og mellomistider. In: Fossen, A.B., ed. Bergensernes Fjellverden. Bergen: Hardanger Forlag. 12-20.

MARCO, H.F. 1932. The identification of the spruces (Picea) by needle structure. M.S. Thesis, College of Forestry, New York State University at Syracuse, New York. 29 p.

- MATTHEWS, J.V., Jr. 1974a. Quaternary environments at Cape Deceit (Seward Peninsula, Alaska): Evolution of a tundra ecosystem. Geological Society of America, Bulletin 85:1353-1385.
- . 1974b. A preliminary list of insect fossils from the Beaufort Formation, Meighen Island, District of Franklin. Geological Survey of Canada, Paper 74-1A:203-206.
- . 1977. Tertiary Coleoptera fossils from the North American Arctic.
 The Coleopterists Bulletin 31:297-308.
- . 1979. Late Tertiary carabid fossils from Alaska and the Canadian Archipelago. In: Erwin, T.L., Ball, G.E., and Whitehead, D.R., eds. Carabid beetles, their evolution, natural history and classification. The Hague: Junk. 425-445.
- . 1987. Plant macrofossils from the Neogene Beaufort Formation on Banks and Meighen islands, District of Franklin. Geological Survey of Canada, Paper 87-1A:73-87.
- . 1989. New information on the flora and age of the Beaufort Formation, Arctic Archipelago, and related Tertiary deposits in Alaska. Geological Survey of Canada, Paper 89-1D:105-111.
- MATTHEWS, J.V., Jr., and OVENDEN, L.E. 1990. Late Tertiary plant macrofossils from localities in arctic/subarctic North America: A review of the data. Arctic 43(4):364-392.
- MOGENSEN, G.S. 1984. Pliocene or early Pleistocene mosses from Kap København, North Greenland. Lindbergia 10:19-26.
- MONTFRANS, H.M. van. 1971. Palaeomagnetic dating in the North Sea Basin. Earth and Planetary Science Letters 11:226-235.
- NELSON, R.E., and CARTER, L.D. 1985. Pollen analysis of a Late Pliocene and Early Pleistocene section from the Gubik Formation of Arctic Alaska. Ouaternary Research 24:295-306.
- NIKITIN, P.A. 1957. Pliocenovye i chetvertichnye flory Voronezhskoj oblasti. Moscow: Nauka. 200 p.
- _____. 1965. Akvitanskaja semennaja flora Lagernogo Sada (Tomsk). Tomsk: Tomsk University. 119 p.

- NIKITIN, V.P. 1976. Flora Mamontovoj gory po semenam i Plodam. In: Saks, V.N., ed. Miocen Mamontovoj Gory. Moscow: Nauka. 131-194. ______. 1979. Nekotorye novye vidy rasteniy v Neogenovych florach Severo-
- Vostoka SSSR. In: Shilo, N.A., and Baranova, Yu.P., eds. Kotinental'nye tretichnye tolshchi Severo-Vostoka Azii. Novosibirisk: Nauka. 125-130.
- REID, C., and REID, E.M. 1915. The Pliocene floras of the Dutch-Prussian border. Mededeelingen van de Rijksopsporing van Delfstoffen 6. 178 p.
- REPENNING, C.A. 1980. Faunal exchanges between Siberia and North America. Canadian Journal of Anthropology 1:37-44.
- . 1983. New evidence for the age of the Gubik Formation Alaskan North Slope. Quaternary Research 19:356-372.
- REPENNING, C.A., BROUWERS, E.M., CARTER, L.D., MARIN-COVICH, L., Jr., and AGER, T.A. 1987. The Berengian ancestry of *Phenacomys* (Rodentia: Cricetidae) and the beginning of the modern Arctic Ocean borderland biota. U.S. Geological Survey, Bulletin 1687. 31 p.
- SHER, A.V., KAPLINA, T.N., GITERMAN, R.E., LOZHKIN, A.V., ARK-HANGELOV, A.A., KISELYOV, S.V., KOUZNETSOV, Y.V., VIRINA, E.I., and ZAZHIGIN, V.S. 1979. Scientific excursion on problem "Late Cenozoic of the Kolyma Lowland." XIV Pacific Science Congress, Tour guide XI. Moscow: Nauka. 115 p.
- TAUXE, L., OPDYKE, N.D., PASINI, G., and ELMI, C. 1983. Age of the Plio-Pleistocene boundary in the Vrica Section, southern Italy. Nature 304:125-129.
- TOLMATCHEV, A.I. 1966. Die Evolution der Pflanzen in arktisch-Eurasien während und nach der quaternären Vereisung. Botanisk Tidsskrift 62:27-36
- WATTS, W.A. 1988. Europe. In: Huntley, B., and Webb, T., III, eds. Vegetation history. Dordrecht: Kluwer Academic Publishers, 155-192.
- ZAGWIJN, W.H., and HAGER, H. 1987. Correlations of continental and marine Neogene deposits in the south-eastern Netherlands and the lower Rhine District. Mededelingen van de Werkgroep voor Tertiaire en Kwartaire Geologie 24:59-78.